

**High spectral
resolution ozone
absorption – Part 1**

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High spectral resolution ozone absorption cross-sections – Part 1: Measurements, data analysis and comparison with previous measurements around 293 K

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Abstract

In this paper we discuss the methodology of taking broadband relative and absolute measurements of ozone cross-sections including uncertainty budget, experimental set-ups, and methods for data analysis. We report on new ozone absorption cross-section measurements in the solar spectral region using a combination of Fourier transform and echelle spectrometers. The new cross-sections cover the spectral range 213–1100 nm at a spectral resolution of 0.02–0.06 nm in the UV-vis and 0.12–0.24 nm in the IR at eleven temperatures from 193 to 293 K in steps of 10 K. The absolute accuracy is better than three percent for most parts of the spectral region and wavelength calibration accuracy is better than 0.005 nm.

The new room temperature cross-sections data are compared in detail with previously available literature data. The temperature dependence of our cross-sections is described in a companion paper.

1 Introduction

Ozone (O_3) is one of the most important atmospheric trace gases as it protects the biosphere from the harmful solar UV radiation and is an important greenhouse gas (United Nations Environment Programme). Global monitoring by ground based and satellite borne instruments plays a unique role in the determination of long-term trends, which is a prerequisite to study air quality, climate chemistry feedback and ozone recovery in response to the Montreal Protocol phasing out ozone depletion substances (Harris et al., 1997).

Nowadays, remote sensing of ozone is done in different spectral regions, mostly in the UV, visible, and thermal IR. The O_3 absorption spectrum in the region 200–1100 nm consists of four absorption bands (Hartley, Huggins, Chappuis, and Wulf). Quantum mechanical dynamics calculations on potential energy surfaces explain the principal features of the O_3 spectrum in the UV/visible/NIR wavelength regions (Banichevich

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et al., 1993; Grebenshchikov et al., 2007; Schinke and McBane, 2010). The assignment of absorption bands from nine electronic states was very comprehensively illustrated using one-dimensional cuts through the potential energy surfaces (Grebenshchikov et al., 2007). However, because of the relatively high density of low-lying states and coupling effects, the quality of the ab initio calculations is still inferior to experimental data accuracy. Results of numerous experimental studies can be found, for example, in the online spectral database of gaseous molecules of the Max-Planck Institute for Chemistry in Mainz (MPI) (Keller-Rudek and Moortgat, 2013), which contains data obtained by more than forty teams covering different spectral regions and temperatures.

The requirement to measure small changes in stratospheric and tropospheric O₃ places strong demands on the accuracy of the O₃ absorption cross-sections and line parameters used in retrievals from remote sensing spectrometers. Besides, accurate measurements are very useful for the assessment of the structure, potential energy diagram, and electronic states of ozone. The following requirements can be postulated:

- spectral resolution of the cross-sections should be at least one order of magnitude better than the instrumental bandwidths of modern remote sensing spectrometers;
- the spectral region measured needs to be sufficiently wide to include spectral channels and windows of as many instruments as possible to provide consistency between the various retrievals;
- data should be available for an adequate temperature range corresponding to that observed in the atmosphere.

The study reported here has addressed these demands. Our dataset provides broad spectral coverage from 213 to 1100 nm with high spectral resolution of 0.02–0.06 nm in the UV-visible (vis) and 0.12–0.24 nm in the near IR (NIR) at 11 temperatures down to 193 K.

Criteria for the choice of a preferred absorption cross-section dataset have been a subject of discussions within the O₃ observing community. A large-scale initiative

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to review and recommend O₃ cross-sections for all commonly used atmospheric remote sounding instruments was started in spring 2009. The ACSO (Absorption Cross Sections of Ozone) committee was established by the Scientific Advisory Group of the Global Atmosphere Watch of the World Meteorological Organization (WMO-GAW) and the International Ozone Commission of the International Association of Meteorology and Atmospheric Sciences (IO3C-IAMAS).

A short summary of several published datasets that are currently used in modern remote sensing applications and are used in our comparisons here is given in the following paragraphs.

The high-resolution broadband data obtained by Bass and Paur (abbreviated as BP) (Bass and Paur, 1984; Paur and Bass, 1984) are used in the standard O₃ total column and profile retrievals using ground-based spectrometers: Brewer (Komhyr and Evans) and Dobson (Scarnato et al., 2009) and satellite spectrometers: SBUV (NOAA, 2007), TOMS (McPeters et al., 1998). These data are also included in the latest version of the high-resolution transmission molecular absorption database HITRAN 2008 (Rothmann et al., 2009); however, a wavelength shift must be applied to obtain optimum results in the Huggins band. This dataset is limited to the spectral region 245–343 nm and the lowest temperature available is 203 K.

The high-resolution broadband data obtained by Brion, Malicet, Daumont (abbreviated as BMD) (Brion et al., 1993, 1998; Daumont et al., 1992; Malicet et al., 1995) are available for the spectral range 195–830 nm at room temperature and for 194.5–520 nm for several lower temperatures down to 218 K. These data are used, for example, in the OMI retrievals (Veefkind and de Haan, 2002) and the most recent SBUV version 8.6 (Bhartia et al., 2012).

The high-resolution temperature-dependent broadband cross-sections obtained previously at the University Bremen team using a Fourier-transform spectrometer in the range 230–830 nm were recorded using total sample pressures of 100 and 1000 mbar (Voigt et al., 2001) at five temperatures.

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The low-resolution broadband cross-sections obtained by the Bremen team using original satellite flight models (FM) spectrometers: SCIAMACHY, GOME, GOME-2 FM3, and GOME-2 FM21 (Bogumil et al., 2003; Burrows et al., 1999a; Chehade et al., 2012) provided pre-flight information about the performance of the instruments. In addition, they are unique source of reference spectral data otherwise not available. These datasets have the advantage that prior knowledge of instrumental slit functions is not needed if used in the ozone retrieval with the same flight-model. However, transformation of these datasets for use with other instruments, having similar but different instrument functions, is not straightforward. Currently the Bogumil et al. (2003) dataset is used for retrievals of the SAGE II, SAGE III (McCormick et al., 1989) and OSIRIS (OSIRIS, 2012) spectra in addition to SCIAMACHY. Weber et al. (2011) considered the impact of absorption cross-section choice on total O₃ retrieval applied to GOME, SCIAMACHY, and GOME2 and discussed the issues associated with resolution matching, wavelength shifts, and scalings. As a consequence the ozone cross-section data for GOME2 FM3 and SCIAMACHY have been adjusted as reported by Chehade et al. (2012, 2013).

Measurements have been performed at single wavelengths in the absorption minimum between the Huggins and Chappuis bands and in the Wulf band by Axson et al. (2011), Anderson and Mauersberger (1992), Enami et al. (2004) and El Helou et al. (2005). These provide additional information for spectral regions, where existing broadband datasets demonstrate strong disagreement or only a single broadband measurement is available.

The ozone absorption at temperature of 296 ± 3 K is the most investigated case, thus a lot of high quality data are available nowadays making a comprehensive analysis possible over a wide spectral region. Data quality assessment is tightly bound to the feasibility of the methods for comparison of datasets characterized by different spectral resolution. In the Sect. 2 of this paper, we consider the experimental techniques for obtaining broadband absolute cross-sections at temperatures ranging from 193 to 293 K. We critically discuss both uncertainties related to the current work and other datasets

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that are currently available. In Sect. 3, we review these methods for every band in UV-vis-IR and provide details on our data analysis and comparisons of this work with published cross-sections at room temperature (293 K). In Sect. 4, we summarize results of our measurements and analysis. In the companion paper (Serdyuchenko et al., 2013), we report in greater detail on the temperature dependence of our new cross-sections.

2 Experimental methods and instruments

2.1 Experimental setup

According to the Beer–Lambert law, a cell containing a characteristic absorber with cross-section $\sigma(\lambda)$ exponentially attenuates initial light intensity I_0 at wavelength λ . Thus, the magnitude of the absorption cross-sections σ is directly proportional to the optical density d , which is given by the natural logarithm of the ratio of initial light intensity to the attenuated intensity:

$$\sigma(\lambda) = d \cdot kT / pL = \ln(I_0/I) \cdot kT / pL, \quad (1)$$

where p is the absorbing gas partial pressure, T is the gas temperature, L is the absorption path length and k is the Boltzmann constant. The light intensity must be weak enough to avoid perturbations in the population of electronic vibrational-rotational states.

The values of the O₃ cross-sections values in UV/visible/NIR cover more than seven orders of magnitude, while the dynamic range of a single measurement typically allows optical density changes of only about one order of magnitude to be measured. In this study, to achieve the required dynamic range, seven or more absorption measurements at different experimental conditions were combined for every temperature.

Experiments were carried out using two absorption cells (Fig. 1), which are cylindrical double-jacketed quartz vessels having a total length of 140 cm and an inner diameter of 5 cm, sealed with evacuated quartz windows at both ends. There is a short

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path across the middle of one of the cells for measurements of strong absorption in the UV. In addition, this cell utilizes White-type multipass optics (White, 1942) to increase the absorption path length for measuring the weak absorption near 380 nm. The cells were designed to achieve diverse combinations of the following parameters: O₃ concentrations from 1 to almost 100 %, total gas pressure of 50–900 mbar, and absorption lengths of 5, 135, 270, and about 2000 cm.

A layer of flowing ethanol, which is insulated from the ambient air by the evacuated outer jacket, surrounds the inner volume of the cells. The outer surface of the cells is wrapped in aluminum foil and insulating foam. Stable cell temperatures from 293 K down to 193 K were maintained by special thermostat systems (Haake CT90 W Phoenix 2 cryostats). The temperature was measured using internal cryostat sensors and external Pt-100 sensors, the latter being installed inside the cells at both ends (not shown in Fig. 1).

Gas supply and vacuum systems (omitted in Fig. 1 to avoid overload of the schematic diagrams) include the rotary vane pumps connected to the cells and several instruments measuring pressure and flow rate (MKS Baratron). The cells were pumped down to pressures below 0.01 mbar before filling with gas. The gas pressure in the ozone generators and the whole oxygen supply system were kept above ambient pressure in order to avoid oxygen dilution in the event of leaks. Ozone-oxygen mixtures were prepared from pure oxygen, by flowing it through an ozone generator, which is in principle a coaxial cylindrical capacitor where electrical field accelerates electrons and causes the dissociation of oxygen molecules. Two types of ozone generators with efficiency up to 10 % and a production rate controllable with an accuracy of about 10 % provided the desired ozone concentration.

For the absolute calibration of the relative optical density spectra, we have chosen a method based on pure O₃ pressure measurements as the most independent and straightforward technique. We performed absolute broadband measurements in spectral regions of relatively large cross-sections in the Huggins and Chappuis bands, using absorption path lengths of 135 and 270 cm, respectively. In this case, we achieved suf-

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ficiently strong absorption with a pure O₃ pressure of about 50 mbar. Pure O₃ was extracted from the ozone-oxygen mixture flow through a quartz trap immersed in a Dewar with liquid nitrogen. Due to the difference of boiling temperatures, oxygen was easily pumped out from the trap while O₃ was condensed on the bottom of the trap. After sufficient O₃ was collected, the trap was warmed up, and evaporated O₃ was fed into the evacuated cells through the Teflon tube of about 1 m length. During all experiments, the leak rate in the cell was below 0.04 mbar h⁻¹.

Due to the instability ozone molecules, absolute measurements were carried out under very careful control of temperature and pressure. Ozone decay during absolute measurements can be taken into account by means of various methods. Two different spectra (*I*₁, *I*₂) can be used, recorded over a short time within which O₃ decay can be assumed negligible and separated by a time interval needed to obtain a measurable O₃ pressure difference *p*₁ – *p*₂ (Griggs, 1968; Molina and Molina, 1986; Yoshino, 1988):

$$\sigma(I_1, I_2, p_1, p_2, k, T, L) = \ln(I_2/I_1) \cdot kT / (p_1 - p_2)L. \quad (2)$$

This way, the absolute value of the absorption cross-section depends only on the pressure difference observed in the cell. However, only very limited number of accumulated spectra and poor signal-to-noise ratio are possible at temperatures above 243 K because of rapid O₃ decay. For temperatures below 243 K, O₃ decay slows down and a measurable pressure difference is achieved only after a fairly long time (more than four hours). Unfortunately such long measurements suffer from lamp intensity drifts. Using this method, we could only achieve an insufficient signal-to-noise ratio.

Therefore, similar to other studies (e.g. El Helou et al., 2005) the actual O₃ pressure *p*(*t*) in our leak-tight cell was obtained by observation of the change in the total pressure *p*_{total}(*t*), i.e. the pressure change Δ*p* after the initial O₃ pressure *p*_{*i*}, the latter measured immediately after the cell was filled:

$$p(t) = p_i - 2\Delta p = 3p_i - 2p_{\text{total}}(t). \quad (3)$$

Acquisition of background spectra, defined as having negligible O₃ absorption, and absorption spectra of O₃ (*I*₀ and *I*, respectively) were separated by the time needed

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(i) to fill the evacuated cell with ozone, typically several minutes, and (ii) to accumulate spectra for sufficient signal-to-noise ratio, typically about 30 min. To control lamp intensity drifts, two background spectra were measured: before filling the cell with ozone and after evacuating the cell. Background spectra were averaged from over about 800 scans, O₃ spectra were recorded in 50 separate files, with each file being an average of 20 scans collected during several minutes. Every file was assigned to the actual O₃ pressure $p(t)$ in the cell, thereby O₃ decay on timescale of several minutes was taken into account:

$$\sigma(t, I, I_0, p, k, T, L) = \ln[I_0/I(t)] \cdot kT/p(t)L. \quad (4)$$

This method yields low-noise spectra, since absorption spectra at a particular pressure are referenced to the stable background spectrum and results are averaged. Unfortunately, the method does not control for any ozone decay prior to start of the measurements. Overestimation of the O₃ concentration would lead to an underestimation of cross-sections. Ozone pressure p_i at the beginning of the measurements was corrected for the amount decayed during filling the cell. Correction was done by means of extrapolation of long time (an hour and more) observations of the total pressure changes in the cell. Depending on the gas temperature in the cell, it was up to several percent. Based on numerous repeated measurements and the quality of extrapolation, we estimated an uncertainty of about 1 % in the knowledge of the initial ozone purity at room temperature.

Some O₃ absorption at 255 nm was observed after the cell was evacuated, which is probably due to an O₃ layer of unknown depth remaining attached to the optical surfaces. This effect could not be neglected during measurements in the Hartley band performed using the short path, therefore, the cross-sections were obtained using oxygen-ozone mixtures in this spectral region (relative measurements). Oxygen-ozone mixtures were also used for weak O₃ absorption measurements in the minimum between Huggins and Chappuis bands and in the Wulf band to avoid risk of explosions. We concatenated the absolute and relative spectra obtained from series of experiments

nal Pt-sensors installed on both ends of the cells and the internal thermostat sensors agreed to within 1 K or better, confirming the uniformity of the temperature in our setups.

The statistical error in optical density arises from the ozone temperature and pressure fluctuations described above and the fluctuation of the light source intensity. The latter is a result of a tradeoff between white noise reduction (lower for long integration) and lamp intensity drift during accumulation and averaging of spectra (lower for short integration). This uncertainty is determined after averaging consequent baseline spectra, recorded without ozone. Obtained standard deviation must be multiplied by square root of 2, because this error contributes both to the baseline and to absorption measurements (I_0 and I). Proper choice of the experiment duration results in uncertainty of the baseline at the level of 0.5 % and better, in best cases down to 0.1 %.

Performance of the optical system (lamps spectra, detector sensitivity curves, etc.) and shape of the absorption spectra lead to a complicated wavelength dependence of the relative uncertainty of the optical density. At identical signal-to-noise ratio, it can change by at least one order of magnitude within a single measurement in parts of the spectrum exhibiting steep slopes, particularly in the Huggins or Wulf bands. Optical densities measured at different spectral regions are given in Table 1 together with lamp stabilities. For simple estimation, we use the statistical uncertainty related to the source stability at optical density of unity. This is valid for almost the entire Hartley and Huggins bands and the 450–750 nm region; however, it strongly underestimates the error bars in the spectral regions with weak absorption between Huggins and Chappuis bands and in the NIR Wulf band. One should keep in mind that quality of the system performance and signal-to-noise ratio typically fall towards edges of the measured spectra pieces.

Table 2 summarizes the errors from different sources for the O₃ cross-sections absolutely measured in the Huggins band and Chappuis band. The systematic uncertainty is below 1 %; the statistical uncertainty at optical density of 1 is 0.2–2 % in UV and in the visible. In the spectral regions characterized by weak absorption, statistical uncertainty can reach up to 20 % and more due to the very low optical density (Table 1).

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The total uncertainty of the absolute measurements propagates into the relative measurements. For relative measurements, additional uncertainties arise from the scaling of these spectra to the absolute calibrated spectra. However, these uncertainties are negligible because differences between the spectra matched during single concatenation in the overlap region are typically much smaller than the estimated experimental uncertainty.

Since information on the total uncertainty is particularly important when comparing different datasets, we shall briefly discuss uncertainties of some published broadband data.

Bass and Paur report 1 % noise during their measurements (Bass and Paur, 1985; Paur and Bass, 1985). This random error needs to be added to the systematic error arising from their absolute calibration. The data were absolutely calibrated for all temperatures using Hearn's value at the Hg lamp line at 253.65 nm, which is $1147 \times 10^{-20} \text{ cm}^2 \text{ molecule}^{-1}$ (Hearn, 1961). Hearn himself states best estimates of the cross-section at 253.65 nm as $(1147 \pm 24) \times 10^{-20} \text{ cm}^2 \text{ molecule}^{-1}$, i.e. 2.1 %. Therefore, the total uncertainty of BP data exceeds 2 %. At other wavelengths of Hg lamp lines in the Hartley band, BP data agree with Hearn's data within 2–3 %.

A systematic uncertainty of at least 2 % in the SCIAMACHY dataset by Bogumil et al. (2003) arises because BP data have been used for absolute calibration. The Bogumil et al. (200) estimated a total uncertainty of about 3 % or better, excluding regions with cross-sections below $10^{-23} \text{ cm}^2 \text{ molecule}^{-1}$ (365–410 nm and longer than 950 nm) and the 305–320 nm spectral region (spectral channel edges). This total uncertainty includes an error arising from the scaling procedure during concatenation. The Bogumil et al. (2003) dataset was obtained using concatenation of several spectra performed from UV to NIR covering a broad region from 230 to 1075 nm. Consecutive concatenations starting from the Hartley band towards longer wavelengths can lead to an accumulated error towards the Chappuis and Wulf bands, because BP data are only available in the Hartley–Huggins bands.

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The other satellite dataset – GOME flight model data – was absolutely calibrated using the titration method (Burrows et al., 1999a). The accuracy stated is 2.6 % with less than 2 % assigned to the lamp intensity drift. Integrated absorption cross-sections recorded with the GOME-FM instrument were used for the absolute calibration of the FTS dataset by Voigt et al. (2001). The spectral wavelength accuracy of the Voigt et al. (2001) data is better than 0.5 pm at 230 nm and 7.2 pm at 850 nm.

BMD team states a total systematic error of 1.5 % in the spectral region 420–830 nm and 1.5–4 % at 350–420 nm and a statistical error (RMS) of 0.9–2 % (Brion et al., 1998). In the earlier works in the UV region BMD team states a total systematic error of 1.3 % in the Hartley band and 1.3–2.5 % in the Huggins band and a statistical error of 0.9–2.2 % (Daumont et al., 1992).

We conclude that accuracy of our data set (2–3 % for most of the spectral regions) is comparable to other broadband data measured so far. In ozone retrievals from remote sensing, an additional error may arise from data convolution with the instrumental slit functions and, therefore, depends on the accurate knowledge of those.

3 Results and analysis

The new room-temperature ozone absorption cross-sections in the entire measured region are shown in Fig. 3 together with the high-resolution BMD dataset (Brion et al., 1998) and lower resolution satellite spectrometer datasets obtained by Burrows et al. (1999a) and Bogumil et al. (2003). Figure 3 also contains cross-section values at single wavelengths obtained by El Helou et al. (2005), Axson et al. (2011), and Anderson et al. (1993).

The red rectangle denotes the spectral window (325–335 nm) used by retrieval algorithms based on the Differential Optical Absorption Spectroscopy (DOAS) in the Huggins band (Weber et al., 2011). The vertical lines schematically mark the single wavelengths used in retrieval algorithms of the spectra obtained by different ground based and remote sensing instruments for detection of O₃, trace gases, clouds and

aerosols (SBUV, OMI, TOMS, OSIRIS, SAGE, Brewer, Dobson). Despite the fact that most channels are present in the UV spectral region, there are important channels in visible and IR as well.

Review by Orphal (2003) provided analysis on broadband data obtained before 2003.

Following the similar approach, we analyzed the new absorption cross-sections at the temperature of 293 K:

- scaling factors between our new data and selected published band-integrated cross-sections are determined;
- a comparison between our results with the published datasets at wavelengths of the He-Ne laser, Hg lamp, and spectral channels of various remote sensing instruments is made;
- scaling factors and wavelength shifts are determined with respect to the high spectral resolution datasets (BMD and BP) and the lower spectral resolution datasets (GOME, GOME-2, SCIAMACHY) in the Hartley, Huggins, and Chappuis bands.

3.1 Band-integrated cross-sections

The integrated absorption cross-sections over an electronic band are only dependent on the number of molecules in the lower state and the transition probability (Burrows et al., 1999a). Absorption cross-section of ozone rapidly decreases away from the band maximum; therefore integrated band intensity is weighted to the largest values of the absorption cross-sections. Consequently, the comparisons of integrated cross-sections reflect mainly the knowledge of the cross-sections close to the maximum of the bands. In addition, the spectral bands overlap (e.g. the high-energy tail of the Huggins band continuously goes over into the Hartley band system). Nevertheless, the method is a valuable test of agreement between broadband absorption cross-sections datasets having different spectral resolution.

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Orphal (2003) proposed integration limits which were found from pragmatic considerations such as the quality of the considered datasets rather than the real band borders, determined by theoretical studies (Grebenshchikov et al., 2007; Banichevich et al., 1993). For comparison purposes, in this work we used the same limits. In the Wulf band, we calculated integrated absorption cross-sections for our new data and Bogumil et al. (2003) data using the range 663–1000 nm (Banichevich et al., 1993).

Table 3 presents the integrated absorption cross-sections obtained from this work, the high spectral resolution BMD dataset, the different satellite instrument datasets and the mean values obtained by Orphal from averaging several data sources (Bass and Paur, 1985; Brion et al., 1998; Voigt et al., 2001; Bogumil et al., 2003; Burrows et al., 1999a). The uncertainty of this mean value is the standard deviation from averaging. Our reported absorption cross-sections agree with all datasets within 1 % in the Hartley and Huggins bands, within 1–4 and 6 % in the Chappuis and Wulf bands. Agreement with high resolution BMD dataset is about 1 % and better.

3.2 Hartley and Huggins bands

The Hartley and Huggins bands (Fig. 4) are especially important for remote sensing of the atmosphere, since they contain spectral channels used by both satellites and ground-based spectrometers; several satellite spectrometers work with the entire Huggins band.

3.2.1 Spectral region near 255 nm

Several datasets in the region around 255 nm are shown in Fig. 4 (top panel). The uncertainty reported by Hearn is about 2 %. In general, our new absorption cross-sections are lower by several percent compared to satellites and BP datasets. Agreement with the BMD dataset is within 1 %. In Table 4, our new data are compared to some published data and the mean values by Orphal (2003) at selected wavelengths, corresponding to Hg lamp and He-Ne lines. Some of the data were interpolated. As

coefficient in the 323–343 nm region. Assuming a linear shift, Orphal reports that BMD is shifted towards shorter wavelengths (-0.014 nm) compared to the Voigt et al. (2001) data.

We compared our data with the published datasets, following a similar approach.

One should take into account that nowadays several slightly different versions of the datasets became available. Below is a short reference and notations, used in our study. The MPI online database (Keller-Rudek and Moortgat) contains experimental BP and BMD datasets for different temperatures at air wavelengths (BP_{exp} and BMD_{exp}) obtained by personal communication in 2000 in 1998, respectively. The same experimental BMD data can be found on the ACSO homepage (http://igaco-o3.fmi.fi/ACSO/cross_sections.html). The ACSO homepage also provides polynomial coefficients (quadratic fits) that have been calculated by Bass and Paur for the temperature dependence of the original data excluding the 218 K at air wavelengths (BP_{calc}). HITRAN 2008 also contains polynomial coefficients (BP_{HITRAN}); however, a few corrections were made on the original BP data: air/vacuum wavelength conversion according to Edlen's equation (Edlen, 1966) and wavelength shift. Using original BMD data at four temperatures (218, 228, 243, and 295 K), Liu et al. (2007) derived polynomial coefficients for temperature dependence for vacuum wavelength scale (BMD_{calc}) making possible the interpolation/extrapolation of the cross-sections to any given temperature.

For comparison we used versions of the BP and BMD datasets from the sources mentioned above as well as Voigt et al. (2001) and Bogumil et al. (2003) datasets from the homepage of the Molecular spectroscopy group of the Institute of Environmental Physics in Bremen, Germany. To avoid altering the original data, when possible, we performed comparisons with high-resolution BP, BMD and Voigt et al. (2001) datasets without resolution matching. For the comparison with Bogumil et al. (2003) dataset, the new high-resolution dataset was convolved with the Gaussian profile with FWHM of 0.2 nm. All datasets were compared on the same grid of vacuum wavelengths in steps of 0.01 nm using Edlen's equation and interpolation.

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The relative difference $\delta\sigma$ between datasets σ_1 and σ_2 was calculated in percent as a function of a scaling factor F and a linear wavelength shift $\delta\lambda$:

$$\delta\sigma(\lambda) = [\sigma_1(\lambda) - F \cdot \sigma_2(\lambda + \delta\lambda)] / [F \cdot \sigma_2(\lambda + \delta\lambda)] \cdot 100\% \quad (5)$$

To check the linearity of the shifts and scalings, we performed the analysis for three different spectral regions: 323–330, 332–340 and 323–340 nm and averaged values to obtain uncertainties. For every spectral region, we calculated the mean over the region differences $\Delta\sigma$ between different pairs of datasets and found the optimum shifts and scaling factors by minimization of these mean differences. In comparisons with our new data, shifts and scaling were applied to the published datasets.

Comparison of the original experimental datasets may be controversial because “minima” in absorption features are sensitive to temperature. Even relatively small difference of ± 5 K influences the absolute values. Our absorption cross-sections and those of Voigt et al. (2001) and Bogumil et al. (2003) were obtained at 293 K, whereas those of BP_{exp} and BMD_{exp} were made at 298 and 295 K, respectively. Approximation of the absorption cross-sections by fitting a second order polynomial seems to be a reasonable solution: calculated for 298 and 295 K BP_{calc} and BMD_{calc} reproduce corresponding experimental BP_{exp} and BMD_{exp} data very well; mean differences are 0.3 % for BP and 0.04 % for BMD, scaling factors are 0.998 for BP and 1 for BMD. Using polynomial coefficients to describe the temperature dependence reduces differences between the datasets, also due to their smoothing effect.

To test capabilities of our technique, we reproduced some results of the analysis done by Orphal. We obtained similar scaling (0.96) and shift ($+0.03 \pm 0.005$ nm) between BP_{exp} and Voigt et al. (2001) data. We found that matching BP_{calc} calculated for 293 K to the Voigt et al. (2001) dataset requires nearly negligible scaling (0.998), but requires shift of $+0.027 \pm 0.004$ nm. Our approach is also capable to reveal the shift of $+0.015 \pm 0.0005$ nm applied to the experimental BP dataset in the HITRAN 2008 version.

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Mean relative differences, shifts, and scaling factors, which match the published datasets to the new data, are shown in Fig. 5a. The relative differences between our absorption cross-sections and other relevant datasets are shown in Fig. 5b. Our absorption cross-sections are slightly lower than all other datasets, but this difference is within the experimental error bars for most of the cases. Among experimental data, our new absorption cross-sections agree best with the Voigt et al. (2001) and Bogumil et al. (2003) datasets (Fig. 5a): scaling factors are 0.99 and 0.995 and mean difference is about 1%. We observe a small wavelength shift of 0.005 ± 0.002 nm between Voigt et al. (2001) and the new data, which exceeds the reported accuracy of the spectral calibration of the FTS spectrometer (0.001 nm). Both measurements were performed under similar conditions using the same instrument in the Bremen laboratory; however, the spectral resolution of the Voigt et al. (2001) data is 5 cm^{-1} , whereas that of the new data is 1 cm^{-1} . The shift between Bogumil et al. (2003) data and the new data is 0.006 ± 0.004 nm. We found that BP_{calc} and $\text{BP}_{\text{HITRAN}}$ datasets agree with our new data within experimental uncertainty and show linear shifts of $+0.023 \pm 0.003$ nm and $+0.008 \pm 0.003$ nm, respectively; shift of the BP_{exp} is $+0.026 \pm 0.002$ nm, which means that original BP data must be shifted towards longer wavelengths to match our new measurements.

We also observe a non-linear wavelength calibration of the BMD dataset in the region 323–340 nm: different shifts were found for 323–330 and 332–340 nm regions. Assuming a mean linear shift, BMD_{calc} is shifted towards longer wavelengths relative to BP, Bogumil et al. (2003) and the new datasets, while absolute values of cross-sections agree within experimental uncertainty (excluding BP_{exp}).

Possible biases between the original datasets in Fig. 5b are masked by deviations of up to 5%. Scaling and wavelength shifts reduce the relative difference to 0.5–2%. The remaining differences can be attributed to noise in the measurements, non-linear wavelength differences, differences in the instrumental slit functions, and slight resolution mismatch between the considered datasets.



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Apart from the analysis in the DOAS window, we compared our dataset with the published datasets at the selected single wavelengths used by different ozone-observing instruments in UV region. We used BP_{calc} and BMD_{calc} data calculated for 293 K using polynomial coefficients and Bogumil et al. (2003) data. Exhaustive comparison involves convolution of the datasets with instruments slit functions, which are not always well documented. For resolution matching purposes, we convolved our new dataset, Bogumil et al. (2003), BMD_{calc} and BP_{calc} datasets to 1 nm using a rectangular slit function for comparing at SBUV, TOMS and Dobson wavelengths and to 0.4 nm for comparison with OMI and Brewer wavelengths. In the pair comparison (new – BMD_{calc} , new – BP_{calc} , and $BP_{\text{calc}} - BMD_{\text{calc}}$) we found that BMD, BP and the new data agree within the accuracy of the measurements (2%) for wavelengths below 320 nm. In general, the new data are close to the BMD dataset, but slightly lower. At 320–340 nm, our new dataset is lower by 2–3 and 2–7% compared to BMD and BP, respectively.

3.3 Chappuis and Wulf bands

Satellites retrievals use the Chappuis and Wulf bands and adjoining Huggins and Chappuis bands for the ozone retrieval in limb and occultation experiments. In addition, accurate O_3 absorption cross-sections in this spectral region are necessary for the retrieval of other trace gases, aerosols and clouds. The relative differences between the new absorption cross-sections and published data are shown in Fig. 6 and discussed below.

3.3.1 Absorption region 350–450 nm

Various broadband datasets show large disagreement in the absorption minimum between the Huggins and Chappuis bands at 350–425 nm and especially near 380 nm as shown in Figs. 6 and 7 (top panel) and in Table 5. Data given in Table 5 were obtained at different spectral resolution as indicated by the numbers in brackets.

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Additional data are available from measurements at single wavelengths recently performed by Axson et al. (2011) using incoherent broadband cavity enhanced absorption spectroscopy. These relative measurements scale directly with the choice of reference cross-section value at 253.65 nm. For absolute calibration, Axson et al. (2011) took the reference value reported by Orphal (Table 5). The reported accuracy of the Axson et al. cross-sections is 4–30 %, with the greatest uncertainty near the minimum absorption.

Our new absorption cross section dataset agrees well with the BMD dataset and shows good agreement with data of Axson et al. (2011) within accuracy limits. There is a significant disagreement between the new dataset and Bogumil et al. (2003) data in the region of minimum absorption around 385 nm and in the IR region at 1020 nm. In both regions, uncertainty of measurements is greater than 30 % as a result of the very weak absorption measured.

3.3.2 Visible region 450–700 nm

Several ozone absorption cross-section datasets are available in the visible region near the maximum of the Chappuis band (Fig. 7, bottom panel). Absolute calibrated data of (Anderson and Mauersberger, 1992) were obtained using a tandem dual-beam spectrometer and pure ozone from the liquid nitrogen trap; reported overall uncertainty of these measurements is below 1 %. Very recent absolute measurements performed by El Helou et al. (2005) using a FT spectrometer were calibrated using pure ozone as well, and have a reported uncertainty of 0.82 %.

The values of cross-sections at the selected wavelengths from the different datasets are listed in Table 6 (in some cases data were interpolated). The relative difference in percent between the new data and some published datasets is shown in Fig. 6.

In general, our new absorption cross section measurements agree well with those obtained by Anderson and Mauersberger (1992) and El Helou et al. (2005) at the selected wavelengths (within about 1 % or better). Our new data agree within 1 % or better for the broad region 510–680 nm with those reported by Burkholder and Taluk-

dar (1994), which were absolutely scaled using the data of Anderson and Mauersberger (1992). Agreement with BMD is also within the accuracy limits. Difference between new data and satellite datasets is about 2–3%, which is almost within their experimental accuracy. W) might have had an impact on accuracy of the absolute calibration of this dataset in this spectral region.

3.3.3 NIR region 700–1100 nm

Some of the datasets available in the near IR region are shown in Fig. 8. Data on selected wavelengths are compared in Table 7. Data of Anderson et al. (1993) obtained with a tandem dual-beam spectrometer were absolutely calibrated using the absolute cross section reported by (Anderson and Mauersberger, 1992). In addition to the El Helou et al. data, there are measurements performed by Enami et al. (2004) at two single wavelengths at 760 nm using cavity ring-down technique (Fig. 8a). The absorption cross-sections of Enami et al. were calibrated using the Hearn's value in the UV with the estimated total uncertainty of the O₃ concentration of $\pm 1.5\%$ and the typical overall uncertainty is about $\pm 2\%$. However, accuracy of the Hearn's value is not included in this error budget.

This region contains an absorption band of oxygen, which is visible, for example, on the Bogumil et al. (2003) dataset, measured in an oxygen/ozone mixture. No traces of oxygen spectrum are visible on the new dataset measured with pure ozone.

At wavelengths longer than 1050 nm ozone absorption decreases rapidly. Because of the experimental limitations, measured absorption in this region was weak and sensitive to the baseline stability. The uncertainty of our measurements around 1000 nm is about 30% and more. Only a few studies on O₃ absorption cross-sections are available in this region. Results from Anderson et al. (1993) and El Helou et al. (2005) show agreement with the new dataset within 5% at wavelengths below 1000 nm and up to 20% around 1040 nm (Figs. 6 and 8b). Compared with Bogumil et al. (2003), our new dataset better resolves the rotational structure of the Wulf band. In general, the new dataset in the “valleys” near 970 and 1010–1040 nm are lower than the Bogu-

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mil et al. (2003) dataset, partly because of the different spectral resolution. Further measurements in this region are of high interest and promising for both O₃ rotational spectra modeling and for use in satellite missions, for example for SAGE II and SAGE III (SAGE III ATBD Team, 2002).

4 Conclusions

The experimental setups with two types of spectrometers (echelle and FTS) made it possible to obtain new broadband ozone absorption cross-sections covering the spectral region from the UV to the NIR, which contains spectral channels of numerous remote sensing instruments. A spectral resolution of at least one order of magnitude better than that of most remote sensing instruments was achieved. Absolute cross-sections were measured using pure ozone in the absorption cell. Special attention was paid to ozone decay before and during the absorption measurements for which corrections were introduced.

Typically, relative datasets are scaled using single reference values in UV region, for example, at 253 nm (Hg emission line). In our work, absolute cross-sections were obtained for broad regions in the UV and visible parts of spectra in the Huggins and Chappuis bands. Relative measurements in the Hartley band, in the absorption minimum region between Huggins and Chappuis bands, and in the Wulf band were concatenated by scaling overlap regions to the absolute cross-sections. Thus, we could avoid error accumulation due to consecutive concatenations across a large spectral range up to the near IR.

From our experience with this work and by evaluating the error budgets of similar measurements reported in the literature, we conclude that 2–3 % accuracy is the best that can be achieved in broadband cross-section measurements including ours so far. Single wavelength measurements with higher accuracy can be of great help for improving absolute calibration. However, such measurements must be performed at several

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temperatures and at different wavelengths across the entire spectral range for reliable independent calibration of concatenated spectra.

Our new ozone cross-sections at room temperature are very close to the high-resolution BMD dataset; the agreement further improves if compared with BMD cross-sections interpolated to 293 K using the quadratic temperature dependence. Compared to the low-resolution satellite flight model data, the new data are systematically lower. However, all considered datasets agree within the reported experimental error bars for almost the entire spectral region when using conventional methods for data comparison. Part of the disagreement between datasets can be a result of the spectral resolution mismatch (particularly in the Huggins band), wavelength calibration errors, and temperature dependence as discussed above. Reported uncertainties should not be a crucial factor on the choice of cross-section data for remote sensing applications. Such a choice should be rather based on tests of the data in retrievals. Differences between cross-sections are not necessarily directly affecting retrieved ozone amounts, e.g. when using differential fitting approaches.

The second part of this study (Serdyuchenko et al., 2013) is devoted to the temperature effect in the new data. The new O₃ absorption cross-sections data set, combined with supporting technical documentation are made available from the homepage of the Molecular Spectroscopy laboratory of the Institute of the Environmental Physics, University of Bremen, Germany (<http://www.iup.uni-bremen.de/gruppen/molspec/databases/index.html>).

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Table 1. Experimental setups for different spectral regions.

Spectral region, nm	Spectrometer, detector	Resolution	Calibration	Absorp. path, cm	Lamp, stability %	Optical density
213–310	Echelle, ICCD	0.018 nm	Relative	5	De, 0.5	0.5–2
310–335	FTS, GaP	1 cm ⁻¹	Absolute	135	Xe, 2	0.1–2
335–350	FTS, GaP	1 cm ⁻¹	Relative	270	Xe, 1	0.1–1
350–450	Echelle, ICCD	0.02 nm	Relative	~ 2000	Xe, 1	0.05–1
450–780	FTS, Si	1 cm ⁻¹	Absolute	270	Tungsten, 0.2	0.05–2
780–1100	FTS, Si	2 cm ⁻¹	Relative	270	Tungsten, 0.2	0.001–0.1

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Table 2. Uncertainty in the absorption cross-section obtained from absolute measurements at 50 mbar and 293 K and path lengths of 135 and 270 cm in the Huggins band and Chappuis band, respectively.

Systematic uncertainty	Statistical uncertainty
Ozone impurity: <ul style="list-style-type: none"> – oxygen purity 99.995 % – leaks < 0.1 %. 	Ozone initial pressure: 1 %
Accuracy of pressure sensors: 0.04 % (0.02 mb)	Fluctuations of pressure: < 0.08 % (< 0.04 mb)
Offset of the temperature sensors: < 0.3 % (1 K)	Fluctuations of temperature: < 0.1 % (< 0.3 K)
Temperature non-uniformity in the system: < 0.3 % (1 K)	Light source stability, relative to optical density OD = 1: 0.3/2.8 % (depending on spectral region, see Table 1).
Cell length: < 0.04–0.07 % (< 1 mm)	
Total: < 0.8 %	Total (RMS): 1–3 %, depending on spectral region

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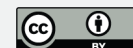
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Table 3. Integrated cross-sections at room temperature (293–298 K), $\text{cm}^2 \text{molecule}^{-1} \times \text{nm}$.

Band, limits, nm		J. Orphal (2003)	<i>This work</i>	BMD (1998)	Bogumil et al. (2003)	Burrows et al. (1999)
Hartley, $\times 10^{-16}$	245–340	$3.55 \pm 0.6 \%$	<i>3.53</i>	3.52	3.56	3.57
Huggins, $\times 10^{-20}$	325–340	$8.30 \pm 0.7 \%$	<i>8.29</i>	8.30	8.31	8.33
Chappuis, $\times 10^{-19}$	410–690	$6.38 \pm 1.6 \%$	<i>6.22</i>	6.29	6.40	6.45
Wulf, $\times 10^{-19}$	633–1000	–	<i>1.04</i>	–	1.10	–

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Table 4. Absolute absorption cross-sections in UV at 295 ± 3 K, $\text{cm}^2 \text{molecule}^{-1} \times 10^{-20}$.

wavelength, nm vacuum (air)	J Orphal (2003)	<i>This work</i>	Hearn (1961)	BMD (1998)	Burrows et al. (1999)	Bogumil et al. (2003)	Petersen et al. (2012) using Hearn/BMD value
244.06	–	946	–	949	964	965	964/950
248.32	–	105.1	–	103.8	106.3	106.3	106.1/104.5
253.73 (253.65)	$1141 \pm 0.9\%$	1120	1147	1131	1150	1147	–
257.34	–	1107	–	1107	1129	1126	1126/1112
289.44 (289.36)	$149 \pm 2.0\%$	151	147	151	151.6	150	–
296.82 (296.73)	$60.3 \pm 1.6\%$	61.1	59.71	61.5	61.4	60.9	–
302.24 (302.15)	$29.2 \pm 1.8\%$	29.6	28.60	29.8	29.9	29.5	–

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Table 5. Absolute absorption cross-sections in minimum absorption region near 380 nm at 295 ± 3 K, $\text{cm}^2 \text{molecule}^{-1} \times 10^{-23}$; typical FWHM, nm is given in brackets.

vacuum wavelength, nm	<i>This work</i>	BMD (1998)	Burows et al. (1999)	Bogumil et al. (2003)	Burkholder et al. (1994)	Axson et al. (2011)
365	4.9 (0.02)	4.74 (0.02)	6.26 (0.17)	5.11 (0.21)	–	3.68 (0.27)
405	1.46 (0.02)	1.47 (0.02)	4.27 (0.17)	2.12 (0.21)		1.51 (0.29)
455	20.6 (0.02)	20.4 (0.02)	22.9 (0.29)	21.4 (0.52)	18.2	18.8 (0.5)

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Table 6. Absolute absorption cross-sections in visible and IR at 295 ± 3 K, $\text{cm}^2 \text{molecule}^{-1} \times 10^{-21}$.

wavelength, nm vacuum (air)	J Orphal, (2003)	<i>This work</i>	BMD (1998)	Burrows et al. (1999)	Bogumil et al. (2003)	Anderson et al. (1992)	Burkholder et al. (1994)	El Helou et al. (2005)
543.67 (543.52)	$3.14 \pm 1.3\%$	<i>3.08</i>	3.12	3.17	3.16	3.075	3.11	3.051
577.12 (576.96)	$4.77 \pm 0.8\%$	<i>4.70</i>	4.77	4.83	4.84	–	4.72	–
594.26 (594.10)	$4.70 \pm 1.2\%$	<i>4.63</i>	4.68	4.76	4.74	4.569	4.66	4.631
604.78 (604.61)	$5.22 \pm 1.0\%$	<i>5.10</i>	5.18	5.24	5.25	5.125	5.12	5.125
612.14 (611.97)	$4.66 \pm 0.7\%$	<i>4.54</i>	4.63	4.69	4.68	4.633	4.52	4.586
632.99 (632.82)	$3.46 \pm 1.2\%$	<i>3.36</i>	3.39	3.51	3.48	3.383	3.37	3.401

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Table 7. Absolute absorption cross-sections in the NIR at 295 ± 3 K, $\text{cm}^2 \text{molecule}^{-1} \times 10^{-22}$.

Vacuum wavelength, nm	<i>This work</i>	BMD (1998)	Burrows et al. (1999)	Bogumil et al. (2003)	Anderson et al. (1993)	Burkholder et al. (1994)	El Helou et al. (2005)
748.721	4.38	4.48	5.19	4.62	–	4.25	4.314
755.21	3.22	3.36	4.04	3.47	3.194	3.26	–
760.21	2.77	2.86	3.52	2.95	2.720	2.80	–
765.21	2.53	2.62	3.33	2.72	2.539	–	–
770.21	2.49	2.58	3.31	2.69	2.509	–	–
779.416	3.15	3.25	3.99	3.37	–	–	3.121
802.22	1.45	1.53	–	1.60	1.462	–	–
817.224	2.20	2.26	–	2.34	2.157	–	2.159
853.234	1.46	–	–	1.56	1.417	–	1.420
877.24	0.377	–	–	0.422	0.368	–	–
889.24	0.510	–	–	0.576	0.511	–	–
898.247	0.638	–	–	0.714	0.620	–	0.621
933.256	0.162	–	–	0.172	0.161	–	–
944.259	0.424	–	–	0.458	0.407	–	0.406
991.841	0.407	–	–	0.376	–	–	0.380
1046.766	0.0773	–	–	0.0723	–	–	0.067

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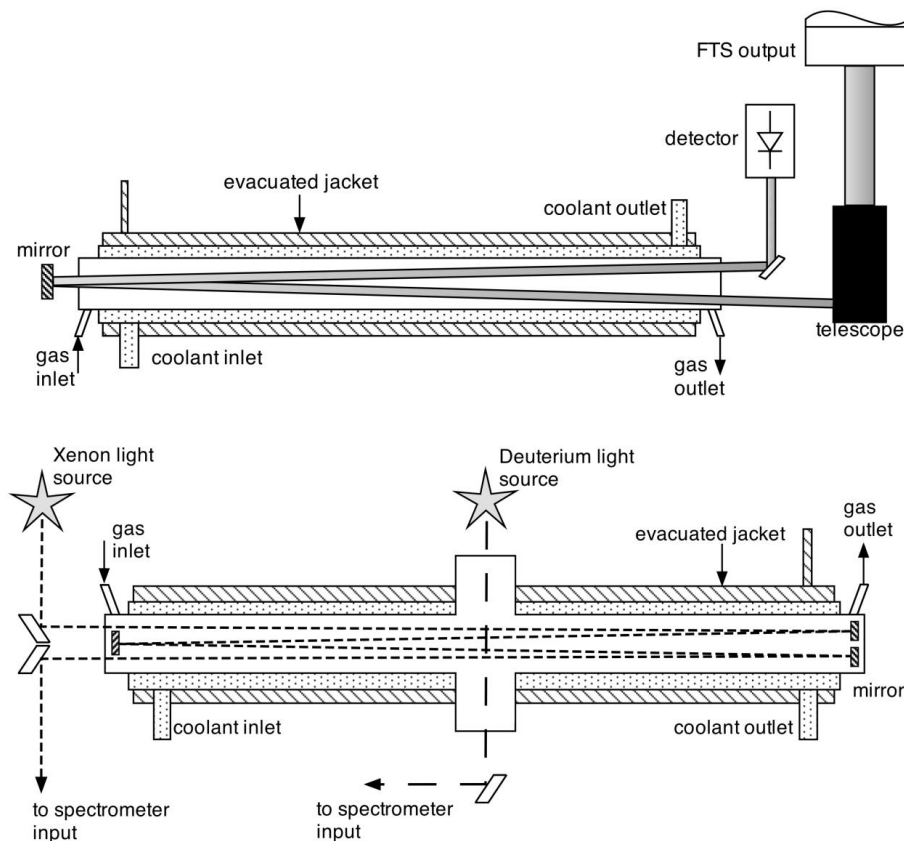


Fig. 1. Experimental setups. Top panel: absorption cell coupled to the FTS. Bottom panel: white absorption cell coupled to the Echelle spectrometer. The shortest path length was achieved using the cross path.

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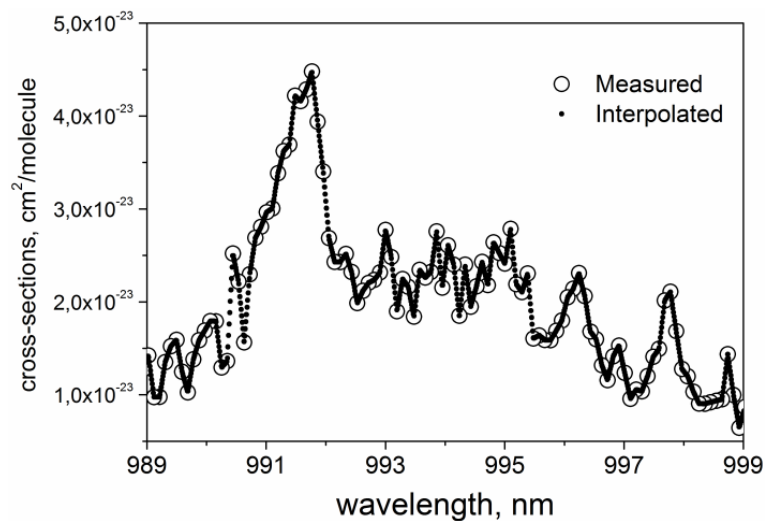


Fig. 2. Final steps in spectra processing: interpolation of the experimental spectra in infrared spectral region.

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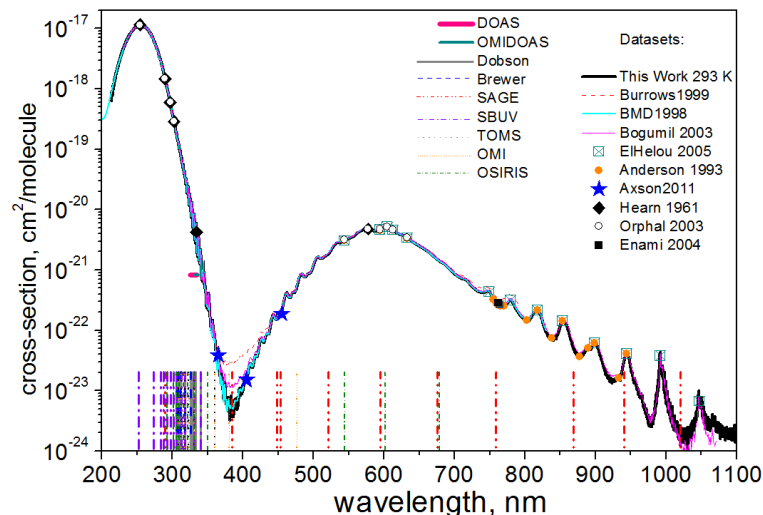


Fig. 3. Ozone absorption cross-section at room temperature. Notations for datasets: black solid line – this work, magenta solid line – Bogumil et al. (2003), light blue solid line – BMD, dash line – Burrows et al. (1999); symbols: cross – El Helou et al. (2005), closed circle – Anderson et al. (1993), star – Axson et al. (2011), open circle – Orphal (2003), diamond – Hearn (1961), closed square – Enami et al. (2004); notations for instrument channels: solid horizontal line – DOAS in Huggins band; vertical lines: grey solid – Dobson, dash – Brewer, dash-dot-dot – SAGE, dash-dot – SBUV, dots – TOMS, short dots – OMI, short dash-dots – OSIRIS.

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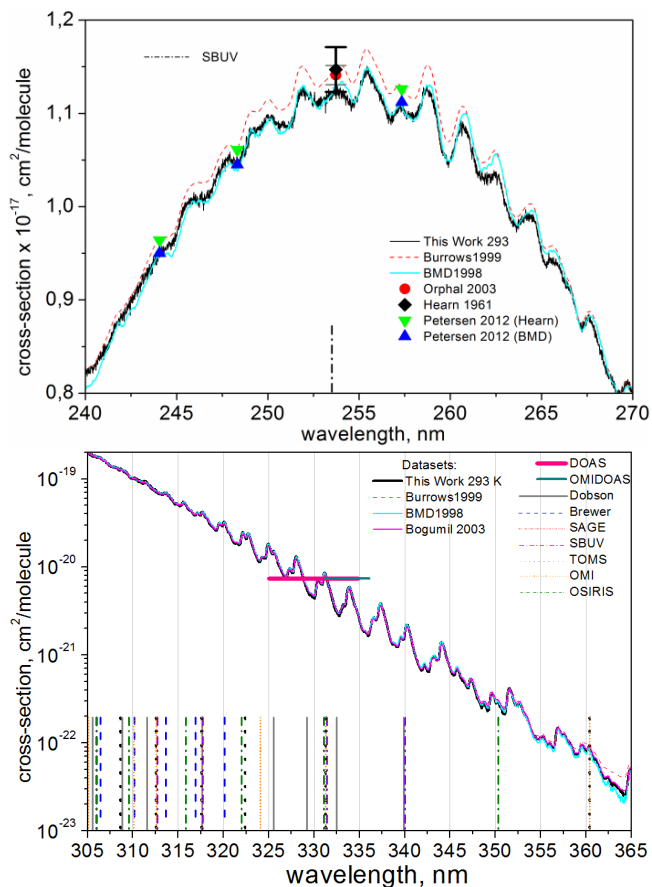


Fig. 4. Ozone absorption cross-section at room temperature in Hartley band (top panel) and Huggins band (bottom panel). Notations for datasets and spectral channels are the same as in Fig. 3. Triangles are data from Petersen et al. (2012).

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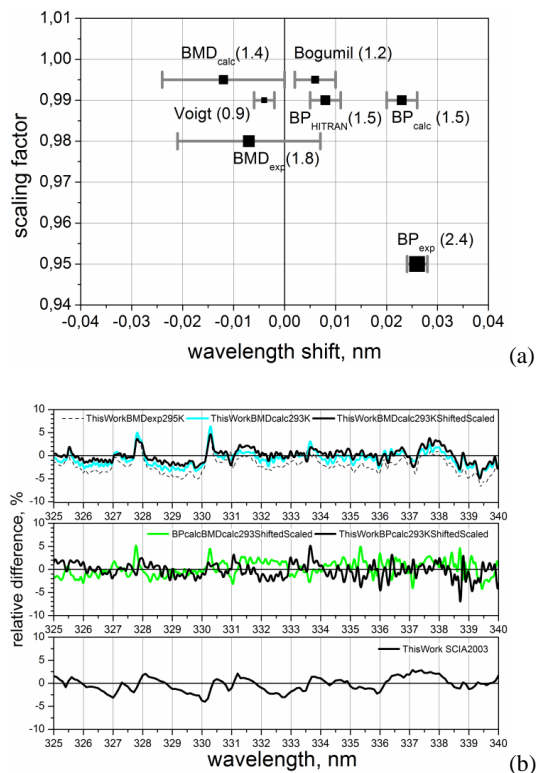


Fig. 5. Comparison of absorption cross-sections in Huggins band: **(a)** scaling factors and wavelength shifts between new data and other datasets, (size of the symbols is proportional to the mean relative difference in percent given in brackets); **(b)** top panel: relative difference between new data and BMD in percent. Dashed line – difference with BMD_{exp}, light line – difference with BMD_{calc}, dark line – difference with BMD_{calc} after wavelength shift and scaling. Middle panel: relative difference between BP_{calc} and BMD_{calc} and BP_{calc} and new data in percent after wavelength shift and scaling. Bottom panel: relative difference between new data and Bogumil et al. (2003) in percent.

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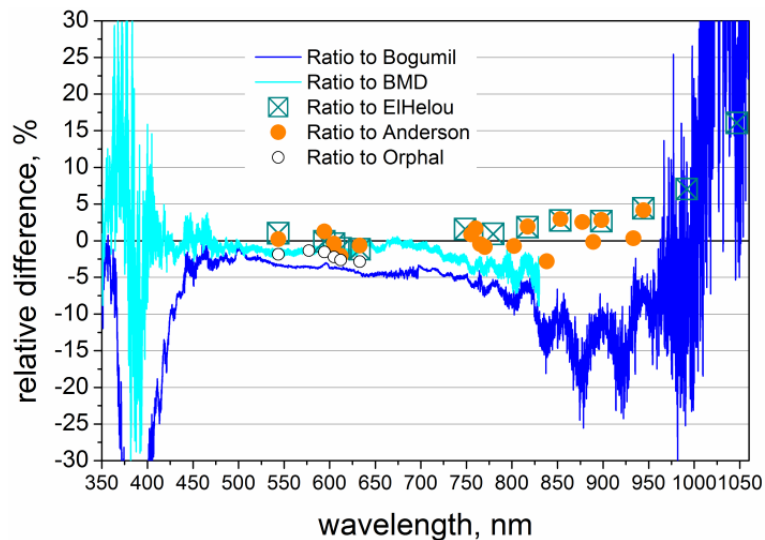


Fig. 6. Relative difference between the new cross-sections and published data in Chappuis and Wulf bands.

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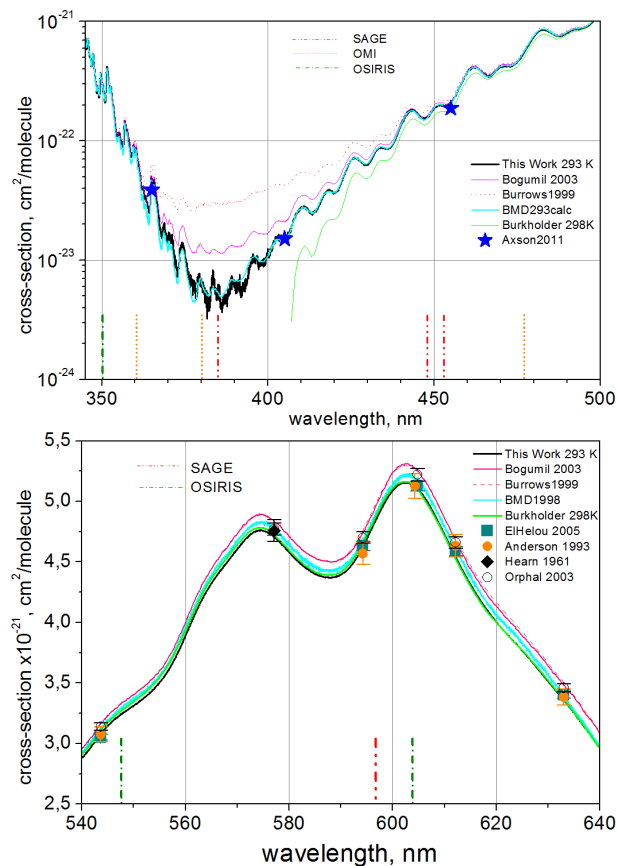


Fig. 7. Ozone absorption cross-sections in the Chappuis band: near absorption minimum at 380 nm (top panel), at the maximum of the Chappuis band (bottom panel). Notations for datasets and spectral channels are the same as in Fig. 3, green solid line – Burkholder and Talukdar (1994).

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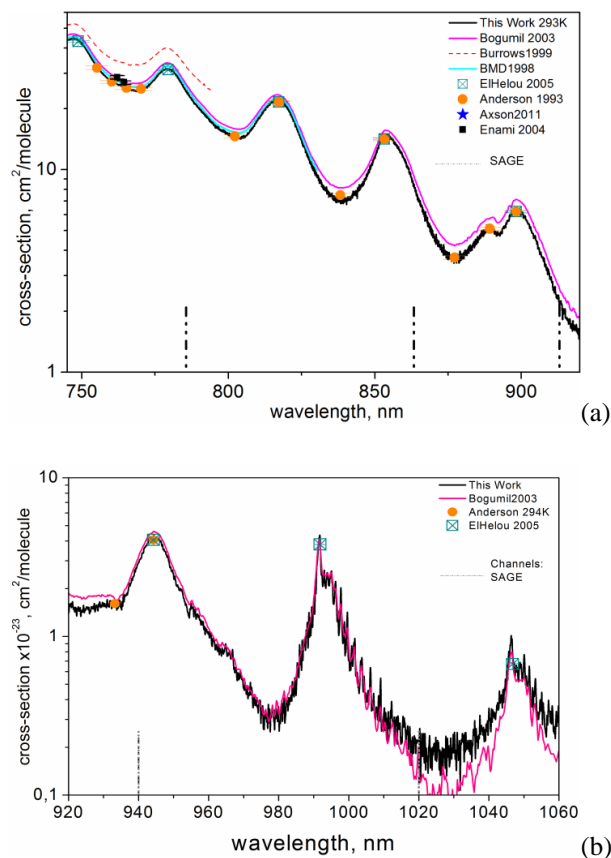


Fig. 8. Ozone absorption cross-sections datasets in the Wulf band: **(a)** 750–900 nm, **(b)** 920–1060 nm. Notations for datasets and spectral channels are the same as in Fig. 3.

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