

Measurement of NO_x and NO_y with a thermal dissociation cavity ring-down spectrometer (TD-CRDS): instrument characterisation and first deployment

Nils Friedrich¹, Ivan Tadic¹, Jan Schuladen¹, James Brooks², Eoghan Darbyshire², Frank Drewnick³, Horst Fischer¹, Jos Lelieveld¹, and John N. Crowley¹

¹Atmospheric Chemistry Department, Max Planck Institute for Chemistry, Mainz, 55128, Germany ²Centre for Atmospheric Science, University of Manchester, Manchester, M13 9PL, UK ³Particle Chemistry Department, Max Planck Institute for Chemistry, Mainz, 55128, Germany

Correspondence: John N. Crowley (john.crowley@mpic.de)

Received: 12 May 2020 – Discussion started: 25 May 2020 Revised: 17 August 2020 – Accepted: 11 September 2020 – Published: 29 October 2020

Abstract. We present a newly constructed, two-channel thermal dissociation cavity ring-down spectrometer (TD-CRDS) for the measurement of NO_x (NO + NO₂), NO_y (NO_x + $HNO_3 + RO_2NO_2 + 2N_2O_5$ etc.), NO_z ($NO_v - NO_x$) and particulate nitrate (pNit). NO_v-containing trace gases are detected as NO₂ by the CRDS at 405 nm following sampling through inlets at ambient temperature (NO_x) or at $850 \,^{\circ}\text{C}$ (NO_{ν}) . In both cases, O₃ was added to the air sample directly upstream of the cavities to convert NO (either ambient or formed in the 850 °C oven) to NO₂. An activated carbon denuder was used to remove gas-phase components of NO_v when sampling pNit. Detection limits, defined as the 2σ precision for 1 min averaging, are 40 pptv for both NO_x and NO_y . The total measurement uncertainties (at 50 %) relative humidity, RH) in the NO_x and NO_y channels are 11%+10 pptv and 16%+14 pptv for NO₇ respectively. Thermograms of various trace gases of the NO₇ family confirm stoichiometric conversion to NO2 (and/or NO) at the oven temperature and rule out significant interferences from NH₃ detection (< 2%) or radical recombination reactions under ambient conditions. While fulfilling the requirement of high particle transmission (> 80% between 30 and 400 nm) and essentially complete removal of reactive nitrogen under dry conditions (> 99 %), the denuder suffered from NO_x breakthrough and memory effects (i.e. release of stored NO_{ν}) under humid conditions, which may potentially bias measurements of particle nitrate.

Summertime NO_x measurements obtained from a ship sailing through the Red Sea, Indian Ocean and Arabian Gulf (NO_x levels from < 20 pptv to 25 ppbv) were in excellent agreement with those taken by a chemiluminescence detector of NO and NO₂. A data set obtained locally under vastly different conditions (urban location in winter) revealed large diel variations in the NO_z to NO_y ratio which could be attributed to the impact of local emissions by road traffic.

1 Introduction

1.1 Atmospheric NO_x and NO_y

Total reactive nitrogen NO_y (= $NO_x + NO_z$) consists of nitrogen oxide, NO; nitrogen dioxide, NO₂ (NO + $NO_2 = NO_x$); and their reservoir species, NO_z (NO₃ + $2N_2O_5 + HNO_3 + HONO + RONO_2 + RO_2NO_2 + XONO_2 +$ $XNO_2 + pNit$), where X is a halogen atom. HCN and NH₃ are generally not considered to be components of NO_y (Logan, 1983).

The formation of both peroxy nitrates (PNs; RO₂NO₂) and alkyl nitrates (ANs; RONO₂) requires the presence of organic peroxy radicals (RO₂), which are formed by processes such as the reaction of OH radicals with volatile organic compounds (VOCs) and oxygen (Reaction R1). RO₂ radicals subsequently react with NO₂ or NO to form peroxy nitrates (PNs; RO₂NO₂) or alkyl nitrates (ANs; RONO₂, Re-

actions R2 and R3). Reaction (R3) competes with the formation of an alkoxy radical (RO) and the oxidation of NO to NO₂ (Reaction R4), which consumes the dominant fraction of RO₂. The branching ratio between these two pathways depends on atmospheric conditions such as pressure and temperature and on the structure and length of the organic backbone (Lightfoot et al., 1992). HNO₃ is produced mainly via the reaction of NO₂ with OH (Reaction R5).

$$OH + RH + O_2 \rightarrow RO_2 + H_2O \tag{R1}$$

 $RO_2 + NO_2 + M \rightarrow RO_2NO_2 + M$ (R2)

$$RO_2 + NO + M \rightarrow RONO_2 + M$$
 (R3)

$$\text{RO}_2 + \text{NO} \rightarrow \text{RO} + \text{NO}_2$$
 (R4)

$$OH + NO_2 + M \rightarrow HNO_3 + M$$
 (R5)

The lifetimes of peroxy nitrates in the low troposphere are mainly governed by the temperature. PNs with an additional acyl group (PANs), such as peroxyacetyl nitrate (PAN), are generally more stable than PNs without an acyl group (e.g. pernitric acid, HO₂NO₂), which are observed only in cold regions (Slusher et al., 2002). Thus, of the peroxy nitrates, only PANs are considered able to act as transportable reservoirs for NO_x. At higher altitudes in the troposphere (above ca. 7 km) photolysis becomes the most important loss process for PAN, while the reaction with OH is negligible throughout the troposphere (Talukdar et al., 1995).

The absence of photolysis reactions and low levels of the OH radical at night-time provide alternative pathways for the formation of NO_z species. NO_2 is oxidised by O_3 to produce the nitrate radical NO_3 , which exists in thermal equilibrium with N_2O_5 (Reactions R6 and R7). The reaction of NO_3 with hydrocarbons represents a night-time source of alkyl nitrates (Reaction R8), and N_2O_5 can be hydrolysed on aqueous aerosol, resulting in the formation of HNO₃ (Reaction R9) and ClNO₂ (Reaction R10) if particulate chloride is available (Finlayson-Pitts et al., 1989).

 $NO_2 + O_3 \rightarrow NO_3 + O_2 \tag{R6}$

$$NO_3 + NO_2 \rightleftharpoons N_2O_5$$
 (R7)

 $NO_3 + R = R(+O_2) \rightarrow \rightarrow RONO_2$ (R8)

$$N_2O_5 + H_2O \rightarrow 2HNO_3 \tag{R9}$$

$$N_2O_5 + Cl^- \rightarrow ClNO_2 + NO_3^- \tag{R10}$$

Nitric acid formation via the reaction of NO₂ and OH (Reaction R5), followed by wet or dry deposition of HNO₃, is considered to be the dominant daytime loss process for atmospheric NO_x (Roberts, 1990), although the reduction of NO_x may result in an increasingly important role for organic nitrates (e.g. in the USA; Romer Present et al., 2020). As some organic nitrates have longer lifetimes than HNO₃, the atmospheric transport of NO_x to remote locations would lead to a more even distribution of NO_x, instead of hotspots in polluted regions close to emission sources. Atmospheric removal processes for ANs include oxidation by OH or O₃ (which may lead to a loss of the nitrate functionality), deposition to the Earth's surface and photolysis. Additionally, partitioning into the aerosol phase is possible for large and multifunctional ANs (Perring et al., 2013). Alkyl nitrates possessing no further functionality (e.g. double bonds or hydroxyl groups) can be unreactive and have long lifetimes (Talukdar et al., 1997). On the global average, RONO₂ has a lifetime of close to 3 h (2.6–3 h) with \sim 30 % being lost by hydrolysis (Zare et al., 2018).

The formation of NO_z in the lower atmosphere reduces the NO_x lifetime, and the partitioning of NO_y into NO_x and NO_z can provide information about the chemical history of an air mass (Day et al., 2002; Wild et al., 2014). In regions impacted by biogenic emissions, the sources and sinks of ANs account for a large fraction of NO_x lost both during the day and night and, thus, control the lifetime of NO_x (Romer et al., 2016; Sobanski et al., 2017).

Laboratory experiments have shown that particulate nitrates (pNits) are formed at high yields in the atmospheric degradation of terpenoids in the presence of NO_x and play an important role in the formation and growth of secondary organic aerosol (SOA; Ng et al., 2017; Ammann et al., 2019). This has been confirmed in field studies, which provide evidence for the partitioning of organic nitrate to the aerosol phase both during day- and night-time (Rollins et al., 2012; Fry et al., 2013; Palm et al., 2017) with the formation of highly functionalised molecules and large contributions (up to 25 %) of particulate organic nitrates to the total aerosol mass (Xu et al., 2015; Lee et al., 2016; Huang et al., 2019).

1.2 Detection of NO_x

Methods for the detection of NO and NO2 include chemiluminescence (CLD), differential optical absorption spectroscopy (DOAS), laser-induced fluorescence (LIF) and cavity ring-down spectroscopy (CRDS). A description and intercomparison of these methods is given in Fuchs et al. (2010), and we restrict the following discussion to an outline of the basic principles. The CLD method detects NO by chemiluminescent emission in its reaction with O3; detection of ambient NO₂ by CLD follows its catalytic or photolytic conversion to NO. The best CLD devices have detection limits for NO and NO₂ in the single-digit parts per trillion by volume (pptv) range (Hosaynali Beygi et al., 2011; Reed et al., 2016; Tadic et al., 2020). Detection of NO₂ via LIF involves photoexcitation in its visible absorption band at wavelengths > 400 nm and detection of fluorescent emission at wavelengths > 600 nm, with detection limits of the order of parts per trillion by volume achieved for an integration time of a few seconds (Day et al., 2002; Javed et al., 2019). The structured spectrum of NO₂ between ≈ 400 and 600 nm is used to detect light absorption by ambient NO₂ by DOAS, using either broadband light sources (long-path DOAS, with a path length of more than a few kilometres) or natural sunlight (Platt et al., 1979; Leser et al., 2003; Pohler et al., 2010; Merten et al., 2011).

The CRDS detection method for NO₂ also utilises its visible absorption spectrum, with high sensitivity being reached by achieving very long path lengths for optical extinction in an optical resonator (see Sect. 2.1). Limits of detection for NO₂ with CRDS of < 20 pptv with a 1 s integration time have been reported (Wild et al., 2014). NO can be detected as NO₂ following its oxidation by O₃ (Reaction R6; Fuchs et al., 2009).

1.3 Detection of NO_y

The first NO_{ν} measurements were based on the conversion of all reactive nitrogen trace gases (apart from NO) to NO on catalytic metal surfaces of gold at $\sim 300\text{--}320\,^\circ\text{C}$ or of molybdenum oxide (MoO) at \sim 350–400 °C (Fahey et al., 1985; Williams et al., 1998), with subsequent CLD detection of NO. Au converters were designed to exclude particulate nitrates, whereas MoO set-ups aimed at a response towards pNit (Williams et al., 1998). In recent years, the thermal decomposition of NO₇ to NO₂ has been employed to detect NO_{z} using inlets held at temperatures high enough (> 650-700 °C) to thermally dissociate the most strongly bound reactive nitrogen trace gas, HNO₃, to NO₂ (Day et al., 2002; Rosen et al., 2004; Wooldridge et al., 2010; Perring et al., 2013; Wild et al., 2014) and/or using multiple inlets at intermediate temperatures (Paul et al., 2009; Paul and Osthoff, 2010; Sadanaga et al., 2016; Sobanski et al., 2016; Thieser et al., 2016). Following thermal decomposition, the NO₂ product can be detected using LIF (Day et al., 2002, 2003; Rosen et al., 2004; Murphy et al., 2006; Wooldridge et al., 2010) or cavity-enhanced absorption spectroscopy (Paul et al., 2009; Wild et al., 2014; Sadanaga et al., 2016; Sobanski et al., 2016; Thieser et al., 2016). These techniques are impacted to various degrees by secondary reactions at high temperatures, including the loss of NO₂ via recombination with α -carbonyl peroxy radicals or reaction with $O(^{3}P)$ atoms (formed by the thermolysis of ambient O_3) and the generation of extra NO_2 from the oxidation of NO via reactions with peroxy radicals (Day et al., 2002; Sobanski et al., 2016; Thieser et al., 2016; Womack et al., 2017). Measures to reduce potential measurement artefacts and avoid excessive data correction include operation at low pressures (Day et al., 2002; Womack et al., 2017) and the addition of surfaces to scavenge peroxy radicals (Sobanski et al., 2016). Nonetheless, data correction may still be necessary and may involve laboratory characterisation and chemical simulation of the chemical reactions within the heated inlet (Sobanski et al., 2016; Thieser et al., 2016).

In this paper we present a two-channel TD-CRDS instrument for the detection of NO_x , NO_y , NO_z and pNit that overcomes these limitations. Compared with the set-ups described by Thieser et al. (2016), the following changes were implemented: (1) the addition of O_3 for NO_x detection, (2) use of an oven that was located directly at the front of the inlet and which was operated at a higher temperature (to detect HNO₃), and (3) the use of a charcoal denuder for separate measurement of pNit and gas-phase NO_z. The addition of O₃ (after the TD inlet) ensures that we detect NO as well as NO₂ and, thus, removes bias caused by factors such as the pyrolysis of O₃ and reactions of O(³P) which reduce NO₂ to NO.

2 Experimental

Our TD-CRDS instrument consists of two identically constructed cavities to monitor NO₂ at 405 nm which are largely unchanged compared to those described by Thieser et al. (2016). In the present set-up, the two cavities are connected to three different inlets. One cavity monitors NO_x via an inlet at ambient temperature, and the second samples air via one of two heated inlets (one equipped with a denuder, see below), thereby monitoring either NO_y or particle nitrate. A schematic diagram (not to scale) of the instrument is given in Fig. 1.

2.1 CRDS operation principals

The optical resonator consists of two mirrors (1 m radius of curvature) with a nominal 0.999965 reflectivity at 405 nm (Advanced Thin Films), which are mounted 70 cm apart. The cavity volumes are defined by Teflon (FEP) coated DURAN glass tubes with an inner diameter of 10 mm.

Under normal operating conditions each cavity samples $3.0 \,\mathrm{L}\,(\mathrm{STP})\,\mathrm{min}^{-1}$ (slpm; standard litres per minute) of ambient air (where STP refers to 0 °C and 1013 hPa). Additional purge flows (0.14 slpm dry synthetic air) are introduced directly in front of each mirror to prevent surface degradation by atmospheric trace gases. The cavities are operated at pressures of 540 to 580 Torr (1 Torr = 1.333 hPa), resulting in a residence time of ~ 1.2 s. A 405 nm laser light (square-wave modulated at 1666 Hz and a 50% duty cycle) is provided by a laser diode (LASER COMPONENTS), the emission of which is coupled into an optical fibre with a Y piece for splitting into both cavities. Temperature and current control of the laser diode are achieved by a Thorlabs ITC502 control unit. The laser emission spectrum is monitored continuously by coupling scattered light from one of the cavity mirrors into a 3648 pixel CCD (charge-coupled device) spectrograph (OMT, ~ 0.1 nm resolution).

The intensity of light exiting the cavity is measured with a photomultiplier (Hamamatsu Photonics), with ring-up and ring-down profiles recorded by a digital oscilloscope (Pico-Scope 3000). NO₂ mixing ratios are derived from the decay constant (k or k_0) describing the exponential decrease in light intensity after the laser has been switched off:

$$[NO_2] = \frac{l}{d} \cdot \frac{1}{\sigma c} \left(k - k_0 \right), \tag{1}$$



Figure 1. Schematic diagram of the TD-CRDS instrument (not to scale). NO_y and pNit are detected via the heated channel, and NO_x is detected via the ambient channel. Ozone is generated via a Pen-Ray lamp (185 nm) and serves to convert NO to NO_2 . TD denotes thermal dissociation, and FC denotes flow controller. The flows listed are those used under normal operating conditions. p denotes a pressure sensor, Ex denotes a membrane pump and exhaust, Amb denotes ambient air and V denotes an electronically switchable PTFE valve. "Filter" is a PTFE filter, 2 µm pore size. O-air refers to zero air. The inset (photo) shows the honeycomb structure of the activated carbon denuder. The critical orifices have diameters of ≈ 0.05 mm. ("slm" is used to refer to standard litres per minute in this figure.)

where *c* is the speed of light, σ is the effective absorption cross section of NO₂ over the emission spectrum of the laser (Vandaele et al., 2002), and *k* and *k*₀ are the decay constant with and without NO₂ present in the cavity respectively. Thus, *k*₀ is defined by the mirror reflectivity and light scattering by the dry, synthetic air.

The ratio l/d accounts for the difference between the physical length of the cavity (l) and the effective optical path length (d) in which NO₂ is present as well as for dilution effects. d is shorter than l due to the purge flows of zero air in front of the mirrors, and a value of $l/d = 1.02 \pm 0.01$ was determined by adding a constant flow of NO₂ and varying the purge-gas flow rate (Schuster et al., 2009; Thieser et al., 2016). k_0 is typically determined every 5 min (for 1 min) by overflowing the inlets with zero air from a commercial zero air generator (CAP 180, Fuhr GmbH) attached to a source of compressed ambient air. PTFE filters (47 mm diameter, 2 µm pore size) prevent particles from entering the cavities. The filter's efficiency, tested with laboratory air containing 1.8×10^3 particles cm⁻³ and a CPC (TSI 3025 A), was > 98 %.

Raw data sets (i.e. ring-down constants) undergo a few basic corrections before further analysis:

- 1. k_0 is interpolated onto the *k* time grid. The first three data points after switching from sampling to zeroing are discarded in order to enable the stabilisation of the zero signal. The remaining data points of each zero cycle are averaged. Finally, a linear interpolation between the averaged k_0 values is performed, allowing for the subtraction of k_0 for each individual data point.
- 2. Depending on conditions of flow, the pressure and the inlet set-up (see Sect. 2.2 and 2.3), changes in flow resistance between the zeroing and sampling periods result in slight changes in the cavity pressure. The resulting change in Rayleigh scattering of the 405 nm light owing to a pressure change of 6.5 Torr was found to be equivalent to a change of ca. 300 pptv in the NO₂ mixing ratio,

which is in accordance with earlier experiments using previous versions of this instrumental set-up (Thieser et al., 2016). We have also used an alternative set-up in which the inlet is overflowed with zero air added close to the tip of the inlet (downstream of the oven); this reduces the pressure difference but has the disadvantage that hot air is blown out of the instrument when zeroing, which may interfere with co-located inlets. The addition of zero air upstream of the quartz inlets would remove this problem, but it would also increase the complexity of the inlet and potentially result in the loss of sticky molecules such as HNO₃.

3. A further correction is associated with the difference in the Rayleigh scattering coefficient between dry air (during zeroing) and humid air (whilst taking ambient measurements). This effect was corrected using the H_2O scattering cross sections reported by Thieser et al. (2016), leading, for example, to a correction of 116 pptv at 70 % RH and 20 °C.

2.2 Detection of NO_x

In order to measure NO (which does not absorb at 405 nm), it is converted to NO₂ by reaction with an excess of ozone (O_3) ; the O₃ was generated by passing zero air over a Hg Pen-Ray lamp emitting at 185 nm, which was housed inside a glass vessel at ca. 980 Torr pressure. The gas stream containing O₃ is split up equally by critical orifices and directed into two identical reaction volumes made of 88 cm long PFA tubing (1/2 inch outer diameter, residence time 1.05 s). The concentration of O_3 (monitored by a commercial monitor, Model 202, 2B Technologies) was optimised in laboratory experiments in which the efficiency of the conversion of NO to NO₂ was varied by changing the flow of air over the Pen-Ray lamp. The maximum concentration of NO₂ (corresponding to 96% of the NO in the gas bottle) was observed when the flow over the Pen-Ray lamp was between 60 and 80 cm³ (STP) min⁻¹ (hereafter sccm), which resulted in 19 ppmv O₃ in the reaction volumes. This result could be confirmed by numerical simulation (see Table S1 in the Supplement) of the reactions involved in the formation and loss of NO₂ when NO reacts with O₃. According to the simulation, the maximum conversion of NO to NO₂ during the 1.05 s residence time occurs between ca. 12 and 20 ppm VO_3 . The conversion efficiency decreases at higher O3 concentrations due to the formation of N₂O₅ and NO₃. The results from the experiments to determine the optimum parameters for O_3 generation are summarised in Fig. S1.

For NO₂, the performance of the instrument was first described by Thieser et al. (2016), who reported a measurement uncertainty of $6\% + (20 \text{ pptv} \times \text{RH}/100)$ that was dominated by uncertainty in the effective cross section of NO₂ and the wavelength stability of the laser diode. The NO_x detection limit of 40 pptv (2σ , 1 min average) for the present instrument (laboratory conditions) was derived from an Allan variance analysis and is worse than that reported by Thieser et al. (2016) (6 pptv at 40 s) due to degradation of the mirror reflectivity. Corrections applied to take humidity and pressure changes into account are discussed in Sect. 2.1. The total uncertainty in NO_y will depend on the uncertainty in the conversion of both gaseous and particulate nitrate to NO_x and, thus, depends on the individual components of NO_y in the air sampled. For purely gaseous NO_y, the major problem is likely to be related to the loss of sticky molecules at the inlet, and we choose to quote a "worst case" uncertainty of 15 %.

2.3 Thermal dissociation inlets: detection of NO_y

The thermal dissociation inlets used to dissociate NO_{ν} to NO₂ are quartz tubes housed in commercial furnaces (Carbolite, MTF 10/15/130). The oven temperature was regulated with a custom-made electronic module, which enabled spatial separation between the heating elements and insulation and the control electronics. The distance between the heated section of the quartz tubes and the point at which air was taken into the inlet was kept short (ca. 30 cm) in order to minimise losses of trace gases with a high affinity for surfaces, especially HNO₃ (Neuman et al., 1999). Experiments characterising the thermal conversion of various trace gases to NO₂ are described in Sect. 3.1. An electronic, PTFE threeway valve (Neptune Research, Inc., type 648T032, orifice diameter 4 mm) under software control switches between the two heated inlets, one of which is equipped with a denuder. Memory effects on the valves' surfaces were not observed for NO₂. Bypassing the valve under normal sampling conditions led to a 0.6 Torr pressure change. The sampling flow through both heated inlets is 3.0 slpm. When sampling ambient air via the denuder, we expect to remove all gas-phase NO_{y} components and, thus, measure only particulate nitrate (pNit). Experiments to characterise the transmission of the denuder for particles and various trace gases are presented in Sect. 3.3.

2.4 Active carbon denuder

The active carbon denuder (Dynamic AQS) has a honeycomb structure with 225 quadratic channels (1 mm × 1 mm) of 10 cm length in a cylindrical form (diameter 3 cm) which is housed inside an aluminium casing with 1/2 inch connections (see Fig. 1). The geometric surface area of the denuder is ~45 cm². Assuming a specific surface area for activated carbon of 1000 m² g⁻¹ (Atsuko et al., 1996), we calculate a BET (Brunauer–Emmett–Teller, Brunauer et al., 1938) surface area of the order of 10^8 cm².

2.5 Chemicals

A stock, liquid sample of PAN in *n*-tridecane (> 98%, Alfa Aesar) was synthesised according to the procedures described by Gaffney et al. (1984) and Talukdar et al. (1995).

Samples of lower concentration (as used in the experiments described below) were produced by diluting the original sample with additional *n*-tridecane. Acetone (> 99 %), isopropyl nitrate (> 98.0 %) and (R)-(+)-limonene (97 %) were obtained from Sigma-Aldrich. An ammonia permeation source (324 ng min⁻¹) was supplied by VICI Metronics. Methanol (> 99.9 %) was acquired from Merck, isoprene (98 %, stabilised) from Acros Organics, and ethanol from Martin and Werner Mundo oHG. Both nitric acid (65 %) and β -pinene (pure) were obtained from Carl Roth. N₂O₅ crystals were synthesised according to Davidson et al. (1978) by reacting NO (5 %) with excess O₃ in a glass reactor. O₃ was produced via electrical discharge through O₂ using a commercial ozone generator (Ozomat Com, Anseros). The crystals were trapped and stored at -78 °C in dry ice and ethanol.

3 Results and discussion

3.1 Trace gas thermograms

The fractional conversion of NO_z to NO_2 in the TD inlets was investigated in a series of experiments in which constant flows of various NO_z trace gases were passed through the heated inlet (bypassing the denuder) while the temperature was varied and NO₂ was monitored. NO_x impurity levels were determined either via the simultaneous operation of the NO_x channel or via the NO_y channel mixing ratio at room temperature, before and after heating the inlet. By inserting a thermocouple into the middle part of the heated section under normal sampling conditions, we were able to show that the temperature of the gas was ≈ 80 °C lower than that indicated by the oven's internal temperature sensor in the 200-300 °C temperature range and about 40 °C lower at a set temperature of 600 °C (see Fig. S3a). We were unable to measure the temperatures of the gas stream at oven temperatures above about 600 °C, and we refer only to the temperature indicated by the internal sensor of the oven throughout the paper.

In the following, we show that the thermograms (plots of fractional dissociation of NO_z to NO_x versus temperature) which we measure with this instrument are broader and are shifted in temperature compared with other examples found in the literature, including those from this laboratory. As this instrument is built for the measurement of NO_y and is not intended for separate measurement of species such as PNs, ANs and HNO₃, overlap of the individual thermograms does not represent a problem.

3.1.1 PAN

A stream of 200 sccm synthetic air was used to elute a constant supply of gaseous PAN from its solution (held at a constant temperature of 0 °C in a glass vessel) into the CRDS inlet. The thermogram is presented in Fig. 2, and the absolute NO₂ mixing ratios are depicted in Fig. S2a. In this experiment the maximum amount of NO₂ observed (at tem-



Figure 2. Thermograms of the NO_z species PAN, iPN (isopropyl nitrate), HNO₃, NH₄NO₃, N₂O₅, HONO, ClNO₂ and the potential interference from NH₃ (without added O₃). The NH₃ fractional conversion is calculated relative to the calibrated output of the permeation source, whereas all other species are calculated relative to the observed mixing ratio at maximum conversion. Error bars are derived from the standard deviations during the averaging intervals. At the set temperature of 850 °C, PAN, iPN, NH₄NO₃, N₂O₅ (x2), HONO and HNO₃ are converted quantitatively to NO₂, while the NH₃ interference is negligible under typical ambient conditions.

peratures > 400 °C) was 2.2 ppbv. At temperatures < 100 °C, there was no measurable thermal decomposition of PAN to NO₂. Increasing the temperature from 100 °C to 300 °C resulted in a sharp increase in NO₂ which flattened off at temperatures > 380 °C. We conclude that PAN is stoichiometrically converted to NO₂ at temperatures above 400 °C in our oven. The steepest part of the isotherm at ~ 200 °C, i.e. 50 % conversion of PAN to NO₂, is therefore shifted by ca. 80 °C compared with those reported in the literature by Wild et al. (2014), Thieser et al. (2016) and Sobanski et al. (2016). This is a consistent feature of our TD ovens and is related to the short time available for thermal decomposition (see below) and a significantly lower gas temperature than indicated by the oven's internal temperature sensor.

3.1.2 Isopropyl nitrate

A 10L stainless-steel canister containing 10.3 ppmv of isopropyl nitrate (iPN) at a pressure of 4 bar N₂ was prepared using a freshly vacuum-distilled liquid sample utilising standard manometric methods. NO_x impurities were ~ 4.7 ppbv, although we note that diluted iPN stored in stainless-steel canisters for periods of several weeks degrades to form NO₂ and HNO₃. The thermogram is displayed in Fig. 2, and the absolute concentrations are shown in Fig. S2b. Based on the mixing ratio of iPN in the canister and the dilution flows, 10.7 ppbv represents (101 ± 11) % conversion. The shaded area around the expected iPN mixing ratio in Fig. S2b signifies the uncertainty of this value based on the propagation of the errors during the manometric and dilution procedures (2% for flow rates, 5% for pressures measured with digital pressure gauges and 10% for the last dilution step using the analogue pressure gauge of the canister).

Between 550 and 850 °C, we observe a weak increase in NO₂ from 10.7 to 11.2 ppbv, which is likely due to small amounts of HNO₃ in the sample. For iPN, the temperature at 50 % conversion is 50 °C higher than those reported by Thieser et al. (2016) and Sobanski et al. (2016). Wild et al. (2014) employed a gaseous mixture of different alkyl nitrates and also observed an initial increase in NO₂ (up to 80 % conversion) for temperatures < 300 °C, followed by a slower increase up to 800 °C. The alkyl nitrates thermogram of Wild et al. (2014) has been included into Fig. S2b to illustrate this behaviour and to facilitate direct comparison.

3.1.3 HNO₃

A custom-made permeation source was used to provide a constant, known flow of HNO₃ (with $\sim 8\%$ NO_x impurity) to the TD-CRDS inlet. The permeation source consisted of a length ($\approx 1 \text{ m}$) of PFA tubing immersed in 66 % HNO₃ solution held at 50 °C through which 100 sccm of dry, zero air was passed. The concentration of HNO₃ and, thus, its permeation rate, $(1.62 \pm 0.2) \times 10^{-4}$ sccm, was derived by measuring the optical extinction of HNO3 at 185 nm using the absorption cross section of Dulitz et al. (2018). The uncertainty is related to the uncertainty in the absorption cross section and the reproducibility of the output. The HNO3 thermogram (Figs. 2 and S2c) has a plateau at temperatures above ≈ 800 °C. In the plateau region of Fig. 2, the HNO₃ mixing ratio measured is 13.0 ± 0.8 ppb, which (within combined uncertainties) is in agreement with the expected value $(15.2 \pm 1.98 \text{ ppbv})$ calculated from the permeation rate and uncertainty in the dilution factor. We cannot rule out some loss of HNO₃ in the tubing connecting the permeation source to the TD-CRDS, although previous studies have shown that irreversible losses are $\sim 5\%$ or less under dry conditions (Neuman et al., 1999). We note that inlet loss of HNO₃ is minimised under ambient sampling conditions, as only a short section (~ 20 cm) of the quartz tubing at ambient temperature is upstream of the heated section in which HNO₃ is converted to NO2. Therefore, our observations are in accord with previous studies that found the complete conversion of HNO₃ to NO₂ in similar set-ups (Day et al., 2002; Di Carlo et al., 2013; Wild et al., 2014; Womack et al., 2017).

3.1.4 N₂O₅

A sample of N₂O₅ was prepared by flowing 80 sccm of synthetic air over N₂O₅ crystals, kept at -70 °C, with further dilution with 20 slpm synthetic air. An 8.5 cm long nylon tube was used to reduce HNO3 impurity. Two distinct dissociation steps can be observed in Fig. 2. The first step, between 50 and 185 °C (in which NO₂ increases to 4.2 ppbv, see Fig. S2d), is due to the dissociation of N_2O_5 to $NO_2 + NO_3$. In the second step, in which NO₃ is dissociated to NO_x between 450 and 800 °C, the NO₂ mixing ratio was 9.2 ppbv. As the amount of N₂O₅ derived from the first dissociation step was in accord with simultaneous measurements of N₂O₅ using a further TD-CRDS set-up (Sobanski et al., 2016), we conclude that some HNO₃ was present in the sample, which was presumably the result of N₂O₅ hydrolysis. Our thermogram is similar to that reported by Womack et al. (2017), who also observed two steps - the first with a plateau at $T > 100 \,^{\circ}\text{C}$ and the second at $T > 650 \,^{\circ}\text{C}$. The shift in temperature (100-150 °C) compared with our results is rationalised in Sect. 3.1.8.

3.1.5 HONO

Gaseous HONO was produced by flowing HCl in air (22 ppbv; relative humidity, RH, ca. 50 %), over a bed of continuously stirred sodium nitrate crystals (Wollenhaupt et al., 2000). In our set-up, the thermal dissociation of HONO to NO starts at $\sim 400 \,^{\circ}$ C and reaches a plateau (6.2 ppbv) between ca. 800 and 850 °C (Fig. 2). We did not have access to independent instrumentation to characterise the concentrations of HONO and potential impurities generated using this method. Previous investigations have reported that HONO thermally dissociates between 450 and 650 °C (Perez et al., 2007) and between 200 and 700 °C (Wild et al., 2014). The reasons for such large divergence in the positions and widths of the thermograms may be partially related to the presence of impurities in the HONO samples used, although the details of the ovens used to thermally dissociated HONO also play an important role as described in see Sect. 3.1.8.

3.1.6 CINO₂

ClNO₂ was generated by passing Cl₂ (33 ppbv in air) over sodium nitrate at room temperature. The thermogram, depicted in Fig. 2, has two steps – one with an apparent plateau at ~ 500 °C and a second at ~ 800 °C. The lower temperature plateau in which ClNO₂ dissociates to NO₂ corresponds to that reported previously (Thaler et al., 2011; Sobanski et al., 2016; Thieser et al., 2016). The observation of further NO₂ formation at higher temperature is consistent with the observations of Wild et al. (2014). We hypothesise, that the second dissociation step might be associated with the presence of ClNO which dissociates to NO (and would therefore not have been detected by instruments that monitor NO₂ rather than NO_x). Even at temperatures > 850 °C, we still see an increase in the NO₂ signal. However, as ClNO is not considered to be an important atmospheric trace gas, this has no repercussions for the deployment of the instrument.

3.1.7 NH₄NO₃ and NaNO₃ particles

NH₄NO₃ and NaNO₃ particles were generated from an aqueous solution (ca. 1 g in 500 mL deionised water) using an atomiser (TSI 3076). The particles were dried prior to size selection (DMA, TSI 3080) and were diluted in a total flow of 6 slpm synthetic air which was split between a condensation particle counter (CPC) and the heated TD-CRDS inlet (after further dilution). The relative thermogram for NH4NO3 is displayed in Fig. 2 and, similar to HNO₃, displays a plateau region at temperatures above 830 °C. The shift in the thermogram when comparing HNO₃ and NH₄NO₃ (which we expect to detect in a two-step process in which NH₄NO₃ first decomposes to HNO₃) may be related to the time required to fully thermally decompose particles (e.g. of 200 nm diameter) containing several million molecules. Particle numbers (in cm^{-3}) detected by the CPC were converted to molar mixing ratios via the diameters and densities of the dry particles $(1.72 \text{ g cm}^{-3} \text{ for NH}_4\text{NO}_3 \text{ and } 2.26 \text{ g cm}^{-3} \text{ for NaNO}_3)$. The fraction of NH_4NO_3 detected as NO_x following passage through the oven is illustrated in Fig. S5, which indicates values between $\approx 60\%$ and 120% depending on particle size. The total uncertainty in the concentration was estimated as 41 %, which includes 10 % uncertainty in the particle diameter (based on measured size distributions of latex calibration particles), 20% uncertainty in the particle number (including the error in the multiple charge correction) and 10 % uncertainty in the density, due to possible differences between single crystal and bulk density. As the particle mass scales to the third order with the particle diameter, the correction for double-charged particles introduces a large uncertainty in the calculated mixing ratios, with the effect being largest in the size range between 100 and 150 nm, which probably explains the lower NH₄NO₃ detection efficiencies in this range. We consider the data obtained at 200 nm to be the most reliable and conclude that, similar to other TD instruments (Womack et al., 2017; Garner et al., 2020), our instrument also detects NH_4NO_3 particles as NO_2 with close to 100 % efficiency. In contrast, our experiments using NaNO3 resulted in much smaller NO_x concentrations despite identical experimental conditions in back-to-back experiments and produced detection efficiencies of close to 25 %. While the inefficient detection of NaNO₃ is consistent with a previous reports suggesting that NaNO₃ would not be detected in TD inlets (Womack et al., 2017), it contrasts strongly with the very recent result of Garner et al. (2020), who observed quantitative conversion of NaNO₃ to NO₂ at 600 °C. This difference may be related to residence times in the heated section of the inlet.

3.1.8 Summary of thermograms

The thermograms obtained by the present instrument deviate from others reported in the literature, with the temperatures required for 50 % dissociation generally being higher by, for example, 80 °C for PAN, 50 °C for iPN and 150 °C for HNO₃ respectively (Day et al., 2002; Wild et al., 2014; Sobanski et al., 2016; Thieser et al., 2016; Womack et al., 2017). This lack of agreement with other set-ups is not unexpected, as the degree of dissociation of a trace gas at any temperature depends not only on the temperature but also on the time over which the molecule is exposed to that temperature (Womack et al., 2017). To illustrate this, based on rate coefficients (related to bond-dissociation energies, BDEs) for the thermal dissociation of PAN (Bridier et al., 1991), iPN (Barker et al., 1977), HNO₃ (Glänzer and Troe, 1974), N₂O₅ (Ammann et al., 2019), CINO₂ (Baulch et al., 1981) and HONO (Tsang and Herron, 1991), we calculated the theoretical 50 % conversion temperature for each molecule as a function of residence time inside the oven (see Fig. S3b). At short residence times the dependence on temperature is very steep (especially for large BDEs) which partially explains the differences between our short heated section inlet and longer ones. However, in practise, we do not know the precise average temperature of the gas at the centre of the oven nor can we characterise the axial and radial gradients in temperature in the quartz tubes; thus, the calculations of fractional dissociation (or complete thermograms) based on bond-dissociation energies are only a rough guide at best. We note that the use of different flows, oven diameters and operational pressures will strongly affect heat transfer from the oven walls to the gas; therefore, reporting the temperature of the external oven wall (as done here and in all reports in the literature) to some extent precludes comparison between different set-ups. The width of the thermograms (i.e. the temperature difference between e.g. 10% and 90% dissociation) will also depend on details of axial and radial temperature gradients in the tubing located within the oven as well as in the downstream section of tubing, which represents a transition regime between oven and room temperature. The impact of temperature gradients inside the quartz tube was explored by calculating the HNO3 thermogram using an Arrhenius expression for its thermal dissociation and the gas residence time within the quartz tube. We initially assumed that all HNO₃ molecules experience the same temperature and then compared this to the situation in which 20 % of the HNO3 molecules experience an 80 °C lower temperature and 20 % experience an 80 °C higher temperature. The resultant thermograms are displayed in Fig. S3c and indicate that the presence of temperature gradients results in an increase in the width of the thermogram from 250 to 350 °C.

The thermograms we report here serve only to determine the temperature needed to ensure the maximum conversion of each trace gas to NO₂. This is achieved in the present set-up with a temperature of 850 °C. Where possible, we have verified that operation at the plateau of the thermogram resulted in quantitative conversion of the traces gases and particles studied, with one exception – NaNO₃ particles. We further note that, in an instrument designed only to measure NO_y , there is no need to ensure separation (in temperature) of the thermograms for different classes of molecules.

3.1.9 Detection of NH₃

As described previously (Wild et al., 2014; Womack et al., 2017) ammonia represents a potential interference in NO_{ν} measurements. In order to quantify this interference, we measured NO₂ formation in air containing 131 ppbv NH₃ delivered by a calibrated permeation source (VICI METRON-ICS, permeation rate 324 ng min^{-1} at $45 \,^{\circ}\text{C}$). The results are summarised in Fig. 2. In NH₃-air mixtures, we observe a small NO₂ signal, increasing slowly at first and then (from \approx 700 °C) rapidly with temperature; the amount of NO₂ observed at 850 °C corresponds to a fractional conversion of NH₃ to NO₂ of 0.006 ± 0.002 . This result is in broad agreement with Wild et al. (2014), who found a conversion efficiency of < 0.01 at 700 °C. In experiments with NH₃ in zero air with relative humidities of 17 %, 31 % and 53 %, we were unable to observe the conversion of NH₃ to NO₂, which is again consistent with the humidity-related suppression of NO_2 formation observed by Wild et al. (2014).

In additional experiments, we investigated the potential influence of ozone on the NH₃ to NO₂ conversion efficiency in zero air containing O₃. The addition of O₃ results in a significant increase in NO₂ with a linear dependence on the O₃ mixing ratio (Fig. 3) and up to 11.4 % conversion of NH₃ to NO_2 at 200 ppbv O_3 . This was not reduced measurably by the addition of water vapour to the air-O3 mixture. In further experiments, we spiked air with the headspace of various organic liquids (acetone, methanol, ethanol, β -pinene, limonene and isoprene). The gas-phase mixing ratios of the organic trace gases were unknown, but the formation of NO₂ was suppressed or completely stopped in each case. A more quantitative investigation was carried out using a known concentration (1 ppmv gas bottle) of isoprene. We found that the addition of 30 ppbv isoprene to zero air (containing 330 ppbv O₃) did not significantly reduce the NH₃-to-NO₂ conversion efficiency under dry conditions, but it decreased it by a factor of 2 when the RH was increased to 50 %.

A tentative chemical mechanism, based partially on Womack et al. (2017), to explain the formation of NO₂ from NH₃ and O₃ at high temperatures and the processes that suppress it is given in Reactions (R11) to (R15). In this scheme, the oxidation of NH₃ is initiated and propagated by O(³P), which is formed from the thermal dissociation of O₃ (Peukert et al., 2013). This leads to the formation of NO and HNO (Reaction R13a and R13b), both of which can be oxidised to NO₂ (Reactions R14 and R15). Forward and reverse rate coefficients for Reaction (R11) indicate that O₃ is converted almost stoichiometrically to O(³P) in the



Figure 3. NH_3 to NO_x conversion in the heated inlet channel of the instrument in the presence of O_3 . The fractional conversion of NH_3 to NO_x is calculated from the 13.1 ppmv of NH_3 from the permeation source.

≈ 10 ms reaction time in the heated inlet. The rate constants (at 1123 K) for the subsequent reactions involving O(³P) are as follows: $k_{12} = 4.4 \times 10^{-13} \text{ cm}^3 \text{ molec.}^{-1} \text{ s}^{-1}$, $k_{13a} = 8.3 \times 10^{-12} \text{ cm}^3 \text{ molec.}^{-1} \text{ s}^{-1}$ and $k_{13b} =$ $7.5 \times 10^{-11} \text{ cm}^3 \text{ molec.}^{-1} \text{ s}^{-1}$ (Cohen and Westberg, 1991). Reaction (R12) converts 0.3% of the initial NH₃ molecules to NH₂ within 10 ms (at 100 ppbv O₃ and 1123 K).

| $O_3 + M \to O(^{\circ}P) + O_2 + M \tag{R1}$ | $(^{3}P) + O_{2} + M$ (R11) |
|---|-----------------------------|
|---|-----------------------------|

$$O(^{3}P) + NH_{3} \rightarrow NH_{2} + OH$$
 (R12)

$$NH_2 + O(^{3}P) \rightarrow NO + H_2$$
 (R13a)

$$\rightarrow$$
 HNO + H (R13b)

$$HNO + O(^{3}P)/OH/H \rightarrow NO + OH/H_{2}O/H_{2}$$
 (R14)

$$NO + O(^{3}P) + M \rightarrow NO_{2} + M$$
(R15)

The experimental results obtained in zero air indicate that reactions involving $O(^{3}P)$ from O_{3} thermolysis can result in the conversion of NH₃ to NO and NO₂. These results could, however, not be reproduced when adding NH₃ to ambient air sampled from outside of the building. In this case, the addition of NH₃ (at 50–60 ppbv O₃) did not result in a measurable increase in NO₂, which was in accord with the observations of Womack et al. (2017). The scavenging of NH₂ radicals and $O(^{3}P)$ by both volatile organic compounds and H₂O provides a likely explanation for this. Womack et al. (2017) also found that the addition of 100 ppbv CO can reduce the conversion of NH₃ to NO_x.

In summary, our experiments indicate that the conversion of NH_3 to NO_2 is suppressed in ambient air samples, or in synthetic air with added VOCs and water. The ambient air used in these experiments was from an urban and polluted environment (typical NO_x levels between 10 and 50 ppby; see Sect. 4.2). As high levels of atmospheric NH₃ are associated with agricultural activity (Langford et al., 1992; Schlesinger and Hartley, 1992) and are often accompanied by high NO_x and VOC levels, the NH₃ interference under these conditions is most likely to be small compared with ambient NO_z levels. Long-term measurements of NH₃ have additionally found a positive correlation between NH₃ concentrations and ambient temperature (Yamamoto et al., 1988; Wang et al., 2015; Yao and Zhang, 2016), with the latter promoting the presence of high levels of biogenic VOCs, such as isoprene (Tingey et al., 1979), which would also help to minimise the NH₃-related interference.

3.2 Bias caused by secondary reactions in the TD ovens

Thermal dissociation techniques coupled to CRD systems for the measurement of organic nitrates suffer bias to different degrees owing to reactions of organic peroxy radicals with NO and NO₂ (Sobanski et al., 2016; Thieser et al., 2016). According to previous studies (Day et al., 2002; Rosen et al., 2004; Thieser et al., 2016), in experiments using iPN at an oven temperature of 450 °C, an overestimation of ANs in the presence of NO is caused by reactions of the initially formed alkoxy radical, C₃H₇O, which results in the formation of both HO₂ and CH₃O₂ (Reactions R16–R21).

| $C_3\Pi_7ONO_2 + M \rightarrow C_3\Pi_7O + NO_2 + M \tag{KIC}$ | (R16 | $_{2} + M$ | NO_2 |)+ | 1_7C | C_3H | $M \rightarrow$ | $)_2 + 1$ | ON | C_3H_7 | (|
|--|------|------------|--------|----|--------|--------|-----------------|-----------|----|----------|---|
|--|------|------------|--------|----|--------|--------|-----------------|-----------|----|----------|---|

$$C_{3}H_{7}O + O_{2} \rightarrow CH_{3}C(O)CH_{3} + HO_{2}$$

$$(R17)$$

$$C_{2}H_{7}O + M_{2} \rightarrow CH_{2} + CH_{2}CHO + M_{2}$$

$$(P18)$$

$$C_{3}H_{7}O + M \rightarrow CH_{3} + CH_{3}CHO + M$$

$$(R18)$$

$$CH_{2} + O_{2} + M \rightarrow CH_{3}O_{2} + M$$

$$(R19)$$

$$CH_{3}O_{2} + NO \rightarrow CH_{3}O + NO_{2}$$
(R20)

$$HO_2 + NO \rightarrow HO + NO_2$$
 (R21)

$$CH_3C(O)O_2NO_2 + M \rightarrow CH_3C(O)O_2 + NO_2 + M \quad (R22a)$$

$$CH_3C(O)O_2 + NO_2 + M \rightarrow CH_3C(O)O_2NO_2 + M \quad (R22b)$$

$$CH_3C(O)O_2 + M \rightarrow CH_3CO + O_2 + M$$
(R23a)

$$CH_3C(O)O_2 \rightarrow CH_2C(O)OOH$$
 (R23b)

$$OH + NO_2 + M \rightarrow HNO_3 + M$$
 (R24)

In order to investigate the potential bias in the measurement of ANs under the present experimental conditions, a set of experiments was conducted in which NO (up to 12 ppbv) was added to various amounts of iPN. The NO mixing ratio was determined by modulating the addition of O₃. The results (Fig. 4a) show that NO₂ derived from the thermal decomposition of iPN increases with the amount of NO added and results in overestimation (factor of ~ 1.6) at 12 ppbv NO, which is consistent with the observations by Thieser et al. (2016). This disappears when 19 ppmv ozone is added in front of the cavity so that NO_x rather than NO₂ is measured (blue data points). This is readily explained by the compensation of the additional NO₂ formed via reactions of NO with RO₂ by an equal loss in NO, which is only detected when introducing O₃. This is illustrated graphically in Fig. S4.



Figure 4. Investigation of bias caused by reactions of NO with HO₂ and RO₂ when measuring iPN. (a) NO varied for two initial iPN mixing ratios in the presence (blue data points) and absence (orange data points) of added O₃. The NO_x background signal from the iPN cylinder was subtracted from the iPN mixing ratios. (b) The investigation of bias caused by the recombination of RO₂ and NO₂ during the thermal decomposition of PAN. In both experiments, the oven temperature was 850 °C. In both plots, the error bars indicate the standard deviation over the averaging interval.

We also explored the potential for bias caused by the recombination of $CH_3C(O)O_2$ and NO_2 (Reaction R22b), following the thermal decomposition of PAN (Reaction R22a). Thieser et al. (2016) reported that PAN was underestimated by a factor 0.45 when adding 10 ppbv NO₂ to an air sample containing PAN at 200 °C. This behaviour was not apparent at 450 °C, which is related to the decomposition (Reaction R23a) or isomerisation (Reaction R23b) of $CH_3C(O)O_2$ at the higher temperature. The results of a similar experiment with our 850 °C inlet are presented in Fig. 4b. In this plot, the measured NO_2 relative to the input PAN is plotted versus the mixing ratio of added NO_2 . For PAN concentrations from 1.5 to 2.2 ppbv, no effect was observed for NO_2 concentrations of up to 10 ppbv. This is consistent with the reaction scheme presented by Thieser et al. (2016) at $450 \,^{\circ}$ C.

A potential source of bias when measuring HNO₃ includes its reformation via the reaction of OH and NO₂ (Reaction R24). Compared with the RO and RO₂ radicals formed in the thermal dissociation of PNs and ANs, OH exhibits a higher affinity for surfaces and is likely to be efficiently removed at the oven wall. Day et al. (2002) estimated that wall losses are the dominant OH sink and that the resulting underestimation of HNO₃ would be < 2% for NO_y levels < 5 ppbv. At our oven temperature, the diffusion coefficient for OH (D_{OH}) can be calculated according to Tang et al. (2014):

$$D_{\rm OH} (1123 \,\mathrm{K}) = D_{\rm OH} (296 \,\mathrm{K}) \cdot \frac{296^{-1.75}}{T}.$$
 (2)

Using an average of the literature values for D_{OH} at room temperature from Ivanov et al. (2007), 165 Torr cm² s⁻¹, and Bertram et al. (2001), 192 Torr cm² s⁻¹, a value of D_{OH} (1123 K) = 1841 Torr cm² s⁻¹ was derived. The maximum rate constant for OH wall loss (assuming laminar flow) can subsequently be approximated according to Zasypkin et al. (1997):

$$k_{\text{wall}} = \frac{D_{\text{OH}} \cdot 3.66}{r^2 \cdot p} \tag{3}$$

With the radius of the oven quartz tube r (0.45 cm) and the pressure p (760 Torr), the maximum value of k_{wall} is 44 s⁻¹. The first-order loss rate coefficient for the reaction of OH with NO₂ is given by $k_{(OH+NO_2)}$ [NO₂], where $k_{(OH+NO_2)}$ is the rate coefficient for the reaction between OH and NO₂ at 1123 K ~ 5×10⁻¹⁴ cm³ molec.⁻¹ s⁻¹ (Ammann et al., 2019) and [NO₂] = 6.5×10^{10} molec. cm⁻³ (the concentration of 10 ppb NO₂ at the pressure and temperature of the oven). The first-order loss rate of OH via reaction with NO₂ is then 3×10^{-3} s⁻¹. Clearly, the efficiency of uptake of OH to the wall would have to be very low in order to reduce the maximum value of 44 s⁻¹ to values that are comparable to the reaction with NO₂, which is very unlikely. We conclude that reformation of HNO₃ via Reaction (R24) will not bias measurements of HNO₃ with the present set-up.

3.3 Denuder characterisation

The efficiency of the removal of NO_y trace gases and the transmission of submicron particles was determined in a series of experiments, which are described below.

3.3.1 Transmission of ammonium nitrate particles (10–414 nm)

In order to characterise the transmission of the denuder for particles of different diameter, a constant flow of particles was generated by passing 3.3 slpm of nitrogen through an atomiser (TSI 3076) containing an aqueous solution of ammonium nitrate. The flow rate (3.3 slpm) was matched to the typical sampling flow through the denuder. A total of 0.28 slpm of the flow was sampled into a scanning mobility particle sizer (SMPS)/CPC system (TSI 3080 and TSI 3025A) to measure the number density and size distribution (10-414 nm) of the ammonium nitrate particles. The flow was delivered either directly to the SMPS/CPC via straight metal tubing (length 27 cm, inner diameter 0.9 cm), for which we assume 100 % particle transmission, or via the denuder. Thus, the ratio of the particle numbers in each size bin represents the size-dependent denuder transmission. As shown in Fig. 5, the transmission of the denuder is > 80 % for particles between 30 and 400 nm in diameter. As expected, some diffusive loss is observed for particles < 20 nm in diameter and loss due to impaction and/or settling is observed for particles > 300 nm. The particle transmission T as a function of particle diameter D can be represented by the following empirical expression:

T(%) =

$$\frac{D(\text{nm}) - 5.79}{0.035 + 0.010 \cdot (D - 5.79) + 1.78 \cdot 10^{-6} \cdot (D - 5.79)^2} \quad (4)$$

The particle transmission through the denuder channels was also calculated using the Particle Loss Calculator (PLC) developed by von der Weiden et al. (2009). The results of this calculation are plotted in Fig. 5. The observed loss of particles smaller than 40 nm is not replicated by the PLC, which was developed for cylindrical piping and not for the square honeycomb shape of the denuder; thus, it does not consider losses due to impact at the finite surface area which the gas and particle flow is exposed to at the entrance to the honeycomb. The PLC does a better job at predicting a reduction in transmission for the largest particles that we measured, indicating a transmission of 74 % at 1 μ m and 45 % at 2 μ m. Therefore, in certain environments, nitrate associated with coarse-mode particles represents a potential bias for TD-CRDS measurements of NO_y.

3.3.2 Efficiency of removal of NO_v trace gases

The efficiency of the removal of trace gases in the denuder under typical flow conditions (3.3 slpm) was investigated for NO, NO₂, PAN, iPN, HONO, N₂O₅, ClNO₂ and HNO₃ as representative NO_y species. The efficiency of removal of each trace gas (generally present at 5–40 ppbv) was determined by measuring its relative concentration when flowing through the denuder (pNit channel) and when bypassing the denuder (NO_y channel). The results (Fig. 6) indicate that, in dry air, all of these trace gases were removed with an efficiency of close to 100 %. However, when the main dilution flow was humidified, RH-dependent breakthrough of NO was observed, with only 60 % stripped from the gas phase at an RH close to 100 %. HONO was removed with 85 % efficiency at an RH of 46 %, and ClNO₂ was removed



Figure 5. Transmission of ammonium nitrate particles through the denuder inlet. Relative transmissions are derived by dividing the number size distribution when sampling through the denuder by a size distribution obtained without the denuder. Error bars are based on the standard deviation of three consecutive measurements with and without the denuder. An aerosol flow of 3.3 slpm was directed through the denuder (diameter 3 cm, see Sect. 2.4), and a DMA subsequently sampled 0.3 slpm from the stream exiting the denuder. The plot also includes a fit of the experimental (solid, black line) data and a theoretical transmission distribution computed with the Particle Loss Calculator (PLC).

with a 75 % efficiency at an RH of 60 %. In contrast, humidification only had a marginal effect on the scrubbing efficiency for NO₂, iPN and HNO₃, for which an efficiency of \geq 95 % was observed. The precise values that the removal efficiencies in Fig. 6 were determined from are listed in Table S2.

In further experiments, we examined the potential for rerelease of NO_{ν} that had previously been stored in the activated carbon substrate of the denuder. In these experiments, in which either NO_x or NO_y was continuously monitored, the denuder was exposed to a flow of 9.5 ppm iPN in dry nitrogen for 90 min during which 2.30×10^{17} molecules of iPN were stripped from the gas phase and deposited onto the denuder. This exposure is equivalent to a month-long exposure to 20 ppb of iPN. The air passing through the denuder was subsequently humidified to an RH of 65 %. The results (Fig. 7a) indicate a high initial (11:40-11:50 UTC) rate of release of NO_x under humid conditions (resulting in a maximum mixing ratio of 2 ppbv at 11:45 UTC). At 11:50 UTC, humidification of the air was stopped and the rate of release of NO_x dropped gradually towards zero. During a second period, in which the air was again humidified (from 11:50 UTC onwards), NO_x was released from the denuder, albeit at a lower rate than during the first humidification. From 12:25 UTC onwards, the oven behind the denuder was heated to 850 °C so that NO_{ν} was added. No significant increase in NO_2 was



Figure 6. Removal efficiency of the denuder for various NO_y trace gases as a function of RH. Units of the inset are identical to the main graph. See Table S2 for the exact values and information on the error determination.

observed, indicating that the trace gas(es) species released from the denuder surface upon humidification are predominantly NO_x. During this experiment, 2.55×10^{15} molecules of NO_x desorbed from the denuder, indicating that the major fraction of iPN molecules remained stored on the denuder surface upon humidification.

Similar denuder exposure experiments were performed with HNO3 and NO2. For HNO3, no evidence for desorption of NO_x or NO_y during exposure to humidified air was observed, whereas NO2 exhibited a similar behaviour to iPN (Fig. 7b). After loading the denuder with 5 sccm from a 0.831 ppm NO₂ gas bottle for 4.8 d (a total of 7.60×10^{17} molecules were deposited as derived from the flow rate, the exposure time and the gas bottle mixing ratio), the effect of passing humidified air through the denuder was to release NO_x, which was observed at concentrations up to \approx 39 ppbv. While the relative humidity was kept constant at close to 100 %, the NO_x released decreased over time so that after $30 \min$, 3.2 ppbv of NO_x could still be detected. By switching the O₃ source off (at \approx 10:45 UTC), the NO₂ measured went to ≈ 0 , indicating that predominantly NO was released (and not NO₂). Integrating these results over time yielded a value of 1.63×10^{16} molecules desorbed NO from the denuder surface in humid air. Qualitatively similar results, i.e. humidity-induced formation and release of NO_x from the denuder, were observed when the denuder was exposed to variable levels of NO_x (i.e. up to 20 ppbv) under dry conditions for periods of weeks.

Clearly, adsorption of water molecules onto the denuder surface can initiate and/or catalyse chemical transformation at the surface that convert stored NO_z into forms which can desorb and be detected as NO_x . To further our understanding



Figure 7. (a) Release of NO_x from the denuder in humid air after exposure to 9.5 ppmv iPN for 1.5 h. Relative humidity was measured before passing through the denuder. The blue shaded area signifies the period in which the inlet oven was heated to 850 °C. Changes in RH are achieved by flowing parts of the zero air stream through deionised water. (b) The release of NO_x from the denuder in humid air after exposure to 0.83 ppmv NO_2 for 4.8 d. O_3 addition was switched off during the blue shaded period.

of the underlying processes that occur upon humidification, a series of experiments were conducted to examine the adsorption of water on the denuder. In these experiments, the denuder was first dried by exposing it to dry air for several hours until the relative humidity of the air exiting the denuder was close to zero. Subsequently, humidified air was passed through the denuder and the RH of air exiting it was continuously monitored. The results of an experiment in which the air was humidified to 68 % are shown in Fig. 8a. After 77 min of exposure to this humidity, the RH of the air exiting the denuder acquired a maximum value of 64 %. After switching back to dry synthetic air (at 09:37 UTC), ~ 60 minutes passed before the RH dropped to values close to zero. In this period, the RH did not decrease monotonically, with the rate of change of relative humidity exiting the denuder revealing a number of discreet steps. Figure 8b plots the derivatives (dRH/dt) of the drying phases of a series of experiments in which the initial RH was varied between 47 % and 75 %. A similar pattern emerges for each experiment with the greatest desorption rates occurring at the beginning of the drying phase followed by a minimum in the desorption rate and a second maximum (at ≈ 15 % relative humidity). This behaviour is a clear indication that H₂O is bound to more than one chemically or physically distinct surface sites on the activated carbon.

By measuring the change in the RH of the air flowing into and out of the denuder, we can derive an equilibrium adsorption isotherm for H₂O at the active carbon surface. An example is given in Fig. 9 where it is also compared to a literature isotherm for the adsorption of water vapour on activated carbon fibre (Kim et al., 2008). The data of Kim et al. (2008) have been scaled by matching the number of adsorbed water molecules at an RH of 65.9% to our observed value at an RH of 67.2 %. The exposure of carbonaceous surfaces to inert gases at high temperatures (2000 °C) reduces the capacity for water uptake, whereas functionalising the carbon surface with oxygen-containing groups (e.g. from HNO₃) enhances the water adsorption capacity (Dubinin et al., 1982; Barton and Koresh, 1983; Liu et al., 2017). Thus, in our experiments, the uptake of gas-phase NO_{y} is expected to generate oxygenated sites on our denuder surface, which, in turn, will influence water uptake and subsequent further trace gas accommodation.

The chemistry leading to the formation of gas-phase NO_x from NO_y trace gases adsorbed at the denuder surface under humid conditions cannot be elucidated in detail with our experimental set-up. However, a strong humidity dependence in the heterogeneous generation of HONO and NO from NO₂ adsorbed on soot particles has been reported (Kalberer et al., 1999; Kleffmann et al., 1999). Formation of HONO from NO_x has also been observed on wet aerosol and ground surfaces in field studies (Lammel and Perner, 1988; Notholt et al., 1992). Previous investigations report the adsorption of NO₂ on activated carbon at ambient or close to ambient temperatures (30–50 °C), followed by its reduction to NO, with the simultaneous oxidation of the carbon surface (Shirahama et al., 2002; Zhang et al., 2008; Gao et al., 2011). These results are consistent with our observation of processes such as the conversion of NO₂ to NO at the denuder surface under humid conditions. In our experiments, we observed that NO₂ was converted to NO (rather than HONO) at the denuder surface. It is possible that the initial step is the formation of HONO (e.g. by the surface-catalysed hydrolysis of NO₂ or iPN) which undergoes further reduction on the surface to NO. The release of HONO and NO has also been observed from soil, after the nitrification of NH_3/NH_4^+ or the reduction of NO_3^- (Su et al., 2011; Pilegaard, 2013; Meusel et al., 2018). Oswald et al. (2013) found compara-



Figure 8. (a) RH of humidified zero air after passing through the denuder. The initial RH was determined by bypassing the denuder before and after the experiment. Zero air was humidified by flowing a fraction of the stream through deionised water stored in a glass vessel. The time at which the experiment was conducted is given on the *x* axis. Until ca. 09:35 UTC, zero air with constant humidity (RH ca. 68 %) was sent through the denuder. Afterwards the denuder was exposed to dry zero air. (b) Derivative of the measured RH during the drying period. The step during the drying phase occurs in a higher RH area when starting the drying from a larger RH value.

ble HONO and NO emission fluxes from non-acidic soils, providing another example of the heterogeneous formation of HONO from other atmospheric nitrogen species, followed by the gas-phase release of NO.

We conclude that the use of this denuder type (and the assumption of complete removal of gaseous NO_y) may potentially result in a positive bias in measurements of particle nitrate owing to the variable breakthrough and release of NO_x



Figure 9. Number of adsorbed water molecules onto the denuder surface at equilibrium versus RH. The red line represents (scaled) results from a study on activated carbon fibre (Kim et al., 2008).

(dependent on the historical exposure of the denuder and relative humidity). Our findings may be applicable (at least in a qualitative sense) to similar denuders using activated carbon surfaces, and careful characterisation of the NO_y components' capacity to adsorb, breakthrough and release should be carried out prior to use in the field.

Reliable surface reactivation techniques for similar denuders would be useful to ensure continuous, efficient scrubbing of NO_y and NO_x and circumnavigate the potential overestimation of pNit. In this regard, attempts to "reactivate" the denuder by cleaning with distilled water, drying at 50 °C and exposure to ca. 300 ppbv O₃ for 1 h did not result in an improvement of the direct NO breakthrough or in the background pNit signal upon humidification. Surface-sensitive spectroscopic investigation of the water-induced transformation of organic and inorganic NO_y to NO_x (and its subsequent release to the gas phase) on denuder surfaces would be useful in resolving these issues.

3.4 Comparison of the TD-CRDS with existing methods for NO_x and NO_y and instruments

In this section, our instrument's performance is compared to other methods for NO_x and NO_y detection. We consider only instruments that measure NO_x and/or NO_y and not those that measure individual trace gases from each family.

 NO_x has traditionally been measured by two-channel CLD instruments, which have one channel for NO and one channel for NO₂ respectively: NO is detected by chemiluminescence from its reaction with O₃, and NO₂ is converted to NO on a molybdenum converter or by photolysis at wavelengths close to 390 nm. Examples of this type of instrument as well as the reported levels of detection (LODs) and overall uncertainties are listed in Table 1. In Table 1, we also list the LOD and uncertainty of a recently developed CRDS set-up which is **Table 1.** Comparison of NO_x and/or NO_y measurements.

| Species | Reference | Method | 2σ level of detection (integration time) | Uncertainty (%) |
|-----------------|------------------------|-------------------------|---|--------------------|
| NO _x | Parrish et al. (2004) | CLD | 20 pptv (1 s) | 10 |
| | Fuchs et al. (2009) | CRDS | 22 pptv (1 s) | 5 |
| | Wild et al. (2014) | CRDS | < 30 pptv (1 s) | 5 |
| | Reed et al. (2016) | CLD lab CLD aircraft | 2.5 pptv (60 s) ~ 1.0 pptv (60 s) | 5 5 |
| | This study | CRDS | 40 pptv (60 s) | 6 |
| NOy | Fischer et al. (1997) | CLD | 200 pptv (6 s) | 25 |
| | Williams et al. (1998) | CLD "BNL" CLD "NOAA" | ~ 50 pptv (1 s) 20 pptv (1 s) | 10 18 |
| | Day et al. (2002) | LIF | \sim 10 pptv (10 s) | < 5 |
| | Parrish et al. (2004) | CLD | 36 pptv (1 s) | 10 |
| | Wild et al. (2014) | CRDS | < 30 pptv (1 s) | 12 |
| | Pätz et al. (2006) | CLD CLD | 51 pptv (1 s) 100 pptv (1s) | 13 9 |
| | This study | CRDS | 40 pptv (60 s) | 15* |

Note: * Refers to gas-phase NO_y only.

similar in principal of operation to the one described here (Fuchs et al., 2009).

 NO_y has frequently been measured by CLD instruments with Au and/or MoO-coated thermal dissociation inlets that reduce NO_z to NO, which is detected as described above for CLD- NO_x instrument. Although such instruments have very low detection limits, they have been shown to be vulnerable to degradation of the NO_y conversion efficiency and suffer from interferences by HCN, CH₃CN and NH₃ (Kliner et al., 1997) as well as loss of NO_y in the inlet (Zenker et al., 1998; Parrish et al., 2004). Table 1 summarises the LODs and uncertainties reported by other NO_y instruments.

Table 1 indicates that the total uncertainty of the present instrument is comparable to those reported for both NO_x and NO_y . However, our present detection limit for both NO_x and NO_y is worse than that reported (for NO_2) for the same instrument in 2016 (Thieser et al., 2016), which is a result of mirror degradation since that study.

4 Application of the instrument in field experiments

4.1 NO_x intercomparison and pNit measurements during the AQABA campaign

The first deployment of the instrument was during the AQABA (Air Quality and climate change in the Arabian Basin) ship campaign in summer 2017. From 31 July to

2 September, the ship "*Kommander Iona*" followed a route from southern France via the Mediterranean Sea, the Suez Channel, the Red Sea, the Arabian Sea and the Arabian Gulf to Kuwait and back. The instrument was located in a container in front of the ship, with the inlet ovens located in an aluminium box on the roof of the container. The (unheated) tips of quartz inlet tubes protruded about 15 cm from the side of the aluminium box. Here, we compare the NO_x and pNit measurements with other measurements of these parameters made during the campaign.

During AQABA, NO_x levels ranged from a few parts per trillion by volume (maritime background) up to several tens of parts per billion by volume in heavily polluted air masses in shipping lanes or in harbours. In Fig. 10a, we compare NO_x measured with the TD-CRDS with the results of a chemiluminescence detector (CLD 790 SR, ECO PHYSICS; Tadic et al., 2020), which measured NO and NO₂. The data points represent 1 min averages for the entire campaign, excluding air masses that were contaminated by the ship's exhaust. Additionally, periods with very high NO_x variability were not included, with a data point being discarded whenever the differences in mean values exceeded 2 ppbv for consecutive data points.

A bivariate fit to the data sets (York, 1966), which incorporates total uncertainties for both instruments (CLD: 8.6 %; TD-CRDS: $11 \% + 20 \text{ pptv} \times \text{RH}/100$) resulted in a slope of 0.996 ± 0.003 and an intercept of -1.3 pptv. This very good agreement serves to underline the general applicability of the



Figure 10. (a) Correlation between the TD-CRDS NO_x measurements (1 min averages) and an independent CLD NO_x instrument from the AQABA campaign. Data obtained during phases of very high NO_x variability have been excluded (see Sect. 4.1). See Fig. S6 for a histogram of the NO_x data points. (b) pNit measurements using the denuder channel (blue data points) during AQABA and comparison with particulate NO_3^- from an AMS (aerosol mass spectrometer). The discrepancy towards the AMS and the correlation with the NO_x mixing ratios indicate a positive bias in the pNit measurements, caused by humidity effects on the denuder surface. OPC (optical particle counter) measurements are added in the lower panel of (b) in order to assess the potential influence of coarse-mode aerosol nitrates.

TD-CRDS in NO_x measurements, even under difficult conditions (e.g. a non-static platform). In this context, we note that the deployment on a ship resulted in a degradation in performance (the LOD was $\approx 100 \text{ pptv}$) owing to the ship's motion, especially in heavy seas, which resulted in drifts in the instrument zero.

Figure 10b shows a ca. 10h time frame with pNit measurements from the denuder channel of the TD-CRDS. Unfortunately, the TD oven of the denuder channel broke down very early in the campaign and was not operational afterwards. The data from the pNit channel are presented along with the NO_x and NO_y measurements as well as particulate nitrate mass concentrations measured by an aerosol mass spectrometer (Aerodyne HR-ToF-AMS; DeCarlo et al., 2006; Brooks et al., 2020). The night-to-day transition is indicated via the NO₂ photolysis rates J_{NO_2} derived from a spectral radiometer (Metcon GmbH). The relative humidity was > 80% throughout the period shown. During the two periods when, apart from some short spikes, NO_x was very low (21:10-23:45 and 02:20-03:40 UTC), TD-CRDS data indicate the presence of 300-400 pptv of pNit, which would then constitute $\sim 80 \%$ NO_v. Such mixing ratios of particulate nitrate are not commensurate with those measured by the AMS, which, on average, are a factor of 6-8 lower. As the AMS does not detect particles larger than ~ 600 nm with high efficiency (Drewnick et al., 2005), the difference could potentially indicate that a significant fraction of the particulate nitrate is associated with coarse-mode aerosol. In the lower panel of Fig. 10b, we plot the coarse-mode aerosol mass concentration determined from measurements of an optical particle counter (OPC) that measures particles between 0.2 and 20 μ m. In the two low-NO_x periods outlined above, the OPC-derived aerosol mass concentrations were between 3 and $5 \mu \text{g m}^{-3}$. If 10% of this coarse-mode aerosol mass concentration was nitrate, which is a typical value in the Mediterranean (Koulouri et al., 2008; Calzolai et al., 2015; Malaguti et al., 2015), this would account for 100-200 pptv of the pNit observed by the TD-CRDS and not by the AMS. However, the time profile of pNit measured by the TD-CRDS is not consistent with those of either the OPC or the AMS and rather resembles the NO_x mixing ratios. This strongly suggests that the large difference between pNit reported by the TD-CRDS and the AMS does not result from the nondetection or detection of supermicron particulate nitrate by the AMS but instead results from the denuder artefacts described in Sect. 3.3.2. This short case study serves to highlight the potential positive bias in denuder-based TD-CRDS measurements of pNit under humid field conditions.

4.2 Ambient NO_x and NO_y measurements in an urban environment (Mainz, Germany)

 NO_x and NO_y mixing ratios were obtained in air sampled outside the Max Planck Institute for Chemistry (MPIC). The MPIC (49°59′27.5″ N 8°13′44.4″ E) is located on the outskirts of Mainz but within 200 m of two busy two- and fourlane roads and within 500 m of additional university buildings as well as commercial and residential areas. The city of Mainz (217 000 inhabitants) is located in the densely populated Rhine–Main area along with the cities of Frankfurt (753 000 inhabitants) and Wiesbaden (278 000 inhabitants); the air in this region is strongly influenced by local pollution. The sampling location was on the top floor of a three-story building (ca. 12 m above ground level). Air was subsampled to the inlets of the instrument from a \sim 1 m long 0.5 inch outer diameter PFA tube that was connected to a membrane pump and flow controller to generate a 20 slpm bypass flow. Aerosol transmission was probably < 100 % in these measurements.

Figure 11a summarises the 8 d of measurement (data coverage 82 %) as a time series for NO_x , NO_y , NO_z , (10 min averages) wind speed (1 h averages) and the NO_z/NO_y ratio. The NO_x and NO_y mixing ratios were highly variable throughout this period, with NO_x mixing ratios between 0.7 and 148.3 ppbv (mean and median values of 22.1 and 6.9 ppbv respectively). Traffic-related morning rush-hour peaks in NO_x were observed on all weekdays (14, 16, 17) and 20 January) between 05:00 and 10:00 UTC. The morning NO_x peak is reduced or absent on the weekends (18 and 19 January). NO_x levels stayed above 50 ppbv for nearly a full day from 18:00 UTC on 16 January until 18:00 UTC on 17 January, which coincides with constantly low wind speeds and the sampling of air masses that were predominantly local, and thus highly polluted. NO₇ mixing ratios were usually between 0.5 and 2.5 ppbv (minimum < LOD, maximum 3.1 ppbv, mean 1.0 ppbv and median 0.9 ppbv), with NO_7/NO_y ratios below 0.5. These values indicate that the air masses had been impacted by recent (local) NO_x emissions.

The NO_z/NO_v ratio can be used as indicator of the degree of chemical processing of an air mass. In Fig. 11b, a median diel profile (including all measurement days) for the NO_z/NO_v ratio from the ambient measurement is shown. The diel profile displays two distinct minima in NO_z/NO_v during the morning and evening rush hours, where NO_z only makes up 5 %–10 % of the total NO_{ν}. This fraction increases up to 15% during midday and up to 25% during nighttime, when emissions of NO_x are reduced. The diel profiles of NO_z/NO_v are strongly influenced by fresh emissions of NO_x . As the measurement location is strongly influenced by traffic, there is a decrease in NO_x (and an increase in NO_z/NO_y) at night-time. Night-time increases in NO₇ (13-14, 15-16, 18-19 and 19-20 January 2020) may also be partially caused by the formation of N2O5 as previously observed (Schuster et al., 2009) and which would have been favoured by the low night-time temperatures ($< 10 \,^{\circ}$ C) in winter.

These measurements serve to illustrate the applicability of our TD-CRDS over a wide range of NO_x and NO_y concentrations under realistic field conditions and in the investigation of processes that transform NO_x into its gas- and particle-phase reservoirs.



Figure 11. (a) Time series of NO_x , NO_y , NO_z , NO_z/NO_y and wind speed from ambient measurements in Mainz, Germany, in January 2020. Highly variable NO_x (between 0 and 150 ppbv) and moderate NO_z (between 0 and 3 ppbv) mixing ratios were observed, identifying the sampled air masses as being dominated by anthropogenic emissions. Wind speed data was obtained from Agrarmeteorologie Rheinland-Pfalz (http://www.wetter.rlp.de, last access: 20 October 2020). (b) Diel profile of the NO_z/NO_y ratio including all measurement days, showing distinct minima during the morning and evening rush hours. Shaded areas signify the time between sunset and sunrise.

5 Conclusions

We report on the development, characterisation and first deployment of a TD-CRDS instrument for the measurement of NO_x, NO_y, NO_z and pNit. Our laboratory experiments suggest that the different gas-phase NO_z species investigated (PAN, iPN, N₂O₅, HONO, ClNO₂ and HNO₃) are converted with near-stoichiometric efficiency to NO_x at a nominal oven temperature of 850 °C. NH₄NO₃ particles of 200 nm in diameter are also detected quantitatively as NO_x, whereas the efficiency of detection of NaNO₃ particles of similar diameter was closer to 25 %. The efficiency of the detection of coarse-mode particles will be further reduced by their lower transmission through the denuder.

The potential for NH₃ to bias NO_y measurements was assessed and was found to be insignificant in ambient air or synthetic air containing VOCs and water. The conversion to NO₂ (by reaction with O₃) of atmospheric NO, as well as NO formed in the heated inlet, circumvents bias resulting from O₃ thermolysis (leading to an NO₂ overestimation) and secondary processes, which are initiated by the thermal dissociation of organic nitrates.

For our activated carbon denuder, we observed > 90 % transmission for ammonium nitrate particles with diameters between 40 and 400 nm. Under humid conditions, the denuder suffered from the direct breakthrough of NO and the re-release of previously stored iPN and NO₂ in the form of NO, indicating a potential bias of pNit measurements using this technique and potentially limiting its deployment to low-NO_x and low-NO_z environments. When using comparable denuders, we recommend regular checks with humidified zero air to characterise potential breakthrough. Our experiments demonstrated that the release of NO_x from the denuder exposed to humid zero air for several hours can decrease to values below 1 ppbv, which, in a first approximation, could be treated as an offset. Cycling between multiple denuders would help to reduce the size of any bias.

The performance of the instrument under field conditions was demonstrated by measurements in Mainz, Germany, and during the AQABA ship campaign. NO_x measurements with the new instrument were in good agreement with those from an established, independent CLD-based instrument.

Data availability. The NO_x data set from the AQABA campaign is available at https://doi.org/10.5281/zenodo.4134659 (Friedrich, 2020). Data used to generate the figures in this paper can be obtained upon request from the corresponding author.

Supplement. The supplement related to this article is available online at: https://doi.org/10.5194/amt-13-5739-2020-supplement.

Author contributions. NF developed the TD-CRDS, performed all laboratory and campaign measurements, evaluated the data sets and wrote the paper. IT and HF provided the AQABA CLD NO_x measurements. JS designed the heated inlet system and performed actinic flux measurements during AQABA. JB, ED and FD provided AMS and OPC measurements from AQABA. JL and JNC designed and supervised the study and the campaigns. JNC, JL and FD contributed to the paper.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. We thank Ezra Wood for providing us with the activated carbon denuder and Chemours for the provision of the FEP sample (FEPD 121) used to coat the cavity walls. This work was supported by the Max Planck Graduate Center with the Johannes Gutenberg-Universität Mainz (MPGC).

Financial support. The article processing charges for this openaccess publication were covered by the Max Planck Society.

Review statement. This paper was edited by Dwayne Heard and reviewed by three anonymous referees.

References

- Atsuko, N., Kazuya, S., Toshiaki, E., Kei-ichi, K., Morinobu, E., and Norifumi, S.: Electronic and Magnetic Properties of Activated Carbon Fibers, Bull. Chem. Soc. Jpn., 69, 333–339, https://doi.org/10.1246/bcsj.69.333, 1996.
- Barker, J. R., Benson, S. W., Mendenhall, G. D., and Goldern, D. M.: Measurements of rate constants of importance in smog, Rep. PB-274530, Natl. Tech. Inf. Serv., Springfield, Va., USA, 1977.
- Barton, S. S. and Koresh, J. E.: Adsorption Interaction of Water with Microporous Adsorbents 1. Water-Vapor Adsorption on Activated Carbon Cloth, J. Chem. Soc. Faraday Trans. I, 79, 1147– 1155, https://doi.org/10.1039/f19837901147, 1983.
- Baulch, D. L., Duxbury, J., Grant, S. J., and Montague, D. C.: Evaluated kinetic data for high temperature reactions. Volume 4 Homogeneous gas phase reactions of halogen- and cyanidecontaining species, J. Phys. Chem. Ref. Data, 10, 1–721, 1981.
- Bertram, A. K., Ivanov, A. V., Hunter, M., Molina, L. T., and Molina, M. J.: The reaction probability of OH on organic surfaces of tropospheric interest, J. Phys. Chem. A, 105, 9415–9421, 2001.
- Bridier, I., Caralp, F., Loirat, H., Lesclaux, R., Veyret, B., Becker, K. H., Reimer, A., and Zabel, F.: Kinetic and Theoretical-studies of the Reactions $CH_3C(O)O_2 + NO_2 + M \leftrightarrow CH_3C(O)NO_2 + M$ between 248 K and 393 K and Between 30-torr and 760-torr, J. Phys. Chem., 95, 3594–3600, 1991.
- Brooks, J., Darbyshire, E., Drewnick, F., Alfarra, R., Struckmeier, C., Fachinger, F., Borrmann, S., Allen, G., and Coe, H.: Characterisation of accumulation mode aerosol across the Middle East and Mediterranean during a ship campaign in the summer of 2017, in preparation, 2020.
- Brunauer, S., Emmett, P. H., and Teller, E.: Adsorption of gases in multimolecular layers, J. Am. Chem. Soc., 60, 309–319, 1938.
- Calzolai, G., Nava, S., Lucarelli, F., Chiari, M., Giannoni, M., Becagli, S., Traversi, R., Marconi, M., Frosini, D., Severi, M., Udisti, R., di Sarra, A., Pace, G., Meloni, D., Bommarito, C., Monteleone, F., Anello, F., and Sferlazzo, D. M.: Characterization of PM₁₀ sources in the central Mediterranean, Atmos. Chem. Phys., 15, 13939–13955, https://doi.org/10.5194/acp-15-13939-2015, 2015.

Cohen, N. and Westberg, K. R.: Chemical Kinetic Data Sheets for High-Temperature Reactions .2., J. Phys. Chem. Ref. Data, 20, 1211–1311, https://doi.org/10.1063/1.555901, 1991.

- Davidson, J. A., Viggiano, A. A., Howard, C. J., Dotan, I., Fehsenfeld, F. C., Albritton, D. L., and Ferguson, E. E.: Rate Constants for Reactions of O_2^+ , NO_2^+ , NO^+ , H_3O^+ , CO_3^- , NO_2^- , and Halide Ions with N₂O₅ at 300 K, J. Chem. Phys., 68, 2085–2087, 1978.
- Day, D. A., Wooldridge, P. J., Dillon, M. B., Thornton, J. A., and Cohen, R. C.: A thermal dissociation laser-induced fluorescence instrument for in situ detection of NO₂, peroxy nitrates, alkyl nitrates, and HNO₃, J. Geophys. Res.-Atmos., 107, 4046, https://doi.org/10.1029/2001jd000779, 2002.
- Day, D. A., Dillon, M. B., Wooldridge, P. J., Thornton, J. A., Rosen, R. S., Wood, E. C., and Cohen, R. C.: On alkyl nitrates, O₃, and the "missing NO_y", J. Geophys. Res.-Atmos., 108, 4501, https://doi.org/10.1029/2003jd003685, 2003.
- DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.: Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, Anal. Chem., 78, 8281–8289, https://doi.org/10.1021/ac061249n, 2006.
- Di Carlo, P., Aruffo, E., Busilacchio, M., Giammaria, F., Dari-Salisburgo, C., Biancofiore, F., Visconti, G., Lee, J., Moller, S., Reeves, C. E., Bauguitte, S., Forster, G., Jones, R. L., and Ouyang, B.: Aircraft based four-channel thermal dissociation laser induced fluorescence instrument for simultaneous measurements of NO₂, total peroxy nitrate, total alkyl nitrate, and HNO₃, Atmos. Meas. Tech., 6, 971–980, https://doi.org/10.5194/amt-6-971-2013, 2013.
- Drewnick, F., Hings, S. S., DeCarlo, P., Jayne, J. T., Gonin, M., Fuhrer, K., Weimer, S., Jimenez, J. L., Demerjian, K. L., Borrmann, S., and Worsnop, D. R.: A new time-of-flight aerosol mass spectrometer (TOF-AMS) – Instrument description and first field deployment, Aerosol Sci. Tech., 39, 637–658, https://doi.org/10.1080/02786820500182040, 2005.
- Dubinin, M. M., Andreeva, G. A., Vartapetyan, R. S., Vnukov, S. P., Nikolaev, K. M., Polyakov, N. S., Seregina, N. I., and Fedoseev, D. V.: Adsorption of Water and the Micropore Structures of Carbon Adsorbents .5. Pore Structure Parameters of Thermally Treated Carbon Adsorbents and the Adsorption of Water-Vapor on These Materials, B. Acad. Sci. Ussr. Ch.+, 31, 2133– 2137, https://doi.org/10.1007/Bf00958379, 1982.
- Dulitz, K., Amedro, D., Dillon, T. J., Pozzer, A., and Crowley, J. N.: Temperature-(208–318 K) and pressure-(18–696 Torr) dependent rate coefficients for the reaction between OH and HNO₃, Atmos. Chem. Phys., 18, 2381–2394, https://doi.org/10.5194/acp-18-2381-2018, 2018.
- Fahey, D. W., Eubank, C. S., Hubler, G., and Fehsenfeld, F. C.: Evaluation of a Catalytic Reduction Technique for the Measurement of Total Reactive Odd-Nitrogen Noy in the Atmosphere, J. Atmos. Chem., 3, 435–468, https://doi.org/10.1007/Bf00053871, 1985.
- Finlayson-Pitts, B. J., Ezell, M. J., and Pitts, J. N. J.: Formation of chemically active chlorine compounds by reactions of atmospheric NaCl particles with gaseous N₂O₅ and ClONO₂, Nature, 337, 241–244, 1989.

- Fischer, H., Waibel, A. E., Welling, M., Wienhold, F. G., Zenker, T., Crutzen, P. J., Arnold, F., Burger, V., Schneider, J., Bregman, A., Lelieveld, J., and Siegmund, P. C.: Observations of high concentrations of total reactive nitrogen (NO_y) and nitric acid (HNO₃) in the lower Arctic stratosphere during the stratosphere-troposphere experiment by aircraft measurements (STREAM) II campaign in February 1995, J. Geophys. Res.-Atmos., 102, 23559–23571, https://doi.org/10.1029/97jd02012, 1997.
- Friedrich, N.: Measurement of NO_x and NO_y with a thermal dissociation cavity ring-down spectrometer (TD-CRDS): Instrument characterisation and first deployment [Data set], Zenodo, https://doi.org/10.5281/zenodo.4134659, 2020.
- Fry, J. L., Draper, D. C., Zarzana, K. J., Campuzano-Jost, P., Day, D. A., Jimenez, J. L., Brown, S. S., Cohen, R. C., Kaser, L., Hansel, A., Cappellin, L., Karl, T., Hodzic Roux, A., Turnipseed, A., Cantrell, C., Lefer, B. L., and Grossberg, N.: Observations of gas- and aerosol-phase organic nitrates at BEACHON-RoMBAS 2011, Atmos. Chem. Phys., 13, 8585– 8605, https://doi.org/10.5194/acp-13-8585-2013, 2013.
- Fuchs, H., Dube, W. P., Lerner, B. M., Wagner, N. L., Williams, E. J., and Brown, S. S.: A Sensitive and Versatile Detector for Atmospheric NO₂ and NO_x Based on Blue Diode Laser Cavity Ring-Down Spectroscopy, Environ. Sci. Technol., 43, 7831– 7836, https://doi.org/10.1021/es902067h, 2009.
- Fuchs, H., Ball, S. M., Bohn, B., Brauers, T., Cohen, R. C., Dorn, H.-P., Dubé, W. P., Fry, J. L., Häseler, R., Heitmann, U., Jones, R. L., Kleffmann, J., Mentel, T. F., Müsgen, P., Rohrer, F., Rollins, A. W., Ruth, A. A., Kiendler-Scharr, A., Schlosser, E., Shillings, A. J. L., Tillmann, R., Varma, R. M., Venables, D. S., Villena Tapia, G., Wahner, A., Wegener, R., Wooldridge, P. J., and Brown, S. S.: Intercomparison of measurements of NO₂ concentrations in the atmosphere simulation chamber SAPHIR during the NO3Comp campaign, Atmos. Meas. Tech., 3, 21–37, https://doi.org/10.5194/amt-3-21-2010, 2010.
- Gaffney, J. S., Fajer, R., and Senum, G. I.: An Improved Procedure for High-Purity Gaseous Peroxyacyl Nitrate Production – Use of Heavy Lipid Solvents, Atmos. Environ., 18, 215–218, https://doi.org/10.1016/0004-6981(84)90245-2, 1984.
- Gao, X. A., Liu, S. J., Zhang, Y., Luo, Z. Y., Ni, M. J., and Cen, K. F.: Adsorption and reduction of NO₂ over activated carbon at low temperature, Fuel Process. Technol., 92, 139–146, https://doi.org/10.1016/j.fuproc.2010.09.017, 2011.
- Garner, N. M., Matchett, L. C., and Osthoff, H. D.: Quantification of Non-refractory Aerosol Nitrate in Ambient Air by Thermal Dissociation Cavity Ring-Down Spectroscopy, Environ. Sci. Technol., 54, 9864–9861, https://doi.org/10.1021/acs.est.0c01156, 2020.
- Glänzer, K. and Troe, J.: Thermal Decomposition of Nitrocompounds in Shock Waves. IV: Decomposition of Nitric Acid, Berichte der Bunsengesellschaft für physikalische Chemie, 78, 71–76, https://doi.org/10.1002/bbpc.19740780112, 1974.
- Hosaynali Beygi, Z., Fischer, H., Harder, H. D., Martinez, M., Sander, R., Williams, J., Brookes, D. M., Monks, P. S., and Lelieveld, J.: Oxidation photochemistry in the Southern Atlantic boundary layer: unexpected deviations of photochemical steady state, Atmos. Chem. Phys., 11, 8497–8513, https://doi.org/10.5194/acp-11-8497-2011, 2011.

- Huang, W., Saathoff, H., Shen, X. L., Ramisetty, R., Leisner, T., and Mohr, C.: Chemical Characterization of Highly Functionalized Organonitrates Contributing to Night-Time Organic Aerosol Mass Loadings and Particle Growth, Environ. Sci. Technol., 53, 1165–1174, https://doi.org/10.1021/acs.est.8b05826, 2019.
- Ammann, M., Cox, R. A., Crowley, J. N., Herrmann, H., Jenkin, M. E., McNeill, V. F., Mellouki, A., Rossi, M. J., Troe, J., and Wallington, T. J.: Evaluated Kinetic and Photochemical Data for Atmospheric Chemistry, Task Group on Atmospheric Chemical Kinetic Data Evaluation, IUPAC, available at: http://iupac. pole-ether.fr/index.html (last access: 28 October 2020), 2019.
- Ivanov, A. V., Trakhtenberg, S., Bertram, A. K., Gershenzon, Y. M., and Molina, M. J.: OH, HO₂, and ozone gaseous diffusion coefficients, J. Phys. Chem. A, 111, 1632–1637, 2007.
- Javed, U., Kubistin, D., Martinez, M., Pollmann, J., Rudolf, M., Parchatka, U., Reiffs, A., Thieser, J., Schuster, G., Horbanski, M., Pöhler, D., Crowley, J. N., Fischer, H., Lelieveld, J., and Harder, H.: Laser-induced fluorescence-based detection of atmospheric nitrogen dioxide and comparison of different techniques during the PARADE 2011 field campaign, Atmos. Meas. Tech., 12, 1461–1481, https://doi.org/10.5194/amt-12-1461-2019, 2019.
- Kalberer, M., Ammann, M., Arens, F., Gaggeler, H. W., and Baltensperger, U.: Heterogeneous formation of nitrous acid (HONO) on soot aerosol particles, J. Geophys. Res.-Atmos., 104, 13825– 13832, 1999.
- Kim, P., Zheng, Y. J., and Agnihotri, S.: Adsorption equilibrium and kinetics of water vapor in carbon nanotubes and its comparison with activated carbon, Ind. Eng. Chem. Res., 47, 3170–3178, https://doi.org/10.1021/ie0713240, 2008.
- Kleffmann, J., Becker, K. H., Lackhoff, M., and Wiesen, P.: Heterogeneous conversion of NO₂ on carbonaceous surfaces, Phys. Chem. Chem. Phys., 1, 5443–5450, https://doi.org/10.1039/a905545b, 1999.
- Kliner, D. A. V., Daube, B. C., Burley, J. D., and Wofsy, S. C.: Laboratory investigation of the catalytic reduction technique for measurement of atmospheric NO_y, J. Geophys. Res.-Atmos., 102, 10759–10776, https://doi.org/10.1029/96jd03816, 1997.
- Koulouri, E., Saarikoski, S., Theodosi, C., Markaki, Z., Gerasopoulos, E., Kouvarakis, G., Makela, T., Hillamo, R., and Mihalopoulos, N.: Chemical composition and sources of fine and coarse aerosol particles in the Eastern Mediterranean, Atmos. Environ., 42, 6542–6550, https://doi.org/10.1016/j.atmosenv.2008.04.010, 2008.
- Lammel, G. and Perner, D.: The Atmospheric Aerosol as a Source of Nitrous-Acid in the Polluted Atmosphere, J. Aerosol Sci., 19, 1199–1202, https://doi.org/10.1016/0021-8502(88)90135-8, 1988.
- Langford, A. O., Fehsenfeld, F. C., Zachariassen, J., and Schimel, D. S.: Gaseous Ammonia Fluxes and Background Concentrations in Terrestrial Ecosystems of the United States, Global Biogeochem. Cy., 6, 459–483, https://doi.org/10.1029/92gb02123, 1992.
- Lee, B. H., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Hallquist, M., Lee, L., Romer, P., Cohen, R. C., Iyer, S., Kurten, T., Hu, W. W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Xu, L., Ng, N. L., Guo, H. Y., Weber, R. J., Wild, R. J., Brown, S. S., Koss, A., de Gouw, J., Olson, K., Goldstein, A. H., Seco, R., Kim, S., McAvey, K., Shepson, P. B., Starn, T., Baumann, K., Edgerton, E. S., Liu, J. M., Shilling, J. E., Miller, D. O., Brune,

W., Schobesberger, S., D'Ambro, E. L., and Thornton, J. A.: Highly functionalized organic nitrates in the southeast United States: Contribution to secondary organic aerosol and reactive nitrogen budgets, P. Natl. Acad. Sci. USA, 113, 1516–1521, https://doi.org/10.1073/pnas.1508108113, 2016.

- Leser, H., Honninger, G., and Platt, U.: MAX-DOAS measurements of BrO and NO₂ in the marine boundary layer, Geophys. Res. Lett., 30, 1537, https://doi.org/10.1029/2002GL015811, 2003.
- Lightfoot, P. D., Cox, R. A., Crowley, J. N., Destriau, M., Hayman, G. D., Jenkin, M. E., Moortgat, G. K., and Zabel, F.: Organic peroxy radicals – kinetics, spectroscopy and tropospheric chemistry, Atmos. Environ., A-Gen., 26, 1805–1961, 1992.
- Liu, L. M., Tan, S. L., Horikawa, T., Do, D. D., Nicholson, D., and Liu, J. J.: Water adsorption on carbon – A review, Adv. Colloid Interf. Sci., 250, 64–78, https://doi.org/10.1016/j.cis.2017.10.002, 2017.
- Logan, J. A.: Nitrogen-Oxides in the Troposphere Global and Regional Budgets, J. Geophys. Res.-Oceans, 88, 78-5-807, https://doi.org/10.1029/JC088iC15p10785, 1983.
- Malaguti, A., Mircea, M., La Torretta, T. M. G., Telloli, C., Petralia, E., Stracquadanio, M., and Berico, M.: Chemical Composition of Fine and Coarse Aerosol Particles in the Central Mediterranean Area during Dust and Non-Dust Conditions, Aerosol Air Qual. Res., 15, 410–425, https://doi.org/10.4209/aaqr.2014.08.0172, 2015.
- Merten, A., Tschritter, J., and Platt, U.: Design of differential optical absorption spectroscopy long-path telescopes based on fiber optics, Appl. Optics, 50, 738–754, 2011.
- Meusel, H., Tamm, A., Kuhn, U., Wu, D., Leifke, A. L., Fiedler, S., Ruckteschler, N., Yordanova, P., Lang-Yona, N., Pöhlker, M., Lelieveld, J., Hoffmann, T., Pöschl, U., Su, H., Weber, B., and Cheng, Y.: Emission of nitrous acid from soil and biological soil crusts represents an important source of HONO in the remote atmosphere in Cyprus, Atmos. Chem. Phys., 18, 799–813, https://doi.org/10.5194/acp-18-799-2018, 2018.
- Murphy, J. G., Day, D. A., Cleary, P. A., Wooldridge, P. J., and Cohen, R. C.: Observations of the diurnal and seasonal trends in nitrogen oxides in the western Sierra Nevada, Atmos. Chem. Phys., 6, 5321–5338, https://doi.org/10.5194/acp-6-5321-2006, 2006.
- Neuman, J. A., Huey, L. G., Ryerson, T. B., and Fahey, D. W.: Study of inlet materials for sampling atmospheric nitric acid, Environ. Sci. Technol., 33, 1133–1136, 1999.
- Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Jimenez, J. L., Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A., Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate radicals and biogenic volatile organic compounds: oxidation, mechanisms, and organic aerosol, Atmos. Chem. Phys., 17, 2103–2162, https://doi.org/10.5194/acp-17-2103-2017, 2017.
- Notholt, J., Hjorth, J., and Raes, F.: Formation of HNO₂ on Aerosol Surfaces During Foggy Periods in the Presence of NO and NO₂, Atmos. Environ. A-Gen., 26, 211–217, 1992.
- Oswald, R., Behrendt, T., Ermel, M., Wu, D., Su, H., Cheng, Y., Breuninger, C., Moravek, A., Mougin, E., Delon, C., Loubet, B., Pommerening-Roser, A., Sorgel, M., Poschl, U.,

Hoffmann, T., Andreae, M. O., Meixner, F. X., and Trebs, I.: HONO Emissions from Soil Bacteria as a Major Source of Atmospheric Reactive Nitrogen, Science, 341, 1233–1235,

- https://doi.org/10.1126/science.1242266, 2013.
 Palm, B. B., Campuzano-Jost, P., Day, D. A., Ortega, A. M., Fry, J. L., Brown, S. S., Zarzana, K. J., Dube, W., Wagner, N. L., Draper, D. C., Kaser, L., Jud, W., Karl, T., Hansel, A., Gutiérrez-Montes, C., and Jimenez, J. L.: Secondary organic aerosol formation from in situ OH, O₃, and NO₃ oxidation of ambient forest air in an oxidation flow reactor, Atmos. Chem. Phys., 17, 5331–5354, https://doi.org/10.5194/acp-17-5331-2017, 2017.
- Parrish, D. D., Ryerson, T. B., Holloway, J. S., Neuman, J. A., Roberts, J. M., Williams, J., Stroud, C. A., Frost, G. J., Trainer, M., Hubler, G., Fehsenfeld, F. C., Flocke, F., and Weinheimer, A. J.: Fraction and composition of NO_y transported in air masses lofted from the North American continental boundary layer, J. Geophys. Res.-Atmos., 109, D09302, doi10.1029/2003jd004226, 2004.
- Pätz, H.-W., Volz-Thomas, A., Hegglin, M. I., Brunner, D., Fischer, H., and Schmidt, U.: In-situ comparison of the NO_y instruments flown in MOZAIC and SPURT, Atmos. Chem. Phys., 6, 2401– 2410, https://doi.org/10.5194/acp-6-2401-2006, 2006.
- Paul, D. and Osthoff, H. D.: Absolute Measurements of Total Peroxy Nitrate Mixing Ratios by Thermal Dissociation Blue Diode Laser Cavity Ring-Down Spectroscopy, Anal. Chem., 82, 6695– 6703, https://doi.org/10.1021/ac101441z, 2010.
- Paul, D., Furgeson, A., and Osthoff, H. D.: Measurements of total peroxy and alkyl nitrate abundances in laboratorygenerated gas samples by thermal dissociation cavity ring-down spectroscopy, Rev. Sci. Instrum., 80, 114101, https://doi.org/10.1063/1.3258204 2009.
- Perez, I. M., Wooldridge, P. J., and Cohen, R. C.: Laboratory evaluation of a novel thermal dissociation chemiluminescence method for in situ detection of nitrous acid, Atmos. Environ., 41, 3993– 4001, https://doi.org/10.1016/j.atmosenv.2007.01.060, 2007.
- Perring, A. E., Pusede, S. E., and Cohen, R. C.: An observational perspective on the atmospheric impacts of alkyl and multifunctional nitrates on ozone and secondary organic aerosol, Chem. Rev., 113, 5848–5870, https://doi.org/10.1021/cr300520x, 2013.
- Peukert, S. L., Sivaramakrishnan, R., and Michael, J. V.: High temperature shock tube studies on the thermal decomposition of O₃ and the reaction of dimethyl carbonate with O-Atoms, J. Phys. Chem. A, 117, 3729–3738, https://doi.org/10.1021/jp400613p, 2013.
- Pilegaard, K.: Processes regulating nitric oxide emissions from soils, Philos. T. Roy. Soc. B, 368, 20130126, https://doi.org/10.1098/rstb.2013.0126, 2013.
- Platt, U., Perner, D., and Patz, H. W.: Simultaneous Measurement of Atmospheric CH₂O, O₃, and NO₂ by Differential Optical-Absorption, J. Geophys. Res.-Oceans, 84, 6329–6335, https://doi.org/10.1029/JC084iC10p06329, 1979.
- Pohler, D., Vogel, L., Friess, U., and Platt, U.: Observation of halogen species in the Amundsen Gulf, Arctic, by active long-path differential optical absorption spectroscopy, P. Natl. Acad. Sci. USA, 107, 6582–6587, https://doi.org/10.1073/pnas.0912231107, 2010.
- Reed, C., Evans, M. J., Di Carlo, P., Lee, J. D., and Carpenter, L. J.: Interferences in photolytic NO₂ measurements: explanation for

an apparent missing oxidant?, Atmos. Chem. Phys., 16, 4707–4724, https://doi.org/10.5194/acp-16-4707-2016, 2016.

- Roberts, J. M.: The atmospheric chemistry of organic nitrates, Atmos. Environ. A-Gen., 24, 243–287, https://doi.org/10.1016/0960-1686(90)90108-y, 1990.
- Rollins, A. W., Browne, E. C., Min, K.-E., Pusede, S. E., Wooldridge, P. J., Gentner, D. R., Goldstein, A. H., Liu, S., Day, D. A., Russell, L. M., and Cohen, R. C.: Evidence for NO_x Control over Nighttime SOA Formation, Science, 337, 1210–1212, 2012.
- Romer, P. S., Duffey, K. C., Wooldridge, P. J., Allen, H. M., Ayres, B. R., Brown, S. S., Brune, W. H., Crounse, J. D., de Gouw, J., Draper, D. C., Feiner, P. A., Fry, J. L., Goldstein, A. H., Koss, A., Misztal, P. K., Nguyen, T. B., Olson, K., Teng, A. P., Wennberg, P. O., Wild, R. J., Zhang, L., and Cohen, R. C.: The lifetime of nitrogen oxides in an isoprene-dominated forest, Atmos. Chem. Phys., 16, 7623–7637, https://doi.org/10.5194/acp-16-7623-2016, 2016.
- Romer Present, P. S., Zare, A., and Cohen, R. C.: The changing role of organic nitrates in the removal and transport of NO_x, Atmos. Chem. Phys., 20, 267–279, https://doi.org/10.5194/acp-20-267-2020, 2020.
- Rosen, R. S., Wood, E. C., Wooldridge, P. J., Thornton, J. A., Day, D. A., Kuster, W., Williams, E. J., Jobson, B. T., and Cohen, R. C.: Observations of total alkyl nitrates during Texas Air Quality Study 2000: Implications for O₃ and alkyl nitrate photochemistry, J. Geophys. Res.-Atmos., 109, D07303, https://doi.org/10.1029/2003jd004227, 2004.
- Sadanaga, Y., Takaji, R., Ishiyama, A., Nakajima, K., Matsuki, A., and Bandow, H.: Thermal dissociation cavity attenuated phase shift spectroscopy for continuous measurement of total peroxy and organic nitrates in the clean atmosphere, Rev. Sci. Instrum., 87, 074102, https://doi.org/10.1063/1.4958167, 2016.
- Schlesinger, W. H. and Hartley, A. E.: A Global Budget for Atmospheric NH₃, Biogeochemistry, 15, 191–211, 1992.
- Schuster, G., Labazan, I., and Crowley, J. N.: A cavity ring down/cavity enhanced absorption device for measurement of ambient NO₃ and N₂O₅, Atmos. Meas. Tech., 2, 1–13, https://doi.org/10.5194/amt-2-1-2009, 2009.
- Shirahama, N., Moon, S. H., Choi, K. H., Enjoji, T., Kawano, S., Korai, Y., Tanoura, M., and Mochida, I.: Mechanistic study on adsorption and reduction of NO₂ over activated carbon fibers, Carbon, 40, 2605–2611, https://doi.org/10.1016/S0008-6223(02)00190-2, 2002.
- Slusher, D. L., Huey, L. G., Tanner, D. J., Chen, G., Davis, D. D., Buhr, M., Nowak, J. B., Eisele, F. L., Kosciuch, E., Mauldin, R. L., Lefer, B. L., Shetter, R. E., and Dibb, J. E.: Measurements of pernitric acid at the South Pole during ISCAT 2000, Geophys. Res. Lett., 29, 2011, https://doi.org/10.1029/2002gl015703, 2002.
- Sobanski, N., Schuladen, J., Schuster, G., Lelieveld, J., and Crowley, J. N.: A five-channel cavity ring-down spectrometer for the detection of NO₂, NO₃, N₂O₅, total peroxy nitrates and total alkyl nitrates, Atmos. Meas. Tech., 9, 5103–5118, https://doi.org/10.5194/amt-9-5103-2016, 2016.
- Sobanski, N., Thieser, J., Schuladen, J., Sauvage, C., Song, W., Williams, J., Lelieveld, J., and Crowley, J. N.: Day and nighttime formation of organic nitrates at a forested mountain site

in south-west Germany, Atmos. Chem. Phys., 17, 4115–4130, https://doi.org/10.5194/acp-17-4115-2017, 2017.

- Su, H., Cheng, Y. F., Oswald, R., Behrendt, T., Trebs, I., Meixner, F. X., Andreae, M. O., Cheng, P., Zhang, Y., and Poschl, U.: Soil Nitrite as a Source of Atmospheric HONO and OH Radicals, Science, 333, 1616–1618, https://doi.org/10.1126/science.1207687, 2011.
- Tadic, I., Crowley, J. N., Dienhart, D., Eger, P., Harder, H., Hottmann, B., Martinez, M., Parchatka, U., Paris, J.-D., Pozzer, A., Rohloff, R., Schuladen, J., Shenolikar, J., Tauer, S., Lelieveld, J., and Fischer, H.: Net ozone production and its relationship to nitrogen oxides and volatile organic compounds in the marine boundary layer around the Arabian Peninsula, Atmos. Chem. Phys., 20, 6769–6787, https://doi.org/10.5194/acp-20-6769-2020, 2020.
- Talukdar, R. K., Burkholder, J. B., Schmoltner, A. M., Roberts, J. M., Wilson, R. R., and Ravishankara, A. R.: Investigation of the loss processes for peroxyacetyl nitrate in the atmosphere: UV photolysis and reaction with OH, J. Geophys. Res.-Atmos., 100, 14163–14173, https://doi.org/10.1029/95jd00545, 1995.
- Talukdar, R. K., Herndon, S. C., Burkholder, J. B., Roberts, J. M., and Ravishankara, A. R.: Atmospheric fate of several alkyl nitrates .1. Rate coefficients of the reactions alkyl nitrates with isotopically labelled hydroxyl radicals, J. Chem. Soc. Faraday T., 93, 2787–2796, 1997.
- Tang, M. J., Cox, R. A., and Kalberer, M.: Compilation and evaluation of gas phase diffusion coefficients of reactive trace gases in the atmosphere: volume 1. Inorganic compounds, Atmos. Chem. Phys., 14, 9233–9247, https://doi.org/10.5194/acp-14-9233-2014, 2014.
- Thaler, R. D., Mielke, L. H., and Osthoff, H. D.: Quantification of nitryl chloride at part per trillion mixing ratios by thermal dissociation cavity ring-down spectroscopy, Anal. Chem., 83, 2761– 2766, https://doi.org/10.1021/ac200055z, 2011.
- Thieser, J., Schuster, G., Schuladen, J., Phillips, G. J., Reiffs, A., Parchatka, U., Pöhler, D., Lelieveld, J., and Crowley, J. N.: A two-channel thermal dissociation cavity ring-down spectrometer for the detection of ambient NO₂, RO₂NO₂ and RONO₂, Atmos. Meas. Tech., 9, 553–576, https://doi.org/10.5194/amt-9-553-2016, 2016.
- Tingey, D. T., Manning, M., Grothaus, L. C., and Burns, W. F.: Influence of Light and Temperature on Isoprene Emission Rates from Live Oak, Physiol. Plant., 47, 112–118, https://doi.org/10.1111/j.1399-3054.1979.tb03200.x, 1979.
- Tsang, W. and Herron, J. T.: Chemical kinetic data base for propellant combustion. I. Reactions involving NO, NO₂, HNO, HNO₂, HCN and N₂O, J. Phys. Chem. Ref. Data, 20, 609–663, 1991.
- Vandaele, A. C., Hermans, C., Fally, S., Carleer, M., Colin, R., Merienne, M. F., Jenouvrier, A., and Coquart, B.: Highresolution Fourier transform measurement of the NO₂ visible and near-infrared absorption cross sections: Temperature and pressure effects, J. Geophys. Res.-Atmos., 107, 4348, https://doi.org/10.1029/2001JD000971, 2002.
- von der Weiden, S.-L., Drewnick, F., and Borrmann, S.: Particle Loss Calculator – a new software tool for the assessment of the performance of aerosol inlet systems, Atmos. Meas. Tech., 2, 479–494, https://doi.org/10.5194/amt-2-479-2009, 2009.
- Wang, S. S., Nan, J. L., Shi, C. Z., Fu, Q. Y., Gao, S., Wang, D. F., Cui, H. X., Saiz-Lopez, A., and Zhou, B.: Atmo-

spheric ammonia and its impacts on regional air quality over the megacity of Shanghai, China, Sci. Rep.-UK, 5, 15842, https://doi.org/10.1038/srep15842, 2015.

- Wild, R. J., Edwards, P. M., Dube, W. P., Baumann, K., Edgerton, E. S., Quinn, P. K., Roberts, J. M., Rollins, A. W., Veres, P. R., Warneke, C., Williams, E. J., Yuan, B., and Brown, S. S.: A measurement of total reactive nitrogen, NO_y, together with NO₂, NO, and O₃ via cavity ring-down spectroscopy, Environ. Sci. Technol., 48, 9609–9615, https://doi.org/10.1021/es501896w, 2014.
- Williams, E. J., Baumann, K., Roberts, J. M., Bertman, S. B., Norton, R. B., Fehsenfeld, F. C., Springston, S. R., Nunnermacker, L. J., Newman, L., Olszyna, K., Meagher, J., Hartsell, B., Edgerton, E., Pearson, J. R., and Rodgers, M. O.: Intercomparison of ground-based NO_y measurement techniques, J. Geophys. Res.-Atmos., 103, 22261–22280, https://doi.org/10.1029/98jd00074, 1998.
- Wollenhaupt, M., Carl, S. A., Horowitz, A., and Crowley, J. N.: Rate coefficients for reaction of OH with acetone between 202 and 395 K, J. Phys. Chem. 104, 2695–2705, 2000.
- Womack, C. C., Neuman, J. A., Veres, P. R., Eilerman, S. J., Brock, C. A., Decker, Z. C. J., Zarzana, K. J., Dube, W. P., Wild, R. J., Wooldridge, P. J., Cohen, R. C., and Brown, S. S.: Evaluation of the accuracy of thermal dissociation CRDS and LIF techniques for atmospheric measurement of reactive nitrogen species, Atmos. Meas. Tech., 10, 1911–1926, https://doi.org/10.5194/amt-10-1911-2017, 2017.
- Wooldridge, P. J., Perring, A. E., Bertram, T. H., Flocke, F. M., Roberts, J. M., Singh, H. B., Huey, L. G., Thornton, J. A., Wolfe, G. M., Murphy, J. G., Fry, J. L., Rollins, A. W., LaFranchi, B. W., and Cohen, R. C.: Total Peroxy Nitrates (ΣPNs) in the atmosphere: the Thermal Dissociation-Laser Induced Fluorescence (TD-LIF) technique and comparisons to speciated PAN measurements, Atmos. Meas. Tech., 3, 593–607, https://doi.org/10.5194/amt-3-593-2010, 2010.
- Xu, L., Suresh, S., Guo, H., Weber, R. J., and Ng, N. L.: Aerosol characterization over the southeastern United States using high-resolution aerosol mass spectrometry: spatial and seasonal variation of aerosol composition and sources with a focus on organic nitrates, Atmos. Chem. Phys., 15, 7307–7336, https://doi.org/10.5194/acp-15-7307-2015, 2015.
- Yamamoto, N., Kabeya, N., Onodera, M., Takahahi, S., Komori, Y., Nakazuka, E., and Shirai, T.: Seasonal-Variation of Atmospheric Ammonia and Particulate Ammonium Concentrations in the Urban Atmosphere of Yokohama over a 5-Year Period, Atmos. Environ., 22, 2621–2623, https://doi.org/10.1016/0004-6981(88)90498-2, 1988.
- Yao, X. and Zhang, L.: Trends in atmospheric ammonia at urban, rural, and remote sites across North America, Atmos. Chem. Phys., 16, 11465–11475, https://doi.org/10.5194/acp-16-11465-2016, 2016.
- York, D.: Least-Squares Fitting of a Straight Line, Can. J. Phys., 44, 1079–1086, https://doi.org/10.1139/p66-090, 1966.
- Zare, A., Romer, P. S., Nguyen, T., Keutsch, F. N., Skog, K., and Cohen, R. C.: A comprehensive organic nitrate chemistry: insights into the lifetime of atmospheric organic nitrates, Atmos. Chem. Phys., 18, 15419–15436, https://doi.org/10.5194/acp-18-15419-2018, 2018.

N. Friedrich et al.: Measurement of NO_x and NO_y with a TD-CRDS

- Zasypkin, A. Y., Grigoreva, V. M., Korchak, V. N., and Gershenson, Y. M.: A formula for summing of kinetic resistances for mobile and stationary media: 1. Cylindrical reactor, Kinet. Catal.+, 38, 772–781, 1997.
- Zenker, T., Fischer, H., Nikitas, C., Parchatka, U., Harris, G. W., Mihelcic, D., Müsgen, P., Pätz, H. W., Schultz, M., Volz-Thomas, A., Schmitt, R., Behmann, T., Weißenmayer, M., and Burrows, J. P.: Intercomparison of NO, NO₂, NO_y, O₃, and RO_x measurements during the Oxidizing Capacity of the Tropospheric Atmosphere (OCTA) campaign 1993 at Izaña, J. Geophys. Res.-Atmos., 103, 13615–13634, https://doi.org/10.1029/97jd03739, 1998.
- Zhang, W. J., Bagreev, A., and Rasouli, F.: Reaction of NO₂ with activated carbon at ambient temperature, Ind. Eng. Chem. Res., 47, 4358–4362, https://doi.org/10.1021/ie800249s, 2008.