



Mie–Raman–fluorescence lidar observations of aerosols during pollen season in the north of France

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Abstract. The multiwavelength Mie–Raman–fluorescence lidar of the University of Lille has the capability to measure three aerosol backscattering coefficients, two extinction coefficients and three linear depolarization ratios, together with fluorescence backscattering at 466 nm. It was used to characterize aerosols during the pollen season in the north of France for the period March–June 2020. The results of observations demonstrate that the presence of pollen grains in aerosol mixture leads to an increase in the depolarization ratio. Moreover, the depolarization ratio exhibits a strong spectral dependence increasing with wavelength, which is expected for the mixture containing fine background aerosols with low depolarization and strongly depolarizing pollen grains. A high depolarization ratio correlates with the enhancement of the fluorescence backscattering, corroborating the presence of pollen grains. Obtained results demonstrate that simultaneous measurements of particle depolarization and fluorescence allows for the separation of dust, smoke particles and aerosol mixtures containing the pollen grains.

by causing allergy-related diseases and contribute to cloud formation by acting as giant cloud condensation nuclei (CCN) (Diehl et al., 2001; Pope, 2010; Griffiths et al., 2012; D’Amato et al., 2014; Steiner et al., 2015; Lake et al., 2017; Mack et al., 2020). To investigate the processes of pollen transport and dispersion, information about the vertical distribution of pollen grains is needed, and this information can be obtained from lidar measurements. Pollen grains are large irregularly shaped particles of complicated morphology (Frenguelli, 2003; Katifori et al., 2010) causing strong depolarization of backscattered laser radiation, which provides a basis for their identification. The first profiling of pollen with depolarization lidar was reported by Sassen (2008, 2011). His measurements over Alaska revealed that the linear depolarization ratio of birch pollen plumes at 694 nm can exceed 30 %. Further studies of pollen with elastic backscatter lidar at 532 nm were reported by Noh et al. (2013a, b) and by Sicard et al. (2016). Their measurements confirmed the high depolarization ratio of pollen grains (particle depolarization ratios as high as 43 % were observed for aerosol mixtures containing *Platanus* and *Pinus* pollen). Moreover, pollen grain backscattering demonstrates a strong diurnal cycle, being the highest near noon. The use of multiwavelength observations increases the capability of the lidar technique for aerosol characterization. In recent studies by Bohlmann et al. (2019) and Shang et al. (2020), measurements performed with Polly^{XT} lidar allowed for the estimation of mean values of lidar ratios (about 45 and 55 sr at 355 and 532 nm, re-

1 Introduction

Pollen grains represent a significant fraction of primary biological particles emitted from the biosphere into the atmosphere in certain seasons and locations (Fröhlich-Nowoisky et al., 2016). There has been a growing interest in pollen study in recent years because they can affect human health

spectively, for birch pollen grains). Lidar measurement of the pollen depolarization ratio at 355, 532 and 1565 nm was also reported in the recent work of Bohlmann et al. (2021), revealing a strong decrease in the depolarization ratio at 355 nm.

Atmospheric biological particles efficiently produce wide-band fluorescence emissions when exposed to UV radiation (Pöhlker et al., 2012; Pan, 2015; Miyakawa et al., 2015), which offers an opportunity for monitoring them with fluorescence lidars. Nowadays, lidar spectrometers based on multi-anode photomultipliers allow simultaneous detection of fluorescence backscattering in 32 spectral bins (Sugimoto et al., 2012; Reichardt, 2014; Reichardt et al., 2017; Saito et al., 2018; Richardson et al., 2019). In particular, such a lidar spectrometer was used in the recent work of Saito et al. (2018) for remote measurement of the fluorescence spectrum of atmospheric pollen grains. The results demonstrate that, for the 355 nm stimulating wavelength, the fluorescence spectra of different pollen grains have maxima in the 400–600 nm range and a intensity peak at around 460 nm.

To achieve the highest sensitivity of fluorescence detection, in many tasks it is preferable to use single-channel monitoring, whereby a part of the fluorescence spectrum is selected with a wideband interference filter (Immler et al., 2005; Rao et al., 2018; Li et al., 2019). In our recent publication (Veselovskii et al., 2020a) we reported the results obtained from a modified Mie–Raman lidar system at the Laboratoire d’optique Atmosphérique with one additional fluorescence channel at 466 nm. Such an approach has proved high sensitivity, allowing for the detection of fluorescence signals from weak aerosol layers and the calculation of the fluorescence backscattering coefficient from the ratio of fluorescence and nitrogen Raman backscatters, thus making it potentially attractive for pollen monitoring.

In the present research we combine the capability of multiwavelength Mie–Raman lidar for providing three backscattering coefficients, two aerosol extinction coefficients and the linear depolarization ratio at three wavelengths with single-channel fluorescence measurements for characterization of aerosol mixtures containing pollen grains. The measurements reported were performed during the March–June 2020 period at the Lille Atmospheric Observation Platform (<https://www-loa.univ-lille1.fr/observations/plateformes.html?p=apropos>, last access: 30 June 2021) hosted by the Laboratoire d’Optique Atmosphérique, University of Lille, Hauts-de-France region.

2 Instrumentation

2.1 Mie–Raman–fluorescence lidar

The measurements were performed using the LILAS (Lille Lidar AtmosphereS) system – a multiwavelength Mie–Raman lidar based on a tripled Nd:YAG laser with a 20 Hz repetition rate and pulse energy of 70 mJ at 355 nm. The

backscattered light is collected by a 40 cm aperture Newtonian telescope. The full geometrical overlap of the laser beam and the telescope field of view (FOV) is achieved at approximately 1000 m. To obtain the information about particles at lower altitudes, some of the measurements were performed at an angle of 30° to the horizon. No overlap correction of lidar signals was performed. The system is designed for simultaneous detection of elastic and Raman backscattering, allowing the so-called $3\beta + 2\alpha + 3\delta$ data configuration, including three particle backscattering coefficients (β_{355} , β_{532} , β_{1064}), two extinction coefficients (α_{355} , α_{532}) and three particle depolarization ratios (δ_{355} , δ_{532} , δ_{1064}). The particle depolarization ratio δ , determined as the ratio of cross- and co-polarized components of the particle backscattering coefficient, was calculated and calibrated the same way as described in Freudenthaler et al. (2009). The description of the system can be found in the recent publication of Hu et al. (2019).

To perform fluorescence lidar measurements, the water vapor Raman channel at 408 nm was replaced by a fluorescence channel, whose spectrum is captured by a wideband filter centered at 466 nm and of 44 nm width (Veselovskii et al., 2020a). The fluorescence measurements were performed during nighttime only. The aerosol extinction and backscattering coefficients at 355 and 532 nm were calculated from Mie–Raman observations (Ansmann et al., 1992), while β_{1064} was derived by the Klett method (Klett, 1985). The fluorescence backscattering coefficient β_F is calculated from the ratio of fluorescence and nitrogen Raman backscattering, as described in Veselovskii et al. (2020a). This approach allows for the evaluation of the absolute values of β_F if the relative sensitivity of the channels is calibrated and the nitrogen Raman scattering cross section is known. We estimate the corresponding uncertainty to be below 50 %. Parameters of detectors were not changed during the campaign, so the uncertainty of relative variations of β_F was significantly lower and was determined by the statistical errors of fluorescence measurements.

To characterize the efficiency of the fluorescence with respect to elastic scattering, the fluorescence capacity,

$$G_F = \frac{\beta_F}{\beta_{532}}, \quad (1)$$

is also used. This parameter depends on the relative humidity (RH), so information about RH is important for data analysis. Radiosonde measurements are used to monitor water vapor, as the water vapor channel is replaced by the fluorescence channel in the current lidar configuration. The closest available radiosonde data are from Herstmonceux (UK) and Beauvecchain (Belgium) stations, located 160 and 80 km away from the observation site, respectively. These radiosonde data are not collocated with the lidar measurements, so only qualitative analysis of humidification effects was possible.

2.2 Pollen in situ sampling

Airborne pollen grains and spores were collected by a Hirst-type volumetric sampler (Lanzoni VPPS 2000). The pollen sampler was located on the campus of the University of Lille (France) on the rooftop of a 20 m high building where the lidar instrument was operated. Ambient air was sampled at a 10 L min^{-1} flow rate, allowing the impaction of pollen and spores on an adhesive strip mounted on a rotating clockwork-driven drum. The impaction surface moves at 2 mm h^{-1} behind the entrance slit, allowing a temporal resolution of 2 h. The adhesive strip was substituted every 7 d after a full rotation of the drum, which is split into seven parts, each corresponding to a day of monitoring. And then they are fixed on a microscope glass slide with gelatin and fuchsine dye. Pollen taxa were identified by light microscopy on the basis of their characteristic shape and size. Airborne pollen concentrations were expressed as a daily and dual hourly number of pollen grains per cubic meter of air.

Figure 1 shows the most abundant pollen taxa for the period from March to June 2020 in Lille. These include *Betula* (54.8 % of total pollen taxa over the period), *Fraxinus* (8.2 %), *Quercus* (5.8 %), *Urticaceae* (4.6 %), *Salix* (4.5 %) and *Cupressaceae* (4.1 %). The same figure also shows the fluorescence backscattering β_F measured by lidar. The results presented are obtained by averaging all available data during the night, and the maximal values in the 500–1000 m height range are shown. The highest fluorescence was observed at the end of March when ash (*Fraxinus*) is the main pollinator. The period of intense birch (*Betula*) pollination (3–15 April 2020) also correlates with high β_F . Strong fluorescence observed for the 5–10 May and 28 May–2 June periods can be due to grass (*Poaceae*) pollen contribution. By the end of June, β_F decreases and becomes comparable with the fluorescence backscattering of background aerosol. From Fig. 1, we can conclude that there is no direct correlation between in situ and fluorescence lidar measurements; thus, pollen grains observed in the boundary layer by the lidar are probably transported from other regions. Comparing lidar and in situ observations, we should also keep in mind that maximum of pollen emission occurs near noon, while lidar measurements were performed in the night.

3 Discrimination of pollen from other types of aerosol

3.1 Specific features of pollen-containing aerosol mixture

In contrast to the observations performed over Alaska (Sassen, 2008, 2011) and in Kuopio, Finland (Bohlmann et al., 2019), where the pollen concentration was high due to boreal forests surrounding, the pollen loading in the north of France is significantly lower. Long-term lidar and sun photometer observations performed at the University of Lille

demonstrate that local aerosol is mainly of continental type, with a predominance of fine-mode particles and a low depolarization ratio. The emission of large pollen grains should lead to strong spectral dependence of the depolarization ratio because the backscattering at 1064 nm is less sensitive to fine background particles than at shorter wavelengths; thus, the particle depolarization ratio at 1064 nm (δ_{1064}) should be more sensitive to the presence of pollen grains compared to δ_{355} and δ_{532} . The particle depolarization ratio δ of the mixture containing background aerosol (b) and pollen (p), with corresponding depolarization ratios δ^b and δ^p , can be calculated as

$$\delta = \frac{\left(\frac{\delta^p}{1+\delta^p}\right)\beta^p + \left(\frac{\delta^b}{1+\delta^b}\right)\beta^b}{\frac{\beta^p}{1+\delta^p} + \frac{\beta^b}{1+\delta^b}}. \quad (2)$$

Here total backscattering $\beta = \beta^b + \beta^p$.

To estimate the dependence of depolarization of the aerosol mixture on the contribution of pollen to the backscattering coefficient $\frac{\beta_{532}^p}{\beta_{532}}$ at 532 nm, a simplified simulation was performed. Assuming that the depolarization ratios of pollen and background aerosol are spectrally independent and that $\delta^p = 30\%$ when $\delta^b = 3\%$, the mixture depolarization ratios δ_{355} , δ_{532} and δ_{1064} were calculated as a function of $\frac{\beta_{532}^p}{\beta_{532}}$ using Eq. (2). For pollen the backscattering Ångström exponents (BAE) $A_{355/532}^\beta = A_{532/1064}^\beta = 0$ were used. The backscattering Ångström exponents of background aerosol are assumed to be the same for both wavelength pairs ($A_{355/532}^\beta = A_{532/1064}^\beta$), and computations were performed for the values of BAE $A^\beta = 1.0, 1.5$ and 2.0 . Results of the simulation are shown in Fig. 2. For a low pollen contribution the depolarization ratio δ_{1064} increases faster with $\frac{\beta_{532}^p}{\beta_{532}}$ compared to δ_{532} and δ_{355} , and the slope of this increase depends on the BAE value. Spectral properties of the real mixture can be more complicated due to possible spectral dependence of both δ^p and δ^b . Information on laboratory-measured spectral dependence of depolarization ratios of pollen is rare. Cao et al. (2010) measured the linear depolarization ratio of several types of pollen in a chamber at 355, 532 and 1064 nm wavelengths. The results demonstrate strong variation of spectral dependence for different taxa, and for most of the samples δ_{532} exceeded both δ_{355} and δ_{1064} . In particular, for birch pollen, depolarization ratios at 355, 532 and 1064 nm are 8 %, 33 % and 28 %. Using the depolarization ratios from Cao et al. (2010) in the analysis, we should also keep in mind that measurements in the chamber were performed at low RH, and depolarization ratios at higher RH may be different.

The Ångström exponents $A_{355/532}^\beta$ and $A_{532/1064}^\beta$ depend on the particle size and the refractive index in different ways (Veselovskii et al., 2015), so their values can be different. Thus, spectral dependence of the depolarization ratio of an aerosol–pollen mixture will differ from the simplified mod-

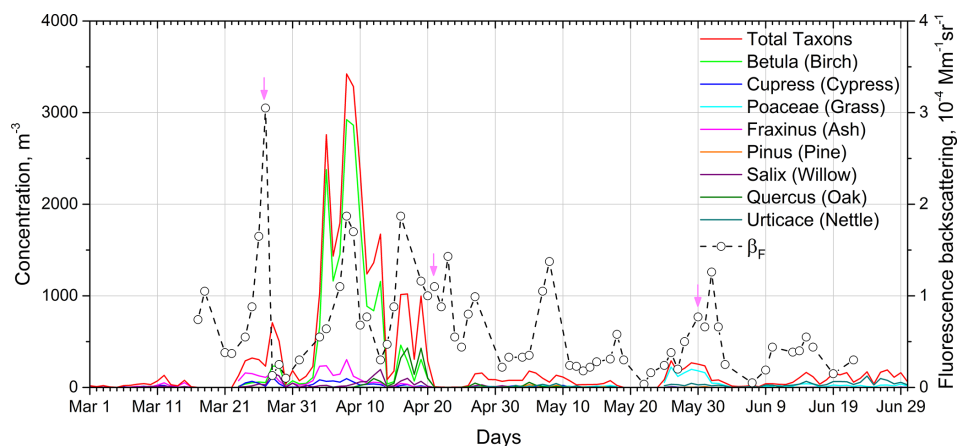


Figure 1. Daily concentration of the most abundant pollen taxa for the period March–June 2020 in Lille from in situ measurements on the rooftop. Open symbols show fluorescence backscattering β_F measured by lidar. Lidar measurements are averaged overnight, and the maximal value in the 500–1000 m range is shown. Vertical arrows mark the sessions, which are further analyzed in Figs. 8, 10 and 11.

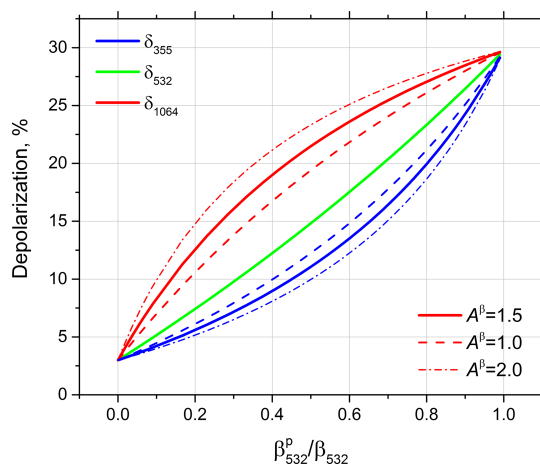


Figure 2. Depolarization ratios at 355, 532 and 1064 nm as a function of pollen contribution to the total backscattering coefficient $\frac{\beta_{532}^p}{\beta_{532}}$. Depolarization ratios of pollen (δ^p) and background aerosol (δ^b) are assumed to be spectrally independent ($\delta^p = 30\%$ and $\delta^b = 3\%$). The backscattering coefficient of pollen is spectrally independent. The backscattering Ångström exponents of background aerosol were assumed to be the same for both pairs of wavelengths ($A_{355/532}^\beta = A_{532/1064}^\beta$), and results are shown for the values $A^\beta = 1.0, 1.5$ and 2.0 .

eling shown in Fig. 2. However, for a moderate contribution of pollen to the total backscattering ($\frac{\beta_{532}^p}{\beta_{532}}$ below ≈ 0.5), the depolarization ratio at 1064 nm should be higher than that at shorter wavelengths. Hence, an increase in the particle depolarization ratio with wavelength can be an indication of the presence of large, irregularly shaped pollen grains in an aerosol mixture.

The presence of pollen should also lead to a decrease in the extinction and backscattering Ångström exponents. The extinction Ångström exponent (EAE) depends mainly on particle size, while BAE is also sensitive to the particle complex refractive index and shape; thus, the measured profiles of EAE and BAE can present significant differences (Veselovskii et al., 2015, 2020b). In our study we analyze the EAE and BAE for the wavelength pair 355 nm and 532 nm ($A_{355/532}^\alpha$ and $A_{355/532}^\beta$) only because the extinction coefficient and backscattering coefficient involved are calculated from Mie–Raman observations.

When analyzing Mie–Raman–fluorescence lidar measurements of pollen-containing aerosol mixtures, numerous factors should be taken into account. These factors include the fluorescence of background aerosol and other non-pollen aerosols that have strong fluorescence capacity: for example, smoke particles. Dust particles can contribute to the increase in the depolarization ratio, and, finally, hygroscopic growth can modify the particle parameters. All these factors will be considered in the following sections.

3.2 Characteristics of background aerosol over the observation site

Long-term observations at the University of Lille demonstrate that aerosol over the observation site is mainly of continental type with a predominance of fine-mode particles. Typical vertical profiles of the background aerosol parameters, observed on 3 June 2020, are given in Fig. 3, showing aerosol elastic and fluorescence backscattering coefficients, lidar ratios, Ångström exponents and depolarization ratios at three wavelengths. The RH from Beauvechain (Belgium) radiosonde observations was below 50 % in the height range considered. Particle depolarization ratios at all three wavelengths are below 7 %, indicating that the contribution of pollen to the total backscattering was low. This agrees with

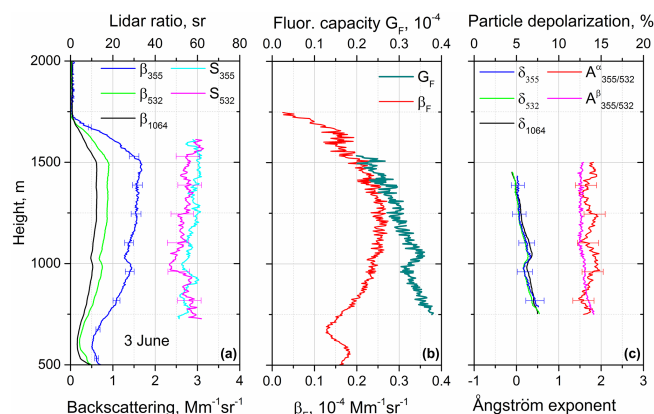


Figure 3. Measurements in the condition of background aerosol predominance. Vertical profiles of (a) backscattering coefficients β_{355} , β_{532} and β_{1064} as well as lidar ratios S_{355} and S_{532} . (b) Fluorescence backscattering β_F and fluorescence capacity $G_F = \beta_F/\beta_{532}$. (c) Particle linear depolarization ratios δ_{355} , δ_{532} and δ_{1064} together with extinction and backscattering Ångström exponents $A_{355/532}^\alpha$ and $A_{355/532}^\beta$ on 3 June 2020 for 20:30–23:00 UTC. Measurements were performed at 30° to the horizon.

the low values of pollen concentration provided by in situ measurements (Fig. 1). The lidar ratios at both wavelengths (S_{355} , S_{532}) are close, varying in the 50–60 sr range, and the fluorescence capacity G_F is below 0.35×10^{-4} . The EAE and BAE ($A_{355/532}^\alpha$, $A_{355/532}^\beta$) are in the 1.5–2.0 range. The presence of pollen should lead to a deviation of the particle intensive parameters, such as the fluorescence capacity, depolarization ratio, EAE and BAE, from the typical values of background aerosol.

3.3 Identification of the smoke particles

During the observation period the smoke elevated layers transported over the Atlantic were frequently detected. Smoke particles are characterized by a high fluorescence cross section (Reichardt et al., 2017; Veselovskii et al., 2020a) and can interfere with pollen fluorescence measurements. The temporal evolution of the range-corrected lidar signal, volume depolarization ratio at 1064 nm and fluorescence backscattering for the smoke episode on the night of 23–24 June 2020 are shown in Fig. 4. During the night the smoke layer, with a low depolarization and high fluorescence, is observed at approximately 5000 m of height. Back trajectories (not shown) indicate that the layer is transported from Canada. Vertical profiles of the particle parameters for this episode are shown in Fig. 5. The lidar ratio is about 50 sr at 355 nm, while the lidar ratio at 532 nm increases within the smoke layer from 60 to 80 sr. This increase in S_{532} occurs simultaneously with decrease in $A_{355/532}^\alpha$ from 1.5 to 0.75, indicating that the particle size inside the layer grows with height. Higher values of S_{532} with respect to S_{355} are typical for aged smoke (e.g., Müller et al., 2005; Nicolae et

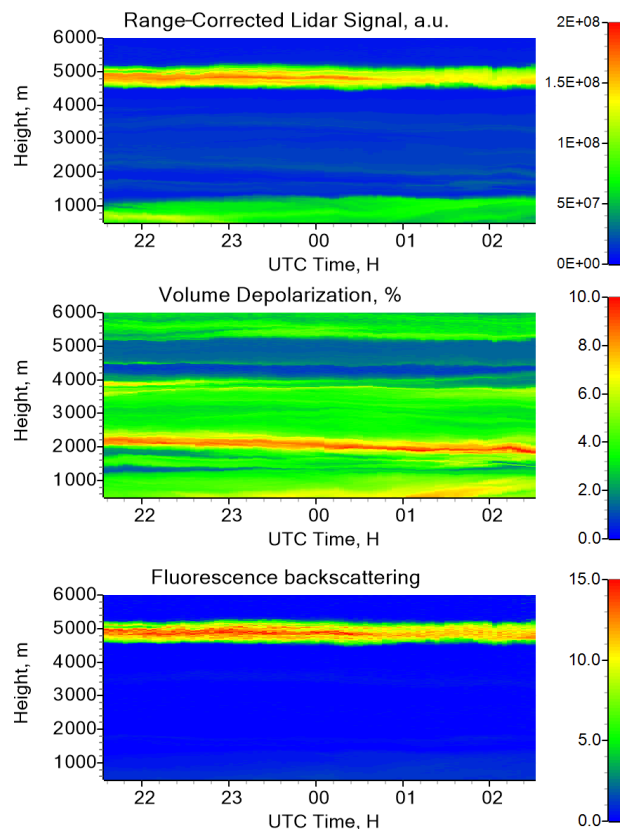


Figure 4. Range-corrected lidar signal at 1064 nm, volume depolarization ratio at 1064 nm and fluorescence backscattering coefficient (in $10^{-4} \text{ Mm}^{-1} \text{ sr}^{-1}$) on 23–24 June 2020.

al., 2013; Hu et al., 2019). The depolarization ratio decreases with wavelength from $\delta_{355} = 10\%$ to $\delta_{1064} = 1.5\%$. A strong spectral dependence of the depolarization ratio and, in particular, low values of δ_{1064} are the features allowing for the identification of smoke layers. We should also recall that an increase in the particle depolarization ratio at 355 nm is more typical for aged smoke layers in the high troposphere (Haarig et al., 2018), though we observed this increase at lower altitudes over Lille during smoke episodes in summer–autumn 2020.

The extinction Ångström exponent $A_{355/532}^\alpha$ in the center of the layer is about 0.75, while $A_{355/532}^\beta$ is about 1.9 and shows no significant variation through the layer. High values of $A_{355/532}^\beta$ compared to $A_{355/532}^\alpha$ represent another feature that will be used for aged smoke discrimination. Smoke fluorescence capacity is high, reaching up to $G_F = 5 \times 10^{-4}$ for the period of observations, and this is one more feature allowing for the separation of smoke from other types of aerosol.

3.4 Identification of the dust particles

The presence of dust particles and pollen in the fine background aerosol leads to some common characteristics in the

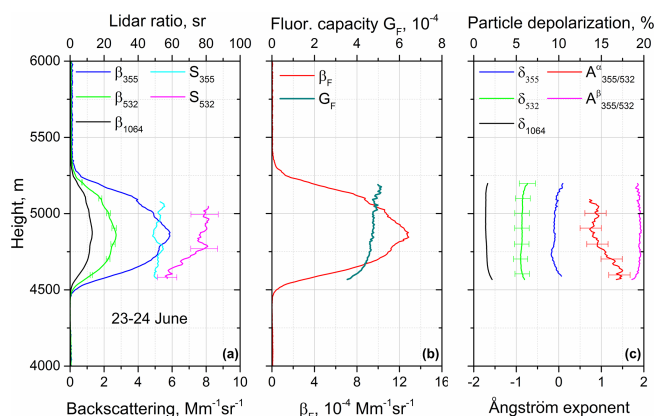


Figure 5. Vertical profiles of (a) backscattering coefficients β_{355} , β_{532} and β_{1064} as well as lidar ratios S_{355} and S_{532} . (b) Fluorescence backscattering coefficient β_F , fluorescence capacity G_F , and (c) particle depolarization ratios δ_{355} , δ_{532} and δ_{1064} together with the extinction and backscattering Ångström exponents $A_{355/532}^\alpha$ and $A_{355/532}^\beta$ on the night of 23–24 June 2020 at 21:30–02:30 UTC.

lidar data, such as decreased Ångström exponents and increased depolarization ratios. However, pollen and dust can be separated by the fluorescence capacity. The vertical profiles of particle parameters during a dust episode on 27 May are shown in Fig. 6. The dust-containing layer extends from 2000 to 7000 m, and the particle depolarization ratios δ_{1064} and δ_{532} in this layer are close to 20 %. These values are lower than the depolarization of pure dust. For example, Freudenthaler et al. (2009) for pure dust provide values of 27 % and 31 % at 1064 and 532 nm wavelengths, respectively; thus, in our case transported dust particles may be mixed with local aerosols. The particle depolarization at 355 nm is not shown in the figure because the scattering ratio in the dust layer was too low to compute δ_{355} reliably. The fluorescence capacity of particles in the dust layer is about 0.1×10^{-4} at 4000 m, which is a factor of 50 lower than G_F of the smoke in Fig. 5. There is also a weak aerosol layer at 1600 m with β_{532} of about $0.035 \text{ Mm}^{-1} \text{ sr}^{-1}$. The fluorescence capacity in this layer is high ($G_F \approx 2.0 \times 10^{-4}$), suggesting that this layer may contain smoke or pollen particles.

3.5 Impact of particle hygroscopic growth

The vertical variation of observed aerosol properties may be a result of particle water uptake, which should be separated from the features related to pollen presence. Figure 7 shows the profiles of the particle parameters for the episode on 15 June 2020 when aerosol hygroscopic growth took place. In the height range 900–1500 m the fluorescence backscattering β_F is stable, while the elastic backscattering β_{532} increases by a factor of 3. Radiosonde measurements in Herstmonceux (UK) in this height range demonstrate an increase in RH from about 75 % to 85 %, while lidar-measured extinc-

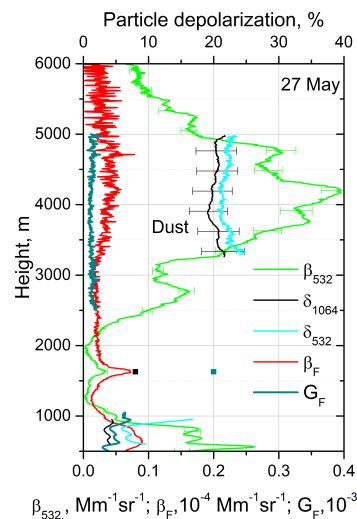


Figure 6. Lidar measurements during a dust episode. Vertical profiles of particle β_{1064} and fluorescence β_F backscattering coefficients, fluorescence capacity G_F , and particle depolarization ratios δ_{1064} and δ_{532} on 27 May 2020 at 21:00–23:00 UTC.

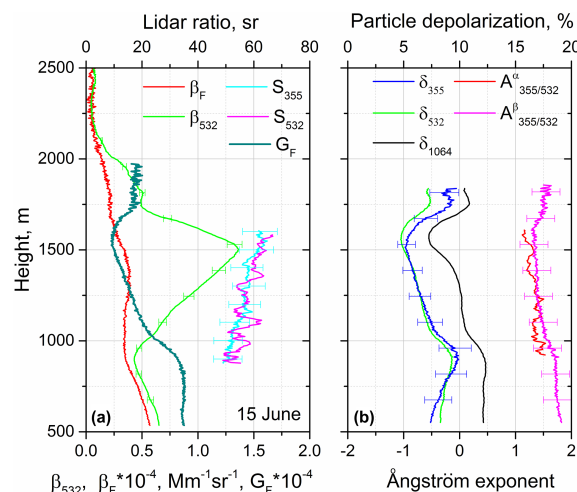


Figure 7. Lidar measurements in the condition of aerosol hygroscopic growth in the 900–1500 m height range. Vertical profiles of (a) particle β_{532} and fluorescence β_F backscattering coefficients, fluorescence capacity G_F , lidar ratios S_{355} and S_{532} , and (b) particle depolarization ratios δ_{355} , δ_{532} and δ_{1064} together with extinction $A_{355/532}^\alpha$ and backscattering $A_{355/532}^\beta$ Ångström exponents measured on 15 June 2020 at 22:00–24:00 UTC.

tion and backscattering Ångström exponents decrease from 1.5 to 1.3, corroborating the occurrence of particle hygroscopic growth. The depolarization ratio δ_{1064} at low altitudes exceeds δ_{355} and δ_{532} , which can be an indication of pollen presence. This is supported by significant fluorescence capacity ($G_F = 0.9 \times 10^{-4}$ at 750 m).

The number of fluorescent particles in the 900–1500 m range does not present significant changes (β_F is stable), so

observed vertical variations, i.e., the decrease in depolarization ratios at all three wavelengths and the increase in lidar ratios S_{355} and S_{532} from 50 to 65 sr, are probably the result of water uptake by the particles. Water uptake does not change the number of fluorescent molecules; however, the fluorescence capacity decreases in the process of hygroscopic growth, so G_F can be a representative parameter of aerosol types only in the condition of low RH.

4 Results of lidar measurements in the presence of pollen

During March–June 2020, we had numerous measurement cases demonstrating features in the profiles of the particle parameters that can be attributed to pollen. For representative cases we have chosen observations with a high depolarization ratio and high fluorescence backscattering. At the same time, we omitted the days with high relative humidity to minimize the impact of hygroscopic growth effects. Below we consider several measurement cases representing different scenarios, in particular the episodes when pollen concentration decreases with height (30–31 May, 1–2 June) and the episodes when pollen grains are well mixed inside the boundary layer (27–28 March and 21 April).

4.1 30–31 May and 1–2 June 2020 observations

The results of lidar measurements during the campaign in many episodes can be interpreted as a decrease in the pollen concentration with height. Vertical profiles of the main particle parameters for two representative cases on the nights of 30–31 May and 1–2 June 2020 are shown in Fig. 8. The atmospheric conditions for these nights were stable, so the profiles presented are averaged over an approximately 5 h interval. The HYSPLIT back-trajectory analysis (Stein et al., 2015) demonstrates that in the 1000–2000 m height range the air masses were transported from northern Europe. At the ground level, grass could be the main pollen contributor for this period, as shown in Fig. 1. On 31 May (at 00:00 UTC) the RH measured by the radiosonde in Herstmonceux (UK) was about 40 % at 500 m, and it increased up to 70 % at 2000 m. On 2 June the RH increased from approximately 40 % to 60 % in the same height range. For both nights the fluorescence backscattering decreases with height, indicating a decrease in the concentration of fluorescent particles (presumably pollen). This decrease in β_F correlates with a decrease in the depolarization ratio at all three wavelengths. Particle depolarization δ_{1064} is the highest (about 15 % at 750 m), while δ_{355} and δ_{532} are significantly lower. Such spectral dependence of the depolarization ratio can be partly due to the contribution of the background aerosol, as follows from the model calculation in Fig. 2. The lidar ratios are available above 1250 m, and for both cases, S_{355} and S_{532} increase with height. It indicates that the lidar ratios of pollen in the two

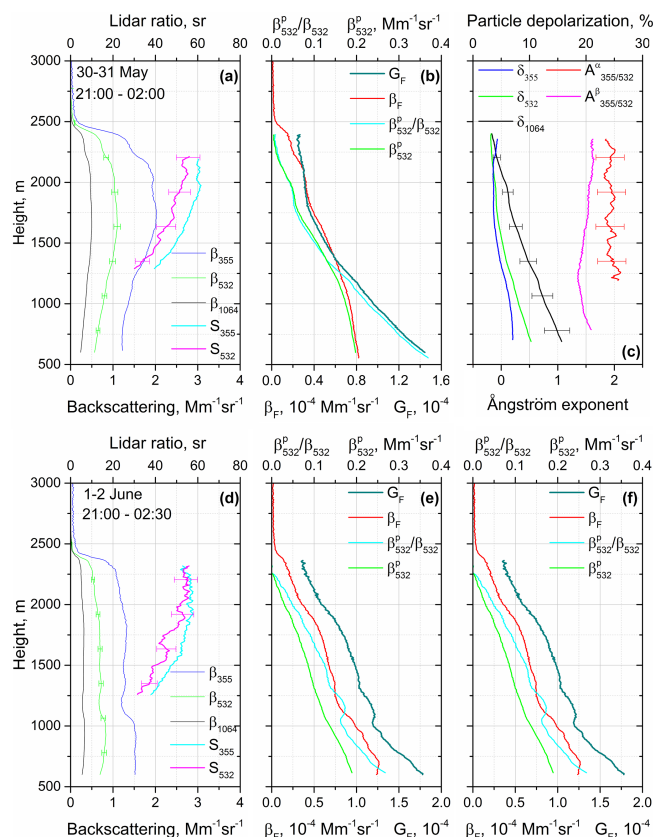


Figure 8. Vertical profiles of (a, d) particle backscattering coefficients β_{355} , β_{532} and β_{1064} as well as lidar ratios S_{355} and S_{532} . (b, e) Fluorescence backscattering coefficient β_F , fluorescence capacity G_F , pollen backscattering coefficient β_{532}^p and its contribution to the total backscattering $\beta_{532}^p/\beta_{532}$. (c, f) Particle depolarization ratios δ_{355} , δ_{532} and δ_{1064} together with extinction $A_{355/532}^\alpha$ and backscattering $A_{355/532}^\beta$ Ångström exponents on (a–c) 30–31 May 2020 at 21:00–02:00 UTC and on (d–f) 1–2 June 2020 at 21:00–02:30 UTC. Profiles of β_{532}^p and $\beta_{532}^p/\beta_{532}$ were computed with the assumption that $\delta_{532}^p = 30\%$. The depolarization ratios of the background aerosol δ_{532}^b are measured or assumed to be 3 % on 30 May and 5 % on 1 May.

considered cases can be quite low: below 40 sr at 355 nm and below 30 sr at 532 nm, considering that the pollen concentration decreases with height, which is inferred from the features of the depolarization ratio and fluorescence backscattering.

The EAE for both nights varied in the 1.75–2.0 range and did not show significant changes with height. The BAE is lower (about 1.5 at 1000 m), and for both nights it shows some increase in the 1250–2250 m range. The BAE, in contrast to EAE, depends strongly on the particle refractive index and shape; thus, it may demonstrate higher sensitivity to the changes in aerosol mixture composition. Recall that backscattering and extinction Ångström exponents are re-

lated as

$$A_{355/532}^{\beta} = A_{355/532}^{\alpha} + \frac{\ln(S_{532}/S_{355})}{\ln(355/532)}. \quad (3)$$

Thus, for $S_{355} > S_{532}$, which has been observed during pollen episodes, the $A_{355/532}^{\beta}$ is lower than $A_{355/532}^{\alpha}$. This is in contrast with smoke episodes in which $S_{355} < S_{532}$ and $A_{355/532}^{\beta} > A_{355/532}^{\alpha}$ (Fig. 5).

If the depolarization ratios of pollen δ^p and background aerosol δ^b are known, the pollen backscattering coefficient β^p can be calculated. Such an approach is widely used for the separation of contributions of dust and smoke particles (Sugimoto and Lee, 2006; Tesche et al., 2009), and the same technique was applied to separate pollen and background aerosol (Noh et al., 2013a; Sicard et al., 2016; Shang et al., 2020). For height-independent depolarization ratios of pollen and background aerosol the pollen backscattering coefficient can be calculated as suggested by Tesche et al. (2009):

$$\beta^p = \beta \frac{(\delta - \delta^b)(1 + \delta^p)}{(\delta^p - \delta^b)(1 + \delta)}, \quad (4)$$

Here β and δ are the backscattering coefficient and particle depolarization ratio of the mixture. The profiles of β_{532}^p and the relative contribution $\frac{\beta_{532}^p}{\beta_{532}}$ are shown in Fig. 8b and e. Computations were performed assuming that height-independent $\delta_{532}^p = 30\%$. For background aerosol, the values $\delta_{532}^b = 3\%$ for 30–31 May and $\delta_{532}^b = 5\%$ for 1–2 June were used. On 30–31 May the contribution of pollen $\frac{\beta_{532}^p}{\beta_{532}}$ at 750 m is estimated as 30 %.

The profiles of β_F and β_{532}^p in Fig. 8b and e behave similarly, decreasing with height. Above 2000 m the decrease in β_F slows down due to the fluorescence of background aerosol. The profiles of the fluorescence capacity G_F and relative contribution $\frac{\beta_{532}^p}{\beta_{532}}$ also demonstrate a good correlation. Thus, both the depolarization and fluorescence techniques lead to the same conclusion: pollen concentration in the boundary layer for the considered episodes decreases with height.

4.2 27–28 March and 21 April 2020 observations

According to the in situ pollen sampling at rooftop level, the maximal pollen content was detected during the birch pollination period on 4–20 April. However, the maximal fluorescence backscattering of lidar data was observed at the end of March when sampling shows an increase in ash (*fraxinus*) pollen emission. The temporal evolution of the range-corrected lidar signal, volume depolarization ratio at 1064 nm and fluorescence backscattering on the night of 27–28 March is shown in Fig. 9. The main part of the aerosol is localized below 2000 m. The back-trajectory analysis demonstrates that the air masses in this episode were transported from eastern Europe. In contrast with Fig. 8,

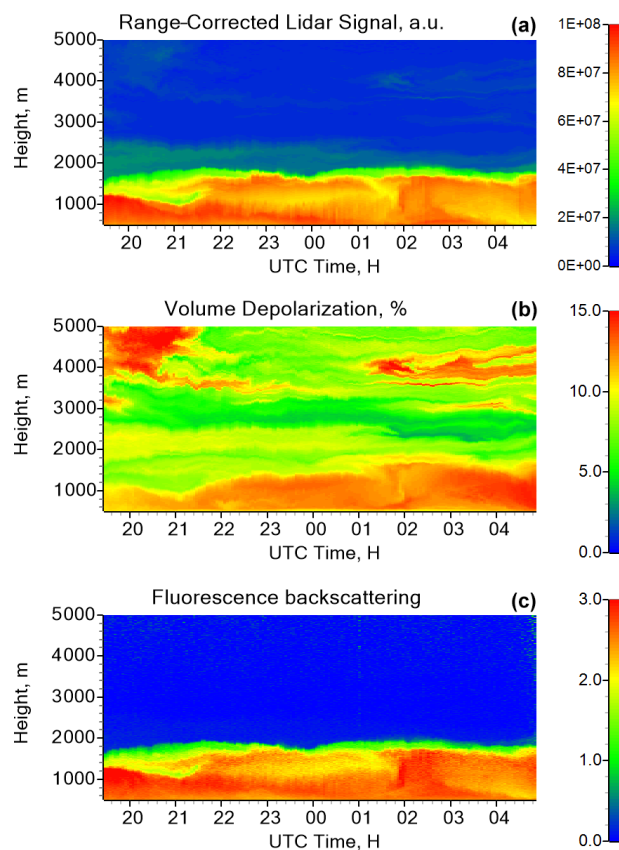


Figure 9. Range-corrected lidar signal at 1064 nm (a), volume depolarization ratio at 1064 nm (b) and fluorescence backscattering coefficient (in $10^{-4} \text{ Mm}^{-1} \text{ sr}^{-1}$, c) measured on 27–28 March 2020.

where fluorescence decreases with height, on 27–28 March the fluorescent particles are rather well mixed inside the PBL (planetary boundary layer). The fluorescence backscattering is high, exceeding $2.5 \times 10^{-4} \text{ Mm}^{-1} \text{ sr}^{-1}$, and the volume depolarization at 1064 nm is about 15 %. The vertical profiles of the particle parameters, averaged for the period 19:20–04:30 UTC, are shown in Fig. 10. In spite of temporal variations of the lidar signal and the fluorescence backscattering inside the PBL (Fig. 9), the averaged overnight profiles of particle parameters are representative. Radiosonde measurements (at both Beauvecchain and Herstmonceux sites) show that RH gradually increased with height from approximately 40 % to 70 % in the 500–1750 m range. The depolarization ratios δ_{532} and δ_{1064} inside the PBL are close, which is in contrast with results in Fig. 8, where δ_{1064} exceeds δ_{532} . This difference can be due to the different types of pollen that were present, which is probably grass in Fig. 8 and birch in Fig. 10. Also, the BAE in Fig. 10 is lower than that in Fig. 8, which decreases the influence of background aerosol on the spectral dependence of the depolarization ratio, as follows from Fig. 2.

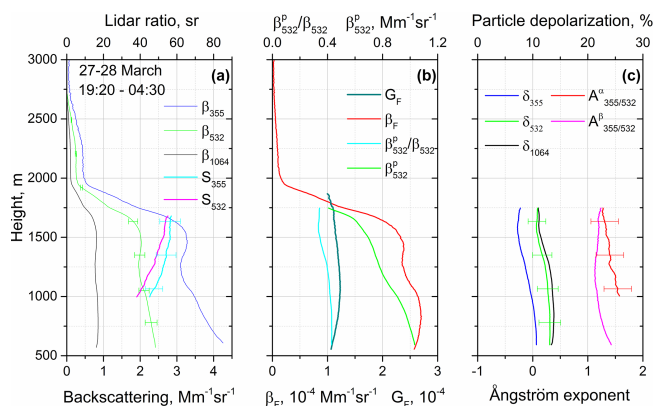


Figure 10. Vertical profiles of (a) particle backscattering coefficients β_{355} , β_{532} and β_{1064} as well as lidar ratios S_{355} and S_{532} . (b) Fluorescence backscattering coefficient β_F , fluorescence capacity G_F , pollen backscattering coefficient β_{532}^p and its contribution to the total backscattering $\frac{\beta_{532}^p}{\beta_{532}}$. (c) Particle depolarization ratios δ_{355} , δ_{532} and δ_{1064} together with extinction $A_{355/532}^\alpha$ and backscattering $A_{355/532}^\beta$ Ångström exponents measured on 27–28 March 2020 at 19:20–04:30 UTC. Profiles of β_{532}^p and $\frac{\beta_{532}^p}{\beta_{532}}$ were computed with the assumption that $\delta_{532}^p = 30\%$ and $\delta_{532}^b = 3\%$.

Both the fluorescence backscattering and depolarization ratios do not demonstrate strong variations inside the 600–1500 m range. The maximum fluorescence capacity exceeds 1.2×10^{-4} , which is significantly higher than G_F for background aerosol in Fig. 3. The profiles of G_F and $\frac{\beta_{532}^p}{\beta_{532}}$ behave reasonably similar, and the slight decrease in $\frac{\beta_{532}^p}{\beta_{532}}$ with height with respect to G_F can be due to dependence of the depolarization ratio of pollen on RH.

Agreement between results obtained from the depolarization and fluorescence techniques in Figs. 8 and 10 corroborates the suggestion that the observed fluorescence is mainly due to the presence of pollen. However, in some episodes the particles with a high fluorescence cross section, other than pollen, could interfere. In particular, such interference occurred in the 20–23 April 2020 period. Figure 11 shows the vertical profiles of particle parameters measured on 21 April. The depolarization ratio $\delta_{1064} = 22\%$ at 750 m was one of the highest during campaign. The RH was low, increasing from 30 % to 45 % in the 800–1500 m range according to Herstmonceux radio sounding. The back-trajectory analysis demonstrates that below 1500 m the air masses are transported from Spain, while at 2000 m the transportation is from northern Europe.

Fluorescence backscattering is stable in the 500–1500 m range, and the fluorescence capacity at 1000 m is about 1.5×10^{-4} , which is a typical value for pollen. However, above 1250 m G_F starts to rise, reaching the value of 2.5×10^{-4} at 1750 m. Such high G_F is more typical for

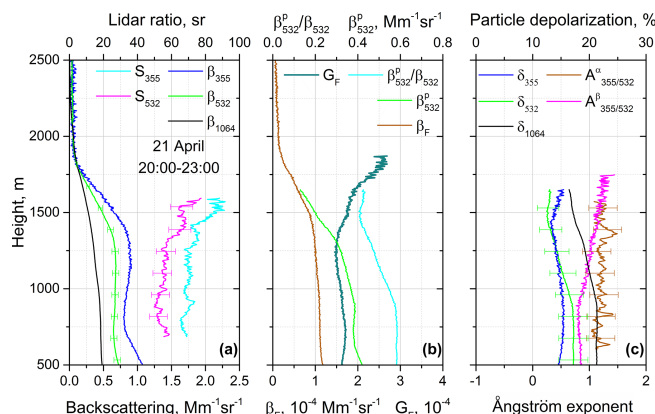


Figure 11. The same particle parameters as in Fig. 10 for 21 April 2020, 20:00–23:00 UTC. Measurements were performed at 30° to the horizon.

smoke particles. The depolarization ratio δ_{355} is about 15 %, and this is higher than corresponding values shown in Figs. 8 and 10, which again may corroborate the presence of smoke particles. We should recall that smoke particles are small, so, in contrast to pollen, their presence influences δ_{355} more strongly than δ_{1064} . Enhanced values of δ_{355} and G_F were observed during the 20–23 April period, indicating the possible presence of biomass burning particles in the aerosol mixture.

4.3 Separation of pollen and smoke layers

During the campaign we observed narrow layers with strong fluorescence. Two examples of such observations, on the nights 13–14 April and 16–27 May 2020, are shown in Fig. 12. The white arrows in this figure point to the fluorescent layers. On 13 April, a weak aerosol layer ($\beta_{532} \approx 0.6 \text{ Mm}^{-1} \text{ sr}^{-1}$ for 23:00–00:00 UTC) is observed at the top of the PBL. This layer demonstrates volume depolarization ratios of about 10 % and high fluorescence backscattering. On 26–27 May a weak layer with high fluorescence backscattering occurs between 3 and 4 km. However, in contrast with the first case, it has a low depolarization ratio, so the layers may have a different nature. Figure 13 shows the vertical profiles of the particle parameters for these two cases. On 13–14 April the fluorescence backscattering below 1000 m is stable, while β_{532} rises, which can be the result of the particle water uptake. Above 1000 m, the depolarization ratio δ_{1064} increases up to 8 %. Results in Fig. 13a are averaged over the 21:15–00:40 UTC temporal interval, but peak values of δ_{1064} between 23:00 and 00:00 UTC exceeded 12 %. Fluorescence backscattering increases simultaneously with the depolarization. The aerosol backscattering coefficient of the fluorescent layer is too low for a reliable calculation of δ_{355} and δ_{532} , so only the profile of δ_{1064} is provided.

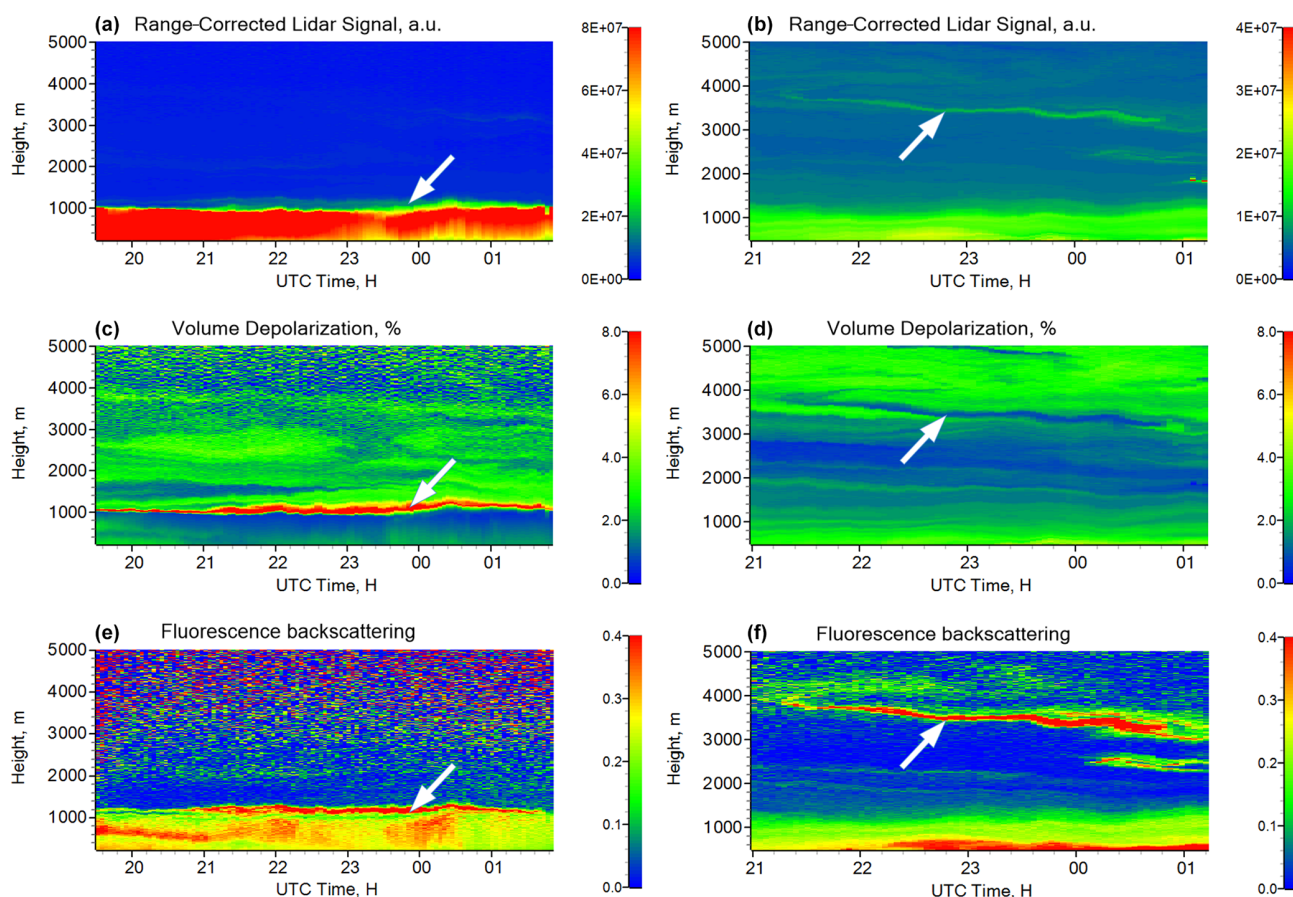


Figure 12. Range-corrected lidar signal at 1064 nm, volume depolarization ratio δ_{1064}^v and fluorescence backscattering coefficient (in $10^{-4} \text{ Mm}^{-1} \text{ sr}^{-1}$) measured on 13–14 April (a, c, e) and 26–27 May 2020 (b, d, f). Arrows point to the fluorescent layers.

On 26–27 May the backscattering coefficient of the fluorescent layer at 3400 m is lower than in Fig. 13a ($\beta_{532} \approx 0.14 \text{ Mm}^{-1} \text{ sr}^{-1}$), so the depolarization ratio δ_{1064} can be calculated only in the center of the layer and is about 2 %, which is significantly lower than that on 13–14 April. However, the fluorescence capacity on 26–27 May is up to 3.5×10^{-4} , which is typical for smoke. Thus, we can conclude that the fluorescent layer on 26–27 May contains smoke particles due to high G_F and low δ_{1064} . On 13–14 April, the fluorescence capacity is significantly lower (about 0.9×10^{-4}) and the depolarization ratio δ_{1064} exceeds 10 %, which is more typical for pollen. Due to low backscattering coefficients of the fluorescent layers in Fig. 12, we are not able to provide a complete set of intensive parameters, such as Ångström exponents and particle depolarization ratios at three wavelengths. However, based on the obtained fluorescence capacities and δ_{1064} values, we conclude that the fluorescent layers probably contain pollen grains in Fig. 12a and smoke particles in Fig. 12b.

4.4 Aerosol classification based on polarization and fluorescence measurements

Table 1 summarizes the results in the campaign, showing the aerosol parameters (particle depolarization and lidar ratios, extinction Ångström exponent, fluorescence backscattering, and capacity) for several days in the March–June 2020 observation period when the contribution of pollen to the total particle backscattering was significant. All available night observations were averaged, and results are given for heights with the highest particle depolarization. Lidar ratios varied approximately in the 40–70 sr range, wherein S_{355} is normally greater than S_{532} . It must be emphasized that pollen lidar ratios may differ for different taxa and that the observed lidar ratios are not attributed to pure pollen but to the aerosol–pollen mixture, so the values provided are influenced by the properties of background aerosol. Moreover, the shape of pollen grains depends on RH (Heidemarie et al., 2018), which may also lead to the variation of pollen lidar ratios. In most of the cases, the depolarization ratio presents strong spectral dependence and increases with wavelength. This spectral dependence is probably the result of mixing

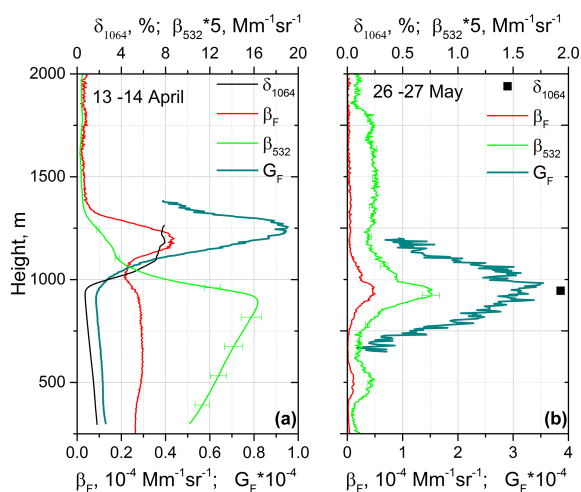


Figure 13. Vertical profiles of elastic β_{532} and fluorescence β_F backscattering coefficients, fluorescence capacity G_F , and particle depolarization ratio δ_{1064} measured on (a) 13–14 April at 21:00–01:00 UTC and (b) 26–27 May 2020 at 23:30–00:40 UTC. Values of β_{532} are multiplied by a factor of 5.

of strongly depolarizing pollen grains with fine background aerosol. The maximal value of observed fluorescence capacity of the pollen–aerosol mixture is 1.6×10^{-4} , which is significantly higher than that of background aerosol but lower than the fluorescence capacity of smoke.

The simultaneous observations of depolarization ratio and fluorescence capacity for different types of aerosol are summarized in Fig. 14. Particle depolarization δ_{532} is plotted versus G_F . The diagram allows for the separation of four types of the particles: (i) dust particles – high δ_{532} and low G_F ; (ii) pollen – high δ_{532} and high G_F ; (iii) smoke – low δ_{532} and high G_F ; and (iv) background aerosol (continental type) – low δ_{532} and low G_F . Points corresponding to the pollen mixture provide an extended pattern because parameters depend on the concentration of pollen in the aerosol mixture. The dust measurements are also scattered because dust over the instrumentation site is transported long-range and mixed with local aerosol. Minimum G_F for dust is about 0.1×10^{-4} , while for smoke maximal G_F is about a factor of 50 higher. The fluorescence capacity depends on the relative humidity, so strong scattering of measurement points can also be partly due to RH variations. Maximal values of G_F for pollen mixture were about 1.5×10^{-4} , and the corresponding depolarization ratios δ_{532} are about 18 %. Thus, assuming that the depolarization ratio of pure pollen is 30 %, we can expect G_F for pure pollen to be about 2.5×10^{-4} , which is comparable with values for smoke.

5 Conclusion

We analyzed the measurements from a multiwavelength Mie–Raman–fluorescence lidar during March–June 2020 in

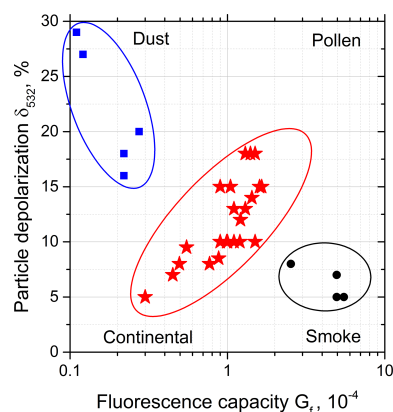


Figure 14. Particle depolarization δ_{532} versus fluorescence capacity G_F . This diagram allows for the identification of dust (blue), smoke particles (black) and aerosol mixtures containing pollen (red).

the north of France to reveal the features that can be attributed to pollen grains. The lidar system allowed us to measure depolarization ratios at three wavelengths simultaneously with the fluorescence backscattering at 466 nm. In numerous episodes during the campaign, high values of the particle depolarization ratio at 1064 nm, exceeding 15 %, were observed. Moreover, the depolarization ratio had strong spectral dependence, being the highest at 1064 nm and lowest at 355 nm, which is expected for big particles of irregular shape mixed with fine, low-depolarizing background aerosol. The increase in particle depolarization correlated with enhancement of the fluorescence backscattering, corroborating the fact that in these episodes we observed aerosol mixtures containing pollen.

The lidar ratios of aerosol–pollen mixtures observed during the campaign varied in a wide range. At low altitudes, where particles presented strong depolarization and fluorescence, in many cases we observed lidar ratios below 40 sr at both wavelengths. However, we also had cases when the lidar ratios at both wavelengths were in the 50–60 sr range. Thus, at the moment we are not able to specify lidar ratios for pure pollen, and additional measurement campaigns in locations with high pollen content are strongly desirable.

Obtained results demonstrate that simultaneous measurements of particle depolarization and fluorescence allows for the separation of dust, smoke particles and pollen grains. Moreover, the fluorescence measurements provide additional information that can be used in aerosol classification schemes. However, further studies are needed to make this technique applicable for quantitative pollen characterization. In the data analysis it is important to account for the process of water uptake by the particles because hygroscopic growth increases the backscattering of background aerosol and influences the pollen grain shape. In the presented lidar configuration, the water vapor channel was absent and radiosonde RH data were not collocated with lidar, which prevented us

Table 1. Lidar-measured aerosol parameters, such as particle depolarization ratios (δ_{355} , δ_{532} , δ_{1064}), lidar ratios (S_{355} , S_{532}), extinction Ångström exponent ($A_{355/532}^\alpha$), fluorescence backscattering coefficient (β_F) and fluorescence capacity (G_F), for several days during the March–June 2020 period when the contribution of pollen to the total particle backscattering could be significant.

Day	Height, m	δ_{355} , %	δ_{532} , %	δ_{1064} , %	S_{355} , sr	S_{532} , sr	$A_{355/532}^\alpha$	$\beta_F \times 10^{-4}$, $\text{Mm}^{-1} \text{sr}^{-1}$	$G_F \times 10^{-4}$
27 March	1150	9	12	13	50	42	1.5	2.5	1.2
7 April	1150	13	13	13	60	60	1.25	1.0	1.3
8 April	1000	11	10	9	50	60	1.0	1.9	1.0
15 April	750	15	15	17	40	40	0.7	0.4	0.9
16 April	1250	15	15	15	–	–	–	0.6	1.6
19 April	650	8	10	14	58	48	1.35	1.2	0.9
20 April	1000	18	18	22	55	45	1.2	0.75	1.3
21 April	750	15	17	22	66	47	1.25	1.1	1.4
22 April	1000	18	18	22	70	55	1.2	0.9	1.5
23 April	1000	6	14	11	53	65	1.25	1.5	1.05
30 May	750	7	10	16	–	–	–	0.8	1.2
1 June	750	7	10	16	–	–	–	1.25	1.5

from performing a quantitative analysis of the hygroscopic effects. Since December 2020, we have recovered the water vapor channel in an upgraded configuration of the lidar. Moreover, we added one more fluorescence channel centered at 549 nm, which will be used in the next pollen campaign in 2021. This additional channel should improve the discrimination of pollen from other aerosols. In the coming campaign we will try to correlate our results with pollen concentrations at different locations in Europe by using a transport model, e.g., SILAM (System for Integrated modelLling of Atmospheric coMposition) (Sofiev et al., 2015). The use of this model should help in the identification of pollen type in our observations.

Data availability. Lidar measurements are available upon request (philippe.goloub@univ-lille.fr).

Author contributions. IV processed the data and wrote the paper. QH and TP performed the measurements. PG supervised the project and helped with paper preparation. MC and NV performed in situ pollen analysis, and MK developed software for data processing.

Competing interests. The authors declare that they have no conflict of interest.

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