



## Supplement of

## Increased aerosol content in the atmosphere over Ukraine during summer 2010

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**Figure S1** shows the number of AOD observations per day at the Kyiv AERONET site during June 1 - August 31, 2010.

**Figures S2–S34** represent HYSPLIT back trajectories of air parcels simulated for 168 hours (7 days) for the Kyiv site at 5 altitudes: 500 m, 1.5, 3, 4, and 5 km. We analyzed possible sources of aerosol from various regions from June 1 to August 31, 2010.

AOD changes in June were defined by both local sources and the transport with air masses mainly from the western areas. According to the weather conditions, at the beginning of June the air in the lower 5 km of the atmosphere was transported to Ukraine from the south, mainly from North Africa and the Mediterranean Sea through the Balkans and the Black Sea. Therefore, the relative maximum AOD during June 1–2 (Fig. **S2-S3**) was caused by local sources and, in addition, by particles transported from southern regions, as mixtures of continental soil dust and urban-industrial aerosol with the sea salt and possibly dust from Sahara. This is testified by both relatively low values of AE (in the Manuscript Fig. 5b), and prevailing coarse mode AOD over fine mode defined by spectral deconvolution algorithm (SDA, O'Neill et al., 2003).

Figure S1. Number of AOD observations per day at the Kyiv AERONET site during summer 2010.



Figure S2. June 1, 2010



During June 8–14 air masses in the altitudes 500 m–5 km were transported to Kyiv from the Atlantic across Scandinavia, the Baltic Sea, circulating over the Western, Central and Eastern Europe (**Fig. S4-S10**). The increase of AOD was caused by aerosol of different types: according to AE and SDA, on June 8–11 fine mode of aerosol prevailed. Further increase of AOD during June 12–13 was caused by the arrival of coarse mode aerosol, which explains the rapid decrease of AE (Fig. 5b in the Manuscript). Coarse aerosol was mainly of local origin, such as city transport and heavy industry, while the fine mode particle likely was brought by air currents from Europe.



NOAA HYSPLIT MODEL Backward trajectories ending at 0000 UTC 09 Jun 10 **GDAS** Meteorological Data ш 30.50 z 50.36 1 at \* Source Meters AGL 6000 5000 4000 3000 4500 3000 1500 500 1500 00 12 00 12 00 12 06/05 06/04 06/03 12 00 12 00 12 00 12 00 06/07 06/06 06/02

Figure S4. June 8, 2010

Figure S5. June 9, 2010



Figure S6. June 10, 2010



Figure S7. June 11, 2010



Figure S8. June 12, 2010



Figure 9. June 13, 2010



Figure S10. June 14, 2010

During June 23–26 AOD increase (Fig. 5a in the Manuscript) cannot be considered as reliable because of the small amount of data (9 measurements in the evening of June 23, 7 measurements on June 25 and 3 measurements on June 26, **Fig. S1**). Herewith, the dynamics of airflow during those days was more diverse (**Fig. S11-S13**). At 500 m air was transported to Kyiv mainly from Kazakhstan through the steppe zone of the ETR and eastern and south-eastern regions of Ukraine; at the altitudes of 1.5–5 km mainly from Turkey and the Caucasus through the Black and Azov seas, often changing direction and heights. In general, the accuracy of derived AOD values, obtained from a rather small number of individual measurements through short-term gaps in the clouds, can not be considered as sufficient. Because the algorithm of detection and elimination of the influence of clouds on measured AOD is not able to distinguish a thin uniform layer consisting of cloud water droplets from aerosol in the sunphotometer field of view. Therefore, we do not interpret relatively high AOD values for June 27–29 as a significant aerosol load over Kyiv. On June 30 the average AOD at 500 nm was equal approximately 0.35. It was computed for almost continuous dataset observed that day. This means that the aerosol content on June 30 increased comparing to the June 16–20.



Figure S11. June 24, 2010



Figure S12. June 25, 2010



Figure S13. June 26, 2010

During June 27–29 the dynamics of tropospheric circulation and the transport of air masses through Kyiv significantly changed (**Fig. S14–S16**). Preceding this period, since the beginning of June, air masses predominantly travelled to Kyiv mainly from the Atlantic throughout the European continent. However, during June 27–29 the path of air mass transport changed as seen in our back trajectory calculations that moved first southward and finally eastward. So that finally from June 29 on, air in above-mentioned altitudes was transported to Kyiv from Kazakhstan and Central Asia, across the Caspian Sea, the steppe zone of the ETR, and eastern regions of Ukraine. In addition, the 500 m back trajectory indicated a rapid change of the air motion from the north-west. On 30 June air motion remained unchanged (**Fig. S17**).



Figure S16. June 29, 2010



According to our analysis of air mass transport and spatio–temporal distribution of the brightness temperature of fires in Eastern Europe, it was seen that the wildfires were not the main source of increased concentrations of aerosol over Kyiv in June. The complex impact of continental particles and aerosol of marine origin was dominating. The relatively high AOD value on July 1 (see **Fig. S1**) was observed based on less than 10 measurements during that day. There were also no more measurements the day after, on July 2. During both days, wind directions at different altitudes varied much. Air entered Kyiv on July 1 at lower altitudes (500 m–1.5 km) passing central regions of the ETR on June 28–30. At 3–5 km altitude air came from the steppe area between the Caspian Sea and the eastern border of Ukraine. A large number of fires burned

during those days (Fig. 2d in the Manuscript). Thus, combustion products in the form of aerosol likely influenced the atmosphere over Kyiv and increased the AOD.

In the morning of July 2 air masses entered Kyiv from the North Atlantic at altitudes of 500 m–3 km; at the end of day also at altitude of 5 km (**Fig. S18**). In the following three days, July 3–5, when washout by rain cleaned the atmosphere in the evening of July 2nd, AERONET-measured AODs were temporarily low, with values below 0.3 at 440 nm. On these days cloudiness decreased significantly and more measurements could be taken into account, thus, decreasing the uncertainty of the derived time-series. Between July 3–5 air masses entered Kyiv from the North Atlantic and Arctic through various parts of Europe, particularly at lower altitudes (500 m–1.5 km) through the ETR (**Fig. S19-S21**).



Figure S18. July 2, 2010



Figure S20. July 4, 2010

NOAA HYSPLIT MODEL Backward trajectories ending at 1100 UTC 03 Jul 10 GDAS Meteorological Data



Figure S19. July 3, 2010

NOAA HYSPLIT MODEL Backward trajectories ending at 1000 UTC 05 Jul 10 GDAS Meteorological Data



Figure S21. July 5, 2010

During July 6–7 weather conditions changed (**Fig. S22-S23**). Air masses now travelled westward at altitudes 0.5–4 km, originating mainly from central Kazakhstan and the ETR. Those air masses crossed the steppe zone of the ETR and south-east Ukraine, where active fires were observed (Fig. 2d in the Manuscript) and caused the relatively high AODs over Kyiv (Fig. 5a in the Manuscript). Another rapid change in weather conditions and transport pathways occurred again on July 8–9: air masses came from the Atlantic coast of Western Europe, travelled eastward causing an AOD decrease over Kyiv (Fig.5a in the manuscript, Fig. **S24-S25**).



Figure S22. July 6, 2010



Figure S23. July 7, 2010



Figure S24. July 8, 2010



Figure S25. July 9, 2010

The aerosol load over Kyiv during July 10–27 (Fig. 5a in the Manuscript) was changing under the influence of air masses mainly coming from the Atlantic and the Arctic Ocean through the ETR, where fires burned (Fig. S26). These air masses crossed Kazakhstan and steppe zones of the ETR and through south-eastern Ukraine. Air masses were also circulating over the Central and Western regions of the ETR, Belarus and Ukraine and other neighboring countries, where at that time vegetation and peat fire were observed (Fig. S27). The number of fires and their intensity were gradually increasing towards end of July (Fig. 2 e,f in the Manuscript). Consequently, the increase of AOD over Kyiv during those days was caused by the accumulation of aerosol particles from the regions of active fires or formed under the influence of combustion products. Intermediate AOD decreases are caused by below cloud scavenging from occasional rains over Kyiv.



Rainy weather on July 27–28 over Kyiv site and a change in the transport pathways caused a significant decrease in AOD during July 29–30. Air masses entered Kyiv in the altitude range 500 m–5 km, coming from North Atlantic, passing Western Europe, the Mediterranean, the Balkans, Romania, Moldova, and western regions of Ukraine (**Fig. S28**).

During July 31–August 1 an increase of AOD over Kyiv was observed (Fig. 5a in the Manuscript), with air masses originating over the Atlantic and traveling throughout Scandinavia, Western and Southern Europe, the Mediterranean, the Balkans, later circulating over Ukraine, the Baltics, Poland, Romania and Moldova. Likely the increase of AOD was observed due to intense biomass burning in Ukraine and neighbouring territories (Fig. 2 g in the Manuscript). On August 2, air masses in the altitude range 500 m–3 km were transported to Kyiv from western and central regions of the ETR, from Kazakhstan throughout North Caucasus and the steppe zone of the ETR, as well as throughout south-eastern Ukraine (**Fig. S29**). In those regions intense forest, peat and vegetation fires were detected at that time (Fig. 2g in the Manuscript). Accordingly, these events were the reason of aerosol pollution over this region and Ukraine in particular till August 18.



Figure S28. July 28, 2010

Between August 2–18 AOD over Kyiv were increasing and reached maximum values on 15 August with about 1.26 at 500 nm, respectively. This is more than factor 2 larger than during the fire period before August 2. For the following 2 days, August 16–17, the AOD level remained high. It increased so much because air masses circulated over the entire region, from Ural and Kazakhstan to Central Europe and the Baltic (**Fig. S30-S31**).



Figure S30. August 7, 2010

Figure S31. August 14, 2010

Figure S29. August 2, 2010

During August 18–21 atmospheric fronts of an active cyclone, which moved from the south of the Baltics to Samara region, caused significant weather change with rains in Eastern Europe. **Figures S32-S34** show back trajectories for the Kyiv site at 00 UTC, 06 UTC, and 12 UTC respectively. Towards 12 UTC the direction of incoming air changed completely from the eastern to western. Aerosol washout cleaned the atmosphere over investigated sites and caused the sharp AOD decrease seen in the Manuscript, Fig. 5a. Until the end of analysing period, i.e. end of August 2010, the average AOD level over Kyiv remained as low as during less polluted days in June 2010.



Figure S32. 00 UTC, August 18, 2010

NOAA HYSPLIT MODEL - NASA/AERONET Run Backward trajectories ending at 1200 UTC 18 Aug 10 GDAS Meteorological Data



Figure S34. 12 UTC August 18, 2010

NOAA HYSPLIT MODEL - NASA/AERONET Run Backward trajectories ending at 0600 UTC 18 Aug 10 GDAS Meteorological Data



Figure S33. 06 UTC August 18, 2010

Data	SSA		Refractive Index				Observations
Date	440nm	870nm	Re 440nm	lm 440nm	Re 870nm	lm 870nm	Number/day
1	2	3	4	5	6	7	8
09.06.2010	0.95	0.92	1.38	0.006	1.38	0.005	1
12.06.2010	0.95	0.93	1.34	0.003	1.41	0.004	2
13.06.2010	0.97	0.97	1.36	0.002	1.41	0.002	3
14.06.2010	0.97	0.98	1.48	0.002	1.50	0.002	1
30.06.2010	0.92	0.91	1.50	0.008	1.52	0.007	1
10.07.2010	0.94	0.90	1.37	0.008	1.39	0.010	1
19.07.2010	0.93	0.87	1.43	0.008	1.46	0.011	4
20.07.2010	0.98	0.90	1.40	0.006	1.44	0.008	2
31.07.2010	0.94	0.89	1.41	0.007	1.43	0.009	3
02.08.2010	0.94	0.95	1.52	0.006	1.55	0.005	4
05.08.2010	0.94	0.92	1.49	0.006	1.53	0.006	1
07.08.2010	0.94	0.92	1.51	0.008	1.53	0.009	8
08.08.2010	0.92	0.89	1.51	0.009	1.56	0.009	4
10.08.2010	0.93	0.87	1.48	0.011	1.50	0.016	3
11.08.2010	0.94	0.90	1.46	0.008	1.49	0.011	8
12.08.2010	0.94	0.91	1.50	0.008	1.52	0.010	3
13.08.2010	0.93	0.88	1.48	0.009	1.51	0.012	10
14.08.2010	0.94	0.89	1.49	0.008	1.52	0.011	3
15.08.2010	0.96	0.95	1.52	0.006	1.53	0.006	8
16.08.2010	0.97	0.95	1.49	0.006	1.51	0.007	10

**Table S35** . Aerosol Single Scattering Albedo (SSA) and Real (Re) and imaginary (Im) parts of refractive index fromAERONET measurements at Kyiv site during June 1-August 31 2010.



**Figure S36**. AOD 532 nm distribution over Ukraine in summer 2010 from CALIOP measurements during 16 day period from June 17 to July 2 (a), from July 3 to 18 (b), from Jul 19 to August 3 (c), and during 12 day period from August 20 to 31 (d). Red numbers at the bottom of the map indicate dates of each daytime track running to north-west and blue numbers at the top of the map indicate the date of each nocturnal track running to south-west.

**Figures S37– S94** represent extinction coefficient of aerosol obtained from CALIPSO for the period August 7-12, 2010. Additional information is added to each figure and was taken directly from the CALIPSO Aerosols Profile Product Level 2.0 Version 3.01 and 3.02 data base. This additional information consists of:

- Date and Time
- Location
- AOD 532 nm
- Extinction Coefficient (plotted on Figures)
- Maximum of the Extinction Coefficient (Extinc.koef.max) and its uncertainty (Extinc.koef.max.Uncert)
- Average of the Extinction Coefficient (Extinc.koef.average)
- Median of the Extinction Coefficient (Extinc.koef.median)
- Heights of top and bottom of integral aerosol layers (Layer Height max, Layer Height min correspondingly)
- Height of Extinction Coefficient maximum (Height of Extinc.koef.max)
- Effective height of aerosol layer (Effective Height of Layer)

\* Average of the Extinction Coefficient (Extinc.koef.average) of integral aerosol layers were computed using data mentioned above.

Effective heights of integral aerosol layers were computed from Extinction Coefficient vertical profile data as

$$H_{Eff} = \frac{\int_{H_{Max}}^{H_{Max}} z \cdot C_{Ext}(z) \cdot dz}{\int_{H_{Max}}^{H_{Max}} C_{Ext}(z) \cdot dz}$$

where *z* is the vertical coordinate,

 $C_{Ext}(z)$  is the Extinction Coefficient,

 $H_{Min'}$ ,  $H_{Max}$  are respectively the bottom and top bounds of the integral aerosol layer.



















Figure S62.	5 <sub>7</sub>
Date: August 12, 2010	
Time(GMT) = 0h 44m 50.9s	4 -
Latitude, deg = 51.9065	
Longitude, deg = 25.5926	
AOD 532 = 0.51	
Extinc.koef.max = 0.96 km <sup>-1</sup>	
Extinc.koef.average = 0.13 km <sup>-1</sup>	
Extinc.koef.median = 0.06 km <sup>-1</sup>	
Layer Height max = 4.22913 km	<b>F</b>
Layer Height min = 0.277296 km	0 +
Height of Extinc.koef.max = 2.133 km	0,0 0,2 0,4 0,6 0,8 1,0
Effective Height of Layer = 2.273 km	Extinction Coefficient, km <sup>-1</sup>
Figure S63.	47.
Date: August 12, 2010	<b>```</b>
Time(GMT) = 0h 45m 13.9s	3
Latitude, deg = 50.5437	
Longitude, deg = 24.9904	
AOD 532 = 0.68	
Extinc.koef.max = 0.96 km <sup>-1</sup>	
Extinc.koef.average = 0.20 km <sup>-1</sup>	
Extinc.koef.median = 0.13 km <sup>-1</sup>	
Layer Height max = 3.810 km	<b>i-</b>
Layer Height min = 0.337 km	
Height of Extinc.koef.max = 3.271 km	0,0 0,2 0,4 0,6 0,8 1,0
Effective Height of Layer = 2.182 km	Extinction Coefficient, km <sup>-1</sup>
Figure S64.	4
Date: August 12, 2010	
Time(GMT) = 0h 45m 42.2s	3
Latitude, deg = 48.8696	5
Longitude, deg = 24.2911	
AOD 532 = 0.65	
Extinc.koef.max = 0.53 km <sup>-1</sup>	
Extinc.koef.average = 0.23 km <sup>-1</sup>	
Extinc.koef.median = 0.21 km <sup>-1</sup>	
Layer Height max = 3.511 km	
Layer Height min = 0.637 km	
Height of Extinc.koef.max = 0.936 km	0,0 0,1 0,2 0,3 0,4 0,5 0,6
Effective Height of Layer = 2.211 km	Extinction Coefficient, km <sup>-1</sup>



















