



21 **ABSTRACT.-** In this work we study the effects of systematic and random errors on the inversion  
22 of multi-wavelength (MW) lidar data, using the well-known regularization technique, to obtain  
23 vertically-resolved aerosol microphysical properties. The software implementation used here was  
24 developed at the Physics Instrumentation Center (PIC) in Troitsk (Russia) in conjunction with  
25 NASA/Goddard Space Flight Center. Its applicability to Raman lidar systems based on  
26 backscattering measurements at three wavelengths (355, 532 and 1064 nm) and extinction  
27 measurements at two wavelengths (355 and 532 nm) has been demonstrated widely. The  
28 systematic error sensitivity is quantified by first determining the retrieved parameters for a given  
29 set of optical input data consistent with three different sets of aerosol physical parameters. Then  
30 each optical input is perturbed by varying amounts and the inversion is repeated. [Using bimodal](#)  
31 [aerosol size distributions](#), we find a generally linear dependence of the retrieved errors in the  
32 microphysical properties on the induced systematic errors in the optical data. For the retrievals of  
33 effective radius, number/surface/volume concentrations and fine mode radius and volume, we  
34 find that these results are not significantly affected by the range of the constraints used in  
35 inversions. But significant sensitivity was found to the allowed range of the imaginary part of the  
36 particle refractive index. Our results also indicate that there exists an additive property for the  
37 deviations induced by the biases present in the individual optical data. This property permits the  
38 results here to be used to predict deviations in retrieved parameters when multiple input optical  
39 data are biased simultaneously as well as to study the influence of random errors on the  
40 retrievals. The above results are applied to questions regarding lidar design, in particular for the  
41 space-borne multi-wavelength lidar under consideration for the upcoming ACE mission.

42

## 43 **1. - INTRODUCTION**

44 The importance of atmospheric aerosol particles on Earth's climate and on environmental  
45 problems is widely recognized. Particularly, the Intergovernmental Panel on Climate Change  
46 2007 (IPCC 2007) [Forster et al., 2007] stated that atmospheric aerosol particles can produce a  
47 negative radiative forcing that is comparable in magnitude, but opposite in sign, to the forcing  
48 induced by the increase in greenhouse gas concentration. However, according to the IPCC,  
49 radiative forcing by atmospheric aerosol particles has greater uncertainties (twice the estimated  
50 value of the forcing) due to the large spatial and temporal heterogeneities of atmospheric  
51 aerosols [e.g. Haywood and Boucher, 2000], the wide variety of aerosol sources [e.g. Dubovik et  
52 al., 2002], the spatial non-uniformity and intermittency of these sources [e.g. Kaufman et al.,  
53 1997], the short atmospheric lifetime of aerosols [e.g. Seinfeld and Pandis, 1998], processes  
54 occurring in the atmosphere [Eck et al., 2010] and aerosol dynamics [e.g. Pérez-Ramírez et al.,  
55 2012].

56 Because of these challenges, the characterization of atmospheric aerosols is being made  
57 through intense observational programs using remote sensing techniques. For example, NASA  
58 has led several space-borne missions to study aerosol properties worldwide (e.g. the MODIS  
59 instrument on the TERRA and AQUA platforms). However, satellite measurements possess  
60 lower temporal resolution than ground-based systems. For example, the AERONET global  
61 network [Holben et al., 1998] is providing large datasets of high temporal resolution ground-  
62 based aerosol measurements at more than 400 locations worldwide. But the aerosol retrievals by  
63 AERONET and by many satellite platforms only provide column-integrated properties. By  
64 contrast, the lidar technique offers vertical profiling of aerosols, from the first lidars in the early

65 1960s to the more sophisticated Raman lidars [Whiteman et al, 1992, Ansmann et al., 1992] or  
66 High Spectral Resolution Lidars (HSRL) [Shiple et al., 1983; Grund and Eloranta, 1991; She et  
67 al., 1992, 2001]. Moreover, the Nd:YAG laser has been used as the transmitter for multi-  
68 wavelength Raman lidar systems (MW) which have permitted the retrieval of the profile of  
69 aerosol microphysical properties [e.g. Müller et al., 2001,2011; Wandinger et al., 2002; Böckman  
70 et al., 2005; Kolgotin and Müller, 2008; Noh et al., 2009; Balis et al., 2010; Alados-Arboledas et  
71 al., 2011; Tesche et al., 2011; Veselovskii et al., 2012; Papayannis et al., 2012; Wanger et al.,  
72 2013; Navas-Guzmán et al., 2013].

73 The first attempts to obtain aerosol microphysical properties from MW Raman lidar  
74 measurements were done at the Institute for Tropospheric Research (IFT) in Leipzig (Germany)  
75 using the regularization technique [Müller et al., 1999a,b; 2000]. The first retrievals done at the  
76 IFT were based on measurements from a complex lidar system providing six backscattering  
77 (355, 400, 532, 710, 800 and 1064 nm) and two extinction (355 and 532 nm) coefficients.  
78 Following these first efforts, a software capability based on the regularization technique was  
79 developed at the Physics Instrumentation Center (PIC) in Troitsk, Russia. The retrieval code  
80 development at PIC has been further advanced and has incorporated a model of randomly-  
81 oriented spheroids for retrieving dust particle properties [Veselovskii et al., 2010]. Müller et al.,  
82 [2001, 2004, 2005] and Veselovskii et al., [2002, 2004] demonstrated the capability of the  
83 regularization technique to retrieve aerosol microphysical properties from a lidar system that  
84 provides just 5 optical signals using a tripled Nd:YAG laser. The optical data provided by this  
85 system were backscatter coefficients ( $\beta$ ) at 355, 532 and 1064 nm and extinction coefficients ( $\alpha$ )  
86 at 355 and 532 nm (hereafter this configuration is referred as  $3\beta + 2\alpha$ ). The inversion procedure  
87 makes use of averaging of the solutions in the vicinity of the minimum of a penalty function, or

88 discrepancy [Veselovskii et al., 2002]. This averaging procedure increases the reliability of the  
89 inversions even when the input optical data are affected by random errors [e.g. Veselovskii et al.,  
90 2002].

91         However, lidar systems are very complex and generally possess both random and  
92 systematic errors. Random errors arise naturally from the measurement process and some  
93 preliminary random error sensitivity studies were done by Müller et al., (1999a,b) and  
94 Veselovskii et al., (2002, 2004). But to date, there is a lack of studies of the effects of systematic  
95 errors on the microphysical inversions. Systematic errors in lidar systems come from many  
96 different sources and need to be considered. From the hardware point of view, systematic errors  
97 can be due to, for example, non-linearity of a photodetector or errors in calibration of the optical  
98 data or the effect of depolarization due to optical imperfections in channels that are sensitive to  
99 polarized light. From the methodological point of view, systematic errors can be caused by, for  
100 example, errors in the assumed atmospheric molecule density profile, the selection of the  
101 reference level (an “aerosol-free” region that may actually contain a small concentration of  
102 particles), or the use of an incorrect extinction-to-backscatter ratio to convert backscatter lidar  
103 measurements to extinction.

104         In general, we expect that systematic errors such as these can affect the retrieval. The aim  
105 of this work, therefore, is to study the sensitivity of microphysical retrievals by the regularization  
106 technique to systematic variations in the input optical data provided by the  $3\beta + 2\alpha$  lidar  
107 configuration. Particularly, we will focus on the study of bimodal size distributions widely found  
108 in nature (e.g. Dubovik et al., 2002). We will show that the results obtained can also be used to  
109 assess the sensitivity of the retrievals to random errors in a new way. The study involves  
110 simulations based on three different bi-modal aerosol size distributions, one with a large

111 predominance of fine mode, another with slight predominance of coarse mode and the last one  
112 with slight predominance of fine mode.

113 The procedure that we used is the following: first the optical data consistent with the  
114 three aerosol size distributions described above are generated using Mie theory. Then the optical  
115 inputs are systematically altered to provide a known amount of systematic error in each of the  
116 individual input data. The inversion code is run using both the biased and unbiased optical data  
117 and the deviations in the retrieved aerosol parameters are quantified. The methodology and the  
118 simulation approach are presented in section 2. Section 3 is devoted to the results. Finally, in  
119 Section 4 we present a summary and conclusions.

120

## 121 **2.- METHODOLOGY AND SIMULATION APPROACH**

### 122 **2.1.- Inversion technique**

123 The optical characteristics of an ensemble of polydisperse aerosol particles are related to  
124 the particle volume distribution via Fredholm integral equations of the first kind as follows  
125 [Müller et al., 1999a; Veselovskii et al., 2002]:

$$126 \quad g_j(\lambda_i) = \int_{r_{min}}^{r_{max}} K_{j,N}(m, r, \lambda_i) n(r) dr \quad (1)$$

127 Where 'j' corresponds either to backscatter ( $\beta$ ) or extinction ( $\alpha$ ) coefficients,  $g_j(\lambda_i)$  are the  
128 corresponding optical data at wavelength  $\lambda_i$ ,  $n(r)$  is the aerosol size distribution expressed as the  
129 number of particles per unit volume between  $r$  and  $r + dr$ , and  $K_{j,N}(m, r, \lambda_i)$  are the number kernel  
130 functions (backscatter or extinction) which are here calculated from Mie theory assuming  
131 spherical particles and depend on particle refractive index 'm', particle radius 'r' and wavelength

132 ‘ $\lambda$ ’. Finally,  $r_{\min}$  and  $r_{\max}$  correspond to the minimum and maximum radius used in the inversion.  
 133 The size distribution in Equation 1 can be written in terms of surface ( $s(r) = 4\pi r^2 n(r)$ ) or volume  
 134 ( $v(r) = (4/3)\pi r^3 n(r)$ ) size distribution. The corresponding kernels are obtained by dividing  
 135  $K_{j,N}(m,r,\lambda_i)$  by  $4\pi r^2$  and  $(4/3)\pi r^3$  respectively, and are thus given by:

$$136 \quad K_{j,S}(m,r,\lambda) = \frac{K_{j,N}(m,r,\lambda)}{4\pi r^2} \quad (2)$$

$$137 \quad K_{j,V}(m,r,\lambda) = \frac{3K_{j,N}(m,r,\lambda)}{4\pi r^3} \quad (3)$$

138  
 139 where  $K_{j,S}(m,r,\lambda)$  and  $K_{j,V}(m,r,\lambda)$  are the surface and volume kernel functions, respectively.  
 140 Generally, the volume kernel functions are used in the retrieval procedure of aerosol  
 141 microphysical properties [Heintzenberg et al., 1981; Qing et al., 1989]. Thus, we perform the  
 142 retrieval of volume size distribution using the volume kernel functions of equation 3. More  
 143 details about the computation of these volume kernel functions from Mie extinction coefficients  
 144 for spherical particles can be found in the references [e.g. Bohren and Huffman, 1983].

145 The regularization technique used here to solve equation 1 has been discussed extensively  
 146 elsewhere [e.g. Veselovskii et al., 2002, 2004, 2005] and thus we provide here only a brief  
 147 overview. The key point is identifying a group of solutions which, after averaging, can provide a  
 148 realistic estimation of particle parameters. Such identification can be done by considering the  
 149 discrepancy ( $\rho$ ) defined as the difference between input data  $g(\lambda)$  and data calculated from the  
 150 solution obtained. The retrieval uses an averaging procedure that consists of selecting a class of  
 151 solutions in the vicinity of the minimum of discrepancy [Veselovskii et al., 2002, 2004]. Such an  
 152 averaging procedure stabilizes the inversion, as the final solution for size distribution and aerosol

153 parameters is an average of a large number of individual solutions near the minimum of  
154 discrepancy [Veselovskii et al., 2002]. In general, we average approximately 1% of the total  
155 number of solutions in arriving at the best estimate of the particle parameters.

156 The inverse problem considered here is under-determined, so constraints on the inversion  
157 are needed. We consider a set of possible values of the particle refractive index as well as a set of  
158 possible radii within a certain size interval. In general, the retrieval result will depend on the  
159 range of parameters considered: the larger the range, the higher the uncertainty of the retrieval as  
160 determined by the spread in the solutions obtained. So the range of parameters should be chosen  
161 reasonably. In our research, the real part of the aerosol refractive index ( $m_r$ ) is allowed to vary  
162 from 1.33 to 1.65 with a stepsize of 0.025, while the imaginary part ( $m_i$ ) varies over the range of  
163 0-0.01 with a stepsize of 0.001. The size interval for the inversions was limited to 0.075 – 5  $\mu\text{m}$   
164 with a stepsize of 0.025  $\mu\text{m}$ . Tests revealed that reducing the stepsize of the different parameters  
165 in the inversion does not decrease the spread in the solution. Therefore, we take the stepsizes  
166 used as adequate for the purposes of the present sensitivity study.

167

## 168 **2.2.- Size distribution for the simulations**

169 For these simulations, we used bimodal aerosol size distributions given as [Veselovskii et  
170 al., 2004]:

$$171 \quad \frac{dn(r)}{d\ln(r)} = \sum_{i=f,c} \frac{N_{t,i}}{(2\pi)^{1/2} \ln \sigma_i} \exp \left[ \frac{(\ln r - \ln r_i^n)^2}{2 (\ln \sigma_i)^2} \right] \quad (4)$$

172



173 Where  $N_{ti}$  is the total particle number of the  $i$ th mode,  $\ln[\sigma_i]$  is the mode width of the  $i$ th mode  
 174 and  $r_i^n$  is the mode radius for the number concentration distribution. The index  $i = f, c$   
 175 corresponds to the fine mode and the coarse mode, respectively. In the retrieval procedure, the  
 176 fine mode is taken to include all particles with radius between 0.075  $\mu\text{m}$  and 0.5  $\mu\text{m}$  while the  
 177 coarse mode includes all particles with radius between 0.5  $\mu\text{m}$  and 5  $\mu\text{m}$ . On the other hand, the  
 178 same distribution can be written for volume concentration  $v(r)$ , which is usually preferred  
 179 because both fine and coarse mode can be easily distinguished. Moreover, the standard  
 180 deviations of  $n(r)$  and  $v(r)$  are the same when using the relationships between radius and  
 181 concentrations for each mode given by [Horvath et al., 1990]:

182

$$r_i^v = r_i^n \exp \left[ \frac{(\ln r - \ln r_i^n)^2}{2(\ln \sigma)^2} \right] \quad (5)$$

183

$$V_{ti} = N_{ti} \frac{4}{3} \pi (r_i^n)^3 \exp \left[ \frac{9}{2} (\ln \sigma)^2 \right] \quad (6)$$

184

185

186 We consider three types of aerosol size distributions for the simulations which we call type I,  
 187 type II and type III. These size distributions are used to approximate real aerosol types found in  
 188 the atmosphere. All types use  $r_f^v = 0.14 \mu\text{m}$ ,  $\ln \sigma_f = 0.4$ ,  $r_c^v = 1.5 \mu\text{m}$  and  $\ln \sigma_c = 0.6$ . [These mode](#)  
 189 [radii and widths are representative of those provided by Dubovik et al., \(2002\) in the AERONET](#)  
 190 [climatology database and are thus considered to represent a large fraction of naturally occurring](#)  
 191 [aerosols](#). The differences between type I, II and III are the ratio of fine and coarse mode ( $V_{tf}/V_{tc}$ ).  
 192 Type I yields  $V_{tf}/V_{tc} = 2$  and represents a distribution with a predominance of fine mode. This  
 193 type can be considered to represent industrial and biomass burning aerosols [e.g. Eck et al.,  
 194 2003; Muller et al., 2004; Schafer et al., 2008]. Type II yields  $V_{tf}/V_{tc} = 0.2$  and corresponds to a

195 slight predominance of the coarse mode over the fine mode [e.g. Smirnov et al., 2002, 2003; Eck  
196 et al., 2005, 2010]. This type is consistent with a mixture of dust/marine aerosol and those of  
197 pollution or biomass burning. Finally, type III yields  $V_{\text{tf}}/V_{\text{tc}} = 1$  and corresponds to a slight  
198 predominance of fine mode over the coarse mode [e.g. Xia et al., 2007; Ogunjobi et al., 2008;  
199 Yang and Wening, 2009; Eck et al., 2009]. This type is representative of predominance of  
200 pollution or biomass-burning but with considerable influence of dust particles. Figure 1  
201 illustrates the three size distributions used. For convenience, the size distributions of Figure 1 are  
202 normalized. Finally, if we would include a strong predominance of coarse mode (e.g. marine or  
203 dust aerosol) in  $3\beta + 2\alpha$  lidar measurements, the effects of polarization and non-sphericity should  
204 be taken into account and previous work indicates that the use of kernel functions for non-  
205 spherical particles can improve the retrievals [Veselovskii et al., 2010] Here, however, our  
206 purpose is to calculate sensitivities due to random and systematic uncertainties so we consider  
207 only spherical (Mie) kernels and thus exclude a distribution with a strong predominance of the  
208 coarse mode.

209

210 [Insert Figure 1 here]

211

212 The simulation consists of generating the three backscattering and two extinction  
213 coefficients for the  $3\beta + 2\alpha$  lidar configuration using Mie theory for the three aerosol size  
214 distributions: type I, type II and type III. These optical data are generated for six different  
215 configurations of aerosol refractive indices ( $m_r$  values of 1.35, 1.45 and 1.55 and  $m_i$  values of  
216 0.005 and 0.01). From previous studies (Muller et al., 1.999; Veselovskii et al., 2002) error in  $m_r$   
217 was initially established as  $\pm 0.05$  while error in  $m_i$  was approximately 50%. Moreover, the

218 AERONET network provides refractive indices with very similar errors (Dubovik et al., 2000).  
219 Thus, the range of refractive indexes proposed for the size distribution is enough to cover most  
220 of the values obtained by AERONET (Dubovik et al., 2002).

221 The regularization inversion is then performed on these data and we obtain the retrieved  
222 microphysical parameters ' $M_{ret}$ '. The next step consists of applying a systematic bias, denoted as  
223  $\Delta\epsilon$ , to one optical datum at a time. The bias varies from -20% to +20% in 8 intervals. For each of  
224 these induced biases, the inversion is performed and a new size distribution and set of  
225 microphysical parameters,  $M_{bias}$ , are then obtained. The comparisons to be performed are  
226 expressed as the percentage difference  $100 * (M_{bias} - M_{ret})/M_{ret}$ . This procedure is applied to each  
227 of the 5 optical data used in the  $3\beta + 2\alpha$  lidar configuration.

228

## 229 **3.- RESULTS**

### 230 **3.1.- Uncertainties in the retrieval of particle refractive index**

231 The  $3\beta + 2\alpha$  lidar configuration permits the retrieval of particle refractive index, both real  
232 ( $m_r$ ) and imaginary ( $m_i$ ) parts [e.g. Veselovskii et al., 2002], by use of the regularization scheme.  
233 But the inverse problem of equation 1 is under-determined and, as already stated, constraints are  
234 needed to permit solutions to be obtained. Particularly, the selection of the range of refractive  
235 indices permitted in the retrieval is important. As commented, we limited the range of  $m_r$   
236 between 1.33 and 1.65 and  $m_i$  from 0.0 up to 0.01. These ranges cover most types of aerosol  
237 particles present in the atmosphere, except for strongly absorbing particles such as black carbon.  
238 Moreover, given that the longest wavelength measurement used here is 1064 nm, the technique  
239 has reduced sensitivity to the coarse mode of the aerosol distribution. Thus, **to stabilize the**

240 retrievals, the maximum radius of the retrieval interval was set to 5  $\mu\text{m}$ . Additionally, the kernel  
241 functions for radius below 0.075 are very near to zero, and thus the minimum radius allowed was  
242 set to 0.075  $\mu\text{m}$ . The behavior of the kernel functions versus wavelength can be consulted, for  
243 example, in Chapter 11 of Bohren and Huffman, 1983.

244 In the analysis that follows, we do not present results on the refractive index sensitivity  
245 analysis. The reason for this is that we found that the retrieval of refractive index is very  
246 sensitive to the range of permitted values for the imaginary part of the refractive index. Changing  
247 the range of permitted values of the imaginary part can change the retrieved refractive index  
248 significantly while not significantly affecting the values of the other retrieved quantities. For  
249 example, computations allowing  $m_i$  to range up to 0.1 provides retrieved values of  $m_i$  of  
250 approximately 0.03, when the values of the input size distributions were 0.01-0.005. Therefore,  
251 recalling that the retrieval is under-determined, we conclude that we can provide reasonable  
252 estimates of the refractive index only with reasonable constraints for  $m_i$ . All these results just  
253 magnify the point that refractive index retrievals are difficult with the MW lidar technique and  
254 that some *a priori* knowledge of the aerosol absorption is helpful to constrain the inversion. A  
255 more detailed discussion about the limitations of the averaging procedure used here to retrieve  
256 accurate values of particle refractive index is in Veselovskii et al., [2013].

### 257 **3.2.- Effects on the retrievals of systematic errors in the optical data.**

258 For the scheme described previously, Figure 2 presents the sensitivity analysis for the  
259 retrieval of effective radius ( $r_{\text{eff}}$ ). Every point corresponds to the mean value of the six different  
260 combinations of aerosol refractive indices used in generating the set of optical data. The error  
261 bars shown are the standard deviations of these mean values. Generally linear patterns are

262 observed for the deviation in retrieved value of  $r_{\text{eff}}$  for differing biases in the input optical data  
 263 for all of types I, II and III aerosols. As the linear patterns pass through the origin, least-squares  
 264 fits of the form  $Y = aX$  were done to the points shown in the plot. Given the definition of  $\Delta r_{\text{eff}} =$   
 265  $r_{\text{eff,bias}} - r_{\text{eff,ret}}$ , positive slopes indicate higher values of  $r_{\text{eff}}$  when the optical data are affected by  
 266 positive biases than when they are not affected by biases, while for negative slopes just the  
 267 opposite occurs. Moreover, Figure 2 reveals the same general patterns for all of types I, II and III  
 268 for each optical channel, with only small changes in the absolute values of the slopes of the  
 269 linear fits. It is quite apparent that the retrievals are more sensitive to biases in the extinction  
 270 coefficients. The lowest sensitivities are to biases in  $\beta(355 \text{ nm})$  and  $\beta(532 \text{ nm})$  while for biases in  
 271  $\beta(1064 \text{ nm})$  the sensitivity of the retrievals is in between those obtained for extinction and  
 272 backscattering coefficients at 355 and 532 nm. Figure 2 also reveals that the linear patterns for  
 273 different optical channels have different signs of the slopes. **Considering the parameters to which**  
 274 **the retrievals are most sensitive, the linear fit of  $\alpha(355 \text{ nm})$  gives negative values of slope ( $a = -$**   
 275  **$1.68 \pm 0.12$  for type I,  $a = -1.74 \pm 0.03$  for type II and  $a = -1.84 \pm 0.04$  for type III ), while for**  
 276  **$\alpha(532 \text{ nm})$  the slopes are positive ( $a = 1.51 \pm 0.04$  for type I,  $a = 1.82 \pm 0.09$  for type II and  $a =$**   
 277  **$1.71 \pm 0.10$  for type III).**

278 [Insert Figure 2 here]

279 The Ångström law, either for the extinction  $\alpha(\lambda) = k\lambda^{-\eta_\alpha}$  or for the backscattering  
 280  $\beta(\lambda) = k\lambda^{-\eta_\beta}$  can be used to help understand the sign of the slopes of Figure 2. For the  
 281 wavelengths used here, the Ångström exponents  $\eta_\alpha$  and  $\eta_\beta$  characterize the spectral features of  
 282 aerosol particles and are related to the size of the particles: Large values of  $\eta_\alpha$  and  $\eta_\beta$  are mainly  
 283 associated with predominance of fine mode particles while low values are associated with a

284 predominance of coarse mode [e.g. Dubovik et al., 2002]. Moreover, many works [e.g. Alados-  
285 Arboledas et al., 2003; O'Neill et al., 2005; Veselovskii et al., 2009] found an inverse  
286 relationship between the Ångström exponent for extinction and the effective radius: large values  
287 of Ångström exponent are associated with low values of  $r_{\text{eff}}$  while just the opposite occurs for  
288 low values of Ångström exponent. Considering this and given that  $\alpha(355 \text{ nm})$  is generally larger  
289 than  $\alpha(532 \text{ nm})$ , a positive bias in  $\alpha(355 \text{ nm})$  increases the spectral difference with  $\alpha(532 \text{ nm})$   
290 and would increase the value of the Ångström exponent and thus would result in a decrease in  
291 the retrieved particle radius. This agrees with the negative slopes of  $\alpha(355 \text{ nm})$  observed in  
292 Figure 2. On the other hand, a positive bias in  $\alpha(532 \text{ nm})$  reduces the spectral difference with  
293  $\alpha(355 \text{ nm})$  and thus serves to decrease  $\eta_{\alpha}$ . Thus, we would expect an increase in the retrieved  
294 particle radius which agrees with the positive slopes observed for  $\alpha(532 \text{ nm})$  in Figure 2. The  
295 slopes of  $\beta(355 \text{ nm})$  and  $\beta(532 \text{ nm})$  possess mostly the same sign as the corresponding extinction  
296 coefficient at each wavelength, and similar logic concerning the relationship of the Ångström  
297 exponent and the particle size given for  $\alpha(355 \text{ nm})$  and  $\alpha(532 \text{ nm})$  can be used to explain this  
298 behavior as well. Finally, for  $\beta(1064 \text{ nm})$  we observe positive slopes ( $a = 0.791 \pm 0.008$  for type  
299 I,  $a = 0.54 \pm 0.07$  for type II and  $a = 0.84 \pm 0.02$  for type III). Positive biases of  $\beta(1064 \text{ nm})$   
300 decrease the spectral difference between  $\beta(355 \text{ nm})$  and  $\beta(532 \text{ nm})$  indicating a decrease of the  
301 Ångström exponent, and thus we would expect an increase in the retrieved particle size which  
302 agrees with the presence of positive slopes in the plot.

303 Figure 3 presents the sensitivity analysis for the retrieval of number concentration (N).  
304 From Figure 3 we again generally observe linear patterns of the deviation in retrieved value of N  
305 for differing biases in the input optical data. Linear fits through the origin in the forms  $Y = aX$   
306 were also performed here. Interestingly, the slopes of the linear fits of the extinction coefficients

307 present opposite signs to those determined for the retrieval of  $r_{\text{eff}}$ , with positive values for  $\alpha(355$   
308 nm) ( $a = 3.09 \pm 0.12$  for type I,  $a = 4.83 \pm 0.22$  for type II  $a = 3.05 \pm 0.13$  for type III) and  
309 negative values for  $\alpha(532$  nm) ( $a = -2.78 \pm 0.17$  for type I,  $a = -4.09 \pm 0.23$  for type II and  $a = -$   
310  $2.61 \pm 0.12$  for type III). Therefore, we see in the retrieved results, for example, that to  
311 compensate for a radius enhancement due to biased input data the retrieval tends to decrease  
312 number density.

313 [Insert Figure 3 here]

314 For the sensitivities of  $r_{\text{eff}}$  and N shown in Figures 2 and 3, the absolute values of the  
315 slopes at  $\alpha(355$  nm) and  $\alpha(532$  nm) are larger than 1 which indicates that the percentage  
316 deviations in the retrieved  $r_{\text{eff}}$  and N using biased data are larger than the percentage bias  
317 imposed on the input optical data. Thus, the accuracy of  $r_{\text{eff}}$  retrievals using  $3\beta + 2\alpha$  lidar is  
318 strongly dependent on the accuracy associated with the extinction coefficients. Other slopes with  
319 absolute value less than 1, as for example those obtained for  $r_{\text{eff}}$  as a function of biases in  $\beta(1064$   
320 nm) ( $0.79 \pm 0.01$  for aerosol type I,  $0.54 \pm 0.07$  for aerosol type II and  $a = 0.84 \pm 0.02$  for type  
321 III) indicate that while the retrieval is still quite sensitive to biases in  $\beta(1064$  nm), the deviations  
322 in the retrieved parameters is less than the magnitude of the biases. Finally, the slopes of  $r_{\text{eff}}$  as a  
323 function of biases in the input data for  $\beta(355$  nm) and  $\beta(532$  nm) are quite small indicating that  
324 biases in these optical parameters have relatively small effects on the retrieval of  $r_{\text{eff}}$ . However,  
325 for the retrieval of number concentration the effects of biases in the backscattering optical data  
326 are not negligible with absolute values of the slopes of the linear fits between 1.3 and 0.3.

327 As with the effective radius and number concentration, we have performed the sensitivity  
328 analysis for the other microphysical parameters obtained from the inversion of  $3\beta + 2\alpha$  lidar data.

329 For these studies, we have also observed generally linear patterns when considering the  
330 differences in the retrieved microphysical parameters as a function of the bias in the input optical  
331 data. Again, the linear patterns pass through the origin and we therefore assumed least-squares  
332 fits of the form  $Y = aX$ . The results of the linear fits for all the parameters are summarized in  
333 Table 1, including also the slopes obtained for  $r_{\text{eff}}$  and  $N$  in Figures 2 and 3, respectively.

334 We note that for some parameters the linear fit possesses different slopes for positive and  
335 negative biases  $\Delta\epsilon$ . For example, in the case of  $r_{\text{eff}}$  for type II,  $\beta(532 \text{ nm})$  has a slope of  $-0.48 \pm$   
336  $0.02$  for positive biases and  $0.02 \pm 0.02$  for negative biases. This is taken into account in Table 1,  
337 where, if there is a difference in slope between positive and negative biases in the input data, the  
338 slopes relating to the positive biases are indicated by (p) while those associated with negative  
339 biases are indicated by (n). We take this difference in slope to be a reflection of the reduced  
340 sensitivity to the coarse mode of the distribution. From Table 1 we observe that the number  
341 concentration is by far the most sensitive parameter to bias in the optical data, particularly to  
342 those biases in  $\alpha(355 \text{ nm})$  and  $\alpha(532 \text{ nm})$ . Moreover, the sensitivities to biases at  $\beta(355 \text{ nm})$  are  
343 generally larger for type I than for type II (absolute values of slopes are larger) with type III  
344 being in the middle. This finding can be explained by the fact that, for the same total volume,  
345 small particles (which predominate in type I) generally provide larger backscattering of light at  
346 the shorter wavelengths (phase function at  $180^\circ$  is larger) [e.g. Mischenko et al., 2000; Liou,  
347 2002; Kokhanovsky 2004].

348 [Insert Table 1 here]

349 From Table 1 the slopes calculated from the linear fits of surface concentration ( $S$ ) as  
350 function of biases in the optical data present the same patterns (sign of slopes) between types I, II



351 and III. The difference in the absolute values of slopes between the three types are then  
352 associated with the differences in the size distribution and with the changes in the kernel  
353 functions. The largest sensitivities of  $S$  are found for biases at  $\alpha(355 \text{ nm})$  (absolute values of  
354 slopes  $\sim 2.0$ ). Sensitivities to biases at  $\alpha(532 \text{ nm})$  (absolute values of slopes between 1.07 and  
355 0.69) are also important for type I, II and III, while the sensitivity associated with  $\beta(355 \text{ nm})$  is  
356 only remarkable for type I (slope of  $-0.73 \pm 0.04$ ). Sensitivities to biases at  $\beta(532 \text{ nm})$  and  
357  $\beta(1064 \text{ nm})$  are quite low (absolute values of slopes below 0.5).

358 Referring back to Table 1, we observe that the volume concentration ( $V$ ) is the retrieved  
359 integrated parameter least affected by bias in the input optical data as indicated by the fact that  
360 most of the slopes have absolute values below 1.0. All aerosol types I, II and III present moderate  
361 sensitivity to biases in  $\beta(355 \text{ nm})$  with slopes of 0.62-0.92. However, we found differences  
362 among these three different aerosol types. For type I aerosols, the retrieval of volume  
363 concentration is most sensitive to biases in  $\beta(355 \text{ nm})$  (slope of -1.39), while for type II aerosols  
364 retrievals are most sensitive to deviations in  $\alpha(532 \text{ nm})$  (slope of 1.18). For type III aerosols the  
365 sensitivities to bias in the optical data are important both at  $\beta(355 \text{ nm})$  (slope of -1.04) and at  
366  $\alpha(532 \text{ nm})$  (slope up to 1.46). These differences among the aerosol types I, II and III  
367 demonstrate the different sensitivities of volume concentration retrievals when the PSD  
368 possesses different weights of fine and coarse mode.

369 As the regularization scheme used here computes the size distribution using the range of  
370 permitted radii of  $0.075 - 5 \mu\text{m}$ , the fine mode part of the distribution (but not the coarse mode)  
371 is completely covered by this inversion window, and thus we study fine mode volume radius and  
372 fine mode volume concentration. Table 1 also shows the sensitivities of these two parameters to

373 biases in the input optical data. From the slopes of the linear fits reported for  $r_{\text{fine}}$ , biases in  
374  $\alpha(355 \text{ nm})$  and  $\alpha(532 \text{ nm})$  produce significant deviations in the retrieval, with absolute values of  
375 the slopes approximately between 1.0 and 1.5, while the deviations in the retrievals created by  
376 biases in other optical parameters are almost negligible. This result would imply that accurate  
377 retrievals of  $r_{\text{fine}}$  can tolerate rather large errors in the backscatter data but not in the extinction  
378 data. The sign of the slopes of  $r_{\text{fine}}$  as a function of  $\alpha(355 \text{ nm})$  and  $\alpha(532 \text{ nm})$  can be explained by  
379 the same reasoning given before for the effective radius: as extinction at 355 nm increases, it  
380 makes the retrieved particle radius decrease. But as  $\alpha(532 \text{ nm})$  increases the retrieved particle  
381 radius increases. On the other hand, for the fine mode volume concentration ( $V_{\text{fine}}$ ), the largest  
382 sensitivities in the retrieval are found to systematic biases at  $\alpha(355 \text{ nm})$ , with slopes of  $1.59 \pm$   
383  $0.05$ ,  $1.66 \pm 0.17$  and  $1.56 \pm 0.06$  for types I, II and III, respectively. For the other optical  
384 parameters, absolute values of the slopes are below 0.5 (except  $\beta(1064 \text{ nm})$  for type I with slope  
385 of  $0.62 \pm 0.03$ ). These dependencies of the sensitivities of  $r_{\text{fine}}$  and  $V_{\text{fine}}$  to biased input data are  
386 associated with the different dependencies of the kernel functions on wavelength and particle  
387 radius (e.g. Chapter 11 of Bohren and Huffman, 1983).

388 At this point we would like to mention that our simulations (graphs not shown for  
389 brevity) showed some departures from the linearity shown in figures 2 and 3 and Table 1 for  
390 systematic errors larger than approximately  $\pm 30\%$ , mainly when the absolute values of the slopes  
391 is larger than 1. We take this to be an indication that biases of approximately  $\pm 30\%$  and larger  
392 can cause the regularization routine to choose a different solution space than the original retrieval  
393 based on data with no errors. On the other hand, up to errors of  $\pm 20\%$ , we find that the same  
394 minimum in the solution space is generally found by the routine so the linear behavior seen in  
395 Figures 2 and 3 is taken to be a characteristic of a stable system that is displaced from its

396 minimum point. Therefore, we selected a threshold value of  $\pm 20\%$  where these results are  
397 applicable and stress that larger errors in the input data can cause significant and unpredictable  
398 deviations in the retrieved results.

399 Finally, we remark that the values given in Table 1 are averaged for the particular size  
400 distributions used here. More simulations performed (graphs not shown for brevity) changing the  
401 fine mode radius between  $0.08 \mu\text{m}$  and  $0.20 \mu\text{m}$ , both for aerosol type I, II and III, revealed the  
402 same average linear patterns as those shown in Figures 2 and 3 and in Table 1. The only  
403 differences observed were in the absolute values of the slopes with values between  $\pm 20\%$ . On the  
404 other hand, no important departures from the linearity observed in Table 1 were found by  
405 changing the widths of the fine mode. Changes in the coarse mode were not tested because of the  
406 difficulty to assess retrievals of the coarse mode with the methodology used here.

407

### 408 **3.2.1. Effects of the constraints used in the retrievals on the sensitivity test results**

409 The sensitivity tests applied to the different sets of data have shown linear dependencies.  
410 The data presented in Table 1 of the linear fits allows the computation of the deviations induced  
411 in retrieved quantities due to biases in the input data in an easy and straightforward way. **But the**  
412 **generality of the results for different constraints in the inversion code needs to be examined.** For  
413 example, the results presented in Table 1 have been based on a maximum radius in the inversion  
414 ( $r_{\text{max}}$ ) of  $5 \mu\text{m}$ . **Although for the aerosol size distributions studied here this  $r_{\text{max}}$  makes the**  
415 **computation more efficient, the selection of  $r_{\text{max}}$  depends on the user and becomes a constraint in**  
416 **the inversion procedure.** Thus, we performed more simulations with  $r_{\text{max}}$  increased to a value of

417 10  $\mu\text{m}$  to study the influence of this change in constraint on the retrieved results. Another  
418 constraint in the inversion that must be checked is the maximum value allowed for  $m_i$ . We  
419 repeated the simulations allowing  $m_i$  to range up to 0.1 (consistent with a very absorbing aerosol  
420 like black carbon). The results of these studies were compared with a baseline retrieval obtained  
421 with  $r_{\text{max}} = 5 \mu\text{m}$  and with maximum value of  $m_i$  of 0.01. To compute the baseline microphysical  
422 parameters, no induced systematic errors were included. We also computed the retrievals using  
423 the new constraints and introducing systematic errors in the optical data as done before.

424 The new simulations performed after changing the constraints for  $r_{\text{max}}$  and maximum  $m_i$   
425 also reveal linear patterns (graphs not shown for brevity). However, these linear patterns do not  
426 pass through the origin implying that there are generally shifts in the retrieved values of the  
427 various parameters due to these changes in constraints. The analysis reveals, though, that the  
428 signs of the slopes of the linear fits remain the same and that very similar deviations in the  
429 retrieved quantities are computed using the linear fits performed, either with the baseline results  
430 or with those retrieved with the different constraints. Therefore, while the selection of exact  
431 value of the constraints for  $r_{\text{max}}$  and  $m_i$  can change the mean values of the different parameters,  
432 the sensitivity to induced biases in the input optical data is generally unchanged by these changes  
433 in constraints.

### 434 **3.2.2. Additive properties of the effects of systematic errors in the optical data**

435 Thus far, the sensitivity tests that have been performed were based on perturbing a single  
436 optical input at a time. But in a real instrument, it is quite possible that two or more input data  
437 might be influenced by biases simultaneously. Therefore, we need to study the effects of the  
438 presence of multiple simultaneous biases in the input data since the existence of such biases

439 would presumably not be known in a real application. In other words, we wish to determine if  
440 the preceding results based on perturbing a single optical input at a time can be generalized to  
441 predict the effects of multiple input data being simultaneously biased. In particular, we will now  
442 test if, when multiple inputs are simultaneously biased, the results from Table 1 can be used to  
443 calculate deviations that can simply be added to determine the total bias. In other words, we now  
444 will test whether the results in Table 1 can be considered additive.

445 To test the additive properties of the results shown in Table 1, we performed a set of  
446 simulations where two or more optical channels were perturbed simultaneously by biases of the  
447 same magnitude, but allowing different signs (over/under estimation). For example, let's assume  
448 that we have systematic errors of absolute magnitude of 5%. Then different combinations of  $\pm 5\%$   
449 are allowed, as for example at  $\alpha_{355}$  and  $\alpha_{532}$ , at  $\alpha_{355}$  and  $\beta_{532}$  or at  $\beta_{355}$ ,  $\beta_{532}$  and  $\beta_{1064}$ . This  
450 procedure was repeated for different sets of biases of magnitude up to 10%. The deviations noted  
451 as "baseline" were computed using the slopes of Table 1 and assuming that the deviations are  
452 additive. We also performed the regularization retrieval with the new set of data affected by two  
453 or more simultaneous biases, called "simulated deviations". Later we computed the differences  
454 in the microphysical properties based on the slopes given in Table 1 and those actually retrieved  
455 running the code with the new biased optical data and characterized the differences. Using this  
456 procedure, we generated for each absolute value of bias a statistical dataset that includes many  
457 different configurations of the different optical channels. Those datasets are analyzed using Box-  
458 Whisker diagrams as shown in Figure 4 for the effective radius.

459 In these box diagrams the mean is represented by an open square. The line segment in the  
460 box is the median. The top limit represents the 75<sup>th</sup> percentile (P75) and the bottom limit the 25<sup>th</sup>  
461 percentile (P25). The box bars are related to the 1<sup>st</sup> (P1) and 99<sup>th</sup> (P99) percentiles, and the

462 crosses represent the maximum and minimum values, respectively. From Figure 4, for biases of  
463 1, 2, 5 and 10%, mean values of the differences in the effective radius are very small: 0.03, 0.34,  
464 0.41 and 1.01 % for type I (Figure 3a) and -0.62, -0.91, -0.49 and -0.18 % for type II (Figure 3b).  
465 Values larger than the 25<sup>th</sup> percentiles (P25) and lower than the 75% percentiles (P75) are found  
466 for the ranges from -1.8% to 1.3% (type I) and from -0.6% to 4.4% (type II). Only two outliers  
467 are found with relative differences greater than 100%. These last occur when all the optical  
468 channels except  $\beta(355\text{ nm})$  are either overestimated or underestimated. But for these particular  
469 cases the baseline deviations are 0.009 % or -0.009 %, while the simulated ones are 0.557 % and  
470 -0.557 % respectively. These small errors are within the uncertainties associated with the  
471 regularization method, and thus these large relative differences are a mathematical artifact  
472 created by dividing by small numbers. Tests have also been performed for the other  
473 microphysical parameters and we also found an additive property in the deviations predicted by  
474 the results shown in Table 1. Furthermore, very similar additive properties were found for  
475 aerosol type III (graph not shown for brevity). Therefore, for the bimodal size distributions used  
476 here that cover most of those size distributions obtained by AERONET, we conclude that the  
477 results of Table 1 can be reliably used to calculate the deviations in retrieved quantities due to  
478 multiple simultaneously biased input data.

479 We take this result to be an indication that, as mentioned earlier, the solutions found by  
480 the inversion technique generally define a local minimum in the multi-dimensional solution  
481 space (e.g. see Figures 1 of Veselovskii et al., 2002, 2012). The linear behavior of the deviations  
482 in the retrieval due to small changes in the input parameters is a characteristic of displacements  
483 from this minimum location. Multiple simultaneous displacements tend also to display this linear  
484 behavior. The results here indicate, therefore, that for biases in the input data of up to

485 approximately 20%, whether for a single channel or multiple ones simultaneously, the solution  
486 space possesses an average linear property and an additive behavior can be assumed. For larger  
487 biases in the optical data (e.g.  $\pm 30\%$ ) the additive property is not assured as under these  
488 circumstances different minima in the solution space may be found by the regularization  
489 algorithm.

490

491 [Insert Figure 4 here]

492

### 493 **3.3.- Application to the sensitivity of retrievals to the presence of random** 494 **errors in the optical data.**

495 Up to this point, we have concerned ourselves only with the effects of systematic biases  
496 in the input optical data on the retrieved quantities. But in lidar systems, random errors are also  
497 present due just to the measurement process itself. Any specific set of 3+2 data affected by  
498 random errors can be considered as a set of biased measurements where the individual biases for  
499 each of the data follow a normal distribution. Given the additive property of the systematic errors  
500 that we have shown, we can assess the effects of random errors in the optical data by generating  
501 random biases in the optical data and computing their deviations in the microphysical parameters  
502 from the values given in table 1. The sensitivities of the regularization technique to those random  
503 errors computed using the procedure just outlined will be compared with previously published  
504 ones [e.g. Müller et al., 1999a,b; Veselovskii et al., 2002, 2004].

505 To assess the sensitivity of the retrievals to random errors we use the additive properties  
506 of the systematic biases just described. The procedure used consists of generating random  
507 numbers distributed in a Gaussian way centered at zero with width according to the value of the

508 random error to study. These random errors are applied to each optical channel of the  $3\beta + 2\alpha$   
509 configuration. This procedure was repeated 50,000 times for each parameter studied. Also, the  
510 initiation of the random number generation is different for each channel to avoid the situation  
511 where all the random numbers are the same in every channel. Finally, we introduced for every  
512 optical data this random number and computed the corresponding error in the retrieved  
513 microphysical parameter using the slopes provided in Table 1. For every set of  $3\beta + 2\alpha$  values,  
514 the final error obtained in the microphysical parameter is the sum of the error obtained for each  
515 channel. The study of the frequency distributions of the final errors for this large number of  
516 simulations yields the effects of random errors. If the frequency distribution is a normal one, the  
517 standard deviation (Full-Width-at-Half-Maximum) provides the final error in the microphysical  
518 parameter. Moreover, if the normal distribution is not centered at zero it demonstrates an  
519 interesting property; that the presence of systematic errors in the retrieved microphysical  
520 property can be induced by random errors in the input optical data. As an illustration, Figure 5  
521 shows the frequency distribution of the differences in the microphysical parameters studied here,  
522 for all aerosol size distributions type I, II and III, where 15% random error is assumed in all the  
523 optical data. Those differences are in percentages and denoted as ‘deviation’ in the ‘x’ axis of the  
524 histograms.

525

526

[Insert Figure 5 here]

527

528 From Figure 5 we observe that the frequency distributions possess the expected Gaussian  
529 shape for all the microphysical parameters. Most of the frequency distributions are centered



530 essentially at zero, although some significant departures from this value are observed in the cases  
531 where random errors can induce systematic biases in the retrieved aerosol microphysical  
532 parameters. The percentage changes in the mean values of the distributions are shown in the  
533 legend. A shift in the mean value due to the presence of random error results for those retrieved  
534 parameters that display a different linear tendency for positive and negative biases in the input  
535 optical data as discussed earlier with respect to Table 1. For example, such departures from zero  
536 are observed for retrievals of  $r_{\text{eff}}$ ,  $N$  and  $V$  for type II aerosols and are approximately -5, 1, and -7  
537 %, respectively. On the other hand, the FWHM –or standard deviations- of the normal  
538 distributions of Figure 5 are representative of the sensitivities to 15% random errors in the  
539 optical data. Generally, there are many similarities in the standard deviations between aerosol  
540 types I, II and III. We observe clearly that  $V$ ,  $r_{\text{fine}}$  and  $V_{\text{fine}}$  exhibit the smallest sensitivity to the  
541 imposed 15% random errors with a 1-sigma spread in the result of approximately 25%. The  
542 effective radius and surface concentration results show moderate sensitivity with 1-sigma values  
543 of ~ 30 – 40 %, while the retrieval of number concentration has the highest sensitivity, with 1-  
544 sigma values of 67.6% for type I, 95.2% for type II and 61.4% for type III. As expected, these  
545 sensitivities to random error track the results of the sensitivities to systematic errors, where the  
546 most sensitive parameter was also found to be number concentration and the least were volume  
547 concentration, fine mode radius and fine mode volume concentration.

548         Using the same procedure as for 15% random error, Table 2 reports the FWHM –or  
549 standard deviations- of normal distributions obtained for other magnitudes of random errors in  
550 the optical data ranging from 5% to 20%. We observe, as expected from the linear functions  
551 involved, that increasing the random uncertainty increases the deviations found in a linear  
552 fashion. Moreover, it is observed again that the largest sensitivities are for  $N$  while the lowest are

553 for  $V$ ,  $r_{\text{fine}}$  and  $V_{\text{fine}}$ . In the same way, Table 3 reports the means of the deviation of every  
554 microphysical property for varying amounts of random uncertainty in the input data. As  
555 mentioned above, the departures of these deviations from zero indicate that random uncertainties  
556 in the input optical data can induce varying amounts of systematic bias in the retrieved  
557 properties. This effect is found more with the type II aerosols that possess a higher fraction of  
558 larger particles. Such a population is more likely to have different slopes in Table 1 due to  
559 positive and negative biases in the input optical data because of the reduced sensitivity of the  
560 MW technique to larger particles. It is this reduced sensitivity to larger particles that, in general,  
561 explains the shifting of the mean values in the retrieved distributions due to varying amounts of  
562 random error in the input data.

563 [Insert Table 2 here]

564 [Insert Table 3 here]

565 Müller et al., [1999a,b] and Veselovskii et al., [2002, 2004] studied 10% random  
566 uncertainties in the optical data in the  $3\beta + 2\alpha$  lidar configurations by introducing random errors  
567 in the optical data and running the regularization code repeatedly. These studies reported that the  
568 retrieved uncertainties were on the order of 25% for  $r_{\text{eff}}$ ,  $V$  and  $S$ , 30% for  $r_{\text{mean}}$  and 70% for  $N$ .  
569 These values are quite similar to those reported in Table 2 for our computations of 10% random  
570 errors. No evaluations for  $r_{\text{fine}}$  and  $V_{\text{fine}}$  were done in the studies of Müller et al., [1999a,b] and  
571 Veselovskii et al., [2002, 2004]. The method shown here for assessing the sensitivity of  
572 retrievals to random errors is generally consistent with these earlier results but permits the  
573 influence of varying amounts of random error to be studied. It also permits the influence of

574 random errors in different input optical channels to be quantified. We will now apply this  
575 capability to the problem of instrument specification.

576

### 577 **3.3.1. Application to instrument specification.**

578 The upcoming space-borne Decadal Survey ACE (Aerosol-Cloud-Ecosystems) mission  
579 of NASA (<http://dsm.gsfc.nasa.gov/ace/>) specifies a High Spectral Resolution Lidar as a core  
580 instrument to measure vertical profiles of aerosol extinction and backscattering worldwide.  
581 These profiles will be used to derive vertically-resolved aerosol microphysical properties such as  
582 effective radius, number concentration or complex refractive index. The system is anticipated to  
583 use the  $3\beta+2\alpha$  configuration and the regularization technique that has been studied here. The first  
584 reports (<http://dsm.gsfc.nasa.gov/ace/>) call for an accuracy of  $\pm 15\%$  for all backscattering and  
585 extinction coefficients, and thus the results presented here can be used to infer the anticipated  
586 uncertainties in the microphysical properties retrieved using the regularization technique on these  
587  $3\beta+2\alpha$  space borne data when all input data possess 15% uncertainties. The results already  
588 presented clearly indicate, however, that for most quantities it is uncertainties in the extinction  
589 coefficients that need to be constrained more carefully than those in the backscattering data.  
590 Volume concentration is an interesting exception to this statement where  $\beta(355\text{ nm})$  for type I  
591 aerosols is the optical parameter requiring the smallest uncertainty budget to help reduce the  
592 uncertainties in retrievals. In this way, the results obtained here can serve as a guide to hardware  
593 designers of multi-wavelength lidar instruments in the sense that if trade-offs need to be made  
594 between the performance of one optical channel versus another, the relative sensitivities shown  
595 in Table 1 can be used to assess which channels would benefit most from decreased uncertainty

596 in the measurements. Another application of the sensitivities derived here is to algorithm  
597 development. Algorithms can introduce systematic uncertainties in the optical data such as  
598 through an incorrect assumption of an aerosol free region, an assumption of the extinction to  
599 backscatter ratio or the use of an estimated molecular profile. The results presented here can be  
600 used to assess the tolerance for both random and systematic errors in the input optical data due  
601 both to instrumentation and to algorithms once uncertainty requirements in the retrieved  
602 quantities are established.

#### 603 **4.- SUMMARY AND CONCLUSIONS**

604 We have presented the results of a study of the sensitivity of the retrievals of aerosol  
605 physical parameters using the regularization technique to systematic and random uncertainties in  
606 the input optical data. We have focused our study on the set of data consisting of three  
607 backscattering coefficients ( $\beta$ ) at 355, 532 and 1064 nm and two extinction coefficients ( $\alpha$ ) at  
608 355 and 532 nm ( $3\beta + 2\alpha$  configuration). These data can be obtained by a lidar system that uses a  
609 Nd:YAG laser and combines backscatter with Raman or HSRL channels. Simulations have been  
610 done for different bimodal aerosol size distributions that are representative of AERONET  
611 climatologies. The values used for aerosol refractive indexes, as well as mode radius and widths  
612 were selected as representative of those climatologies as well. The selected aerosol bimodal size  
613 distributions include one with fine mode predominance (type I), another with predominance of  
614 coarse mode but with significant presence of fine mode (type II) and another with predominance  
615 of fine mode but with significant presence of coarse mode (type III). Optical data consistent with  
616 these bimodal size distributions were generated using Mie theory. Retrievals were performed  
617 using these baseline optical data. The optical data were then perturbed by systematic biases in the

618 range  $\pm 20\%$  to study the effects of biases on the retrieved parameters. This threshold value of  
619  $\pm 20\%$  is enough for many practical lidar applications. As the problem of the inversion of  
620 microphysical properties is under-determined, constraints are needed which, in principle, can  
621 influence the values retrieved by the algorithm. Particularly, we have found that the range of  
622 radius and refractive index used in the inversion did not have a large influence on the  
623 sensitivities of the different microphysical particles. However, our results showed that the  
624 maximum value of  $m_i$  allowed in the retrieval had a significant influence on the value of  
625 refractive index retrieved, supporting earlier results that indicate significant uncertainties in the  
626 retrieval of refractive index using the  $3\beta+2\alpha$  MW lidar technique studied here.

627         The microphysical parameters studied included effective radius ( $r_{\text{eff}}$ ) and volume ( $V$ ),  
628 surface ( $S$ ) and number ( $N$ ) concentration. Also, as the inversion window ranged from 0.075 to 5  
629  $\mu\text{m}$ , we were able to study the fine mode of the aerosol size distribution (0.075-0.5  $\mu\text{m}$ )  
630 separately, and thus we have also presented the results for both fine mode radius ( $r_{\text{fine}}$ ) and  
631 volume ( $V_{\text{fine}}$ ). From these sensitivity tests, the percentage deviations of the microphysical  
632 parameters as function of biases in the optical data presented linear patterns. Generally, these  
633 linear patterns presented the same sign of slopes for aerosol type I, II and III and the largest  
634 sensitivities were observed for biases in the extinction coefficients  $\alpha(355 \text{ nm})$  and  $\alpha(532 \text{ nm})$ .  
635 Moreover, the largest sensitivities were found for  $N$ , while the least affected parameters were  $V$ ,  
636  $r_{\text{fine}}$  and  $V_{\text{fine}}$ .

637         An important result is that we have found an additive property for the deviations induced  
638 by the biases in the optical data. This implies that if, for example, several optical data are  
639 simultaneously affected by systematic errors, the total deviation in the retrieved quantity can be

640 well approximated by the sum of those deviations computed when each optical input was biased  
641 separately. From this additive property, we have been able to compute the effects of random  
642 errors in the optical data. The largest sensitivities to random errors were found for  $N$ , while the  
643 lowest were obtained for  $V$ ,  $r_{\text{fine}}$  and  $V_{\text{fine}}$ . Moreover, we have found some systematic differences  
644 in the mean retrieved microphysical properties when the retrievals are affected by random errors  
645 in the input optical data. The presence of these systematic differences is associated with the  
646 different behavior (although with linear patterns) between positive and negative biases in the  
647 input optical data and is due to a reduced sensitivity of the retrieval to the coarse part of the size  
648 distribution.

649       The results presented here cannot be generalized to every possible size distribution as we  
650 only focused on bimodal size distributions representative of those obtained by AERONET.  
651 Studies of the sensitivities of the microphysical retrieval to errors in the optical data for other  
652 size distributions such as, for example, one showing tri-modal behavior are still needed although  
653 the results presented here for three differing bi-modal distributions leads one to expect that  
654 similar results would be obtained for tri-modal distributions as well. The tests performed here  
655 showed that the average linearity of the sensitivities in the retrieval to random errors in the input  
656 data can be useful for a wide range of lidar applications, and thus can be used to establish  
657 acceptable error budgets in optical data if maximum permissible errors in the retrieved quantities  
658 can be established. Therefore, the values given here for the sensitivities of the microphysical  
659 properties to systematic errors in the optical data can be useful for many lidar applications. For  
660 example, for the Decadal Survey ACE mission a multi-wavelength lidar is planned. Among their  
661 measurement requirements is that the accuracy of the optical data be  $\pm 15\%$ . If these uncertainties  
662 are taken to be all random, we were able to use the results here to estimate that this implies an

663 uncertainty in the retrieved microphysical properties by the regularization technique of ~40% for  
664  $r_{\text{eff}}$ , ~85% for N, ~25% for S, ~20% for V, 16% for  $r_{\text{fine}}$  and  $V_{\text{fine}}$  respectively. The results also  
665 permit assessing the deviations in the retrievals if the biases in the optical data are systematic and  
666 exist in only one or more channels. In this way, trade-off decisions can be made between the  
667 retrieval requirements and the hardware configuration of a lidar system taking into account the  
668 different sensitivities of the retrievals to biases in the optical data of different channels. We hope  
669 these results aid the future design of multi-wavelength lidar systems intended for retrieval of  
670 aerosol microphysical properties.

671

672

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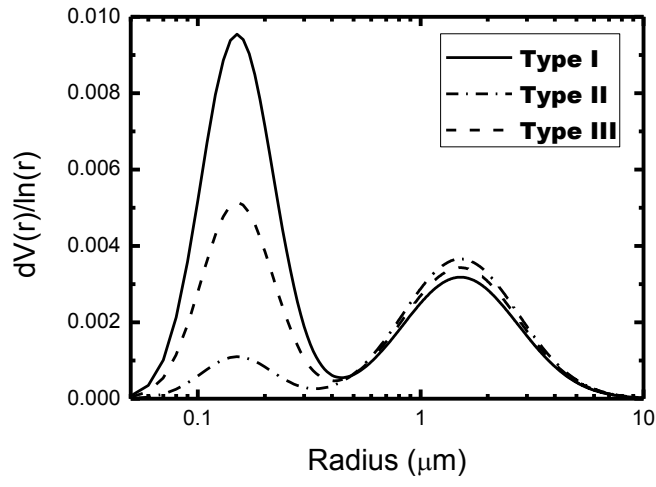
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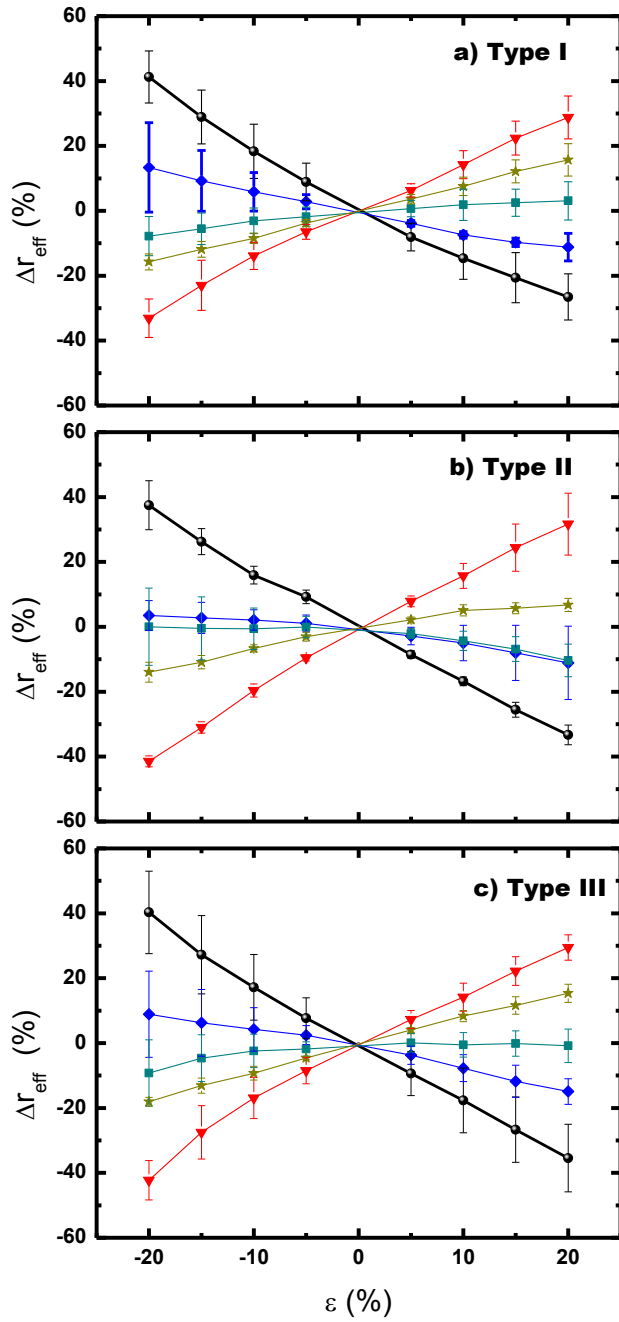
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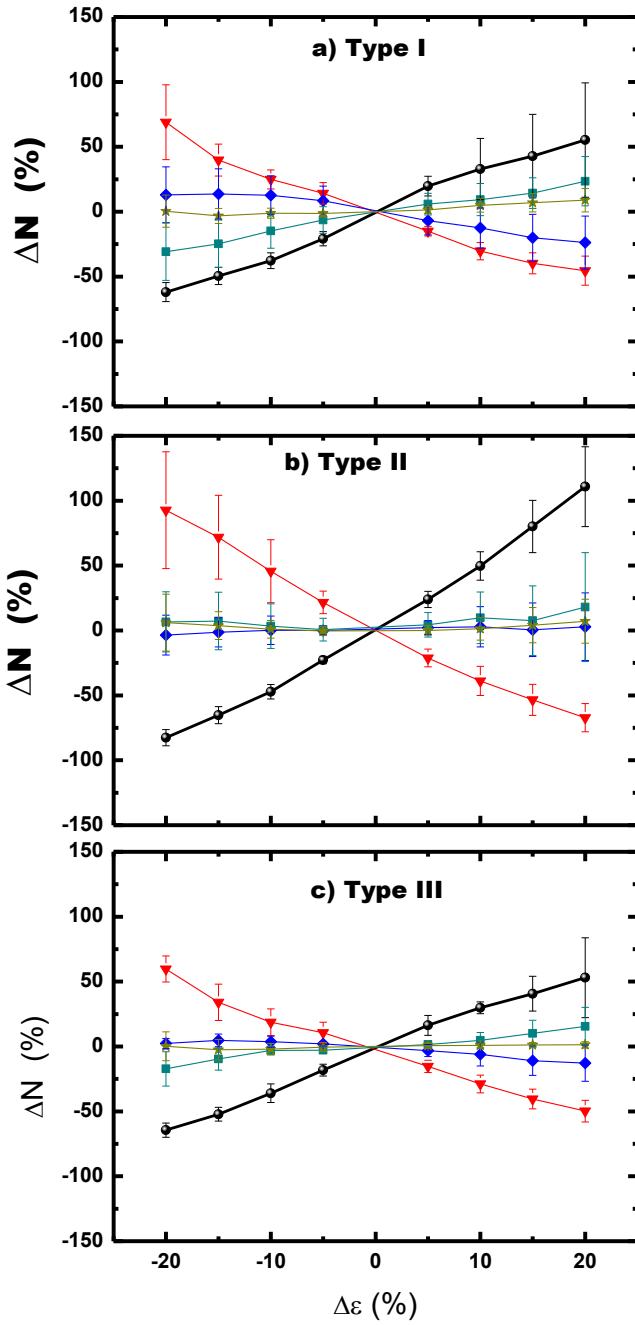
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**Figure 1:** Normalized size distributions used for computing the simulated optical data. The ratio between the volume of fine and coarse mode,  $V_{tc}/V_{tc}$ , is 2 for type I, 0.2 for type II and 1 for type III.

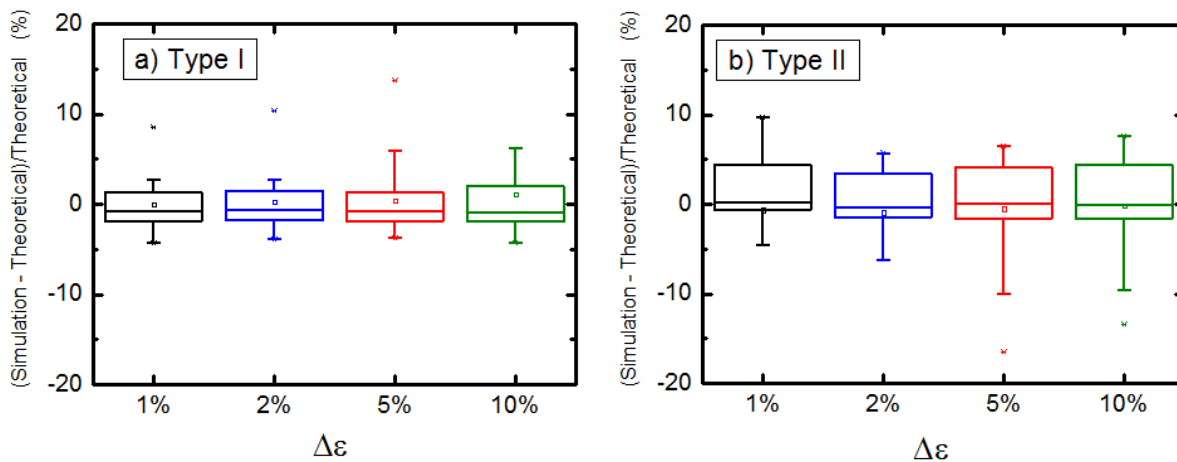


**Figure 2:** Percentage deviation of the effective radius as a function of systematic bias in the optical data ( $\epsilon$ ). a) Type I. b) Type II. C) Type III.

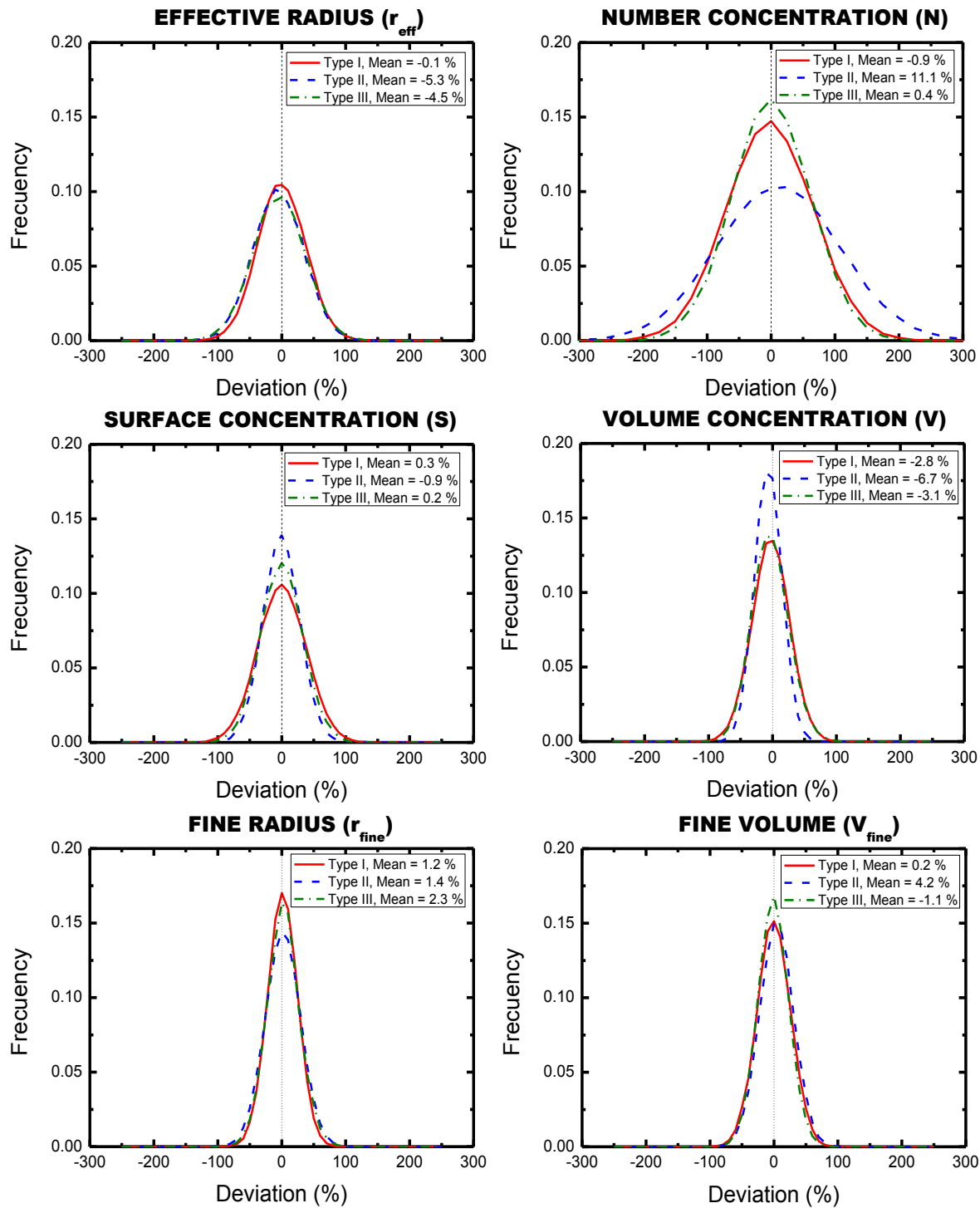


**Figure 3:** Percentage deviation of the number concentration as a function of systematic bias in the optical data ( $\epsilon$ ). a) Type I. b) Type II. c) Type III.





**Figure 4:** For the effective radius, Box-Whisker diagrams of the differences between the theoretical deviations computed with the slopes of table 1 and the simulated deviations. At least two optical channels have been simultaneously perturbed by biases of the same magnitude although different combinations of over/under estimations are allowed. In these box diagrams the mean is represented by an open square. The line segment in the box is the median. The top limit represents the 75<sup>th</sup> percentile (P75) and the bottom limit the 25<sup>th</sup> percentile (P25). The box bars are related to the 1<sup>st</sup> (P1) and 99<sup>th</sup> (P99) percentiles, and the crosses represent the maximum and minimum values respectively. We used biases in the optical data of 1% (black diagrams), 2% (blue diagrams), 5% (red diagrams) and 10% (green diagrams).



**Figure 5:** Frequency distributions of the different microphysical parameters for 15% random errors in the optical data using 50000 random samplings of the systematic error sensitivities shown in Table 1. The ‘x’ axis represents the difference between microphysical parameters with no errors in the input optical data and those affected by random errors in the optical data. Random errors were simulated by a normal distribution centred at zero and with standard deviation of 15%. The random number generator is initialized at different values for each of the 5 optical data used in the  $3\beta + 2\alpha$  lidar configuration. The mean value of the deviation between the microphysical parameter affected by random error and that unaffected by random error is included in the legend.

		$\frac{\Delta r_{eff}(\%)}{\Delta \varepsilon(\%)}$	$\frac{\Delta N(\%)}{\Delta \varepsilon(\%)}$	$\frac{\Delta S(\%)}{\Delta \varepsilon(\%)}$	$\frac{\Delta V(\%)}{\Delta \varepsilon(\%)}$	$\frac{\Delta r_{fine}(\%)}{\Delta \varepsilon(\%)}$	$\frac{\Delta V_{fine}(\%)}{\Delta \varepsilon(\%)}$
$\alpha(355 \text{ nm})$	<b>Type I</b>	$-1.68 \pm 0.12$	$3.09 \pm 0.12$	$2.08 \pm 0.05$	$0.26 \text{ (p)} / 0.77 \text{ (n)} \pm 0.07$	$-0.99 \pm 0.11$	$1.59 \pm 0.05$
	<b>Type II</b>	$-1.74 \pm 0.03$	$4.83 \pm 0.22$	$1.77 \pm 0.04$	$-0.37 \text{ (p)} / 0.35 \text{ (n)} \pm 0.05$	$-1.27 \pm 0.17$	$1.66 \pm 0.17$
	<b>Type III</b>	$-1.84 \pm 0.04$	$3.04 \pm 0.13$	$1.95 \pm 0.05$	$-0.47 \text{ (p)} / 0.77 \text{ (n)} \pm 0.04$	$-0.64 \text{ (p)} / -1.51 \text{ (n)} \pm 0.07$	$1.56 \pm 0.06$
$\alpha(532 \text{ nm})$	<b>Type I</b>	$1.51 \pm 0.04$	$-2.78 \pm 0.17$	$-1.07 \pm 0.08$	$0.44 \pm 0.12$	$1.17 \pm 0.04$	$-0.28 \pm 0.05$
	<b>Type II</b>	$1.82 \pm 0.09$	$-4.09 \pm 0.23$	$-0.69 \pm 0.03$	$1.18 \pm 0.17$	$1.28 \pm 0.07$	$-0.44 \pm 0.04$
	<b>Type III</b>	$1.71 \pm 0.10$	$-2.61 \pm 0.12$	$-0.92 \pm 0.07$	$1.46 \pm 0.08 \text{ (p)} / 0.77 \text{ (n)} \pm 0.02$	$0.98 \text{ (p)} \pm 0.01 / 1.46 \text{ (n)} \pm 0.01$	$-0.20 \pm 0.04$
$\beta(355 \text{ nm})$	<b>Type I</b>	$-0.63 \pm 0.02$	$-1.25 \pm 0.04 \text{ (p)} / -0.85 \pm 0.15 \text{ (n)}$	$-0.73 \pm 0.04$	$-1.39 \pm 0.04$	$-0.01 \text{ (p)} / -0.06 \text{ (n)} \pm 0.01$	$-0.62 \pm 0.03$
	<b>Type II</b>	$-0.54 \text{ (p)} / -0.18 \text{ (n)} \pm 0.01$	$0.19 \text{ (p)} / 0.12 \text{ (n)} \pm 0.04$	$-0.22 \text{ (p)} / -0.04 \text{ (n)} \pm 0.02$	$-0.48 \pm 0.10$	$0.33 \text{ (p)} / 0.06 \text{ (n)} \pm 0.03$	$0.26 \text{ (p)} / -0.01 \text{ (n)} \pm 0.01$
	<b>Type III</b>	$-0.76 \text{ (p)} / -0.43 \text{ (n)} \pm 0.01$	$-0.44 \pm 0.08$	$-0.47 \pm 0.06$	$-1.04 \pm 0.08$	$0.10 \pm 0.01$	$-0.39 \text{ (p)} / -0.19 \text{ (n)} \pm 0.01$
$\beta(532 \text{ nm})$	<b>Type I</b>	$0.27 \pm 0.04$	$1.3 \pm 0.09$	$0.50 \pm 0.03$	$0.77 \pm 0.05$	$-0.05 \text{ (p)} / -0.22 \text{ (n)} \pm 0.03$	$0.22 \pm 0.02$
	<b>Type II</b>	$-0.48 \text{ (p)} / 0.02 \text{ (n)} \pm 0.02$	$0.79 \pm 0.11 \text{ (p)} / -0.37 \pm 0.05 \text{ (n)}$	$0.05 \pm 0.02$	$-0.38 \text{ (p)} / 0.03 \text{ (n)} \pm 0.03$	$-0.11 \pm 0.02$	$-0.11 \text{ (p)} / -0.34 \text{ (n)} \pm 0.01$
	<b>Type III</b>	$-0.03 \text{ (p)} / 0.38 \text{ (n)} \pm 0.05$	$0.70 \pm 0.06$	$0.30 \pm 0.03$	$0.48 \pm 0.07$	$-0.16 \pm 0.01$	$0.02 \pm 0.02$
$\beta(1064 \text{ nm})$	<b>Type I</b>	$0.79 \pm 0.01$	$0.37 \pm 0.05$	$0.17 \pm 0.02$	$0.92 \pm 0.04$	$-0.17 \pm 0.01$	$-0.04 \pm 0.01$
	<b>Type II</b>	$0.54 \pm 0.07$	$0.29 \text{ (p)} / -0.25 \text{ (n)} \pm 0.05$	$0.04 \pm 0.02$	$0.60 \pm 0.05$	$-0.28 \pm 0.02$	$-0.15 \text{ (p)} / -0.34 \text{ (n)} \pm 0.02$
	<b>Type III</b>	$0.84 \pm 0.02$	$0.07 \pm 0.03$	$0.08 \pm 0.02$	$0.92 \pm 0.03$	$-0.26 \pm 0.01$	$-0.19 \pm 0.01$

**TABLE 1:** Percentage deviations in the aerosol microphysical properties as a function of systematic errors in the optical data  $\varepsilon$ . Particularly, the slopes 'a' of the linear fits  $Y = aX$  are presented, where 'X' is the systematic bias in the optical data and Y is the corresponding deviation in the microphysical properties. All these fits presented linear determination coefficient  $R^2 > 0.90$ . For the cases when there is a difference in slope between positive and negative biases in the input data, the slopes relating to the positive biases are indicated by (p) while those associated with negative biases are indicated by (n).

**TABLE 2:** Standard deviations of the frequency distributions of the deviation induced in the microphysical parameters due to random errors in the optical data.

Random Errors (%)	$r_{\text{eff}}$			N			S			V			$r_{\text{fine}}$			$V_{\text{fine}}$		
	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III
<b>5</b>	12.5	13.1	13.7	22.5	31.8	20.5	12.5	9.5	11.2	9.8	7.2	9.5	7.7	9.2	8.4	8.7	8.8	8.1
<b>10</b>	24.9	26.2	27.2	45.0	63.6	40.8	25.1	19.1	22.3	19.6	14.4	19.0	15.5	18.4	16.8	17.4	17.6	16.1
<b>15</b>	37.2	39.2	40.8	67.6	95.2	61.4	37.7	28.5	33.4	29.5	21.5	28.5	23.3	27.6	25.3	26.1	26.3	24.1
<b>20</b>	50.0	52.6	54.8	90.1	127.3	82.1	50.2	38.2	44.6	39.3	28.8	38.0	31.1	36.9	33.8	34.9	35.2	32.2
<b>10*</b>	25*			70*			25*			25*			-----			-----		

\*From the previous work of Muller et al., (1999a,b) and Veselovskii et al., (2002, 2004).

**TABLE 3:** Mean of the differences (in percentages) in the retrieved microphysical parameters due to varying amounts of random error in the optical data.

Random Errors (%)	$r_{\text{eff}}$			N			S			V			$r_{\text{fine}}$			$V_{\text{fine}}$		
	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III	Type I	Type II	Type III
<b>5</b>	0.0	-1.7	-1.6	-0.8	3.5	0.2	0.0	-0.4	0.1	-1.1	-2.3	-1.1	0.4	0.5	0.7	0.0	1.4	-0.3
<b>10</b>	0.0	-3.5	-3.0	-1.4	7.1	0.1	0.1	-0.7	0.0	-1.9	-4.4	-2.3	0.9	1.1	1.5	0.1	2.8	-0.8
<b>15</b>	-0.1	-5.3	-4.5	-1.9	11.1	0.4	0.3	-0.9	0.2	-2.8	-6.7	-3.1	1.2	1.4	2.1	0.2	4.2	-1.0
<b>20</b>	-0.3	-7.2	-5.6	-2.3	15.2	-0.6	0.6	-1.0	-0.4	-3.8	-9.0	-4.5	1.5	1.8	3.3	0.4	5.8	-1.9