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Interactive Comment

Interactive comment on "Water vapor total column measurements using the Elodie Archive at Observatoire de Haute Provence from 1994 to 2004" by A. Sarkissian and J. Slusser

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1) Trend Calculation

1.a) Section 5 Chapter 5 p1083 line 11Sentence A negative trend of $-0.44 pm 0.24 \pm 10^{22}\cmolecule \times cm^{-2}\ per 10 years is obtained after removing the sine fit.$

is replaced by

A negative slope of $-0.44 pm 0.24 \times 10^{22} om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope of -0.44 pm 0.24 om characteristic slope of -0.44 pm 0.24 times 10^{22} om characteristic slope om characteristic slope of -0.44 pm 0.24 times 10^{22} om charact$



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it is larger than two times its error. This is in agreement with the review of measured trends over Europe made by \cite{morland}. This negative slope might be introduced by the high summer values in 1994 and 1995. We believe that more detailled trend calculation, including seasonal trends calculations, could be made after we extend this analysis to the full Elodie Archive and to the Sophie Archive.

Also in caption of fig 7:

A negative trend of $-0.44 pm 0.24 \times 10^{22} \$ molecule $\ cm^{-2}\$ per 10 years is obtained after removing the sine fit.

is replaced by

A negative slope of $-0.44 pm 0.24 \times 10^{22} \$ molecule $\ \$ per 10 years is obtained after removing the sine fit.

1.b) Validation

p1084 line 5

The calibration of the lidar gives by few percent better agreement with our measurements than with radiosondes, but we need more co-located and simultaneous observations for validation.

is replaced by

The calibration procedure uses the variance between radiosonde profile and lidar profiles between 2 and 8 km calibrated with our total column values on available days of lidar observation. This variance varies from 0.4 at 2 km to nearly 1 at 5 km, decreasing to 0.6 at 8 km is due to the natural time variability of atmospheric water vapor. The decrease of this variance by few percent observed when using Elodie water vapor data gives information on the validity of our measurements, but we need more co-located and simultaneous observations for better validation.

2) Minor suggestions, all agreed

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p. 1077 I14 its involvement in stratospheric ozone equilibrium in the troposphere

is replaced by

its involvement in stratospheric ozone equilibrium due to the contribution of the troposphere-stratosphere exchanges

p.1081 |15

The differential optical absorption spectroscopy (DOAS) method is used commonly because using scattered light increases the path length up to 20 air masses.

is replaced by

The differential optical absorption spectroscopy (DOAS) method is used commonly for tiny absorptions and because of geometrical considerations, using scattered light (zenith sky light at twilight) increases the path length up to 20 air masses.

p.1083 |18

Secondly, saturation effects on strong absorption lines in the \$5 \nu\$ range are not negligible

is replaced by

Secondly, saturation effects on strong absorption lines in the polyad \$5 \nu\$ range at 592 nm are not negligible

Interactive comment on Atmos. Meas. Tech. Discuss., 2, 1075, 2009.

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