



# Supplement of

## Horizontal distribution of tropospheric NO<sub>2</sub> and aerosols derived by dual-scan multi-wavelength multi-axis differential optical absorption spectroscopy (MAX-DOAS) measurements in Uccle, Belgium

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Fitting	Reference	Molecule	Reference
window/ nm	wavelength/		
	nm		
330 - 361	343	BrO (223 K)	Fleischmann et al. (2004)
350 - 370	360	$O_4$	Finkenzeller, H. (private communication)
		NO <sub>2</sub> (298 K)	Vandaele et al. (1998) with $I_o$ correction (SCD of $10^{17}$
			molec.cm <sup>-2</sup> )
		O <sub>3</sub> (223 K)	Serdyuchenko et al. (2014) with $I_0$ correction (SCD of $10^{20}$ molecules/cm <sup>2</sup> )
		HCHO (297 K)	Meller and Moortgat (2000)
		O <sub>3</sub> (243 K)	Pre-orthogonalized Serdyuchenko et al. (2014) with $I_0$
			correction (SCD of 10 <sup>20</sup> molecules/cm <sup>2</sup> )
		Ring	Pseudo cross-section according to Chance and Spurr
			(1997) and normalized as in Wagner et al. (2009)
		Polynomial	Order 3 (4 coefficients)
360-383.5	380	BrO	Fleischmann et al. (2004)
		O4	Finkenzeller, H. (private communication)
		NO <sub>2</sub> (298 K)	Vandaele et al. (1998) with $I_o$ correction (SCD of $10^{17}$ molec.cm <sup>-2</sup> )
		NO <sub>2</sub> (220 K)	Pre-orthogonalized Vandaele et al. (1998) with $I_0$ correction
		Ring	Pseudo cross-section according to Chance and Spurr
			(1997) and normalized as in Wagner et al. (2009)
		Polynomial	Order 5 (6 coefficients)
420-460	447	NO <sub>2</sub> (298 K)	Vandaele et al. (1998) with $I_o$ correction (SCD of $10^{17}$
450 - 490	477		molec.cm <sup>-2</sup> )
		O4(293 K)	Thalman and Volkamer (2013)

Table S1. DOAS settings for  $O_4$  and  $NO_2$  in the six different fitting intervals.

		O <sub>3</sub> (223 K)	Serdyuchenko et al. (2014) with $I_0$ correction (SCD of $10^{20}$ molecules/cm <sup>2</sup> )
		H <sub>2</sub> O	HITRAN (Rothman et al., 2013)
		NO <sub>2</sub> (220 K)	Pre-orthogonalized Vandaele et al. (1998) with $I_{0} \label{eq:Interm}$ correction
		Ring	Pseudo cross-section according to Chance and Spurr
			(1997) and normalized as in Wagner et al. (2009)
		Polynomial	Order 3 (4 coefficients)
510 - 540.1	530	NO <sub>2</sub> (298 K)	Vandaele et al. (1998) with $I_o$ correction (SCD of $10^{17}$ molec.cm <sup>-2</sup> )
		O4 (293 K)	Thalman and Volkamer (2013)
		O <sub>3</sub> (223 K)	Serdyuchenko et al. (2014) with $I_0$ correction (SCD of $10^{20}$ molecules/cm <sup>2</sup> )
		H <sub>2</sub> O	HITRAN (Rothman et al., 2013)
		NO <sub>2</sub> (220 K)	Pre-orthogonalized Vandaele et al. (1998) with $I_0$ correction
		Ring	Pseudo cross-section according to Chance and Spurr
			(1997) and normalized as in Wagner et al. (2009)
		Polynomial	Order 2 (3 coefficients)



Figure S1. Simulated L<sub>NO2</sub> as a function of the RAA for different MLH<sub>NO2</sub> values (from left to right panel: MLH<sub>NO2</sub> equal to 500 m, 1000 m, and 1500 m), wavelengths, one SZA value (30°), and one AOD value (0.3).



Figure S2. Simulated L<sub>NO2</sub> as a function of the SZA for different MLH<sub>NO2</sub> values (from left to right panel: MLH<sub>NO2</sub> equal to 500 m, 1000 m, and 1500 m), wavelengths, one RAA value (60°), and one AOD value (0.3).



Figure S3. Seasonal near-surface NO<sub>2</sub> concentration grids as estimated over Brussels by the RIO air-quality model. The black square shows the MAX-DOAS position, the black polygon the National Airport, the black dots the NO<sub>2</sub> hotspots, and the black line represents the Brussels Ring motorway.



Figure S4. Simulated differential effective light path of O<sub>4</sub> dSCDs (L<sub>04</sub> sim.) and NO<sub>2</sub> dSCDs (L<sub>NO2</sub> sim.) as a function of wavelength for different AOD scenarios.



Figure S5. Maps of hourly averaged NO<sub>2</sub> near-surface concentration horizontal profiles per azimuthal direction for June 28, 2019. The wind direction is shown with the black arrow. The black square shows the MAX-DOAS instrument location, the black polygon the National Airport, the black dots the NO<sub>2</sub> hotspots emitting more than 10 kg of NO<sub>x</sub> per hour (Emission Inventory of the Belgian Interregional Environment Agency, 2017), and the black line represents the Brussels Ring road.



Figure S6. Maps of hourly averaged near-surface aerosol extinction coefficient horizontal profiles per azimuthal direction for June 28, 2019. The wind direction is shown with the black arrow. The black square

shows the MAX-DOAS instrument location, the black polygon the National Airport, the black dots the NO<sub>2</sub> hotspots emitting more than 10 kg of NO<sub>x</sub> per hour (Emission Inventory of the Belgian Interregional Environment Agency, 2017), and the black line represents the Brussels Ring road.



Figure S7. Seasonal histograms of the Probability Density Function (PDF) of the calculated RMS (%) between measured and retrieved NO<sub>2</sub> near-surface concentrations of the horizontal retrieval (see Table 3).



Figure S8. Seasonal histograms of the Probability Density Function (PDF) of the degrees of freedon (DOFS) of the horizontal retrieval (see Table 3).

## S1 MAX-DOAS horizontal NO2 distribution versus air quality model data

The RIO air quality model over the Brussels-Capital Region is used to validate the retrieved near-surface NO<sub>2</sub> concentration horizontal profiles.

RIO is a land-use regression model based on the interpolation of the hourly NO<sub>2</sub> near-surface concentrations measured by the in-situ telemetric air quality network in Brussels (Hooyberghs et al., 2006; Janssen et al., 2008). RIO model provides hourly NO<sub>2</sub> concentration maps on a 4x4 km<sup>2</sup> spatial resolution.

To compare the RIO and MAX-DOAS near-surface NO<sub>2</sub> concentrations, the following steps have been followed:

- 1. One year (March 2018 December 2018) of RIO and MAX-DOAS observations have been used.
- 2. The MAX-DOAS near-surface NO2 concentrations are hourly averaged in each azimuthal direction.
- The MAX-DOAS near-surface NO2 concentrations on the horizontal segment crossing a RIO pixel (i.e., 4x4 km<sup>2</sup>) and located inside each pixel are averaged,
- 4. Pixel-to-pixel comparison between the two datasets has been performed.

Figure S9 shows the seasonal correlation plots for March 2018 – December 2018. The higher correlation coefficients are found during spring (R=0.82) and autumn (R=0.75), while lower correlations are reported in summer (R=0.64) and winter (R=0.51).

During all seasons, the slope values are quite small (s in the range of 0.25-0.50). Given the fact that the RIO model is based in in-situ observations, this finding is excepted. The interested reader is referred to Dimitropoulou et al. (2020).



Figure S9. Seasonal scatter plots of near-surface NO<sub>2</sub> VMRs derived from the dual-scan MAX-DOAS and RIO observations over the Brussels-Capital Region.

### S2 Investigation of the a priori NO<sub>2</sub> profile shape and clouds in TROPOMI NO<sub>2</sub> retrievals

Three additional comparisons were conducted in this study. First, a TROPOMI tropospheric  $NO_2$  column product (DDSv2 product) with an improved FRESCO-S cloud retrieval was tested. As discussed in Dimitropoulou et al. (2020), clouds can significantly affect tropospheric  $NO_2$  VCD retrievals from satellite observations. The dataset is available for four different periods in 2018 – 2019 (see Sect. 3). Fig. S10a and S10b show that the slope value increases by about 29% (equal to 0.49 instead of 0.38 for the baseline product), as well as the correlation coefficient between both datasets (R equal to 0.73 instead of 0.467. This is in agreement with the TROPOMI Routine Operations Consolidated Validation Report (ROCVR; <u>https://mpc-vdaf.tropomi.eu/</u>), where the use of the improved FRESCO-wide resulted in a bias reduction with respect to ground-based  $NO_2$  data.

Secondly, a new TROPOMI data product covering the November 2018 to February 2020 period is used. In this product, the coarse TM5-MP a priori NO<sub>2</sub> profiles are replaced by NO<sub>2</sub> profile shapes from the CAMS regional CTM ensemble at a spatial resolution of  $0.1^{\circ} \times 0.1^{\circ}$  (S5P-CAMS product; Douros et al., in preparation; Ialongo et al., 2019; Tack et al., 2021). As can be seen in Fig. S10c and S10d, using a spatially finer a priori NO<sub>2</sub> vertical profile improves slightly the slope value, which is equal to 0.68 (instead of 0.60 for the baseline TROPOMI product). This represents an increase of the slope by about 13%. This finding indicates that part of the TROPOMI underestimation of tropospheric NO<sub>2</sub> columns is caused by inadequate a priori profiles in the TROPOMI retrievals for urban conditions. On the other hand, the fact that the slope value is still lower than unity, even when CAMS regional a priori profiles are used, indicate that other factors contribute to the TROPOMI underestimation or that CAMS profiles are still sub-optimal, as suggested by results obtained when applying MAX-DOAS profiles to TROPOMI (see below).

Finally, the impact of the a priori profile in the TROPOMI NO<sub>2</sub> retrieval is investigated using MAX-DOAS profile data. For this test, TROPOMI NO<sub>2</sub> columns are recalculated, similarly as in Dimitropoulou et al. (2020), using daily median MAX-DOAS vertical profiles derived in the main azimuthal direction by applying the MMF inversion algorithm. Those TROPOMI NO<sub>2</sub> columns are then compared to the horizontally-resolved MAX-DOAS data, as in Sect. 6.4. Figure S11 presents the comparison results per season. When comparing it with Fig. 23, we find that the change in the NO<sub>2</sub> vertical profile shape improves the slope value in the comparison with ground-based observations. Except for winter, the slopes are improved (slopes in the 0.68 - 1.02 range) due to an increase of the recalculated TROPOMI columns. This result confirms once again that the a priori profile in the TROPOMI retrieval is a key player in the TROPOMI underestimation of tropospheric NO<sub>2</sub> columns in urban conditions, as already stated in previous studies (see e.g. Dimitropoulou et al., 2020; Ialongo et al., 2019; Tack et al., 2021). The present study suggests that in urban conditions, the NO<sub>2</sub> profile shapes from the CAMS regional CTM ensemble are not the most suitable a priori information that can be applied in the TROPOMI retrieval. This finding is in agreement with the recent study of van Geffen et al. (2021b).



Figure S10. Scatter plots between the tropospheric NO<sub>2</sub> columns derived from the dual-scan MAX-DOAS instrument and the TROPOMI pixels over Brussels. The left plots are for the baseline TROPOMI dataset ((a) and (c) panels), while the right plots correspond to two new versions of TROPOMI datasets ((b): improved FRESCO-S cloud product – DDSv2 product; (d) NO<sub>2</sub> a priori profiles from the CAMS regional CTM ensemble – S5P-CAMS product).



Figure S11. Seasonal scatter plots between the horizontally-averaged MAX-DOAS NO<sub>2</sub> VCDs and TROPOMI NO<sub>2</sub> columns recalculated using median daily MAX-DOAS vertical profiles as a priori information.

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