

**Measurements of BC
in snow**

J. P. Schwarz et al.

Assessing recent measurement techniques for quantifying black carbon concentration in snow

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Abstract

We evaluate the performance of the Single Particle Soot Photometer (SP2) and the Integrating Sphere/Integrating Sandwich Spectrophotometer (ISSW) in quantifying the concentration of refractory black carbon (BC) in snow samples. We find that the SP2 can be used to measure BC mass concentration in snow with substantially larger uncertainty (60 %) than for atmospheric sampling (<30 %). Achieving this level of accuracy requires careful assessment of nebulizer performance and SP2 calibration with consideration of the fact that BC in snow tends to larger sizes than typically observed in the atmosphere. Once these issues are addressed, the SP2 is able to measure the size distribution and mass concentration of BC in the snow. Laboratory comparison of the SP2 and the Integrating Sphere/Integrating Sandwich Spectrophotometer (ISSW) revealed significant biases in the estimate of BC concentration from the ISSW when test samples contained dust or non-absorbing particulates. These results suggest that current estimates of BC mass concentration in snow and ice using either the SP2 or the ISSW may be associated with significant underestimates of uncertainty.

1 Introduction

The contribution of absorbing materials to the ratio of light scattering to light extinction by snow (i.e., snow albedo) is a critical parameter influencing the planetary albedo and, hence, the radiative balance of the Earth system (e.g., Flanner et al., 2007). A priori modeling of the impact of snow albedo on the radiation budget is a challenging task because snow albedo depends on many complex factors including snow structure (“grain size”, which depends on the temperature history of the snow), amount and composition of contaminants, snow depth, and characteristics of the underlying surface (including the presence of tall grasses or trees). Within this complicated framework scientists have been working towards identifying the incremental contribution of a particular contaminant – combustion-generated black carbon (BC) – on the snow’s absorption of light.

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Measurements of BC
in snow**

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The Single Particle Soot Photometer (SP2, Schwarz et al., 2008) is an aerosol instrument with the ability to quantify the mass concentration and size distribution of BC in ambient aerosol without significant sensitivity to other materials, and thus is a natural choice for application to measurements of BC-in-snow. McConnell et al. (2007), Bisiaux et al. (2011a), Bisiaux et al. (2011b), and Kaspari et al. (2011) have used it to measure BC in ice cores, and Ohata et al. (2011) have applied it to BC in rainwater. However, applying the SP2 to determining BC size distribution and concentration in water requires a careful assessment of additional uncertainties introduced by the need to aerosolize the liquid, and by the potential presence of much larger BC particles than are typically observed in the atmosphere. The assessment provided here is for snow samples, but we expect it to translate well to measurements of BC in ice.

The Integrating Sphere/Integrating Sandwich Spectrophotometer (ISSW, Grenfell et al., 2011) is a new instrument based on combining and modifying elements of previous instrument designs (including the Integrating Sandwich, Clarke et al., 1987). The ISSW measures the wavelength-resolved absorption of particulate material collected on filters from melted snow samples. The BC concentration of the filtrate is determined by quantifying the fraction of absorption associated with the wavelength (λ) dependence believed to be characteristic of BC aerosol, λ^{-1} . The remainder of the absorption is attributed to other (non-BC) particulate constituents. The ISSW's estimates of BC concentration rely on conversion of measured absorption to BC mass using a set of calibration standards prepared in the laboratory.

The ISSW has been used to chart BC loadings in a pan-Arctic snow study, which attributed at least half of solar absorption by particulate contamination to BC (Doherty et al., 2010). The ISSW has not been directly compared against independent assessments of BC concentration in snow, nor has it previously been tested with laboratory standards containing non-BC materials in addition to BC.

Here we apply the SP2 to quantification of BC concentration and size distribution in snow (Sect. 2.1), briefly describe the ISSW (Sect. 2.2), present preparation of laboratory samples (Sect. 2.3), and then compare SP2 and the ISSW measurements of

laboratory standards including BC, dust, and scattering particulate in Sect. 3. Section 4 summarizes our conclusions and their implications.

2 Experimental

2.1 SP2 technique for BC-in-snow

5 The SP2 uses an intense intra-cavity laser (1064 nm) to heat the refractory BC component of individual aerosol particles to vaporization. At vaporization the peak visible thermal radiation emitted by the BC is measured, to constrain the color-temperature of the hot material (and thus BC composition), and to provide a measure of its mass. This detection method is not affected by the mixing state or morphology of the particles for most of the BC mass observed in the ambient atmosphere (Cross et al., 2010; Moteki and Kondo, 2010). Mass loadings in air are quantified by integrating the individual BC masses associated with individual particles detected within a given amount of air. We convert the measured mass of BC in a particle to a volume-equivalent diameter (VED) by assuming a void-free density of 1.8 g cm^{-3} .

15 SP2 measurements of BC-in-snow require consideration of two issues not directly associated with the SP2's use for quantifying BC in air. First, the efficiency with which the particles in liquid are mixed with air and then transported to the SP2 laser must be known; in general, this is a function of particle size. Second, the SP2 range of detection must be extended to larger sizes than typically encountered in the atmosphere, because the BC size distribution in snow samples can extend to substantially larger size than those found in the typical ambient atmosphere. This requires calibration of the mass-vs-thermal emission relationship to large mass, as well as statistical analysis of the impact of large, rare BC particles on the mass integral. Section 2.1.1 presents the approach used to calibrate the SP2 for large BC particles, Sect. 2.1.2 discusses aerosolization of particulates in liquids such as melted snow, and Sect. 2.1.3 presents the size distribution of BC observed in snow melt and its impact on SP2 uncertainties.

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



2.1.1 Calibration of the SP2

The SP2 was calibrated with fullerene soot (Alpha Aesar, Inc., Wardhill, MA, USA, Lot# F12S011) over a range of 0.5–40 fg (corresponding to ~ 80–350 nm volume-equivalent diameter (VED) assuming a density of 1.8 g cm^{-3}) selected with a DMA. This is a range that typically covers >75 % of ambient accumulation mode mass. The calibration of mobility diameter to BC mass for this material over this range was provided by averaging the results of Moteki and Kondo (2010) and Gysel et al. (2011). The calibration was very close to purely linear over this range.

For larger BC particles, the empirical power-law dependences in Moteki and Kondo (2010) were used. The Moteki and Kondo dependences relate SP2 detection signal to BC mass raised to a power that is anti-correlated to the material's density, and were measured up to near 600 nm VED. A power of 0.9 was chosen as intermediate between that of the ambient BC observed in Tokyo during that study (0.95) and that of the densest BC identified by Moteki and Kondo as formed of agglomerations of primary particles (0.85). A sensitivity study using the more extreme power values showed a range of 20 % in the inferred BC mass for the larger size distributions observed in snow, which extend to 2000 nm VED (Sect. 2.1.3). Because this extrapolation extends so far in mass space, we double the range of the inferred mass to estimate the uncertainty associated with the contribution of large BC masses to total BC mass. Hence the total uncertainty we associate with the calibration of the SP2 is 50 %, driven primarily by a 30 % uncertainty in the range of SP2 response to different BC materials (a very conservative estimate based on the results of Moteki and Kondo (2010) and Gysel et al. (2011)) added in quadrature with 40 % uncertainty due to the extrapolation of SP2 calibration to large sizes. Direct calibration using an Aerosol Particle Mass Analyzer (Kanomax USA Inc, Andover, NJ) or Couette Particle Mass Analyzer (Olfert et al., 2005), which were not available for this study, could reduce this uncertainty.

The SP2 was operated with laser intensity ~700 nW/(220-nm PSL). Laser intensity was measured with polystyrene latex spheres as described in Schwarz et al. (2010a).

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Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



This intensity was sufficient to insure proper detection of BC over the mass range studied here.

2.1.2 Aerosolization of Particulates in Snow Melt

The SP2 can only sample aerosol, which we generated from liquid samples with a Collison-type nebulizer (“CNEB”) that was built in-house and is similar to a commercial TSI Collison nebulizer #3076 (TSI Inc., Shoreview, MN). The nebulizer uses a stream of pressurized air expanding through a critical orifice to shear liquid into small droplets that are dried out, releasing any particulate contained in the droplets to the air. The liquid samples were pumped into the nebulizer with a peristaltic pump at rates of order 0.005 cm s^{-1} . The length of the peristaltic pump line was varied to test for losses of BC in the liquid phase. No dependence on line length was observed, yet there were clear signs that when very concentrated mixtures were passed through the line (for example to provide BC for size selection with a differential mobility analyzer), residues of BC on the line were released into the next solutions sampled. This suggests whatever the loss/resuspension of BC to the pump line occurs, its most significant impact is likely on contaminating BC measurements. To avoid these issues and minimize the BC contamination, we used clean, new peristaltic pump lines for sampling snow and laboratory concentration samples after more concentrated solutions, and observed contamination in the range equivalent to $0.05\text{--}0.5 \text{ ng g}^{-1}$ BC concentration. The lower end of this range was limited by stray BC particles in the SP2 laser head, rather than by BC contained in the air stream from the nebulizer. Clearly, depending on the concentration in an unknown sample, care must be taken to establish that the background is suitably low.

To assess the efficiency with which particles in liquid were nebulized and transported to the SP2, we used polystyrene latex sphere (PSL) concentration standards (Poly-science Inc, Warrington, PA, USA) in sizes of 220, 356, 505, 771, 1025, and 1537 nm diameter to test both the Collison nebulizer and an ultrasonic nebulizer, the U5000AT (U5000AT, CETAC Technologies, Omaha, NE, USA), which was used in previous SP2

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Measurements of BC
in snow**

J. P. Schwarz et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

determinations of BC in ice and rainwater. PSLs have a density of 1.05 g cm^{-3} . PSL-number concentration standards in water were generated using the solid fraction weight provided by the manufacturer and checked against laser transmission spectroscopy (LTS, Li et al., 2010), which confirmed the concentrations to 7 %, excepting the 220 nm standard, which was used without confirmation from the LTS. The standards were diluted to $\sim 10^8 \text{ particles cm}^{-3}$ for use with the nebulizer.

The nebulization efficiency at each size was determined by measuring, with the SP2, the airborne concentration of the PSL particles. As the PSL concentration in the liquid standard was known, this is essentially a measure of how much liquid the nebulizer aerosolized into a given volume of air. For the particular setup of flows/liquid pumping rate we used, this efficiency was of order $1 \mu\text{g-liquid cm}^{-3}$ for 505 nm PSLs. Only particles within the appropriate size range of the PSLs were counted in the aerosol state (i.e., agglomerations or PSL fragments were not included). Aerosol particles outside the appropriate size range were very scarce, but the 505 nm PSL standard was particularly clean. The absolute value of the efficiency at a given size is not fundamentally important (so long as it was known), hence this size was used to provide the relative reference in Fig. 1, which shows the nebulization efficiency of other PSL sizes with respect to that for 505 nm.

The nebulization efficiency was parameterized as a function of PSL diameter for eventual correction of BC mass mixing ratios measured with the same nebulizer. For sizes below 500 nm, it was modeled as unity. We then fit a line to the relative efficiencies for sizes larger than 500 nm, fixing the intercept at 500 nm to unity. This produced the “best-estimate” efficiency curve in Fig. 1. To test sensitivity to the correction to this parameterization, we also produced a line intercepting 1 at 350 nm, and roughly tracing the lowest relative efficiencies at larger sizes (dashed line in Fig. 1). This is the “worst-case” efficiency curve.

The U5000AT was found to aerosolize PSLs larger than $\sim 500 \text{ nm}$ with very poor efficiency (i.e., $\sim < 10\%$) with respect to smaller particles, and to produce BC size distributions that had essentially no contribution from BC particles larger than $\sim 500 \text{ nm}$

VED. As we show later, there can be significant BC mass contributions in sizes larger than this, hence the U5000AT was not used here, and we recommend in depth examination of its behavior to those intending to apply it to aerosolizing BC from a liquid sample.

Application of the nebulization efficiency curves, which were based on PSL measurements, to correct observed BC mass distributions was conducted by assuming that the nebulization efficiency was a function of particle stopping distance, which is proportional to particle diameter squared, and linearly proportional to the particle density. The PSL density and sizes were well known. For BC, we adopted a value for the effective density of 0.8 g cm^{-3} , which is the average density measured between 400–900 nm mobility diameter for six BC materials in Moteki and Kondo (2011); these materials did not include the glassy carbon spheres from that study, which are clearly not related to BC structure in the ambient, and have much higher density. Thus, the SP2 determination of BC mass was used in conjunction with the assumed effective density to calculate the stopping distance of the particles. Our results (Sect. 3) suggest that the presence of non-BC materials internally mixed on BC in nebulized, melted snow will not significantly affect the net particle mass or density; this may not always be the case. Five size distributions of BC from first-melt snow samples were scaled with both the best-estimate and worst-case efficiency curves (discussed in Sect. 2.1.3, the best-estimate average of which are shown in Fig. 2). On average, the best-estimate nebulization efficiency parameterization lead to a 25 % shift upwards in total BC mass, and the worst case produced an average correction of nearly 50 %. Changing only the assumed density from 0.8 to 0.9 g cm^{-3} for the BC material had essentially no impact on our best-estimate of the correction due to imperfect nebulization efficiency. Hence, we suggest that an uncertainty estimate of 30 % associated with the size-dependent nebulization efficiency determination is reasonable for this system and snow samples. This uncertainty (30 %) added in quadrature to the 50 % uncertainty due to the BC calibration results in a total SP2 uncertainty estimate of 60 %. Note that the uncertainty

**Measurements of BC
in snow**

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



associated with quantifying BC mass concentration in liquid in cases of smaller BC size distribution than those seen here would likely be associated with smaller uncertainty.

2.1.3 BC size distributions in snow

The impact of the nebulization efficiency curves on the BC mass observed by the SP2 depends on the underlying size distribution of the particulate material in the liquid, which we assume here is not shifted from that of the unmelted snow during melting in the laboratory. Note that the size distribution itself, in addition to the BC concentration in snow, may be of scientific interest, for example because it can affect BC mass absorption efficiency, and hence snow albedo. We find that freeze/thaw cycles affect BC size distributions in water, hence it is also possible that the size distribution can give information about the thermal history of the snow, or about the mechanisms by which snow forms.

As a first estimate of the likely range of BC size to be found in snow, we examined fresh local snow and field samples from the Arctic. The local snow was collected within 60 km of Denver, CO in both semi-rural and rural areas. The local snow samples were collected in pre-chilled glass containers, then thawed in a microwave for immediate sampling with the SP2, followed by additional testing. Figure 2 shows the the average mass distribution calculated from five local first-melt snow samples. The size distributions directly observed by the SP2 (not shown) were individually scaled with the best-estimate CNEB efficiency curve as discussed in Sect. 2.1.1.

The field samples were collected, then kept frozen until they were melted rapidly in the laboratory with a microwave oven, usually within a few days of being gathered in the field. They were then decanted into high-density polypropylene (HDPE) bottles and immediately deep-frozen for storage before transport. During transport, these samples were partially thawed for a period of approximately 1 day, then refrozen for storage. Due to this somewhat extended processing, we do not present quantitative results from the Arctic samples.

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Measurements of BC
in snow**

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



For sampling in the SP2, the field samples were thawed in a sonicator where they were agitated for 20 min, and then immediately sampled with the SP2. After sampling, they were first acidified to 0.5 M nitric acid, as per Kaspari et al. (2011), then sonicated for another 20 min before resampling with the SP2. There was a slight reduction in the mass fraction of super-micron BC particles after the acidification step, likely due to the acid helping break apart coagulated BC fragments. If the BC size distribution in snow is of interest, acidification should not be carried out.

Many of the snow samples had significant BC mass contributions at larger sizes than in typical ambient air (where typically >90% of the BC mass is below 600 nm VED). The peak diameter of the mass distribution of the sample average shown in Fig. 2 were similar to that for typical atmospheric BC but some of the samples contributed more strongly to the larger-BC size fraction. Hence it appears likely that there is significant variability in the size distribution of BC in snow. We use the local first-melt snow average distribution to calculate a reasonable uncertainty analysis of the SP2 technique for BC in snow (Sect. 2.1.2).

Testing with laboratory standards as well as with fresh local snow showed that individual freeze/thaw cycles cause the agglomeration of a small fraction of BC mass into larger sizes, without dramatically shifting the underlying BC size distribution. It follows that the most accurate measurement of the BC size distribution will be from the first melt of the particular snow sample in question. Laboratory standards were tested up to high concentrations of BC (Table 1, discussed in Sect. 2.3), and the BC size distribution was independent of BC concentration. This indicates that the large BC particles observed in snowmelt are not merely artifacts of the nebulization process.

The fact that the BC mass distributions peak well below the range where the nebulizer fails to aerosolize particles suggests that mass contributions of large BC agglomerations are not likely to cause significant underestimation in our treatment. This assumption, of course, may not hold for all samples, in which case this technique would only provide a lower limit on BC concentration. Clearly, the uncertainty associated with the SP2 determination of BC concentration in water will depend on the

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



characteristics of the individual nebulizer used, as well as the underlying size distribution of BC-containing particles, and the technique used to calibrate the nebulization efficiency. For example, if a gravimetric standard is used to calibrate the nebulizer, then errors may occur if the size distribution of the experimental sample is not the same of that of the standard, depending on the nebulization efficiency size dependence of the nebulizer. Note that our measurements suggest that BC tends not to be internally mixed with much non-BC material in aerosolized melted snow (Sect. 3).

In typical ambient air samples, only a small correction is required to scale the observed mass to the mass of BC within the entire accumulation mode because the detection range of the SP2 is typically only slightly narrower than that of BC in the ambient (Schwarz et al., 2010b), and the assumption that negligible BC mass exists in sizes below the SP2 detection threshold is generally valid. This is not true of very fresh emissions, which tend to smaller sizes than reliably detectable by the SP2. Due to the increased uncertainties associated with SP2 measurements of BC larger than typically seen in the ambient air, any method one can use to break particles into smaller sizes will improve the concentration measurement (so long as they are not broken down to sizes below the practical SP2 detection range). We assume, based on the clear decrease in BC mass concentration towards the lower size limit of the SP2 that there is negligible BC mass below the SP2 detection range in all the snow samples. We found that acidification of the field samples to 0.5 mol l^{-1} with 70 % nitric acid (following Kaspari et al., 2011, to release BC from container walls), followed by sonication for 20 min broke down some of the large BC agglomerations (Fig. 2). However, measurements with the SP2 indicated that in high-density polyethylene bottles, BC losses in samples at room temperature for even one day were significant (highly variable, but of order 50 %), and acidification failed to restore most BC lost to walls. BC losses in glass containers at room temperature were much more limited.

A large number, of order 10 000, BC-containing particles need be sampled to reduce the statistical uncertainty associated with the incremental detection of individual large particles; these are rare but may contain a significant fraction of the total BC mass. In

**Measurements of BC
in snow**

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



snow samples in which BC is very dilute, yet which include large particles, this could require long integration times. A straightforward way to estimate the statistical uncertainty associated with these types of measurements is to compare the concentrations measured twice over equal integration times. For the ambient samples tested here, we found that subsequent 5 min integrations were typically sufficient to ensure that $\leq 20\%$ statistical uncertainty was associated with this effect in our measurements, although this was generally insufficient to generate a reasonable BC size distribution.

2.2 Experimental technique for the ISSW

The ISSW was described in detail by Grenfell et al. (2011), with additional uncertainty analysis in (Doherty et al., 2010). Briefly, the ISSW spectrally resolves light absorption by snow contaminants collected on a filter for wavelengths 420–730 nm. Absorption associated with BC and non-BC material is differentiated by assuming absorption by BC scales with wavelength as λ^{-1} and absorption by the non-BC constituents scales as $\lambda^{-\alpha}$, where 1.0 and α are the absorption Ångström exponents (AAEs) of the BC and non-BC constituents, respectively. The value of α used is selected based on chemical analysis of the snow melt water; if the chemical signature of combustion aerosol is dominant, $\alpha = 5$ is used (Doherty et al., 2010). If high iron content is present, the snow is assumed to have dominant non-BC absorption contributions from mineral dust, and thus is analyzed with a value of α of 3.5. If further chemical data are available, assumptions about iron and organic carbon spectrally-resolved mass absorption efficiencies can be used to estimate α .

The fraction of absorption attributed to BC is converted to a mass concentration of BC using a set of gravimetrically determined synthetic BC standards that have a mass absorption coefficient (MAC) of $6.3 \text{ m}^2 \text{ g}^{-1}$. If the MAC of the unknown BC on the filter is higher (lower) than this value, then the estimated BC mass values will be too high (low). A correction for filter undercatch of 15% is applied to field samples (Doherty et al., 2010).

The total uncertainty associated with the ISSW BC concentration determination for ambient snow is estimated as 40 %, which is the sum, in quadrature, of 11 % for instrumental uncertainty, 15 % for undercatch uncertainty, 17 % for calibration material MAE uncertainty, and 30 % for uncertainty in the AAE of non-BC material.

2.3 Laboratory standards

Gravimetric standards of BC (using the fullerene soot used to calibrate the SP2) and test dust (ISO-12103-1, A2 Fine Test Dust, Powder Technology, Inc., Burnsville, MN, USA), in addition to PSL concentration standards were used to test the SP2 and ISSW. Mixtures of dust or BC in deionized water were allowed to settle for several days to remove large particles, and then the top liquid was slowly drawn off for use and drying/weighing. For the dust standards, some settled material was included in the drawn-off mixture. Note that allowing large BC particles to settle out of solution resulted in a BC size distribution for the standards that was narrower and smaller than that observed in snow (Fig. 2). Thus the laboratory samples do not inform us about uncertainties in the correction calculated for the SP2 due to the size dependence of the nebulizer efficiency (described in Sect. 2.1.2).

For the BC standard three samples were dried to determine the BC concentration by weighing the solid residues, with a standard error of 6 %. Impurities in the dilution water contributed less than 1 % to the dry weights. Master solutions of the three materials (BC, PSLs and dust) were diluted and mixed to create a wide variety of mixtures and concentrations of either pure BC, a mixture of BC + PSL or BC + dust, and BC-free dilutions of pure PSLs or pure dust. The concentrations of dust tested roughly spanned the range observed in ambient snow samples from north and northeast China. The range of PSL concentrations was similar to that of the dust standards

The laboratory samples were prepared, placed in glass bottles, immediately characterized with the SP2 in Colorado, and then transported to the ISSW at the University of Washington for filtration within 24 h. These samples were not frozen. Some duplicate

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



samples were retested with the SP2 after this amount of time, and were found to match the fresh samples well within expected uncertainties.

For the laboratory samples, ISSW filter undercatch was much more significant than for ambient field samples, likely because of the difference in size between the fullerene soot gravimetric standards and ambient snow BC. Undercatch was $38 \pm 3\%$ based on refiltering postfilter liquid using finer-mesh ($0.2 \mu\text{m}$) nuclepore filters, and analyzing with the ISSW. The ISSW laboratory results are scaled to remove this artifact. The fact that this correction is quite different from that applied to field samples suggests that the ISSW undercatch correction may be sensitive to the particulate size distribution in snow.

To explore the possible magnitude of a low bias in the SP2 BC mass determination due to BC internal mixing with large amounts of non-BC material (which would affect nebulization efficiency), two analyses were performed. First, a test using laboratory samples bounded the number-fraction of BC particles in solution that became attached to dust particles during a single freeze/thaw cycle; only $\sim 1\%$ did. Second, the number fraction of BC particles associated with detectable (by the SP2) amounts of non-BC material was calculated for field samples (as in Schwarz et al., 2008). On the whole, the fraction showed that the BC was predominantly bare, and that the acidification process did not result in any measurable additional coatings on the BC. These results indicate that this mechanism is unlikely to broadly bias the SP2 results.

3 Laboratory results

Figure 3 shows the results of measurements of BC and non-BC concentration standards. For pure BC-in-water samples, the SP2 and ISSW produced values that correlated strongly with each other with very little scatter. This correlation revealed common mode variations in the gravimetric value at the level of $\sim 10\%$ from the mean that we attribute to inaccuracies in diluting the samples. These variations do not affect our conclusions. The SP2 values matched, within uncertainties, the gravimetrically determined

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Measurements of BC
in snow**

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



concentration of BC. The ISSW values were biased high by a factor $1.40 \pm 9\%$, likely due to differences in the optical properties of the fullerene material used in the gravimetric standards and the ISSW calibration material (Monarch 71). A mass absorption coefficient of $8.84 \text{ m}^2 \text{ g}^{-1}$ for the fullerene soot, which is well within range associated with BC, would correspond to this 40%. Use of an incorrect assumed value of the AAE for BC (in this case the fullerene) will also produce a bias in the ISSW BC estimates. The pure fullerene samples have an ISSW-measured AAE of 1.2 ± 0.2 , not the assumed 1.0. Accounting for this, however, increased the bias by an additional $\sim 15\%$.

The SP2 proved to have a small positive offset associated with dust content, and none with PSLs. The SP2 sensitivity to dust was caused by incandescent signals associated with small amounts of dust that produced visible light signals through heating in the laser. In real snow samples with low BC concentrations, it would be difficult to separate these signals from those of BC. In the most contaminated dust samples ($\sim 50 \mu\text{g g}^{-1}$ dust), this led to a $15 \text{ ng-BC/g-H}_2\text{O}$ positive offset in the SP2 BC determination. Dust composition could affect this sensitivity. It is possible that this effect could be reduced via stringent discrimination of particle microphysical properties (including color temperature and scattering cross-section normalized to incandescent light signal) coupled with additional validation on well-controlled laboratory samples.

The ISSW results revealed positive biases associated with contamination of the BC samples with both PSLs and with dust. For Fig. 3, the ISSW values were calculated using an α (non-BC AAE) of 4.4, based on unambiguous ISSW measurements of pure dust-in-water samples at several concentrations, and hence represent a “best-case” evaluation; use of this α results in a lack of dependence of the relative bias on the amount of BC present. The high bias was as high as a factor 3 for the most dust-contaminated samples. Non-BC AAE was varied from 3.5–5.0 to test for the ISSW sensitivity to this parameter. The former is based on measurements of ambient dust AAE, which range from approximately 2 to 4 (e.g., Fialho et al., 2005, 2006; Alfaro et al., 2004; Bergstrom et al., 2007; Yang et al., 2009). The latter is an appropriate α for combustion aerosol (Doherty et al., 2010). For $\alpha = 3.5$, lower than the appropriate

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



4.4, the relative bias between estimated BC concentration and gravimetric BC became strongly dependent on the BC concentration in the sample, with non-physical negative BC concentration estimates at high dust : BC (Tables 1 and 2). With assumed $\alpha = 5.0$, larger than the appropriate 4.4, a purely positive bias with a stronger dependence on dust : BC ratio was observed. These high biases are larger than expected based on the uncertainty analysis from Doherty et al. (2010). For pure BC standards this range of α resulted in only a 5 % range in estimated BC concentration.

The monodisperse nature of the PSL-containing test standards produced strong wavelength dependencies in the inferred ISSW absorption, hence these results are not quantitatively presented. However, the fact that scattering artifacts caused an “absorption” signal in the ISSW indicates that the ISSW has a positive artifact associated with non-absorbing particulate matter. The size of the excursions suggest that purely scattering particles could significantly bias affect ISSW absorption measurements. This sensitivity to light scattering may contribute to the two separate biases that appear to be affecting the ISSW results: the bias of $\sim 40\%$ in the pure BC (fullerene soot) samples that may be due to the MAC of the gravimetric standard material; and a high relative bias that is a function of non-BC aerosol concentration. The use of the optimal values of α for these laboratory samples indicate that even detailed information about the absorption spectra of non-BC contaminants is insufficient to remove the ISSW sensitivity to the presence of other aerosol constituents.

4 Conclusions

We find that BC in snow tends to be present at larger sizes than in the typical ambient atmosphere. This has implications for determinations of BC mass content in snow, as well as possible implications for our understanding of the roles of BC in affecting snow albedo and in evincing snow history. Due in part to the large sizes of BC observed in snow, use of the SP2 for BC-in-liquid measurements requires careful characterization of nebulization efficiency as a function of particle size, as well, potentially, as

**Measurements of BC
in snow**

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



5 extrapolation of instrument response to very large BC masses. The inherent uncer-
tainties associated with these complications are unlikely to be amenable to substan-
tial reduction, yet are still small enough that the SP2 technique can contribute mean-
ingful information about BC content in liquid. We estimate that the SP2 can provide
a measurement of BC-in-snow with 60% systematic uncertainty, dominated by cali-
bration uncertainty and limiting assumptions about the size distribution of BC in the
snowmelt. A necessary assumption is that BC mass attached to larger particulate ma-
terials ($> \sim 2 \mu\text{m}$) in the snow mass, or in sizes smaller than the accumulation mode is
negligible. This assumption is supported by the observed BC mass size distributions
10 shown in the paper, and by tests indicating that BC does not tend to be associated
with large amounts of non-BC material in individual particles. The SP2 can provide the
size distribution of BC in snow, within the range of sizes aerosolized with reasonable
efficiency, and within the calibration range applicable to the SP2. Due to these issues,
previous SP2 determinations of BC concentration in snow or ice have may be asso-
15 ciated with uncertainties significantly larger than estimated. We recommend careful
assessment of the size dependence of the nebulization efficiency when using the SP2
to study the size distribution or mass concentration of BC in liquids.

The SP2 has a small positive artifact due to the difficulty of separating mass contribu-
tions of the uncommon non-BC materials that incandesce like BC. Based on laboratory
20 tests with dust, this artifact is estimated to be a $+15 \text{ ng g}^{-1}$ offset in the most highly
contaminated samples, equivalent to the most dust-laden snow samples from a recent
field sampling campaign in Central China.

Laboratory testing of the ISSW revealed biases associated with both non-BC light-
absorbing and purely scattering particulate in water. For dust, this bias was propor-
25 tional to the amount of contaminant, with a value $\sim \pm 25\%$ at low dust concentration,
and, at the highest dust concentrations, peaking at a factor 3 and offsetting BC con-
centration values by up to $\sim 600 \text{ ng-BC/g-H}_2\text{O}$. This bias is inconsistent with the current
uncertainty estimate for the ISSW. The fact that the ISSW responded to non-absorbing
contaminants indicates that the single-scatter albedo (SSA) of the aerosol can change

**Measurements of BC
in snow**

J. P. Schwarz et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



ISSW BC concentration determinations. This suggests that the size distribution and mixing state of material on the filter may impact ISSW results. Additionally, we are not able to distinguish the contribution of this scattering artifact to the biases associated with dust mixed into BC gravimetric standards; the observed high bias in ISSW-derived BC in the presence of dust contaminants may be due entirely to scattering sensitivity.

These results support recent concern about the relative contribution of BC to total absorption in snow. Results with the ISSW in the Arctic (Doherty et al., 2010), which themselves suggested that non-BC absorbers play a significant role in snowpack light absorption by particulates, may need to be revised further to a point where BC is not the primary absorbing material in the snow.

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**Measurements of BC
in snow**

J. P. Schwarz et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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**Measurements of BC
in snow**

J. P. Schwarz et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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Measurements of BC in snow

J. P. Schwarz et al.

Table 1. Laboratory sample information. Labels beginning with “B” refer to pure BC in water; those beginning with “D” refer to samples that include dust with or without BC. Samples with PSLs are not included here (as discussed in the text). The ISSW results were calculated for $\alpha = 4.4$.

Sample label	Gravimetric BC (ngg^{-1})	Gravimetric dust ($\mu\text{g g}^{-1}$)	SP2 BC (ngg^{-1})	ISSW BC (ngg^{-1})	Absorp. Å	SP2: grav.	ISSW: grav.
B1	1192	0		1889	1.3		1.59
B2	1192	0	1110	1797	1.2	1.07	1.51
B3	335	0	356	467	1.5	0.94	1.39
B4	746	0	835	1165	1.3	0.89	1.56
B5	1068	0	1004	1268	1.2	0.94	1.19
B6	496	0	462	569	1.4	0.93	1.15
B7	217	0	214	244	1.5	0.99	1.12
B8	62.3	0	54	73	1.5	0.86	1.18
B9	12.2	0	16	9	1.7	1.27	0.76
D1	60.6	1.6	59	73	2.1	0.97	1.21
D2	66.5	5.3	68	114	2.5	1.02	1.71
D3	57.5	49.2	84	168	3.8	1.46	2.92
D4	454	2.5	426	648	1.5	0.94	1.43
D5	478	7.4	440	832	1.7	0.92	1.74
D6	436	52.1	440	1232	2.5	1.01	2.83
D7	0	8.2	3	3	4.3	n/a	n/a
D8	0	54.1	14	-32	4.5	n/a	n/a

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Measurements of BC in snow

J. P. Schwarz et al.

Table 2. The ratio of ISSW measured concentration to the gravimetric concentration (“grav”) calculated for different values of α for the dust-containing samples.

Sample label	Gravimetric	Ratio dust: BC (10^3)	ISSW: grav.		
	BC concentration (ngg^{-1})		$\alpha = 3.5$	$\alpha = 4.4$	$\alpha = 5$
D1	60.6	0.03	1.01	1.21	1.30
D2	66.5	0.08	1.21	1.71	1.92
D3	57.5	0.86	-2.43	2.92	5.16
D4	454	0.01	1.34	1.43	1.46
D5	478	0.02	1.58	1.74	1.81
D6	436	0.12	2.02	2.83	3.16

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Measurements of BC in snow

J. P. Schwarz et al.

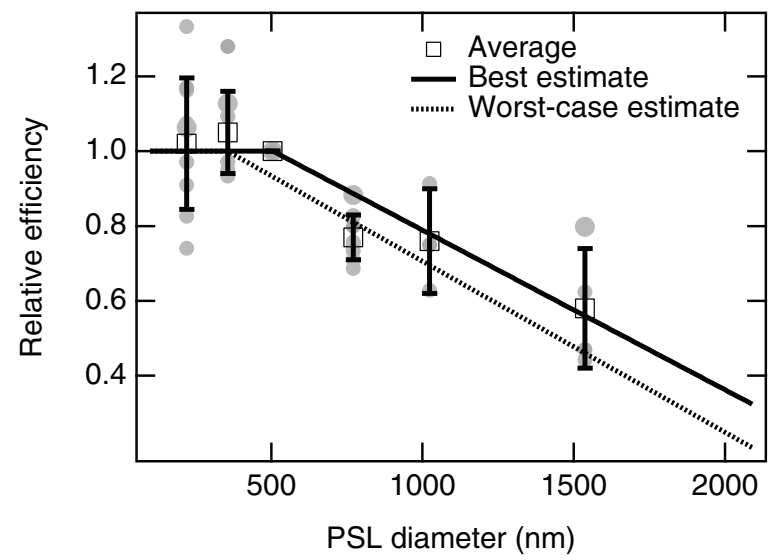


Fig. 1. Relative transmission efficiency of PSLs in liquid to the SP2 after aerosolization with a collision-type nebulizer. The results are scaled to the transmission of 505 nm PSLs.

Title Page

Abstract	Introduction
Conclusions	References
Tables	Figures

⏪ ⏩
⏴ ⏵
Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Measurements of BC in snow

J. P. Schwarz et al.

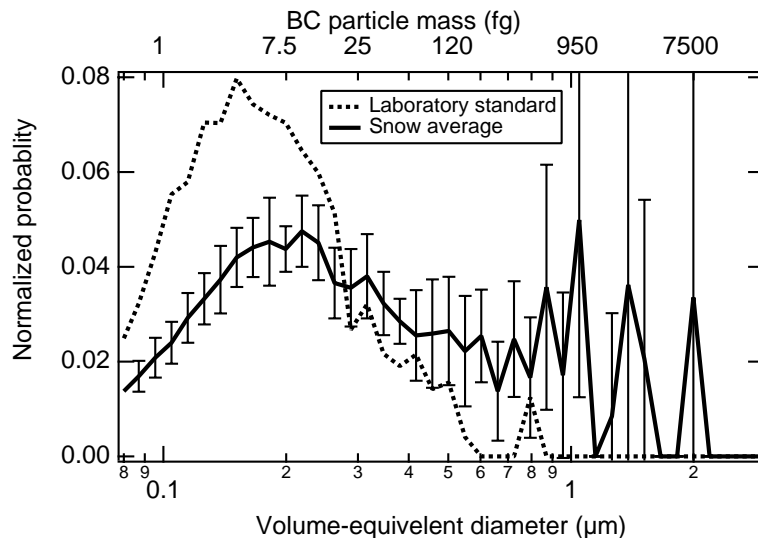


Fig. 2. Mass distributions from five first-melt snow samples from Colorado measured with the SP2 (heavy line). The whiskers indicate the standard deviation about the mean. Also shown is the mass distribution of the laboratory BC standard, which is similar to that typically observed in the ambient atmosphere (dashed). The vertical axis is $dM/d \log(D)$ with all distributions normalized to an area of 1. These mass distributions were scaled with the best-estimate nebulizer efficiency curve to represent the actual size distribution of the BC in the liquid.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



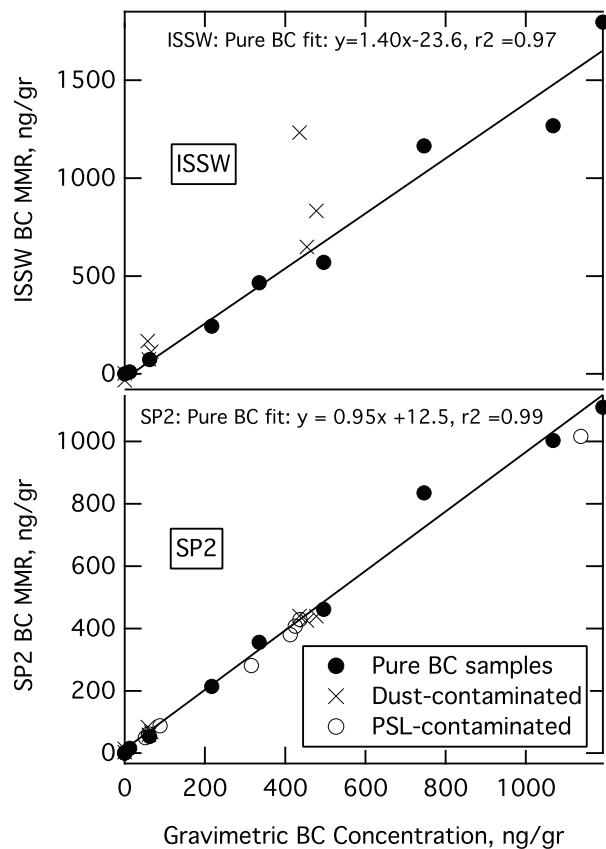


Fig. 3. Top: ISSW determinations of laboratory mixtures of fullerene soot and test dust with water; increasing concentrations of dust lead to increasing deviation from the pure-BC trend. Bottom: SP2 determinations for these samples, and additionally for samples contaminated with PSLs. The solid lines represent best-fit lines only to uncontaminated BC-in-water samples.

Measurements of BC in snow

J. P. Schwarz et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

◀ ▶

◀ ▶

Back Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

