



Remote sensing of PM_{2.5} during cloudy and nighttime periods using ceilometer backscatter

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Abstract. Monitoring PM_{2.5} (particulate matter with aerodynamic diameter $d \leq 2.5 \mu\text{m}$) mass concentration has become of more importance recently because of the negative impacts of fine particles on human health. However, monitoring PM_{2.5} during cloudy and nighttime periods is difficult since nearly all the passive instruments used for aerosol remote sensing are not able to measure aerosol optical depth (AOD) under either cloudy or nighttime conditions. In this study, an empirical model based on the regression between PM_{2.5} and the near-surface backscatter measured by ceilometers was developed and tested using 6 years of data (2006 to 2011) from the Howard University Beltsville Campus (HUBC) site. The empirical model can explain ~ 56 , ~ 34 and ~ 42 % of the variability in the hourly average PM_{2.5} during daytime clear, daytime cloudy and nighttime periods, respectively. Meteorological conditions and seasons were found to influence the relationship between PM_{2.5} mass concentration and the surface backscatter. Overall the model can explain ~ 48 % of the variability in the hourly average PM_{2.5} at the HUBC site when considering the seasonal variation. The model also was tested using 4 years of data (2012 to 2015) from the Atmospheric Radiation Measurement (ARM) Southern Great Plains (SGP) site, which was geographically and climatologically different from the HUBC site. The results show that the empirical model can explain ~ 66 and ~ 82 % of the variability in the daily average PM_{2.5} at the ARM SGP site and HUBC site, respectively. The findings of this study illustrate the strong need for ceilometer data in air quality monitoring under cloudy and nighttime conditions. Since ceilometers are used broadly over the world, they may provide an important supplemental

source of information of aerosols to determine surface PM_{2.5} concentrations.

1 Introduction

The adverse impacts of high PM_{2.5} (particulate matter with aerodynamic diameter $d \leq 2.5 \mu\text{m}$) mass concentration on human health have been found from epidemiological studies around the world (Samet et al., 2000; Pope et al., 2009; Krewski et al., 2009). PM_{2.5} concentration has been found to be associated with cardiopulmonary disease, lung cancer, and an increased morbidity and mortality (Schwartz et al., 1996; Gent et al., 2003, 2009; Dominici et al., 2006; Bell et al., 2007; Franklin et al., 2007; Slama et al., 2007; Pope et al., 2002; Miller et al., 2007; Lepeule et al., 2012). As an official norm to stand for fine-particle abundance, PM_{2.5} mass concentrations are monitored widely by the US Environmental Protection Agency (EPA) through in situ instruments at surface monitoring sites. However, the number of EPA monitoring sites is limited. Therefore, remote sensing of PM_{2.5} from ground stations and satellites is desirable, allowing for fuller coverage of PM_{2.5} concentration between the EPA surface sites.

Aerosol optical depth (AOD) plays an important role in the remote sensing of PM_{2.5} since it has a good relationship with PM_{2.5} concentration. However, most measurements of AOD which are derived from passive remote-sensing techniques are only available under daytime and clear-sky conditions. Remote sensing of PM_{2.5} during either cloudy or nighttime periods is very rare. Different from passive instruments which measure column-integrated AOD, active

instruments like advanced lidars have the capacity to provide the vertical distribution of aerosol backscatter coefficient even under cloudy conditions or at nighttime. However advanced lidar networks are rare due to the complexity and cost. Instead, ceilometers which are simple, automatically operating single-wavelength lidars are used broadly all over the world. Ceilometers were originally developed for cloud-based height retrieval. With the improvement of accuracy and power, the potential capabilities of ceilometers to detect mixing layer height and aerosol optical properties have been explored recently (Münkel et al., 2007; Markowicz et al., 2008; Heese et al., 2010; Tsaknakis et al., 2011; Wiegner and Geiß, 2012). Another distinct advantage of ceilometers is their small overlap distance, which makes them suitable to detect aerosol information near the surface. PM_{2.5} concentration is an index of fine-particle mass concentration near the surface, while AOD is the integration of aerosol extinction in the total atmospheric column. So using aerosol backscatter near the surface has an inherent advantage in the remote sensing of PM_{2.5} concentration.

There are extensive studies investigating the PM_{2.5}–AOD relationship by the use of either an empirical statistical method (Engel-Cox et al., 2004; Liu et al., 2005, 2009; Gupta et al., 2006; Koelemeijer et al., 2006; Gupta and Christopher, 2008; Paciorek et al., 2008; Di Nicolantonio et al., 2009; Schaap et al., 2009; Lee et al., 2012; Sorek-Hamer et al., 2013; Strawa et al., 2013; Chudnovsky et al., 2014; Hu et al., 2013, 2014; Ma et al., 2014) or a chemical transportation model (Liu et al., 2004; Van Donkelaar et al., 2006, 2010; Kessner et al., 2013; Xu et al., 2015). In these studies, aerosol vertical distributions are estimated based on model simulation or under an assumption that aerosols are well mixed within the boundary layer and then decrease exponentially with height. Recently Li et al. (2016) developed an algorithm combining the backscatter measured from ceilometers with AOD for the PM_{2.5} retrieval. That work showed the capability of the ceilometer to improve PM_{2.5} estimation by introducing measurements of aerosol optical properties near the surface. Although there are a plenty of studies on PM_{2.5} estimation, studies on the remote sensing of PM_{2.5} during either cloudy or nighttime periods are rare due to the limitation of measurements of AOD.

In this study, to estimate PM_{2.5} under cloudy or during night periods, we developed a regression model based on the relationship between PM_{2.5} and the ceilometer backscatter under different meteorological conditions. The model is tested and validated against the 6 years (2006–2011) of ground-based observations of ceilometer backscatter, PM_{2.5}, AOD and meteorological conditions at the Howard University Beltsville Campus (HUBC) site and the 4 years (2012–2015) of data from the Atmospheric Radiation Measurement (ARM) Southern Great Plains (SGP) site. The data and model are described in Sect. 2. The results of the testing and evaluation of the model are illustrated in Sect. 3. The discussion is given in the last section.

2 Data and model

2.1 Data

In this study, the data were obtained from the HUBC site in Beltsville, MD, which is situated in a rural–suburban transition region between Washington, DC and Baltimore, MD urban centers. The site has a wide range of collocated instruments to observe atmospheric radiation, aerosol, cloud properties, meteorological conditions and air quality (Li et al., 2016), which makes the HUBC site suitable for PM_{2.5} investigations.

The backscatter data were provided by a Vaisala CT25k ceilometer, which is a single-lens lidar system equipped with a pulsed near-infrared diode laser (905 nm). As a commercial ceilometer, the CT25k provides a range-corrected attenuated backscatter coefficient, but the raw data are not available to the customer, which limits the access of the correction process. However, it has been shown that the signal reduction due to the near-field problem can be compensated for well with the manufacturer's correction (Markowicz et al., 2008). The unique single-lens design gives full overlap of the transmitter and receiver field of view at an altitude of 0 m (Münkel et al., 2007), which allows CT25k ceilometers to obtain high signal-to-noise ratio for lidar return signals at a low altitude. The working wavelength of CT25k ceilometers is ~905 nm where water vapor absorption exists (Wiegner et al., 2014; Wiegner and Gasteiger, 2015). However, water vapor impacts on backscatter retrieval are smaller than ~2% for 905 nm ceilometers under midlatitude climatology (Wiegner and Gasteiger, 2015) within a short distance from the surface to the height of 150 m. Given the small attenuation within a short distance, the attenuated backscatter coefficient below 150 m can be reasonably taken as a measure of backscatter coefficient when there is no rain or fog. The vertical resolution of the CT25k is 30 m. Since we are interested in the PM_{2.5} concentration near the surface, we only use the first five layer backscatter measurements from the CT25k ceilometer to estimate PM_{2.5} concentrations. The choosing of 150 m is arbitrary, but the sensitivity test showed that the retrieval results are quite similar for the different heights from 90 to 300 m (Li et al., 2016). It worth noting that the CT25k is an uncalibrated instrument of which the response may change in time. That change may induce differences in retrieving attenuated backscatter at different times, especially in retrieving backscatter at high altitude where the signal-to-noise ratio is small (Kotthaus et al., 2016). To estimate the impacts of ceilometer response changing on backscatter measurements, we compared 3 years (2007, 2008, 2009) of yearly average ceilometer backscatter profiles under nighttime clear conditions (PM_{2.5} < 15 µm, relative humidity < 40%). The method is similar to the method used in Kotthaus et al. (2016) to illustrate impacts of background signal and cosmetic shift on ceilometer-reported signals. The largest difference of the yearly average backscatter

below 150 m among the 3 years is found within 10 % for the CT25k ceilometer.

The near-surface meteorological conditions – including temperature (T), relative humidity (RH), pressure, wind speed (W) and wind direction – are provided by a nearby 31 m micrometeorological tower, and the AOD observations and cloud optical depth (COD) are retrieved from a Multifilter Rotating Shadowband Radiometer (MFRSR). The details of the MFRSR and the corresponding retrieval algorithms are introduced in Harrison et al. (1994), Harrison and Michalsky (1994) and Min and Harrison (1996). The hourly average PM2.5 are measured by a Met One BAM 1020 (beta ray attenuation monitor) from the collocated Maryland Department of the Environment (MDE) monitor station (Li et al., 2016).

In this study, hourly average data were used for all the data sets. Precipitation and fog cases were screened out by using cloud effective radius larger than 15 μm , microwave-radiometer-measured liquid water path larger than 200 g m^{-2} , ceilometer-derived cloud layer lower than 200 m and relative humidity larger than 95 %.

2.2 Model

For a ceilometer, the energy observed is a function of backscattering coefficient

$$P(x) = \frac{P_0 A \eta O(x) C \Delta t}{2x^2} \beta(x) T^2(x), \quad (1)$$

where $P(x)$ and P_0 are the received and emitted powers from a ceilometer; A and η are the area of the receiver and its efficiency, respectively; and x is the range from receiver to scattering volume. $O(x)$ is overlap function, C is light speed, Δt is the laser pulse duration and $T(x)$ is the transmittance of the atmosphere between receiver and scattering volume. $\beta(x)$ is the backscattering coefficient, which can be separated into two components:

$$\beta(x) = \beta^m(x) + \beta^a(x). \quad (2)$$

where $\beta^m(x)$ and $\beta^a(x)$ denote the backscattering by molecules and aerosols, respectively. The aerosol backscattering can be derived from the total backscattering coefficient as the molecule scattering is well modeled by Rayleigh scattering. For the backscattering at the near-infrared wavelength, the contribution from molecules can be disregarded due to the rapidly decreased Rayleigh scattering with wavelength, so $\beta(x)$ is taken as $\sim \beta^a(x)$ in this study.

With the assumption that aerosol size distribution is bimodal lognormal and aerosol particles are spherical, Li et al. (2016) illustrated that both the extinction and PM2.5 can be expressed in terms of particle volume concentration (cv_i)

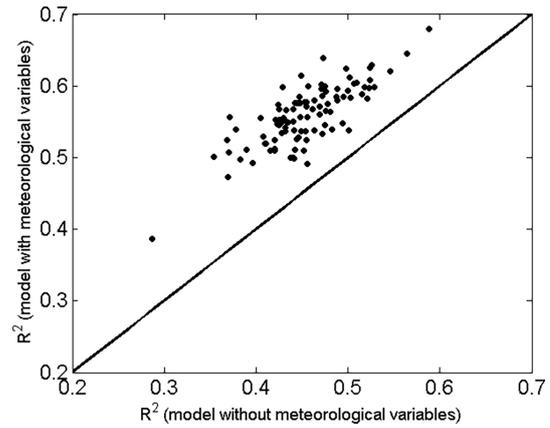


Figure 1. Comparison of R^2 out of the 100 independent cross-validations for the model without meteorological variables (Eq. 11) and the model with meteorological variables (Eq. 12) based on all the available daytime clear-sky cases at the HUBC site.

Table 1. Parameters based on the best fitting of 100 independent tests for Eq. (9).

Best-fitting parameters	a_0	a_1	b_1
Daytime clear	−97.61	66.95	0.14
Daytime cloudy	−100.00	94.02	0.05
Nighttime	−100.00	85.70	0.08

for each mode as

$$\text{ext}(\lambda) = \sum_{i=1}^2 cv_i h(R_i \sigma_i m \lambda), \quad (3)$$

$$\text{PM2.5} = \sum_{i=1}^2 cv_i g(R_i \sigma_i \rho), \quad (4)$$

where $h(R_i \sigma_i m \lambda)$ and $g(R_i \sigma_i \rho)$ are the integral functions of volume-concentration-normalized aerosol size distribution; c is the total particle volume concentration; v_i is the fraction of volume concentration for each mode i ; R_i and σ_i are the geometric mean radius and the standard deviation of aerosol size distribution, respectively; λ is the wavelength; m is the refractive index; and ρ is the particle mass density. The relationship between the aerosol backscattering coefficient $\beta^a(\lambda)$ and the extinction coefficient $\text{ext}(\lambda)$ at the wavelength λ is usually expressed by a lidar ratio (K):

$$K = \frac{\text{ext}(\lambda)}{\beta^a(\lambda)}. \quad (5)$$

From Eqs. (3), (4) and (5) the relationship between $\beta^a(\lambda)$ and PM2.5 can be expressed by

$$\text{PM2.5} = F \beta^a(\lambda). \quad (6)$$

Table 2. Parameters based on the best fitting of 100 independent tests for Eq. (10).

Best fitting parameters	c_0	c_1	c_2	c_3	c_4	d_1	d_2
Daytime clear	-10.50	3.49	-2.92	0.06	-0.11	0.07	0.55
Daytime cloudy	-14.49	12.86	-7.20	0.10	-0.49	0.12	0.32
Nighttime	-1.38	0.74	-0.13	0.029	-0.20	0.68	0.64

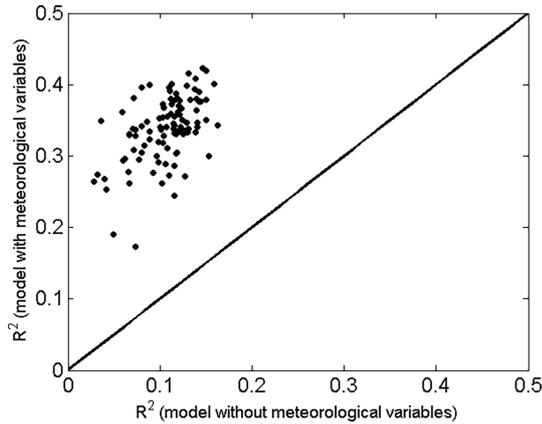


Figure 2. Same as Fig. 1 but just for cases under daytime cloudy conditions.

where

$$F = K \frac{\sum_{i=1}^2 v_i f(R_i, \sigma_i, m, \lambda)}{\sum_{i=1}^2 v_i g(R_i, \sigma_i, \rho)}, \quad (7)$$

The PM2.5/backscatter ratio F only depends on aerosol size and composition. Given that the variation of aerosol size and composition could be associated with the meteorological conditions and the assumption that aerosols mixed well near the surface, an empirical model based on the relationship between PM2.5 and the backscatter near the surface is proposed as

$$PM_{2.5} = a_0 + \left(a_1 + a_2 f(\text{RH}) + \sum_{i=1}^n a_{2+i} M_i \right) \left(\int_0^z \beta(x, \lambda) dx \right)^{b_2} + \varepsilon, \quad (8)$$

where the hygroscopic grow factor is expressed as

$$f(\text{RH}) = \frac{1}{(1 - \text{RH})^{b_1}}.$$

RH is relative humidity; M_1 through M_n are the meteorological factors including surface temperature, wind speed, wind direction and surface pressure; z is height; a_0 through a_{2+n} ,

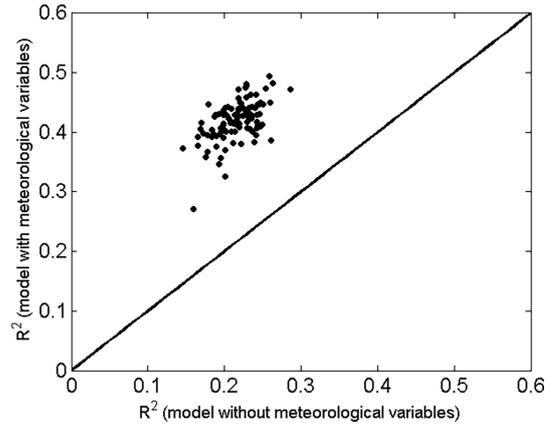


Figure 3. Same as Fig. 1 but just for cases during nighttime periods.

b_1 and b_2 are the regression coefficients; and ε is the error term. In the following part, we will test the model performance without considering the meteorological variables. In that case, Eq. (8) can be expressed as

$$PM_{2.5} = a_0 + a_1 \left(\int_0^z \beta(x, \lambda) dx \right)^{b_1} + \varepsilon. \quad (9)$$

When we test the model including the impacts from observations of surface T , RH and W , Eq. (8) can be expressed as

$$PM_{2.5} = c_0 + \left(c_1 + c_2 \times \frac{1}{(1 - \text{RH})^{d_1}} + c_3 T + c_4 \times W \right) \left(\int_0^z \beta(x, \lambda) dx \right)^{d_2} + \varepsilon. \quad (10)$$

3 Results

Overfitting can occur when a regression model is too complex. The overfitted model describes random error or noise instead of the underlying relationship. To test and evaluate the model, cross-validations (CVs) are implemented on the 6 years of hourly average measurements at the HUBC site under the different conditions including daytime clear, daytime cloudy and nighttime periods. For the cross-validation, we randomly select 90 % of the data as a training data set, use the remaining 10 % to test the models and repeat the procedure 100 times to avoid random bias and misleading R^2 induced by overfitting. Cross-validations are conducted for each model under each condition.

3.1 Simulation results under different sky conditions

Under daytime clear-sky conditions when AOD measurements from the MFRSR are available (no cloud, daytime),

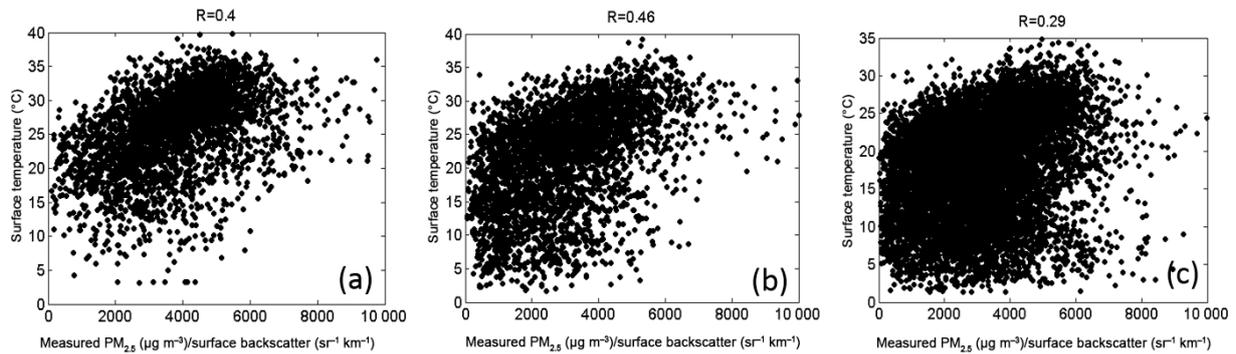


Figure 4. The relationship between surface temperature and PM2.5/backscatter ratio for (a) daytime clear-sky cases, (b) daytime cloudy cases and (c) nighttime cases.

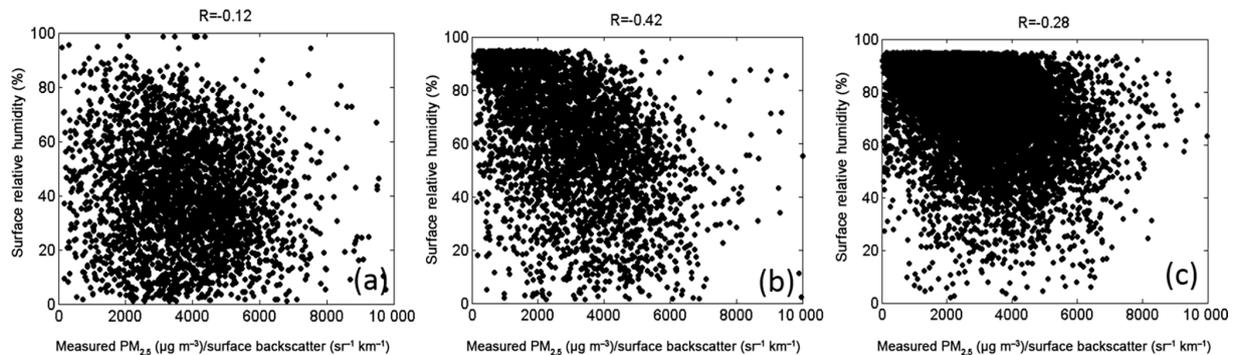


Figure 5. The relationship between surface relative humidity and PM2.5/backscatter ratio for (a) daytime clear-sky cases, (b) daytime cloudy cases and (c) nighttime cases.

the average CV R^2 out of the 100 times random cross-validations for the model (Eq. 10) is 0.56 (Fig. 1) with a root mean square error (RMSE) of $6.12 \mu\text{g m}^{-3}$. This result is close to that of the nonlinear model which combines both AOD and the ceilometer backscattering ($\text{CVR}^2 = 0.60$, $\text{RMSE} = 5.83 \mu\text{g m}^{-3}$) developed by Li et al. (2016) and performed much better than that of the model using AOD only ($\text{CVR}^2 = 0.40$, $\text{RMSE} = 7.14 \mu\text{g m}^{-3}$; Li et al., 2016). Without considering the meteorological conditions (Eq. 9), the average CVR^2 of the model is 0.45 (Fig. 2), which is better than that of the model using AOD only (Li et al., 2016) but not as good as the model including meteorological variables. Based on the fitted parameters (the parameters of the best fitting are shown in Tables 1, 2) from the 100 independent cross-validations (10 % of the total data), the average correlation coefficient between all the in situ measured PM2.5 under daytime clear-sky conditions and the simulated PM2.5 from the model without meteorological variables is 0.68 and increased to 0.76 when meteorological variables were included (Eq. 10).

Remote sensing of AOD is commonly based on the measurements of spectral extinction of solar radiation due to aerosol scattering and absorption in the atmospheric column from passive instruments. However most passive instruments

cannot readily discern AOD from COD under cloudy conditions. So any PM2.5 remote-sensing method relying on passive AOD measurements cannot retrieve PM2.5 under cloudy conditions. However, measurements of backscatter under cloudy conditions are still available for ceilometers, which can help to determine the near-surface aerosol extinction when upper-layer clouds exist.

Under daytime cloudy conditions, the average CVR^2 of the model without meteorological variables is only 0.11 (Fig. 2), which means only around 11 % of the variability in the hourly PM2.5 can be explained by the model. When meteorological factors are considered, the model can explain 34 % of the variability. Based on the fitted parameters of the 100 independent cross-validations, the average correlation coefficient between all the in situ measured PM2.5 under daytime cloudy conditions and the simulated PM2.5 from the model without meteorological variables is only 0.34, and it improved to 0.59 when meteorological variables were included in the model.

During nighttime periods, passive measurement relying on solar radiation is not available, but active instruments like ceilometers are still able to measure regardless of solar radiation and have better signal-to-noise ratio because of the absence of background sunlight contamination. During night-

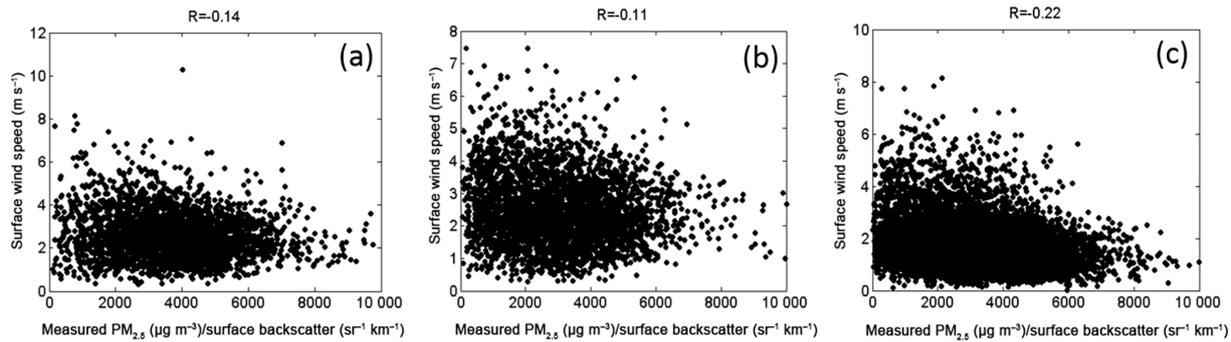


Figure 6. The relationship between surface wind speed and PM_{2.5}/backscatter ratio for (a) daytime clear-sky cases, (b) daytime cloudy cases and (c) nighttime cases.

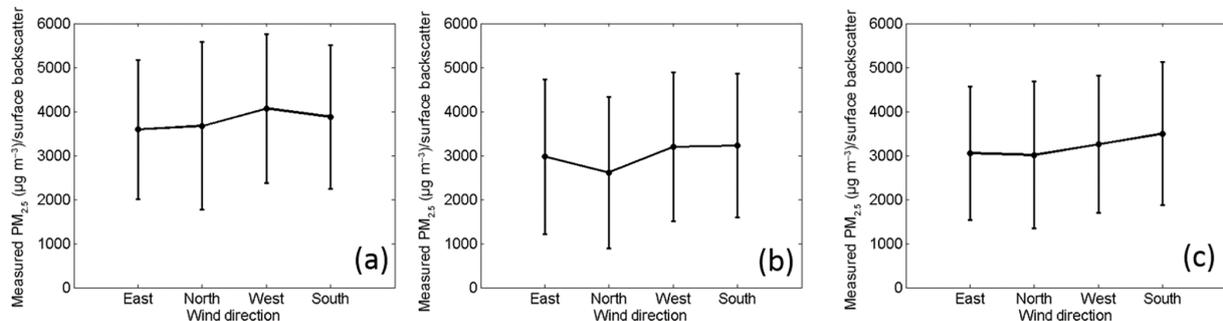


Figure 7. Average PM_{2.5}/backscatter ratio with standard deviation at four direction ranges – east (315 to 45°), north (45 to 135°), west (135 to 225°) and south (225 to 315°) – for (a) daytime clear cases, (b) daytime cloudy cases and (c) nighttime cases.

time periods, the average CVR² out of the 100 independent cross-validations for the model without meteorological variables is 0.21, while the average CVR² for the model with meteorological variables is 0.42 (Fig. 3). In this study, measurements under clear sky and cloudy sky were not separated during nighttime periods. Based on the fitted parameters of the 100 independent tests, the average correlation coefficient between all the in situ measured PM_{2.5} during nighttime and the simulated PM_{2.5} from the model without and with meteorological variables was 0.47 and 0.65, respectively.

3.2 Impacts from meteorological variables

The previous results showed that without considering meteorological factors the model predicting ability largely decreased, especially under cloudy and nighttime conditions. Remote sensing of PM_{2.5} using backscattering coefficients is based on the relationship between PM_{2.5} and aerosol backscatter which is determined by aerosol physical and chemical properties. Aerosol physical and chemical characteristics are sensitive and dependent on meteorological conditions that can impact aerosol transportation, hygroscopic growth and aerosol nucleation/creation. Therefore, meteorological conditions can be potentially used to estimate aerosol characteristics when the direct observations are not available. So taking into account the variations of meteorological con-

ditions may largely improve the model which is based on the regression between PM_{2.5} and backscattering coefficients.

To investigate impacts from different meteorological factors on PM_{2.5} remote sensing, the relationship between each meteorological variable and the PM_{2.5}/backscatter ratio were analyzed in three data categories: daytime clear (AOD measurements are available), daytime cloudy and nighttime (Figs. 4–7). Among the meteorological variables, temperature was found to have the most prominent positive correlation with the PM_{2.5}/backscatter ratio. The correlation coefficients are equal to 0.4, 0.46 and 0.29 under daytime clear, daytime cloudy and nighttime conditions, respectively (Fig. 4). In the eastern United States, sulfate dominates the aerosol chemical composition (Hand et al., 2012), and sulfate concentrations are expected to increase with increasing temperature due to faster SO₂ oxidation. Fine particles have smaller backscatter coefficients due to the smaller size index based on Mie theory (Wiscombe, 1980) than larger particles with the same PM_{2.5} mass concentration. So at the HUBC site, the increase of temperature associated with the high PM_{2.5}/backscatter ratio could be due to the increase of fine particles.

As opposed to the surface temperature, it is shown that the surface relative humidity had a prominent negative association with the PM_{2.5}/backscatter ratio. The correlation coefficient is equal to -0.12 , -0.42 and -0.28 under the

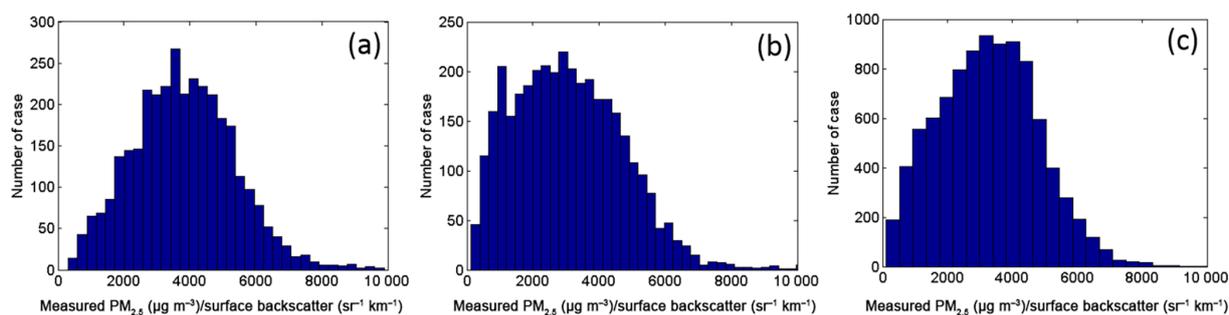


Figure 8. The number distribution of PM_{2.5}/backscatter ratio for (a) daytime clear cases, (b) daytime cloudy cases and (c) nighttime cases.

Table 3. Cross-validation tests of the model with different meteorological variables included.

Test	R^2 (RMSE)	95 % confidence intervals of R^2 (of RMSE)
Test1: model including all available meteorological variables	0.43 (6.70)	0.421–0.429 (6.672–6.736)
Test2: model without surface temperature	0.37 (7.01)	0.367–0.375 (6.981–7.044)
Test3: model without relative humidity	0.39 (6.91)	0.385–0.393 (6.880–6.946)
Test4: model without wind speed	0.37 (7.01)	0.368–0.375 (6.978–7.043)
Test5: model without wind direction	0.42 (6.71)	0.420–0.428 (6.683–6.742)
Test6: model without surface pressure	0.42 (6.71)	0.421–0.429 (6.674–6.738)
Test7: model not including any meteorological variable	0.21 (7.88)	0.203–0.209 (7.846–7.914)

daytime clear, daytime cloudy and nighttime conditions, respectively (Fig. 5). Under high-relative-humidity conditions there can be significant variations in the aerosol optical properties due to the aerosol hygroscopic growth effect. In the eastern United States, the dominant aerosols are composed of ammonium sulfate aerosols for which the ambient size will increase with the increase of the relative humidity due to hygroscopic growth. That can result in the decrease of the PM_{2.5}/backscatter ratio due to the increase of the aerosol extinction cross section while the aerosol dry mass is relatively invariant. It should be noted that the correlation coefficient is -0.12 for the cases under daytime clear conditions, while it is -0.42 under daytime cloudy conditions. Chu et al. (2015) showed that the effect of hygroscopic growth on extinction is more prominent when the relative humidity is larger. Under nighttime conditions, including both the clear and cloudy situations, the correlation coefficient is -0.28 .

A negative association is also found between the wind speed and PM_{2.5}/backscatter ratio under all the three conditions (Fig. 6). That may be explained by the association of higher PM_{2.5} concentrations with more stagnant, weaker wind conditions (Tai et al., 2010). Based on the average PM_{2.5}/backscatter ratio at four wind direction ranges – east (315 to 45°), north (45 to 135°), west (135 to 225°) and south (225 to 315°) – the variation of the mean PM_{2.5}/backscatter ratio at the four different wind directions was found to be small (within 10 %) compared to the standard deviation (~ 50 % of the mean value) at the HUBC site

(Fig. 7). The association of the surface pressure with the PM_{2.5}/backscatter ratio was found to be weak, with the correlation coefficient equal to -0.05 (not shown). The distributions of PM_{2.5}/backscatter ratio under the three conditions are shown in Fig. 8. Statistically, the PM_{2.5}/backscatter ratio under daytime clear-sky conditions is larger than that under daytime cloudy or nighttime conditions.

Figures 4–7 show the potential impacts of meteorological factors on model prediction. However, some information possibly overlaps among the different meteorological variables. To investigate the contribution of each meteorological variable to improving the model predicting power, the model was tested with different meteorological variable combinations. For each test, the cross-validation was randomly repeated 100 times based on all the available cases, including the daytime clear, daytime cloudy and nighttime periods.

Table 3 demonstrates the average CVR², RMSE and 95 % confidence intervals for each test. It is shown that without the information of surface temperature, relative humidity or wind speed the average CVR² of the model decreases from 0.43 to 0.37, 0.39 or 0.37, respectively. In other words, adding the variable of surface temperature, relative humidity or wind speed into the model can bring in additional information which may improve the model prediction capability regarding PM_{2.5}.

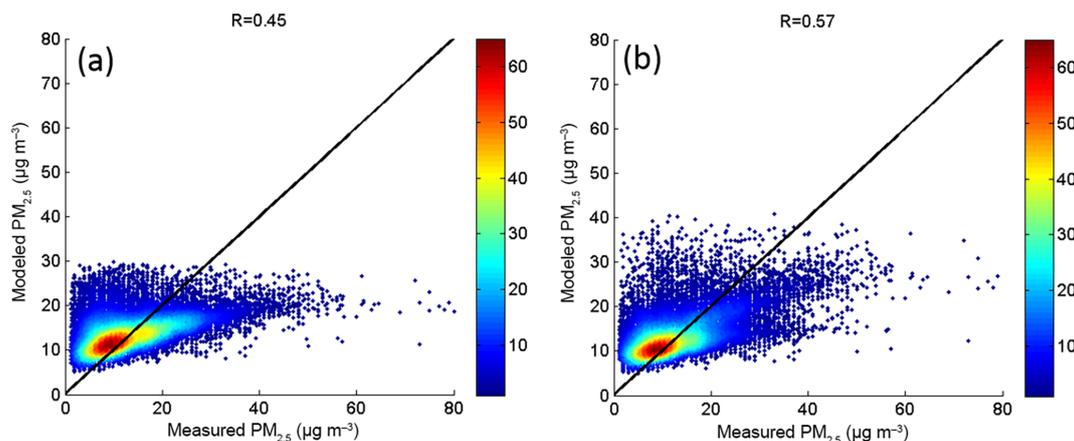


Figure 9. Comparison of measured PM_{2.5} and modeled PM_{2.5} when meteorological variables are not taken into account. (a) The model is non-seasonally fitted, and (b) the model is seasonally fitted. The colors stand for the number density of the points.

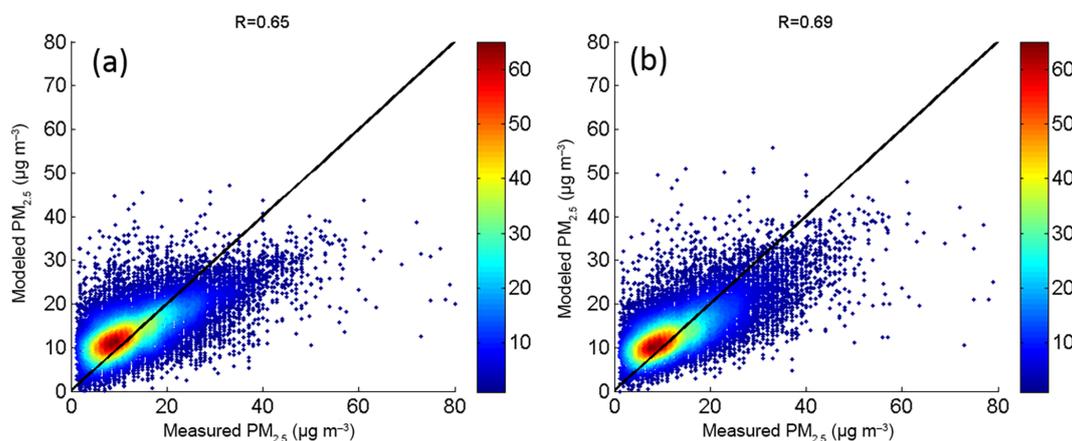


Figure 10. Same as Fig. 9 but with all meteorological variables taken into account.

3.3 Seasonally fitting

Besides meteorological factors, the seasonal variations of aerosol physical and chemical properties could impact the PM_{2.5}/backscatter ratio and then PM_{2.5} retrievals. To investigate the impacts of seasonal variations on PM_{2.5} retrievals, we fit the model seasonally and compared that performance with the model fitted on all the data without considering seasonal variation. Just as in the previous section, the cross-validations were implemented for each test. The parameters of the fitting with the median CVR² out of the 100 independent cross-validations for each fitting method are used to calculate the correlation between the in situ measurements of PM_{2.5} and the simulated PM_{2.5}. When meteorological variables were not considered, the simulated PM_{2.5} from the model with the seasonally fitted parameters had a much stronger association with the in situ measured PM_{2.5} ($R = 0.57$) than the model with the non-seasonally fitted parameters ($R = 0.45$; Fig. 9). When meteorological variables were taken into account, the correlation coefficient between

the simulation and the in situ measurements of PM_{2.5} for the model with the seasonally fitted parameters and the model with the non-seasonally fitted parameters was 0.69 and 0.65, respectively (Fig. 10), and the average CVR² was 0.48 and 0.43. The meteorological conditions have seasonal variation, so taking into account meteorological variables in the model can mitigate downside impacts of ignoring seasonal variations of aerosol properties on PM_{2.5} prediction.

3.4 Test in a different region

Given that aerosol types, aerosol compositions and meteorological conditions could be different in a different region, the model was tested based on the observations at the ARM SGP site, which is located in Oklahoma, USA. The site is in a rural area with fewer anthropogenic aerosols than the HUBC and the DC area. The ARM SGP site is the largest and most extensive climate research field site in the world. A newer version of the Vaisala Ceilometer CL31 has been used instead of the Vaisala CT25K since 2010 at the ARM SGP site.

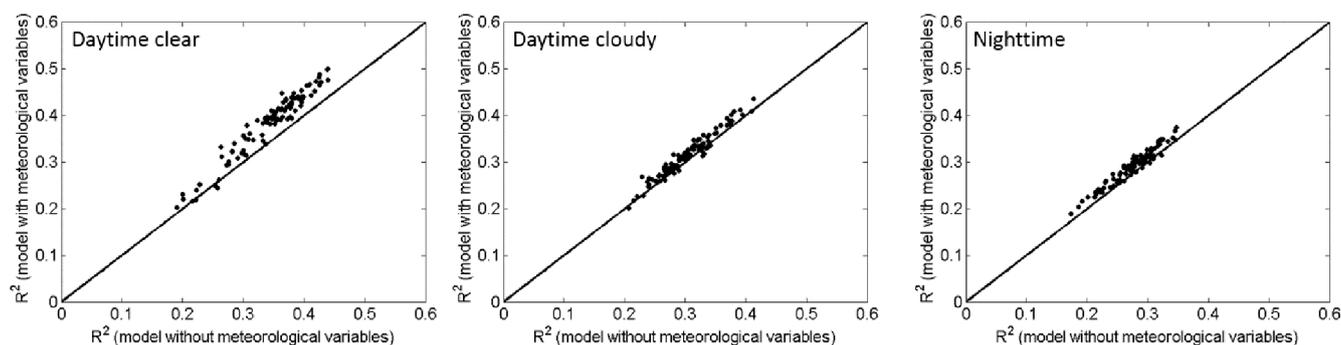


Figure 11. Comparison of cross-validation R^2 of the model without meteorological variables and model with meteorological variables during daytime clear, daytime cloudy and nighttime periods with the data from ARM SGP site.

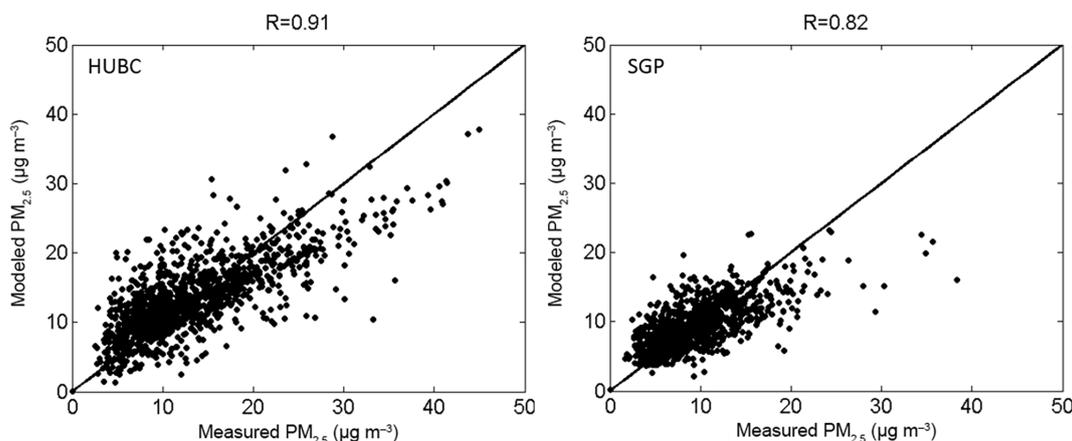


Figure 12. Comparison of daily average PM_{2.5} between in situ measurements and model simulation at the HUBC site and ARM SGP site.

The Vaisala CL31 has the same laser system, wavelength and measurement range as CT25k but improved spatial and temporal resolution and algorithms for cloud amount and mixing layer height detection (Münkel et al., 2007). To estimate the possible response changing of the CL31 ceilometer in different years, we compared the same 3 years (2012, 2013, 2014) of yearly average ceilometer backscatter profiles under nighttime clear conditions as used for the CT25K. The largest difference of the yearly average backscatter under nighttime clear conditions among the 3 years (2012, 2013 and 2014) is found within 3 % for the backscatter below 150 m. The principle of the algorithm (Eqs. 1 to 10) is applicable to most lidars with small overlap distance. So it is possible to use the model with the Vaisala CL31. In the test, we used 4 years (from 2012 to 2015) of measurements of Vaisala CL31 ceilometer backscatter and the surface meteorological conditions provided by the ARM SGP site and the FRM/FEM PM_{2.5} mass concentration from the nearest EPA site (36.697° N and 97.081° W; Air Quality System Data Mart, available via <http://www.epa.gov/airdata>). The same cross-validation procedure was implemented in the measurements at the ARM SGP site under daytime clear, daytime cloudy and nighttime periods. For the hourly average PM_{2.5},

the cross-validation results (Fig. 11) show that the performance of the model with meteorological variables (Eq. 10) at the ARM SGP site was not as good as that of the HUBC site, but the model without meteorological variables (Eq. 9) performed better at the ARM SGP site than at the HUBC site during daytime cloudy and nighttime periods. That could be due to the different aerosol type and composition which are associated with the hygroscopic growth of aerosols at the SGP area and the DC area. When the model (Eq. 12) is applied on the daily average PM_{2.5}, the average CVR² out of the 100 independent cross-validations is 0.82 and 0.66 at the HUBC site and ARM SGP site, respectively. That means the model (Eq. 12) can explain ~ 82 and ~ 66 % of the variability in daily average PM_{2.5} at the HUBC site and ARM SGP site, respectively. The correlation coefficient between the in situ measurements of PM_{2.5} and the simulation based on the fitted parameters of the fitting with the median CVR² out of the 100 independent cross-validations is 0.91 and 0.82, respectively (Fig. 12).

4 Discussion

Remote sensing of PM_{2.5} is generally based on AOD measurements due to its strong relationship with PM_{2.5}. For nearly all the passive instruments, the measurements of AOD rely on solar radiation. Ceilometers are compact, low-cost and unattended operational lidars and have been broadly used around the world. Although their laser power is relatively lower, the advantages of the small overlap distance and unattended and continuous operation make ceilometers suitable for remote sensing of aerosols near the surface. Moreover, the measurements of ceilometers do not rely on solar radiation, which makes them capable of retrieving aerosols during cloudy or nighttime periods.

In this study, an empirical model based on the regression between PM_{2.5} concentrations and ceilometer backscatter measurements was developed and tested with 6 years of observations at the HUBC site. The empirical model can explain ~ 56, ~ 34 and ~ 42 % of the variability in the hourly average PM_{2.5}, respectively, during the daytime clear, daytime cloudy and nighttime periods. During the daytime clear periods the prediction capability was close to that of the model combining AOD and backscatter (explain ~ 60 % of the variability) developed by Li et al. (2016), while during the daytime cloudy or nighttime period only the empirical model, which is independent of AOD, is available for the PM_{2.5} retrieval.

The impacts of meteorological conditions on the relationship between the in situ measured PM_{2.5} and the ceilometer-measured backscatter were analyzed. The prominent positive relationship found between the surface temperature and the PM_{2.5}/backscatter ratio could be due to the faster SO₂ oxidation under higher temperature given that the dominant aerosol chemical composition is sulfate in the eastern United States. The measured relative humidity showed a significant negative association with the PM_{2.5}/backscatter ratio, which could be due to hygroscopic growth of aerosols. The wind speed also showed a negative association with the PM_{2.5}/backscatter ratio, but the relationship between the measured wind direction and PM_{2.5}/backscatter ratio was found to not be obvious at the HUBC site. However, it is noteworthy that wind direction can be related to aerosol transportation and is usually associated with aerosol concentration and type. Although there was no significant association of the wind direction with the PM_{2.5}/backscatter ratio at the HUBC site, wind direction impacts could be significant at other places where transported aerosols like dust are found near the surface. Aerosol properties usually vary seasonally due to the seasonally varied meteorological conditions, large-scale transportation and local emission of anthropogenic and natural aerosols. Taking into account the meteorological conditions in the model can to some extent mitigate the seasonal impacts on the PM_{2.5} retrieval, and conducting the seasonal fitting can further improve the model predicting capability. Overall, the model with the seasonally fitted pa-

rameters can explain ~ 48 % of the variability in the hourly PM_{2.5} including during daytime clear, daytime cloudy and nighttime periods at the HUBC site. Aerosol physical and chemical characteristics which are associated with aerosol dry mass and optical properties could vary at different locations. So a test was implemented based on the observations from the ARM SGP site, which is geographically and climatologically different from the HUBC site. The results show that the impacts of meteorological conditions on the retrieval of PM_{2.5} using the ceilometer backscatter at the ARM SGP site are not as prominent as those at the HUBC site. That could be due to the different aerosol types in the SGP area and the DC area. In addition, the model parameters could be different for different aerosol types or in different climatic regions. That is because the relationship between PM_{2.5} and aerosol backscatter is related to aerosol types and sizes (Li et al., 2016), and the relationship between meteorological conditions and aerosols (i.e. size, composition) could vary with variation of aerosol types or climatic regions. Overall, the regression model using the ceilometer backscatter with meteorological variables could explain around 66 and 82 % of the variability in the daily average PM_{2.5} at the ARM SGP site and the HUBC site, respectively. It is worth noting that both the instrument hardware-related background signals and software-related artifacts could impact attenuated backscatter profiles observed by ceilometers. Further processing of attenuated backscatter profiles is needed to get accurate attenuated backscatter observations from ceilometers especially under low signal-to-noise ratio situations (Kotthaus et al., 2016). In this study, we only use the attenuated backscatter at low altitude, where both the Vaisala CT25k and CL31 have high signal-to-noise ratio for lidar return signals and hourly average will also decrease noise. So the potential background signals and systematic artifacts should have small impacts on the regression model performance. However, the parameters of the regression model could be different for different instruments.

The most important objectives of this study were to develop an algorithm for remote sensing PM_{2.5} during cloudy and nighttime periods by using ceilometer-measured backscatter. Retrievals of PM_{2.5} during cloudy or nighttime periods are very rare based on current remote-sensing methods. A large number of ceilometers have been used over the world, especially in the Europe and United States. The exploitation of the ceilometer on PM_{2.5} remote sensing could provide important information for air quality purpose, especially in helping to improve PM_{2.5} forecast over a larger area and can help fill the gaps among the EPA stations. Moreover that will largely increase the monitoring of air quality during cloudy or/and nighttime periods.

Data availability. EPA data: <https://aqs.epa.gov/api>. ARM SGP site: <http://www.archive.arm.gov/discovery>. Contact Siwei Li: siwei.li@howard.edu.

Competing interests. The authors declare that they have no conflict of interest.

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