

Characterization and mitigation of water vapor effects

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Characterization and mitigation of water vapor effects in the measurement of ozone by chemiluminescence with nitric oxide

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Abstract

Laboratory experiments were conducted to investigate the effects of water vapor on the reaction of nitric oxide with ozone in a chemiluminescence instrument used for fast response and high sensitivity detection of atmospheric ozone. Water vapor was introduced into a constant level ozone standard and both ozone and water vapor signals were recorded at 10 Hz. The presence of water vapor was found to reduce, i.e. quench the ozone signal. A correction factor was determined to be $4.15 \pm 0.14 \times 10^{-3}$, which corresponds to a 4.15 % increase in the measured ozone signal per 10 mmol mol⁻¹ co-sampled water vapor. An ozone-inert water vapor permeable membrane (Nafion dryer) was installed in the sampling line and was shown to remove the bulk of the water vapor mole fraction in the sample air. At water vapor mole fractions above 25 mmol mol⁻¹, the Nafion dryer removed over 75 % of the water vapor in the sample. This reduced the ozone signal correction from over 11 % to less than 2.5 %. The Nafion dryer was highly effective at reducing the fast fluctuations of the water vapor signal (more than 97 %) while leaving the ozone signal unaffected, which is a crucial improvement for minimizing the interference of water vapor fluxes on the ozone flux determination by the eddy covariance technique.

1 Introduction and background

Recent developments in instrumentation for ambient air ozone measurements have enabled direct observations of open ocean atmospheric ozone concentrations and fluxes. The measurement of ozone is based on the chemiluminescence reaction of ozone (O₃) and nitric oxide (NO) (Reaction R1), which emits light between 600 nm < λ < 2800 nm that is detected with a photomultiplier tube (PMT):



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5 The excited nitrogen dioxide (NO_2^*) reaches equilibrium through photoemission (Reaction R2). NO_2^* can also react with a molecule through collisional energy transfer, reducing it to the ground state and effectively quenching the signal (Reaction R3). The chemiluminescence signal resulting from the reaction of nitric oxide and ozone is sensitive to several other atmospheric molecules such as H_2 , CO_2 , and H_2O (Matthews et al., 1977). An earlier study did not find an effect of water vapor at 75 % saturation when compared to 0 % saturation on the O_3 -NO chemiluminescence reaction (Fontijn et al., 1970). Subsequently, Matthews et al. (1977) found that water vapor is more than ten times more effective at quenching the chemiluminescence signal than molecular hydrogen and more than three times more effective than carbon dioxide, which makes

15 water the primary interferent of this ozone measurement under ambient air conditions. In contrast to the O_3 -NO chemiluminescence measurement, instruments based on the reaction of ozone and ethylene reported an increase in ozone signal with water vapor (Kleindienst et al., 1993). This was determined to be due to a second compound being formed in the presence of water vapor that generates chemiluminescence.

20 Instead of correcting for the quenching effect of water vapor, some instruments were configured to supply a flow of water vapor to the reaction chamber to keep the effect of water vapor constant, complicating the operation of the system (Ridley and Grahek, 1990). Another proposed method to account for the quenching effect of water was to approximate the reduction in the ozone signal as a function of the water vapor mole fraction and apply a correction factor (Lenschow et al., 1981; Ridley et al., 1992):

$$\text{O}_3 = \text{O}_{3\text{m}}(1 + \alpha r) \quad (1)$$

where O_3 is the corrected ozone mole fraction, $\text{O}_{3\text{m}}$ is the measured ozone volumetric mole fraction in nmolmol^{-1} , α is the correction factor, and r is the water vapor

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mole fraction (expressed as the ratio of moles of water vapor to moles of dry air in mmol mol^{-1} , which is equivalent to parts per thousand). Lenschow et al. (1981) reported the α correction factor as $5 \times 10^{-3} \pm 1 \times 10^{-3}$ and the work of Ridley et al. (1992) further refined the value to $4.3 \times 10^{-3} \pm 0.3 \times 10^{-3}$. For example, for a typical equatorial region open ocean atmospheric water vapor mole fraction of 30 mmol mol^{-1} the correction accounts to 15 % when using the correction factor of 5×10^{-3} . A correction of this magnitude was applied by Williams et al. (2006) in their chemiluminescence measurement of ozone. Previous work has not detailed if and how much the correction factor is dependent on instrument configuration and operational conditions, or if this correction is universally applicable. Prior to the experiments described here, the correction factor had not been determined for our particular custom-built fast-response ozone instrument (FROI). Previous work with this instrument had therefore selected $\alpha = 5 \times 10^{-3}$ according to Lenschow et al. (1981), which resulted in up to a 25 % correction for determining the atmospheric ozone mole fraction (Lang, 2008; Bariteau et al., 2010; Helmig et al., 2012b).

A benefit of the fast response time and high sampling frequency of a chemiluminescence ozone instrument is the ability to define surface fluxes in combination with a sonic anemometer by the eddy covariance technique. Applying a correction to the ozone signal to account for the water vapor influences is particularly critical for these eddy covariance calculations as these are susceptible to interferences from the total atmospheric water vapor mole fraction and the water vapor flux. Our FROI has been deployed for ozone flux determination to locations vastly ranging in water vapor content, from the dry arctic to the equatorial ocean (Bariteau et al., 2010; Helmig et al., 2012a; Helmig et al., 2012b). Reynolds averaging of the corrected ozone signal in Eq. (1) and the vertical component of the wind vector results in the following equation for the water vapor corrected ozone flux:

$$F_{\text{O}_3} = (1 + \alpha \bar{r}) F_{\text{O}_{3\text{m}}} + \overline{\alpha \text{O}_{3\text{m}} w' r'}, \quad (2)$$

were studied in more depth, with a critical examination of the applicability of the correction factors determined in the earlier work of Lenschow et al. (1981) and Ridley et al. (1992).

2 Instrumentation and methodology

Ozone was measured by a custom-built FROI with a precision sufficient to resolve small changes in ozone mole fractions at a high temporal resolution. The FROI has a sensitivity of ~ 2000 counts s^{-1} ppbv $^{-1}$ and a background noise of 900 counts s^{-1} . Details and a schematic of the FROI have been published by Bariteau et al. (2010) (see Fig. 1 in this reference for a schematic of the FROI). Sample air was pulled through a Teflon[®] (PFA, perfluoroalkoxy copolymer) line controlled to 1.5 L min $^{-1}$ by a mass flow controller (MFC). All ozone sample tubing was 0.64 cm outer diameter Teflon[®] tubing. Nitric oxide reactant gas flowed through stainless steel tubing and was controlled at 3 mL min $^{-1}$. The sample and NO were mixed in a 44 cm³ gold-plated reaction chamber. The reaction chamber temperature was maintained at 30 °C by a heater and temperature controller. An integrated PMT housing Peltier cooler maintained the PMT temperature at -30 °C (Hamamatsu, Model C10372, Japan), essential to reach low noise and high sensitivity levels. The reaction chamber pressure was controlled to 18 Torr by a pressure controller (UPC 1300, Coastal Instruments) downstream of the reaction chamber, which asserted that the instrument response was insensitive to fluctuations in the sample delivery flow rate. Photons were counted by a PMT (Hamamatsu Photonics K. K., Shizuoka, Japan) with a cutoff filter (RG-610, Newport Industrial Glass, Stanton, CA) removing radiation with wavelengths less than 600 nm. The FROI was calibrated against a commercial UV absorption instrument (Model TEI 49i, Thermo Scientific, Franklin, MA, USA). This UV-instrument was referenced against the ozone standard at the Global Monitoring Division (GMD), National Oceanic and Atmospheric Administration (NOAA), Boulder, Colorado.

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took place behind the NOAA David Skaggs Research Center in Boulder, CO in October 2008. The footprint of the sampling location consisted of a small parking lot surrounded by surface vegetation. The same FROI and Nafion drying system setup were used in this experimental setup. Water vapor was measured by two LI-7500 (LI-COR Inc., Lincoln, NE, USA) hygrometers. These hygrometers were converted to closed path instruments by inserting the calibration tube between the sapphire-glass windows. The FROI and the two LI-CORs were housed in a container for weather protection. Ambient air was drawn through a 23 m sampling line with an inlet located at 4 m height on a meteorological tower. A Teflon[®] membrane filter (5 μm, Millipore, Billerica, MA, USA) was used during ambient air measurements to prevent contamination of the tubing due to air pollutants. The air passed through one LI-COR, then through the Nafion dryer followed by the other LI-COR, before sampling by the FROI. Prior to the experiment, an inter-comparison of both LI-CORs was conducted to determine the offset between the instruments. The ambient air ozone mole fraction was $\sim 39 \text{ nmol mol}^{-1}$ and the water vapor mole fraction varied between 4 and 6 mmol mol^{-1} .

3 Results and discussion

3.1 Effects of water vapor on the chemiluminescence ozone signal

The water vapor mole fraction was varied across different ozone levels in order to determine the appropriate correction factor, α , for this instrument and to evaluate how the correction factor compares with previously reported results for other instruments. To determine the correction factor, a re-write of Eq. (1) is used, shown as:

$$O_{3,0} = O_{3,r}(1 + \alpha r) \quad (3)$$

where $O_{3,0}$ is the ozone signal in counts s^{-1} , when the water vapor mole fraction is $< 0.1 \text{ mmol mol}^{-1}$, and $O_{3,r}$ is the ozone signal at a water vapor mole fraction r . At

enrichment would consequently cause an increase in the FROI signal of 2%. Figure 6 displays the inferred ozone enrichment as a function of the water vapor content, as measured upstream of the Nafion dryer, ranging from 0.3–2% under the water vapor mole fractions applied here.

5 The ozone signal that is restored when using the Nafion dryer was determined by comparing results from three different cases. Case 1 is the sample flow containing 30 nmol mol⁻¹ of ozone in dry air, < 0.1 mmol mol⁻¹ of water vapor, without the sample passing through the Nafion dryer. Case 2 is a humidified sample containing 30 nmol mol⁻¹ of ozone, a water vapor mole fraction of 6.04 mmol mol⁻¹, without passing through the Nafion dryer. Case 3 is for a sample flow with the Nafion dryer installed, containing 30 nmol mol⁻¹ of ozone, 26.5 mmol mol⁻¹ of water vapor upstream of the Nafion dryer and 6.04 mmol mol⁻¹ of water vapor downstream of the Nafion dryer. In Cases 2 and 3, the amount of water vapor entering the FROI reaction chamber is very similar, at ~ 6 mmol mol⁻¹. In theory, the ozone signal from Case 2 should be equal to the ozone signal from Case 1 after correcting for the quenching effect, and Case 3 should agree to Case 1 after correcting for the enrichment and quenching.

10 For Case 1, the FROI signal was 60 645 counts s⁻¹ (Table 1). For Case 2, the corrected ozone signal was determined from the measured 59 135 counts s⁻¹ by using Eq. (4), $\alpha = 4.15 \times 10^{-3}$, and $r = 6.04$ mmol mol⁻¹ of water vapor. This yields a corrected ozone signal of 60 617 counts s⁻¹. The ozone signal for Case 3 required corrections for both enrichment and quenching. The difference in water vapor mole fractions upstream and downstream of the Nafion dryer was 20.1 mmol mol⁻¹ of water vapor, which corresponded to 2.01% of the total molecules in the sample flow being removed by the Nafion dryer. The measured ozone signal was 60 267 counts s⁻¹ corresponding to an ozone signal of 59 079 counts s⁻¹ after this correction. In order to account for the quenching effect, Eq. (4) was applied, with $O_{3m} = 59 079$ counts s⁻¹, $\alpha = 4.15 \times 10^{-3}$, and $r = 6.4$ mmol mol⁻¹ of water vapor. This calculation resulted in a corrected ozone signal of 60 648 counts s⁻¹.

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With these considerations, the three cases gave close agreement, with the difference between the three cases of less than 32 counts s^{-1} (or 0.02 nmol mol^{-1} of ozone), which is well within the precision of the FROI. This consistency confirms the correctness of the determined quenching effect, developed correction algorithms, and the efficiency of the Nafion dryer in mitigating the quenching effects in the FROI ozone detection.

3.4 Reduction of atmospheric water vapor high frequency signals

The high sampling frequency of the FROI and LI-COR allowed for the investigation of high frequency behavior of the ozone and water signal with use of the Nafion dryer, specifically the reduction of water vapor fluctuations that determine the water vapor flux in Eq. (3). The following analyses are based on the experiments conducted on the mesa behind the NOAA-ESRL building.

The water vapor power spectrum distributions with and without the Nafion dryer shown in Fig. 7a illustrate that the Nafion dryer was very efficient in damping the high frequency water vapor signal. The water vapor spectrum obtained without the Nafion dryer has both low and high frequency contributions. White noise was seen at frequencies higher than 2 Hz. The water vapor spectrum with the Nafion dryer installed has its primary contribution in the lower frequency range and a reduction of the higher frequencies when compared to the water vapor signal without the Nafion dryer. The ratio of the integrals of the power spectra showed a 77% reduction of water vapor mole fraction, which confirmed the amount of water vapor removed as seen at the highest water vapor levels in Fig. 5.

The frequency response spectrum in Fig. 7c shows the coherency between the water vapor with and without a Nafion dryer. The coherency is the ratio of the cospectrum between the two water vapor signals and the square root of the product of the power spectra. A coherency value of 1 is representative of a high correlation between two signals at a given frequency. The water vapor signals have high coherency between 10^{-3} and 10^{-2} Hz, a decrease between 10^{-2} and 10^{-1} Hz, and display low coherence

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Table 1. Comparison of measured ozone signals (mean of 15 min data) at 30 nmol mol^{-1} (in counts s^{-1}) for Cases 1, 2 and 3, before and after applying each correction term.

	Case 1	Case 2	Case 3
Water Vapor Mole Fraction	$< 0.1 \text{ mmol mol}^{-1}$ No Nafion Dryer	$6.04 \text{ mmol mol}^{-1}$ No Nafion Dryer	$26.5 \text{ mmol mol}^{-1}$ Nafion Installed
Measured Counts	60 645 ^a	59 135	60 267
Corrected for Enrichment	N/A ^b	N/A	59 079
Corrected for Quenching	N/A	60 617	60 648

^a Calculated from a raw count of 60 975 after correcting for the dry air flow biases of MFC 3 and MFC 5.

^b Not applicable.

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Table A1. Averaged loss of ozone signal at the 3 different ozone levels tested (30, 60, and 100 ppbv).

H ₂ O	Ozone signal loss \pm 95 % confidence interval
6.2 ‰	2.5 \pm 0.2 %
12.0 ‰	4.7 \pm 0.4 %
17.9 ‰	7.1 \pm 0.7 %
23.0 ‰	9.2 \pm 0.6 %
27.1 ‰	11.2 \pm 0.9 %

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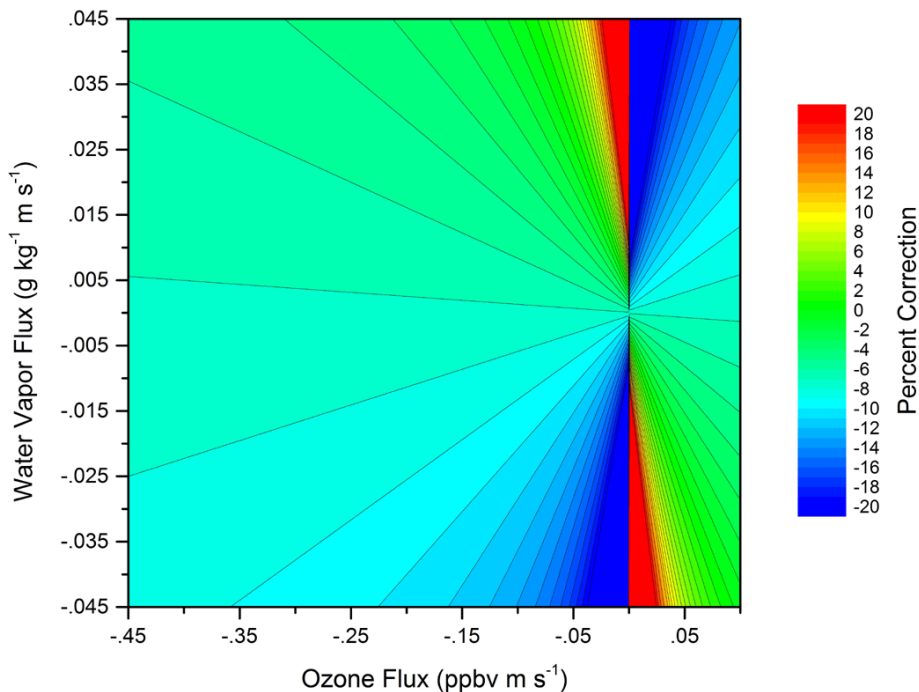


Fig. 1. Isopleths of the correction to be applied to the measured ozone flux as a function of water vapor flux. When the ozone fluxes and water vapor fluxes are in the same direction, the measured ozone flux has a positive error, as seen in the negative correction factor to be applied for quadrants 1 and 3. When the ozone flux and water vapor flux are in opposite directions, there is a negative error, requiring a positive correction to the measured ozone flux, as seen in quadrants 2 and 4.

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Fig. 2. Schematic of the laboratory setup. The red box in the upper-left of the figure shows where the sample air was humidified. Ambient air was scrubbed through a zero-air generator and run through a drying agent to remove any excess water vapor. The flow was varied through mass flow controller (MFC) 1 to produce dry air and MFC 2 to produce humid air. The humidifier was a Nafion membrane containing liquid water in the inner tube and the sample flow through the outer shell. Excess flow was released through the vent with a flow restrictor. MFC 3 controlled the flow to 8 L min^{-1} . This air was mixed with ozone-enriched air from the TEI 49i ozone generator (red box in upper-right of figure). Sample air was provided from a tank of dry breathing air. The flow through the ozone generator was controlled to 1 L min^{-1} . The Nafion drying system, FROI and LI-COR are shown at the lower portion of the figure. Switching valves directed the flow through or around the Nafion dryer. MFCs 5 and 6 controlled the flow to the FROI and LI-COR and were set at 1.5 L min^{-1} . All data were collected on the data acquisition computer housed in the FROI.

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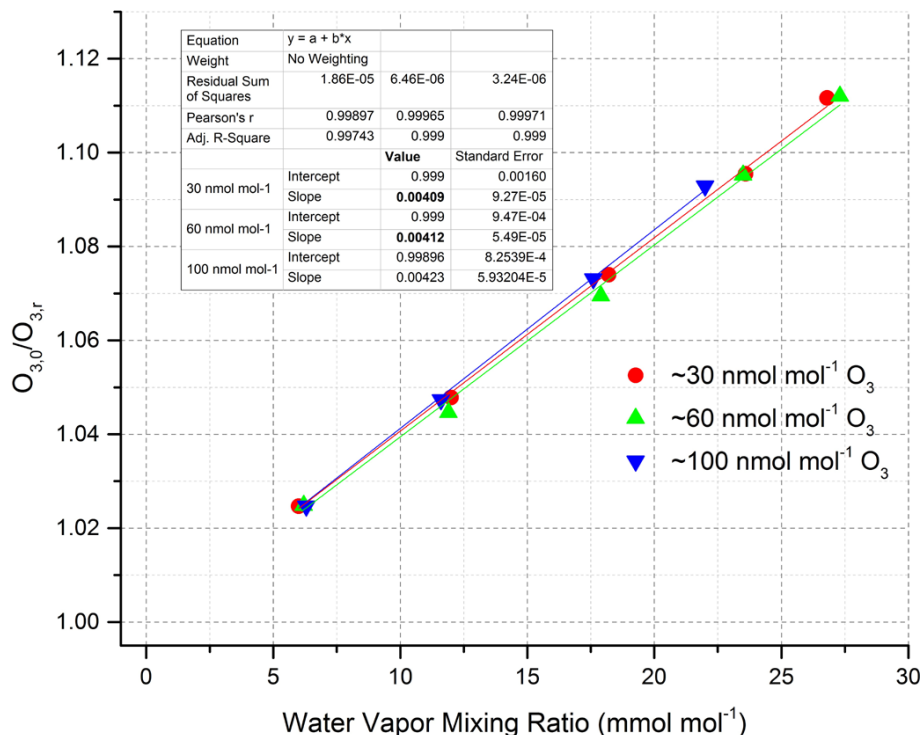


Fig. 3. Ratio of ozone signal at water vapor level less than 0.1 mmol mol⁻¹ ($O_{3,0}$) to ozone signal at water vapor level r ($O_{3,r}$) vs. water vapor mole fraction. The points are color-coded by the amount of ozone generated by the TEI 49i. The solid lines represent results from linear regression analyses. The slope results from the linear regression analyses are shown in the table insert.

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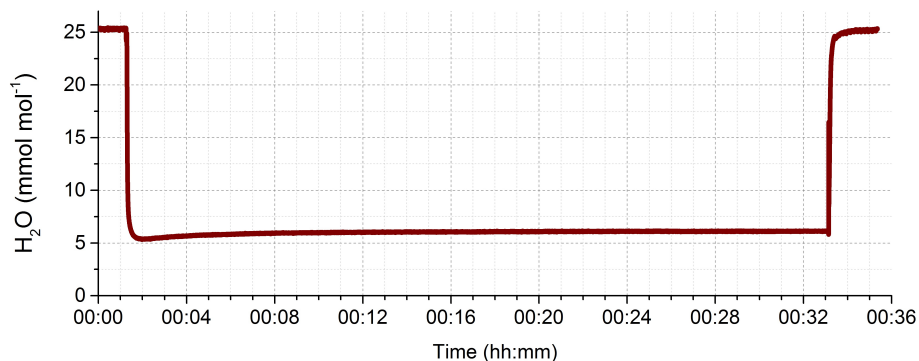


Fig. 4. The water vapor signal before, during and after switching the Nafion dryer into the sample flow.

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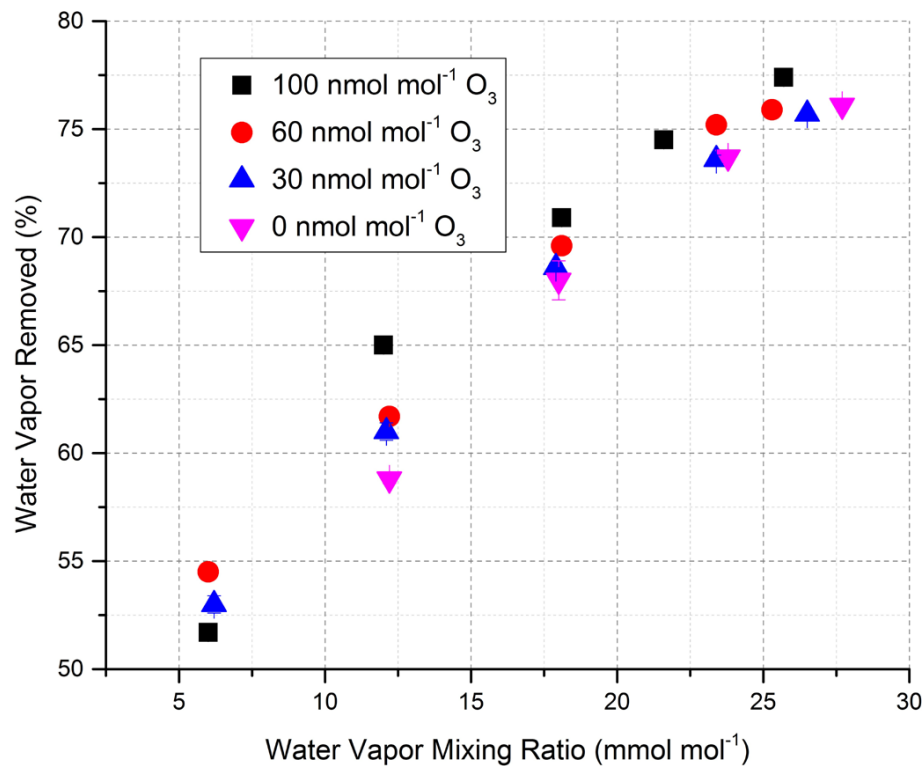


Fig. 5. Fraction of water vapor removed by the Nafion dryer vs. water vapor mole fraction upstream from the Nafion dryer. Data are color-coded by ozone level.

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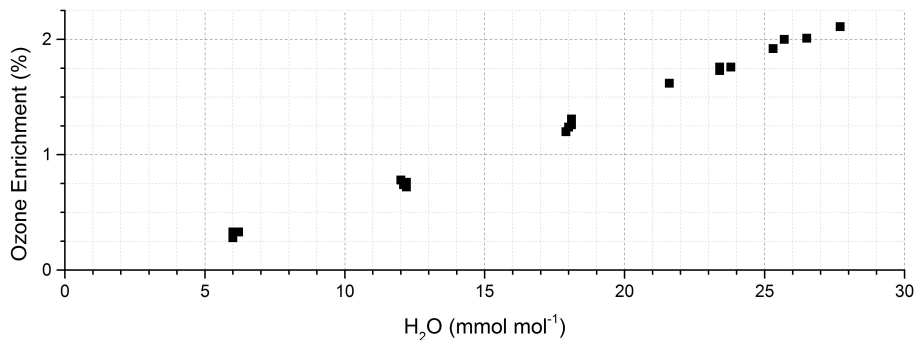


Fig. 6. Increase in the ozone signal from the removal of water vapor molecules by the Nafion dryer as a function of the water vapor mole fraction in the sample air, using the drying efficiency ratios shown in Fig. 5.

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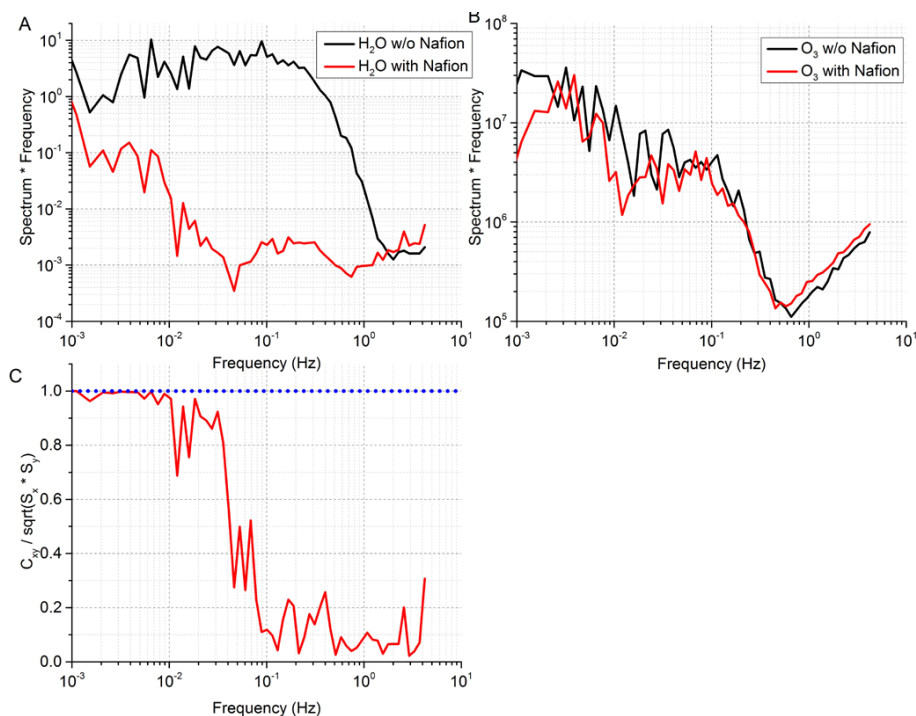


Fig. 7. Spectral distribution plots using two hours of data with the same x-axis range for all three plots. **(A)** Power spectra of ambient water vapor with (red) and without (black) the Nafion drying system. **(B)** Power spectra of the ozone signal with (red) and without (black) the Nafion drying system. **(C)** Coherence spectral distribution of the ambient water vapor signal.

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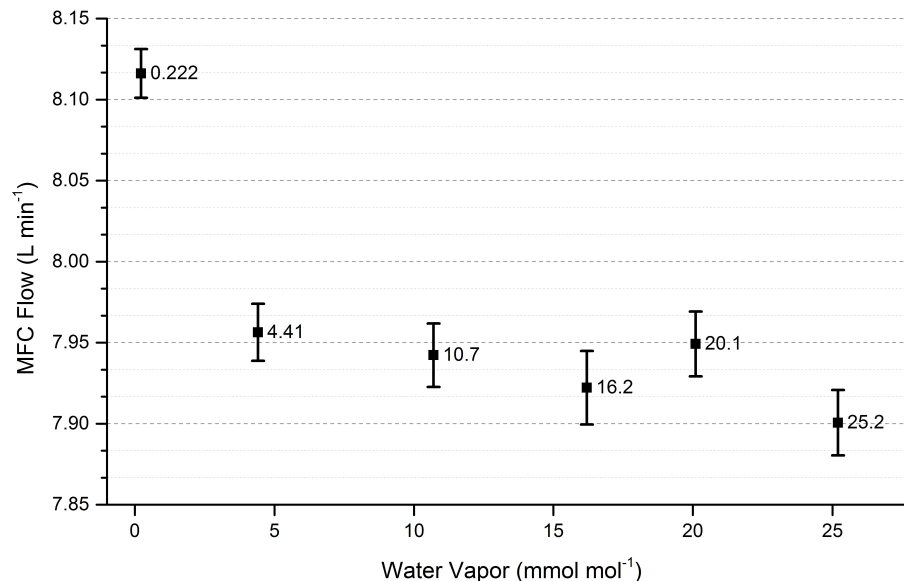


Fig. A1. Flow rate determined with a bubble meter, corrected for ambient pressure and temperature, against water vapor mole fraction, for MFC 3 (Tylan FC-2900) operated at a constant set point. Each point shows the mean flow rate and the error bars represent the standard error with a sample size of 20. The numbers to the right of each point correspond to the water vapor content determined with the LICOR.

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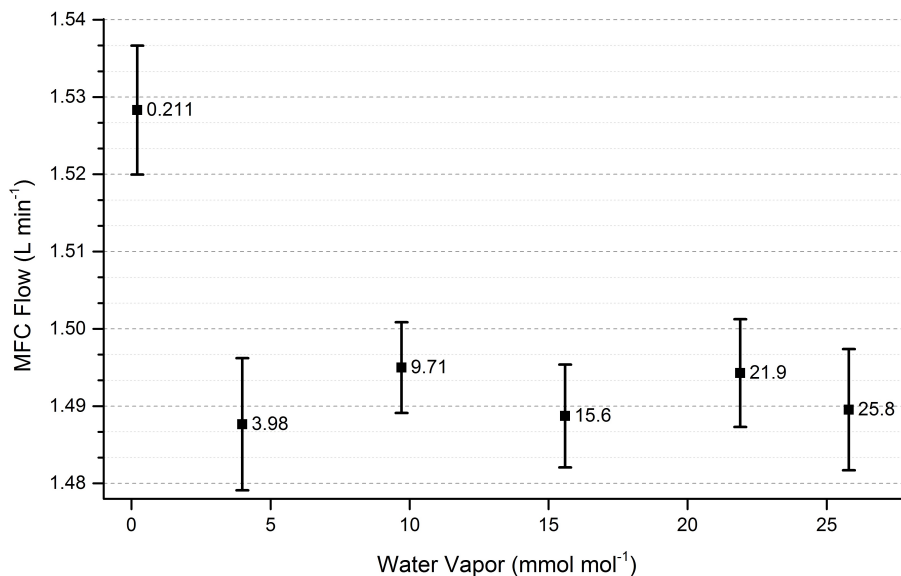


Fig. A2. Flow rate determined with bubble meter, corrected for ambient pressure and temperature, against water vapor mole fraction, for MFC 5 (Tylan FC-260) operated at a constant set point. Each point shows the mean flow rate and the error bars represent the standard error with a sample size of 20. The numbers to the right of each point correspond to the water vapor content determined with the LICOR.

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