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Laboratory evaluation of the effect of nitric acid uptake on frost point hygrometer performance

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Paper

Discussion

Discussion Paper

Discussion

Printer-friendly Version

Interactive Discussion



AMTD

Effect of HNO₃ uptake on frost point hygrometry

3, 3725-3745, 2010

T. Thornberry et al.

Introduction **Abstract**

Title Page

Conclusions References

> Tables **Figures**

Back Close

Full Screen / Esc

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Chilled mirror hygrometers (CMH) are widely used to measure water vapour in the troposphere and lower stratosphere from balloon-borne sondes. Systematic discrepancies among in situ water vapour instruments have been observed at low water vapour mixing ratios (<5 ppm) in the upper troposphere and lower stratosphere (UT/LS). Understanding the source of the measurement discrepancies is important for a more accurate and reliable determination of water vapour abundance in this region. We have conducted a laboratory study to investigate the potential interference of gas-phase nitric acid (HNO₃) with the measurement of frost point temperature, and consequently the water vapour mixing ratio, determined by CMH under conditions representative of operation in the UT/LS. No detectable interference in the measured frost point temperature was found for HNO₃ mixing ratios of up to 4 ppb for exposure times up to 150 min. HNO₃ was observed to co-condense on the mirror frost, with the adsorbed mass increasing linearly with time at constant exposure levels. Over the duration of a typical balloon sonde ascent (90-120 min), the maximum accumulated HNO₃ amounts were comparable to monolayer coverage of the geometric mirror surface area, which corresponds to only a small fraction of the actual frost layer surface area. This small amount of co-condensed HNO3 is consistent with the observed lack of HNO3 interference in the frost point measurement because the CMH utilizes significant reductions (>10%) in surface reflectivity by the condensate to determine H_2O .

1 Introduction

Water vapour is the most important greenhouse gas in the atmosphere. Changes in its atmospheric abundance are a principal feedback mechanism in the radiative forcing of climate (Held and Soden, 2000; Soden and Held, 2006; Trenberth et al., 2007; Dessler et al., 2008). Knowledge of the distribution of water vapour and how it is changing as the earth system changes is especially important in the upper troposphere

AMTD

Discussion Paper

Discussion Paper

Discussion Paper

Discussion Paper

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ◀ ▶I

Full Screen / Esc

Back

Printer-friendly Version

Close

Interactive Discussion



3726

Discussion Paper

Discussion Paper

Back Full Screen / Esc

Printer-friendly Version

Interactive Discussion



and lower stratosphere (UT/LS) where water vapour plays critical roles in atmospheric radiative balance, cirrus cloud formation, and photochemical processes (Forster and Shine, 2002; Gettelman et al., 2004; Solomon et al., 2010). Dehydration processes reduce water vapour mixing ratios to part per million (ppm) values in the tropical UT/LS. The microphysics related to dehydration and cirrus cloud particle nucleation and growth are poorly understood at present (Jensen et al., 2005; Peter et al., 2006), limiting the ability to accurately model the dehydration process and, hence, our ability to estimate the influence of climate change on the UT/LS water vapour distribution. Part of this limitation derives from the uncertainties in available water vapour measurements in the 1 to 10 ppm range typical of this region of the atmosphere. Over the past several decades, there has been considerable disagreement (~25-100%) among in situ water vapour measurements by different instruments made at low mixing ratios in the UT/LS. The reasons for the observed differences in measured water vapour in the UT/LS are actively under investigation. A recent, chamber-based intercomparison of the principal in situ UT/LS water vapour instruments (AguaVIT, https://aquavit.icg.kfa-juelich.de/AquaVit/AquaVitWiki) revealed far less disagreement, indicating the differences observed in the atmosphere are likely related to the use of these instruments on moving platforms or interferences from other atmospheric trace species.

One of the principal instruments currently used to measure atmospheric water vapour is the chilled mirror hygrometer (Oltmans et al., 2000a). A 30-yr record of balloon-borne CMH measurements provides important constraints on the evolving trend of stratospheric water vapour abundances, which have an important role in climate change (Oltmans et al., 2000b; Rosenlof et al., 2001; Solomon et al., 2010). Observed decadal differences in the trend analysis of less than 1 ppm require high confidence in the calibration stability and the absence of measurement interferences in the CMH technique. The CMH technique utilizes the principle of condensation of water vapour as a liquid or solid on a chilled mirror. The temperature of the mirror is controlled in order to maintain a constant optical scattering from a thin dew (in the

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page Introduction **Abstract** Conclusions References **Tables Figures** Close

Discussion Pape

lower troposphere) or frost (above the lower troposphere) layer in contact with water vapour in an ambient sample flowing continuously past the mirror. An optical detector measures the specular reflectivity of the condensate-covered mirror, while a feedback controller regulates the mirror temperature such that the measured reflectivity of the mirror remains constant. Under this condition, the condensate layer on the mirror is maintained in approximate thermal equilibrium with the water vapour in the air passing over the mirror. The mirror temperature is then assumed equal to the ambient dew point or frost point temperature. The water vapour partial pressure is calculated from this measurement using a formulation of the Clausius-Clapeyron equation for the saturation pressure over water or ice, and subsequently the water vapour mixing ratio is calculated by combining with the measured ambient pressure. In this work we focus only on frost as the condensate, which is applicable to measurements at the low

temperatures in the UT/LS.

Chilled mirror hygrometry provides an absolute measurement of water vapour with an accuracy that relies only on the calibration of the temperature sensor that measures the temperature of the mirror. This is true, however, only when two conditions are valid: 1) the surface reflectivity is inversely correlated with the amount of condensate, and 2) the saturation vapour pressure of the condensate layer remains that of pure hexagonal ice. A change in the composition of the surface of the frost by incorporation of other compounds could potentially invalidate one or both conditions, which would then affect the determination of ambient water vapour. There are several trace gases in the UT/LS less abundant than water vapour that are likely to interact with the frost layer. One is HNO₃, with mixing ratios ranging from less than 0.1 parts per billion (ppb) to several ppb (Neuman et al., 2001; Popp et al., 2009). A number of laboratory studies have shown that HNO₃ adsorbs readily to ice surfaces (Zondlo et al., 1999; Abbatt, 2003; Ullerstam et al., 2005) and forms a thermodynamically stable, nitric acid hydrate condensed phase, such as nitric acid trihydrate (HNO₃·3H₂O, NAT) in the presence of water (Worsnop et al., 1993). HNO₃ has been observed in cirrus cloud particles (e.g., Popp et al., 2004; Voigt et al., 2006) and contrail particles (Schnäuble et al., 2009) and

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables

Figures

I◀











Printer-friendly Version



proposed to explain the increase in the relative humidity over ice observed in cirrus clouds (Gao et al., 2004). Szakáll et al. (2001) observed during an AIDA chamber study that the addition of nitrogen dioxide (0.01 Pa of NO $_2$, corresponding to more than 0.5 ppm) to the cold (202 K) chamber containing 5.5 ppm O $_3$ resulted in an increase in the frost point temperature measured by a chilled mirror hygrometer and, hence, an overestimate of the water vapour mixing ratio. In their study, the CMH signal was compared with two other independent water-vapour measurement techniques that provided evidence that the chamber water vapour remained constant while the CMH response changed. They hypothesized that adsorption of HNO $_3$, formed from N $_2$ O $_5$ hydrolysis on sulphuric acid aerosols present in the chamber, led to conversion of the hygrometer frost to NAT. NAT on the mirror would raise the apparent frost point because it has a condensation temperature a few degrees higher than ice (Tolbert and Middlebrook, 1990).

In the present study, we report a laboratory evaluation of the interference of ${\rm HNO_3}$ in CMH frost point measurements under pressure, flow, and ${\rm HNO_3}$ conditions representative of operation in the UT/LS.

2 Experimental

2.1 Chilled mirror hygrometer

The chilled mirror hygrometers evaluated in this study are a reconfiguration of the Cryogenic Frostpoint Hygrometer (CFH) sonde instrument. The CFH is a small (400 g) balloon-borne instrument that has been used extensively to measure water vapour profiles between the surface and the middle stratosphere at a number of locations around the globe. The stated accuracy in the determination of the frost point temperature is $\pm 0.5\,^{\circ}$ C (Vömel et al., 2007), which corresponds to approximately $\pm 0.2\,$ ppm under UT/LS conditions. The CFH units used in this laboratory evaluation (Fig. 1) were copies of the modified version previously employed in the AquaVIT chamber study.

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

l4 ►FI

■ Back Close

Full Screen / Esc

Printer-friendly Version



3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.



The modified housing is made of stainless steel and has a similar geometry to that of the balloon-borne instrument, but is hermetically sealed to allow sample flow at UT/LS pressures (<200 hPa). The optical path of the laboratory instrument has the light emitting diode (LED) source and infrared (IR) photodiode detector arranged at smaller (~40°) angles to the mirror surface. This arrangement allows for imaging of the mirror surface with a microscope camera through a sapphire window mounted directly opposite the mirror. The LED, IR photodiode detector, and electronics were used unmodified from a standard CFH instrument.

The dimensions of the mirror, mirror stem and heater coil are identical to the balloon instrument as is the location of the mirror thermistor. However, a mirror mount was made of thin-walled (0.127 mm) stainless-steel tube 1.6 cm long, which was glued to the edge of the mirror disk and acted both as a seal and thermal separation from the housing, since the body of the housing remained at room temperature. The low thermal conductivity of the tube assures that the thermal properties of the mirror assembly are similar to those of the balloon instrument. A PCTFE (Kel-F) sleeve was placed around the mirror mount to insulate it from the sample flow. The cold finger, the part of the mirror assembly that is immersed in the cryogen to provide mirror cooling is longer than that of the balloon instrument, since the fittings for connection to the flow system necessitated a larger separation between the cryogen reservoir and the sensor housing. The cooling rate of the mirror, which is largely determined by the dimensions of the thin mirror stem, remained unchanged. All key operational parameters of the laboratory version, including proportional, integral, and differential values of the temperature control loop, are identical to those in the balloon instrument. The laboratory CFH was operated using liquid nitrogen as the cryogen instead of trifluoromethane (CHF₃) as used with the balloon version. A Teflon sleeve and foam collar were placed over the cold finger to reduce the heat transfer rate into the cryogen. This insulation largely compensated for the colder temperature of the liquid nitrogen as indicated by mirror heater currents and duty cycles that are comparable to those observed in flight instruments.

2.2 HNO₃ measurement

The $\mathrm{NO_y}$ instrument used here to measure $\mathrm{HNO_3}$ is described in detail elsewhere (Fahey et al., 1989). The $\mathrm{NO_y}$ instrument has been used previously in the stratosphere during several field campaigns (e.g., Fahey et al., 1989; Kawa et al., 1992). $\mathrm{HNO_3}$ and other $\mathrm{NO_y}$ compounds are measured by conversion to NO in a 300 °C gold tube catalyst in the presence of CO (Fahey et al., 1985), followed by detection of the NO using $\mathrm{NO/O_3}$ chemiluminescence. The sensitivity of the instrument was determined during each experiment by adding NO upstream of the catalytic converter from a gas standard with a known mixing ratio. Background signals in the chemiluminescence detector were subtracted prior to calculation of $\mathrm{HNO_3}$ mixing ratios. The overall uncertainty in the $\mathrm{HNO_3}$ measurement is 20%. There is no interference in the $\mathrm{NO_y}$ measurement from water vapour mixing ratios in the range used here.

2.3 Experimental setup

The experimental setup used to evaluate the potential interference of gas-phase HNO $_3$ on the performance of chilled mirror hygrometers is shown in Fig. 2. The system consists of two identical hygrometers (described above) connected in series, sources of low mixing ratios of water vapour and HNO $_3$, the NO $_y$ instrument, and the pressure control and pumping unit. The flow system manifold consisted of 1.58 cm ID stainless steel tube sections connected with ISO-KF16 fittings. A distance of 23.5 cm separated the two CFH mirrors. A mass flow controller (Tylan FC-261) was used to deliver a stable flow of synthetic air (Scott Specialty Gases, Zero Air) through the system. The water vapour content of the zero air was determined to be less than 0.3 ppm for all cylinders used. The total gas flow rate through the manifold in most experiments was ~2000 standard cm 3 min $^{-1}$ (sccm), but experiments were conducted with flows between 1200 and 6000 sccm. Pressures were measured using 100 and 1000 Torr capacitance manometers (MKS, Inc. 626B Baratron). The pressure sensor used to monitor the pressure of the gas flowing past the frost point mirrors was attached to the

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀











Full Screen / Esc

Printer-friendly Version



flow tube about 20 cm downstream of the second CFH instrument. The flow tube remained 1.58 cm inner diameter from the frost point hygrometers to the pressure sensor and no significant pressure gradient is expected under this configuration.

Two methods were used to add water vapour to the manifold sample flow. The first was zero air flow (20 sccm) past a temperature-controlled water vapour permeation tube (Kin-Tek Laboratories, HRT 010.00). The permeation source provided 2 to 4 µg min⁻¹ of water vapour for temperatures between 60 and 80 °C, respectively. When added to the main sample flow, water vapour mixing ratios were adjustable in the range of 1 to 4 ppm, depending on the sample flow rate. The second method was to add a small flow of room air into the manifold through a needle valve. This allowed higher mixing ratios of water vapour (up to 20 ppm) in the flow system. Given the water vapour content of lab air (>6000 ppm), this flow was at most a few sccm, ≪1% of the total system flow.

HNO₃ was added to the sample flow manifold using a 10 sccm flow of zero air past a HNO₃ permeation source (Kin-Tek Laboratories, SRT-002.00-2022/40) operated at either room temperature or 40 °C. HNO₃ permeation sources are known to emit a small amount of NO₂ or other NO_v species along with the HNO₃. The non-HNO₃ emission was determined by measuring the NO_{ν} signal after passing the permeation source flow through a plug of nylon wool, which completely scrubs HNO₃ from the flow. The residual (non-HNO₃) signal, which was ~5% of the NO_v signal without the nylon wool in line, was subtracted from the measured NO_v signal to determine the HNO₃ mixing ratio during experiments. The HNO₃ addition point was located between the two CFH units in order to expose only the downstream unit (Fig. 2). The sampling point for the NO_v instrument was located 15 cm downstream of the second CFH mirror. The inner surfaces of the flow tubes were coated from the HNO₃ addition point between the hygrometers to a distance past the NO_v sampling point with a thin film of FluoroPel (Cytonix Inc., PFC 504A-FS) to minimize HNO₃ adsorption on the walls. HNO₃ transmission through the flow system was determined by comparing the NO_v signal with HNO₃ added to the flow system with that observed when the HNO₃ flow was added to the NO_v instrument **AMTD**

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ✓ ▶I

Close

Full Screen / Esc

Back

Printer-friendly Version



sample flow immediately in front of the catalyst. After 10–20 min, typical transmission values were $\geq 85\%$. HNO $_3$ addition to the flow system was controlled using a vacuum line connected to the HNO $_3$ flow line immediately upstream of the manifold. When the valve in this auxiliary line was closed, flow past the HNO $_3$ permeation tube was added to the manifold flow. When the valve in this auxiliary line was opened, the HNO $_3$ flow was diverted away from the flow system to the vacuum pump along with a small amount of air from the flow system. Opening and closing this valve resulted in prompt addition or removal of HNO $_3$ in he flow system.

One operational difference between a CFH operating in the atmosphere on a balloon and in our laboratory experiments is that in the former the temperatures of the sample flow and housing walls are near ambient while in the latter they remain at room temperature. This difference is not expected to significantly affect the results of our experiments because it is the mirror surface that controls the temperature of absorbing and desorbing molecules. This result follows from a consideration that the molecular mean free path (<1 μ m) is much smaller than the thickness of the laminar-flow thermal boundary layer (~1 mm) over the mirror surface and much smaller than the characteristic mirror dimension (1 cm).

3 Results and discussion

3.1 Frost point interference

HNO $_3$ exposure tests were conducted first by allowing both hygrometers to achieve a stable frost layer at low water vapour conditions (1–5 ppm) in the absence of HNO $_3$ and determining the NO $_y$ instrument background. HNO $_3$ was then added to the manifold and the behaviour of the two CFH instruments was compared to evaluate the possible interference. A typical experimental run lasting \sim 4 h is shown in Fig. 3. In the experiment, the permeation source produced a constant background water vapour mixing ratio of 2.3 ± 0.2 ppm in the sample flow as measured consistently by both CFH

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables

l∢ ÞI

Figures

■ Back Close

Full Screen / Esc

Printer-friendly Version



Effect of HNO₃ uptake on frost point hygrometry

3, 3725-3745, 2010

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I

I

I

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

units. The absolute value determined from the CFH measurements agreed to within 10% of the value calculated from the manufacturer's stated permeation rate. The higher water vapour values shown in Fig. 3, up to 4.3 and 11 ppm, were produced by periodically opening a valve (shown in Fig. 2) to admit a small flow of air from the room. The two CFH units exhibited slightly different transient responses to the rapid increases and decreases in water vapour, but reported the same value (within 0.3 ppm) upon reaching steady state after a few minutes.

Just before 2 h in the time series in Fig. 3, the HNO₃ diversion valve was closed, thereby exposing the downstream CFH unit, #2, to HNO₃. The HNO₃ mixing ratio in the sample flow rapidly reached 3.5 ppb and then continued to increase gradually over the next 2 h. The long rise time for HNO3 is caused by the strong adsorption of HNO₃ to all exposed surfaces, particularly those not covered by the Fluoropel coating. The variations observed in the HNO₃ mixing ratio are due to slight changes in HNO₃ transmission through the flow system that arise from changes in surface wetness and mirror temperature, and mirror frost evaporation and condensation related to rapid changes in water vapour. The frost point temperature reported by CFH#2 never deviated significantly from that of CFH#1, even after a 2.5-h exposure of up to 4 ppb of HNO₃. Furthermore, the transient and steady state responses of CFH#2 to the periods of increased water vapour were unchanged following HNO₃ exposure. A typical CFH balloon flight to 30 km altitude typically requires less than 2 h and HNO₃ is typically <0.5 ppb in the middle and upper troposphere and gradually increases to 4 ppb within a few kilometres above the tropopause (Popp et al., 2009). Thus, based on several experimental runs similar to that shown in Fig. 3, over a range of UT/LS conditions, there is no interference (<0.1 °C frost point change) due to ambient HNO₃ in CMH balloon sonde measurements.

3.2 HNO₃ surface coverage

 ${\rm HNO_3}$ is known to deposit readily to ice surfaces at temperatures similar to those of the CFH mirrors in these experiments in both the laboratory (Abbatt, 2003) and

Back Full Screen / Esc

Printer-friendly Version

Close

Interactive Discussion



atmosphere (Popp et al., 2004; Voigt et al., 2006). The lack of any observed interference of HNO₃ on the measured frost point temperature raises the question of the amount of HNO₃ that accumulates on the frost during exposure. By rapidly heating the CFH mirror and vaporizing the frost layer after a certain HNO₃ exposure time, any adsorbed or co-condensed HNO₃ on the mirror is released into the sample flow and detected by the NO_v instrument. Separate tests showed that mirror heating did not desorb appreciable HNO₃ from surfaces of the insulated mirror mount (Fig. 1). A similar rapid mirror evaporation is typically programmed to occur during balloon ascent of the CFH instrument to ensure the condensate layer is frost and not supercooled liquid (Vömel et al., 2007). The resulting pulses of HNO₃ measured by the NO_v instrument can be clearly seen in the plot in Fig. 4. Examination of the NO_v trace in Fig. 4 reveals that the magnitude of the desorbed HNO₃ pulses is a function of the exposure time. The mass of adsorbed HNO₃ was determined by integrating the pulse measured by the NO_v instrument. Figure 5 shows the measured mass of desorbed HNO₃ plotted against the exposure expressed as the product of the HNO₃ mixing ratio and exposure time. HNO₃ accumulated approximately linearly with exposure at mixing ratios of up to 4 ppb over time periods up to 3 h. Laboratory studies have shown HNO₃ monolayer densities to be approximately 1–2×10¹⁴ molecules cm⁻² (Abbatt, 2003), so the highest observed HNO₃ mass loadings (~8 ng) in our experiments represent roughly monolayer coverage of the geometric surface area of the mirror (0.37 cm²). There is no information on the specific surface area (cm² cm⁻²) of the frost, but based on electron micrographs recently published by Leu and Keyser (2009) of frost layers formed by water vapour deposition at similar temperatures, it is likely to be ≫1. This implies that only a small fraction of the surface area of the frost is affected by HNO₃ adsorption and, therefore, no significant interference in the CFH measurement would be expected.

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

H₂O adsorption and desorption on the frost layer are close to steady state when the water vapour concentration is constant above the mirror surface and the CFH mirror temperature is at a nearly constant value; under these conditions the frost layer is neither growing nor shrinking significantly. In contrast, HNO₃ adsorption and desorption are clearly not in steady state, as evidenced by the time dependence of HNO3 mass loading shown in Fig. 5. The thermodynamic state of the adsorbed HNO₃ is not known for the frost layer in our experiments. However, the mirror temperature and H₂O and HNO₃ partial pressures present in these experiments fall thermodynamically within the region of stable NAT formation (Hanson and Mauersberger, 1988; Tolbert and Middlebrook, 1990; Worsnop et al., 1993; Zondlo et al., 1999). The estimated small fractional HNO₃ coverage of the frost surface area suggests that while NAT may exist locally on the frost surface or within the frost layer, the total HNO₃/H₂O molar ratio on the mirror is much less than 0.33, which would be required for complete conversion of the frost to NAT. The observed linear increase in HNO₃ adsorbed mass in the first hours of exposure indicates that the molar ratio may continue to increase with time until it reaches 0.33 in the frost or perhaps in a layer at the frost surface, which would bring HNO₃ adsorption and desorption into steady state. If the NAT formation on the surface of the frost were to inhibit further formation of water ice, the CFH would then attempt to control at the NAT frost temperature, which is slightly higher than the frost temperature under pure H₂O conditions (Tolbert and Middlebrook, 1990) and the mirror temperature would then lead to an overestimate of the water vapour partial pressure. However, due to the low mixing ratio of HNO₃ in our case, the adsorption rate of NAT is likely to be too small to allow NAT to effectively control the mirror reflectivity. The ppb-min exposure of HNO₃ required to reach NAT stoichiometry is much larger than that realized in this laboratory evaluation and, hence, is much larger than is relevant for CFH balloon-borne measurements. The interpretation of the interference in the reading of a commercial

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

Discussion Paper

Discussion Paper

Discussion Paper

▼Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



3736

frost point instrument observed in the AIDA chamber was that HNO₃ had accumulated

Discussion Paper

in sufficient amounts to convert the frost to NAT (Szakáll et al., 2001). The results presented here cannot confirm or disprove their interpretation since their experiment was conducted with much higher HNO₃ levels.

4 Conclusions

Two CFH frost point hygrometers and an NO_v instrument were used to investigate the potential interference of gas-phase nitric acid (HNO3) at atmospherically-relevant mixing ratios in measurement of low ppm water vapour mixing ratios by chilled mirror hygrometers under pressure and flow conditions representative of operation in the UT/LS. No detectable change (<±0.1 °C) in the measured CFH frost point temperature was found with additions of up to 4 ppb HNO₃ for periods exceeding 2 h. HNO₃ that had adsorbed or co-condensed on the mirror with H2O was detected during rapid frost layer evaporation using the NO_v instrument. Our measured HNO₃ loadings of the CFH mirror surface indicate that over a typical balloon-borne CFH ascent period (90-120 min), condensed HNO₃ mass increased linearly at a rate of 0.033 ng min⁻¹ ppb⁻¹. The observed mass loading of ~8 ng following 120 min exposure at 2 ppb HNO₃ corresponds to approximately monolayer coverage of the geometric mirror surface area, but likely much less than monolayer coverage of the frost surface area. We conclude that for typical atmospheric conditions, it is unlikely that HNO3 alone could cause a significant error in water vapour measured by a balloon-sonde chilled-mirror hygrometer such as the CFH. The potential for HNO₃ to interfere with CMH water vapour measurement during extended sampling in the lower stratosphere from aircraft or other balloon platforms is likely small, but would require further study to critically assess.

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AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

1 1

Back Close

Full Screen / Esc

Printer-friendly Version



Discussion Paper

AMTD

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

- Title Page Introduction **Abstract** Conclusions References **Tables Figures**
- Discussion Paper Full Screen / Esc

Back

Printer-friendly Version

Close

- Interactive Discussion

Abbatt, J. P. D.: Interactions of atmospheric trace gases with ice surfaces: adsorption and

References

reaction, Chem. Rev., 103, 4783-4800, 2003.

- Fahey, D. W., Kelly, K. K., Ferry, G. V., Poole, L. R., Wilson, J. C., Murphy, D. M., Loewenstein, M., and Chan, K. R.: In situ measurements of total reactive nitrogen, total water, and aerosol in a polar stratospheric cloud in the Antarctic, J. Geophys. Res., 94, 11299-11315, 1989.
- Gao, R. S., Popp, P. J., Fahey, D. W., Marcy, T. P., Herman, R. L., Weinstock, E. M., Baumgardner, D. G., Garrett, T. J., Rosenlof, K. H., Thompson, T. L., Bui, P. T., Ridley, B. A., Wofsy, S. C., Toon, O. B., Tolbert, M. A., Kärcher, B., Peter, T., Hudson, P. K., Weinheimer, A. J., and Heymsfield, A. J.: Evidence that nitric acid increases relative humidity in low-temperature cirrus clouds, Science, 303, 516-520, 2004.
- Hanson, D. R., and Mauersberger, K.: Laboratory studies of the nitric acid trihydrate: Implications for the south polar stratosphere, Geophys. Res. Lett., 15, 855 (1988).
- Held, I. M. and Soden, B. J.: Water vapour feedback and global warming, Annu. Rev. Energ. Env., 25, 441–475, 2000.
 - Jensen, E. J., Smith, J. B., Pfister, L., Pittman, J. V., Weinstock, E. M., Sayres, D. S., Herman, R. L., Troy, R. F., Rosenlof, K., Thompson, T. L., Fridlind, A. M., Hudson, P. K., Cziczo, D. J., Heymsfield, A. J., Schmitt, C., and Wilson, J. C.: Ice supersaturations exceeding 100% at the cold tropical tropopause: implications for cirrus formation and dehydration, Atmos. Chem. Phys., 5, 851-862, doi:10.5194/acp-5-851-2005, 2005.
 - Kawa, S. R., Fahey, D. W., Kelly, K. K., Dye, J. E., Baumgardner, D., Gandrud, B. W., Loewenstein, M., Ferry, G. V., Chan, K. R.: The Arctic polar stratospheric cloud aerosol - aircraft measurements of reactive nitrogen, total water, and particles, J. Geophys. Res., 97, 7925-7938, 1992.
 - Neuman, J. A., Gao, R. S., Fahey, D. W., Holecek, J. C., Ridley, B. A., Walega, J. G., Grahek, F. E., Richard, E. C., McElroy, C. T., Thompson, T. L., Elkins, J. W., Moore, F. L., and Ray, E. A.: In situ measurements of HNO₃, NO₄, NO₇, NO, and O₃ in the lower stratosphere and upper troposphere, Atmos. Environ., 35, 5789-5797, 2001.
- 30 Oltmans S., K. Rosenlof, H. A. Michelsen, et al.: Chapter 2: Data Quality, in: SPARC Assessment of Upper Tropospheric and Stratospheric Water Vapour, edited by: Kley, D., Russell, III, J. M., and Phillips, C., 2000a.

- AMTD
 - 3, 3725-3745, 2010
- Effect of HNO₃ uptake on frost point hygrometry
 - T. Thornberry et al.
- Title Page

 Abstract Introduction

 Conclusions References

 Tables Figures
 - Full Screen / Esc

Back

Printer-friendly Version

Close

- Interactive Discussion
 - © O

- Oltmans, S. J., Vömel, H., Hofmann, D. J., Rosenlof, K. H., and Kley, D.: The increase in stratospheric water vapor from balloonborne, frostpoint hygrometer measurements at Washington DC, and Boulder, Colorado, Geophys. Res. Lett., 27, 3453–3456, 2000b.
- Peter, T., Marcolli, C., Spichtinger, P., Corti, T., Baker, M. B., and Koop, T.: When dry air is too humid, Science, 314, 1399–1402, 2006.
- Popp, P. J., Marcy, T. P., Gao, R. S., Watts, L. A., Fahey, D. W., Richard, E. C., Oltmans, S. J., Santee, M. L., Livesey, N. J., Froidevaux, L., Sen, B., Toon, G. C., Walker, K. A., Boone, C. D., and Bernath, P. F.: Stratospheric correlation between nitric acid and ozone, J. Geophys. Res., 114, D03305, doi:10.1029/2008JD010875, 2009.
- Popp, P. J., Gao, R. S., Marcy, T. P., Fahey, D. W., Hudson, P. K., Thompson, T. L., Kärcher, B., Ridley, B. A., Weinheimer, A. J., Knapp, D. J., Montzka, D. D., Baumgardner, D., Garrett, T. J., Weinstock, E. M., Smith, J. B., Sayres, D. S., Pittman, J. V., Dhaniyala, S., Bui, T. P., and Mahoney, M. J.: Nitric acid uptake on subtropical cirrus cloud particles, J. Geophys. Res., 109 (D06302), doi:10.1029/2003JD004255, 2004.
- Rosenlof, K. H., Oltmans, S. J., Kley, D., Russel III, J. M., Chiou, E.-W., Chu, W. P., Johnson, D. G., Kelly, K. K., Michelsen, H. A., Nedoluha, G. E., Remsberg, E. E., Toon, G. C., and McCormick, M. P.: Stratospheric water vapor increases over the past helf-century, Geophys. Res. Lett., 28, 1195–1198, 2001.
 - Schnäuble, D., Voigt, C., Kärcher, B., Stock, P., Schlager, H., Krämer, M., Schiller, C., Bauer, R., Spelten, N., de Reus, M., Szakáll, M., Borrmann, S., Weers, U., and Peter, Th.: Airborne measurements of the nitric acid partitioning in persistent contrails, Atmos. Chem. Phys., 9, 8189–8197, doi:10.5194/acp-9-8189-2009, 2009.

20

- Soden, B. J. and Held, I. M.: An assessment of climate feedbacks in coupled ocean-atmosphere models, J. Climate, 19, 3354–3360, 2006.
- Szakáll, M., Bozoki, Z., Kraemer, M., Spelten, N., Moehler, O., and Schurath, U.: Evaluation of a photoacoustic detector for water vapor measurements under simulated tropospheric/lower stratospheric conditions, Environ. Sci. Technol., 35, 4881–4885, 2001.
 - Tolbert, M. A. and Middlebrook, A. M.: Fourier transform infrared studies of model polar stratospheric cloud surfaces' growth and evaporation of ice and nitric acid/ice, J. Geophys. Res., 95, 22423–22431, 1990.
 - Trenberth, K. E., Jones, P. D., Ambenje, P., Bojariu, R., Easterling, D., Klein Tank, A., Parker, D., Rahimzadeh, F., Renwick, J. A., Rusticucci, M., Soden, B., and Zhai, P.: Observations: Surface and Atmospheric Climate Change, in: Climate Change 2007: The Physical Science

- Basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change, edited by: Solomon, S., Qin, D., Manning, M., Chen, Z., Marquis, M., Averyt, K. B., Tignor, M., and Miller, H. L., Cambridge University Press, Cambridge, UK and New York, NY, USA, 2007.
- Ullerstam, M., Thornberry, T., and Abbatt, J. P. D.: Uptake of gas-phase nitric acid to ice at low partial pressures: evidence for unsaturated surface coverage, Faraday Discuss., 130, 211–226, 2005.
 - Voigt, C., Schlager, H., Ziereis, H., Kärcher, B., Luo, B. P., Schiller, C., Krämer, M., Popp, P. J., Irie, H., and Kondo, Y., Nitric acid in cirrus clouds, Geophys. Res. Lett., 33, doi:10.1029/2005GL025159, 2006.
 - Vömel, H., David, D. E., and Smith, K.: Accuracy of tropospheric and stratospheric water vapour measurements by the cryogenic frost point hygrometer: Instrumental details and observations, J. Geophys. Res., 112, D08305, doi:10.1029/2006JD007224, 2007.
 - Worsnop, D. R., Fox, L. E., Zahniser, M. S., and Wofsy, S. C.: Vapor pressures of solid hydrates of nitric acid: implications for polar stratospheric clouds. Science, 259, 71–74, 1993.
 - Zondlo, M. A., Barone, S. B., and Tolbert, M. A.: Uptake of HNO₃ on ice under upper tropospheric conditions, Geophys. Res. Lett., 24, 1391–1394, 1997.

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ← ▶I

Back

Full Screen / Esc

Close

Printer-friendly Version



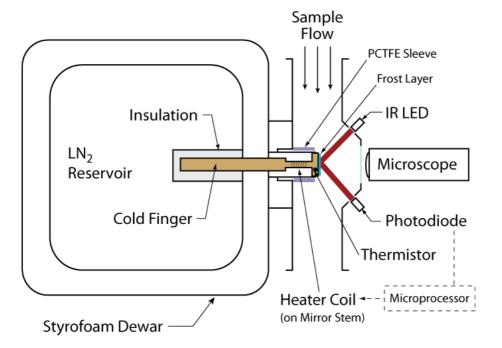


Fig. 1. Schematic of the laboratory CFH instruments used for the HNO₃ experiments. Drawing not to scale.

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures













Full Screen / Esc

Printer-friendly Version



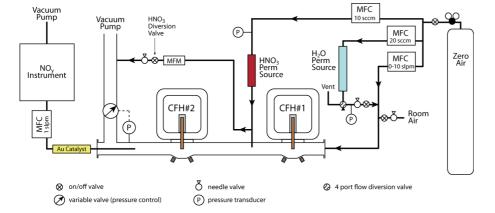


Fig. 2. Experimental set-up for the HNO_3 -frost point tests showing the arrangement in the manifold of the hygrometer mirrors, HNO_3 addition point, and NO_y sampling inlet. MFC is a mass flow controller operating at the specified flow or within the given flow range; MFM is a mass flow meter. Drawing not to scale.

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

l**∢** ▶l

Back Close

Full Screen / Esc

Printer-friendly Version



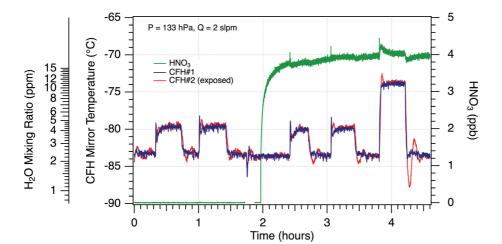


Fig. 3. Time series showing the reported frost point temperatures from both CFH instruments prior and subsequent to the addition of HNO_3 to the sample flow and the HNO_3 mixing ratio in the flow system as measured with the NO_y detector. Increases in observed frost point temperature (water vapour mixing ratios) were due to addition of small flows of room air to the sample flow. Frost point temperatures measured by the exposed CFH before and after HNO_3 addition agreed to within $0.1\,^{\circ}C$.

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I ◀ ▶I

Back

Full Screen / Esc

Printer-friendly Version

Close

Interactive Discussion



3743

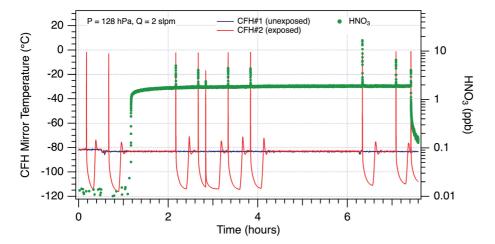


Fig. 4. Time series of CFH mirror temperatures and HNO_3 mixing ratios. The stable mirror temperature corresponds to a water vapour mixing ratio of 2.5 ppm. The pulses visible in the CFH#2 mirror temperature are rapid increases used to evaporate the frost layer. Corresponding pulses were observed in HNO_3 as measured by the NO_y instrument when the frost layer had been exposed to HNO_3 . The negative excursions in the CFH#2 temperature occur when the frost layer is being reformed following removal.

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

I

▶I

Back Close

Full Screen / Esc

Printer-friendly Version



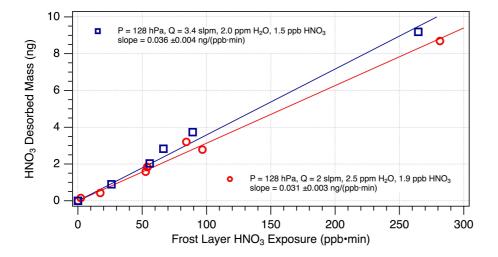


Fig. 5. Mass of HNO_3 measured during mirror frost layer evaporation events plotted against the frost layer exposure to HNO_3 expressed as the time integrated mixing ratio. The legend shows the experimental conditions of pressure, sample flow (Q), and water vapour and HNO_3 mixing ratios, along with the derived growth rate from the best fit line with a zero intercept.

3, 3725-3745, 2010

Effect of HNO₃ uptake on frost point hygrometry

T. Thornberry et al.

Title Page

Abstract Introduction

Conclusions References

Figures

Tables



Back Close

Full Screen / Esc

Printer-friendly Version

