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## Novel SO<sub>2</sub> spectral evaluation scheme

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# Novel SO<sub>2</sub> spectral evaluation scheme using the 360–390 nm wavelength range

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## Abstract

Differential Optical Absorption Spectroscopy (DOAS) is a well established spectroscopic method to determine trace gases in the atmosphere. During the last decade, passive DOAS, which uses solar radiation scattered in the atmosphere as a light source, has become a standard tool to determine SO<sub>2</sub> column densities and emission fluxes from volcanoes and other large sources by ground based as well as satellite measurements. For the determination of SO<sub>2</sub> column densities, the structured absorption of the molecule in the 300–330 nm region (due to the A<sup>1</sup>B<sub>1</sub>←X<sup>1</sup>A<sub>1</sub> transition) is used. However, there are several problems limiting the accuracy of the technique in this particular application. Here we propose to use an alternative wavelength region (360–390 nm) due to the spin-forbidden a<sup>3</sup>B<sub>2</sub>←X<sup>1</sup>A<sub>1</sub> transition for the DOAS evaluation of SO<sub>2</sub> in conditions where high SO<sub>2</sub> column densities prevail. We show this range to have considerable advantages in such cases, in particular when the particle content of the plume is high and when measurements are performed at large distances from the area of interest.

## 1 Introduction

Developed since the mid-seventies (Noxon, 1975; Perner et al., 1976; Platt et al., 1979), today Differential Optical Absorption Spectroscopy (DOAS) is a well established spectroscopic method to determine trace gases in the atmosphere. DOAS is typically applied in the UV-visible wavelength region and uses the narrow band absorption features of molecules for their identification; a detailed description can e.g. be found in Platt and Stutz (2008). During the last decade, the passive DOAS technique, which uses solar radiation scattered in the atmosphere as a light source (see e.g. Hönninger et al., 2004a; Platt and Stutz, 2008), has become a standard tool to determine SO<sub>2</sub> emission fluxes from volcanoes world wide (e.g. Galle et al., 2003, 2010; Horton et al., 2006; McGonigle, 2007) and has started to replace the Correlation Spectroscopy

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(COSPEC) developed in the 1960ties, (e.g. Millan et al., 1969). Moreover, DOAS offers the possibility to analyse for other volcanic gases, like BrO (see e.g. Bobrowski et al., 2003), as well. This was possible due to the technological development of cheaper and smaller spectrometers, sensitive and fast multichannel array detectors, powerful computers, and efficient algorithms for the analysis of differential absorption spectra.

The DOAS method is usually applied to weak absorbers with optical densities below approximately 0.05. However, when measuring SO<sub>2</sub> absorption in volcanic plumes and considering the commonly used evaluation range between 300 and 325 nm, optical densities can exceed unity. Such strong absorbers can complicate the DOAS retrieval algorithm when using high resolution reference spectra from literature, which have to be convoluted with the instrumental line-shape of the employed spectrometer. This well known phenomenon (see e.g. Platt and Stutz, 2008) is caused by a non-linearity in the relationship between the measured optical density and the column density occurring at high optical densities when the spectral resolution of the instrument is insufficient to fully resolve the absorption bands, as is usually the case with the compact spectrometers employed for volcanic surveillance. The problem can be solved by using an iterative approach like the Iterative Maximum A Posteriori DOAS technique (IMAP-DOAS, see Frankenberg et al., 2005 for more detailed information), but such methods require high computation power and are rarely applied today.

A further problem associated with the commonly evaluated wavelength range below 330 nm, where the strong absorptions due to the A<sup>1</sup>B<sub>1</sub>←X<sup>1</sup>A<sub>1</sub> transition of the SO<sub>2</sub> molecule are located, is caused by the rather low intensity of scattered sunlight (mostly due to ozone absorption in the stratosphere) and therefore unfavourably high photon shot noise. Additionally for this UV wavelength region where the measured intensities are generally low, compared to the longer wavelength UV also entering the instrument, stray-light can contribute a significant fraction of the measured signal, leading to a considerable underestimation of the true optical density. This is particularly severe for the frequently applied compact spectrometers (e.g. from Ocean Optics). Correction of this effect is not at all straightforward (Van Roozendaal et al., 2006).

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Apart from the above mentioned spectroscopic issues, it has only recently been shown that the accuracy of remote sensing techniques measuring volcanic SO<sub>2</sub> in the UV-visible wavelength region using scattered solar radiation as a light source is often largely dependent on our knowledge of the radiative transfer in and around the measured volcanic plume (Kern et al., 2009). The results show that this especially applies to situations in which SO<sub>2</sub> is a strong absorber, as the optical path of incident radiation is significantly influenced by absorption in the plume in such cases. There are three problems related to radiation transport:

1) Radiation is scattered in the atmosphere both, by air molecules (mostly Rayleigh scattering) and by aerosol particles. While aerosol scattering is only weakly dependent on the wavelength  $\lambda$  (approximately proportional to  $\lambda^{-1.3}$ ), Rayleigh scattering is proportional to  $\lambda^{-4}$ . This fact leads to a wavelength dependent mean free photon path in the atmosphere. Photons of shorter wavelengths travel significantly shorter mean paths between scattering events than photons of longer wavelengths. Thus, radiation of shorter wavelengths is more likely to be scattered between instrument and plume than radiation of longer wavelengths. This so-called “radiation dilution” effect was first observed in DOAS measurements by Mori et al. (2006), who realized that the SO<sub>2</sub> column density (CD) retrieved from measured spectra depended on the wavelength range in which the evaluation took place. Especially for long distances between instrument and plume (i.e. several km), evaluations at shorter wavelengths yielded lower SO<sub>2</sub> column densities than those performed at longer wavelengths.

2) Aside from the reduction of the apparent SO<sub>2</sub> CD due to radiation dilution, multiple scattering inside the plume can also enhance the SO<sub>2</sub> CD. Depending on the particular situation (i.e. amount and properties of aerosols in the plume and in the air surrounding the plume), the enhancement of the effective photon path due to multiple scattering can be of the same order of magnitude as the above described dilution. Therefore, the measured CD can be either over- or underestimated in any particular measurement.

3) In cases of strong absorption, the observed scattered light has only penetrated the outermost layers of the SO<sub>2</sub> plume. In extreme cases, the sensitivity for the whole

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plume can become essentially zero ( $AMF \sim 0$  leading to a strong underestimation of the true CD), because at the wavelength where the SO<sub>2</sub> absorptions are observed the plume is completely dark. In reality, this effect is, however, typically reduced due to the presence of aerosols, and a small fraction of the observed light is scattered by aerosols inside the plume. Then the SO<sub>2</sub> absorption signal is still weak but not zero.

In fact, the effect that not all of the detected radiation has passed through the region of interest in a straight line is a well known issue in optical remote sensing. Typically, the concept of air mass factor (AMF) is used to describe the phenomenon. The AMF gives the ratio of the measured column density to the theoretical column density along a straight line through the region of interest (see Kern et al., 2009 for a more specific definition). Therefore, an AMF of less than unity indicates that the average light path inside the region of interest (in this case the volcanic plume) is shorter than the column along a straight path through the plume.

Here we propose a new approach to the problem using an alternative wavelength region for the DOAS evaluation of SO<sub>2</sub>, which can have considerable advantages under certain conditions, in particular when high SO<sub>2</sub> column densities prevail or when measurements at large distances from the area of interest are necessary.

## 2 The new approach

In cases of high SO<sub>2</sub> column densities, the relatively weak system of absorption lines corresponding to the spin-forbidden  $a^3B_2 \leftarrow X^1A_1$  transition of the SO<sub>2</sub> molecule can be used to quantitatively detect this species. Compared to the standard evaluation range of approximately 310 to 325 nm (due to the  $A^1B_1 \leftarrow X^1A_1$  transition), which has been used by all hitherto reported UV-spectroscopic SO<sub>2</sub> measurements, the absorption lines of this system are found in the longer UV wavelength region of about 350 to 390 nm, which can have a number of advantages. In the following we refer to the two evaluation windows as the “long wave UV” and “short wave UV” windows, respectively. The differential absorption cross section of SO<sub>2</sub> is about two to three orders of

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magnitude lower in the long wave UV window, therefore the above mentioned problems of non-linearity are avoided. While the weaker absorption cross section also leads to a lower sensitivity for SO<sub>2</sub>, this effect is in part compensated by a higher light intensity at longer wavelengths, which leads to considerably better signal/noise ratios. In the following, two examples of the SO<sub>2</sub> evaluation in the proposed long wave UV window of 360–390 nm are presented in comparison to an evaluation between 315 and 325 nm. As one example, we discuss a spectrum recorded by a ground based instrument at Mt. Etna Volcano in June 2008, as the other we report satellite measurements of the eruption of Kasatochi volcano in August 2008.

Also a theoretical model study was performed, which investigates the influence of radiation dilution on remote sensing measurements in the proposed long wave UV wavelength region around 375 nm in comparison to the standard range around 320 nm using the backward Monte-Carlo radiative transfer model McArtim (Deutschmann et al., 2008) to simulate the propagation of radiation between the sun and a simulated instrument located on the ground at various distances from the volcanic plume or of the backscattered sun-light to a satellite instrument. The results show that the transition from short wave UV to long wave UV leads to a decrease in the radiation dilution effect, and that the AMF is typically closer to unity for the new retrieval window if the distance between instrument and plume is larger than approximately 1 km. Also in the case of satellite remote sensing, the new proposed fit region prevents underestimation of SO<sub>2</sub> column densities in cases of high SO<sub>2</sub> loads.

### 3 Sample measurements and evaluations in the short wave UV and long wave UV windows

#### 3.1 Ground based measurement example

For the ground based measurement example, a Miniature MultiAxis-DOAS (Mini-MAX-DOAS) instrument (described in detail e.g. by Hönninger et al., 2004b; Bobrowski

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et al., 2007; Bobrowski and Platt, 2007) was set up on the North East crater rim of Mt. Etna volcano (Italy) on 10 June 2008. The miniaturised MAX-DOAS system consists of an airtight aluminium box containing a temperature stabilized commercial miniature spectrometer (OceanOptics Inc., USB2000) covering the spectral range from 289 nm to 442 nm, a small telescope (fused silica lens,  $f=40$  mm,  $d=20$  mm, field of view approximately  $0.6^\circ$ ) coupled to the spectrometer by a quartz fibre bundle (4 fibres with  $200\ \mu\text{m}$  core diameter), and electronics for temperature regulation and stepper motor control. A geared stepper motor is attached to one side of the box and turns it to point the telescope at different elevation angles between  $0$  and  $150^\circ$ . Automatic data acquisition was performed by the DOASIS software package (Kraus et al., 2001) running on a notebook computer. The distance to the volcanic plume was about 100 m. The plume visually had a white appearance, probably caused by the large amount of condensation of emitted water vapour. Aside from the plume, a cloudless sky was observed.

The wavelength to pixel mapping (WPM) as well as the instrumental line shape (instrument function) were determined by taking a spectrum of a low-pressure mercury emission lamp. By using the known mercury spectral line positions, a second order polynomial describing the WPM was calculated. The WPM was later adjusted using the Fraunhofer structure of the measured spectra compared to a literature solar emission spectrum (Kurucz et al., 1984).

The software WinDoas V2.10 from IASB (Belgium Institute for Space Aeronomy, Fayt and Van Roozendaal, 2001) was used to derive the slant column densities (SCD) of SO<sub>2</sub> from the recorded spectra using a standard DOAS fit. A background spectrum recorded shortly after each plume measurement was taken as a Fraunhofer reference spectrum (FRS) to remove the Fraunhofer lines. Care was taken to avoid absorption structures of volcanic species in the FRS. All measured spectra were corrected for spectrometer stray light by subtracting a constant offset in such a way as to force the intensity measured below 298 nm, where ozone absorption is sufficiently large to block all incident solar radiation, to zero. In addition, a “Ring” spectrum was calculated to

compensate for the effect of Raman scattering in the atmosphere (Bussemer, 1993).

SO<sub>2</sub> was first analyzed in the typical short wave UV region of 315 to 325 nm. Apart from the SO<sub>2</sub> reference spectrum (Bogumil et al., 2003 – 293 K), two O<sub>3</sub> reference spectra measured at 223 and 246 K (Voigt et al., 2001), a Ring spectrum (to remove artefacts caused by inelastic scattering in the atmosphere, Ring effect, see Bussemer 1993) and the Fraunhofer reference spectrum were included in the fitting process. A 3rd order polynomial was applied to remove broad-band structures, and a shift and 1st order squeeze were allowed for the measured spectra to compensate possible wavelength calibration errors.

A sample evaluation result is shown in Fig. 1a. The SO<sub>2</sub> absorption features are easily detectable, but several unexplained structures remain in the residual spectrum, with fairly high amplitudes ( $2.7 \times 10^{-2}$  RMS) and features that do not appear to be randomly distributed (as would be expected of pure noise). Also, the Ring spectrum is not very well captured by the fit and might be one of the main causes for the high residual structures. The slant column density of SO<sub>2</sub> was determined to  $(1.6 \pm 0.1) \times 10^{19}$  molec/cm<sup>2</sup>.

In a second step, SO<sub>2</sub> was evaluated in the long wave UV window between 360 and 390 nm, where 5 weak absorption bands are located. As in the short wave UV evaluation besides SO<sub>2</sub> and two O<sub>3</sub> reference spectra taken from the same source and for the same temperatures as above, an O<sub>4</sub> reference spectrum (Herman et al., 1999), NO<sub>2</sub> at 246 K (Voigt et al., 2002), a Ring spectrum, and a Fraunhofer reference spectrum was included in the DOAS fit. Again, a 3rd order polynomial was applied to remove any broad-band structures, and a shift and 1st order squeeze were allowed for the measured spectra.

The DOAS evaluation results obtained in the long wave UV window for the same spectrum as evaluated before are shown in Fig. 1b. SO<sub>2</sub> could also be clearly identified in this wavelength region, and in this case  $(3.2 \pm 0.4) \times 10^{19}$  molec/cm<sup>2</sup> were found. The higher value obtained here is likely a result of a lower radiation dilution, as discussed in the section “Model Studies” below. The amplitude of the residual structures ( $9.7 \times 10^{-4}$  RMS) is almost an order of magnitude lower than in the previous evaluation,

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and the features appear to be of random nature. The causes for the lower amplitude of the fit residual are (1) elimination of non-linear effects caused by high optical densities (see Frankenberg et al., 2005); (2) reduced influence of spectrometer stray light in this well illuminated region of the detector; (3) the absence of a wavelength dependent air mass factor (see “Model Studies”), and (4) the lower photon noise in the long wave UV window due to more incident radiation (see below).

### 3.2 Satellite measurement example

As a second example we present the evaluation of satellite data from an overpass of the GOME-2 instrument after the eruption of the Kasatochi volcano (Aleutian Islands) on the 8 August 2008. The GOME-2 (Global Ozone Monitoring Experiment-2) on board of the MetOp satellite is the first of a series of three identical instruments. MetOp-A was launched into a sun synchronous polar orbit at 800 km altitude in October 2006. The satellite crosses the equator at 09:30 LT. GOME-2 is a 4 channel UV/Vis grating spectrometer observing the Earth’s atmosphere in nadir viewing geometry (including up to  $\pm 50^\circ$  difference to nadir). It measures both the radiance component of the light reflected by the Sun-illuminated Earth’s atmosphere and the direct Sun light, covering the wavelength region of 240–790 nm at moderate spectral resolution of 0.2–0.4 nm. With a pixel size of 80×40 km (240×40 km for the back scan), GOME-2 observes about 4 times smaller ground pixels than its predecessor GOME on ERS-2. By scanning the earth surface with a scan-width of 1920 km, global coverage can be achieved within 1.5 days (EUMETSAT, 2005).

As in the first example, WinDoas V2.10 was used to derive the SO<sub>2</sub> SCD from the recorded spectra using a DOAS fit (spectral range 360–390 nm). For the Fraunhofer reference spectrum, the Sun Mean Reference Spectrum (SMR) of GOME-2 for that day was taken (containing no atmospheric absorptions) and again a “Ring” spectrum (Wagner et al., 2009) was calculated to compensate for the effect of Raman scattering.

For the short wave UV SO<sub>2</sub> evaluation of the satellite data, a slightly different fit range than for the ground-based measurements was chosen, covering the region between

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312.1–324 nm. Besides the SO<sub>2</sub> reference spectrum (Bogumil et al., 2003 – 293 K), an O<sub>3</sub> cross section (221 K, see Gür et al., 2005), the SMR, the Ring spectrum (calculated from the SMR) and an inverse SMR spectrum were included into the fitting process (the inverse SMR spectrum should compensate the effect of possible spectrograph straylight). Since the atmospheric light paths within the selected fitting window change systematically with wavelength (e.g. Van Roozendael et al., 2006), not only the original ozone absorption cross section but also a second one (the original cross section scaled with a fourth order polynomial in wavelength) was included in the fitting algorithm. It should be noted that, in principle, additional ozone absorption cross sections for other temperatures might be included into the fit. However, for the selected case study, no significant change of the retrieved SO<sub>2</sub> results and the magnitude of the residual were found if additional temperatures were considered. A 5th order polynomial was applied to remove the broad-band structures and a slight wavelength shift was allowed for the measured spectra.

In Fig. 2a, the short wave UV fit results for the evaluation of one of the recorded spectra including the Kasatochi SO<sub>2</sub> plume are shown. The SO<sub>2</sub> signal can be clearly identified with a determined slant column density of  $(2.2 \pm 0.3) \times 10^{18}$  molec/cm<sup>2</sup>. As in the ground-based example, the residual ( $5 \times 10^{-2}$  RMS) shows unexplained systematic structures that are caused by the fact that the SO<sub>2</sub> AMF depends strongly on wavelength. Thus, the SO<sub>2</sub> absorption structures can not be well described by simply including a single SO<sub>2</sub> cross section in such cases.

The long wave UV SO<sub>2</sub> evaluation of the satellite data included, besides the above mentioned reference spectra for the short wave UV range, reference spectra for O<sub>4</sub> (Burkholder et al., 1990 – 298 K) and NO<sub>2</sub> (Vandaele, 1997 – 294 K). Again a spectral shift for the measured spectra was allowed to correct for possibly occurring instrumental changes.

The results for the long wave UV evaluation can be seen in Fig. 2 (left panel), showing a significantly higher SO<sub>2</sub> slant column density of  $(4.7 \pm 0.1) \times 10^{19}$  molec/cm<sup>2</sup> and are thus exceeding the value of the short wave UV evaluation by a factor of about 20. The

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now quite random looking residual structures are up to 50 times smaller than those obtained in when the short wave fit range ( $2 \times 10^{-3}$  RMS).

The spatial distribution of the SO<sub>2</sub> plume on 8 August for the different evaluation ranges is presented in Fig. 3a–c. In Fig. 3a the plume can be clearly identified by the significantly enhanced SO<sub>2</sub> SCD values obtained in the short wave UV evaluation. For the long wave UV fit region, the plume can also be clearly identified in Fig. 3b (notice the SCD values are about a factor of 20 higher), however showing in principle the same shape. However, some interesting discrepancies remain. While the highest SO<sub>2</sub> columns are measured in the north-east and south-west regions of the plume when using the short wave UV window, the maximum SO<sub>2</sub> column is found in the south-east when using the long wave UV range. As we show in the section “Model Studies”, this is likely a result of a decreased sensitivity for high SO<sub>2</sub> columns that occurs in short wave UV when SO<sub>2</sub> absorption becomes so strong that radiation can no longer penetrate the volcanic plume (see Fig. 6). In order to illustrate the advantages of the long wave UV evaluation more clearly, the fit results for the previous used short wave region are shown in Fig. 3c on the same scale as used for the long wave results.

## 4 Discussion

The absolute SO<sub>2</sub> cross section (Fig. 4a) in the longer wavelength region is about three orders of magnitude smaller than in the commonly used shorter wavelength range. Nevertheless, SO<sub>2</sub> shows a characteristic high frequency structure in both regions, and the narrow band (differential) absorption features used for the DOAS evaluation are only about two orders of magnitude smaller in the long wave UV window (see Fig. 4b, c).

The disadvantage with regard to sensitivity and detection limit that this weaker system of absorption lines brings with it are largely compensated by a significantly better signal to noise ratio in the long wave UV region and a much lower residual (see above): The diffuse solar irradiance incident at the Earth’s surface is about one order

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of magnitude higher in the long wave UV region than in the short wave UV window. As the noise of a DOAS measurement is typically dominated by photon noise (the magnitude of which is proportional to the square root of the measured intensity), measurements in the long wave UV window provide a 3 times better signal to noise ratio from photoelectron statistics for a given measurement integration time. Another factor of two is gained by using a wider wavelength region for the long wave UV wavelength evaluation, which adds another factor of 1.4 to the signal to noise ratio. Moreover, the quantum efficiency of typical compact spectrometer detectors is better in the long wave UV and, due to reduced distortion of the band shape (see discussion above), the amplitude of the residuum is lower in the long wave UV range.

In summary, the overall loss in sensitivity when switching to the longer wavelength window (long wave UV) can be estimated at about an order of magnitude. While this may seem high, a considerable fraction of volcanic SO<sub>2</sub> emissions measurements are conducted under conditions in which the lower sensitivity is not an issue, and the potentially improved accuracy of evaluations in the long wave UV region is arguably more important.

## 5 Model studies

### 5.1 Ground based geometry

A model study was conducted to investigate the influence of radiation transport on ground-based remote sensing measurements performed in the long wave UV wavelength region around 375 nm in comparison to the standard short wave UV range around 320 nm. In this study, the backward Monte-Carlo radiative transfer model McArtim (Deutschmann et al., 2008) was used to simulate the propagation of radiation between the sun and a simulated instrument located on the ground at various distances from the centre of a volcanic plume.

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The geometry of the model run corresponds to the measurements at Mt. Etna on 10 June 2008. A volcanic plume with a 450 m diameter circular cross section was simulated 250 m above the ground, the DOAS instrument, was located 350 m from the plume center and looked towards the plume axis. While the plume contained SO<sub>2</sub> and aerosols, an aerosol-free ambient atmosphere was assumed, an assumption that is justified by the observed clear sky conditions above and around the plume on the date of the measurement (see Sect. 3.1). The model atmosphere exhibited a typical stratospheric ozone profile to account for the influence of ozone on radiative transfer in the upper atmosphere.

Using the radiative transfer model as a forward model for the measurement results, an iterative problem inversion was performed in which the SO<sub>2</sub> concentration and the aerosol extinction coefficient inside the simulated plume were varied until the modelled spectral intensity, when divided by the modelled spectral intensity in a viewing direction outside the plume and after applying the logarithm, matched the measured optical density. The details of this procedure are subject of a separate manuscript in preparation.

The best fit, both, to the measured spectrum between 306 and 324 nm as well as to the measured column density at 375 nm was achieved for a plume SO<sub>2</sub> concentration of  $4.1 \times 10^{11}$  molec/cm<sup>3</sup>, which corresponds to a straight column of  $1.85 \times 10^{19}$  molec/cm<sup>2</sup> and an aerosol extinction coefficient of approximately 80 km<sup>-1</sup>. Therefore, these values are assumed to be the best estimate for the true conditions inside the volcanic plume.

However, a standard DOAS evaluation will not necessarily give these values. As described by Kern et al. (2009), a naïve DOAS approach can yield considerable errors, especially if the measurements are conducted at a large distance from the plume or if the plume contains a high aerosol load. As an aerosol extinction coefficient of 80 km<sup>-1</sup> corresponds to an approximate visibility of only 50 m, a value that is not unrealistic considering the high opacity of the plume caused by condensation of water vapour, considerable errors can be expected in standard DOAS retrievals in this example as well. One indication of this fact is the factor of 2 difference between the result of the evaluation in the short wave UV window compared to that of the long wave UV range.

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The question which of the two evaluation ranges generally gives the better results cannot be answered without taking into account the distance to the plume. In the example measurement at Mt. Etna, the instrument was located next to the active crater and therefore the distance to the plume was exceptionally small (assumed to be 350 m in the model run). In this case, the short wave UV seems to give the better result, as the retrieved CD of  $1.6 \times 10^{19}$  molec/cm<sup>2</sup> is closer to the radiative transfer model inversion result of  $1.85 \times 10^{19}$  molec/cm<sup>2</sup> than the  $3.2 \times 10^{19}$  molec/cm<sup>2</sup> obtained in the long wave UV range. However, this situation changes rapidly if the measurements are conducted farther away from the plume. Figure 5 shows the simulated air mass factor (AMF) for a measurement of the above described plume as a function of distance between instrument and plume. Both 320 and 375 nm are shown, it should however be noted that the simulations performed at 320 nm are not necessarily indicative of the radiative transfer in the entire region between 300 and 330 nm, as the AMF is strongly dependent on wavelength in this region (see Kern et al., 2009).

The first data points shown in Fig. 5 are the results of model runs conducted for the exact geometry used for the radiative transfer inversion, i.e. the distance to the plume center was 350 m. In this geometry, the long wave UV range yields an AMF of almost 1.75, thus indicating a light path extension by 75% inside the plume due to multiple scattering. At 320 nm, the light path is also enhanced by multiple scattering, but the strong SO<sub>2</sub> absorption leads to an AMF of about 0.75. For this geometry, a standard DOAS retrieval in the short wave UV range is therefore expected to slightly underestimate the true straight CD, while a retrieval in the long wave UV will overestimate the CD by 75%.

However, as the instrument is moved farther from the plume in the simulation, the AMF rapidly decreases in both cases, an effect that is caused by radiation entering the line of sight between plume and instrument, thus diluting the signal. The AMF at 320 nm drops below 0.5 at 750 m, and continues to decrease as the distance is increased. At 375 nm, the situation is much better. Here, the AMF drops to approximately unity at 1 km distance and then declines more slowly, thus staying close to one for a large range

of distances from the plume.

The results of this model study show that the novel evaluation range is better suited for DOAS measurements of SO<sub>2</sub> in volcanic plumes with high SO<sub>2</sub> and aerosol loads when the instruments are located at distances larger than approximately 1 km from the plume. While the exact AMF of such measurements will depend on the conditions inside the plume as well as the aerosol load in the ambient atmosphere, the described example shows that the dilution effect is considerably less prominent for the long wave UV evaluation region than for wavelengths around 315 nm.

## 5.2 Satellite geometry

The GOME-2 observations of the Kasatochi eruption were also simulated using the full spherical Monte Carlo radiative transfer model McArtim (Deutschmann et al., 2008). To simplify the simulations, fixed values for the viewing geometry were assumed (SZA: 50°, viewing angle: -90°). The layer of enhanced SO<sub>2</sub> concentrations was assumed to be between 10 and 11 km (constant SO<sub>2</sub> concentration within this layer). Below the SO<sub>2</sub> layer, a cloud layer extending from 5 to 6 km (OD: 20, single scattering albedo: 1, asymmetry parameter: 0.85) was included in the simulations. Two basic model runs were performed: the first one without aerosols and the second including aerosols within the SO<sub>2</sub> layer (aerosol optical depth: 3, asymmetry parameter: 0.68, single scattering albedo: 0.9). For both runs the SO<sub>2</sub> concentration was varied from zero to 5 × 10<sup>14</sup> molec/cm<sup>3</sup>, and the corresponding SO<sub>2</sub> SCDs were calculated for both of the regarded wavelengths (315 and 375 nm).

The results are presented in Fig. 6, where the modelled SO<sub>2</sub> SCDs for 315 nm are plotted versus those at 375 nm. In addition to the modelled SO<sub>2</sub> SCDs also the measurements for 8 August 2008 are included in the figure. In both model runs a linear relationship between the SO<sub>2</sub> SCDs for both wavelengths is only found for rather small SO<sub>2</sub> concentrations (about <1 × 10<sup>13</sup> molec/cm<sup>3</sup>, corresponding to about 2.5 × 10<sup>18</sup> molec/cm<sup>2</sup>). For higher concentrations, the SO<sub>2</sub> SCDs for 315 nm increase

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much slower compared to those for 375 nm, and eventually they even decrease if the SO<sub>2</sub> concentration is further increased. This is a result of the very strong light absorption caused by SO<sub>2</sub> inside the volcanic plume. If aerosols are present, this effect is less pronounced (and for very high SO<sub>2</sub> concentrations even higher SO<sub>2</sub> SCDs for 315 nm are found compared to the non-aerosol case), because more light is scattered back by the aerosol particles to the satellite instrument from the layer of enhanced SO<sub>2</sub>. The comparison with the measured SO<sub>2</sub> SCDs clearly indicates that it is not possible to describe the satellite observations from 8 August 2008 without aerosols present in the volcanic plume. Good agreement for the upper bound of measured SO<sub>2</sub> SCDs for 315 nm is found for an assumed aerosol optical depth of about 3. The scatter of the measured data indicates that the aerosol load was different (and mostly higher than 3) in different parts of the observed plume.

## 6 Conclusions

An alternative evaluation range for the DOAS retrieval of SO<sub>2</sub> was proposed. It was shown that using the long wave UV window of 360–390 nm has a number of advantages in situations in which high SO<sub>2</sub> column densities are encountered. In particular, problems with non-linearity in the DOAS evaluation, stray light in the spectrometer, and signal dilution by scattering could be avoided or at least significantly reduced.

Although the absolute SO<sub>2</sub> absorption cross section in the long wave UV range is about three orders of magnitude smaller than around 320 nm, the high frequency (differential) absorption features are only about two orders of magnitude smaller. Because the intensity of incident solar radiation on the ground is about an order of magnitude higher in the long wave UV region, the smaller sensitivity towards SO<sub>2</sub> can in part be compensated by a lower photon noise level (about a factor of 4.2). However, since the quality of volcanic SO<sub>2</sub> measurements using the DOAS technique are rarely limited by instrument sensitivity but rather by retrieval accuracy, the suggested long wave UV evaluation region can be advantageous in many situations.

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spectral region to constrain the radiative transfer calculation in situations where this is possible will greatly improve the retrieval accuracy.

Finally, the implications of this study go beyond the measurement of SO<sub>2</sub> in volcanic plumes alone. While somewhat less sensitive, the novel evaluation range can also be applied to data evaluations of industrial plume measurements or urban pollution in general. The long wave UV range could be could, for example, also be applied to satellite data over highly polluted regions like China.

In summary, the proposed long wave UV retrieval range not only represents an additional, independent retrieval technique that can be used to validate measurements evaluated in the short wave UV, it also gives additional information which, if interpreted correctly, can further increase our understanding of radiative transfer in and around volcanic plumes.

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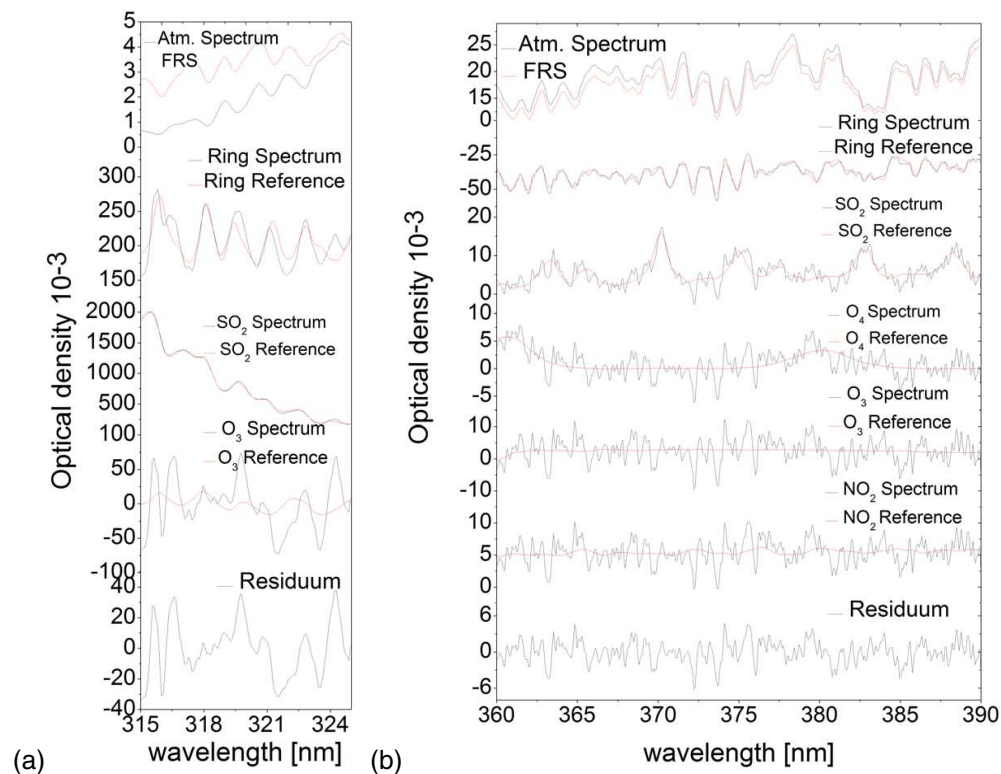
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**Fig. 1.** Example of the DOAS SO<sub>2</sub>-fit. Solid lines: measured spectra, dotted lines: reference spectra, the atmospheric spectrum was taken inside the volcanic cloud of Etna ( $\alpha=9^\circ$ ) on 10 June 2008 at 11:22 LT **(a)** evaluation in the short wave UV window 315–325 nm **(b)** evaluation in the long wave UV window 360–390 nm.

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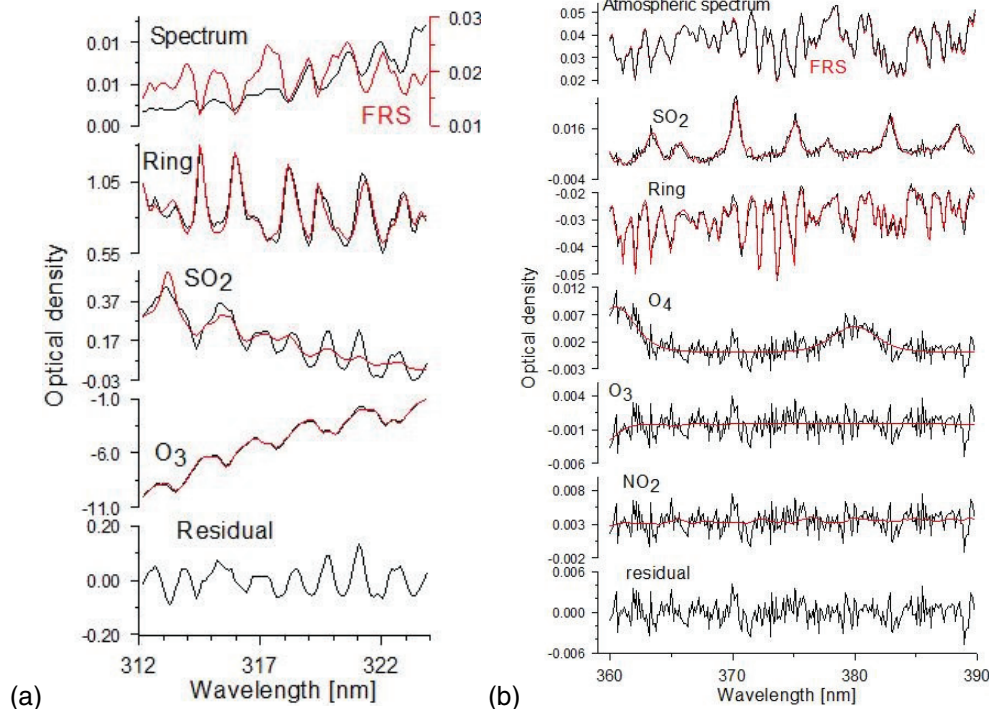
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**Fig. 2.** Example of a DOAS SO<sub>2</sub>-Fit. Solid lines: measured spectra, red lines: reference spectra, the atmospheric spectrum was taken on the 8 August 2008 after the eruption of the Kasatochi volcano **(a)** evaluation in the short wave UV window 312.1–324 nm **(b)** evaluation in the long wave UV window 360–390 nm.

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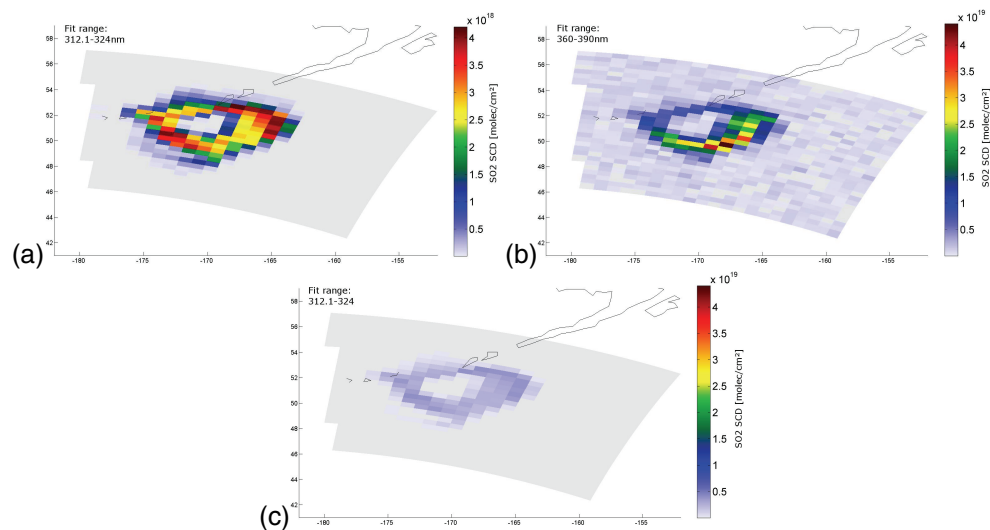
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**Fig. 3.** SO<sub>2</sub> plume after the Kasatochi eruption on 8 August 2008 recorded by GOME-2 **(a)** evaluation in the short wave UV window 312.1–324 nm (max. SO<sub>2</sub> SCD  $\approx 2.2 \pm 0.3 \times 10^{18}$  mol/cm<sup>2</sup>), **(b)** evaluation in the long wave UV window 360–390 nm (max. SO<sub>2</sub> SCD  $\approx 4.7 \pm 0.1 \times 10^{19}$  mol/cm<sup>2</sup>), **(c)** evaluation in short wave UV window using the same color scale as shown in (b).

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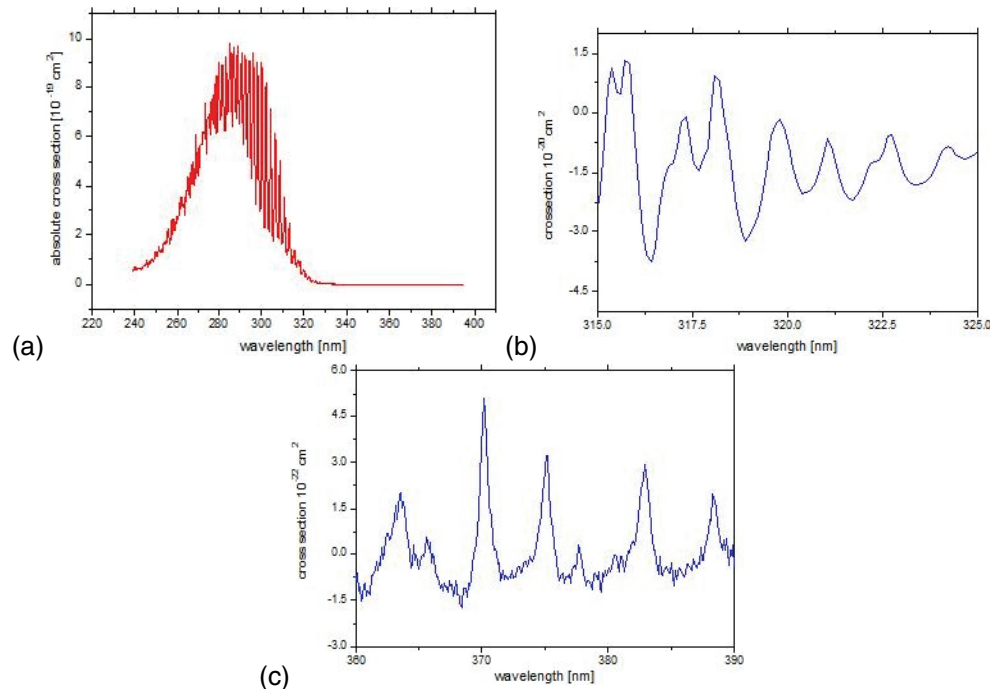
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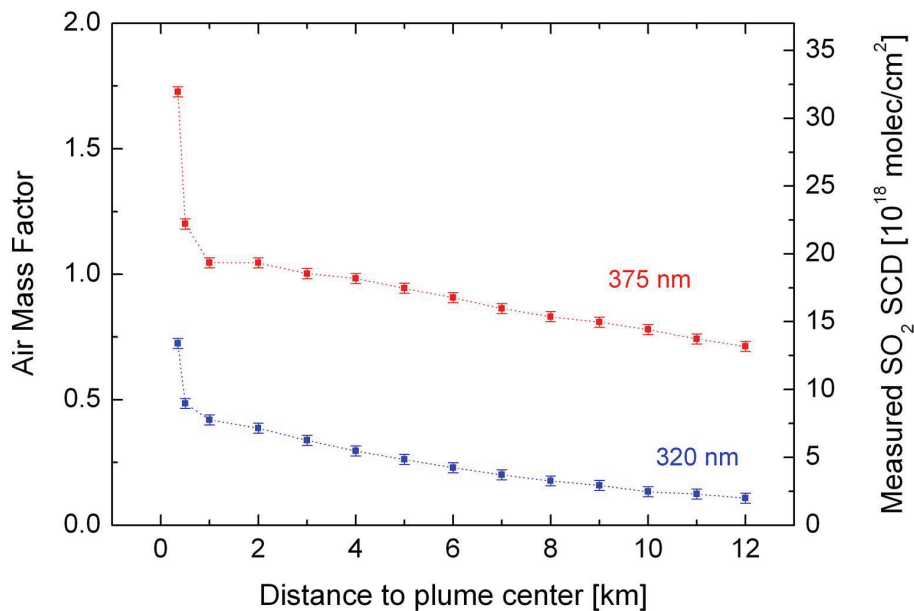


**Fig. 4.** SO<sub>2</sub> cross section for 273 K from Bogumil et al. (2003) **(a)** absolute cross section, **(b)** differential cross section high pass filtered for the wavelength range 315–325 nm **(c)** differential cross section high pass filtered for the wavelength range 360–390 nm. The differential absorption cross sections are  $\approx 4.5 \times 10^{-20}$  cm<sup>2</sup>/molec and  $6 \times 10^{-22}$  cm<sup>2</sup>/molec for the short wave UV and long wave UV regions, respectively.

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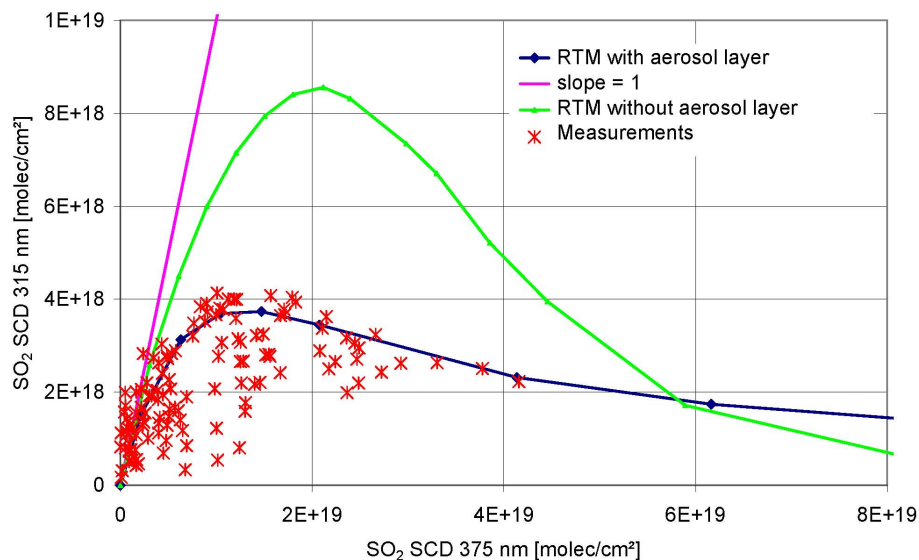


**Fig. 5.** Simulated air mass factors (AMF) for remote sensing measurements of a volcanic plume with high aerosol load performed at 320 (shown in blue) and 375 nm (shown in red). The distance between the volcanic plume and the remote sensing instrument was varied between 1 and 12 km.

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**Fig. 6.** Simulated SO<sub>2</sub> SCDs for 315 nm versus SO<sub>2</sub> SCDs at 375 nm for satellite remote sensing of the Kasatochi SO<sub>2</sub> plume on 8 August 2008 for two cases: without aerosols or with aerosol inside the plume (optical depth of 3). In addition to the modelled SO<sub>2</sub> SCDs also the GOME-2 measurements are included in the figure. Only if aerosols are present in the plume, simulations and measurements are in agreement.

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