

Intercomparison of peroxy radical measurements obtained at atmospheric conditions by laser-induced fluorescence and electron spin resonance spectroscopy

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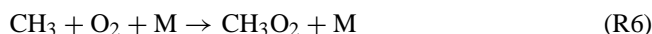
Abstract. Measurements of hydroperoxy radical (HO₂) and organic peroxy radical (RO₂) concentrations were performed by two different techniques in the atmospheric simulation chamber SAPHIR in Jülich, Germany. The first technique was the well-established Matrix Isolation Electron Spin Resonance (MIESR), which provides absolute measurements with a time resolution of 30 min and high accuracy (10%, 2σ). The other technique, ROxLIF, has been newly developed. It is based on the selective chemical conversion of RO_x radicals (HO₂ and RO₂) to OH, which is detected with high sensitivity by laser-induced fluorescence (LIF). ROxLIF is calibrated by quantitative photolysis of water vapor at 185 nm and provides ambient measurements at a temporal resolution of 1 min and accuracy of 20% (2σ). The measurements of HO₂ and RO₂ obtained by the two techniques were compared for two types of atmospheric simulation experiments. In one experiment, HO₂ and CH₃O₂ radicals were produced by photooxidation of methane in air at tropospheric conditions. In the second experiment, HO₂ and C₂H₅O₂ were produced by ozonolysis of 1-butene in air at dark conditions. The radical concentrations were within the range of 16 to 100 pptv for HO₂ and 12 to 45 pptv for RO₂. Good agreement was found in the comparison of the ROxLIF and MIESR measurements within their combined experimental uncertainties. Linear regressions to the combined data set yield slopes of 1.02±0.13 (1σ) for RO₂ and 0.98±0.08 (1σ) for HO₂ without significant offsets. The results confirm the calibration of the ROxLIF instrument and demonstrate that it can be applied with good accuracy for measurements of atmospheric peroxy radical concentrations.

1 Introduction

Peroxy radicals play an important role in the photochemistry of the troposphere (e.g., Finlayson-Pitts and Pitts, Jr., 2000). The major source of HO₂ and RO₂ (R = alkyl group) at day-time is the reaction of photochemically formed OH radicals with carbon monoxide and hydrocarbons (RH), respectively:



Another source of peroxy radicals is the photolysis of organic carbonyl compounds, for example, of acetaldehyde:

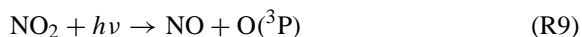


At night many atmospheric hydrocarbons can be oxidized by photolabile nitrate radicals (NO₃), leading to alkyl radicals and subsequently RO₂ (Atkinson and Arey, 2003). Peroxy radicals are also formed in the oxidation of olefins by their reaction with ozone, which can be another significant nighttime source of RO₂ and HO₂ (e.g., Geyer et al., 2003; Kanaya et al., 2007).

One reason for the importance of peroxy radicals is their ability to produce tropospheric ozone. In an NO_x (= NO + NO₂) rich environment, ozone is mainly destroyed and produced by photochemical cycling of NO and NO₂ without a net change of O₃ once steady state is achieved.



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However, if NO is oxidized by peroxy radicals, the subsequent NO₂ photolysis (R9) results in a net production of tropospheric ozone:



Furthermore, the peroxy radical Reactions (R11)–(R13) constitute an important source of OH in the atmosphere. For these reasons, peroxy radical concentration measurements are essential for quantifying (1) photochemical ozone production rates, and (2) radical budgets of OH, HO₂ and RO₂ in field studies.

Atmospheric measurements of peroxy radicals are a challenging task. Owing to their high reactivity, their maximum mixing ratios are small in the troposphere, on the order of 10 pptv during daytime (e.g., Monks, 2005). Thus, measurement instruments must have a high detection sensitivity. Reliable measurements are further complicated by the fact that the short-lived radicals can be lost by surface reactions in inlet systems of instruments. Lastly, calibrations of the instruments are difficult, because no permanently stable calibration mixtures exist for atmospheric peroxy radicals. If required, quantitative amounts of radicals must be prepared in specially designed flow reactors at the time of calibration.

Various techniques have been reported for the measurement of RO₂ and HO₂: (1) matrix isolation and electron spin resonance spectroscopy, MIESR (Mihelcic et al., 1985, 1990), (2) peroxy radical chemical amplification, PERCA (Cantrell et al., 1984; Hastie et al., 1991; Burkert et al., 2001; Green et al., 2006) and (3) chemical ionization mass spectrometry, ROxMas (Reiner et al., 1997) and PerCIMS (Edwards et al., 2003). These techniques differ from each other with respect to the chemical speciation of the measured radicals. MIESR is most versatile, because it can distinguish between NO, NO₂, HO₂, CH₃C(O)O₂ and RO₂. The latter represents the sum of alkyl peroxy radicals, $\text{RO}_2 = \sum_i \text{RO}_2^{(i)}$, where *i* denotes different organic peroxy radical species. PERCAs measure RO_x (= RO₂ + RO + HO₂ + OH), which is a good proxy for RO₂ + HO₂, while OH and RO make a negligible contribution to RO_x at atmospheric conditions. The CIMS based techniques measure RO_x and distinguish the contributions of RO₂ and HO₂ by chemical modulation.

Another class of instruments applies laser-induced fluorescence (LIF) for measurement of OH and HO₂ (Hard et al., 1984; Hofzumahaus and Holland, 1993; Stevens et al., 1994; Brune et al., 1995; Kanaya et al., 2001; Creasey et al., 2003; Dusanter et al., 2008; Butler et al., 2008). These instruments sample ambient air by expansion into a low pressure volume, where OH is detected spectroscopically by laser excited fluorescence at 308 nm. This concept is also known as fluorescence assay with gas expansion, FAGE (Hard et al., 1984).

The detection of HO₂ in a FAGE instrument involves chemical conversion of HO₂ to OH by reaction with NO, with subsequent LIF detection of the OH product molecules.

Recently, we have further developed our LIF technique at Forschungszentrum Jülich to enable separate measurements of OH, HO₂ and RO₂ concentrations with one instrument (Fuchs et al., 2008). The setup involves one measurement channel for the specific detection of OH, while a second measurement channel (ROxLIF) is used for detection of RO_x. The ROxLIF concept involves a two-stage chemical conversion of RO_x into OH, which is then detected by laser induced fluorescence. Modulation of the chemical conversion conditions allows the distinction between HO₂ and RO₂.

In this study, we have tested the performance and calibration of the new ROxLIF instrument by a measurement comparison against the MIESR technique. The latter was chosen as a reference, because MIESR can distinguish between HO₂ and RO₂ radicals, and yields absolute measurements with high accuracy (Mihelcic et al., 1985, 1990). The intercomparison was performed in the atmospheric simulation chamber SAPHIR (Simulation of Atmospheric PHotochemistry In a large Reaction chamber) in Jülich. It provides a large volume of homogeneously mixed air, to which trace gases can be added at tropospheric concentration levels. For instrumental comparisons, it ensures that the instruments sample air of the same composition and at natural conditions, as was demonstrated for OH (Schlosser et al., 2007) and oxygenated volatile organic compounds (Apel et al., 2008). A further advantage of the use of a chamber is that experiments can be carried out at controlled conditions, allowing, for example, to keep radical concentrations constant at different levels. In this paper, we report the results of the measurement comparison between ROxLIF and MIESR for two sets of simulation experiments: (1) photochemical oxidation of methane by OH as an example for daytime production of peroxy radicals, and (2) ozonolysis of 1-butene as an example for a nighttime source of peroxy radicals.

2 Instrumental description

2.1 SAPHIR chamber and instrumentation

SAPHIR is designed to investigate photochemical processes under ambient conditions with atmospheric concentrations of trace gases. Detailed descriptions, focussing on particular experiments in SAPHIR, were given in the papers by Bohn and Zilken (2005), Rohrer et al. (2005), and Wegener et al. (2007). Therefore, only a short description is given here.

The reaction chamber consists of a double-walled FEP film of 125 μm to 250 μm thickness arranged in a cylindrical shape (inner diameter 5 m, length 20 m, volume 270 m³). The film is chemically inert and it has an acceptable light transmission over the whole wavelength range of ultraviolet and visible sunlight that reaches the ground. The chamber

can be exposed to sunlight or can be kept dark by employing a fast shutter system.

The space between both films of the chamber wall is flushed permanently with clean nitrogen to avoid uncontrolled diffusion of ambient trace gases into the inner volume of the chamber. SAPHIR is operated at ambient temperature and at a pressure which is slightly (30–50 hPa) above ambient pressure. Before experiments take place, the gas inside the chamber is exchanged by clean synthetic air at a maximum flow rate of 500 m³/h until the concentrations of trace gases are below the detection limits of the attached measurement instruments. Trace gases can be injected into the chamber and are mixed to the chamber air by a fan within a few minutes. Air which is consumed by measurement devices or is lost by chamber leakages is constantly replaced with clean synthetic air, in this study at a rate of approximately 7 m³/h.

If the reaction chamber is exposed to sunlight, two chemical compounds, nitrous acid (HONO) and formaldehyde (HCHO) are produced photolytically on the surface of the Teflon film (Rohrer et al., 2005; Karl et al., 2004). Although the detailed mechanism of these processes is not known, the source strengths of HONO and HCHO emissions into the chamber are well characterized by previous experiments. The quantification of these processes is important, because the following photolysis of HONO and HCHO in the gas phase constitutes a source of OH and HO₂, respectively.

In this work, parameters like temperature, pressure, relative humidity and experimental flows in the chamber were routinely monitored during all experiments. The solar actinic flux was measured by a spectroradiometer and used for the calculation of photolysis frequencies (Bohn and Zilken, 2005). NO, NO₂, and O₃ concentrations were measured by chemiluminescence detectors (Eco Physics, CLD 780TR). A gas chromatograph (Chrompack, Chrompack VOCair) measured the concentrations of non-methane hydrocarbons. In one set of experiments, 1-butene was ozonolyzed for peroxy radical generation in the chamber. Here, 1-butene was measured by the GC instrument and simultaneously monitored by the NO chemiluminescence detector in the so called “zero mode”, which serves to quantify and subtract interference signals from the reaction of alkenes with ozone. The interference signal of the chemiluminescence detector was calibrated against the GC measurements and provided a quasi-continuous measurement of 1-butene, while the time resolution of the GC was about 20 min.

2.2 ROxLIF technique

The LIF instrument applies one measurement channel for OH (Holland et al., 2003) and a second channel (ROxLIF) for the measurement of RO₂ and HO₂. ROxLIF is a newly developed technique, which has been described in detail elsewhere (Fuchs et al., 2008). It applies chemical conversion in combination with LIF for radical detection and can be operated in two chemical modes (RO_x and HO_x mode).

In brief, ambient air is sampled through an orifice (1 mm diameter) into a differentially pumped flow reactor, which is kept at a reduced pressure of 25 hPa. In the RO_x mode, NO and CO are added as reagents to the sampled air, in order to convert atmospheric RO_x to HO₂. NO (0.7 ppmv) is responsible for the successive conversion of atmospheric RO₂ to HO₂ and HO₂ to OH (Reactions R11–R13). In contrast to previous detection of HO₂ by LIF, where RO₂ conversion was limited by the short reaction time and dilution of oxygen required in Reaction (R12), the reaction time of 0.6 s is long enough and the oxygen concentration high enough for quantitative conversion. The large amount of CO (0.17%) ensures that all HO_x (=HO₂+OH), from ambient air and from RO₂ conversion, becomes HO₂ (Reactions R1 and R2). This has the advantage that potential HO_x loss by reactions of the highly reactive OH becomes negligible in the reactor. In the HO_x mode, the NO reagent is turned off and only CO is added. At this condition, only atmospheric HO_x is converted to HO₂. The alternating measurement in the HO_x and RO_x mode allows to distinguish between HO_x and RO₂ radical concentrations.

At the exit of the reactor, the central gas flow is transferred through a nozzle into a detection chamber at even lower pressure (3.5 hPa). Here, additional NO is injected into the sampled gas, causing nearly complete chemical conversion of HO₂ to OH (Reaction R13), which is detected by LIF. A pulsed narrow-bandwidth UV laser (308 nm) is used to excite the OH radicals at the rovibronic line Q₁(3) of the A²Σ⁺ ν′=0–X²Π ν″=0 transition, and the resulting resonance fluorescence is measured by gated photon counting. The laser is tuned periodically on and off resonance to distinguish the OH fluorescence signal from non-resonant background signals.

The calibration of the ROxLIF measurements is based on the quantitative photolysis of water vapor in air at a wavelength of 185 nm. The photolysis produces equal concentrations of OH and HO₂ and is widely applied for calibration of HO_x instruments (Aschmutat et al., 1994; Schultz et al., 1995; Heard and Pilling, 2003). In the present work, a modified concept is used for peroxy radicals (Schultz et al., 1995; Hanke et al., 2002; Qi et al., 2006; Fuchs et al., 2008). The photolytically generated OH is completely converted to HO₂ by adding excess CO (Reactions R1 and R2) in order to determine the HO₂ sensitivity. To calibrate the ROxLIF instrument for RO₂, a corresponding hydrocarbon (RH) is added to the calibration gas, resulting in the complete conversion of OH to RO₂ (Reactions R3 and R4). In the present work only methane was used for RO₂ calibrations. This simplified procedure is justified as the ROxLIF detection sensitivity is essentially the same for the two radical species (CH₃O₂ and C₂H₅O₂), which were generated in the SAPHIR experiments reported below (Fuchs et al., 2008). The similar detectivity is plausible, since alkyl peroxy radicals show similar reaction rate constants towards their reaction with NO (Atkinson and Arey, 2003), which is the rate limiting step in

the RO₂ to HO₂ conversion in the flow reactor of the instrument. The estimated accuracy of calibration is approximately 20% (2σ) limiting the accuracy of ROxLIF measurements. The sensitivity of the instrument is corrected for fluorescence quenching by water vapor, but there is no chemical interference from water vapor as observed in PERCA instruments (Mihele and Hastie, 1998). If sampling highly polluted air, interferences such as thermal decomposition of atmospheric peroxy nitrates could reduce the accuracy, but such conditions were not encountered in the current experiments. The detection limits (SNR=2) are 0.1 pptv for both HO₂ and RO₂ at a time resolution of 1 min (Fuchs et al., 2008).

In the reported experiments, the flow reactor of the instrument extended into the SAPHIR chamber. The sampling point was 87 cm above the floor of the chamber. The outer surfaces of the reactor (except the inlet region) were covered with a Teflon bag in order to minimize possible radical losses on instrumental surfaces close to the inlet.

2.3 MIESR technique

MIESR was developed as a measurement technique for atmospheric peroxy radicals at the Forschungszentrum Jülich by Mihelcic et al. (1985, 1990) and is the only one of its kind. It has the capability to measure simultaneously the concentrations of HO₂, RO₂, CH₃C(O)O₂, NO₃ and NO₂ in the same sample of air.

The MIESR technique is an offline method. Ambient air is collected by gas expansion through a nozzle into a vacuum chamber. The radicals are trapped from the gas jet in a polycrystalline D₂O matrix which is formed in situ on a cold finger at a temperature of 77 K. The collection time for an air sample is 30 min. The samples are stored and measured under vacuum at 77 K. At this condition, the ice matrix prevents chemical interactions between reactive molecules and allows to store the sample up to several weeks without loss of radicals. Speciation and quantification of the trapped radicals are achieved in the laboratory by electron spin resonance spectroscopy. The spectra are analyzed by a nonlinear fitting procedure (Mihelcic et al., 1990). HO₂, CH₃C(O)O₂, NO₃ and NO₂ can be individually measured, whereas the spectra of organic peroxy radicals overlap and provide a measure of RO₂. The MIESR technique is an absolute technique and achieves a detection limit of 2 pptv for peroxy radicals at a time resolution of 30 min (Mihelcic et al., 2003). The main advantage of this method is that no calibration of the instrument is required, making it useful as a reference method for other atmospheric RO_x measurement techniques (Platt et al., 2002). The accuracy of measurements is 5%.

In the reported experiments, the MIESR sampling unit was mounted in a flange at the bottom of the SAPHIR chamber. In this setup, the upward oriented inlet nozzle of the sampling unit was about 2 cm above the chamber floor, which consists of a Teflon film. The horizontal distance between the MIESR sampling unit and the ROxLIF reactor was approximately

3 m. Turbulent mixing of the chamber air ensures that the relatively short distance between sampling point and chamber surface does not lead to gradients in the radical concentrations, which would affect the intercomparison (Schlosser et al., 2007).

3 Results

Two types of experiments were carried out in SAPHIR which provided atmospheric radical concentrations. The first experiment was performed at daylight and involved the photochemical oxidation of methane, while the second experiment involved ozonolysis of 1-butene at dark conditions. Raw data were evaluated independently by the operators of the two instruments without communicating results before final data were exchanged.

3.1 Methane photooxidation experiment

After purging the chamber with dry synthetic air overnight, water vapor was introduced into the dark chamber in the morning. High purity water was provided by a Milli-Q-water device (Millipore Corp., Milli-Q Gradient A10), vaporized and flushed into the chamber together with a high flow of synthetic air, until a water vapor mixing ratio of about 0.7% by volume (relative humidity: 50%) was achieved. Methane (Air Liquide, purity 5.5) was introduced into the chamber, resulting in an initial large mixing ratio of approximately 0.5% by volume. Fast mixing in the chamber was achieved by a fan, which was operated for the entire duration of the experiment. After the gas addition, the shutter system of the chamber was opened at 08:00 UTC for a duration of 5.5 h. The experiment was stopped at 14:30 UTC. The ROxLIF instrument was running continuously during the whole period, while the MIESR instrument took three air samples of 30 min integration time, starting at 09:00, 10:00 and 11:00 UTC.

The peroxy radical measurements of the ROxLIF and MIESR instrument are shown in Fig. 1, together with the HONO photolysis frequency, $j(\text{HONO})$, and the mixing ratio of NO. Dashed vertical lines indicate the opening and closing of the shutter system. As soon as the shutter was opened and sunlight entered the chamber, HO₂ and RO₂ concentrations measured by ROxLIF increased during the first minutes up to 15–20 pptv. Both, HO₂ and RO₂, showed similar diurnal variations and reached approximately the same peak concentrations of 40 pptv at 12:30 UTC. The concentrations measured by ROxLIF agree very well with the simultaneous measurements by MIESR during the time of measurement overlap (9:00–11:30 UTC). Short-term fluctuations of the radical concentrations measured by ROxLIF were mainly correlated with variations of the solar UV radiation represented by $j(\text{HONO})$. The OH radical concentration (not shown here) remained always below the detection limit ($<1 \times 10^6 \text{ cm}^{-3}$) of the LIF instrument. After the shutter of

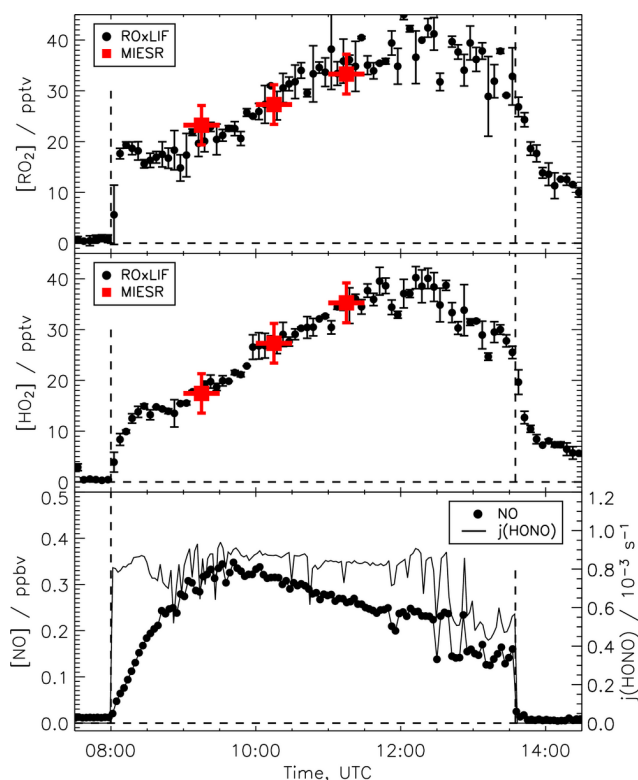


Fig. 1. Comparison of peroxy radical concentrations measured simultaneously by ROxLIF (5 min average values) and MIESR during the methane photooxidation experiment (upper panel: RO₂, middle panel: HO₂). Vertical error bars represent the 1 σ precisions of the measurements. Horizontal bars denote the 30 min sampling time intervals of the MIESR measurements. The bottom panel shows mixing ratios of NO measured by chemiluminescence and the HONO photolysis frequency measured by a spectroradiometer. Dashed vertical lines indicate the opening and closing of the shutter system of the chamber.

the chamber had been closed, the photolytic radical production stopped, followed by a decay of the radical concentrations. The mixing ratios of NO_x and O₃ (not shown here) increased continuously over the course of the experiment, starting at zero concentrations at 08:00 UTC and reached maximum values of 2.8 ppbv and 65 ppbv at 13:30 UTC, respectively.

3.2 1-butene ozonolysis experiment

The ozonolysis experiment was performed at dark conditions, keeping the shutter system of the chamber closed all the time. After purging the chamber overnight with clean synthetic air, ozone was generated by a silent discharge ozonizer and introduced into the chamber. The amount of added O₃ was accurately controlled by a magnetic valve in the inlet line, which was opened and closed by an electronic timer. The ozone mixing ratio was increased in four steps

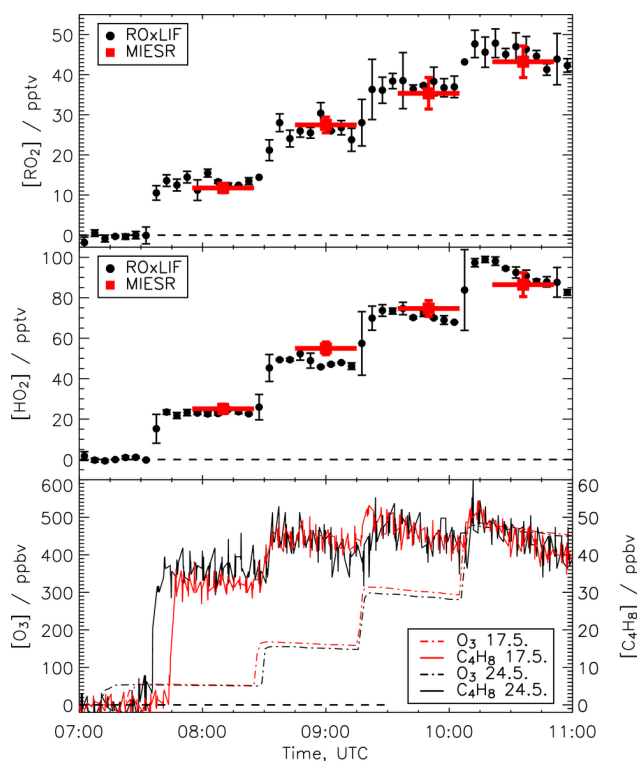


Fig. 2. Comparison of peroxy radical concentrations measured by ROxLIF (5 min average values) and MIESR during the 1-butene ozonolysis experiments on 24 May 2006 and 17 May 2006, respectively (upper panel: RO₂, middle panel: HO₂). Vertical error bars represent the 1 σ precisions of the measurements. Horizontal bars denote the 30 min sampling time intervals of the MIESR measurements. The bottom panel shows the corresponding concentrations of ozone and 1-butene during the two experimental runs.

from 50 ppbv to a maximum of nearly 500 ppbv (Fig. 2). The water vapor mixing ratio was about 0.04% by volume all the time. In addition, 480 ppmv CO were initially introduced to the chamber, serving as a scavenger for OH and stabilized Criegee radicals. After the first addition of ozone, 1-butene (Linde AG, purity 3.5) was injected into the chamber, yielding an initial concentration of 36 ppbv. The slow decrease of 1-butene during the following experiment was balanced by further additions of 1-butene (each corresponding to an incremental increase of 12 ppbv), added at the same time when the ozone concentration was increased. The complete experiment was performed twice, once on 17 May 2006 and a second time on 24 May 2006, at nearly the same chemical conditions (see bottom panel of Fig. 2).

Nearly constant mixing ratios of HO₂ and RO₂ were produced in the chamber at five different concentration levels reaching from zero to a maximum of 100 pptv of HO₂ and 45 pptv of RO₂. The stepwise increase of concentrations was caused by the incremental additions of O₃ and 1-butene. In the first experimental run (17 May 2006), the

MIESR instrument collected four samples, each at a different level of peroxy radical concentrations for 30 min. The sampling was started approximately 15 min after each addition of O₃ and 1-butene, in order to ensure that the air in the chamber was homogeneously mixed and the chemical system was equilibrated. The ROxLIF instrument was running in parallel, but due to a technical failure of its flow reactor, measurements became flawed and had to be discarded. The ROxLIF measurements were repeated in the second experimental run on 24 May 2006, where the chemical conditions from 17 May 2006 were closely reproduced. This time the ROxLIF system worked properly, however, the MIESR sampler was no longer available, because it had been moved to the laboratory for analysis of the previously taken samples. Thus, the ROxLIF and MIESR data shown in Fig. 2 were not obtained simultaneously. However, as the chemical conditions were nearly the same, a comparison of the two data sets is possible (see discussion below). Based on this assumption, the results demonstrate a good agreement between the ROxLIF and MIESR technique for measurement of both HO₂ and RO₂ in the ozonolysis experiments.

4 Discussion

4.1 Peroxy radical production by methane oxidation

The chamber experiments were designed to produce HO₂ and one particular RO₂ species in each type of experiment. In the methane photooxidation experiment, traces of methane and water vapor were exposed to solar radiation in synthetic air at ambient temperature and pressure. It is expected that this mixture produces CH₃O₂ and HO₂ as the only peroxy radicals in the SAPHIR chamber at the given conditions. As soon as sunlight enters the chamber, i.e. when the shutter system is opened, production of OH is started by photolysis of HONO, which is formed heterogeneously at the sunlit chamber walls (see Sect. 2.1). The HONO photolysis is also a source for NO in the gas phase, which starts to build up when the chamber is illuminated. The high concentration of methane (0.5%) ensures that OH reacts quickly and exclusively to CH₃O₂, which is further converted by the available NO to HO₂. Recycled OH from the reaction of HO₂ and NO is immediately consumed by methane again. This sequence of radical reactions explains why the observed concentrations of CH₃O₂ and HO₂ exhibit a similar time dependence and show similar concentrations during the course of the experiment. HO₂ is formed predominantly from CH₃O₂, and both, HO₂ and CH₃O₂, are depleted mainly by reaction with NO, which has nearly the same rate constant for both radical species (Sander et al., 2006). Thus, approximately equal steady-state concentrations are expected for both radical species.

The high methane concentration in the chamber led to a reduction in the ROxLIF sensitivity, because OH radicals pro-

duced in the conversion reactor partly regenerated CH₃O₂ radicals. Model calculations of the reaction kinetics in the reactor show a reduction of the sensitivity of 4% for RO₂ and 9% for HO₂ for a methane mixing ratio of 0.5%. Since methane was not measured during this experiment and the initial mixing ratio was only estimated from its addition before the experiment started, this effect was not taken into account in the data evaluation and adds to the uncertainty in the ROxLIF measurements.

4.2 Peroxy radical production by 1-butene ozonolysis

In the dark ozonolysis experiments, ozone and 1-butene were mixed in 1 atm of dry synthetic air, which was additionally loaded with CO as an OH scavenger. The ozonolysis involves a complex system of chemical reactions, which yield various products (Grosjean et al., 1996; Grosjean and Grosjean, 1997; Olzman et al., 1997; Hasson et al., 2001; Wegener et al., 2007). In the first step, ozone is cycloadded to the alkene and forms an ozonide, which rapidly decomposes to either formaldehyde or propanal and a corresponding excited Criegee intermediate, C₂H₅CHOO* or HCHOO*, respectively. The two types of Criegee intermediates undergo fast decomposition reactions, yielding OH, HO₂ and C₂H₅O₂ radicals and stable byproducts (e.g., C₂H₆, CO, CO₂) or can be stabilized by collisions. The stabilized Criegee intermediates react with the large excess of CO initially injected into the chamber to corresponding aldehydes and CO₂ (Atkinson, 1997; Hatakeyama et al., 1994). The large amount of CO also converts OH completely to HO₂ and prevents possible reactions of OH with 1-butene and its ozonolysis products. Thus, the chemical system is expected to produce C₂H₅O₂ and HO₂ as the only peroxy radicals in the chamber.

Interferences of the ROxLIF measurement from Criegee intermediates are not expected in these experiments. The excited Criegee radicals from the ozonolysis reactions are short-lived intermediates in the formation process of the peroxy radicals, and as their chemical lifetime is very much shorter, their concentrations are negligible compared to those of the peroxy radicals (Jenkin et al., 1997). Stabilized Criegee radicals also do not play a role, because they are efficiently scavenged and converted to non-radical products by the high concentration of CO in the SAPHIR chamber and in the ROxLIF converter (Atkinson, 1997; Hatakeyama et al., 1994).

Peroxy radical concentrations are controlled by the concentrations of the precursors ozone and 1-butene, which slowly react with each other on a time scale of several hours, at a rate constant of $\sim 1 \times 10^{-17} \text{ cm}^3 \text{ s}^{-1}$ (Atkinson and Arey, 2003). While the radical production is rate limited by this reaction, the radical destruction in the absence of NO is determined by the much faster recombination reactions of the peroxy radicals (HO₂+HO₂, HO₂+RO₂, and RO₂+RO₂). The overall reaction kinetics has two interesting consequences. First, as 1-butene and ozone are depleted

slowly, the resulting steady-state concentrations of the peroxy radicals stay nearly constant over the 1 h observation periods applied in the ozonolysis experiments (Fig. 2). Second, the radical concentrations can be easily reproduced by adjustment of the ozone and alkene concentrations.

The chemical behavior of the reaction system and the fact that the two ozonolysis experiments (17 and 24 May 2006) were performed in clean synthetic air and at dark conditions, avoiding variable photolytical radical production, is the basis for the reproducibility of the experiments. In fact, the amount of reagents for radical production and the timing of their incremental additions were almost the same during both runs (see Fig. 2, bottom panel). The maximum difference in the measured ozone concentrations was 3% during the time intervals, where MIESR and ROxLIF measurements are compared. The corresponding concentrations of 1-butene differed by 12% after the first addition, but were equal for the following concentration levels of the experiments within the precision of the 1-butene measurement. For these reasons, it can be assumed that the time series of the relevant peroxy radical concentrations were essentially the same during both runs. Based on this assumption, it is justified to compare the measurements by ROxLIF and MIESR for both HO₂ and RO₂ in the ozonolysis experiments.

4.3 Measurement intercomparison

For quantitative analysis, the ROxLIF measurements are averaged over the sampling time period of the MIESR instrument and are compared in scatter plots versus the MIESR measurements in Fig. 3.

The HO₂ data in the right panel show a linear relationship with little scatter and follow closely a 1:1 line. A linear fit, which takes into account the statistical errors of both instruments ("FitExy" procedure in Press et al. (1992)), yields a slope of 0.98 ± 0.08 , an intercept of (-1.2 ± 3.8) pptv and a linear correlation coefficient R^2 of 0.98. The scatter of the individual data points around the regression line can be explained by their statistical precision. This is confirmed by the goodness-of-fit probability q of 0.65. This fitting parameter indicates a linear relationship of the data within their statistical errors, if $q > 0.1$ (Press et al., 1992). The intercept of the regression line is not significant and can be explained by its statistical error.

The scatter plot for RO₂ in the left panel of Fig. 3 shows a similar good correlation as for HO₂. The data follow a linear relationship of nearly unity slope and exhibit a correlation coefficient $R^2 = 0.97$. No significant difference can be seen between the data points from the methane photooxidation and 1-butene ozonolysis experiments. This confirms that the ROxLIF instrument, which was calibrated for CH₃O₂ (see Sect. 2.2), is equally sensitive to CH₃O₂ and C₂H₅O₂. The regression line for the combined RO₂ data exhibits a slope of 1.02 ± 0.13 , an intercept of (0.5 ± 4.0) pptv and a parameter q of 0.93. As for HO₂, no significant offset is observed.

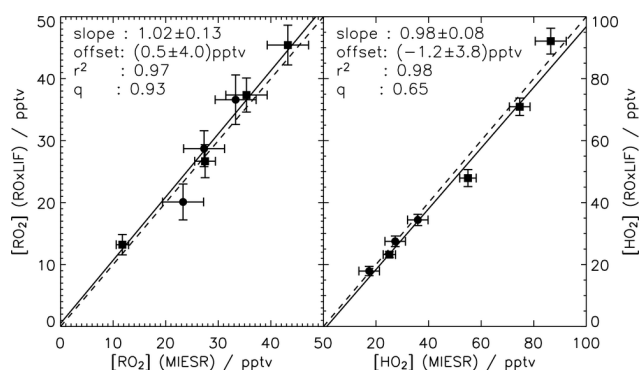


Fig. 3. Correlation of peroxy radical data measured by ROxLIF and MIESR (left panel: RO₂, right panel: HO₂). Square symbols represent data from the ozonolysis experiments and circles from the methane photooxidation experiment. For these presentations the ROxLIF data were averaged over the sampling time interval of MIESR. Error bars are 1 σ standard deviations of average ROxLIF values or 1 σ measurement errors of MIESR. The black lines represent linear fits that take into account the errors of both measurement instruments (Press et al., 1992). The corresponding fitting parameters (slope and offset, correlation coefficient r^2 and q as the probability for a linear fit) are given with their 1 σ errors in the legends. The dashed lines represent 1:1 lines.

The slopes of the regression lines for HO₂ and RO₂ indicate very good agreement of the two measurement techniques within their combined uncertainties. The uncertainties are determined by the calibration accuracy of 20% (2σ) of the ROxLIF instrument (Fuchs et al., 2008) and the 10% (2σ) measurement error of MIESR (Mihelcic et al., 2003). As a result, the measurement comparison confirms the calibration of the ROxLIF technique, here demonstrated for HO₂, CH₃O₂ and C₂H₅O₂.

The present work is the first, in which two measurement instruments have been compared for speciated peroxy radicals at atmospheric conditions. Earlier studies have either compared measurements of atmospheric HO₂ or total peroxy radicals. HO₂ measurements were first compared between LIF and MIESR during the BERLIOZ field campaign (Platt et al., 2002; Mihelcic et al., 2003). Good agreement was observed as indicated by a slope of 1.03 ± 0.08 and a correlation coefficient R^2 of 0.88 (Platt et al., 2002). The same campaign offered also an opportunity to compare RO_x measurements by a PERCA instrument and MIESR. A regression line with a slope of 1.07 and a correlation coefficient $R^2 = 0.91$ demonstrated good agreement for these measurements (Platt et al., 2002). In another more recent study, HO₂ measurements were compared between a PerCIMS and LIF instrument in two phases: 1. by mutual exchange of calibration sources, 2. by ambient air measurements (Ren et al., 2003). In the calibration intercomparison, very good agreement was found within 2% for PerCIMS sampling from the LIF calibration

source and within 4% for LIF sampling from the PerCIMS calibration source. Good agreement was also obtained in a side-by-side intercomparison of the ambient HO₂ measurements. Here, the regression line yielded a slope of 0.96 and a correlation coefficient $R^2=0.85$. It is noteworthy that the calibrations of both techniques applied water vapor photolysis at 185 nm, which is also the basis for the ROxLIF calibration in the present work. MIESR, however, which was used as a reference in the present and previous intercomparisons, is an independent technique requiring no calibration. The result of the present intercomparison is therefore important support for the general concept of water photolysis as a tool for peroxy radical calibrations.

5 Conclusions

An intercomparison of speciated measurements of concentrations of HO₂ and two alkyl peroxy radicals, CH₃O₂ and C₂H₅O₂, has been performed successfully between the fundamentally different ROxLIF and MIESR techniques. Very good agreement of the two instruments was observed during two types of simulation experiments carried out in the atmosphere simulation chamber SAPHIR. This facility ensured that both instruments sampled well mixed air of the same chemical composition and allowed the design of experiments, in which only one specific type of RO₂ radical is produced together with HO₂. In the photooxidation of methane, concentrations of CH₃O₂ and HO₂ were prepared, which were dependent on the strength of the natural sunlight. In the 1-butene ozonolysis experiments, nearly constant radical concentrations of C₂H₅O₂ and HO₂ were generated, which were entirely controlled by the chosen concentrations of the reactants ozone and 1-butene.

The good agreement of the measurements by the newly developed ROxLIF and the well established MIESR technique strengthens the confidence in the measurement and calibration of the new instrument. The ROxLIF technique has advantages over the MIESR technique by offering a higher time resolution and a better detection sensitivity. The SAPHIR chamber offers the potential to serve as a reference standard for peroxy radical concentrations. As has been demonstrated in this work, the ozonolysis of 1-butene allows the controlled production of HO₂ and C₂H₅O₂ at dark conditions. The capability to create constant radical concentrations is of advantage for intercomparison of instruments, which need a relative long sampling time. The peroxy radical concentrations obtained from the 1-butene ozonolysis have been quantified in this work and can be reproduced at any time in the chamber for future quality assurance tests.

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References

- Apel, E. C., Brauers, T., Koppmann, R., Bandowe, B., Boßmeyer, J., Holzke, C., Tillmann, R., Wahner, A., Wegener, R., Brunner, A., Jocher, M., Ruuskanen, T., Spirig, C., Steigner, D., Steinbrecher, R., Gomez Alvarez, E., Müller, K., Burrows, J. P., Schade, G., Solomon, S. J., Ladstätter-Weissenmayer, A., Simmonds, P., Young, D., Hopkins, J. R., Lewis, A. C., Legreid, G., Reimann, S., Hansel, A., Wisthaler, A., Blake, R. S., Ellis, A. M., Monks, P. S., and Wyche, K. P.: Intercomparison of oxygenated volatile organic compound measurements at the SAPHIR atmosphere simulation chamber, *J. Geophys. Res.*, 113, D20307, doi:10.1029/2008JD009865, 2008.
- Aschmutat, U., Heßling, M., Holland, F., and Hofzumahaus, A.: A tunable source of hydroxyl (OH) and hydroperoxy (HO₂) radicals: In the range between 10⁶ and 10⁹ cm⁻³, in: *Physico-chemical behaviour of atmospheric pollutants*, edited by G. Angeletti and G. Restelli, European Commission, Brussels, 811–816, 1994.
- Atkinson, R.: Gase-phase tropospheric chemistry of volatile organic compounds: 1. alkanes and alkenes, *J. Phys. Chem. Ref. Data*, 26, 217–219, 1997.
- Atkinson, R. and Arey, J.: Atmospheric Degradation of Volatile Organic Compounds, *Chem. Rev.*, 103, 4605–4638, 2003.
- Bohn, B. and Zilken, H.: Model-aided radiometric determination of photolysis frequencies in a sunlit atmosphere simulation chamber, *Atmos. Chem. Phys.*, 5, 191–206, 2005, <http://www.atmos-chem-phys.net/5/191/2005/>.
- Brune, W. H., Stevens, P. S., and Mather, J. H.: Measuring OH and HO₂ radicals in the troposphere by laser-induced fluorescence at low pressure, *J. Atmos. Sci.*, 52, 3328–3336, 1995.
- Burkert, J., Andrés Hernández, M.-D., Stöbener, D., Weissenmayer, M., and Kraus, A.: Peroxy radical and related trace gas measurements in the boundary layer above the Atlantic Ocean, *J. Geophys. Res.*, 106, 5457–5477, 2001.
- Butler, T. M., Taraborrelli, D., Brühl, C., Fischer, H., Harder, H., Martinez, M., Williams, J., Lawrence, M. G., and Lelieveld, J.: Improved simulation of isoprene oxidation chemistry with the ECHAM5/MESSy chemistry-climate model: lessons from the GABRIEL airborne field campaign, *Atmos. Chem. Phys.*, 8, 4529–4546, 2008, <http://www.atmos-chem-phys.net/8/4529/2008/>.
- Cantrell, C. A., Stedman, D. H., and Wendel, G. J.: Measurement of atmospheric peroxy radicals by chemical amplification, *Anal. Chem.*, 56, 1496–1502, 1984.
- Creasey, D. J., Evans, G. E., Heard, D. E., and Lee, J. D.: Measurements of OH and HO₂ concentrations in the Southern Ocean marine boundary layer, *J. Geophys. Res.*, 108, 4475, doi:10.1029/2002JD003206, 2003.
- Dusanter, S., Vimal, D., and Stevens, P. S.: Technical note: Measuring tropospheric OH and HO₂ by laser-induced fluorescence at low pressure. A comparison of calibration techniques, *Atmos. Chem. Phys.*, 8, 321–340, 2008, <http://www.atmos-chem-phys.net/8/321/2008/>.
- Edwards, G. D., Cantrell, C. A., Stephens, S., Hil, B., Goyea, O., Shetter, R. E., Mauldin, R. L., Kosciuch, E., Tanner, D. J., and Eisele, F. L.: Chemical ionization mass spectrometer instrument for the measurement of tropospheric HO₂ and RO₂, *Anal. Chem.*, 75, 5318–5327, 2003.

- Finlayson-Pitts, B. J. and Pitts Jr., J. N.: Chemistry of the upper and lower atmosphere, Academic Press, 2000.
- Fuchs, H., Hofzumahaus, A., and Holland, F.: Measurement of tropospheric RO₂ and HO₂ radicals by a laser-induced fluorescence instrument, *Rev. Sci. Instr.*, 79, 084104, doi:10.1063/1.2968712, 2008.
- Geyer, A., Bächmann, K., Hofzumahaus, A., Holland, F., Konrad, S., Klüpfel, T., Pätz, H. W., Schäfer, J., and Platt, U.: Nighttime formation of peroxy and hydroxyl radicals during the BERLIOZ campaign: Observations and modeling studies, *J. Geophys. Res.*, 108, 8249, doi:10.1029/2001JD000656, 2003.
- Green, T. J., Reeves, C. E., Fleming, Z. L., Brough, N., Rickard, A. R., Bandy, B. J., Monks, P. S., and Penkett, S. A.: An improved dual channel PERCA instrument for atmospheric measurements of peroxy radicals, *J. Environ. Monit.*, 8, 530–536, 2006.
- Grosjean, E. and Grosjean, D.: Gas phase reaction of alkenes with ozone: Formation yields of primary carbonyls and biradicals, *Environ. Sci. Technol.*, 31, 2421–2427, 1997.
- Grosjean, E., Bittencourt de Andrade, J., and Grosjean, D.: Carbonyl products of the gas-phase reaction of ozone with simple alkenes, *Environ. Sci. Technol.*, 30, 975–983, 1996.
- Hanke, M., Uecker, J., Reiner, T., and Arnold, F.: Atmospheric peroxy radicals: ROXMAS, a new mass-spectrometric methodology for the speciated measurements of HO₂ and Σ RO₂ and first results, *Int. J. Mass. Spectrosc.*, 213, 91–99, 2002.
- Hard, T. M., O'Brian, R. J., Chan, C. Y., and Mehrabzadeh, A. A.: Tropospheric free radical determination by FAGE, *Environ. Sci. Technol.*, 18, 768–777, 1984.
- Hasson, A. S., Ho, A. W., Kuwata, K. T., and Paulson, S. E.: Production of stabilized Criegee intermediates and peroxides in the gas phase ozonolysis of alkenes: 2. asymmetric and biogenic alkenes, *J. Geophys. Res.*, 106, 34143–34154, 2001.
- Hastie, D. R., Weissenmayer, M., Burrows, J. P., and Harris, G. W.: Calibrated chemical amplifier for atmospheric RO_x measurements, *Anal. Chem.*, 63, 2048–2057, 1991.
- Hatakeyama, S., and Kobayashi, H., and Akimoto, H.: Gas-phase oxidation of SO₂ in the ozone-olefin reactions, *J. Phys. Chem.*, 88, 4736–4739, 1994.
- Heard, D. E. and Pilling, M. J.: Measurement of OH and HO₂ in the Troposphere, *Chem. Rev.*, 103, 5163–5198, 2003.
- Hofzumahaus, A. and Holland, F.: Laser-induced fluorescence based detection system for measurement of tropospheric OH using 308 nm excitation at low pressure, *Proc. SPIE*, 1715, 163–173, 1993.
- Holland, F., Hofzumahaus, A., Schäfer, J., Kraus, A., and Pätz, H.-W.: Measurements of OH and HO₂ radical concentrations and photolysis frequencies during BERLIOZ, *J. Geophys. Res.*, 108, 8246, doi:10.1029/2001JD001393, 2003.
- Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: The tropospheric degradation of volatile organic compounds: A protocol for mechanism development, *Atmos. Environ.*, 31, 81–104, 1997.
- Jenkin, M. E., Saunders, S. M., Wagner, V., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile organic compounds, *Atmos. Chem. Phys.*, 3, 181–193, 2003, <http://www.atmos-chem-phys.net/3/181/2003/>.
- Kanaya, Y., Sadanaga, Y., Nakamura, K., and Akimoto, H.: Development of a ground-based LIF instrument for measuring HO_x radicals: Instrumentation and calibration, *J. Atmos. Chem.*, 38, 73–110, 2001.
- Kanaya, Y., Cao, R., Akimoto, H., Fukuda, M., Komazaki, Y., Yokouchi, Y., Koike, M., Tanimoto, H., Takegawa, N., and Kondo, Y.: Urban photochemistry in central Tokyo: 1. observed and modeled OH and HO₂ radical concentrations during the winter and summer 2004, *J. Geophys. Res.*, 112, D06301, doi:10.1029/2007JD008670, 2007.
- Karl, M., Brauers, T., Dorn, H.-P., Holland, F., Komenda, M., Poppe, D., Rohrer, F., Rupp, L., Schaub, A., and Wahner, A.: Kinetic study of the OH-isoprene and O₃-isoprene reaction in the atmosphere simulation chamber SAPHIR, *Geophys. Res. Lett.*, 31, L05117, doi:10.1029/2003GL019189, 2004.
- Mihelcic, D., Müsgen, P., and Ehhalt, D. H.: An improved method of measuring tropospheric NO₂ and RO₂ by matrix isolation and electron spin resonance, *J. Atmos. Chem.*, 3, 341–361, 1985.
- Mihelcic, D., Volz-Thomas, A., Pätz, H. W., Kley, D., and Mihelcic, M.: Numerical analysis of ESR spectra from atmospheric samples, *J. Atmos. Chem.*, 11, 271–297, 1990.
- Mihelcic, D., Holland, F., Hofzumahaus, A., Hoppe, L., Müsgen, P., Pätz, H. W., and Moortgat, G. K.: Peroxy radicals during BERLIOZ at Pabstthum: Measurements, radical budgets and ozone production, *J. Geophys. Res.*, 108, 9–1, 2003.
- Mihele, C. M. and Hastie, D. R.: The sensitivity of the radical amplifier to ambient water vapor, *Geophys. Res. Lett.*, 25, 1911–1913, 1998.
- Monks, P. S.: Gas-phase radical chemistry in the troposphere, *Chem. Soc. Rev.*, 34, 376–395, 2005.
- Olzman, M., Kraka, E., Cremer, D., Gutbrod, R., and Andersson, S.: Energetics, kinetics and product distributions of the reactions of ozone with ethene and 2,3-dimethyl-2-butene, *J. Phys. Chem. A*, 101, 9421–9429, 1997.
- Platt, U., Alicke, B., Dubois, R., Geyer, A., Hofzumahaus, A., Holland, F., Martinez, M., and Stutz, J.: Free radicals and fast photochemistry during BERLIOZ, *J. Atmos. Chem.*, 42, 359–394, 2002.
- Press, W. H., Teukolsky, S. A., Vetterling, W. T., and Flannery, B. P.: Straight-line data with errors in both coordinates, in: *Numerical recipes in C*, Cambridge University Press, Cambridge, 666–670, 1992.
- Qi, B., Takami, A., and Hatakeyama, S.: A calibration method for measurement of small alkyl organic peroxy radicals by chemical amplifier, *Anal. Sci.*, 22, 1091–1093, 2006.
- Reichert, L., Andrés Hernández, M. D., Stöbener, D., Burkert, J., and Burrows, J. P.: Investigation of the effect of water complexes in the determination of peroxy radical ambient concentrations: Implications for the atmosphere, *J. Geophys. Res.*, 108, 4017, doi:10.1029/2002JD002152, 2003.
- Reiner, T., Hanke, M., and Arnold, F.: Atmospheric peroxy radical measurement by ion molecule reaction-mass spectroscopy: A novel analytic method using amplifying chemical conversion to sulfuric acid, *J. Geophys. Res.*, 102, 1311–1326, 1997.
- Ren, X., Edwards, G. D., Cantrell, C. A., Leshner, R. L., Metcalf, A. R., Shirley, T., and Brune, W. H.: Intercomparison of peroxy radical measurements at a rural site using laser-induced fluorescence and peroxy radical chemical ionization mass spectrometer (PerCIMS) techniques, *J. Geophys. Res.*, 108, 4605, doi:10.1029/2003JD003644, 2003.

- Rohrer, F., Bohn, B., Brauers, T., Brüning, D., Johnen, F.-J., Wahner, A., and Kleffmann, J.: Characterisation of the photolytic HONO-source in the atmosphere simulation chamber SAPHIR, *Atmos. Chem. Phys.*, 5, 2189–2201, 2005, <http://www.atmos-chem-phys.net/5/2189/2005/>.
- Sander, S. P., Friedl, R. R., Golden, D. M., Kurylo, M. J., Moortgart, G. K., Keller-Rudek, H., Wine, P. H., Ravishankara, A. R., Kolb, C. E., Molina, M. J., Finlayson-Pitts, B. J., Huie, R. E., and Orkin, V. L.: Chemical kinetics and photochemical data for use in atmospheric studies, NASA JPL Publication, 06-2, 2006.
- Schlosser, E., Bohn, B., Brauers, T., Dorn, H., Fuchs, H., Häsel, R., Hofzumahaus, A., Holland, F., Rohrer, F., Rupp, L. O., Siese, M., Tillmann, R., and Wahner, A.: Intercomparison of two hydroxyl radical measurement techniques at the atmosphere simulation chamber SAPHIR, *J. Atmos. Chem.*, 56, 187–205, 2007.
- Schultz, M., Heitlinger, M., Mihelcic, D., and Volz-Thomas, A.: Calibration source for peroxy radicals with built-in actinometry using H₂O and O₂ photolysis at 185 nm, *J. Geophys. Res.*, 100, 18811–18816, 1995.
- Stevens, P. S., Mather, J. H., and Brune, W. H.: Measurement of tropospheric OH and HO₂ by laser-induced fluorescence at low pressure, *J. Geophys. Res.*, 99, 3543–3557, 1994.
- Wegener, R., Brauers, T., Koppmann, R., Rodríguez Bares, S., Rohrer, F., Tillmann, R., Wahner, A., Hansel, A., and Wisthaler, A.: Simulation chamber investigation of the reactions of ozone with short-chained alkenes, *J. Geophys. Res.*, 112, D13301, doi:10.1029/2006JD007531, 2007.