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Supplement of

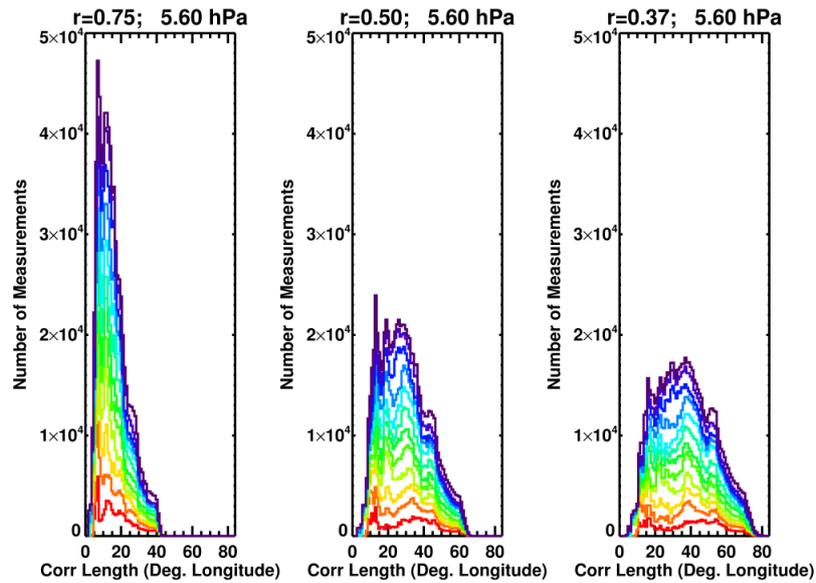
Model-based climatology of diurnal variability in stratospheric ozone as a data analysis tool

Stacey M. Frith et al.

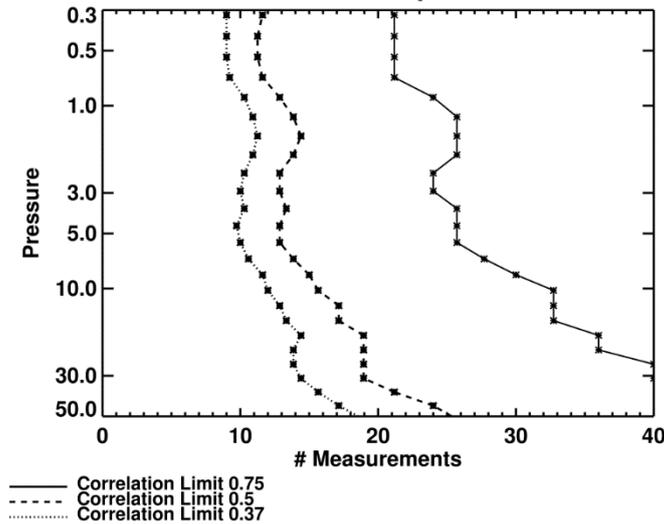
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Supplemental Figures



Independent Number of Measurements for Error Computation



5 Figure S1. Top: Distribution of “correlation length” in degrees longitude, based on longitudinally-lagged correlations at each model grid point. The correlation length is the lag (in degrees longitude) at which the correlation drops below the threshold. We compared threshold correlations of 0.75 (left panel), 0.5 (center panel), and 0.37 (right panel). Correlations are computed for the 15th day of each month in order to sample at different seasons. Different colors show the distributions for successive months. Bottom: The median of the distributions for each correlation threshold, plotted as a function of pressure, converted to number of independent measurements (corr. length/360). Based on this we assume 12 independent measurements (30° longitudinal sampling) at all levels each day. Thus, we use 12 * 60 days=720 independent measurements for computing standard error of the mean.

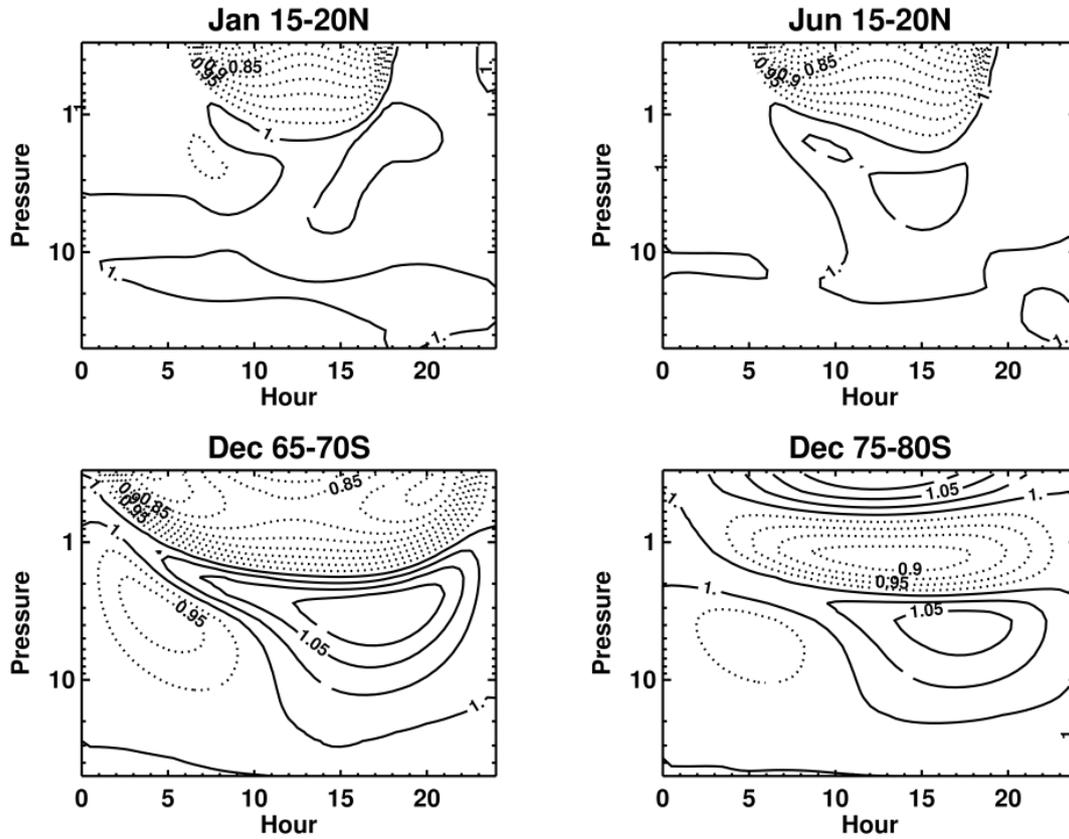


Figure S2. Same as Figure 1 but for January 15-20° N (top left); June 15-20° N (top right); December 65-70° South (bottom left); and December 75-80° S (bottom right). The top two panels are directly comparable to the diurnal cycle in Parish et al. [2014]. The bottom panels show the hemispheric symmetry of the polar summer signal as compared to the June Northern Hemisphere results shown in Figure 1.

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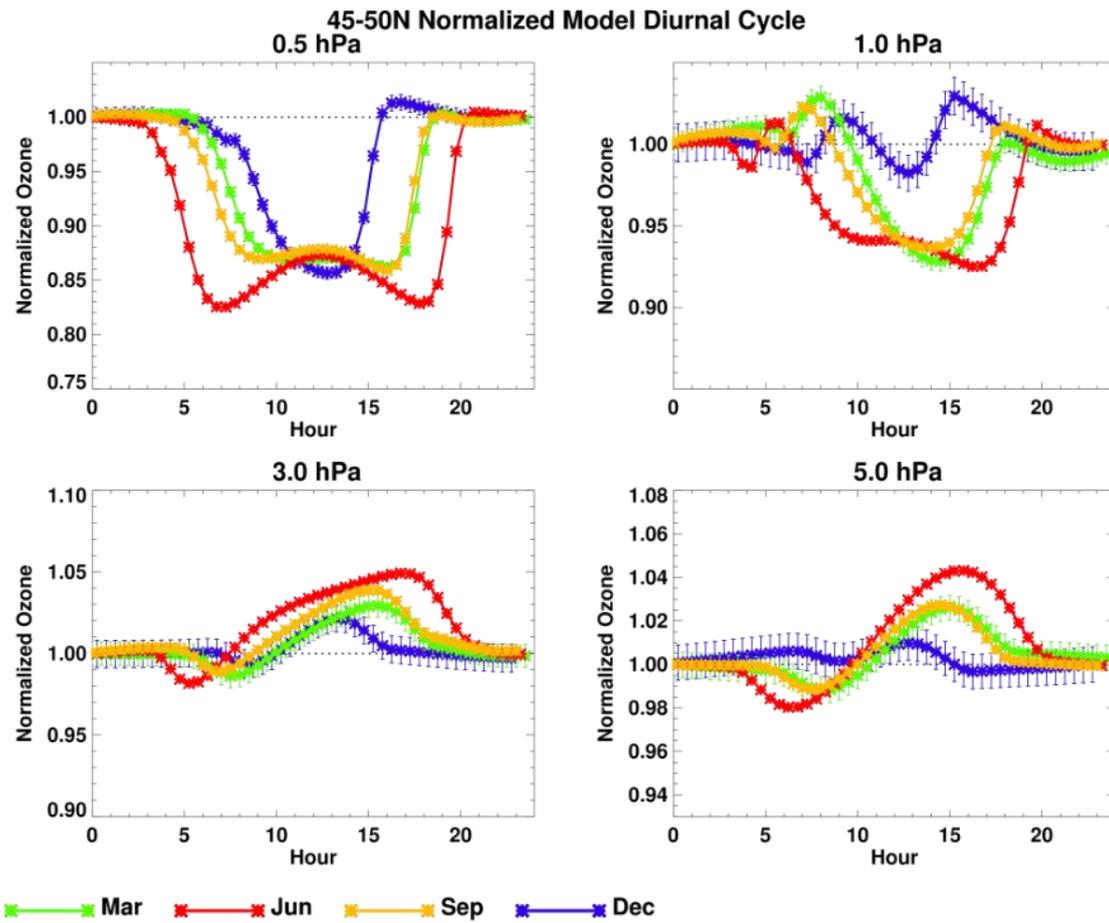


Figure S3. Same as Figure 2 but for 45-50° N. The 5 hPa result (bottom right panel) is directly comparable to Schanz et al. [2014; Figure 2].

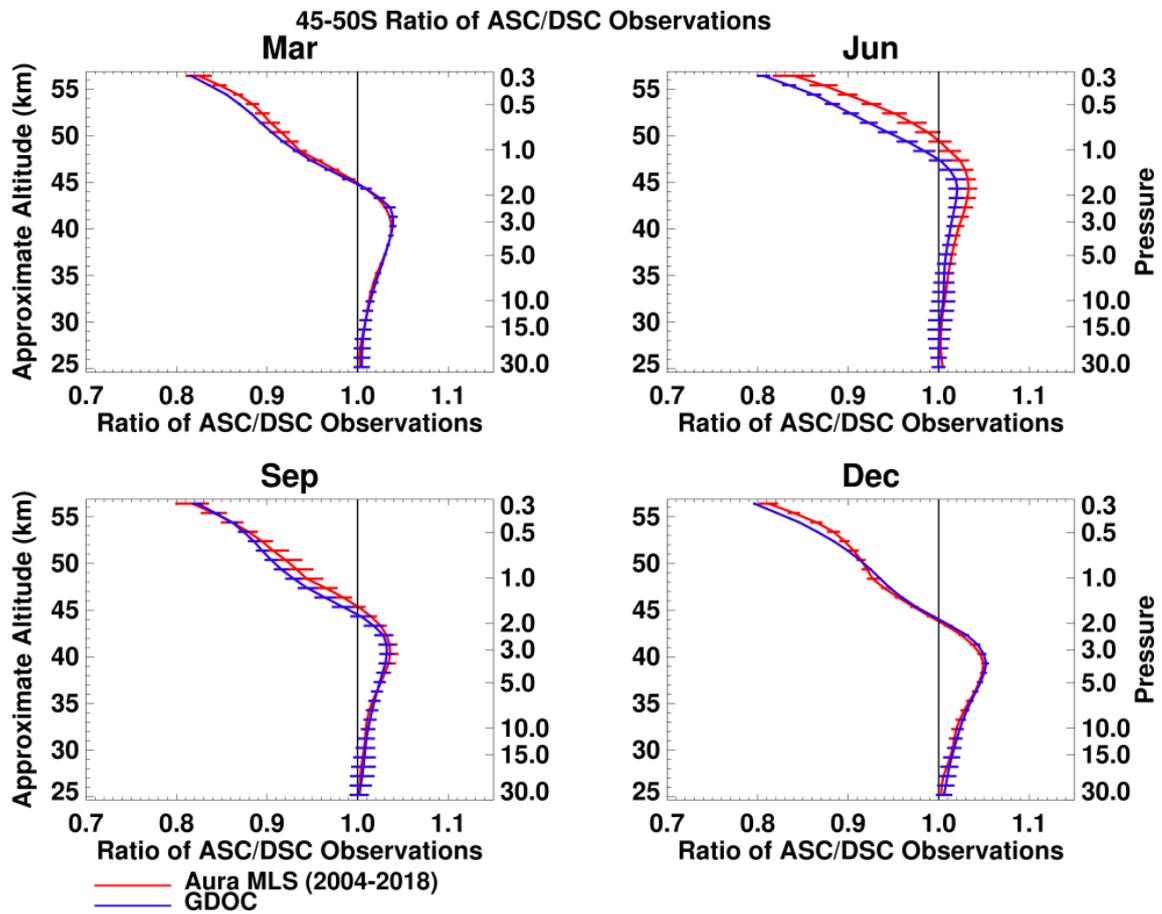


Figure S4. Same as Figure 6 but for 45-50° S.

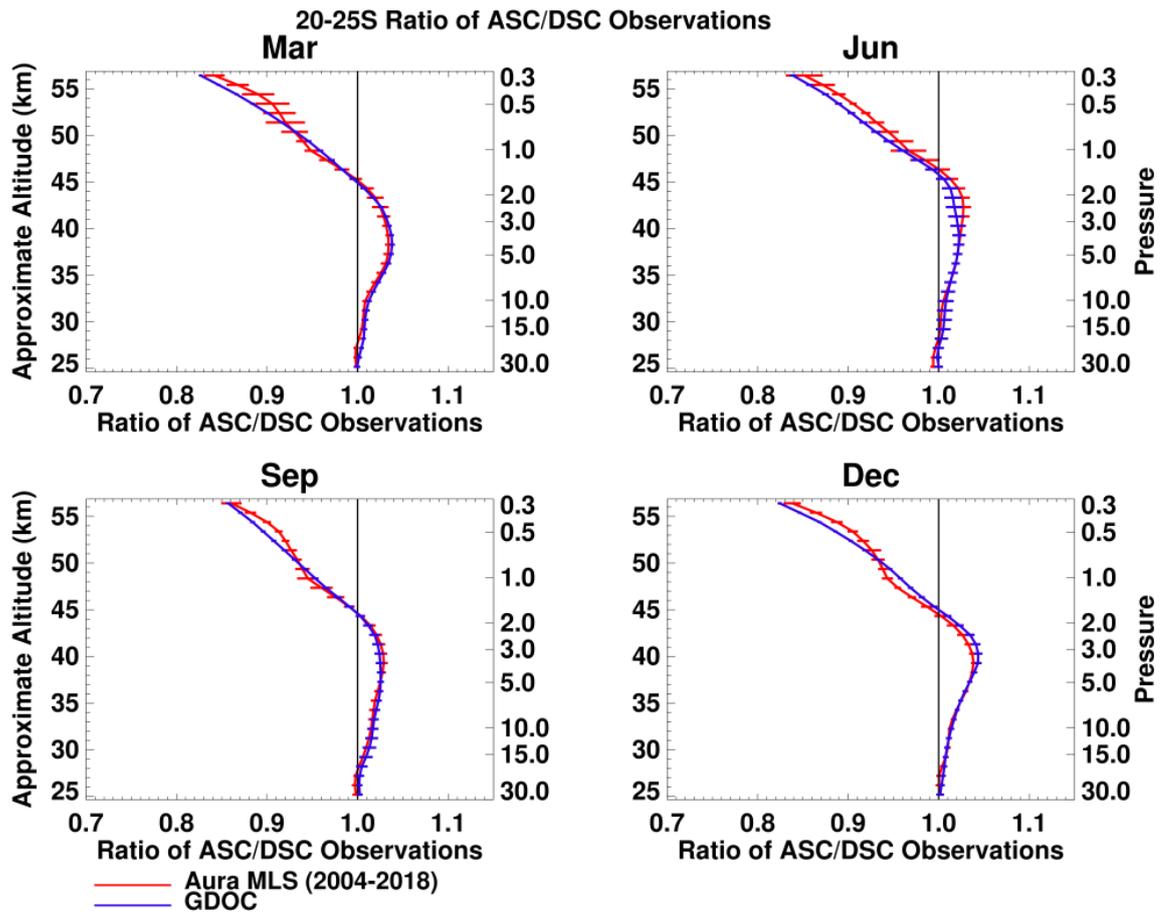


Figure S5. Same as Figure 6 but for 20-25° S.

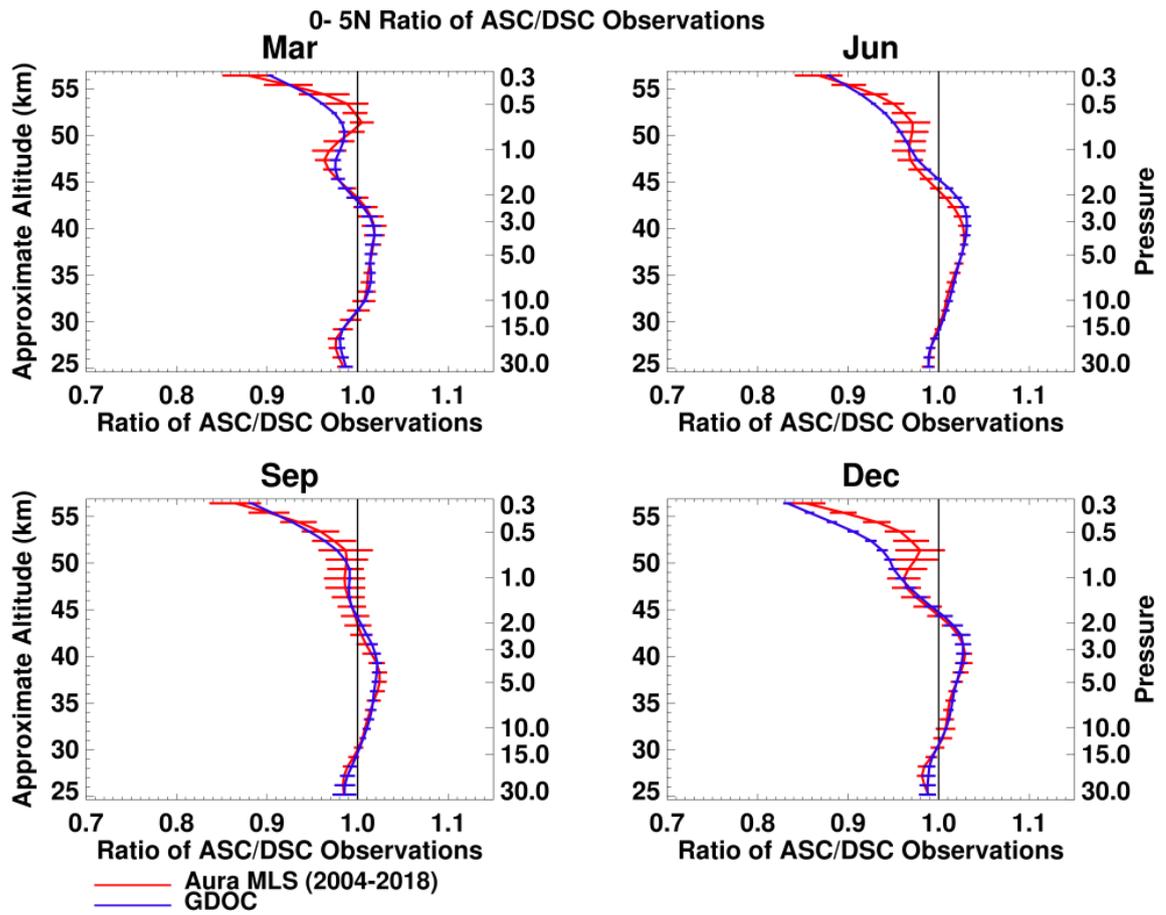


Figure S6. Same as Figure 6 but for 0-5° N.

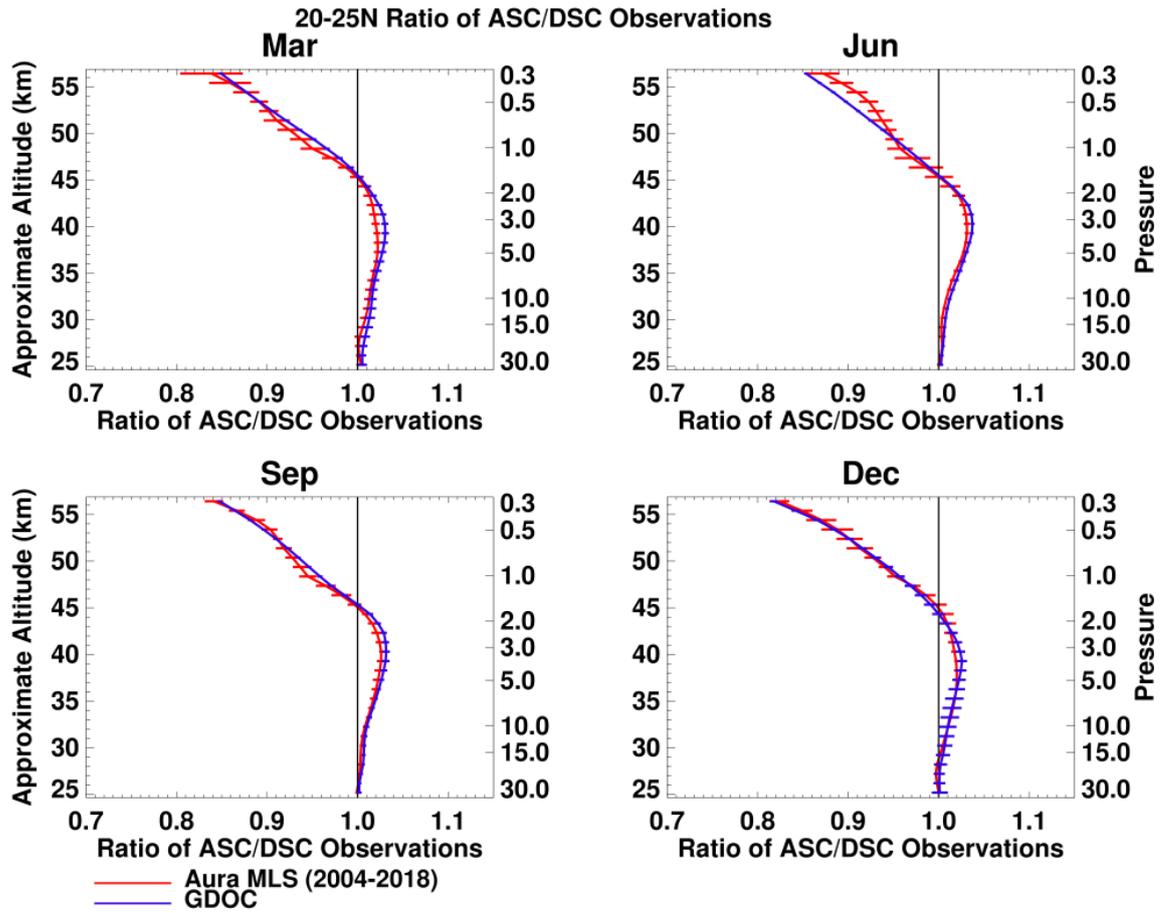


Figure S7. Same as Figure 6 but for 20-25° N.

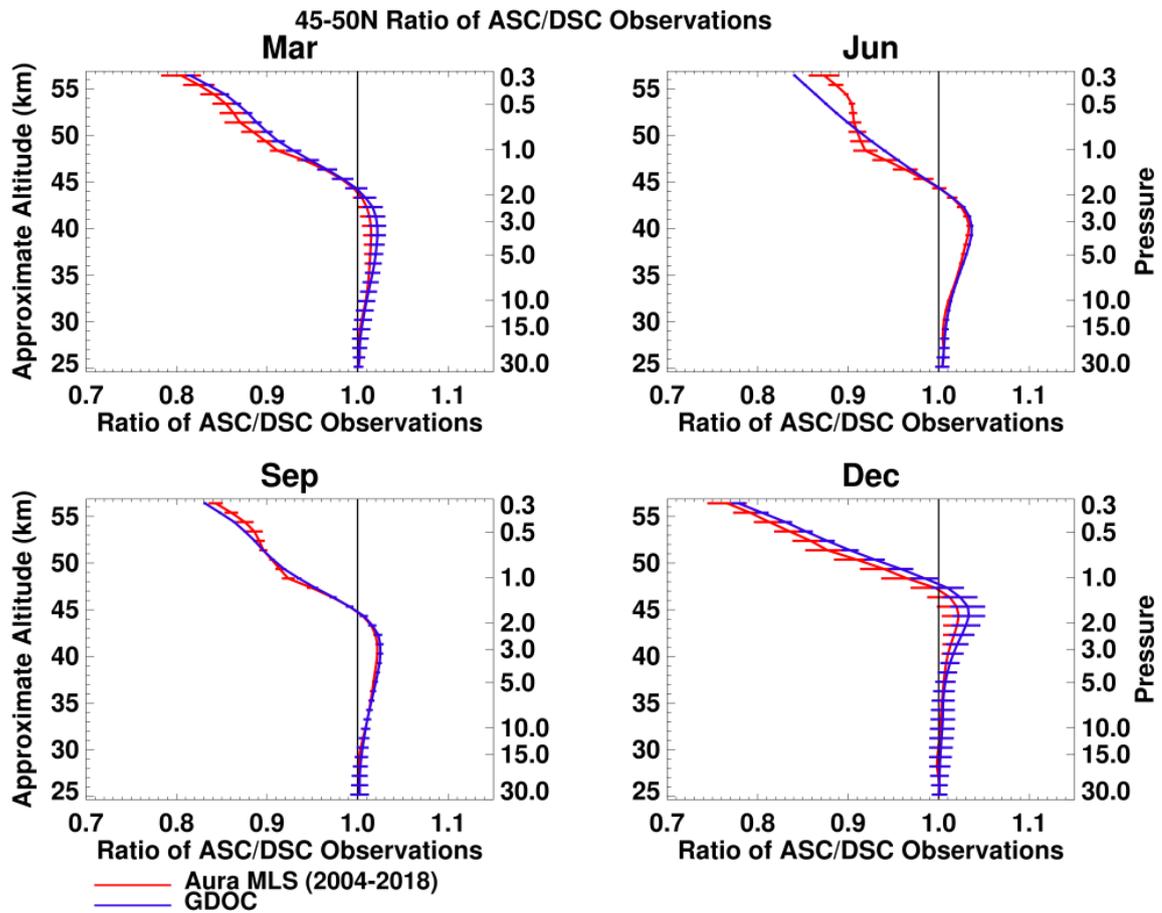
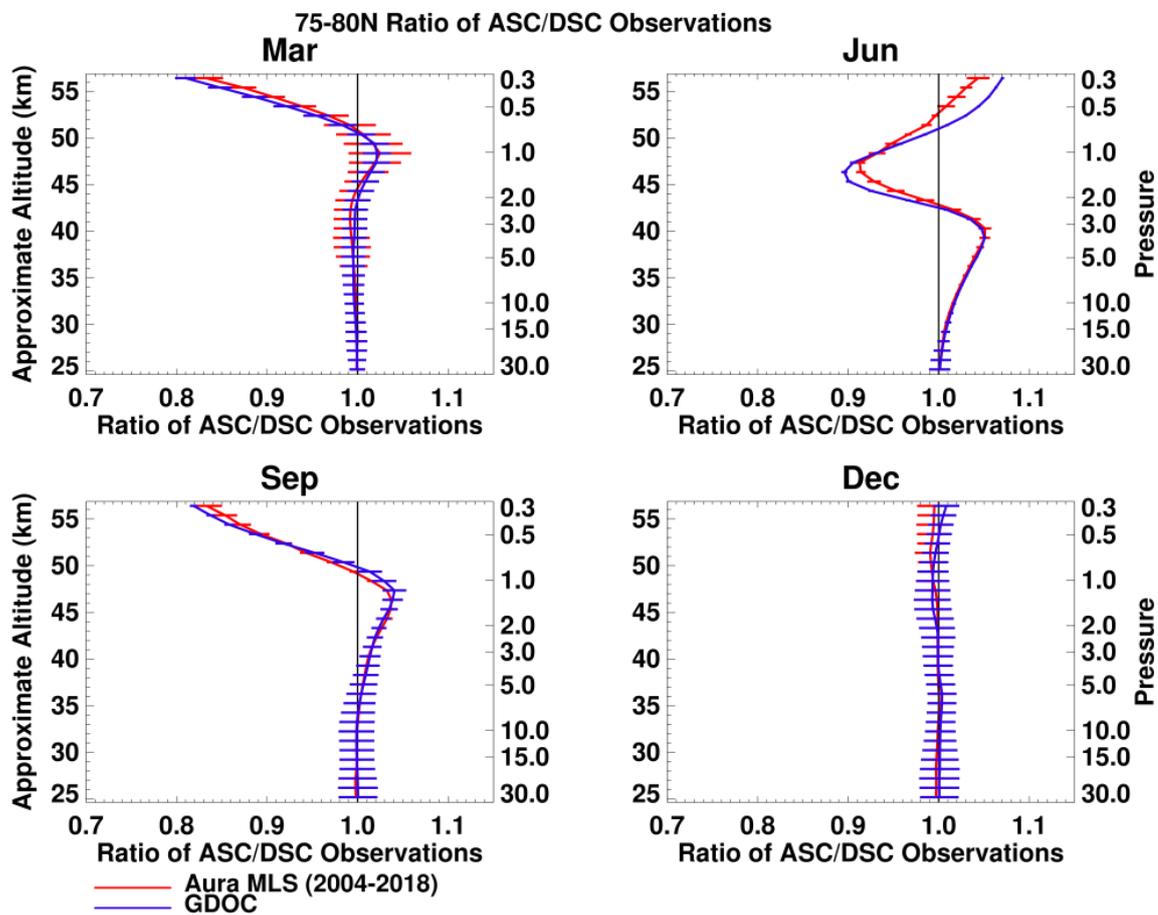
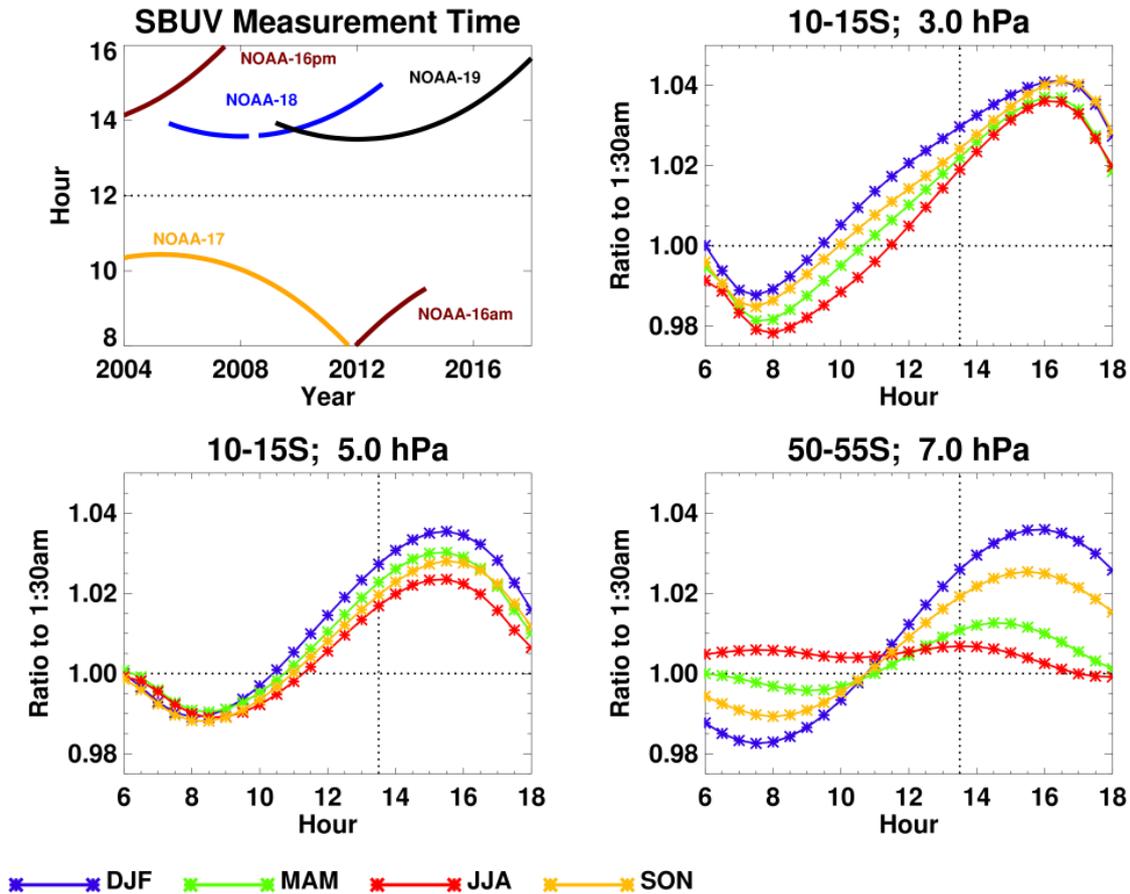


Figure S8. Same as Figure 6 but for 45-50° N.



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Figure S9. Same as Figure 6 but for 75-80° N.



5 Figure S10. GEOS-GMI diurnal climatology (GDOC) as a function of month for the latitude range and pressure level of the SBUV

 time series comparisons relative to Aura MLS in Figures 9 to 11. The first panel shows the slow drift of the SBUV measurement

 local solar time as a result of satellite orbit drift. Measurements are taken primarily in the 2 to 4 pm time range and 8 to 10 am time

 range. In the upper right panel (10-15° S; 3 hPa) there is an offset between the morning and afternoon measurements as well as a

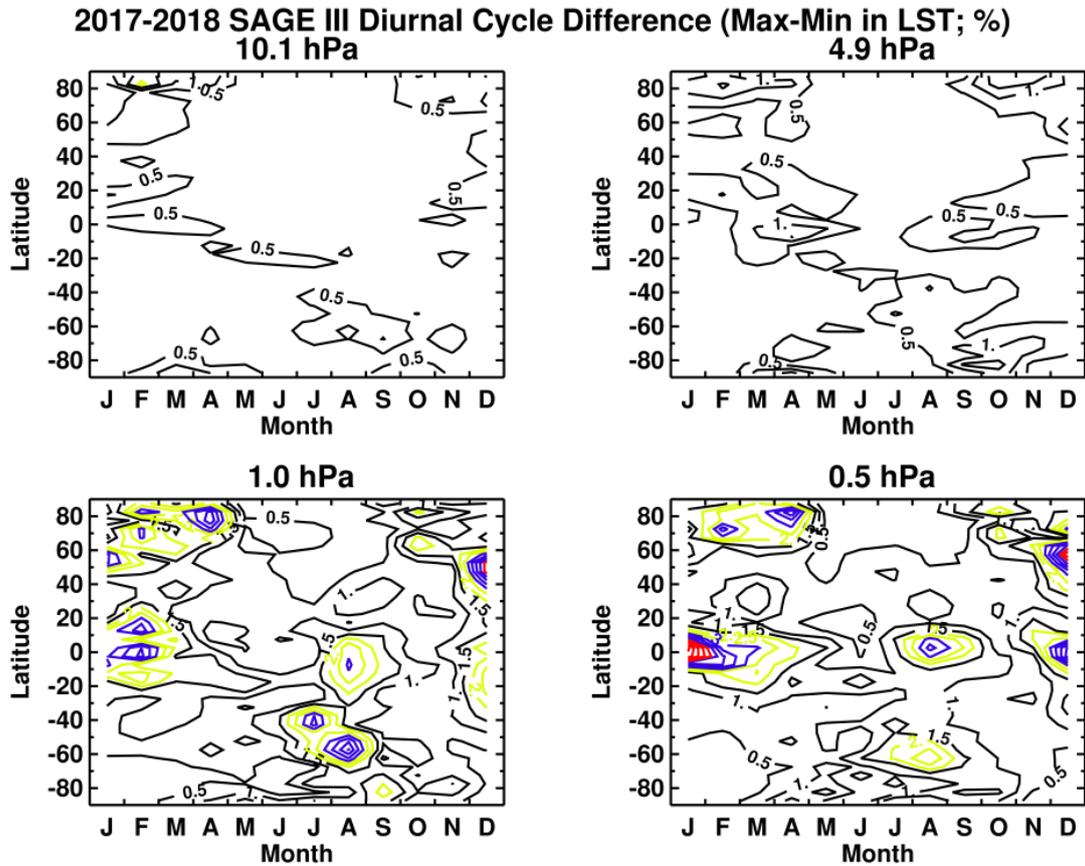
 gradient from 8 to 10am and 2 to 4pm, which translates to a diurnally-induced drift in the SBUV time series. In the lower left panel

 (10-15° S; 5 hPa) there is little variation between 8 to 10am and 2 to 4pm, but only an offset between the morning and afternoon

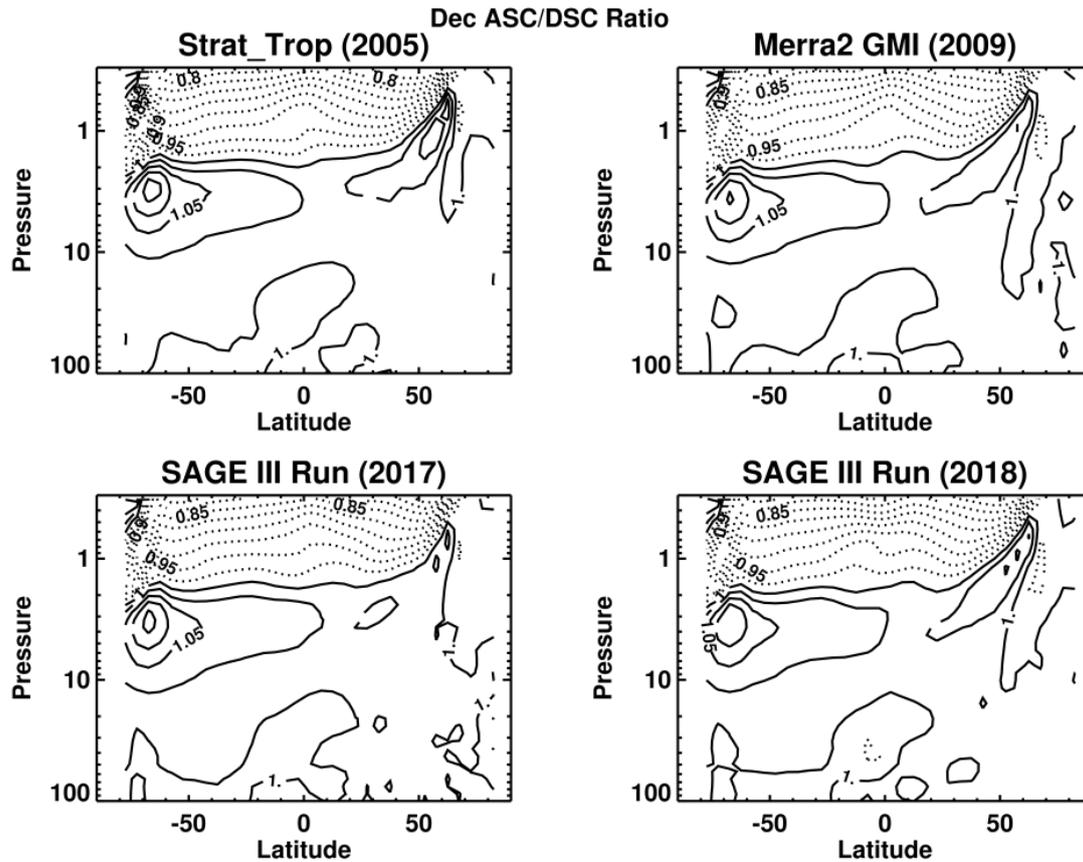
 10 measurements. Finally, in the bottom right panel (50-55° S; 7 hPa) there is little diurnal cycle in winter (JJA) but a strong diurnal

 cycle in summer (DJF), which imparts a diurnally-induced seasonal-scale variation in the SBUV record. Vertical dotted line indicates

 1:30pm Aura MLS measurement time for reference.



5 **Figure S11.** Contour plot of the differences between the diurnal climatology computed from 2017 output and from 2018 output. Differences are expressed as the amplitude of the max-min differences in local solar time and provide information on interannual variability in the model diurnal signal. Contour intervals are 0.5 (5%). Variations of 2% or greater are colored in yellow, 3% or greater in blue, and greater than 5% in red. Below 5 hPa the differences are generally less than 1%. At higher levels there are sporadic instances of larger differences (3-5%) in the tropics (also seen in Figure S10) and at higher latitudes. As more years of model output become available, we will be able to better characterize interannual variability in the model.



5 Figure S12. Contour plot of diurnal climatology as derived from four model realizations. The ratio of the day to night values corresponding to Aura MLS measurement times are shown with a ratio contour interval of .025 (2.5%). The day to night ratio of ozone is similar in all models, despite a number of model modifications and different year simulations. We note that the GEOSCCM simulation used in Parrish et al. [2014] is included here (Strat_Trop 2005) only to demonstrate that model improvements have not drastically changed the model diurnal signal, but that this particular simulation is outdated. The final climatology is the average of outputs from the SAGE III 2017 and 2018 model runs.