



Benchmarking K_{DP} in rainfall: a quantitative assessment of estimation algorithms using C-band weather radar observations

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Abstract. Accurate and precise K_{DP} estimates are essential for radar-based applications, especially in quantitative precipitation estimation and radar data quality control routines. The accuracy of these estimates largely depends on the post-processing of the radar's measured Φ_{DP} , which aims to reduce noise and backscattering effects while preserving fine-scale precipitation features. In this study, we evaluate the performance of several publicly available K_{DP} estimation methods implemented in open-source libraries such as Py-ART (the Python ARM (atmospheric radiation measurement) Radar Toolkit) and *wradlib* and the method used in the Vaisala weather radars. To benchmark these methods, we employ a polarimetric self-consistency approach that relates K_{DP} to reflectivity and differential reflectivity in rain, providing a reference self-consistent K_{DP} (K_{DP}^{sc}) for comparison. This approach allows for the construction of the reference K_{DP} observations that can be used to assess the accuracy and robustness of the studied K_{DP} estimation methods. We assess each method by quantifying uncertainties using C-band weather radar observations, where the reflectivity values ranged between 20 and 50 dBZ.

Using the proposed evaluation framework, we were able to define optimized parameter settings for the methods that have user-configurable parameters. Most of these methods showed a significant reduction in the estimation errors after the optimization, with respect to the default settings. We have found significant differences in the performance of the studied methods, where the best-performing methods showed smaller normalized biases in the high reflectivity values (i.e., ≥ 40 dBZ) and overall smaller normalized root-mean-square errors across the range of reflectivity values.

1 Introduction

The specific differential phase (K_{DP}) plays an important role in many weather radar applications, particularly in hydrometeor classification (Höller et al., 1994; Vivekanandan et al., 1999; Liu and Chandrasekar, 2000; Zrnić et al., 2001; Keenan, 2003; Lim et al., 2005; Tessendorf et al., 2005; Marzano et al., 2007; Dolan and Rutledge, 2009; Park et al., 2009; Snyder et al., 2010; Al-Sakka et al., 2013; Dolan et al., 2013; Thompson et al., 2014; Bechini and Chandrasekar, 2015; Grazioli et al., 2015; Wen et al., 2015; Besic et al., 2016; Ribaud et al., 2019) and quantitative precipitation estimation (QPE) (Sachidananda and Zrnić, 1987; Chandrasekar et al., 1990; Ryzhkov and Zrnić, 1995, 1996; May et al., 1999; Bringi and Chandrasekar, 2001; Bringi et al., 2006; Matrosov et al., 2006; Giangrande and Ryzhkov, 2008; Bringi et al., 2011; Cifelli et al., 2011; Wang et al., 2013; Figueras i Ventura and Tabary, 2013; Chen and Chandrasekar, 2015; Chen et al., 2017; Thompson et al., 2018; Zhang et al., 2020), and is used in data assimilation for numerical weather prediction models (Thomas et al., 2020; Du et al., 2021) and in hydrological applications (Brandes et al., 2002; Ryzhkov et al., 2005a; Vulpiani et al., 2012; Li et al., 2023; Cremonini et al., 2023). Compared to radar power variables; i.e., the reflectivity factor at horizontal polarization (Z_H) and differential reflectivity (Z_{dr}), K_{DP} offers advantages in terms of accuracy, resilience, and reliability due to its immunity to radar miscalibration, attenuation (Bringi and Chandrasekar, 2001; Illingworth, 2004; Ryzhkov and Zrnić, 2019), and partial beam blockage (Zrnić and Ryzhkov, 1996). It has also proven successful in hydrometeor classifi-

cation routines (Lim et al., 2005; Park et al., 2009; Grazioli et al., 2015; Tiira and Moisseev, 2020), especially in the detection of graupel (Dolan and Rutledge, 2009; Oue et al., 2015) and small melting hail (Kumjian et al., 2019), and in the dendritic growth zone and the processes within (Kennedy and Rutledge, 2011; Andrić et al., 2013; Schneebeli et al., 2013; Moisseev et al., 2015; Kumjian and Lombardo, 2017). The ability of K_{DP} to accurately estimate heavy rainfall, differentiate hydrometeor types, and overcome attenuation in precipitation makes it an invaluable operational and research radar variable.

Despite its advantages, accurate estimation of K_{DP} from the radar-measured differential phase (Φ_{DP}) remains challenging. Mathematically, K_{DP} is half of the range derivative of Φ_{DP} , which measures the phase shift between horizontally and vertically polarized signals as they propagate through precipitation. This phase shift (Φ_{DP}) is influenced by hydrometeor concentration, shape, orientation, and composition (Kumjian, 2018). However, Φ_{DP} is not typically smooth and does not monotonically increase along the rain path; it contains fluctuations due to noise (ϵ) and the backscattering differential phase (δ_{HV}) (Ryzhkov and Zrnić, 1996; Ryzhkov and Zrnić, 1998). Excessive filtering of Φ_{DP} to remove ϵ can lead to the loss of fine-scale precipitation features, affecting the accuracy of K_{DP} estimates, especially in light precipitation (Huang et al., 2017). In heavier precipitation, δ_{HV} causes spikes in Φ_{DP} , especially at higher radar frequencies, further complicating accurate K_{DP} estimation (Bringi and Chandrasekar, 2001).

To address these challenges, various methods have been developed to post-process Φ_{DP} and derive K_{DP} (Hubbert et al., 1993; Hubbert and Bringi, 1995; Ryzhkov et al., 2005c; Wang and Chandrasekar, 2009; Otto and Russchenberg, 2011; Maesaka et al., 2012; Vulpiani et al., 2012; Schneebeli and Berne, 2012; Giangrande et al., 2013; Schneebeli et al., 2014; Huang et al., 2017; Reinoso-Rondinel et al., 2018; Wen et al., 2019). Basic approaches include median filters and moving windows, while more advanced methods use regression techniques and self-consistency constraints based on Z_H or Z_{dr} . Many of these methods are now available in open-source Python libraries such as the Python ARM (atmospheric radiation measurement) Radar Toolkit (Py-ART; Helmus and Collis, 2016) and *wradlib* (Heistermann et al., 2013). For this study, some of the most popular methods based on Maesaka et al. (2012), Vulpiani et al. (2012), Giangrande et al. (2013), and Schneebeli et al. (2014) were selected for analysis. Additionally, the K_{DP} product implemented by Vaisala in the IRIS software (Vaisala, 2017), based on Wang and Chandrasekar (2009), was also included in our analysis. Each algorithm has its own data requirements, mathematical approach, and optimizing parameters, raising the question of which method performs optimally under varying parameter settings and rainfall intensities.

Recent studies show that K_{DP} estimates can vary significantly depending on the algorithm and the optimizing param-

eters used. Reimel and Kumjian (2021) evaluated the errors in several methods using synthetic K_{DP} profiles and found that no single algorithm was optimal across all rainfall conditions. Instead, performance varied according to the complexity of the rain profile and the parameters selected. They identified *kdp_maesaka* (Py-ART's implementation of the Maesaka et al., 2012, method) and *phase_proc_lp* (Py-ART's implementation of the Giangrande et al., 2013, method) as particularly versatile. However, Reimel and Kumjian (2021) used synthetic data, which may miss some of the effects present in radar observations of rainfall (e.g., δ_{HV}). More recently, Li et al. (2023) compared *kdp_maesaka* and *phase_proc_lp* in an extreme summer rainfall event, finding that fine-tuning the methods played a key role in retrieving the most accurate K_{DP} estimate. Despite these insights, the performance of and uncertainties in most methods of rainfall observations remain largely unexplored.

The goal of this study is to evaluate the performance of publicly available K_{DP} estimation methods on real rainfall observations and quantify their uncertainties as a function of reflectivity intensities. To achieve this, we employ a benchmarking K_{DP} , herein K_{DP}^{sc} , computed from measured Z_H and Z_{dr} , and use self-consistency relations in rain. In rainfall observations, the polarimetric radar variables are not independent, but one can be computed in terms of the others via the self-consistency relations (Aydin et al., 1987; Scarchilli et al., 1993). These relations have proven successful in hydrometeor classification (Aydin and Giridhar, 1992) and radar calibration correction (Gorgucci et al., 1992) routines. For instance, Aydin and Giridhar (1992) showed that the hydrometeors can be classified based on their proximity to clusters around self-consistency curves between polarimetric variables. At nearly the same time, Gorgucci et al. (1992) noted the self-consistency of Z_H , Z_{dr} , and K_{DP} in rainfall and proposed a method to calibrate Z_H and correct Z_H -rainfall estimates, benchmarking against K_{DP} -rainfall estimates. Thereafter, several methods linking the polarimetric variables via self-consistency relations have been widely used to calibrate Z_H (Goddard et al., 1994; Scarchilli et al., 1996; Vivekanandan et al., 2003; Ryzhkov et al., 2005b; Gourley et al., 2009). In this study, K_{DP}^{sc} is computed using the consistency relation linking K_{DP} to Z_H and Z_{dr} , which was first noted by Goddard et al. (1994) and described in Gourley et al. (2009), requiring thorough selection and filtering of Z_H and Z_{dr} . K_{DP}^{sc} computed from quality controlled Z_H and Z_{dr} measurements provides a solid benchmark against which to compare the methods' performance, to select optimal parameters, and to quantify the associated uncertainties.

This paper is organized as follows. Section 2 describes the radar and disdrometer data, shows the evaluation framework, and introduces the K_{DP} estimation methods. Section 3 presents and discusses the parameter optimization and performance evaluation of the methods, and Sect. 4 summarizes the findings.

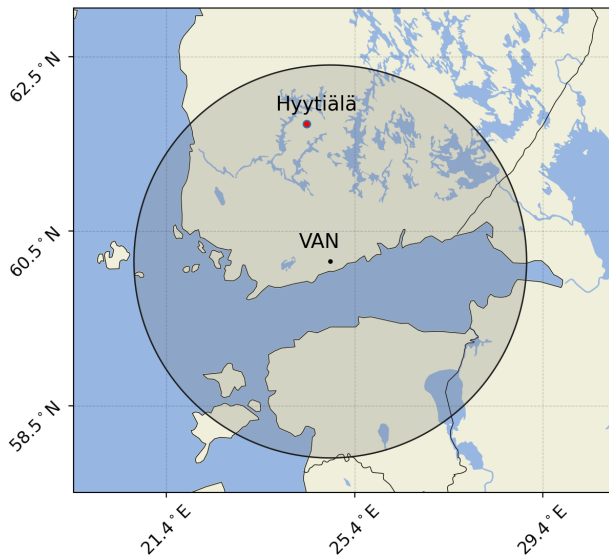


Figure 1. Map showing the locations of FMI's Vantaa radar (VAN) and Hyytiälä's research station where the drop size distribution (DSD) data were collected. The shaded area is a circle with a 250 km radius corresponding to the spatial coverage of the radar.

2 Data and methods

2.1 Radar and disdrometer data

This study evaluates the performance of K_{DP} estimation methods using real rainfall data. The dataset was collected from the Finnish Meteorological Institute (FMI) C-band Vantaa radar, located near Helsinki, Finland (see Fig. 1). The radar recorded various quantities, including Z_H , Z_{dr} , Φ_{DP} , K_{DP} , the cross-correlation coefficient (ρ_{HV}), and the hydrometeor classification product available in IRIS (Vaisala, 2017) and based on the methodology described by Chandrasekar et al. (2013). The spatial resolution of the radar is 500 m in range and 1° in azimuth, with scans performed every 5 min, and the data were collected with an elevation angle of 0.7° . The dataset spans June to September during the years 2017 to 2019, capturing precipitation events with variable rainfall intensities and spatial extents. The raw radar dataset as well as the post-processed K_{DP} estimates are available from the link provided in Aldana (2024).

To ensure data quality, only periods when the Vantaa radar had calibration errors within 1 dB were selected. The calibration was verified by (i) identifying periods where solar flux estimates from Vantaa radar estimates aligned consistently with Dominion Radio Astrophysical Observatory (DRAO) estimates (Huuskonen and Holleman, 2007; Tapping, 2013; Holleman et al., 2022) and (ii) selecting radar scans within the periods where Z_H -calibration offsets were within 1 dB, following the absolute calibration procedure outlined by Gourley et al. (2009). Z_{dr} bias was estimated and corrected during these periods by computing the offset

between observed and self-consistent Z_{dr} derived from observed Z_H , as described in Hickman (2015), and by computing the average for several cases.

The performance of the K_{DP} estimation methods is benchmarked against the self-consistent K_{DP}^{sc} computed from measured Z_H and Z_{dr} and by using self-consistency relations in rain. The self-consistency relations, which link the polarimetric radar variables, were derived by fitting radar variables computed using the open-source library, PyTMatrix (Leinonen, 2014). PyTMatrix provides a simple interface for T-matrix electromagnetic scattering calculations (Waterman, 1965; Mishchenko et al., 2000), requiring the user to provide drop size distribution (DSD) data and setting parameters such as temperature, the radar wavelength band, and the raindrop shape model. The parameters used for the T-matrix calculations were 10°C , C-band, and Thurai et al. (2007), respectively, and the DSD data provided were collected by an optical Parsivel disdrometer (Moisseev, 2024) located in Hyytiälä, Finland (see Fig. 1).

The Parsivel disdrometer records the number of particles and their velocity at 1 min intervals, sorting the data into 32 bins depending on the particle's size (i.e., equivalent volume diameter) and 32 additional bins depending on the particle's fall velocity. From the number of particles and the size and velocity classes, the Parsivel disdrometer computes the precipitation type, which was used to filter out non-liquid particles. Observations were further limited to times when the 30 min average 2 m temperature exceeded 2°C to ensure liquid rain. Following the filtering procedure suggested by Leinonen et al. (2012) to reduce statistical errors, only those measurements with at least 100 counts in two consecutive bins and positive counts in at least four consecutive bins were retained. The disdrometer dataset, covering June to September from 2014 to 2019, provided a robust basis for deriving average summer-season DSD parameters such as the mean volume diameter (D_0) and intercept (N_w) and shape (μ) parameters. These parameters showed strong agreement with those reported by Leinonen et al. (2012) in a climatological study of Finland. From the derived DSD parameters (N_w , D_0 , and μ), the polarimetric radar variables were computed and used to derive the self-consistency relation, defining the framework to evaluate the K_{DP} estimation methods.

2.2 K_{DP} evaluation framework

The performance of the K_{DP} estimation methods is evaluated using K_{DP}^{sc} as a benchmark. This quantity is calculated from each radar-measured tuple (Z_H , Z_{dr}), following a relationship of the form (Goddard et al., 1994; Illingworth and Blackman, 2002; Gourley et al., 2009)

$$K_{DP}^{sc} = z_H \times 10^{-5} \times (a_1 + a_2 \times Z_{dr} + a_3 \times Z_{dr}^2 + a_4 \times Z_{dr}^3), \quad (1)$$

where $z_H = 10^{0.1 \times Z_H}$ represents Z_H in linear units ($\text{mm}^6 \text{m}^{-3}$), and Z_{dr} is in decibels (dB). The coefficients

used in this relation are $a_1 = 6.78$, $a_2 = -2.65$, $a_3 = 0.562$, and $a_4 = -0.0624$. The coefficients align well with those reported by Gourley et al. (2009), which employed the raindrop shape models by Brandes et al. (2002) and Thurai and Bringi (2005).

To ensure the accuracy and robustness of the K_{DP}^{sc} estimates used in the method assessment, it was crucial to quality control the Z_H and Z_{dr} data. Radar observations of rain are often affected by non-meteorological measurements, resonance effects, and hail contamination (Bringi and Chandrasekar, 2001; Kumjian, 2013; Ryzhkov and Zrníc, 2019). To address these issues, the following filtering steps were applied.

- *Noise filtering.* A minimum threshold of 0.97 was applied to ρ_{HV} .
- *Non-meteorological observation filtering.* The hydrometeor classification product from IRIS (Vaisala, 2017), based on Chandrasekar et al. (2013), was used to exclude gates classified as non-meteorological.
- *δ_{HV} reduction.* Gates with $Z_{dr} > 3.5$ dB were excluded (Bringi and Chandrasekar, 2001; Gourley et al., 2009).
- *Non-liquid rain filtering.*
 - Only radar scans from the warm months (June–September) were selected.
 - Gates not classified as rain by the hydrometeor classification product were excluded.
 - Hail contamination was addressed by removing gates with $Z_H \geq 50$ dBZ.
 - Observations from the melting layer and above were suppressed by masking gates further than 70 km (see last dashed ring in Fig. 2) from the radar in the radial direction. The distance was manually set by identifying gates with melting layer signatures (Giangrande et al., 2008; Boodoo et al., 2010).

In addition to addressing noise and non-liquid rain measurements, K_{DP}^{sc} estimates are affected by attenuation in Z_H and differential attenuation in Z_{dr} , particularly in cases of heavy rainfall, of extended propagation paths through rain (hereafter rain paths) (Zrníc and Ryzhkov, 1996; Carey et al., 2000; Bringi and Chandrasekar, 2001; Kumjian, 2013), and when the radar’s antenna radome is wet (Blevins, 1965; Kurri and Huuskonen, 2008). To mitigate these effects, radar scans when there was rain on top of the radar within the past 20 min were discarded. Then, for the remaining cases, attenuation in heavy precipitation or extended rain paths was addressed by flagging the radar gates when suspected attenuation of at least 1 dB was detected. The attenuation in range gates was inferred using a standard method that linearly relates the losses in Z_H and Z_{dr} to increases in $\Delta\Phi_{DP}$ (Ryzhkov and Zrníc, 1995; Carey et al., 2000; Bringi and Chandrasekar,

2001; Gourley et al., 2009). $\Delta\Phi_{DP}$ corresponds to the total span of Φ_{DP} along the radial within a rain path. A rain path was defined as a set of consecutive gates with rain features extending at least 20 km in the radial direction. For C-band radar, a minimum threshold of 12° in $\Delta\Phi_{DP}$ indicates attenuation of at least 1 dB (Carey et al., 2000). In this study, a threshold of 10° was used, meaning that gates within rain paths featuring $\Delta\Phi_{DP} \geq 10^\circ$ were flagged as attenuated.

An example of the filtering procedure applied to a radar scan is shown in Fig. 2. This figure demonstrates the effects of the filtering process and the attenuation considered on the chosen data samples.

Following the filtering process, the dataset comprised 652 624 quality controlled gates from 70 radar scans. Figure 3 presents a histogram of the data proportions across different Z_H values, showing the highest percentage of data between 30 and 35 dBZ, with a sharp decrease from 35 to 50 dBZ. The stacked bars indicate the percentages of attenuated and non-attenuated gates, with the ratio of attenuated to non-attenuated data increasing with greater Z_H .

2.3 K_{DP} estimation methods

This section provides an overview of the K_{DP} estimation methods selected for this study. The selection criteria focused on the availability of these methods in widely used open-source libraries, such as Py-ART (Helmus and Collis, 2016) and *wradlib* (Heistermann et al., 2013). At the time of this study, Py-ART version 1.17.0 included the following methods: *kdp_maesaka*, *kdp_vulpiani*, *phase_proc_lp*, and *kdp_schneebeli*. *wradlib* version 2.0.3 included *kdp_from_phidp* and *phidp_kdp_vulpiani*. However, *phidp_kdp_vulpiani* was excluded from our analysis, as it is based on the same method proposed by Vulpiani et al. (2012) that is already represented in Py-ART by *kdp_vulpiani*. Additionally, *kdp_iris*, a method based on Wang and Chandrasekar (2009) and implemented by Vaisala in the IRIS software (Vaisala, 2017), was included. Table 1 summarizes the key features of the selected methods, and a brief description of the methods is provided below.

- a. *kdp_maesaka*. Developed by Maesaka et al. (2012) and available in Py-ART, this method estimates non-negative K_{DP} from liquid-precipitation measurements. It addresses the issue of negative K_{DP} estimates observed in exclusively liquid-precipitation regions when using classical methods based on iterative filtering and local linear regression. Maesaka et al. (2012) identified that negative K_{DP} were caused by noise in Φ_{DP} during weak precipitation and by δ_{HV} during heavy precipitation. The method restricts K_{DP} to positive values and assumes that Φ_{DP} is a monotonically increasing function with range, which is already unfolded.
- b. *kdp_vulpiani*. Developed by Vulpiani et al. (2012) and available in Py-ART, this method estimates K_{DP} for

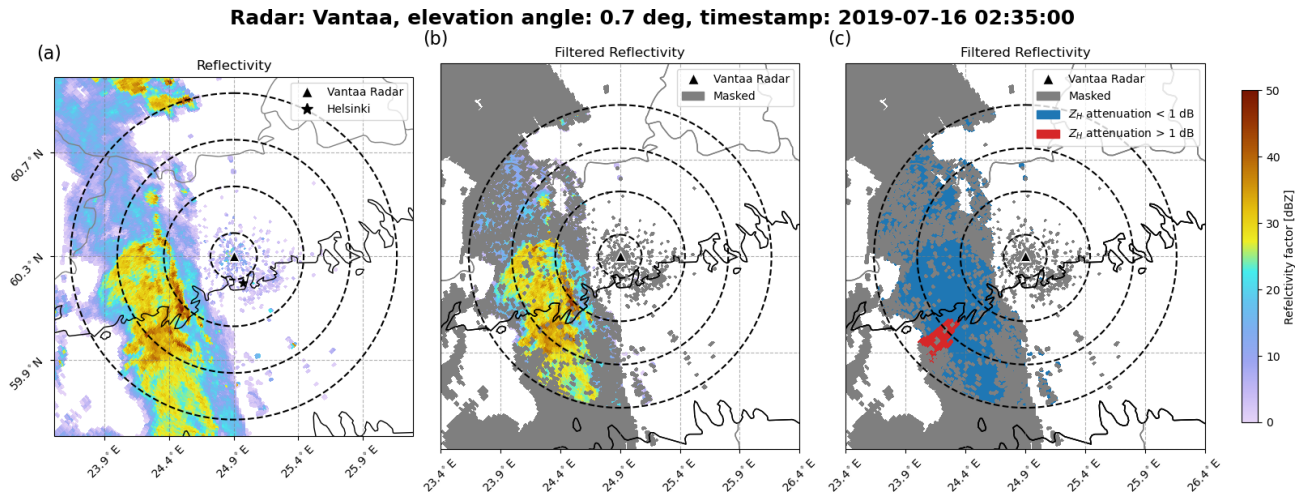


Figure 2. Example of a Vantaa radar scan during a precipitation event on 16 July 2019 at an elevation angle of 0.7° . Panel (a) shows measured Z_H ; panel (b) shows filtered Z_H with masked gates in gray; panel (c) shows the same as panel (b) but with attenuated gates marked in red and non-attenuated gates marked in blue. Dashed rings represent radial distances of 10, 30, 50, and 70 km from the radar.

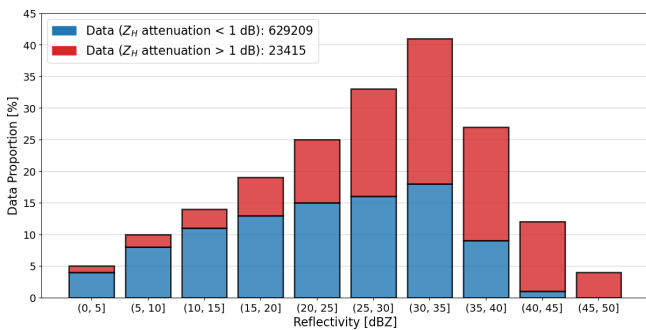


Figure 3. Proportion of data across Z_H intervals of 5 dBZ. Attenuated data are represented by red bars, and non-attenuated data are represented by blue bars. The legend indicates the total number of gates with suspected attenuation of at least 1 dB (red) and less than 1 dB (blue).

any type of precipitation. It uses a multistep moving-window range derivative approach to obtain K_{DP} . It calculates a K_{DP} profile from the range derivative of a noise-reduced, offset-corrected, and unfolded Φ_{DP} profile. At each window, K_{DP} is compared to thresholds representing unrealistic K_{DP} values within precipitation, correcting possible aliasing with the minimum threshold.

c. *phase_proc_lp*. Developed by Giangrande et al. (2013) and available in Py-ART, this method estimates non-negative K_{DP} from liquid-precipitation measurements. It uses a linear-programming (LP) method to enforce monotonic behavior in Φ_{DP} , restricting K_{DP} to positive values. It extracts δ_{HV} from Φ_{DP} , and it uses self-consistency constraints to bound K_{DP} estimates based on measured Z_H . The method requires quality con-

trolled Z_H and allows user-defined thresholds to exclude hail and the setting of the environmental 0°C level to exclude mixed-phase particles.

- d. *kdp_from_phidp*. Implemented in *wradlib* (Heistermann et al., 2013) and based on Vulpiani et al. (2012), this method estimates K_{DP} for any type of precipitation. It computes range-wise differentiation of Φ_{DP} over a user-defined window size length, defaulting to seven gates for a range resolution of 1 km. Unlike *kdp_vulpiani*, it allows the selection of the method for range gate differentiation, albeit without supporting multiple iterations, prioritizing speed over phase unfolding and noise issues in Φ_{DP} .
- e. *kdp_schneebeli*. Developed by Schneebeli et al. (2014) and available in Py-ART, this method estimates K_{DP} for any type of precipitation. It selects the best-averaged K_{DP} profile from forward and backward propagation Kalman-filtered estimates. The Kalman filters are applied twice to each range gate state (accounting for forward and backward propagation) multiple times, recalculating the covariance matrices each time to yield unique states, and the best estimate is selected.
- f. *kdp_iris*. Implemented in the Vaisala software IRIS (Vaisala, 2017) and based on Wang and Chandrasekar (2009), this method estimates K_{DP} for any type of precipitation. It computes K_{DP} adaptively through piecewise regression and a regularization framework that minimizes both smoothness in Φ_{DP} and regression errors. The regularization adapts based on range variations in K_{DP} and ρ_{HV} measurements, preserving steep Φ_{DP} changes in high-intensity precipitation while reducing variations in low-intensity precipitation.

Table 1. List of the K_{DP} methods studied, with key features.

Method	Source	Data pre-requisites	Precipitation type	Mathematical approach (constraints)	Tested parameters
kdp_maesaka	Py-ART	Unfolded ϕ_{DP}	Liquid	Variational	Clpf
kdp_vulpiani	Py-ART	Pre-filtered Ψ_{DP}	Any	Moving window	windsize, n_iter
phase_proc_lp	Py-ART	Unattenuated Z_H	Liquid	Linear programming ($K_{DP}(Z_H)$)	self_const, coef, window_len
kdp_from_phidp	<i>oradlib</i>	No NaN values	Any	Moving window	winlen, dr
kdp_schneebeli	Py-ART	Pre-filtered Ψ_{DP}	Any	Kalman filter	–
kdp_iris	IRIS	–	Any	Adaptive regression	–

3 Results

3.1 Parameter optimization of methods

All the methods except `kdp_iris` are available in open-source libraries and feature user-configurable parameters to improve the K_{DP} estimates. However, two methods are excluded from the optimization: `kdp_schneebeli` and `kdp_iris`. In `kdp_schneebeli`, the error covariance matrices of the measurements (`rcov`) and state transitions (`pcov`) require a large ensemble of stochastic simulated rainfall fields to be derived. Since such information is not available to us, we use the method with default settings. In `kdp_iris`, the end user has no effect on the derivation of K_{DP} . Instead, at the FMI, we use the K_{DP} product as it comes from the IRIS software (Vaisala, 2017). Therefore, the optimization focuses on the `kdp_maesaka`, `kdp_vulpiani`, `phase_proc_lp`, and `kdp_from_phidp` methods, and in this section, we quantify the errors under varying parameter settings and select the optimal values.

First, a qualitative analysis is provided using of K_{DP} vs. Z_H scatterplots, illustrating the relationship between estimated K_{DP} (y axis) and Z_H (x axis) and benchmarking against K_{DP}^{sc} (dashed black line). Then, the errors in each method as a function of parameter setting and Z_H are provided. To achieve this, the dataset was divided into six 5 dB intervals ranging from 20 to 50 dBZ; we then computed the root-mean-square error (RMSE) and mean error (herein bias) for each interval and normalized by the mean K_{DP}^{sc} from each interval. The optimal parameters were selected based on the smallest averaged normalized RMSE (herein NRMSE) in the last three Z_H intervals (i.e., 35–50 dBZ), prioritizing the accuracy of K_{DP} estimates in high-intensity precipitation.

The settings tested for `kdp_maesaka`, `kdp_vulpiani`, `phase_proc_lp`, and `kdp_from_phidp` are summarized in Table 2, which indicates the tested values, the default value(s) used in the implementation, and the optimal value(s) found in this study.

3.1.1 Py-ART's Maesaka method

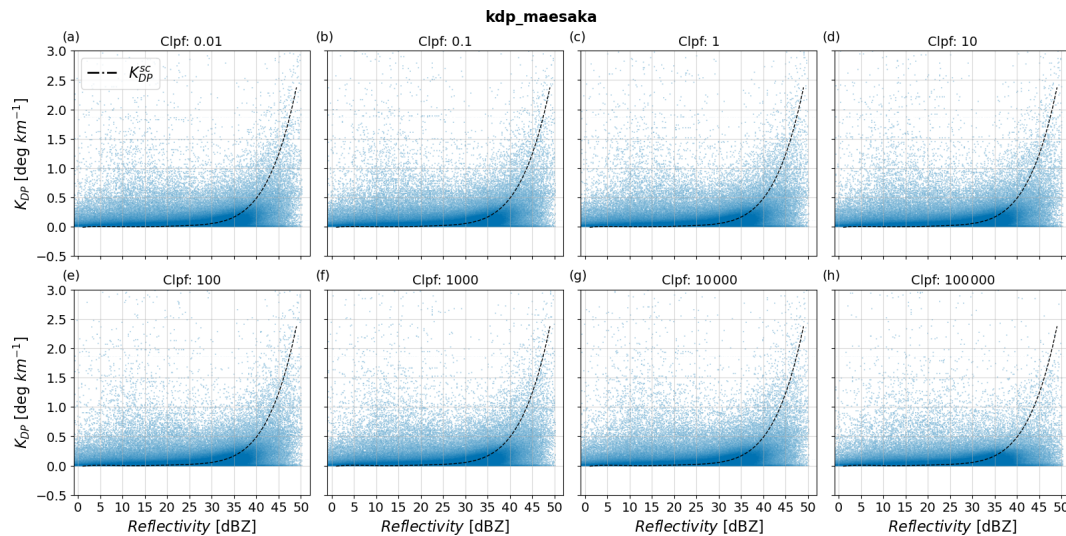
Py-ART's implementation of the Maesaka et al. (2012) method, `kdp_maesaka`, features the optimizing parameter

`Clpf`, which regulates the low-pass filter in Φ_{DP} . The low-pass filter controls the degree of smoothing of Φ_{DP} , with higher `Clpf` values producing smoother Φ_{DP} profiles. In `kdp_maesaka`, the default value of `Clpf` is 1.0, and this value is scaled by the range resolution of the radar to match the resolution of the constraints applied to Φ_{DP} . The scaling is proportional to the fourth power of the range resolution of the radar, and if we were to compare to the values used in Maesaka et al. (2012), a value of 1.0 corresponds to 10^{10} for the Vantaa radar's range resolution of 500 m. In Maesaka et al. (2012), `Clpf` values from 10^9 to 10^{13} were tested on one rainfall case using a 250 m range resolution X-band radar. Their results show that values closer to 10^{13} suppressed fine-scale precipitation features while producing a smooth and clean K_{DP} , whereas values closer to 10^9 preserved fine-scale features while including substantially more noise. These results lead us to test values from 10^8 to 10^{15} , corresponding to 10^{-2} and 10^5 in `kdp_maesaka` and accounting for the Vantaa radar's range resolution. Figure 4a–h show scatterplots of K_{DP} estimates using `kdp_maesaka` as a function of Z_H for different `Clpf` values. All scatterplots show overall accurate and precise K_{DP} estimates within the Z_H range of 0–30 dBZ. This result implies that the subset of `Clpf` values studied produces sufficiently smoothed Φ_{DP} to reduce the impact of noise in light precipitation. However, the effects of excessive smoothing are observed in the range of 40–50 dBZ, where K_{DP} noticeably underestimates K_{DP}^{sc} . By comparing the scatterplots from `Clpf` = 10^{-2} to `Clpf` = 10^5 in the Z_H interval of 40–50 dBZ, the underestimation of K_{DP} is stronger with increasing `Clpf`.

To capture the influence of `Clpf` on the errors when estimating K_{DP} as a function of precipitation intensity, Fig. 5a and b show NRMSE and the normalized bias of K_{DP} estimates with varying `Clpf`. The smaller and more consistent NRMSEs in regions of $Z_H \geq 35$ dBZ in Fig. 5a indicate that `kdp_maesaka` reaches stable solutions for all `Clpf` values tested. However, `Clpf` of 10^5 showed the largest variability when transitioning from lowest to highest Z_H among the values tested, producing the largest NRMSE for $Z_H \geq 35$ dBZ and the lowest otherwise. The underestimation of K_{DP} using 10^5 is evidenced in Fig. 5b for $Z_H \geq 35$ dBZ, where the re-

Table 2. Summary of the parameter settings for each of the optimized methods.

Method	Parameter(s) tested	Tested values	Default	Optimal
kdp_maesaka	Clpf	$\{10^{-2}, 10^{-1}, 10^0, 10^1, 10^2, 10^3, 10^4, 10^5\}$	10^0	10^{-2}
kdp_vulpiani	windsize	{2, 6, 10, 14}	10	10
	n_iter	{2, 6, 10, 14}	10	2
phase_proc_lp	window_len	{5, 10, 15, 20, 25, 30, 35, 40}	35	5
kdp_from_phidp	winlen	{3, 7, 11}	7	11
	dr	{0.5, 1, 2, 4}	1.0	2

**Figure 4.** Scatterplots of estimated K_{DP} from kdp_maesaka as a function of reflectivity and for various values of Clpf. Panels (a)–(h) show results with Clpf values from 10^{-2} to 10^5 . The dashed black line corresponds to K_{DP}^{SC} .

sults were the most negatively biased. The biases from the remaining parameters were equally consistent and smaller.

Our results show that larger values of Clpf lead to larger errors due to oversmoothing of Φ_{DP} . Overall, kdp_maesaka performs consistently when precipitation intensities reach 35 dBZ. The Clpf yielding the smallest 35–50 dBZ averaged NRMSE was 10^{-2} .

3.1.2 Py-ART's Vulpiani method

Py-ART's implementation of the Vulpiani et al. (2012) method, kdp_vulpiani, features two optimizing parameters: windsize (the number of gates used for estimating K_{DP}) and n_iter (the number of re-estimations of K_{DP} per window). Higher values of these parameters result in smoother Φ_{DP} profiles. Reimel and Kumjian (2021) found various parameter combinations that worked well depending on precipitation complexity, leading us to test combinations from 2 to 14 for both parameters. Figure 6a–p show scatterplots comparing the performance of kdp_vulpiani for different values of windsize and n_iter when estimating K_{DP} . Figure 6a

shows the scatter of K_{DP} using the largest settings tested, whereas Fig. 6p shows the results for the smallest. Each row holds windsize constant, while each column holds n_iter constant. In the scatterplot from Fig. 6a with windsize = 14 and n_iter = 14, the data are predominantly clustered under K_{DP}^{SC} for $Z_H \geq 35$ dBZ, indicating underestimation of K_{DP} . For $Z_H < 35$ dBZ, this parameter setting produces accurate and precise results. The scatterplots for smaller setting values, i.e., towards Fig. 6p, are slightly more accurate, albeit significantly less precise; the scatterplot from Fig. 6p with windsize = 2 and n_iter = 2 shows a wider spread of K_{DP} data for all Z_H values, although with slightly enhanced clustering of data around K_{DP}^{SC} for $Z_H \geq 35$ dBZ. These results indicate a trade-off between precision and accuracy when varying windsize and n_iter from 14 to 2. In particular, larger settings favored precision while degrading accuracy, and smaller settings favored accuracy with degraded precision.

To further analyze the trade-off between accuracy and precision when varying windsize and n_iter in kdp_vulpiani, Fig. 7a–b show the NRMSE and normalized bias of K_{DP} estimates with varying windsize and n_iter as a function of Z_H .

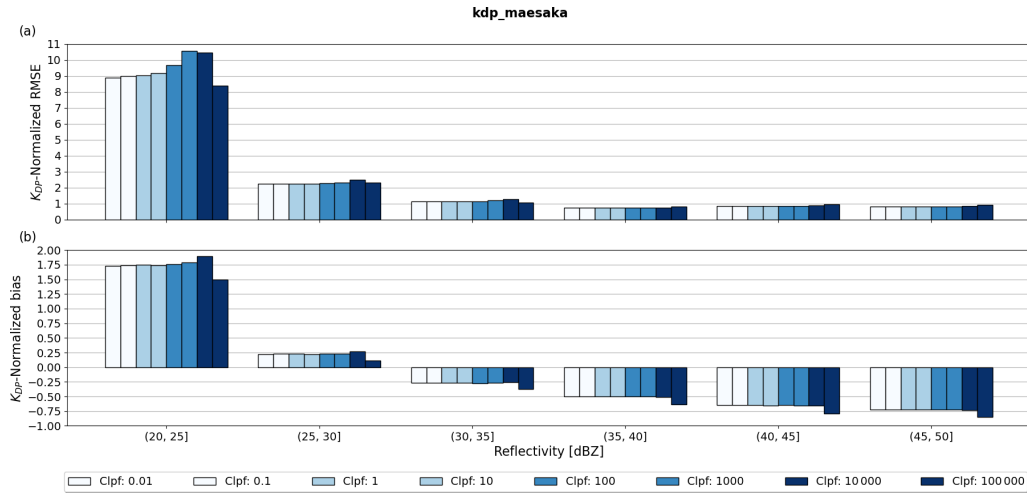


Figure 5. Panel (a) shows RMSE normalized by the interval-averaged K_{DP}^{SC} of $kdp_maesaka$ relative to K_{DP}^{SC} as a function of reflectivity and for various values of Clpf; panel (b) shows the same as panel (a) but for the normalized bias metric.

Figure 7a shows that a window size of 2 yielded the worst performance, implying that the gain in accuracy by including fine-scale fluctuations in Φ_{DP} is not enough to compensate for the increased errors due to the inclusion of outliers. On the other hand, a window size of 14 shows good performance across the entire Z_H range. However, the predominantly negative normalized bias of a window size of 14 relative to the smaller counterparts in Fig. 7b indicates that the larger window size leads to more underestimation of K_{DP} than lower window size values. The consistent errors when varying n_iter in Fig. 7a indicate that this parameter setting does not impact the performance of $kdp_vulpiani$ as strongly as window size does, especially in low Z_H . However, results from Fig. 7b suggest that smaller n_iter significantly reduces the underestimation of K_{DP} estimates when the window size is large. Our results strongly resemble those reported in Reimel and Kumjian (2021), indicating that a smaller number of iterations and moderate window sizes significantly enhance the performance of $kdp_vulpiani$. In particular, among the RMSE heat maps of $kdp_vulpiani$ shown in Reimel and Kumjian (2021), window size = 10 and $n_iter = 2$ produced the best results, coinciding with the smallest 35–50 dBZ averaged NRMSE in this study.

3.1.3 Py-ART’s linear programming method

Py-ART’s implementation of an LP method proposed in Giangrande et al. (2013), `phase_proc_lp`, allows the user to tune the window length to smooth Φ_{DP} , `window_len`, and two intertwined parameters constraining the K_{DP} output via self-consistency relations: `self_const` and `coef`. The former is the weight of the self-consistency constraint and the latter is the exponent in the self-consistency relation linking K_{DP} to Z_H , which is given in Giangrande et al. (2013) as aZ_H^b but is expressed in `phase_proc_lp` as $(10^{0.1 \times Z_H})^{coef} / self_const$. Since information about the expected K_{DP} was known be-

forehand, given by K_{DP}^{SC} , we provided the method with the optimal values of `self_const` = 10^4 and `coef` = 0.914. In this way, the parameter optimization of `phase_proc_lp` was focused solely on `window_len` variations.

The parameter `window_len` defines the window length for smoothing of the LP-processed Φ_{DP} field before K_{DP} is estimated. The default setting of this parameter is 35, indicating a smoothing window length of 17.5 km for a range resolution of 500 m. To include finer-scale precipitation features (e.g., ~ 2.5 km), `phase_proc_lp` was tested with `window_len` values ranging from 5 to 40. Figure 8a–h show scatterplots comparing the performance of `phase_proc_lp` for different settings of `window_len` estimating K_{DP} . Each panel from Fig. 8a to h shows K_{DP} estimated using window lengths from 5 to 40 in intervals of 5. The scatterplot from Fig. 8a with `window_len` = 5 shows data points predominantly clustered around K_{DP}^{SC} across the entire Z_H range, indicating strong correlation between K_{DP} and K_{DP}^{SC} . Even in high Z_H ranges (i.e., ≥ 35 dBZ), the tight correlation between K_{DP} and K_{DP}^{SC} holds, indicating high accuracy and precision of K_{DP} in the presence of heavy precipitation. The accuracy and precision of K_{DP} relative to K_{DP}^{SC} decreases progressively when `window_len` increases, indicated by the spreading and downward shifting of the K_{DP} estimates relative to K_{DP}^{SC} . Especially for the range $Z_H \geq 35$ dBZ, the scatterplots from Fig. 8e to h, with `window_len` from 25 to 40, respectively, show substantial underestimation of K_{DP} relative to K_{DP}^{SC} , indicating stronger oversmoothing of Φ_{DP} for larger values of `window_len`. Comparing the scatterplots, `window_len` = 5 undoubtedly shows the best performance of `phase_proc_lp`. This result agrees with `phase_proc_lp` `window_len` experiments by Li et al. (2023) in an extremely heavy precipitation event, where small `window_len` yielded the best performance. Compared to the `phase_proc_lp` experiments by Reimel and Kumjian (2021), our results suggest that smaller

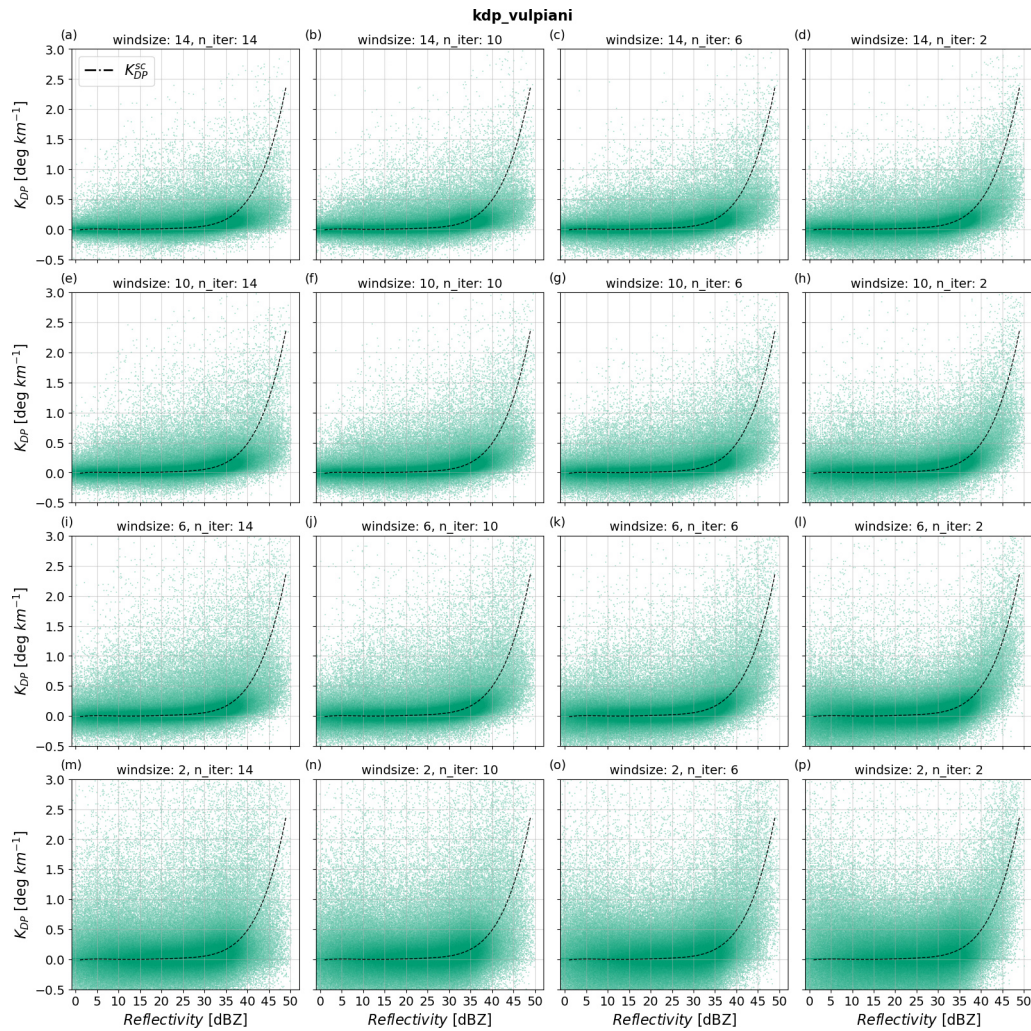


Figure 6. Scatterplots of estimated K_{DP} from `kdp_vulpiani` as a function of reflectivity and for various values of `window_size` and `n_iter`. Panels (a)–(p) show results with (`window_size`, `n_iter`) tuple values from (14, 14) to (2, 2), decreasing `window_size` with increasing rows. The dashed black line corresponds to K_{DP}^{SC} .

`window_len` produce overall more accurate K_{DP} estimates. However, the influence of the self-consistency constraints proposed in Giangrande et al. (2013) plays a key role in this aspect; if optimal self-consistency constraints are not provided or do not match theoretical expectations, the precision and accuracy in K_{DP} significantly decreases, and larger `window_len` values compensate for this by oversmoothing Φ_{DP} (see Appendix A for results of the performance of `phase_proc_lp` with very little influence of self-consistency constraints).

To further investigate the effects of `window_len` on the performance of `phase_proc_lp`, Fig. 9a–b show the NRMSE and normalized bias of K_{DP} estimates with varying `window_len` as a function of Z_H . In agreement with the patterns observed in the scatterplots in Fig. 8, `window_len` = 5 produced the best performance compared to other parameter settings. Interestingly, even in light precipitation (e.g.,

$Z_H < 30$ dBZ), smaller values of `window_len` produced the best NRMSE metrics, indicating that larger `window_len` do not further improve the precision of `phase_proc_lp`. Instead, larger `window_len` enhanced the bias of K_{DP} relative to K_{DP}^{SC} , as shown in Fig. 9b. The parameter `window_len` = 5 produced undoubtedly the best metrics for `phase_proc_lp`, and it was selected as the optimal parameter.

3.1.4 *oradlib*'s Vulpiani method

oradlib's implementation of the Vulpiani et al. (2012) method, `kdp_from_phidp`, features two optimizing parameters: `winlen` (the number of gates used to reconstruct Φ_{DP}) and `dr` (the gate length resolution in km). We tested `winlen` values from 3 to 11 and `dr` values from 0.5 to 4. Figure 10a–l show scatterplots of K_{DP} estimates using `kdp_from_phidp` when varying the settings of `winlen` and `dr`. Each row of

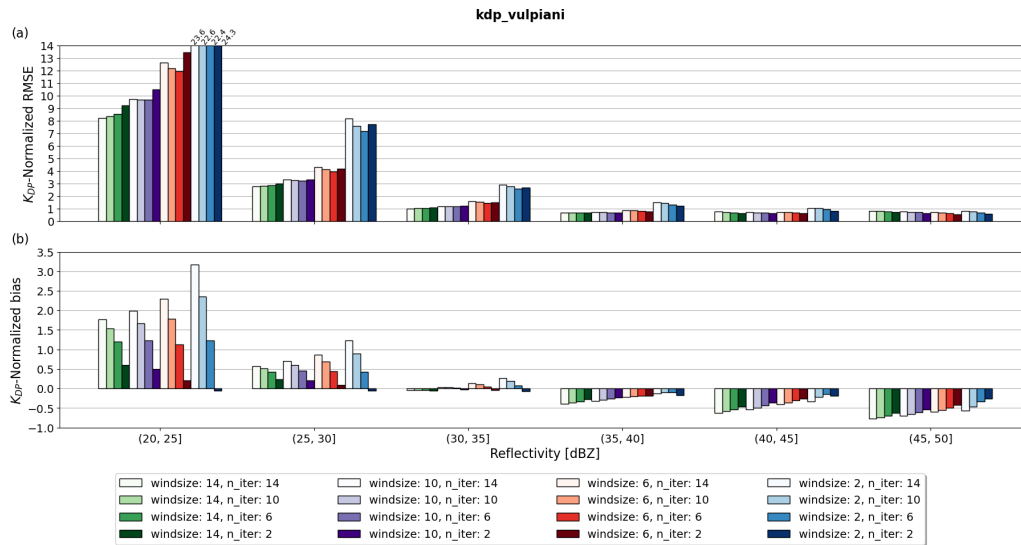


Figure 7. Panel (a) shows RMSE normalized by the interval-averaged K_{DP}^{SC} of `kdp_vulpiani` relative to K_{DP}^{SC} as a function of reflectivity and for various values of `windsize` and `n_iter`; panel (b) shows the same as panel (a) but for the normalized bias metric.

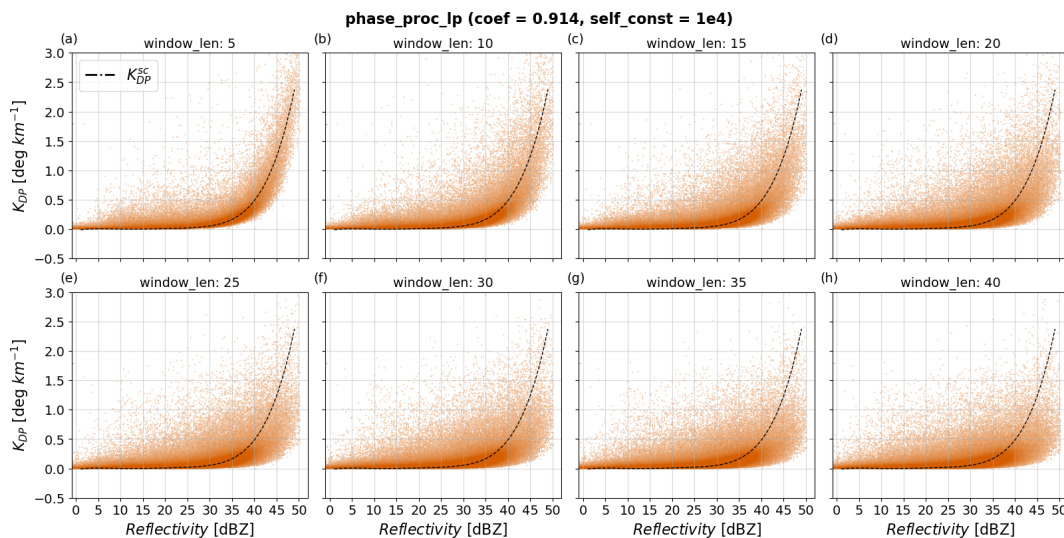


Figure 8. Scatterplots of estimated K_{DP} from `phase_proc_lp` as a function of reflectivity and for various values of `window_len`. Panels (a)–(h) show results with `window_len` values from 5 to 40 when fixing `coef` at 0.914 and `self_const` at 10^4 . The dashed black line corresponds to K_{DP}^{SC} .

scatterplots holds `winlen` constant, while decreasing `dr` from left to right. Similarly, each column of scatterplots holds `dr` constant, while decreasing `winlen` from top to bottom. The scatterplot in Fig. 10a, with `winlen` = 11 and `dr` = 4, shows K_{DP} clustered predominantly around 0° km^{-1} across the entire Z_H range, indicating substantial oversmoothing of Φ_{DP} . Even for $Z_H \geq 30$ dBZ, the noticeable underestimation of K_{DP} relative to K_{DP}^{SC} indicates that `kdp_from_phidp` is not able to capture signatures of heavy precipitation for large `winlen` and `dr` settings. Moving towards the scatterplot in Fig. 10d, a smaller `dr` enhances the accuracy of

`kdp_from_phidp`, particularly for $Z_H \geq 30$ dBZ. However, the gain in accuracy comes together with a loss in precision in K_{DP} estimates, indicated by the wider spread of the data. In addition, decreasing `dr` makes `kdp_from_phidp` more prone to the inclusion of outliers, illustrated by data points with $K_{DP} > 1^\circ \text{ km}^{-1}$, even for $Z_H \leq 20$ dBZ. The scatterplots in Fig. 10e–h follow the same behavior as in the first row except for a wider spread of data, suggesting that decreasing `winlen` while holding `dr` constant overall reduces the precision of `kdp_from_phidp`. When moving from Fig. 10e to h, the accuracy of K_{DP} estimates increases while precision de-

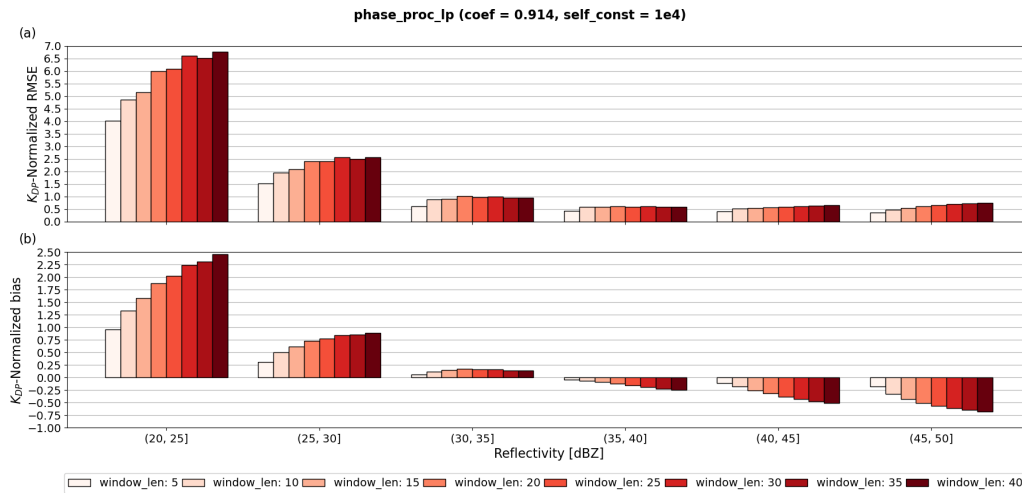


Figure 9. Panel (a) shows RMSE normalized by the interval-averaged K_{DP}^{SC} of *phase_proc_lp* relative to K_{DP}^{SC} as a function of reflectivity and for various values of *window_len*; panel (b) shows the same as panel (a) but for the normalized bias metric.

creases with decreasing *dr*. In the last row, i.e., from Fig. 10i to l, K_{DP} estimates are the most scattered for the same *dr*, indicating a loss in precision of *kdp_from_phidp* when reducing *winlen*. The scatterplot in Fig. 10l with the smallest parameter settings tested (*winlen* = 3 and *dr* = 0.5) resembles a scatterplot of random noise with no significant clustering of data, suggesting extremely poor correlation relative to K_{DP}^{SC} . Comparing the scatterplots row-wise and column-wise, decreasing *winlen* or *dr* significantly degrades the precision of the method. However, the effect on the accuracy is more complex; simultaneously setting *winlen* and *dr* to large values leads to substantial underestimation of K_{DP} , whereas small values lead to noisy K_{DP} . These results suggest that the effects of varying *winlen* and *dr* on the performance of *kdp_from_phidp* are strongly intertwined, requiring more analysis of the trade-off between accuracy and precision offered by variations in these parameters.

To analyze the trade-off between accuracy and precision when using *winlen* and *dr* in *kdp_from_phidp*, Fig. 11a–b show the NRMSE and normalized bias of K_{DP} estimates with varying *winlen* and *dr* as a function of Z_H . Even though Fig. 11a has been clipped at 5.0, it is important to note the significantly high values when using the smallest *dr* (97.6, 135.4, and 278.6 for *winlen* of 11, 7, and 3, respectively). The predominantly higher NRMSE values with the smallest *dr* indicate that the precision of *kdp_from_phidp* reduces significantly with *dr* < 1 for any *winlen* tested. An exception occurs in the Z_H interval (45, 50] dBZ, where the smallest *dr* yield the best metrics due to slight improvements in the accuracy. Despite the limited amount of data within this Z_H interval (see Fig. 3), the clustering of K_{DP} around K_{DP}^{SC} in Fig. 10 and the small normalized biases in Fig. 11b suggest that accuracy improved slightly for the smallest *dr*. The smaller NRMSE with high *dr* in Fig. 11a is counterbalanced by the predominantly larger negative bias for larger *dr* in Fig. 11b. This im-

plies that larger *dr* values in *kdp_from_phidp* lead to the underestimation of K_{DP} for all *winlen* tested. As a conclusion, combining large *winlen* with smaller *dr* produces the best performance for heavier precipitation (i.e., $Z_H > 30$ dBZ), whereas combining large *winlen* with larger *dr* produces the best results for light precipitation. Overall, small values of *winlen* reduce the precision significantly in the method without improving accuracy. The parameter setting with the smallest 35–50 dBZ averaged NRMSE was *winlen* = 11 and *dr* = 2.

3.2 Performance assessment of methods relative to K_{DP}^{SC}

The performance of the methods is analyzed qualitatively in Sect. 3.2.1 and quantitatively in Sect. 3.2.2. For these analyses, we used the parameter-optimized *kdp_maesaka*, *kdp_vulpiani*, *phase_proc_lp*, and *kdp_from_phidp* and included *kdp_schneebeli* and *kdp_iris*.

3.2.1 Qualitative assessment

We qualitatively assessed the precision and accuracy of the estimated K_{DP} using scatterplots of K_{DP} vs. Z_H for each method. Figure 12 shows six scatterplots comparing the performance of *kdp_maesaka*, *kdp_vulpiani*, *phase_proc_lp*, *kdp_from_phidp*, *kdp_schneebeli*, and *kdp_iris* at estimating K_{DP} . Each scatterplot illustrates the relationship between estimated K_{DP} (*y* axis) relative to Z_H (*x* axis) against benchmarking K_{DP}^{SC} (dashed black line). For the parameter-optimized methods in Fig. 12a–d, the optimal parameter selected is indicated in the panel title together with the method’s name. Comparing the scatterplots, *phase_proc_lp* demonstrates the highest accuracy and precision, evidenced by the data narrowly clustered around K_{DP}^{SC} across the entire Z_H range. The *kdp_from_phidp*

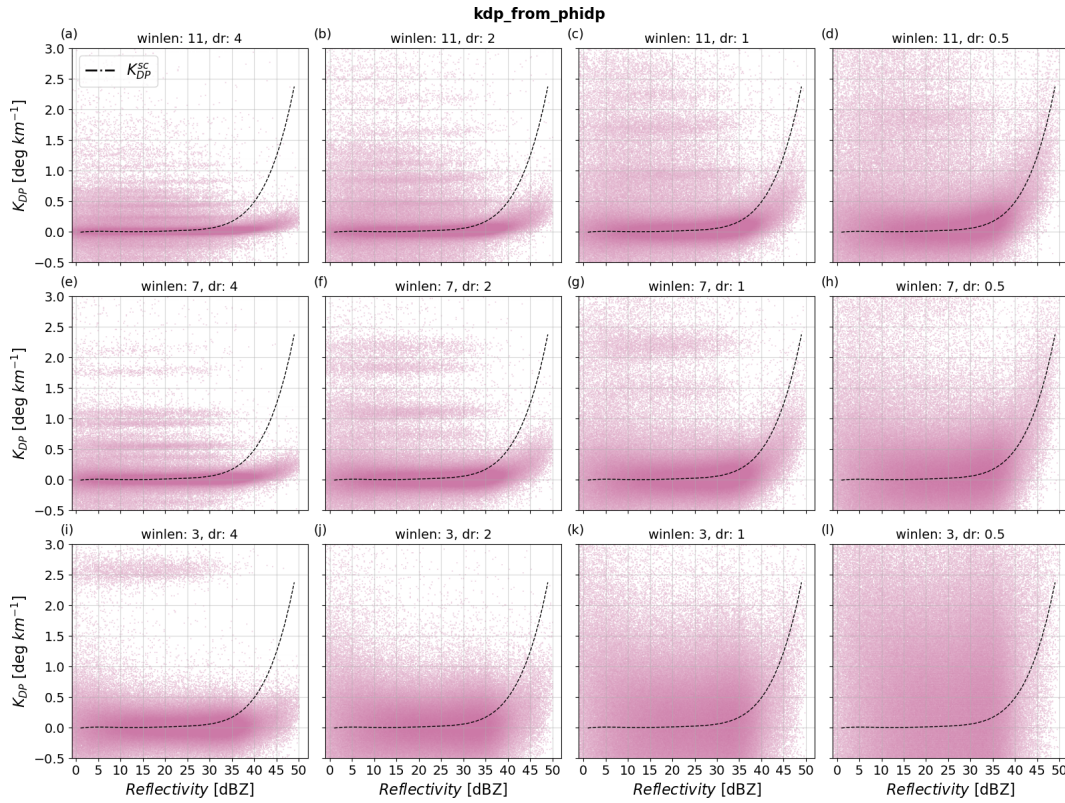


Figure 10. Scatterplots of estimated K_{DP} from `kdp_from_phidp` as a function of reflectivity and for various values of `winlen` and `dr`. Panels (a)–(l) show results with (`winlen`, `dr`) tuple values from (11, 4) to (3, 0.5); `winlen` decreases in intervals of 4 per row, whereas `dr` decreases by half per column. The dashed black line corresponds to K_{DP}^{SC} .

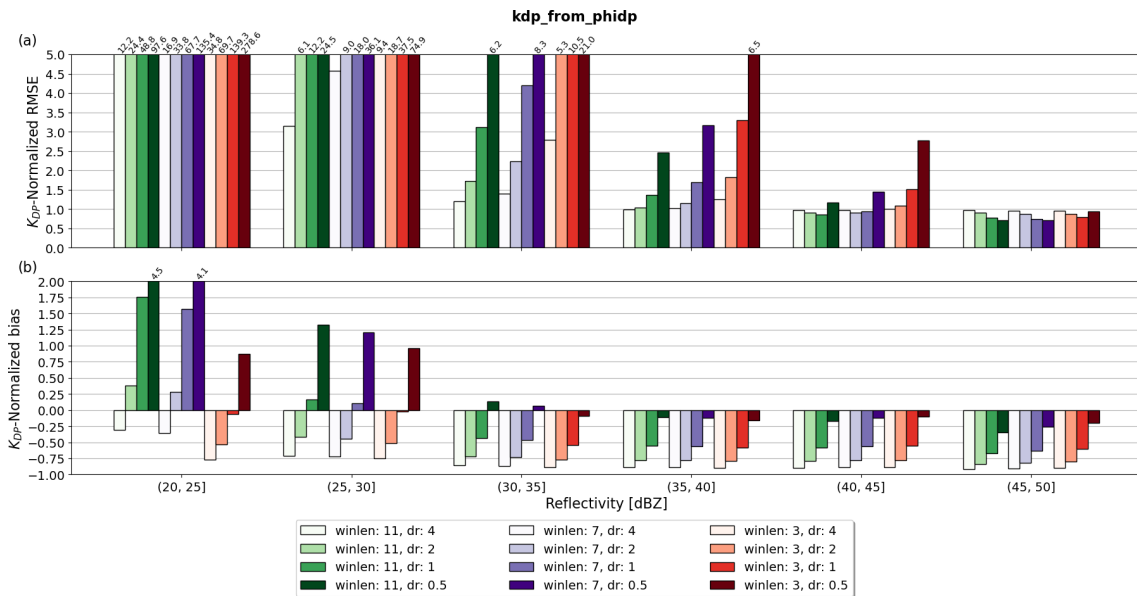


Figure 11. Panel (a) shows RMSE normalized by the interval-averaged K_{DP}^{SC} of `kdp_from_phidp` relative to K_{DP}^{SC} as a function of reflectivity and for various values of `winlen` and `dr`; panel (b) shows the same as panel (a) but for the normalized bias metric. The numbers on top of the bars indicate the values of the metric exceeding the y-axis limit selected.

and `kdp_schneebeli` methods show the least accuracy and precision, with a broader spread and more outliers, particularly when $Z_H < 30$ dBZ. For higher Z_H values, even though `kdp_from_phidp` shows better precision but worse accuracy than `kdp_schneebeli`, these two methods strongly underestimate K_{DP} , evidenced by the predominant clustering of K_{DP} estimates below 0.5°km^{-1} . The `kdp_maesaka` method shows less scattering of K_{DP} estimates compared to `kdp_from_phidp` and `kdp_schneebeli`, indicating higher precision and accuracy, particularly for $Z_H < 30$ dBZ. However, for $Z_H \geq 30$ dBZ, the performance of `kdp_maesaka` deteriorates rapidly, as shown by the broader spread and significant underestimation of K_{DP} relative to K_{DP}^{sc} . The `kdp_vulpiani` and `kdp_iris` methods show moderate performance, with better accuracy and precision than `kdp_from_phidp`, `kdp_schneebeli`, and `kdp_maesaka` but worse performance than `phase_proc_lp`. Between the `kdp_vulpiani` and `kdp_iris` methods, `kdp_vulpiani` shows better correlation of K_{DP} estimates with K_{DP}^{sc} for $Z_H \geq 35$ dBZ, indicating higher accuracy in heavier precipitation. However, `kdp_iris` shows less scattering across the entire Z_H range, indicating overall higher precision than `kdp_vulpiani`. The `kdp_vulpiani`, `kdp_from_phidp`, `kdp_schneebeli`, and `kdp_iris` methods include negative K_{DP} values, which should not be expected in rain observations. These negative estimates predominantly show up in lighter precipitation (i.e., $Z_H < 30$ dBZ), indicating that they are most likely produced by noise in Φ_{DP} . However, the inclusion of negative K_{DP} estimates is useful, for instance, in the detection of snow crystals, allowing `kdp_vulpiani`, `kdp_from_phidp`, `kdp_schneebeli`, and `kdp_iris` to be used in a wider range of applications compared to `kdp_maesaka` and `phase_proc_lp`. The relatively high accuracy and precision of `kdp_iris` and `kdp_vulpiani`, together with the inclusion of negative K_{DP} estimates, leave these two methods as candidates well-suited for QPE and calibration and hydrometeor classification routines.

3.2.2 Quantitative assessment

The quantitative assessment of the methods was achieved through the metrics of NRMSE and normalized bias and complemented with statistics from the Wasserstein distance (WD) (Ramdas et al., 2015). The WD measures the similarity between two cumulative distributions, given in this study by the K_{DP} estimated by each method and K_{DP}^{sc} . On the one hand, NRMSE and normalized bias computed as a function of Z_H , allow the assessment of the relative accuracy and precision of the methods based on precipitation intensities. The WD, on the other hand, estimated from each radar scan independently and with statistics over the entire set of scans, allows the assessment of the relative consistency and robustness of the methods.

Figure 13a–b show the NRMSE and normalized bias of estimated K_{DP} for each method. Overall, `phase_proc_lp` shows

the best performance, as evidenced by the lowest NRMSE values in Fig. 13a and moderately low bias in Fig. 13b across all Z_H intervals. In contrast, `kdp_schneebeli` shows the worst performance among the methods, indicated by the highest NRMSE values and moderately high bias across all Z_H intervals. The `kdp_from_phidp` method shows substantially higher NRMSE values than `kdp_maesaka`, `kdp_vulpiani`, `phase_proc_lp`, and `kdp_iris` but is significantly smaller than `kdp_schneebeli`, particularly for the smallest Z_H values. The relatively small bias of `kdp_from_phidp` when NRMSE values are substantially high is explained by the positive-to-negative symmetrical spread of K_{DP} estimates around the x axis, indicating poor precision. Additionally, the persistently negative and large normalized bias of this method relative to the other methods indicates that `kdp_from_phidp` underestimates K_{DP} the most. The `kdp_maesaka`, `kdp_vulpiani`, and `kdp_iris` methods have moderate NRMSE values, performing better than `kdp_schneebeli` and `kdp_from_phidp` but not as well as `phase_proc_lp`. Among these three methods, `kdp_maesaka` has the smallest NRMSE values for $Z_H \leq 35$ dBZ but the largest when $Z_H \geq 40$ dBZ. The relatively large positive bias of `kdp_maesaka` when $Z_H < 30$ is a direct consequence of the exclusion of negative K_{DP} estimates. However, the persistently larger negative bias of `kdp_maesaka` relative to `kdp_vulpiani` and `kdp_iris` when $Z_H \geq 30$ dBZ indicates stronger underestimation of K_{DP} and thus less accuracy. These results indicate that in comparison to other methods, `kdp_maesaka` performs slightly better in light precipitation (i.e., $Z_H < 30$ dBZ) but worse in heavier precipitation. Between `kdp_vulpiani` and `kdp_iris`, `kdp_iris` shows overall smaller NRMSEs and normalized bias, indicating higher accuracy and precision than `kdp_vulpiani`.

Complementary to the NRMSE and normalized bias metrics, we evaluated the consistency and robustness of the methods using the Wasserstein distance (WD). The WD was computed for each radar scan independently using the `wasserstein_distance` module from SciPy (Virtanen et al., 2020). Then, the statistics from the estimated WD values for all scans were visualized and analyzed using boxplots. Figure 14 consists of two panels comparing the WD boxplots of the methods. Figure 14a compares the WD for all methods, including `kdp_schneebeli`, which presented a significantly large WD. Figure 14b presents the same data as (a) but excludes `kdp_schneebeli` to better compare the remaining methods. Each boxplot summarizes the statistics of estimated WDs by showing the median (dashed black line), interquartile ranges (IQR), $1.5 \times \text{IQR}$ (whiskers), and outliers (crosses). The insights provided by the boxplots in this analysis are twofold. First, a WD median closer to 0 indicates higher similarity between the cumulative distributions of a method's estimated K_{DP} and that from K_{DP}^{sc} , ultimately indicating higher accuracy. Second, a narrower IQR indicates less variability in a method's performance between scans, indicating higher consistency.

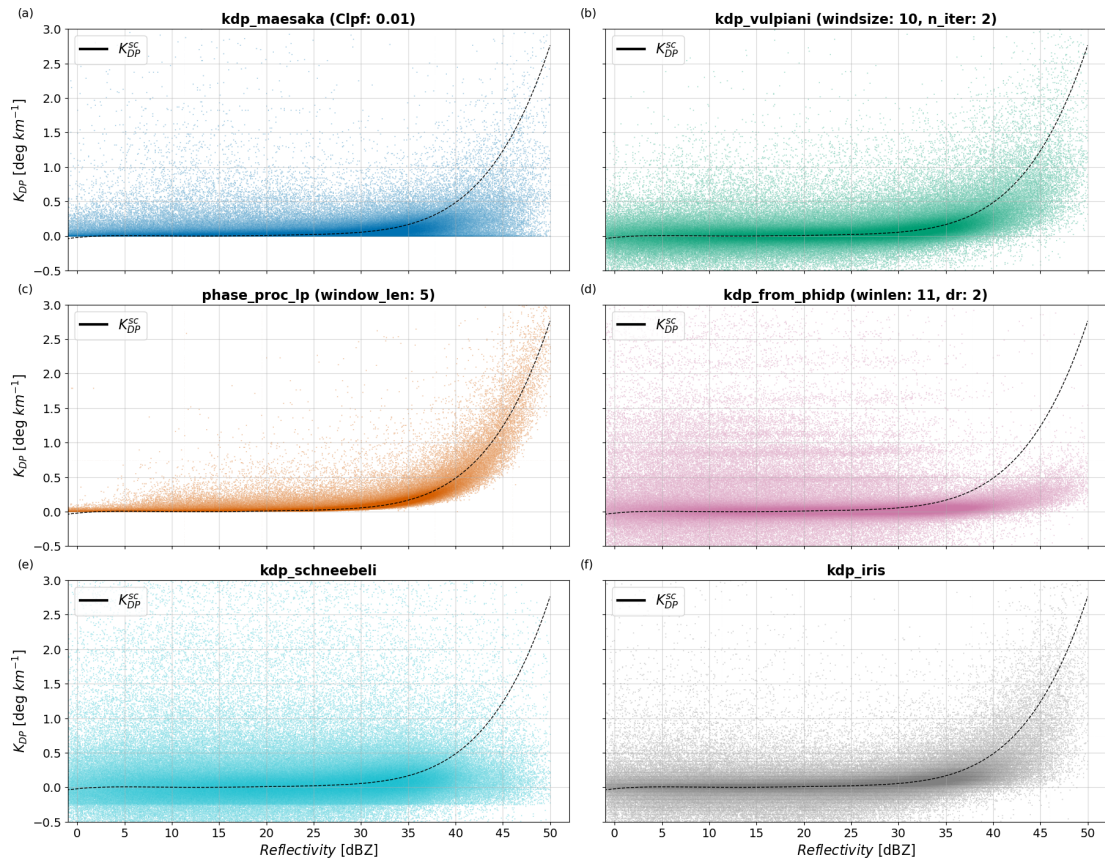


Figure 12. Scatterplot of estimated K_{DP} from each parameter-optimized method relative to K_{DP}^{SC} as a function of reflectivity. Panels (a)–(f) show `kdp_maesaka`, `kdp_vulpiani`, `phase_proc_lp`, `kdp_from_phi_dp`, `kdp_schneebeli`, and `kdp_iris`, respectively. The dashed black line corresponds to K_{DP}^{SC} .

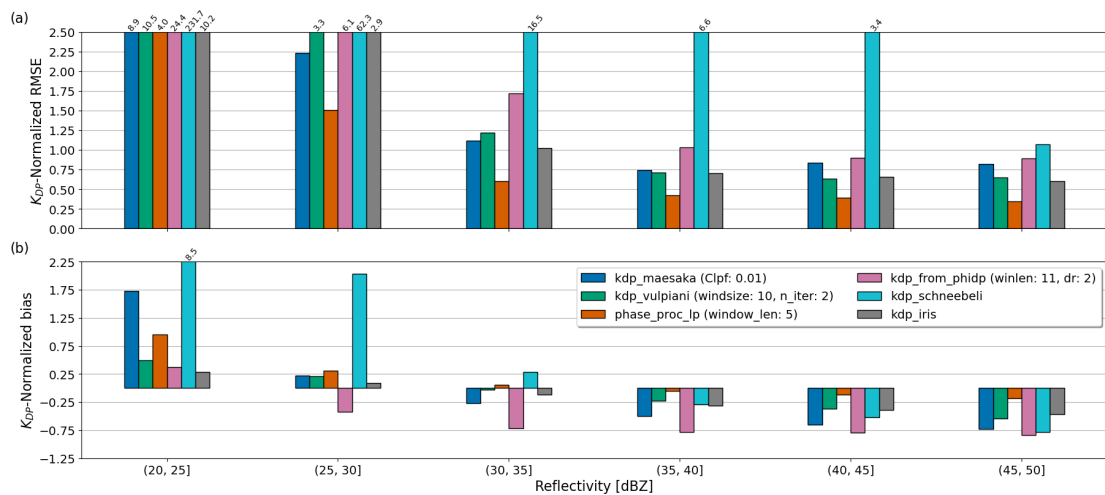


Figure 13. Panel (a) shows the bias of estimated K_{DP} from each parameter-optimized method relative to K_{DP}^{SC} as a function of reflectivity; panel (b) shows the same as panel (a), but the bias is normalized by interval-averaged K_{DP}^{SC} . The numbers on top of the bars indicate the values of the metric exceeding the y-axis limit selected.

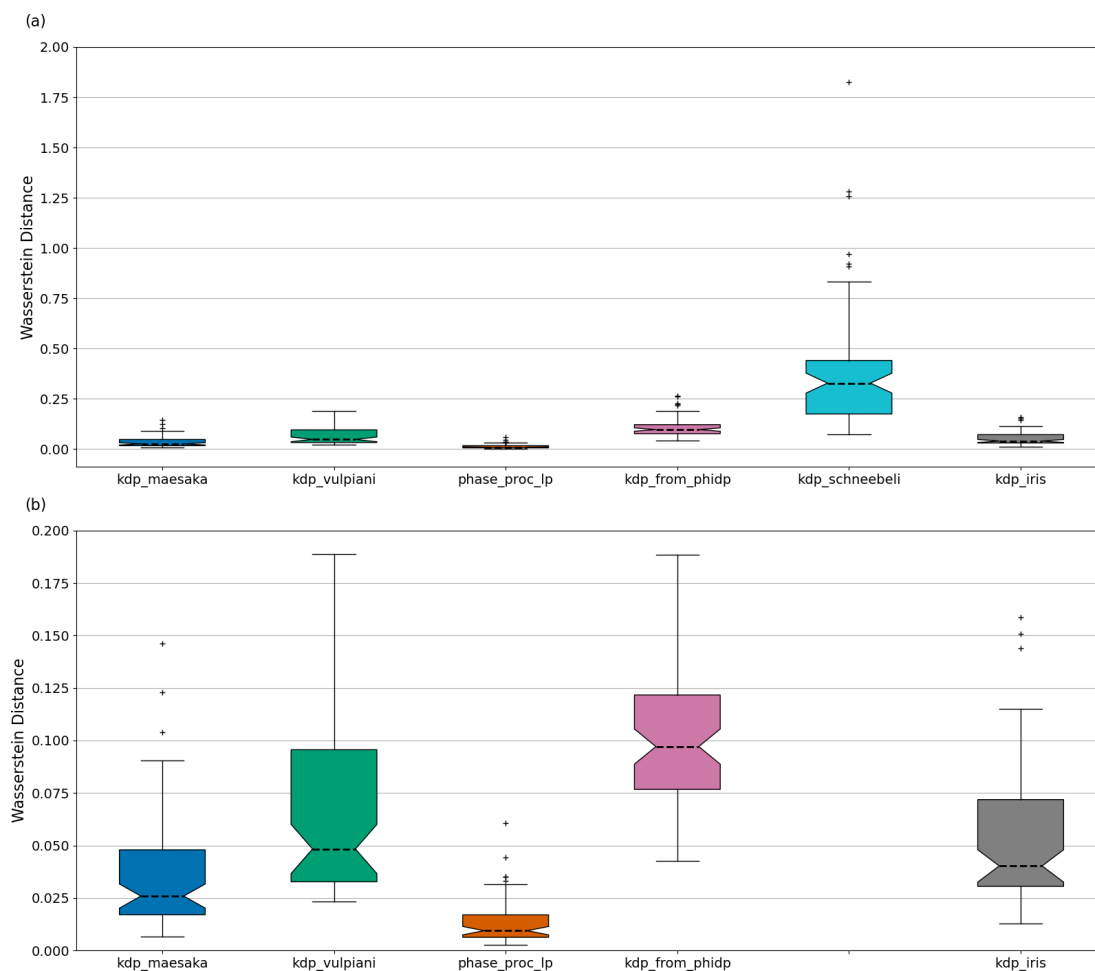


Figure 14. Panel (a) shows the boxplot of the WD computed for each parameter-optimized method; panel (b) shows the same as panel (a) but excluding kdp_schneebeli for better visualization of the better-performing methods. The boxplots display the WD median (dashed black line), IQRs (boundaries of the box), $1.5 \times$ IQR (whiskers), and the outliers (black crosses).

In Fig. 14a, the x axis lists six methods: kdp_maesaka, kdp_vulpiani, phase_proc_lp, kdp_from_phidp, kdp_schneebeli, and kdp_iris. The y axis measures the WD values ranging from 0 to 2. The boxplot for the kdp_schneebeli method shows the largest WD, with a median of 0.33, an IQR from 0.18 to 0.45, and several outliers. The other methods (kdp_maesaka, kdp_vulpiani, phase_proc_lp, kdp_from_phidp, and kdp_iris) have median WD values ranging from 0.0 to 0.1, with smaller IQRs and fewer outliers. In Fig. 14b, the kdp_schneebeli method is excluded, allowing for a clearer comparison of the kdp_maesaka, kdp_vulpiani, phase_proc_lp, kdp_from_phidp, and kdp_iris methods. The y axis is rescaled to range from 0.0 to 0.2 for better visualization. The phase_proc_lp method has the lowest WD median at 0.01, with a narrow IQR from 0.008 to 0.018. The kdp_from_phidp method has a significantly larger WD median of 0.098 and an IQR from 0.077 to 0.122. The kdp_maesaka and kdp_iris methods have WD medians of 0.026 and 0.041, respectively, with moderate IQRs and

few outliers. The kdp_vulpiani method has a moderate WD median of 0.049 but a noticeably wider IQR from 0.033 to 0.096 when compared to kdp_maesaka, phase_proc_lp, kdp_from_phidp, and kdp_iris.

The large WD median of kdp_schneebeli indicates that it performs worse compared to the other methods, overshadowing the performance differences among the remaining methods. Additionally, the large IQR of kdp_schneebeli implies that the method does not perform consistently, thus reducing its reliability. The phase_proc_lp method demonstrates the best and most consistent performance, with the lowest WD median and narrowest IQR. These results additionally indicate that the distribution of K_{DP} estimated from phase_proc_lp is the closest to K_{DP}^{SC} . It is important to remember here that phase_proc_lp is supported by self-consistency relations constraining the K_{DP} estimates based on Z_H observations, ultimately enhancing its accuracy and stability. The moderate IQR and significantly larger WD median of kdp_from_phidp indicate that its performance is

consistent, albeit less accurate relative to the other methods. The `kdp_vulpiani` method, in turn, has a moderate WD median but relatively larger IQR, indicating better accuracy than `kdp_from_phidp`, although less consistency. The `kdp_maesaka` and `kdp_iris` methods show similar consistency and accuracy, evidenced by their relatively low WD medians and moderate IQRs. These findings suggest that while `kdp_schneebeli` is the least accurate and consistent, the performance among the remaining methods varies, with `phase_proc_lp` presenting the highest robustness, provided that the method with quality controlled Z_H and optimized self-consistency settings is used.

3.3 Consistency analysis of K_{DP} retrievals

Each method has a unique combination of mathematical approaches, data requirements, and constraints (see Table 1), indicating uniqueness in the K_{DP} fields produced. The similarity or dissimilarity of these outputs is not clearly visible from the metrics computed or from the scatterplots displayed in Sect. 3.2. To answer this question, we study the consistency among methods using the K_{DP} vs. K_{DP} correlation plots shown in Fig. 15. Each scatterplot in Fig. 15 shows the relationship between K_{DP} estimated by a method (y axis) with respect to K_{DP} estimated by a different method (x axis), and the Pearson correlation coefficient (R) is shown in the upper-left corner of each scatterplot. The axes range from -0.5 to 3.0 km^{-1} to include negative K_{DP} estimates. This part of the analysis does not require any ground-truth framework, allowing the use of the entirety of the radar dataset, i.e., including the attenuated observations (see red data in Fig. 3).

In Fig. 15, the scatterplot of `kdp_iris` against `kdp_vulpiani` shows the best correlation among the methods, illustrated by the data significantly clustered along the diagonal and corroborated by the highest R of 0.66. The `kdp_iris` and `kdp_vulpiani` methods correlate similarly with `phase_proc_lp`, indicated by the second-highest R of 0.65 for both. In relation to `kdp_maesaka`, the consistencies of `kdp_iris` and `kdp_vulpiani` are rather moderate, whereas in relation to `kdp_from_phidp` and `kdp_schneebeli`, they are significantly poorer. Among the methods, `kdp_schneebeli` correlates the least with any of the other methods, evidenced by the data widely spread along the axes and showing negligible clustering along the diagonal. In particular, `kdp_schneebeli` against `kdp_from_phidp` shows the worst consistency, with $R = 0$ and the majority of the data clustered around the x and y axes. The `phase_proc_lp` method correlates moderately with `kdp_maesaka`, with an $R = 0.41$, although the scatterplot does not exhibit any particular pattern or clustering of data along the diagonal. Relative to `kdp_from_phidp`, `phase_proc_lp` shows significantly lower R despite the clear data correlation off of the diagonal. However, the small R value becomes evident when observing the dense clustering of data around 0 km^{-1} for `phase_proc_lp`. This result indicates that the consistency be-

tween `kdp_from_phidp` and `phase_proc_lp` is highly influenced by the negative K_{DP} estimates in `kdp_from_phidp` that are mapped to 0 km^{-1} in `phase_proc_lp`. Overall, the scatterplots show that `kdp_from_phidp` underestimates K_{DP} relative to the other methods. The `kdp_maesaka` method shows no significant correlation with any method, with the largest R being 0.41 relative to both `phase_proc_lp` and `kdp_vulpiani`.

4 Conclusions

In this study, we conducted a comprehensive evaluation of several K_{DP} estimation methods using C-band weather radar data, with a focus on their performance in rainfall observations. We employed a self-consistency framework that links Z_H and Z_{dr} observations, with K_{DP} as the basis for our evaluations. This approach allows for the construction of the reference K_{DP} observations that can be used to assess the accuracy and robustness of the K_{DP} estimation methods studied. The use of the self-consistency framework requires rather strict quality control, which is described in the paper. In this way, our study focuses on the performance of the methods in highly idealized rainfall observations.

Some (four out of six) of the K_{DP} estimation methods have user-configurable parameters. Using the evaluation framework proposed, we were able to define optimized parameter settings. Most of the methods showed significant improvement in the performance after the optimization.

By comparing the relative performance of the estimation methods over the range of rain intensities, as characterized by the radar Z_H values, we have found significant differences in the performance of the methods evaluated. Overall, implementations of the Giangrande et al. (2013), Vulpiani et al. (2012), and Wang and Chandrasekar (2009) methods exhibited the lowest NRMSE and normalized biases over the range of Z_H values studied, from 20 to 50 dBZ.

Our comparative analysis revealed that while the implementation of the Giangrande et al. (2013) method stands out for its high accuracy and precision, its performance is heavily dependent on the self-consistency constraint provided. Without proper optimization of the self-consistency relation, linking of Z_H and K_{DP} , and quality control of Z_H , even the best window length setting for this method can lead to suboptimal results, i.e., higher RMSE and K_{DP} underestimation at higher Z_H values. It should be noted, however, that the reference framework and the Giangrande et al. (2013) method use self-consistency relations to determine K_{DP} , and, therefore, the uncertainties are correlated, and part of the reported performance is caused by this dependence. The implementations of Vulpiani et al. (2012) and Wang and Chandrasekar (2009) showed good performance and do not require the use of other radar variables, which potentially make them less sensitive to radar data quality issues, such as calibration and attenuation.

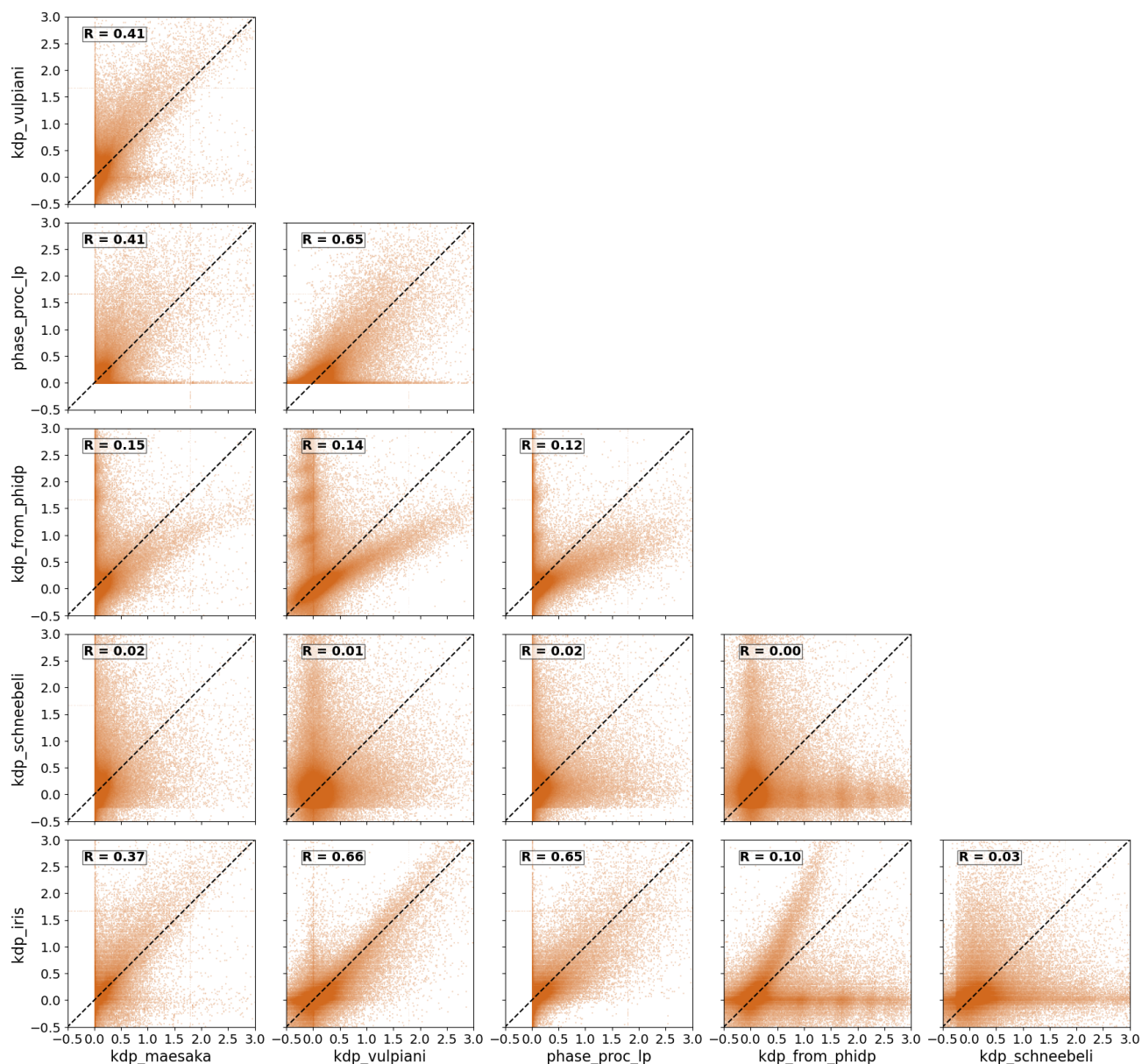


Figure 15. Correlation between the K_{DP} estimation methods. Each scatterplot shows the relationship between two different methods without repetition, and no method is compared to itself. The x and y axes represent the K_{DP} estimated by a method, in units of $^{\circ}\text{km}^{-1}$. Each panel shows the Pearson correlation coefficient between the two methods compared.

An additional qualitative comparison of the performance of the methods was carried out by computing correlations of derived K_{DP} values from the dataset that also included attenuated radar observations. The correlation between K_{DP} values estimated using different methods is not very high. The highest correlation values (0.65–0.66) were observed between the Giangrande et al. (2013), Vulpiani et al. (2012), and Wang and Chandrasekar (2009) methods. This indicates that uncertainty between different precipitation estimates could stem from the differences in the K_{DP} methods used.

The study is based on a self-consistency framework that limits its use to the cases where no significant attenuation is observed. Additionally, the scope of our study is limited to the Finnish climatology and a single radar frequency, namely C-band radar observations. Despite these limitations, our findings offer valuable guidance for the use of K_{DP} estimation methods in rainfall observations. These results have significant implications for both operational radar networks and hydrometeorological research, where the accuracy, precision, and stability of K_{DP} estimates are crucial.

Appendix A: Influence of the self-consistency constraint on phase_proc_lp

Figure A1 shows the same scatterplots as Fig. 8, with K_{DP} estimated from phase_proc_lp using self_const of 10^6 instead of 10^4 . The motivation behind this was to study the performance of phase_proc_lp with little influence from self-consistency constraints. In Giangrande et al. (2013), the non-negativity condition in K_{DP} estimates is ensured by restricting the \mathbf{b} vectors: $\mathbf{b} \geq 0$. In addition, to produce more realistic K_{DP} estimates, they introduced the self-consistency relation $K_{DP}(Z_H) = aZ_H^b$ to bound the estimates based on observed Z_H , requiring that the user provide quality controlled data. The restriction of the \mathbf{b} vectors becomes $\mathbf{b} \geq aZ_H^b$, which in phase_proc_lp is implemented as $\mathbf{b} \geq (10^{0.1 \times Z_H})^{coef} / \text{self_const}$. Therefore, a self_const value that is 2 orders of magnitude larger was used in this study to test the performance of phase_proc_lp with a significantly reduced influence of self-consistency constraints. The scatterplots show K_{DP} data clustered around K_{DP}^{sc} up to 35 dBZ. Beyond this threshold, precision and accuracy decay significantly regardless of the window length. However, in scatterplots with larger window lengths, K_{DP} data are less scattered across the entire Z_H range and are only slightly less accurate after 35 dBZ.

To further investigate the effects of the self-consistency constraint on phase_proc_lp, Fig. A2a–b show the normalized RMSE and bias of K_{DP} (estimated with self_const = 10^6) relative to K_{DP}^{sc} . Interestingly, the normalized RMSE in Fig. A2a behaves inversely to the normalized RMSE in Fig. 9, whereas normalized bias shows similar behavior for both. The opposite behaviors in normalized RMSE results indicate that window length has a strong impact on the performance of phase_proc_lp, depending on whether adequate self-consistency settings were provided; if so, smaller window lengths yield better performance by capturing fine-scale precipitation features, especially in heavy precipitation. In the opposite case, larger window lengths yield better performance by oversmoothing Φ_{DP} , thus reducing the impact of noise at the expense of losing fine-scale precipitation features. The oversmoothing effect from larger window lengths in K_{DP} is also implied from the normalized bias shown in Fig. A2b; larger window lengths produced the largest absolute biases at both extremes of the Z_H range. In addition, even though the normalized bias shows similar behavior for self_const = 10^6 and self_const = 10^4 , the latter produces larger differences between window lengths, indicating that the high accuracy and precision of phase_proc_lp predominates in smaller window lengths, provided that there are adequate self-consistency constraints and quality controlled Z_H .

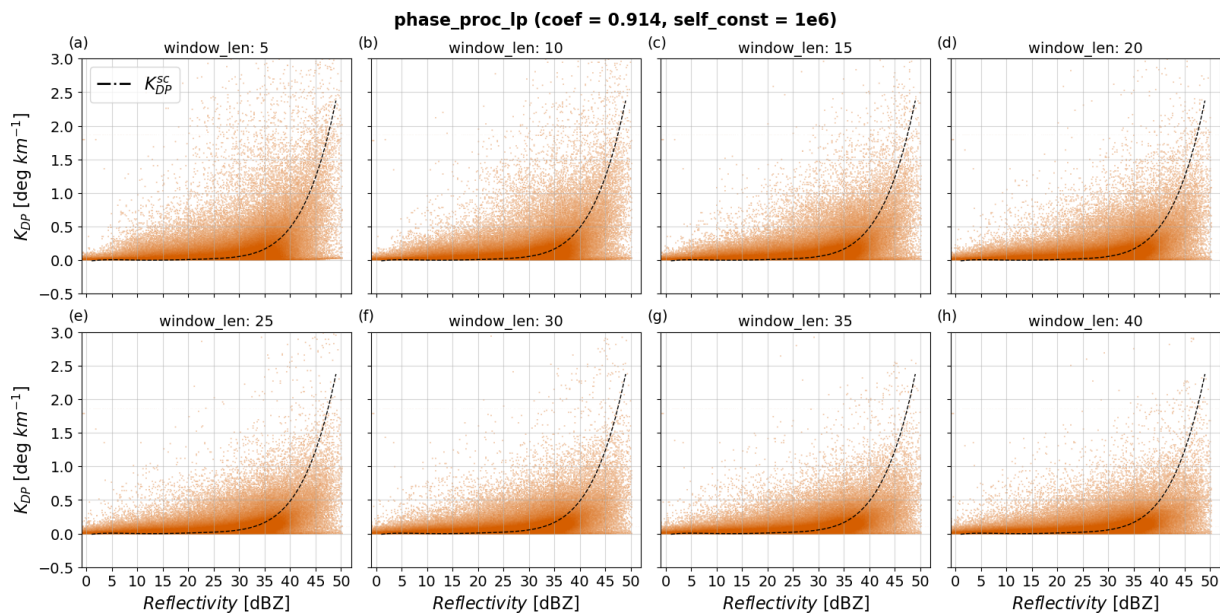


Figure A1. Scatterplots of estimated K_{DP} from phase_proc_lp as a function of reflectivity and for various values of window_len. Panels (a)–(h) show results with window_len values from 5 to 40 when fixing coef to 0.914 and self_const to 10^6 . The dashed black lines correspond to K_{DP}^{sc} .

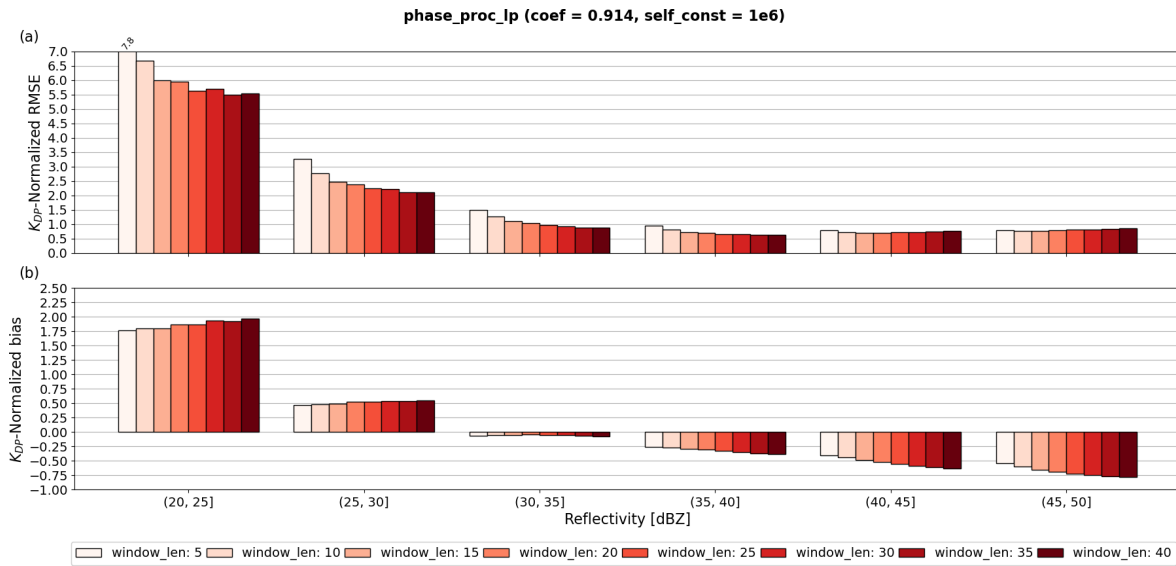


Figure A2. Panel (a) shows RMSE normalized by the interval-averaged K_{DP}^{SC} of phase_proc_lp relative to K_{DP}^{SC} as a function of reflectivity and for various values of window_len; panel (b) shows the same as panel (a) but for the normalized bias metric.

Data availability. The raw radar data and K_{DP} dataset can be accessed via the link at <https://doi.org/10.57707/fmi-b2share.4126c5db27d24ddea10d5c3163ff95a> (Aldana, 2024). This includes the raw radar data and the K_{DP} -processed data used to analyze the K_{DP} estimation methods. The data have been processed using Python and include the following.

- The radar folder includes several subfolders, such as yyyy/mm/dd/iris/raw/VAN. The VAN subfolder includes the raw radar with plan position indicators (PPIs) observed by Vantaa radar at an elevation angle of 0.7 for a specific time: yyyyymmddHHMM_VAN.PPI3_B.raw. This data can be read with Py-ART (Helmus and Collis, 2016, <https://doi.org/10.5334/jors.119>).
- The folder K_{DP_data} includes five .hdf5 files storing tables containing information about date (in pandas numerical value). It requires transformation to a date–time object), Z (in dBZ), Z_{dr} (in dB), attenuated gate (Boolean), theoretical or self-consistent K_{DP} (in $^{\circ} km^{-1}$), or computed K_{DP} (in $^{\circ} km^{-1}$) from a given method for different settings. The method is indicated in the name of the file as kdp_method_scatter.hdf5, where the method can be one of the following.
 - iris_sch refers to table containing K_{DP} from the iris software (used in the Finnish Meteorological Institute) and K_{DP} computed from Py-ART’s implementation of Schneebeli et al. (2014, <https://doi.org/10.1109/TGRS.2013.2287017>). These two methods were computed together because only one K_{DP} output was retrieved. They do not feature any user-configurable parameters to test.

- mae refers to table containing K_{DP} computed from Py-ART’s implementation of Maesaka et al. (2012, http://www.meteo.fr/cic/meetings/2012/ERAD/extended_abs/QPE_233_ext_abs.pdf). The columns correspond to K_{DP} computed by varying the parameter Clpf.
- vulpiani refers to a table containing K_{DP} computed from Py-ART’s implementation of Vulpiani et al. (2012, <https://doi.org/10.1175/JAMC-D-10-05024.1>). The columns correspond to K_{DP} computed by varying the parameters windsize and n_iter.
- pplp refers to table containing K_{DP} computed from Py-ART’s implementation of Giangrande et al. (2013, <https://doi.org/10.1175/JTECH-D-12-00147.1>). The columns correspond to K_{DP} computed by varying the parameter windowlen.
- wradlib refers to table containing K_{DP} computed from *wradlib*’s implementation of Vulpiani et al. (2012, <https://doi.org/10.1175/JAMC-D-10-05024.1>). The columns correspond to K_{DP} computed by varying the parameters winlen and dr.

The disdrometer dataset to obtain the DSD parameters can be accessed via the link provided in Moisseev (2024, <https://hdl.handle.net/21.12132/3.69dddc0004b64b32>).

Author contributions. MA conducted the investigation process, collected the data, and performed the formal analysis of the data and visualization; MA, SP, and DM designed the methodology; SP, AL, MK, and DM formulated the research goals and aims; AL and DM provided data; MA prepared the paper draft; and MA, SP, AL, MK, and DM reviewed, commented on, and edited the paper.

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