

# Atmospheric N<sub>2</sub>O and CH<sub>4</sub> total columns retrieved from low-resolution Fourier transform infrared (FTIR) spectra (Bruker VERTEX 70) in the mid-infrared region

Minqiang Zhou<sup>1,2</sup>, Bavo Langerock<sup>2</sup>, Mahesh Kumar Sha<sup>2</sup>, Christian Hermans<sup>2</sup>, Nicolas Kumps<sup>2</sup>, Rigel Kivi<sup>3</sup>, Pauli Heikkinen<sup>3</sup>, Christof Petri<sup>4</sup>, Justus Notholt<sup>4</sup>, Huilin Chen<sup>5,6</sup>, and Martine De Mazière<sup>2</sup>

<sup>1</sup>CNRC & LAGEO, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing, China

<sup>2</sup>Royal Belgian Institute for Space Aeronomy (BIRA-IASB), Brussels, Belgium

<sup>3</sup>Space and Earth Observation Centre, Finnish Meteorological Institute, Sodankylä, Finland

<sup>4</sup>Institute of Environmental Physics, University of Bremen, Bremen, Germany

<sup>5</sup>Centre for Isotope Research (CIO), Energy and Sustainability Research Institute Groningen (ESRIG),

University of Groningen, Groningen, the Netherlands

<sup>6</sup>Joint International Research Laboratory of Atmospheric and Earth System Sciences, School of Atmospheric Sciences, Nanjing University, Nanjing, China

Correspondence: Minqiang Zhou (minqiang.zhou@mail.iap.ac.cn, minqiang.zhou@aeronomie.be)

Received: 12 January 2022 – Discussion started: 3 June 2022

Revised: 14 September 2023 - Accepted: 10 October 2023 - Published: 21 November 2023

Abstract. Nitrous oxide (N<sub>2</sub>O) and methane (CH<sub>4</sub>) are two important greenhouse gases in the atmosphere. In 2019, midinfrared (MIR) solar absorption spectra were recorded by a Bruker VERTEX 70 spectrometer and a Bruker IFS 125HR spectrometer at Sodankylä, Finland, at spectral resolutions of 0.2 and  $0.005 \text{ cm}^{-1}$ , respectively. The N<sub>2</sub>O and the CH<sub>4</sub> retrievals from high-resolution MIR spectra have been well investigated within the Network for the Detection of Atmospheric Composition Change (NDACC) but not for MIR spectra gathered with instruments operating at low spectral resolution. In this study, N<sub>2</sub>O and CH<sub>4</sub> retrieval strategies and retrieval uncertainties from the VERTEX 70 MIR lowresolution spectra are discussed and presented. The accuracy and precision of the VERTEX 70 N<sub>2</sub>O and CH<sub>4</sub> retrievals are assessed by comparing them with the coincident 125HR retrievals and AirCore measurements. The relative differences between the N2O total columns retrieved from 125HR and VERTEX 70 spectra are  $-0.3 \pm 0.7 (1\sigma)$  % with a correlation coefficient (R) of 0.93. Regarding the CH<sub>4</sub> total column, we first used the same retrieval microwindows for 125HR and VERTEX 70 spectra, but there is an underestimation in the VERTEX 70 retrievals, especially in summer. The relative differences between the CH<sub>4</sub> total columns retrieved from the 125HR and VERTEX 70 spectra are  $-1.3\pm$ 1.1  $(1\sigma)$ % with a R value of 0.77. To improve the VER-TEX 70 CH<sub>4</sub> retrievals, we propose alternative retrieval microwindows. The relative differences between the CH<sub>4</sub> total columns retrieved from the 125HR and VERTEX 70 spectra in these new windows become  $0.0 \pm 0.8 (1\sigma)$  %, along with an increase in the R value to 0.87. The coincident AirCore measurements confirm that the VERTEX 70 CH<sub>4</sub> retrievals using the latter window choice are better, with relative mean differences between the VERTEX 70 CH<sub>4</sub> retrievals and Air-Core measurements of -1.9% for the standard NDACC microwindows and of 0.13 % for the alternative microwindows. This study provides insight into the N<sub>2</sub>O and CH<sub>4</sub> retrievals from the low-resolution  $(0.2 \text{ cm}^{-1})$  MIR spectra observed with a VERTEX 70 spectrometer, and it demonstrates the suitability of this kind of instrument for contributing to satellite validation, model verification, and other scientific campaigns with the advantage of its transportability and lower cost compared to standard NDACC-type Fourier-transform infrared (FTIR) instruments.

## 1 Introduction

Nitrous oxide (N<sub>2</sub>O) and methane (CH<sub>4</sub>) are two important atmospheric greenhouse gases, which contribute greatly to global climate change (IPCC, 2013). The global mean N<sub>2</sub>O concentration has been increasing continuously since the pre-industrial era and reached up to 332.87 ppb in 2020, as observed by National Oceanic and Atmospheric Administration (NOAA; https://gml.noaa.gov/ccgg/trends\_n2o/, last access: 1 January 2023) surface measurements (Hall et al., 2007). Moreover, the NOAA surface measurements show that the global annual growth rate of N<sub>2</sub>O has been increasing recently. The N<sub>2</sub>O annual growth rate has been about  $1.0 \text{ ppb yr}^{-1}$  during the last 5 years (2015–2020) compared to that of 0.9 ppb yr<sup>-1</sup> from 2010 to 2015. The increasing N<sub>2</sub>O concentration also contributes to stratospheric ozone depletion (Park et al., 2012). Tian et al. (2020) used both bottomup and top-down methods to estimate the atmospheric N<sub>2</sub>O emissions between 1980 and 2016 and highlighted that it is urgent to mitigate human-induced N2O emissions, as their measured growth rate exceeds some of the highest projected emission scenarios. According to the surface in situ measurements (WMO, 2020), the global mean CH<sub>4</sub> concentration kept increasing from 1984 to 1999 and remained stable between 1999 and 2006, but it started to increase again after 2007. It is not fully understood what caused the variation in CH<sub>4</sub> during the last few decades, as the sources of atmospheric CH<sub>4</sub> are diverse and variable (Saunois et al., 2020). Many factors can affect the atmospheric CH<sub>4</sub> concentration, such as natural wetland and fossil fuel emissions (Kirschke et al., 2013), tropical biogenic sources (Schwietzke et al., 2016), biomass burning emissions (Worden et al., 2017), and the atmospheric OH level (Turner et al., 2017).

Mid-infrared (MIR) solar absorption spectra (2400-3000 cm<sup>-1</sup>) recorded by ground-based Fourier-transform infrared (FTIR) spectrometers are used to retrieve the N<sub>2</sub>O and CH<sub>4</sub> vertical profiles and total columns within the Network for the Detection of Atmospheric Composition Change Infrared Working Group (NDACC-IRWG) (De Mazière et al., 2018). The NDACC N<sub>2</sub>O and CH<sub>4</sub> measurements are widely used for understanding their temporal variations (Angelbratt et al., 2011; Zhou et al., 2018), for satellite validation (Barret et al., 2021; Sha et al., 2021; Zhou et al., 2023), and for model comparison and verification (Bader et al., 2017; Zhou et al., 2019b). The NDACC-IRWG sites are typically operated with either a Bomem DA8, a Bruker IFS 120 M/HR, or a Bruker IFS 125 M/HR instrument to record the MIR spectra with a high spectral resolution of 0.003-0.005 cm<sup>-1</sup>. The lower-spectral-resolution spectrometers, such as the Bruker VERTEX 70, Bruker EM27/SUN, and Bruker IRcube with a spectral resolution of  $0.16-0.5 \text{ cm}^{-1}$ , have advantages such as lower cost, higher temporal resolution, and transportability.

Within the Fiducial Reference Measurements for Ground-Based Infrared Greenhouse Gas Observations (FRM4GHG) project, carbon dioxide (CO<sub>2</sub>), methane (CH<sub>4</sub>), and carbon monoxide (CO) total columns retrieved from the nearinfrared (NIR) solar spectra recorded by these lowerspectral-resolution spectrometers are proved to be comparable with the data products retrieved from the Total Carbon Column Observing Network (TCCON) highspectral-resolution measurements (Sha et al., 2020). Intercomparisons between low- and high-resolution MIR spectra and retrievals have been investigated in the past as well; e.g., Taylor et al. (2008) used two FTIR instruments with differing spectral resolutions  $(0.004 \text{ cm}^{-1} \text{ for the Bomem DA8})$ and  $0.2 \text{ cm}^{-1}$  for the ABB Analytical DA5) over a period of 4 months in the summer of 2005, but they pointed out that relatively poor correlations were observed for the N2O and CH4 column observed with high and low spectral resolutions, with the median differences being -0.36 % for N<sub>2</sub>O and 3.7 % for CH<sub>4</sub>.

The Bruker VERTEX 70 instrument is capable of measuring NIR and MIR spectra quasi-simultaneously. However, to our knowledge, the retrievals of N<sub>2</sub>O and CH<sub>4</sub> from the MIR spectra observed by the Bruker VERTEX 70 instrument have not yet been investigated. To better understand the N<sub>2</sub>O and CH<sub>4</sub> retrievals from the MIR spectra observed by lower-spectral-resolution measurements, we operated the Bruker VERTEX 70 and the Bruker 125HR next to each other at Sodankylä in 2019 (March-November). The spectral resolutions of the Bruker VERTEX 70 and the Bruker IFS 125HR are 0.2 and  $0.005 \text{ cm}^{-1}$ , respectively. Further details of the instrumental setup are provided in Sect. 2. The retrieval is performed with the SFIT4 v0.9.4.4 code, updated from SFIT4 (Pougatchev et al., 1995; Hase et al., 2004), which is commonly used in the NDACC-IRWG community. The retrieval strategies and retrieval uncertainties of N2O and CH<sub>4</sub> are discussed in Sect. 3. In Sect. 4, the N<sub>2</sub>O and CH<sub>4</sub> retrievals derived from the VERTEX 70 are compared to those from the 125HR measurements following NDACCrecommended retrieval strategies. In addition, the CH<sub>4</sub> retrievals are also compared to the coincident AirCore profiles. Finally, conclusions are drawn in Sect. 5.

## 2 Site description and instrument configuration

The FTIR measurements used in this study are carried out at the Sodankylä facility of the Finnish Meteorological Institute (FMI), north Finland ( $67.4^{\circ}$  N,  $26.6^{\circ}$  E; 180 m a.s.l.). The FMI site is located in a boreal forest region and is about 6 km south of the downtown. Due to the high latitude, the site is located in the polar night in boreal winter, and the atmospheric components above the site are affected by polarvortex conditions in winter and spring (Kivi et al., 2007; Ostler et al., 2014).

A Bruker IFS 125HR FTIR instrument has been operated at Sodankylä since 2009. The 125HR initially only recorded NIR spectra ( $4000-15000 \text{ cm}^{-1}$ ) with indium gallium ar-

	VERTEX 70	125HR	
Detector	InSb		
Spectral range	$2200 \text{ to } 3500 \text{ cm}^{-1}$		
Spectral resolution	$0.2{ m cm}^{-1}$	$0.005{ m cm}^{-1}$	
Number of scans	12	2	
Duration of one measurement	52 s	194 s	

**Table 1.** The main characteristics of the spectra observed by theVERTEX 70 and 125HR at Sodankylä used in this study.

senide (InGaAs) and silicon (Si) detectors and has been affiliated to the TCCON network (Kivi and Heikkinen, 2016; Wunch et al., 2015). In 2011, the 125HR started measuring MIR spectra ( $2200-3500 \text{ cm}^{-1}$ ) using an indium antimonide (InSb) detector as well, with a spectral resolution of  $0.005 \text{ cm}^{-1}$ . The NIR and MIR spectra at Sodankylä have been observed quasi-simultaneously since then. Profiles of N<sub>2</sub>O and CH<sub>4</sub> have been successfully retrieved from the MIR spectra using the SFIT4 code following the NDACC-IRWG guide, and the products are supporting the Copernicus Atmosphere Monitoring Service (https://cams27.aeronomie. be/, last access: 1 January 2023) and the TROPOMI satellite validation (Sha et al., 2021).

During the ESA-funded FRM4GHG project, several portable FTIR instruments (Bruker EM27/SUN, Bruker IRcube, and Bruker VERTEX 70) were operated between 2017 and 2019 close to the Bruker 125HR instrument at Sodankylä. A schematic diagram of the VERTEX 70 measurement is shown in Fig. 1. In 2017 and 2018, only NIR spectra were recorded by the VERTEX 70 because the main objective of this campaign was to study the performance of  $CO_2$ , CH<sub>4</sub>, and CO retrieved from the NIR spectra observed by the portable FTIR instruments, and we compared them to the TCCON NIR measurements (Sha et al., 2020). However, it is also important to understand the performance of N<sub>2</sub>O and CH<sub>4</sub> retrievals derived from the MIR spectra observed by the portable FTIR and to compare them with the NDACC MIR measurements performed at a high spectral resolution. Therefore, at the beginning of 2019, we added an InSb detector to the VERTEX 70 instrument and performed 1 year of measurements to record MIR spectra. Note that the implementation of MIR measurements for the IRcube and EM27/SUN needs extra effects and has not been done yet. Table 1 lists the main characteristics of the MIR spectra observed by the VERTEX 70 and 125HR instruments; more information about the instruments and the FRM4GHG Sodankylä campaign can be found in Sha et al. (2021).

## 3 FTIR retrievals

 $N_2O$  and  $CH_4$  partial column profiles are derived from the width and shape of their absorption lines in the solar spectra (*y*) observed by the ground-based FTIR instrument, modeled



**Figure 1.** The Bruker VERTEX 70 system operated at Sodankylä during the FRM4GHG campaign, including the weather station and the solar tracker. The orange arrows shows the light path coming from the sun till it enters the FTIR instrument. BIRA-IASB is the Royal Belgian Institute for Space Aeronomy.

as

$$\mathbf{y} = F(\mathbf{x}, \mathbf{b}) + \boldsymbol{\varepsilon},\tag{1}$$

where F is the forward atmospheric radiation transfer model, x is the state vector (the retrieved parameters), b is the forward-model parameters but not retrieved, and  $\varepsilon$  represents the measurement noise and forward-model uncertainty. In this study, the SFIT4 algorithm is used to simulate the observed spectrum and to minimize the residual between the simulated and the observed spectra by looking for the optimal state vector using an inverse model (optimal estimation model or Tikhonov regularization). The inverse problem is solved using the Levenberg–Marquardt Gauss–Newton iteration in the SFIT4 algorithm, and the retrieved state vector can be related to the actual state of the atmosphere by

$$\boldsymbol{x}_r = \boldsymbol{x}_a + \mathbf{A}(\boldsymbol{x}_t - \boldsymbol{x}_a) + \boldsymbol{\varepsilon}, \tag{2}$$

where  $x_a$  and  $x_t$  are the a priori and the true state vectors, respectively. **A** is the averaging kernel matrix (AVK), representing the sensitivity of the retrieved vertical profile of the target species to the true profile.  $\boldsymbol{e}$  is the retrieval uncertainty. The trace of the **A** is the degrees of freedom (DOFs), indicating the number of independent pieces of information (Rodgers, 2000). Following Eq. (2), the retrieved total column can be written as

$$TC_{r} = TC_{a} + A_{col} \cdot (PC_{t} - PC_{a}) + \varepsilon, \qquad (3)$$

where  $TC_a$  and  $TC_r$  are the a priori and retrieved total columns, respectively.  $PC_a$  and  $PC_t$  are the a priori and true partial columns, respectively.  $A_{col}$  is the column averaging kernel (CAVK), representing the sensitivity of the retrieved total column to the true partial column.

## 3.1 Retrieval strategy

## 3.1.1 N<sub>2</sub>O

The N<sub>2</sub>O retrieval strategy for the high-spectral-resolution measurements from the 125HR is well harmonized within the NDACC-IRWG community (Zhou et al., 2019b). Four microwindows (2481.3-2482.6, 2526.4-2528.2, 2537.85-2538.8, 2540.1–2540.7 cm<sup>-1</sup>) including the N<sub>2</sub>O absorption lines are used to retrieve the N<sub>2</sub>O vertical profile (Fig. 2), and the HITRAN2008 spectroscopy is adopted (Rothman et al., 2009). To reduce the effect from the interfering species, HDO, CO<sub>2</sub>, and CH<sub>4</sub> columns are retrieved simultaneously together with the N<sub>2</sub>O vertical profile. The a priori profiles of the target species and CO<sub>2</sub> and CH<sub>4</sub> interfering species are created from the Whole Atmosphere Community Climate Model (WACCM) v4 monthly means between 1980 and 2020. Due to its high spatial-temporal variability, the temperature vertical profile and the a priori profile of HDO is derived from the National Centers for Environmental Prediction (NCEP) 6-hourly reanalysis data with interpolation to the measurement time. In this study, the same four windows are used for VERTEX 70 and 125HR N<sub>2</sub>O retrievals. The signal-to-noise ratio (SNR) of the spectra is calculated from the observed spectra, and then the measurement uncertainty  $S_{\epsilon}$  is calculated as  $1/\text{SNR}^2$  for the diagonal values and 0 for the off-diagonal values. The regularization matrix **R** for the N<sub>2</sub>O retrieval is created with the Tikhonov  $L_1$  method  $\mathbf{R} = \alpha L_1^T L_1$  (Tikhonov, 1963), and the same regulation matrix is applied to VERTEX 70 and 125HR retrievals. To determine the regularization strength  $\alpha$ , we tune it so that the retrieved DOF is similar to the DOF obtained when retrieving N<sub>2</sub>O with the optimal estimation method (Steck, 2002) and where the a priori covariance matrix  $(S_a)$  is derived from the WACCM monthly means. The DOF of N2O profile retrievals from 125HR spectra is about 2.5 derived by the optimal estimation method (OEM), so we set the  $\alpha$  to 4000 to obtain a similar DOF with the Tikhonov method.

## 3.1.2 CH<sub>4</sub>

First, we apply the standard NDACC retrieval strategy for CH<sub>4</sub> from 125HR spectra. Six microwindows (2611.6–2613.35, 2613.7–2615.4, 2835.55–2835.8, 2903.82–2903.925, 2914.7–2915.15, and 2941.51–2942.22 cm<sup>-1</sup>; bands 1–6) are used (Sepúlveda et al., 2014; Zhou et al.,



Figure 2. The four microwindows used for  $N_2O$  retrieval. In each band, the transmittances from atmospheric species, together with the VERTEX 70 typical spectra, are shown in the upper part of each panel, and the typical residuals for the VERTEX 70 (black) and 125HR (grey) spectra at Sodankylä are presented in the lower part of each panel.

2018), and the columns of H<sub>2</sub>O, HDO, and CO<sub>2</sub> are retrieved together with the CH<sub>4</sub> profile. Figure 3 shows the typical transmittances from the atmospheric species for CH<sub>4</sub> retrievals at Sodankylä, together with the fitting residuals for VERTEX 70 and 125HR spectra in each window. However, there are only four and two samples in wavenumber space in bands 3 and 4, respectively, for the VERTEX 70 lowspectral-resolution spectra, and the signals in bands 3 and 4 for the VERTEX 70 low-spectral-resolution spectra are 2-3 times weaker as compared to the high-spectral-resolution spectra. Therefore, we also apply two other relatively broad windows, including stronger CH<sub>4</sub> absorption lines, to provide more information. As a result, we apply two window settings for VERTEX 70 CH<sub>4</sub> retrieval (Table 2): (i) band 1-band 6 (same as the 125HR), named V1, and (ii) replacing band 3 and band 4 with band 3V2 (2875.15–2875.85 cm<sup>-1</sup>) and band 4V2 (2898.3-2898.93 cm<sup>-1</sup>), named V2. Regarding the spectroscopic data, several line lists have been tested for CH<sub>4</sub> retrievals within the NDACC-IRWG community. Currently, ATM19 and HITRAN01 are two recommended

	V1	V2
Retrieval windows $(cm^{-1})$	2611.6-2613.35 (band 1)	2611.6–2613.35 (band 1)
	2613.7-2615.4 (band 2)	2613.7–2615.4 (band 2)
	2835.55-2835.8 (band 3)	2875.15-2875.85 (band 3V2)
	2903.82-2903.925 (band 4)	2898.3-2898.93 (band 4V2)
	2914.7-2915.15 (band 5)	2914.7–2915.15 (band 5)
	2941.51-2942.22 (band 6)	2941.51–2942.22 (band 6)
Notes	Same as 125HR bands 1–6	Replace bands 3 and 4 with bands 3V2 and 4V2

Table 2. The retrieval windows of the two VERTEX 70 CH<sub>4</sub> retrieval strategies.

line lists (Chesnokova et al., 2020). In this study, we take the atmospheric line list ATM19 for CH<sub>4</sub> retrieval. The fitting residuals in the 3V2 and 4V2 bands are comparable to other windows (bands 1, 2, 5, 6) for VERTEX 70. In fact, we have tested many microwindows in addition to bands 3V2 and 4V2, but these two bands offer us the best result (close to 125HR retrievals; see Sect. 4.2). Note that the observed spectrum with a high spectral resolution (125HR) is not well fitted in band 3V2. Therefore, we do not apply the V2 retrieval strategy for 125HR spectra. As with the N<sub>2</sub>O retrieval, the a priori profiles of CH<sub>4</sub> and CO<sub>2</sub> are created as the average WACCM monthly mean between 1980 and 2020, and the a priori profiles of H<sub>2</sub>O and HDO are derived from the NCEP 6-hourly reanalysis data with interpolation to the measurement time. The SNR of the spectra is calculated from the observations. The regularization matrix  $\mathbf{R}$  for the  $CH_4$  retrieval is also created with the Tikhonov  $L_1$  method, with  $\alpha = 10000$ , leading to a mean DOF of 2.1 for 125HR retrievals.

#### 3.2 Retrieval vertical sensitivity

## 3.2.1 N<sub>2</sub>O

Figure 4 shows both the CAVK and the AVK together with the a priori and retrieved N<sub>2</sub>O vertical profiles for 125HR and VERTEX 70. The mean and standard deviation (SD) of the DOF are  $2.6 \pm 0.1$  for 125 HR and  $1.6 \pm 0.1$  for VER-TEX 70. As expected, more vertical information is derived from the high-spectral-resolution spectra (125HR) as compared to the low-spectral-resolution spectra (VERTEX 70). The typical AVKs show that the retrieved N<sub>2</sub>O profile from the 125HR has a good sensitivity between the surface and 25 km, while the retrieved N<sub>2</sub>O profile from the VERTEX 70 is mainly sensitive to the vertical range between the surface and 20 km. The CAVK of 125HR is close to 1 in the whole vertical range and hardly varies with solar zenith angle (SZA). The CAVK of VERTEX 70 is close to 1 in the whole vertical range at a large SZA ( $> 70^{\circ}$ ), but it starts deviating from 1 above 20 km at a relatively small SZA. Since the partial column of N2O above 20 km makes up only about 1.5% of the total column, such deviation in the CAVK of VERTEX 70 does not affect the retrieved  $N_2O$  total column strongly, and the smoothing error of the  $N_2O$  total column is very small (see Sect. 3.3). The retrieved  $N_2O$  vertical profiles from the 125HR and VERTEX 70 are similar but not exactly the same. Due to the limited vertical information, we will focus on the partial columns or the total column instead of the vertical profile (see Sect. 4).

#### 3.2.2 CH<sub>4</sub>

The CAVK, AVK, and retrieved profiles from the 125HR and VERTEX 70 (V1 and V2) are shown in Fig. 5. The vertical sensitivities of the retrieved CH<sub>4</sub> total column from the 125HR and VERTEX 70 V1 are similar, since they use the same retrieval windows. There is no obvious SZA dependence in the CAVK for the 125HR CH<sub>4</sub> retrievals. However, for both VERTEX 70 V1 and VERTEX 70 V2 retrievals, the CAVK varies with SZA, especially above 10 km. The mean DOF from 125HR CH<sub>4</sub> retrievals is 2.1, which is about twice than that from the VERTEX 70 V1 retrievals (DOF  $\sim$  1.0). As for the N<sub>2</sub>O retrievals, more vertical information can be obtained with a higher spectral resolution. The CAVK of the VERTEX 70 V2 retrieval shows that the retrieved CH<sub>4</sub> total column has better sensitivity in the lower troposphere as compared to the VERTEX 70 V1 retrieval, and the mean DOF is 1.3. The vertical information obtained in the VER-TEX 70 V2 retrieval is higher than that in the VERTEX 70 V1 retrieval. Still, the DOFs indicate that we must focus on the CH<sub>4</sub> total columns for both the VERTEX 70 V1 and the VERTEX 70 V2 retrievals.

### 3.3 Retrieval uncertainty

According to Rodgers (2000), the retrieval uncertainty of  $x_r$  can be written as

$$\boldsymbol{x}_{\mathrm{r}} - \boldsymbol{x}_{\mathrm{t}} = (\mathbf{A} - \mathbf{I})(\boldsymbol{x}_{\mathrm{t}} - \boldsymbol{x}_{\mathrm{a}}) + \mathbf{G}_{\mathrm{y}}\mathbf{K}_{\mathrm{b}}(\boldsymbol{b}_{\mathrm{t}} - \boldsymbol{b}) + \mathbf{G}_{\mathrm{y}}\boldsymbol{\epsilon}_{\mathrm{noise}}, \quad (4)$$

where  $\mathbf{G}_{\mathbf{y}}$  is the contribution matrix, **K** is the Jacobian matrix,  $\hat{\boldsymbol{b}}$  is the best estimate of the forward-model parameter,  $\boldsymbol{b}_{t}$  is the true forward- model parameter, and  $\epsilon_{\text{noise}}$  is the noise of the observed spectra. The items on the right side of Eq. (4) are the smoothing error (first), the model parameter error (second), and the measurement error (third).



Figure 3. The six microwindows (bands 1–6) used for 125HR and VERTEX 70 V1 CH<sub>4</sub> retrieval, together with the other two microwindows used for VERTEX 70 V2 CH<sub>4</sub> retrieval. In each band, the transmittances from atmospheric species, together with the VERTEX 70 typical spectra, are shown in the lower part of each panel, and the typical residuals for the VERTEX 70 (black) and 125HR (grey) spectra at Sodankylä are presented in the upper part of each panel.

**Table 3.** The relative uncertainties (%) of N<sub>2</sub>O total column retrievals from 125HR and VERTEX 70 spectra at Sodankylä. "–" indicates zero uncertainty contribution. See text for the interpretation of SD ( $\pm$  1 h noon). The time zone throughout the text and figures where applicable is local time.

	125H	łR	VERTEX 70		
	Systematic	Random	Systematic	Random	
Smoothing	0.1	0.1	0.1	0.2	
Spectroscopy	3.5	-	3.6	-	
Temperature	0.3	0.1	0.3	0.1	
SZA	0.3	1.2	0.3	1.2	
zshift	0.2	0.2	0.2	0.2	
Measurement	-	0.1	-	0.9	
Total	3.6	1.5	3.7	1.8	
SD ( $\pm 1$ h noon)		1.2		2.0	

Regarding the smoothing error, the systematic uncertainty of  $(\mathbf{x}_t - \mathbf{x}_a)$  for both N<sub>2</sub>O and CH<sub>4</sub> is set to 5 %, and the random uncertainty of  $(x_t - x_a)$  is created from the covariance matrix from the WACCM monthly means between 1980 and 2020. The important model parameters are the spectroscopy, temperature, SZA, and zero-level shift (zshift). The systematic uncertainties of the N2O and CH4 line intensities are set to 3%, and the systematic uncertainties of the spectroscopic temperature and pressure broadening parameters are set to 5% (Rothman et al., 2009; Gordon et al., 2017). It is assumed that there is no random uncertainty associated with the spectroscopic parameters. The systematic and random uncertainties of the temperature profiles are evaluated by calculating the mean differences and SD, respectively, of ERA5 and NCEP temperature profiles in 2019. The systematic and random uncertainties of SZA are set to 0.1 % and 0.5 %, respectively. The systematic and random uncertainties of zshift are both set to 0.1 % (Zhou et al., 2016; García et al., 2021). We assume that the measurement noise has only a random uncertainty component and evaluate it as  $\epsilon_{noise} = 1/SNR$ , with SNR the calculated signal-to-noise ratio of the spectra.

The systematic and random retrieval uncertainties of the retrieved N<sub>2</sub>O and CH<sub>4</sub> columns retrieved from 125HR and VERTEX 70 spectra are then evaluated according to Eq. (4) with the above assumptions and are listed in Tables 3 and 4, respectively. The systematic and random retrieval uncertainties of the  $N_2O$  total column are 3.6% and 1.5% for retrievals from 125HR spectra and 3.7% and 1.8% for retrievals from VERTEX 70 spectra. The systematic and random retrieval uncertainties of the CH4 total column are 3.9 % and 1.6% for 125HR spectra retrievals, 4.1% and 2.3% for VERTEX 70 spectra V1 retrievals, and 4.3 % and 2.4 % for VERTEX 70 spectra V2 retrievals. For both N<sub>2</sub>O and CH<sub>4</sub> retrievals, the dominant systematic uncertainties come from the spectroscopy. According to our estimations, the random uncertainties in N2O retrievals between 125HR and VER-TEX 70 are dominated by the SZA and measurement er-



**Figure 4.** N<sub>2</sub>O retrievals: the column averaging kernels' variation with solar zenith angles (in units of  $(mol cm^{-2}) (mol cm^{-2})^{-1}$ ), the typical averaging kernel matrix (in units of  $(vmr vmr^{-1}) (vmr vmr^{-1})^{-1}$ , where "vmr" denotes the volume mixing ratio), and the a priori profile (black) and retrieved profiles (mean: solid red line;  $1\sigma$ : red-shaded area) between the surface and 50 km for 125HR retrievals (**a–c**) and for VERTEX 70 retrievals (**d–f**).

Table 4. Same as Table 3 but for the relative uncertainties (%) of  $CH_4$  total columns from 125HR and VERTEX 70 V1 and V2 retrievals at Sodankylä.

	125HR		VERTEX 70 V1		VERTEX 70 V2	
	Systematic	Random	Systematic	Random	Systematic	Random
Smoothing	0.1	0.3	0.2	0.5	0.1	0.1
Spectroscopy	3.6	-	3.6	-	3.6	-
Temperature	1.0	0.3	1.7	1.6	1.5	1.5
SZA	0.3	1.2	0.2	1.0	0.2	1.0
zshift	0.9	0.9	1.0	1.0	1.4	1.4
Measurement	_	0.1	-	0.6	-	0.6
Total	3.9	1.6	4.1	2.3	4.3	2.4
SD ( $\pm 1$ h noon)		1.9		2.4		2.6

ror, while the random uncertainties in CH<sub>4</sub> retrievals from 125HR and VERTEX 70 spectra are dominated by the SZA, temperature, zshift, and measurement error. To assess the estimated random uncertainty, we calculate the mean of all the standard deviations (SDs) of the retrieved N<sub>2</sub>O or CH<sub>4</sub> total columns within  $\pm 1$  h around noon (at least two available retrievals) on each measurement day. We find a good agreement with the estimated random uncertainties for both 125HR and VERTEX 70 retrievals, indicating that our estimated random uncertainties are reasonable. The estimated random uncertainties of retrievals from 125HR spectra are smaller than those from VERTEX 70 spectra for both N<sub>2</sub>O

and CH<sub>4</sub> retrievals, and this is confirmed by the SDs of the retrieved total columns.

#### 4 Results and discussions

## 4.1 N<sub>2</sub>O

### 4.1.1 Total column

The time series of all the individual  $N_2O$  total columns retrieved from the 125HR and VERTEX 70 are shown in Fig. 6, together with the relative differences ((VER-



**Figure 5.** CH<sub>4</sub> retrievals: the column averaging kernels' variation with solar zenith angles (in units of  $(mol cm^{-2})(mol cm^{-2})^{-1}$ ), the typical averaging kernel matrix (in units of  $(vmr vmr^{-1})(vmr vmr^{-1})^{-1}$ ), and the a priori profile (black) and retrieved profiles (mean: solid red line;  $1\sigma$ : red-shaded area) between the surface and 50 km for 125HR retrievals (**a–c**), for VERTEX 70 V1 (**d–f**), and for VERTEX 70 V2 (**g–i**) retrievals.

TEX - 125HR / 125HR × 100 %) between the coincident hourly means. The mean difference is -0.3% ( $1\sigma = 0.7\%$ ), which is within the retrieval uncertainties of 125HR and VERTEX 70 (Table 2). To reduce the impact of the vertical sensitivities from the 125HR and VERTEX 70, we smooth the retrieved N<sub>2</sub>O 125HR profile by the VERTEX 70 averaging kernel. The results with and without smoothing correction are very close to each other, with a relative difference of less than 0.1 %, because the smoothing error is very small. Low-N<sub>2</sub>O total columns in spring are observed by both 125HR and VERTEX 70 measurements. There is no clear seasonal variation in the relative differences. Figure 7 shows that the Pearson correlation coefficient (R) between the coincident daily means is 0.93. The correlation coefficient R between the H<sub>2</sub>O total columns and the relative differences is -0.02 and between the SZAs and the relative differences is 0.05, thus demonstrating that there are no water vapor or SZA dependences in the relative differences.

## 4.1.2 Partial columns

Apart from the total column, we also compare the 125HR and VERTEX 70  $N_2O$  retrievals in two partial columns (lower layer: 0–6 km; upper layer: 6–25 km). The two partial columns are selected as we can get a DOF of 0.8 in each layer for the VERTEX 70  $N_2O$  retrievals. The DOFs of the 125HR  $N_2O$  are 1.0 and 1.5 for the lower and upper partial columns, respectively.

The correlation coefficients between the 125HR and VER-TEX 70 coincident  $N_2O$  partial column daily means are 0.86 and 0.94 for the lower and upper layers, respectively (not shown). The time series of the  $N_2O$  partial columns retrieved M. Zhou et al.: CH<sub>4</sub> and N<sub>2</sub>O FTIR retrievals with low-resolution spectra



**Figure 6.** The time series of the  $N_2O$  total columns retrieved from the 125HR and VERTEX 70, together with the differences between their coincident hourly means. The time format here and in Figs. 8 and 9 is year-month.

from the 125HR and VERTEX 70, together with their relative differences, are shown in Fig. 8. Seasonal variations in N<sub>2</sub>O partial columns observed by 125HR and VERTEX 70 retrievals are consistent in both layers. The low N<sub>2</sub>O values in the upper layer due to the polar vortex in spring (Zhou et al., 2019b) are well captured by 125HR and VERTEX 70 retrievals. The mean  $(\pm SD)$  of the differences between the 125HR and VERTEX 70 is  $0.0 \pm 1.9\%$  in the partial column between 6 and 25 km and  $-0.5 \pm 1.7$  % between the surface and 6 km. The mean differences in the partial columns are close to that in the total column. However, the SDs of the differences in the partial columns are larger than that in the total column. Moreover, it is found that there is a seasonal variation in the relative differences for both layers; e.g., the 125HR-retrieved N<sub>2</sub>O is larger/lower than the VERTEX 70retrieved N<sub>2</sub>O in spring/summer between 6 and 25 km. As the seasonal variations in the relative differences in the two layers are opposed, they compensate for each other in the total column, for which there is almost no seasonal variation in the relative difference. Keep in mind that the DOFs of the two partial columns (0-6 km and 6-25 km) of N2O derived from the VERTEX 70 spectra are both 0.8, indicating that neither layer is truly independent. In summary, the partial columns (0-6 km and 6-25 km) of N<sub>2</sub>O can be derived from the VER-TEX 70 spectra but with a larger uncertainty as compared to the 125HR retrievals.

## 4.2 CH<sub>4</sub>

## 4.2.1 Total column

Since the mean DOF of the VERTEX 70 CH<sub>4</sub> retrievals is relatively small (1.0 for V1 and 1.3 for V2), we only look at the CH<sub>4</sub> total columns retrieved from the 125HR and VERTEX 70 spectra. As for the previous comparisons, the 125HR CH<sub>4</sub> profile is smoothed with the VERTEX 70 AVK to reduce the impact of the different vertical sensitivities of 125HR and VERTEX 70 retrievals. The time series of the retrieved CH<sub>4</sub> total columns from the 125HR and VERTEX 70 V1 and V2 are shown in Fig. 9. The mean difference between the 125HR and VERTEX 70 V1 CH<sub>4</sub> total columns is  $-1.3 \pm 1.1$  %. The VERTEX 70 V1 CH<sub>4</sub> total column is much lower than the 125HR-retrieved CH<sub>4</sub> in summer. The mean difference between the 125HR and VERTEX 70 V2 CH<sub>4</sub> total columns is  $0.0 \pm 0.8$  %.

The scatterplots (Fig. 10) confirm that the VERTEX 70 V2 CH<sub>4</sub> total columns are much closer to the 125HR retrievals as compared to the VERTEX 70 V1 results. The correlation coefficients between 125HR and VERTEX 70 coincident CH<sub>4</sub> total column daily means are 0.77 for VERTEX 70 V1 and 0.87 for VERTEX 70 V2. The relative differences between 125HR and VERTEX 70 V1 CH<sub>4</sub> total columns vary with the SZA (R = 0.42) and H<sub>2</sub>O total column (R = -0.76). The relative differences between 125HR and VERTEX 70 V2 CH<sub>4</sub> total columns have almost no SZA dependence (R = -0.03). Besides, the linear regression between the relative differences and SZA total column show that the slope of the regression line for V2 is very close to zero as compared to the V1 value of 0.41  $\%^{\circ 1}$ . In addition, the correlation with the H<sub>2</sub>O total column is strongly reduced (R = -0.51). The linear regression between the relative differences and H<sub>2</sub>O total column shows that the slope of the regression line for V2 is  $-1.6 \times 10^{-23}$  % molec. cm<sup>-2</sup>, which is about 2.5 times less that for V1 with  $-3.9 \times 10^{-23}$  % molec. cm<sup>-2</sup>.

#### 4.2.2 Comparison with AirCore measurements

AirCore consists of a long tube system that ascends with a balloon to reach the mid-stratosphere ( $\sim 25-30$  km). It starts collecting the surrounding air when descending from the high altitude to the surface. After landing, the sampled air inside the long tube is passed through a gas analyzer and represents a vertical profile of the atmospheric components if the analysis is performed before the air becomes mixed by diffusion. For more information about the AirCore system, we refer to Karion et al. (2010). AirCore launches have been regularly operated by FMI at Sodankylä since 2013, and the data have been used to compare with coincident FTIR and satellite measurements (Zhou et al., 2019a; Tu et al., 2020; Sha et al., 2020). There were six AirCore measurements in 2019 performed in the framework of the FRM4GHG campaign. For the moment, only  $CO_2$ ,  $CH_4$ , and CO are observed by the AirCore system at Sodankylä.

On the AirCore launch days, the daily mean FTIRretrieved CH<sub>4</sub> columns and the relative differences between the FTIR-retrieved CH<sub>4</sub> total column and AirCore measurements are listed in Table 5. There is no VERTEX 70 measurement on 28 June 2019 and no 125HR measurement on 1 August 2019. To account for the limited vertical sensitivity of the FTIR retrievals, we smooth the AirCore profile with the corresponding FTIR averaging kernels (Rodgers and Connor,



Figure 7. (a) The correlation plot between the coincident 125HR- and VERTEX 70-retrieved N<sub>2</sub>O daily mean total columns, colored by the measurement month. The error bar represents the daily SD. (b, c) The variations in the hourly relative differences with the SZA and water vapor total column. *R* is the correlation coefficient, and *N* is the number of coincident days.

**Table 5.** The daily mean FTIR-retrieved CH<sub>4</sub> total columns ( $10^{19}$  molec. cm<sup>-2</sup>) and their relative differences against smoothed AirCore, (FTIR – AirCore) / AirCore × 100 %.

Date	125HR	Difference	VERTEX 70 V1	Difference	VERTEX 70 V2	Difference
10 April 2019	3.922	0.54 %	3.916	-0.01 %	3.920	0.12 %
28 June 2019	3.803	0.17~%				
24 July 2019	3.884	0.01~%	3.780	-3.26%	3.886	0.21 %
1 August 2019			3.817	-0.99%	3.877	0.82%
28 August 2019	3.933	1.15 %	3.804	-2.85%	3.862	-0.65~%
9 September 2019	3.937	0.44 %	3.844	-2.67~%	3.907	-0.16%
Mean ± SD	$3.90\pm0.05$	$0.46 \pm 0.39$ %	$3.83\pm0.05$	$-1.90 \pm 1.2$ %	$3.89 \pm 0.02$	$0.13 \pm 0.50$ %

$$\boldsymbol{x}_{\text{core,s}} = \boldsymbol{x}_{a} + \mathbf{A}(\boldsymbol{x}_{\text{core}} - \boldsymbol{x}_{a}), \tag{5}$$

where  $x_{core,s}$  is the smoothed AirCore profile,  $x_{core}$  is the extended AirCore profile,  $x_a$  is the a priori profile of the FTIR retrieval, and **A** is the FTIR averaging kernel matrix (125HR, VERTEX 70 V1 or V2). Note that the original AirCore profile is extended with a scaled FTIR a priori profile above its maximum height to obtain the  $x_{core}$ . The scaling factor is calculated by the ratio of the observed AirCore CH<sub>4</sub> and FTIR a priori CH<sub>4</sub> mean mole fractions between 20 km and the AirCore profile's maximum height.

The mean relative differences between the FTIR-retrieved CH<sub>4</sub> total columns and the smoothed AirCore measurements are 0.46 %, -1.90 %, and 0.13 % for 125HR, VERTEX 70 V1, and VERTEX 70 V2, respectively. Consistent with the results shown in Fig. 9, the VERTEX 70-V1-retrieved CH<sub>4</sub> total columns are generally close to the 125HR retrievals in spring, but they become much lower than the 125HR retrievals in summer. As an example, the VERTEX 70-V1-

retrieved CH<sub>4</sub> daily mean is 3.26 % less than the smoothed AirCore on 24 July 2019. Several tests have been carried out to understand why there is an underestimation in VER-TEX 70 V1 retrievals on that day: (i) using ERA5 reanalysis data instead of NCEP for the water vapor and temperature a prior profiles; (ii) using the coincident AirCore measurement instead of WACCM as the a priori CH<sub>4</sub> profile; (iii) using the VERTEX 70-V2-retrieved CO<sub>2</sub>, H<sub>2</sub>O, and HDO instead of WACCM as the a priori interfering species; and (iv) using HITRAN2001 or HITRAN2016 instead of ATM19 for the spectroscopy. However, all these parameters only have a small impact on the VERTEX 70-V1-retrieved CH<sub>4</sub> total column, with a relative difference of less than 0.2 %. So far, it remains unclear to us why the VERTEX 70 V1 retrievals underestimate the total columns as observed by the coincident AirCore measurements, but a similar bias is observed with respect to the 125HR retrievals. For the moment, we recommend using the V2 CH<sub>4</sub> retrieval strategy for analyzing the VERTEX 70 MIR low-spectral-resolution spectra.



**Figure 8.** Same as Fig. 6 but for  $N_2O$  in two partial columns (**c**, **d**: 0-6 km; **a**, **b**: 6-25 km).



**Figure 9.** Same as Fig. 6 but for the comparison between 125HR and VERTEX 70 V1 CH<sub>4</sub> total columns ( $\mathbf{a}$ ,  $\mathbf{b}$ ) and between 125HR and VERTEX 70 V2 CH<sub>4</sub> total columns ( $\mathbf{c}$ ,  $\mathbf{d}$ ).

Figure 11 shows the CH<sub>4</sub> total columns retrieved from the 125HR and VERTEX 70 V2, together with the Air-Core measurements on these six dates. We fit the coincident AirCore and FTIR data pairs using y = ax. The slope is  $0.9952 \pm 0.0041$  for 125HR and  $0.9990 \pm 0.0048$  for VER-TEX 70 V2. In terms of the vertical profiles, the mean relative difference between the smoothed AirCore and 125HRretrieved CH<sub>4</sub> is within 2% for all heights. However, the mean relative difference between the smoothed AirCore and VERTEX 70-V2-retrieved CH<sub>4</sub> is about -3.0% in the lower troposphere and larger than 5.0% in the stratosphere. The low DOF value (1.3) suggests that we could only get the total column of CH<sub>4</sub> from the VERTEX 70 V2 retrieval, and the AirCore measurements confirm that the CH<sub>4</sub> vertical profile of the 125HR retrieval is closer to the observed one as compared to the VERTEX 70 V2 retrieval.

## 4.3 N<sub>2</sub>O vs. CH<sub>4</sub>

The vertical profile shapes of N<sub>2</sub>O and CH<sub>4</sub> are similar: a high but relatively constant mole fraction in the troposphere and decreasing with altitude in the stratosphere. Most of the N<sub>2</sub>O and CH<sub>4</sub> emissions are near the surface, so the mole fractions of these two species are high in the troposphere. The N<sub>2</sub>O and CH<sub>4</sub> molecules are transported from the troposphere to the stratosphere via the Brewer-Dobson circulation. The long lifetimes of N<sub>2</sub>O and CH<sub>4</sub> cause a compact correlation between their abundances in the stratosphere, as has been confirmed by balloon, aircraft, and satellite measurements (Herman et al., 1998; Sawa et al., 2015; Wang et al., 2014). The 125HR and VERTEX 70 measurements show that the N<sub>2</sub>O and CH<sub>4</sub> total columns have a positive correlation (Fig. 12), with an R value larger than 0.9. To reduce the sampling uncertainty, we select the days for which both 125HR and VERTEX 70 retrievals are available (83 in total). A larger scatter exists in the N<sub>2</sub>O and CH<sub>4</sub> total columns from the VERTEX 70 as compared to 125HR, because of the larger random uncertainties in VERTEX 70 retrievals (Tables 3 and 4). The slope of the linear regression ( $\mathbf{y} = a\mathbf{x} + b$ ) is 5.27 ± 0.20 from the 125HR retrievals and  $4.91 \pm 0.29$  from the VERTEX 70 retrievals. The uncertainty in the slope is calculated by the bootstrapping method (Fox, 2008) using the daily SD as the random variability. The difference in the slopes between the N<sub>2</sub>O and CH<sub>4</sub> total columns observed by the 125HR and VERTEX 70 measurements is within their combined uncertainties. To conclude, the high correlations between N<sub>2</sub>O and CH<sub>4</sub> total columns at Sodankylä are confirmed by the 125HR and VERTEX 70 retrievals.

#### 5 Conclusions

In this study, the atmospheric abundances of  $N_2O$  and  $CH_4$  are retrieved from low-resolution (0.2 cm<sup>-1</sup>) MIR solar ab-



Figure 10. Same as Fig. 7 but for the coincident 125HR and VERTEX 70 V1 (a-c) or V2 (d-f) CH<sub>4</sub> retrievals.



**Figure 11.** Upper panels: the time series of the CH<sub>4</sub> total columns retrieved from the 125HR (light green) and VERTEX 70 V2 (cyan), together with their corresponding smoothed AirCore total columns (triangle) on the six dates. The rightmost plot is the scatterplot between the smoothed AirCore and FTIR CH<sub>4</sub> total columns. Note that there is no VERTEX 70 measurement on 28 June 2019 and no 125HR measurement on 1 August 2019. On 28 August 2019, the smoothed AirCore (with 125HR and VERTEX 70 V2) CH<sub>4</sub> total columns are very close to each other, so we can only see one triangle. Lower panels: the CH<sub>4</sub> vertical profiles from 125HR and VERTEX 70 V2 retrievals, together with the AirCore measurements and the smoothed AirCore profiles. The rightmost plot shows the mean and SD of the relative differences between the smoothed AirCore and FTIR CH<sub>4</sub> vertical profiles.



Figure 12. The correlation plot between the N<sub>2</sub>O and CH<sub>4</sub> total column daily means from the 125HR (a) and VERTEX 70 V2 (b). The error bar is the daily SD. N is the number of days, R is the correlation coefficient, and a is the slope of the linear fit.

sorption spectra recorded by the Bruker VERTEX 70 spectrometer in 2019 at Sodankylä and are compared with standard NDACC retrievals from coincident 125HR spectra with a spectral resolution of  $0.005 \text{ cm}^{-1}$ . Moreover, the CH<sub>4</sub> retrievals from the 125HR and VERTEX 70 spectra are also compared with the coincident in situ AirCore measurements.

We find that the standard NDACC retrieval strategy works well for the VERTEX 70 N2O retrieval as well. The relative differences between the N2O total columns retrieved from the 125HR and VERTEX 70 spectra are  $-0.3 \pm 0.7$  %. The correlation between the N2O total columns retrieved from 125HR and VERTEX 70 spectra is 0.93, and the relative differences are almost independent of SZA or water vapor content. However, the retrieval uncertainties of the N2O total columns from the VERTEX 70 spectra are larger than those of the total columns retrieved from 125HR spectra due especially to a higher measurement noise. What's more, the mean DOF of the N2O retrievals from the VERTEX 70 spectra is 1.6, which is less than the DOF of 2.6 obtained for the retrievals from the 125HR spectra. This agrees with the expectation that less vertical information can be derived from the spectra with a lower spectral resolution. Inspection of the  $N_2O$  partial columns (0–6 and 6–25 km) confirms that this vertical information is captured by the VERTEX 70 measurements, albeit with a larger uncertainty than what is obtained from the 125HR measurements.

Unlike N<sub>2</sub>O, the recommended NDACC retrieval strategy does not perform well for the CH<sub>4</sub> retrievals from VER-TEX 70 spectra. Using the same retrieval windows (V1 retrieval strategy), the relative differences between the CH<sub>4</sub> total columns retrieved from the 125HR and VERTEX 70 spectra are  $-1.3 \pm 1.1$ %, and they show a seasonal variation: in spring the VERTEX 70-V1-retrieved CH<sub>4</sub> total columns are close to the ones retrieved from the 125HR spectra, but in summer they are about 3.0% smaller than the 125HR retrievals. In addition, these relative differences vary with the SZA and water vapor content. The DOF of V1-retrieved CH<sub>4</sub> profiles from VERTEX 70 spectra is 1.0. The correlation coefficient between the CH<sub>4</sub> total columns retrieved from 125HR and VERTEX 70 spectra is 0.77. The underestimation of the VERTEX 70-V1-retrieved CH<sub>4</sub> columns in summer is confirmed by the coincident AirCore profiles. In this study, we propose to replace two microwindows of the recommended NDACC retrieval strategy (2835.55-2835.8 and 2903.82–2903.925 cm<sup>-1</sup>) with two alternative ones, namely 2875.15-2875.85 and 2898.3-2898.93 cm<sup>-1</sup>, including stronger CH<sub>4</sub> absorption lines, to perform the CH<sub>4</sub> profile retrievals from the VERTEX 70 spectra (V2 retrieval strategy). The DOF of VERTEX 70 V2 CH<sub>4</sub> retrievals is 1.3, providing more vertical information as compared to VERTEX 70 V1 retrievals. The relative differences between the CH<sub>4</sub> total columns retrieved from 125HR and VER-TEX 70 spectra are  $0.0 \pm 0.8$  %, with a much reduced seasonal variation. The correlation coefficient between the CH<sub>4</sub> total columns retrieved from 125HR and VERTEX 70 spectra increases to 0.87. The relative differences between the CH<sub>4</sub> total columns from the coincident AirCore and VER-TEX 70 V2 retrievals are  $0.13 \pm 0.50$  %.

The NIR vs. MIR and low- vs. high-resolution performance differences comprise an interesting measurement technique topic to characterize, including how performance differences have evolved over time, with various technological and retrieval advancements. In the framework of the ESA FRM4GHG 2.0 project, we intend to design a compact solar tracker system and enclosure to make the Bruker VER-TEX 70 even more convenient to transport and operate autonomously. More campaigns will be carried out to better understand the performances of the CH<sub>4</sub> and N<sub>2</sub>O VERTEX 70 retrievals with different conditions, such as at a humid or a low-latitude site. Such kind of portable FTIR instruments will help to fill the gaps of the NDACC network for CH<sub>4</sub> and N<sub>2</sub>O observations and contribute effectively to satellite validation, model verification, and other scientific studies. In the future, it is also planned to investigate the retrieval of additional tropospheric species from the Bruker VERTEX 70 low-resolution MIR spectra.

*Data availability.* The Bruker 125HR retrievals at Sodankylä are publicly available at the NDACC archive (http://www.ndacc.org, InfraRed Working Group, 2023). The VERTEX 70 retrievals and AirCore profiles are available upon request to the authors.

*Author contributions.* MZ wrote the manuscript. MZ, MKS, BL, JN, and MDM designed the experiment. CH, RK, PH, and CP contributed to the FTIR measurements at Sodankylä. RK and HC provided the AirCore measurements. All the authors read and commented on the manuscript.

*Competing interests.* At least one of the (co-)authors is a member of the editorial board of *Atmospheric Measurement Techniques*. The peer-review process was guided by an independent editor, and the authors also have no other competing interests to declare.

*Disclaimer.* Publisher's note: Copernicus Publications remains neutral with regard to jurisdictional claims made in the text, published maps, institutional affiliations, or any other geographical representation in this paper. While Copernicus Publications makes every effort to include appropriate place names, the final responsibility lies with the authors.

Acknowledgements. The authors would like to thank all FRM4GHG and FRM4GHG 2.0 colleagues for the measurement campaign and discussions in between.

*Financial support.* This study is supported by the National Natural Science Foundation of China (42205140) and the European Space Agency projects FRM4GHG and FRM4GHG 2.0, which received research funding from ESA's FRM program under grant agreement nos. 4000117640/16/I-LG and 4000136108/21/I-DT-Ir.

*Review statement.* This paper was edited by Dietrich G. Feist and reviewed by Dietrich G. Feist and one anonymous referee.

## References

- Angelbratt, J., Mellqvist, J., Blumenstock, T., Borsdorff, T., Brohede, S., Duchatelet, P., Forster, F., Hase, F., Mahieu, E., Murtagh, D., Petersen, A. K., Schneider, M., Sussmann, R., and Urban, J.: A new method to detect long term trends of methane (CH<sub>4</sub>) and nitrous oxide (N<sub>2</sub>O) total columns measured within the NDACC ground-based high resolution solar FTIR network, Atmos. Chem. Phys., 11, 6167–6183, https://doi.org/10.5194/acp-11-6167-2011, 2011.
- Bader, W., Bovy, B., Conway, S., Strong, K., Smale, D., Turner, A. J., Blumenstock, T., Boone, C., Collaud Coen, M., Coulon, A., Garcia, O., Griffith, D. W. T., Hase, F., Hausmann, P., Jones, N., Krummel, P., Murata, I., Morino, I., Nakajima, H., O'Doherty, S., Paton-Walsh, C., Robinson, J., Sandrin, R., Schneider, M.,

Servais, C., Sussmann, R., and Mahieu, E.: The recent increase of atmospheric methane from 10 years of ground-based NDACC FTIR observations since 2005, Atmos. Chem. Phys., 17, 2255–2277, https://doi.org/10.5194/acp-17-2255-2017, 2017.

- Barret, B., Gouzenes, Y., Le Flochmoen, E., and Ferrant, S.: Retrieval of Metop-A/IASI N<sub>2</sub>O Profiles and Validation with NDACC FTIR Data, Atmosphere, 12, 219, https://doi.org/10.3390/atmos12020219, 2021.
- Chesnokova, T., Makarova, M., Chentsov, A., Kostsov, V., Poberovskii, A., Zakharov, V., and Rokotyan, N.: Estimation of the impact of differences in the CH<sub>4</sub> absorption line parameters on the accuracy of methane atmospheric total column retrievals from ground-based FTIR spectra, J. Quant. Spectrosc. Ra., 254, 107187, https://doi.org/10.1016/j.jqsrt.2020.107187, 2020.
- De Mazière, M., Thompson, A. M., Kurylo, M. J., Wild, J. D., Bernhard, G., Blumenstock, T., Braathen, G. O., Hannigan, J. W., Lambert, J.-C., Leblanc, T., McGee, T. J., Nedoluha, G., Petropavlovskikh, I., Seckmeyer, G., Simon, P. C., Steinbrecht, W., and Strahan, S. E.: The Network for the Detection of Atmospheric Composition Change (NDACC): history, status and perspectives, Atmos. Chem. Phys., 18, 4935–4964, https://doi.org/10.5194/acp-18-4935-2018, 2018.
- Fox, J.: Applied Regression Analysis and Generalized Linear Models, SAGE Publications, http://books.google.ch/books?id= GKkn3LSSHFsC (last access: 1 January 2023), 2008.
- García, O. E., Schneider, M., Sepúlveda, E., Hase, F., Blumenstock, T., Cuevas, E., Ramos, R., Gross, J., Barthlott, S., Röhling, A. N., Sanromá, E., González, Y., Gómez-Peláez, Á. J., Navarro-Comas, M., Puentedura, O., Yela, M., Redondas, A., Carreño, V., León-Luis, S. F., Reyes, E., García, R. D., Rivas, P. P., Romero-Campos, P. M., Torres, C., Prats, N., Hernández, M., and López, C.: Twenty years of ground-based NDACC FTIR spectrometry at Izaña Observatory – overview and long-term comparison to other techniques, Atmos. Chem. Phys., 21, 15519–15554, https://doi.org/10.5194/acp-21-15519-2021, 2021.
- Gordon, I., Rothman, L., Hill, C., Kochanov, R., Tan, Y., Bernath, P., Birk, M., Boudon, V., Campargue, A., Chance, K., Drouin, B., Flaud, J.-M., Gamache, R., Hodges, J., Jacquemart, D., Perevalov, V., Perrin, A., Shine, K., Smith, M.-A., Tennyson, J., Toon, G., Tran, H., Tyuterev, V., Barbe, A., Császár, A., Devi, V., Furtenbacher, T., Harrison, J., Hartmann, J.-M., Jolly, A., Johnson, T., Karman, T., Kleiner, I., Kyuberis, A., Loos, J., Lyulin, O., Massie, S., Mikhailenko, S., Moazzen-Ahmadi, N., Müller, H., Naumenko, O., Nikitin, A., Polyansky, O., Rey, M., Rotger, M., Sharpe, S., Sung, K., Starikova, E., Tashkun, S., Auwera, J. V., Wagner, G., Wilzewski, J., Wcisło, P., Yu, S., and Zak, E.: The HITRAN2016 molecular spectroscopic database, J. Quant. Spectrosc. Ra., 203, 3–69, https://doi.org/10.1016/j.jqsrt.2017.06.038, hITRAN2016 Special Issue, 2017.
- Hall, B. D., Dutton, G. S., and Elkins, J. W.: The NOAA nitrous oxide standard scale for atmospheric observations, J. Geophys. Res. Atmos., 112, D09305, https://doi.org/10.1029/2006JD007954, 2007.
- Hase, F., Hannigan, J., Coffey, M., Goldman, A., Höpfner, M., Jones, N., Rinsland, C., and Wood, S.: Intercomparison of retrieval codes used for the analysis of high-resolution, groundbased FTIR measurements, J. Quant. Spectrosc. Ra., 87, 25–52, https://doi.org/10.1016/j.jqsrt.2003.12.008, 2004.

- Herman, R. L., Scott, D. C., Webster, C. R., May, R. D., Moyer, E. J., Salawitch, R. J., Yung, Y. L., Toon, G. C., Sen, B., Margitan, J. J., Rosenlof, K. H., Michelsen, H. A., and Elkins, J. W.: Tropical entrainment time scales inferred from stratospheric N<sub>2</sub>O and CH<sub>4</sub> observations, Geophys. Res. Lett., 25, 2781–2784, https://doi.org/10.1029/98GL02109, 1998.
- InfraRed Working Group: NDACC datasets, NDACC, http://www. ndacc.org, last access: 1 January 2023.
- IPCC: Climate change 2013: The physical science basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Sweeney, Karion. A., C., Tans, P., and Newberger. AirCore: T.: An innovative atmospheric sampling system, J. Atmos. Ocean. Tech., 27, 1839-1853, https://doi.org/10.1175/2010JTECHA1448.1, 2010.
- Kirschke, S., Bousquet, P., Ciais, P., Saunois, M., Canadell, J. G., Dlugokencky, E. J., Bergamaschi, P., Bergmann, D., Blake, D. R., Bruhwiler, L., Cameron-Smith, P., Castaldi, S., Chevallier, F., Feng, L., Fraser, A., Heimann, M., Hodson, E. L., Houweling, S., Josse, B., Fraser, P. J., Krummel, P. B., Lamarque, J.-F., Langenfelds, R. L., Le Quéré, C., Naik, V., O'Doherty, S., Palmer, P. I., Pison, I., Plummer, D., Poulter, B., Prinn, R. G., Rigby, M., Ringeval, B., Santini, M., Schmidt, M., Shindell, D. T., Simpson, I. J., Spahni, R., Steele, L. P., Strode, S. A., Sudo, K., Szopa, S., van der Werf, G. R., Voulgarakis, A., van Weele, M., Weiss, R. F., Williams, J. E., and Zeng, G.: Three decades of global methane sources and sinks, Nat. Geosci., 6, pages 813– 823, https://doi.org/10.1038/ngeo1955, 2013.
- Kivi, R. and Heikkinen, P.: Fourier transform spectrometer measurements of column CO<sub>2</sub> at Sodankylä, Finland, Geosci. Instrum. Method. Data Syst., 5, 271–279, https://doi.org/10.5194/gi-5-271-2016, 2016.
- Kivi, R., Kyrö, E., Turunen, T., Harris, N. R., von der Gathen, P., Rex, M., Andersen, S. B., and Wohltmann, I.: Ozonesonde observations in the Arctic during 1989-2003: Ozone variability and trends in the lower stratosphere and free troposphere, J. Geophys. Res.-Atmos., 112, D08306, https://doi.org/10.1029/2006JD007271, 2007.
- Ostler, A., Sussmann, R., Rettinger, M., Deutscher, N. M., Dohe, S., Hase, F., Jones, N., Palm, M., and Sinnhuber, B.-M.: Multistation intercomparison of column-averaged methane from NDACC and TCCON: impact of dynamical variability, Atmos. Meas. Tech., 7, 4081–4101, https://doi.org/10.5194/amt-7-4081-2014, 2014.
- Park, S., Croteau, P., Boering, K. A., Etheridge, D. M., Ferretti, D., Fraser, P. J., Kim, K. R., Krummel, P. B., Langenfelds, R. L., Van Ommen, T. D., Steele, L. P., and Trudinger, C. M.: Trends and seasonal cycles in the isotopic composition of nitrous oxide since 1940, Nat. Geosci., 5, 261–265, https://doi.org/10.1038/ngeo1421, 2012.
- Pougatchev, N. S., Connor, B. J., and Rinsland, C. P.: Infrared measurements of the ozone vertical distribution above Kitt Peak, J. Geophys. Res., 100, 16689, https://doi.org/10.1029/95JD01296, 1995.
- Rodgers, C. D.: Inverse Methods for Atmospheric Sounding Theory and Practice, Series on Atmospheric Oceanic and Planetary Physics, vol. 2, World Scientific Publishing Co. Pte. Ltd, Singapore, https://doi.org/10.1142/9789812813718, 2000.

- Rodgers, C. D. and Connor, B. J.: Intercomparison of remote sounding instruments, J. Geophys. Res., 108, 46–48, https://doi.org/10.1029/2002JD002299, 2003.
- Rothman, L. S., Gordon, I. E., Barbe, A., Benner, D. C., Bernath, P. F., Birk, M., Boudon, V., Brown, L. R., Campargue, A., Champion, J. P., Chance, K., Coudert, L. H., Dana, V., Devi, V. M., Fally, S., Flaud, J. M., Gamache, R. R., Goldman, A., Jacquemart, D., Kleiner, I., Lacome, N., Lafferty, W. J., Mandin, J. Y., Massie, S. T., Mikhailenko, S. N., Miller, C. E., Moazzen-Ahmadi, N., Naumenko, O. V., Nikitin, A. V., Orphal, J., Perevalov, V. I., Perrin, A., Predoi-Cross, A., Rinsland, C. P., Rotger, M., Šimečková, M., Smith, M. A., Sung, K., Tashkun, S. A., Tennyson, J., Toth, R. A., Vandaele, A. C., and Vander Auwera, J.: The HITRAN 2008 molecular spectroscopic database, J. Quant. Spectrosc. Ra., 110, 533–572, https://doi.org/10.1016/j.jqsrt.2009.02.013, 2009.
- Saunois, M., Stavert, A. R., Poulter, B., Bousquet, P., Canadell, J. G., Jackson, R. B., Raymond, P. A., Dlugokencky, E. J., Houweling, S., Patra, P. K., Ciais, P., Arora, V. K., Bastviken, D., Bergamaschi, P., Blake, D. R., Brailsford, G., Bruhwiler, L., Carlson, K. M., Carrol, M., Castaldi, S., Chandra, N., Crevoisier, C., Crill, P. M., Covey, K., Curry, C. L., Etiope, G., Frankenberg, C., Gedney, N., Hegglin, M. I., Höglund-Isaksson, L., Hugelius, G., Ishizawa, M., Ito, A., Janssens-Maenhout, G., Jensen, K. M., Joos, F., Kleinen, T., Krummel, P. B., Langenfelds, R. L., Laruelle, G. G., Liu, L., Machida, T., Maksyutov, S., McDonald, K. C., McNorton, J., Miller, P. A., Melton, J. R., Morino, I., Müller, J., Murguia-Flores, F., Naik, V., Niwa, Y., Noce, S., O'Doherty, S., Parker, R. J., Peng, C., Peng, S., Peters, G. P., Prigent, C., Prinn, R., Ramonet, M., Regnier, P., Riley, W. J., Rosentreter, J. A., Segers, A., Simpson, I. J., Shi, H., Smith, S. J., Steele, L. P., Thornton, B. F., Tian, H., Tohjima, Y., Tubiello, F. N., Tsuruta, A., Viovy, N., Voulgarakis, A., Weber, T. S., van Weele, M., van der Werf, G. R., Weiss, R. F., Worthy, D., Wunch, D., Yin, Y., Yoshida, Y., Zhang, W., Zhang, Z., Zhao, Y., Zheng, B., Zhu, Q., Zhu, Q., and Zhuang, Q.: The Global Methane Budget 2000-2017, Earth Syst. Sci. Data, 12, 1561-1623, https://doi.org/10.5194/essd-12-1561-2020, 2020.
- Sawa, Y., Machida, T., Matsueda, H., Niwa, Y., Tsuboi, K., Murayama, S., Morimoto, S., and Aoki, S.: Seasonal changes of CO<sub>2</sub>, CH<sub>4</sub>, N<sub>2</sub>O, and SF<sub>6</sub> in the upper troposphere/lower stratosphere over the Eurasian continent observed by commercial airliner, Geophys. Res. Lett., 42, 2001–2008, https://doi.org/10.1002/2014GL062734, 2015.
- Schwietzke, S., Sherwood, O. A., Bruhwiler, L. M. P., Miller, J. B., Etiope, G., Dlugokencky, E. J., Michel, S. E., Arling, V. A., Vaughn, B. H., White, J. W. C., and Tans, P. P.: Upward revision of global fossil fuel methane emissions based on isotope database, Nature, 538, 88–91, https://doi.org/10.1038/nature19797, 2016.
- Sepúlveda, E., Schneider, M., Hase, F., Barthlott, S., Dubravica, D., García, O. E., Gomez-Pelaez, A., González, Y., Guerra, J. C., Gisi, M., Kohlhepp, R., Dohe, S., Blumenstock, T., Strong, K., Weaver, D., Palm, M., Sadeghi, A., Deutscher, N. M., Warneke, T., Notholt, J., Jones, N., Griffith, D. W. T., Smale, D., Brailsford, G. W., Robinson, J., Meinhardt, F., Steinbacher, M., Aalto, T., and Worthy, D.: Tropospheric CH<sub>4</sub> signals as observed by NDACC FTIR at globally distributed sites and comparison to

GAW surface in situ measurements, Atmos. Meas. Tech., 7, 2337–2360, https://doi.org/10.5194/amt-7-2337-2014, 2014.

- Sha, M. K., De Mazière, M., Notholt, J., Blumenstock, T., Chen, H., Dehn, A., Griffith, D. W. T., Hase, F., Heikkinen, P., Hermans, C., Hoffmann, A., Huebner, M., Jones, N., Kivi, R., Langerock, B., Petri, C., Scolas, F., Tu, Q., and Weidmann, D.: Intercomparison of low- and high-resolution infrared spectrometers for groundbased solar remote sensing measurements of total column concentrations of CO<sub>2</sub>, CH<sub>4</sub>, and CO, Atmos. Meas. Tech., 13, 4791–4839, https://doi.org/10.5194/amt-13-4791-2020, 2020.
- Sha, M. K., Langerock, B., Blavier, J.-F. L., Blumenstock, T., Borsdorff, T., Buschmann, M., Dehn, A., De Mazière, M., Deutscher, N. M., Feist, D. G., García, O. E., Griffith, D. W. T., Grutter, M., Hannigan, J. W., Hase, F., Heikkinen, P., Hermans, C., Iraci, L. T., Jeseck, P., Jones, N., Kivi, R., Kumps, N., Landgraf, J., Lorente, A., Mahieu, E., Makarova, M. V., Mellqvist, J., Metzger, J.-M., Morino, I., Nagahama, T., Notholt, J., Ohyama, H., Ortega, I., Palm, M., Petri, C., Pollard, D. F., Rettinger, M., Robinson, J., Roche, S., Roehl, C. M., Röhling, A. N., Rousogenous, C., Schneider, M., Shiomi, K., Smale, D., Stremme, W., Strong, K., Sussmann, R., Té, Y., Uchino, O., Velazco, V. A., Vigouroux, C., Vrekoussis, M., Wang, P., Warneke, T., Wizenberg, T., Wunch, D., Yamanouchi, S., Yang, Y., and Zhou, M.: Validation of methane and carbon monoxide from Sentinel-5 Precursor using TCCON and NDACC-IRWG stations, Atmos. Meas. Tech., 14, 6249-6304, https://doi.org/10.5194/amt-14-6249-2021, 2021.
- Steck, T.: Methods for determining regularization for atmospheric retrieval problems, Appl. Optics, 41, 1788–1797, https://doi.org/10.1364/AO.41.001788, 2002.
- Taylor, J. R., Wunch, D., Midwinter, C., Wiacek, A., Drummond, J. R., and Strong, K.: An extended intercomparison of simultaneous ground-based Fourier transform infrared spectrometer measurements at the Toronto Atmospheric Observatory, J. Quant. Spectrosc. Ra., 109, 2244–2260, https://doi.org/10.1016/j.jqsrt.2008.03.011, 2008.
- Tian, H., Xu, R., Canadell, J. G., Thompson, R. L., Winiwarter, W., Suntharalingam, P., Davidson, E. A., Ciais, P., Jackson, R. B., Janssens-Maenhout, G., Prather, M. J., Regnier, P., Pan, N., Pan, S., Peters, G. P., Shi, H., Tubiello, F. N., Zaehle, S., Zhou, F., Arneth, A., Battaglia, G., Berthet, S., Bopp, L., Bouwman, A. F., Buitenhuis, E. T., Chang, J., Chipperfield, M. P., Dangal, S. R. S., Dlugokencky, E., Elkins, J. W., Eyre, B. D., Fu, B., Hall, B., Ito, A., Joos, F., Krummel, P. B., Landolfi, A., Laruelle, G. G., Lauerwald, R., Li, W., Lienert, S., Maavara, T., MacLeod, M., Millet, D. B., Olin, S., Patra, P. K., Prinn, R. G., Raymond, P. A., Ruiz, D. J., van der Werf, G. R., Vuichard, N., Wang, J., Weiss, R. F., Wells, K. C., Wilson, C., Yang, J., and Yao, Y.: A comprehensive quantification of global nitrous oxide sources and sinks, Nature, 586, 248–256, 2020.
- Tikhonov, A. N.: Solution of Incorrectly Formulated Problems and the Regularisation Method, Soviet. Math. Dokl., 4, 1035–1038, 1963.
- Tu, Q., Hase, F., Blumenstock, T., Kivi, R., Heikkinen, P., Sha, M. K., Raffalski, U., Landgraf, J., Lorente, A., Borsdorff, T., Chen, H., Dietrich, F., and Chen, J.: Intercomparison of atmospheric CO<sub>2</sub> and CH<sub>4</sub> abundances on regional scales in boreal areas using Copernicus Atmosphere Monitoring Service (CAMS) analysis, COllaborative Carbon Column Observing Network (COCCON) spectrometers, and Sentinel-5 Precur-

sor satellite observations, Atmos. Meas. Tech., 13, 4751–4771, https://doi.org/10.5194/amt-13-4751-2020, 2020.

- Turner, A. J., Frankenberg, C., Wennberg, P. O., and Jacob, D. J.: Ambiguity in the causes for decadal trends in atmospheric methane and hydroxyl, P. Natl. Acad. Sci. USA, 114, 5367–5372, https://doi.org/10.1073/PNAS.1616020114, 2017.
- Wang, Z., Deutscher, N. M., Warneke, T., Notholt, J., Dils, B., Griffith, D. W. T., Schmidt, M., Ramonet, M., and Gerbig, C.: Retrieval of tropospheric column-averaged CH<sub>4</sub> mole fraction by solar absorption FTIR-spectrometry using N<sub>2</sub>O as a proxy, Atmos. Meas. Tech., 7, 3295–3305, https://doi.org/10.5194/amt-7-3295-2014, 2014.
- WMO: WMO Greenhouse Gas Bulletin, World Meteorological Organization, WMO Bull., ISSN 2078-0796, 2020.
- Worden, J. R., Bloom, A. A., Pandey, S., Jiang, Z., Worden, H. M., Walker, T. W., Houweling, S., and Röckmann, T.: Reduced biomass burning emissions reconcile conflicting estimates of the post-2006 atmospheric methane budget, Nat. Commun., 8, 2227, https://doi.org/10.1038/s41467-017-02246-0, 2017.
- Wunch, D., Toon, G. C., Sherlock, V., Deutscher, N. M., Liu, C., Feist, D. G., and Wennberg, P. O.: The Total Carbon Column Observing Network's GGG2014 Data Version, Tech. rep., California Institute of Technology, Pasadena, CA, https://doi.org/10. 14291/tccon.ggg2014.documentation.R0/1221662, 2015.
- Zhou, M., Vigouroux, C., Langerock, B., Wang, P., Dutton, G., Hermans, C., Kumps, N., Metzger, J.-M., Toon, G., and De Mazière, M.: CFC-11, CFC-12 and HCFC-22 ground-based remote sensing FTIR measurements at Réunion Island and comparisons with MIPAS/ENVISAT data, Atmos. Meas. Tech., 9, 5621–5636, https://doi.org/10.5194/amt-9-5621-2016, 2016.
- Zhou, M., Langerock, B., Vigouroux, C., Sha, M. K., Ramonet, M., Delmotte, M., Mahieu, E., Bader, W., Hermans, C., Kumps, N., Metzger, J.-M., Duflot, V., Wang, Z., Palm, M., and De Mazière, M.: Atmospheric CO and CH<sub>4</sub> time series and seasonal variations on Reunion Island from ground-based in situ and FTIR (NDACC and TCCON) measurements, Atmos. Chem. Phys., 18, 13881–13901, https://doi.org/10.5194/acp-18-13881-2018, 2018.
- Zhou, M., Langerock, B., Sha, M. K., Kumps, N., Hermans, C., Petri, C., Warneke, T., Chen, H., Metzger, J.-M., Kivi, R., Heikkinen, P., Ramonet, M., and De Mazière, M.: Retrieval of atmospheric CH<sub>4</sub> vertical information from ground-based FTS near-infrared spectra, Atmos. Meas. Tech., 12, 6125–6141, https://doi.org/10.5194/amt-12-6125-2019, 2019a.
- Zhou, M., Langerock, B., Wells, K. C., Millet, D. B., Vigouroux, C., Sha, M. K., Hermans, C., Metzger, J.-M., Kivi, R., Heikkinen, P., Smale, D., Pollard, D. F., Jones, N., Deutscher, N. M., Blumenstock, T., Schneider, M., Palm, M., Notholt, J., Hannigan, J. W., and De Mazière, M.: An intercomparison of total column-averaged nitrous oxide between ground-based FTIR TC-CON and NDACC measurements at seven sites and comparisons with the GEOS-Chem model, Atmos. Meas. Tech., 12, 1393– 1408, https://doi.org/10.5194/amt-12-1393-2019, 2019b.
- Zhou, M., Ni, Q., Cai, Z., Langerock, B., Jiang, J., Che, K., Wang, J., Nan, W., Liu, Y., and Wang, P.: Ground-Based Atmospheric CO<sub>2</sub>, CH<sub>4</sub>, and CO Column Measurements at Golmud in the Qinghai Tibetan Plateau and Comparisons with TROPOMI/S5P Satellite Observations, Adv. Atmos. Sci., 40, 223–234, https://doi.org/10.1007/s00376-022-2116-0, 2023.