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# Retrieval of three-dimensional small scale structures in upper tropospheric/lower stratospheric composition as measured by GLORIA

M. Kaufmann<sup>1</sup>, J. Blank<sup>1,\*</sup>, T. Guggenmoser<sup>1</sup>, J. Ungermann<sup>1</sup>, A. Engel<sup>7</sup>, M. Ern<sup>1</sup>, F. Friedl-Vallon<sup>2</sup>, D. Gerber<sup>4</sup>, J. U. Grooss<sup>1</sup>, G. Guenther<sup>1</sup>, M. Höpfner<sup>2</sup>, A. Kleinert<sup>2</sup>, Th. Latzko<sup>2</sup>, G. Maucher<sup>2</sup>, T. Neubert<sup>9</sup>, H. Nordmeyer<sup>2</sup>, H. Oelhaf<sup>2</sup>, F. Olschewski<sup>3</sup>, J. Orphal<sup>2</sup>, P. Preusse<sup>1</sup>, H. Schlager<sup>8</sup>, H. Schneider<sup>9</sup>, D. Schuettemeyer<sup>6</sup>, F. Stroh<sup>1</sup>, O. Suminska-Ebersoldt<sup>2</sup>, B. Vogel<sup>1</sup>, C. M. Volk<sup>3</sup>, W. Woiwode<sup>2</sup>, and M. Riese<sup>1</sup>

<sup>1</sup>Institute of Energy and Climate Research, Stratosphere (IEK-7), Research Centre Jülich, 52425 Jülich, Germany

<sup>2</sup>Institute for Meteorology and Climate Research, Karlsruhe Institute of Technology, Karlsruhe, Germany

<sup>3</sup>Physics Department, University of Wuppertal, 42097 Wuppertal, Germany

<sup>4</sup>RAL Space, STFC Rutherford Appleton Laboratory, Harwell Oxford, Didcot, OX11 0QX, UK

<sup>6</sup>European Space Agency, Mission Science Division, 2200AG Noordwijk, the Netherlands

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<sup>7</sup>Institute for Atmospheric and Environmental Sciences, Goethe University Frankfurt, 60438 Frankfurt/Main, Germany

<sup>8</sup>German Aerospace Center, Institute of Atmospheric Physics, Atmospheric trace gases, 82234 Oberpfaffenhofen-Wessling, Germany

<sup>9</sup>Central Institute of Engineering, Electronics and Analytics (ZEA-2) – Electronic Systems, Research Centre Jülich, Jülich, Germany

\*now at: Google Inc., Dublin, Ireland

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Correspondence to: M. Kaufmann (m.kaufmann@fz-juelich.de)

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Tomographic measurements are performed in the so called dynamics mode. In this mode, GLORIA operates at low optical path differences to obtain more spectra at a given time. During ESSenCe GLORIA recorded spectra at a maximum optical path difference of 1.6 cm in 2.8 s for a pair of forward/backward spectra, yielding a spectral sampling of  $0.3125 \text{ cm}^{-1}$ . During TACTS/ESMVal dynamics mode spectra were recorded at a maximum optical path difference of 0.8 cm in 2 s for a pair of forward/backward spectra, yielding a spectral sampling of  $0.625 \text{ cm}^{-1}$ . Another measurement mode is the so called chemistry mode with a spectral sampling of  $0.0625 \text{ cm}^{-1}$  and 12 s measurement time. GLORIA data recorded in chemistry mode are analyzed by Woiwode et al. (2014). Individual GLORIA images contain  $128 \times 48$  pixel/spectra. The horizontal resolution of the detector array is used to analyze and filter cloud fractions. After filtering, the spectra are averaged horizontally on a  $128 \times 1$  grid for further processing. Different viewing angles of the same air mass are obtained by pointing the instrument from  $48$  to  $118^\circ$  ( $4^\circ$  steps) with respect to the aircraft nose in a short time frame. The data obtained in this way can be combined and processed in a tomographic way to get a three-dimensional picture of the atmosphere.

### 3 Data processing

When operating the GLORIA instrument during a field measurement, raw interferograms and house keeping data are written to a mass storage for later processing. The data rate is about  $1 \text{ Gbits}^{-1}$ . In the laboratory, this data is processed in several steps, including different calibration procedures and the retrieval of atmospheric temperature, trace gases, etc. These steps and the corresponding data level products are described below.

### 3.1 Level 0 and Level 1

The Level 0 and Level 1 processing transforms the measured raw interferograms into radiometrically and spectrally calibrated spectra. The core of the Level 0 processing is the resampling of the interferograms. The raw interferograms are sampled on a time-equidistant grid. Before Fourier transformation, they have to be interpolated onto a space-equidistant grid, taking velocity variations of the interferometer drive into account. The Level 0 processing also includes a quality check of the data, spike detection and correction, non-linearity correction, correction of phase errors, and the spectral calibration. The Level 1 processing comprises the Fourier transform converting the interferograms to complex spectra and the radiometric calibration. Calibration spectra are generated from blackbody and quasi deep space measurements using a two-point calibration. Details about the Level 0 and Level 1 processing of GLORIA data are given in Kleinert et al. (2014).

### 3.2 Level 2

Level 2 processing encompasses the derivation of atmospheric or instrumental parameters from calibrated radiance measurements. This is performed in an iterative way by minimizing the difference between a forward model mapping an atmospheric state to radiances as they would have been observed by the instrument and real GLORIA measurements. This minimization is performed by the so-called inversion model.

#### 3.2.1 Forward modeling

The forward model consists of a radiative transfer and an instrument model. The radiative transfer model is based on a combination of the emissivity growth approximation (EGA; Gordley and Russell, 1981) and the Curtis–Godson approximation (CGA; Curtis, 1952; Godson, 1953). Both methods are based on pre-calculated emissivities of homogeneous gas cells for a variety of atmospheric conditions and pre-defined spectral

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ranges. The spectral ranges are either individual spectral points or integrated spectral windows combining several spectral points (Riese et al., 1997, 1999). The optical paths have been calculated by means of the line-by-line Reference Forward Model (RFM, Dudhia, 2000) utilizing the HITRAN-2012 spectral database (Rothman et al., 2013). The speed up of this method is up to three orders of magnitude in comparison to a line-by-line model and the uncertainties arising from the approximations are generally less than one percent (Marshall et al., 1994; Francis et al., 2006; Ungermann et al., 2013).

The second part of the forward model is the instrument model. It maps various line of sight radiances onto the detector array. To consider field of view effects, several line of sight radiances are typically combined to simulate one or more detector pixel.

The retrieval of atmospheric quantities requires the calculation of the Jacobian of the forward model (see Sect. 3.2.2). This is done by algorithmic differentiation (Lotz et al., 2011; Ungermann et al., 2011), which is highly efficient in our case, because the computational costs for calculating the Jacobian matrix are only a constant multiple of a single execution of the forward model.

### 3.2.2 Inverse modeling

The reconstruction of an atmospheric state compatible with a sequence of GLORIA measurements is a non-linear inversion problem. This is solved by minimizing a cost function  $\sigma_y$  describing the difference between the forward model  $F: \mathbb{R}^n \rightarrow \mathbb{R}^m$  of an atmospheric state  $\mathbf{x} \in \mathbb{R}^n$  (temperature, trace gas abundance, etc.) and the radiance measurement  $\mathbf{y} \in \mathbb{R}^m$ :

$$\sigma_y = (F(\mathbf{x}) - \mathbf{y})^T \mathbf{S}_e^{-1} (F(\mathbf{x}) - \mathbf{y}) \quad (1)$$

The difference is weighted by the inverse measurement error/error covariance matrix  $\mathbf{S}_e^{-1}$ . In most cases, the solution is not unique and/or under-constrained. Therefore it





purpose, the Russian M55 Geophysica was chosen, which has a ceiling altitude of about 21 km. The aircraft is highly flexible in terms of operations, ground weather as well as stratospheric conditions. The campaign base was Kiruna in northern Sweden.

The first part of the ESSenCe campaign consisted of ground based measurements inside and outside of the hangar Arena Arctica to assure the interaction of all subsystems of the GLORIA instrument after transport and final integration. Several certification and calibration measurements were conducted, including electromagnetic tests of GLORIA when mounted on Geophysica.

The second part of the campaign consisted of flights with full instrumentation under stratospheric conditions. Flight patterns were planned such that the functionality and performance of GLORIA could be tested under all major conditions that are expected during later research flights. This includes particularly electromagnetic compatibility considerations (under flight conditions), the thermal behavior of the sensor and the gimbal frame, the detector and instrument optics cooling system, the pointing system, and the interferometer control system.

The dynamic conditions of the Arctic upper troposphere and lower stratosphere during the ESSenCe campaign were mostly undisturbed. At the beginning of the campaign in early December 2011 a low pressure system over the Norwegian Sea was slowly approaching Scandinavia. East of this low, mid-latitude air was moved polewards, thereby being lifted isentropically. Connected to the lows frontal system high clouds appeared south of this tongue of air (Fig. 1, blue lines). On the western rim of the low pressure system air from the lower stratosphere was descending deep into the troposphere, coiling up cyclonically. A few days later, the polar vortex strengthened and showed very low core temperatures with values well below 190 K in the lower stratosphere. Due to large-scale disturbances the vortex center was moved towards the European sector and air from the vortex edge was moved polewards over Siberia, thus forming a broad intrusion of air with low potential vorticity reaching from Siberia across the north pole to the Norwegian Sea.

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## 4.2 GLORIA operations

Two GLORIA flights were performed during ESSenCe. Flight 1 took place on 11 December, 2011, 11:00–15:00 UTC, and Flight 2 on 16 December, 2011, 14:00–17:50 UTC, respectively. Although the instrument encountered some technical problems, partly related to the extremely cold temperatures in the polar vortex, GLORIA took nominal atmospheric measurements including calibration for more than one hour during each flight. During Flight 1, all measurements were taken in chemistry mode while during Flight 2, both chemistry and dynamics mode measurements were performed.

Due to interferometer slide velocity variations during the first flight, these data were utilized for engineering purposes, only. The functionality of the gimbal frame was partly lost in both flights after 1–2 h of flight time. For the remaining flight time it was not possible to compensate the image rotation anymore. The loss of pointing and image stabilization capabilities impacts data quality significantly such that Level 0 to Level 2 processing was restricted to the first measurement period of the 2nd flight. The flight path of the Geophysica aircraft and GLORIA instrument operations during the 2nd flight are illustrated in Fig. 1.

During both ESSenCe flights one of the internal blackbodies and quasi deep space spectra are used for a two point radiometric calibration. The latter exhibit remnants of atmospheric radiation, namely signatures of  $\text{CO}_2$  below  $750\text{ cm}^{-1}$ ,  $\text{O}_3$  around  $1050\text{ cm}^{-1}$ , and  $\text{CH}_4$  and  $\text{N}_2\text{O}$  around  $1300\text{ cm}^{-1}$ , which affect the radiometric accuracy. These spectral windows are not used in the retrieval.

A typical GLORIA raw data image is illustrated in Fig. 2. It demonstrates the ability of the instrument to resolve spatial structures in the atmosphere as low as 150 m in horizontal and vertical direction. Constantly high signals in the lower two thirds of the image indicate gray body emissions from clouds or aerosol layers at these altitudes. In contrast to cloud sensing, the observation and quantification of trace gas filaments or temperature fluctuations does not require a horizontal resolution as provided by

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individual GLORIA pixel. Therefore only the mean value for each detector row is used in the subsequent processing. No averaging in the vertical direction of the image is applied. Broken pixel (deviation of more than two times the standard deviation from the line average) or pixel pointing to clouds are removed. Entire detector rows are ignored if more than 75 % of the pixels are marked as broken or cloudy. Weighting of individual pixels is based on the noise analysis of blackbody spectra. Typical GLORIA dynamics mode spectra for the binned data are illustrated in Fig. 3. The ro-vibrational emission lines of CO<sub>2</sub> at 792 cm<sup>-1</sup>, CFC-11 at 850 cm<sup>-1</sup>, and HNO<sub>3</sub> at 890 cm<sup>-1</sup> are clearly visible in the data.

Figure 4 illustrates the horizontal scan pattern of the GLORIA dynamics mode measurements during ESSenCe. As stated above, the line of sight is panned from 48° to 118° with respect to the airplane nose. GLORIA dynamics mode data were recorded from 14:58 to 15:10 UTC yielding to 121 individual images at about 2200 spectral points.

### 4.3 GLORIA Level 2 data analysis

#### 4.3.1 Retrieval setup

The retrieval setup for GLORIA depends on the instrument performance and atmospheric conditions. It consists of the choice of spectral windows utilized for the radiative transfer forward model, the a priori and regularization parameters utilized in the constrained global fit method, and the error covariance matrix of the measurements.

The selection of appropriate spectral windows for the GLORIA dynamics mode retrieval is non-trivial, because the spectral resolution of this mode does not allow to separate individual trace gas emission lines. In addition, the radiometric calibration procedure utilizing “quasi” deep space spectra (in contrast to “real” deep space spectra as observed from a satellite) causes a non-uniform error pattern in the wavelength domain. Therefore the selection of spectral windows is performed by an automatic algorithm maximizing the information gain obtained by the measurements. The figure

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of merit for information gain is determined by the instrumental uncertainties and the a priori knowledge (Kullback and Leibler, 1951). This selection is performed by a newly developed genetic algorithm optimizing the spectral window size and position simultaneously (Blank, 2013). In contrast to previous work (von Clarmann and Echle, 1998; Dudhia et al., 2002), this method does not require pre-selected broader spectral ranges and is able to optimize several spectral windows simultaneously. The latter is most relevant here due to the limited spectral resolution of the GLORIA dynamics mode measurements.

For ESSenCe, the main retrieval targets are temperature,  $O_3$ ,  $HNO_3$  and CFC-12. The spectral windows as selected by the genetic algorithm to retrieve these quantities are illustrated in Fig. 5. There are twice as many spectral windows than main retrieval targets to consider cloud/aerosol effects as well as other contaminant gases. A priori data for the constrained global fit are taken from the MIPAS reference atmospheres (Remedios et al., 2007). Some species (such as  $CO_2$ ) were updated to 2011 conditions. Pressure and temperature data were taken from the ECMWF ERA-Interim data set (Dee et al., 2011). The penalty of zero order regularization (Eq. 2) was suppressed by a factor of three in comparison to an optimal estimation method assuming the standard deviations of the MIPAS climatology. For temperature and pressure, an uncertainty of 3 K and 1 % is assumed, respectively. Weighting factors for the first order regularization in the vertical direction are inversely proportional to the standard deviations of the MIPAS climatology multiplied by a characteristic length scale of 0.25 km for temperature and CFC-12 and 1 km for  $O_3$  and  $HNO_3$ , respectively. The horizontal weighting factor is 100 km for all species.

The characterization of the GLORIA measurement error covariance matrix is work in progress. The main source of uncertainty is instrument noise, approximated by a diagonal matrix assuming an uncertainty of  $25 \text{ nW}/(\text{cm}^{-2} \text{ sr cm}^{-1})$  and 2 %, respectively. The transportation and mounting of the GLORIA instrument into the aircraft, as well as thermal stress during a research flight may cause a misalignment between the GLORIA attitude control system and the interferometer's line of sight. Therefore, it is mandatory

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to verify the vertical pointing of the instrument and apply some corrections, if necessary. This is done by simulating the CO<sub>2</sub> Q-branch emissions at 792.5 cm<sup>-1</sup> adopting ECMWF temperature data for the entire flight. Assuming that the real atmosphere during a research flight has a mean temperature similar to ECMWF data, a single elevation angle correction is retrieved. In the case of ESSenCe Flight 2 this correction is 0.1955°. Since the dominant uncertainties of the resulting line of sight elevation angles are not expected to change during the flight, small and medium scale structures in atmospheric temperature data can still be resolved by the measurement. The effect of the remaining pointing uncertainty of 0.023° is considered in the error budget of the retrieval targets (see Sect. 4.3.2).

### 4.3.2 Results, uncertainties and comparison to other measurements

Typical limb radiance altitude profiles for a few spectral windows are illustrated in Fig. 6. The intensities are generally increasing towards lower altitudes, as expected. The dynamical range given by maximum and minimum observed radiances in the considered spectral range varies between 50 % and 500 % for tangent altitudes between 9 km and 17 km, respectively.

Corresponding retrieval results for the main targets of the GLORIA dynamics mode are illustrated in Fig. 7. The most prominent feature in the retrieved quantities is a pronounced enhancement of O<sub>3</sub> and HNO<sub>3</sub> around an altitude of 14 km, which may point to a filamentary structure in the stratosphere. This enhancement extends several hundred kilometers horizontally, as indicated by the two-dimensional cross-sections measured by GLORIA (Fig. 8). The uncertainties of HNO<sub>3</sub> and O<sub>3</sub> (Fig. 9) are in the order of 0.25 ppbv and 0.1 ppbv, respectively and are clearly dominated by detector noise. This means, that the trace gas enhancement at 14 km is significant, whereas the small scale structures below this altitude are within the noise error. The vertical resolution (Fig. 10) is about 500 m for HNO<sub>3</sub> and O<sub>3</sub>, and somewhat worse for temperature (750 m) and CFC-12 (600–1400 m). The vertical gradient of CFC-12 dynamics mode data as well as its (comparatively low) vertical resolution does not allow to resolve the filament at

14 km, whereas the high vertical resolution of O<sub>3</sub> and HNO<sub>3</sub> should give a realistic picture of this structure.

To validate these GLORIA measurements the retrieved data are compared to several other instruments. These are the two limb emission sounders mounted on the same airplane as GLORIA, namely the MIPAS-STR (Woiwode et al., 2012) and MARSCHALS (Millimetre-wave Airborne Receivers for Spectroscopic Characterization in Atmospheric Limb Sounding, Castelli et al., 2013) instruments. Furthermore, two in situ datasets obtained during the ascent and descent of the second ESSenCe flight are used for the comparison, namely measurements of the Whole Atmosphere Sampler (WAS; Laube et al., 2013) and the High Altitude Gas AnalyzeR (HAGAR; Volk et al., 2000; Werner et al., 2010) datasets. In addition, GLORIA measurements are compared to the satellite borne limb emission measurements of MIPAS (Fischer et al., 2007) on ESA's Environmental Satellite (Envisat) and the Earth Observing System (EOS) Microwave Limb Sounder (EOS-MLS, data version v03.33; Froidevaux et al., 2008) on NASA's EOS AURA satellite. The MIPAS data were derived with the retrieval processor of the Karlsruhe Institute of Technology; data version is V5R\_O3\_221 and V5R\_HNO3\_221 for ozone and HNO<sub>3</sub> (von Clarmann et al., 2009) and V5R\_CFC-11\_221 for CFC-11 (Kellmann et al., 2012). Finally, the GLORIA data are compared to simulations of the Chemical Lagrangian Model of the Stratosphere (CLaMS; Konopka et al., 2007; Grooß et al., 2014, and references therein). CLaMS runs with full stratospheric chemistry (including heterogeneous chemistry and sedimentation) and is driven by ERA-Interim reanalysis data (for the detailed model setup, see Vogel et al., 2014). The simulations cover an altitude range from the surface up to the 900 K potential temperature with a horizontal resolution of approximately 100 km and a maximum vertical resolution of about 400 m around the tropopause. The simulations were initialized on 1 November 2011 (for ESSenCe) and on 1 May 2012 (for the TACTS/ESMVal campaign), respectively, using satellite measurements and tracer tracer correlations following Grooß et al. (2014).

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The mismatch in distance between the comparison data and GLORIA's tangent points is typically a few hundred kilometers (Fig. 4). Since this is larger than the spatial extent of the GLORIA dynamics mode measurements and since the atmospheric scene is relatively homogeneous in the horizontal direction (Fig. 8), the entire dynamics mode sequence from 14:58 to 15:10 UTC was averaged for comparison with the other data sets.

The satellite instruments' vertical profiles exhibit a much coarser altitude resolution than the GLORIA data. This effect is taken into account by convolving the GLORIA data  $x_{\text{gloria}}$  with the averaging kernel of these instruments ( $A_{\text{mipas}}$  and  $A_{\text{mls}}$ ) considering the respective a priori data ( $x_{\text{mipas}}^{\text{apr}}$  and  $x_{\text{mls}}^{\text{apr}}$ ):

$$x_{\text{gloria}}^{\text{conv}} = x_{\text{mipas/mls}}^{\text{apr}} + A_{\text{mipas/mls}} (x_{\text{gloria}} - x_{\text{mipas/mls}}^{\text{apr}})$$

Figures 11–13 show a comparison of GLORIA retrieval results with the different data sets described above. The left panel of these figures illustrates the data “as it is”, and the right panel shows differences between GLORIA and the other data sets, where the effect of the averaging kernels was taken into account for the calculation of the differences. It is obvious, that the effect of the averaging kernel is rather large for the comparison with the satellite instruments, because the vertical resolution of the satellite data sets is several kilometers in the lower stratosphere.

GLORIA  $\text{O}_3$  data show an excellent agreement with MARSCHALS data, which were recorded a few minutes (14:52–14:56 UTC) before GLORIA dynamics mode measurements. MARSCHALS  $\text{O}_3$  abundance was retrieved from the Band B data (297–305 GHz). Differences between the two datasets are a few percent and smaller than the combined error bars. Most notable is the observation of the  $\text{O}_3$  enhancement at 14 km in both datasets. Technical improvements of the MARSCHALS instrument and a new retrieval setup (Gerber, 2014) allowed the derivation of such a fine structure from MARSCHALS data for the first time.  $\text{O}_3$  observations of MIPAS-STR exhibit similar absolute values, but the local enhancement at 14 km is not resolved due to the



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coarser altitude resolution of this dataset. Differences to the satellite borne datasets of EOS-MLS are typically less than 15 % and somewhat larger for MIPAS-Envisat, but both datasets cannot resolve the ozone enhancement at 14 km due to their broad averaging kernels. CLaMS model simulations do not show this localized enhancement either, which is consistent with the non-existence of this structure in the ERA-interim data at a resolution of  $1^\circ \times 1^\circ$  as utilized for this model run. However, modified potential vorticity data (Müller and Günther, 2003) based upon ECMWF operational analyzes at a much higher spatial resolution ( $0.1^\circ \times 0.1^\circ$ , not shown here) reveals increased values at an altitude of 13.5 km at the location of the GLORIA measurements. This indicates that two different air masses get close together and model simulations at higher spatial resolution are needed to resolve such structures.

Mixing ratios of  $\text{HNO}_3$  measured from GLORIA and MIPAS-STR show similar values at most altitudes as well – differences are typically less than 15 %, except for the 14 km region, where GLORIA data is enhanced by more than 1 ppbv. MIPAS-Envisat  $\text{HNO}_3$  values (spatial separation 350–550 km) are 20–50 % larger than the other two datasets, which may be due to the differences in the tangent point location and viewing geometry.

For CFC-12, GLORIA data shows a small high bias of about 5 % in the entire altitude region in comparison to collocated MIPAS-STR data. This bias is also observed in comparison to the CFC-12 data of the WAS and HAGAR in-situ instruments measured during the ascent and descent of the aircraft. CLaMS results are in good agreement with the in-situ measurements. Although this difference between GLORIA and the in-situ or model data is within the error margins of the GLORIA instrument, it may point so some instrumental issues to be solved in future missions.



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the antarctic polar vortex (Fig. 14). This region is characterized by strong gradients in potential vorticity, ozone, and  $\text{HNO}_3$  mixing ratios. At the edge of the polar vortex, HALO performed a hexagonal flight track, such that the GLORIA lines of sight are directed to the center of the hexagon. At this location ( $45^\circ \text{S}$ ,  $20^\circ \text{E}$ ) the horizontal winds are sheared and tilted as a consequence of a distortion of the jet stream at lower altitudes (Fig. 15). Wind speeds are in the order of  $45 \text{ m s}^{-1}$ .

To retrieve the atmospheric composition within the air volume contained by the hexagon, an appropriate atmospheric grid was defined. The size of the grid cells is based on the vertical and horizontal field of view of the GLORIA instrument, namely 125 m in the vertical and 7.5 km in the horizontal direction, respectively, resulting in about  $2 \times 10^6$  individual cells. The size of the grid cells is increasing towards the boundaries of the air-volume.

The horizontal and vertical regularization applied to the retrieval targets is similar to ESSenCe, as is the selection of spectral windows. The reduced spectral resolution and uncertainties in the instrument calibration led to the elimination of a few spectral windows describing the aerosol background and some “secondary” gases such as PAN or  $\text{ClONO}_2$ , which cannot be retrieved with high confidence for this atmospheric situation. The most relevant difference to ESSenCe is the omission of the spectral window at  $845 \text{ cm}^{-1}$  showing strong CFC-11 emissions, which turned out to be contaminated by instrument emissions from the optical window of GLORIA (Kleinert et al., 2014). Cloud indices (Spang et al., 2004) observed during the hexagonal flight were 4–6 between 9 km and 14 km, indicating an almost cloud free atmosphere.

Retrieved temperature and constituent fields exhibit a pronounced intrusion of stratospheric air into the troposphere at about  $45^\circ \text{S}$ ,  $20^\circ \text{E}$  (Fig. 16). This intrusion is observed between an altitude of 12.5 km and 14.5 km and extends about 500 m vertically. The horizontal extent is about 200 km and non-uniform in the longitudinal direction. Its stratospheric nature is indicated by enhanced  $\text{HNO}_3$  and  $\text{O}_3$  – at least by a factor of two compared to the air parcels just below or above the filament. This enhancement is significantly larger than the uncertainty of these species. Temperature data does not

show this filamentary structure and varies by less than 2 K in the vicinity of this structure. CFC-12 does not show a pronounced structure as well due to its weak vertical gradient at these altitudes.

Using three-dimensional tomography the horizontal resolution is about 20 km in the core of the measured atmospheric region (Fig. 17). In this region, each atmospheric grid cell is measured from at least 4 different directions. The vertical resolution of the atmospheric fields is 200–300 m and somewhat larger than for the data obtained from the linear flight tracks during ESSenCe due to the increased instrument performance of GLORIA during the TACTS/ESMVal campaign. Since the calculation of retrieval diagnostics such as averaging kernel and spatial resolution is computationally very demanding (it breaks the sparsity of the computational problem; see, e.g., Ungermann et al., 2010), this information has not been calculated for all atmospheric grids cells considered in the retrieval.

A preliminary comparison to the CLaMS simulation for this flight (Fig. 14) indicates that the model resolution of 100 km (horizontally) is in turn not sufficient to resolve this intrusion of air as observed by the tomographic sounding.

## 6 Conclusions

The first two deployments of GLORIA demonstrated the readiness of this instrument to measure upper tropospheric/lower stratospheric composition at high spatial resolution. The retrieval of temperature, O<sub>3</sub>, HNO<sub>3</sub> and CFC-12 flagged no major problems, except for a small bias in CFC-12 data. Increased effort in the radiometric calibration procedure will reduce this bias in future data versions. During ESSenCe GLORIA already demonstrated its ability to measure intrusions of stratospheric air on vertical scales as low as a few hundred meters. However, the outstanding capability of the GLORIA instrument can only come into full effect when the airplane carrier flies a closed path around the air volume to be observed as it was performed several times during the TACTS/ESMVal

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We thank all members of the GLORIA instrument team for their large efforts in the development of the GLORIA instrument and the conduction of the ESSenCe and TACTS/ESMVal campaigns.

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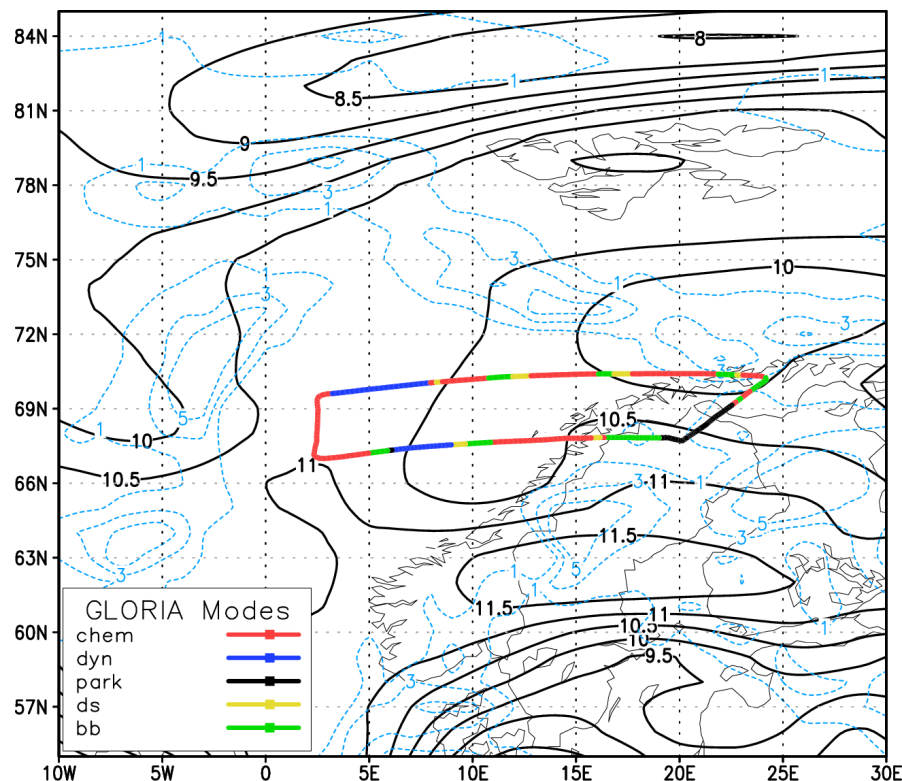
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**Fig. 1.** Potential vorticity (in PVU, black solid lines) on 370 K isentropic surface (at approx. flight level) and cirrus clouds in lower levels (shown by cloud ice water content [ $10^5 \text{ kg kg}^{-1}$ ] on 280 K, blue lines) and the track of the Geophysica aircraft for ESSenCe Flight 2 of 16 December 2011. “dyn” and “chem” mark dynamics and chemistry mode measurements of GLORIA, respectively. “park” indicates that GLORIA is in parking position, “ds” and “bb” indicate deep space and blackbody calibration measurements, respectively.

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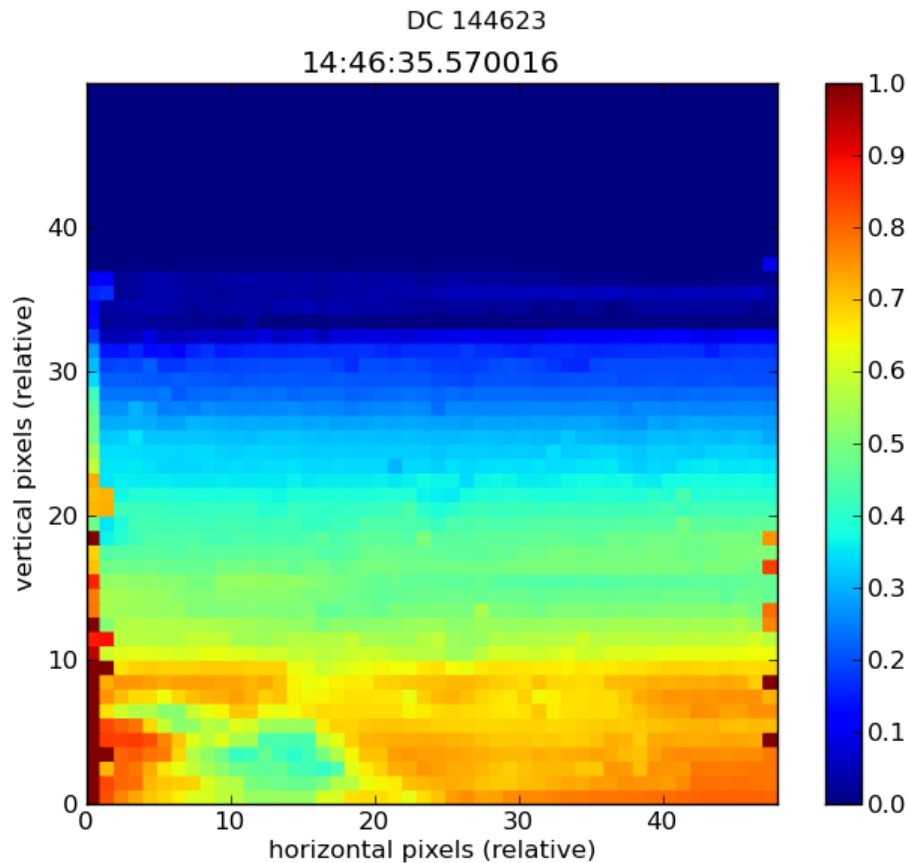
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**Fig. 2.** Direct current signal (arb. units) for one GLORIA image, before the Fourier transformation is applied. Exposure time is 40 ms and the horizontal extent is about 7.5 km at 10 km tangent altitude.

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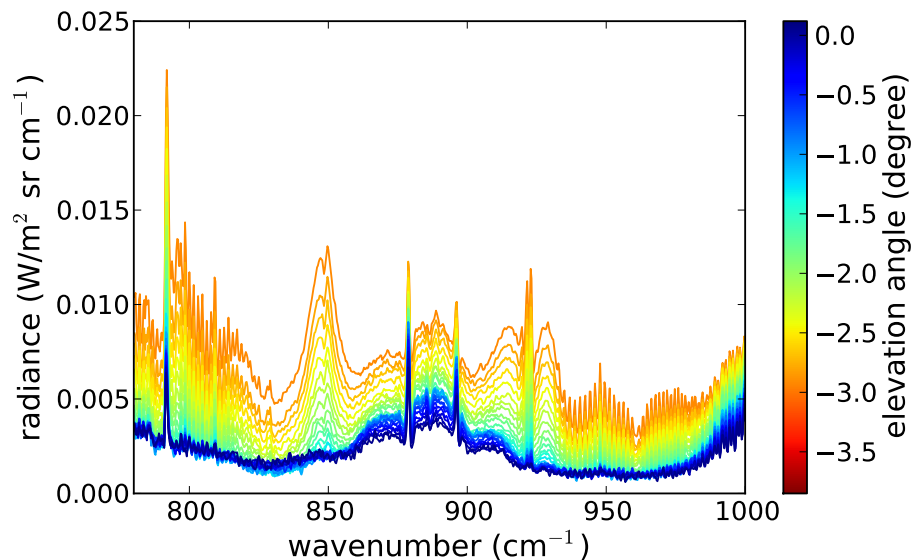
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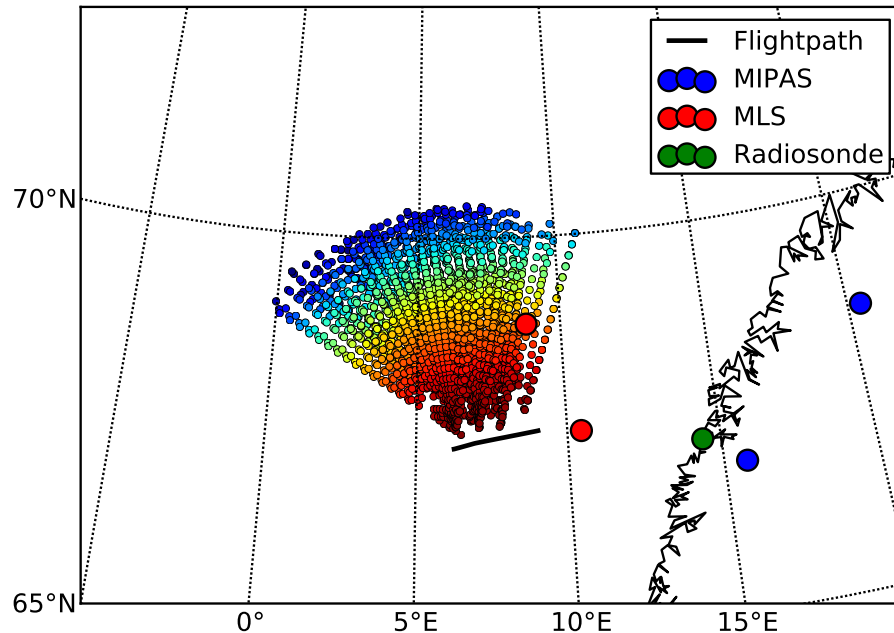
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**Fig. 3.** Spectra obtained by averaging over a detector row for a single GLORIA dynamics mode measurement (Image 72, measured at 15:04:49 UTC at 67° N, 7° E) after pixel weighting and filtering. Corresponding tangent altitudes are 9 km (red) to 15 km (dark blue).

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**Fig. 4.** GLORIA tangent points of dynamics mode measurements during ESSenCe Flight 2. Tangent points closest (furthest) from the flight path (black line) are at highest (lowest) altitudes. Corresponding tangent altitudes are 4 km (red) to 15 km (blue). Thick colored dots mark the geolocation of collocated MIPAS-Envisat and EOS-MLS tangent points and radiosonde positions. MIPAS-Envisat measurements were made on 16 December 2011 at 15:21 and 16:09 UTC, and EOS-MLS measurements at 11:40 and 11:42, respectively.

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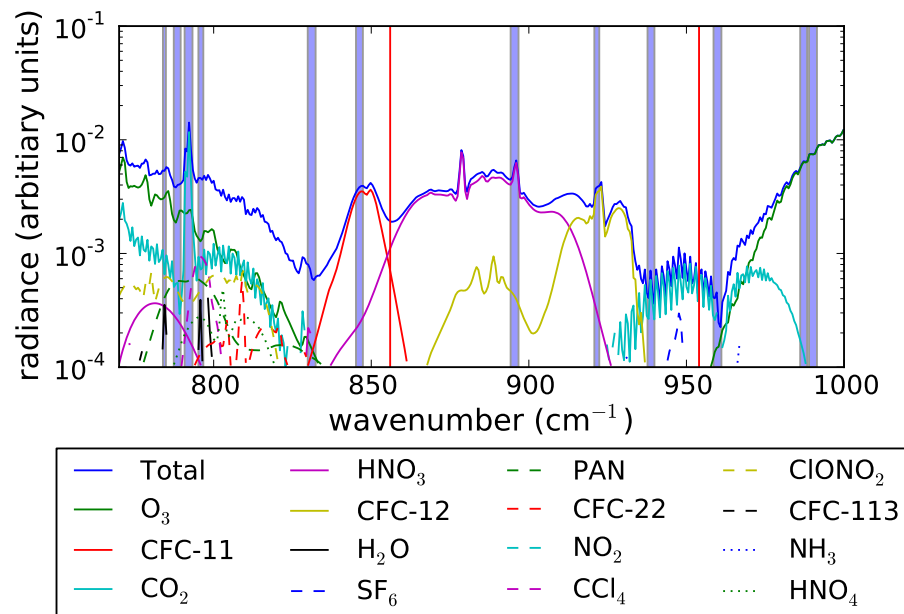
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**Fig. 5.** Simulated GLORIA dynamics mode spectra for 16 km tangent altitude applying the MIPAS reference atmosphere for mid latitudes. The spectral windows utilized in the retrieval are marked by blue vertical boxes.

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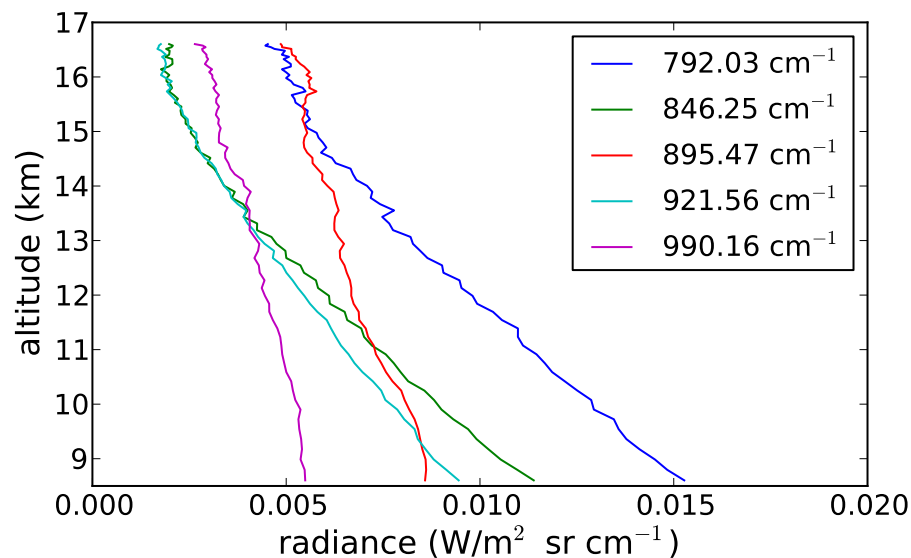
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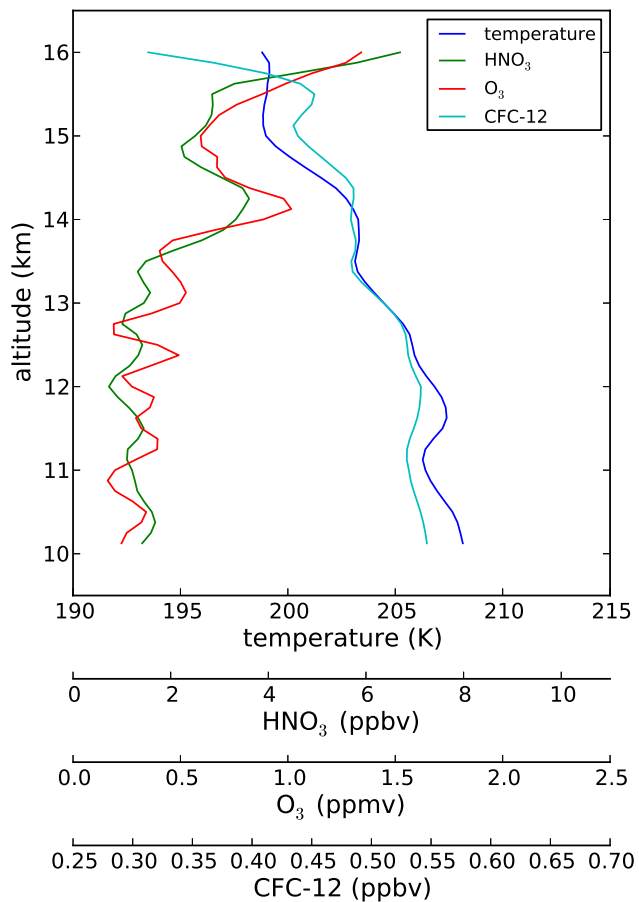
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**Fig. 6.** GLORIA integrated radiances for some spectral windows as selected by the genetic algorithm. The data is for Image 72 at 67° N and 7° E on 16 December, 15:05 UTC.



**Fig. 7.** Retrieved temperature, HNO<sub>3</sub>, O<sub>3</sub> and CFC-12 for GLORIA Image 72.

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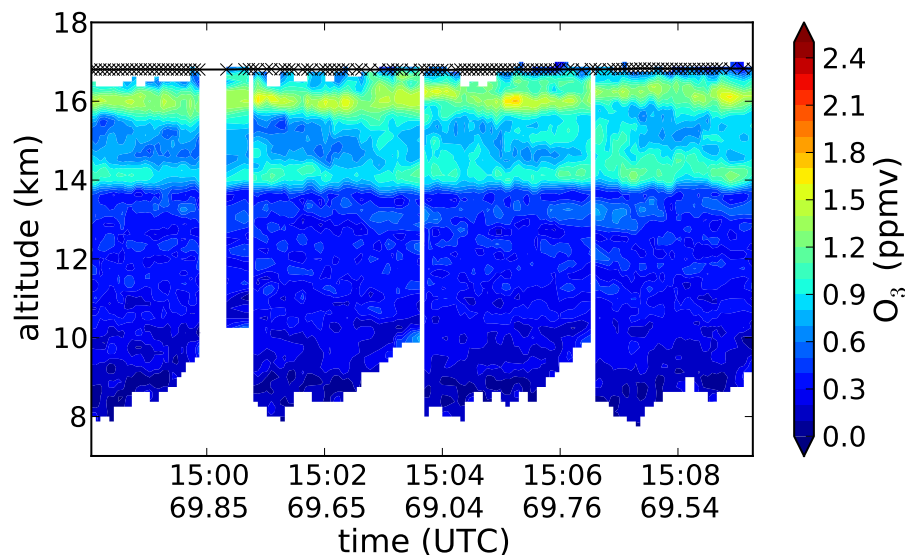
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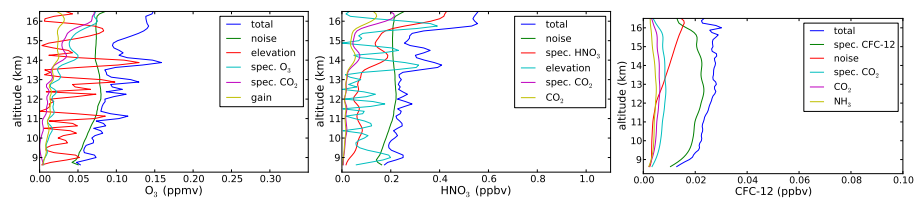
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**Fig. 8.** Retrieved ozone volume mixing ratio from GLORIA dynamics mode data for 16 December 2011. Universal time and geographic latitude (in degrees) label the  $x$  axis. The figure shows four horizontal sweeps starting at  $48^\circ$  and ending at  $118^\circ$  with respect to the airplane nose.

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**Fig. 9.** Error budget for  $\text{HNO}_3$ ,  $\text{O}_3$  and CFC-12 for GLORIA Image 72.

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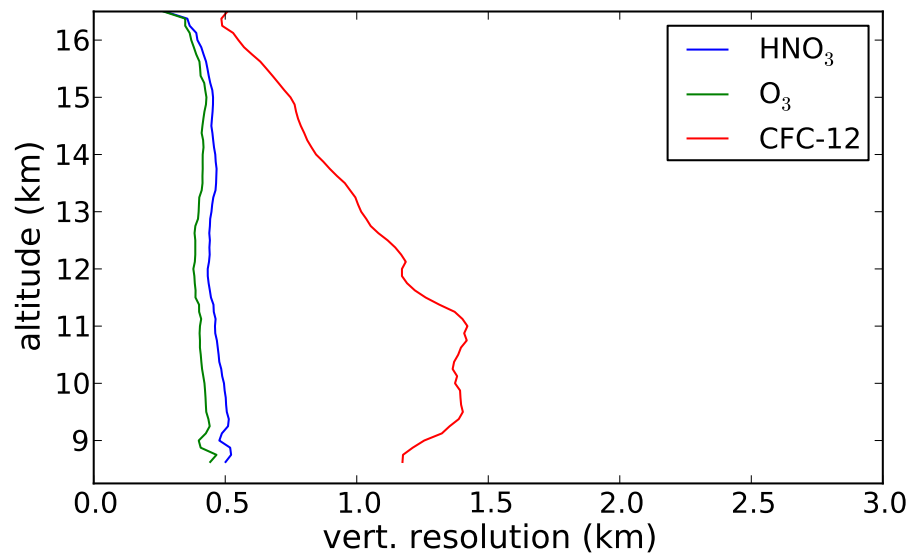
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**Fig. 10.** Vertical resolution for HNO<sub>3</sub>, O<sub>3</sub> and CFC-12, for GLORIA Image 72.

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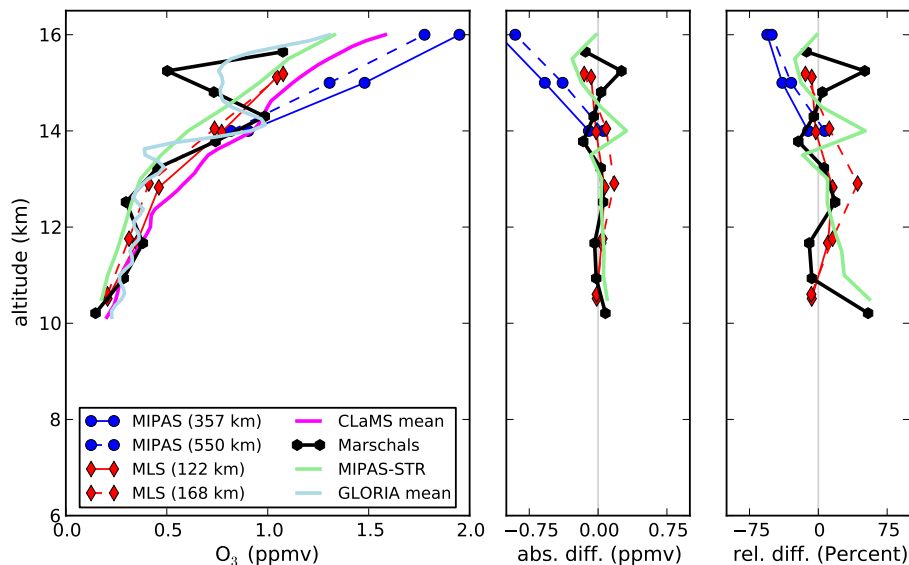
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**Fig. 11.** Comparison of  $O_3$  abundance between GLORIA dynamics mode measurements and various collocated datasets.

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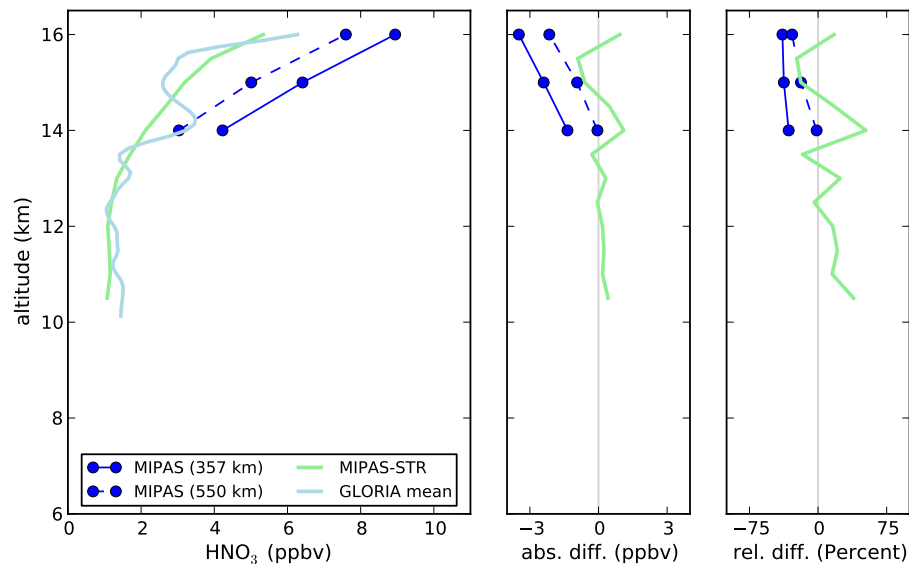
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**Fig. 12.** Comparison of HNO<sub>3</sub> abundance between GLORIA dynamics mode measurements and collocated datasets.

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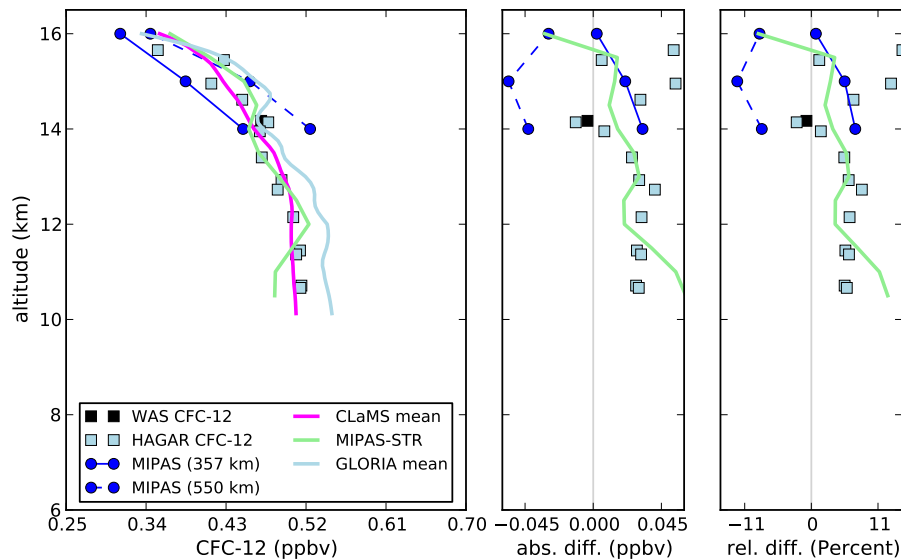
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**Fig. 13.** Comparison of CFC-12 abundance between GLORIA dynamics mode measurements and collocated datasets.

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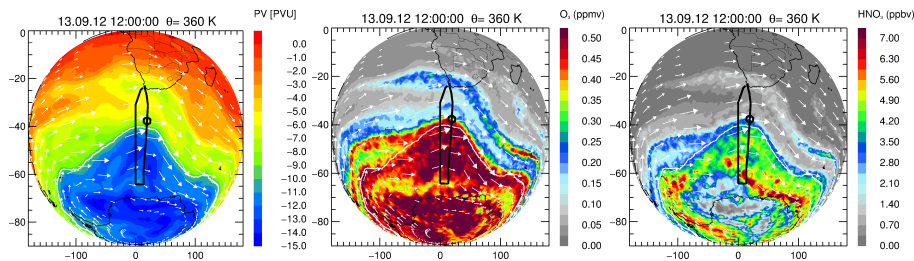
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**Fig. 14.** ECMWF potential vorticity (left) at a potential temperature of 360 K (about 12 km altitude) for the TACTS/ESMVal flight on 13 September 2012. The flight track is indicated by a black line. At 45° S, 20° E a tomographic flight pattern (approx. altitude about 12.5 km) was performed. The middle and right plots show  $O_3$  and  $HNO_3$  volume mixing ratios at 360 K as simulated by the CLaMS model, respectively. The vortex edge is calculated as defined by Nash et al. (1996) and is marked by a white contour line. The horizontal winds are indicated by white arrows.

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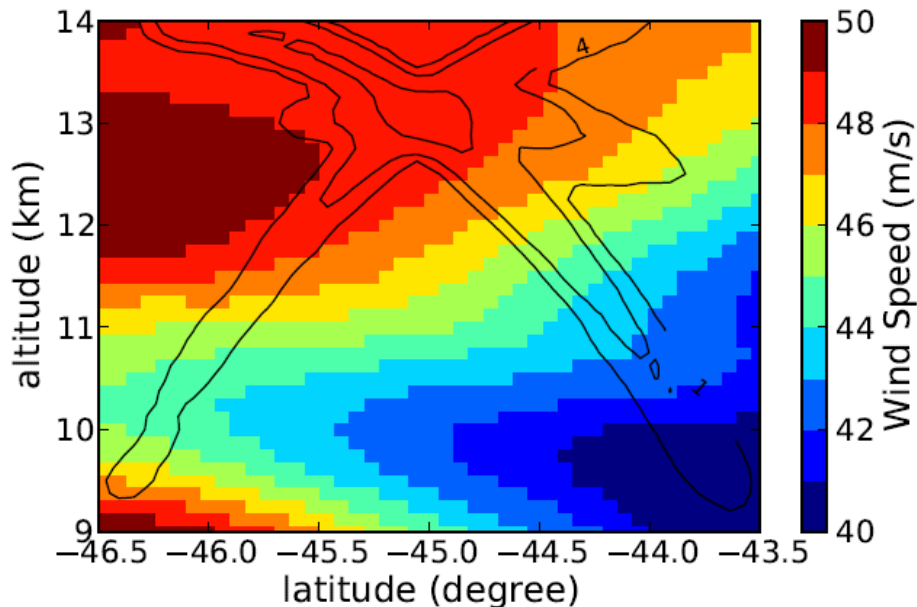
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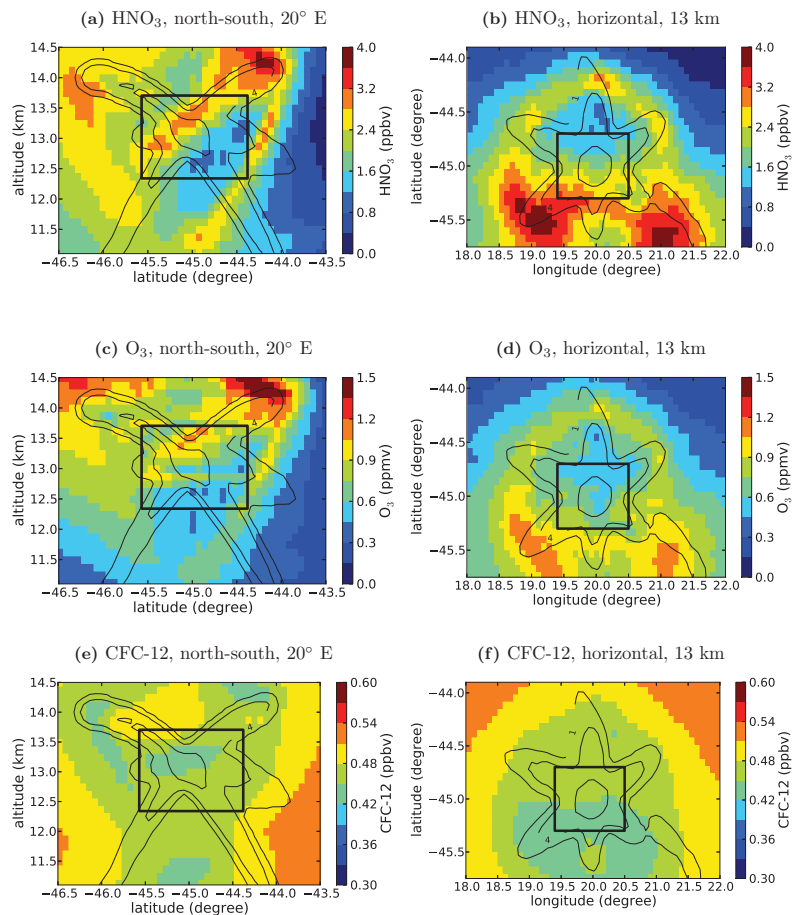


**Fig. 15.** Horizontal wind speeds at 20° E for 13 September 2012, 14:00 UTC. Contour lines (at 2, 4, 16) indicate the number of tangent points per grid cell; highest numbers are at the center of the plot.

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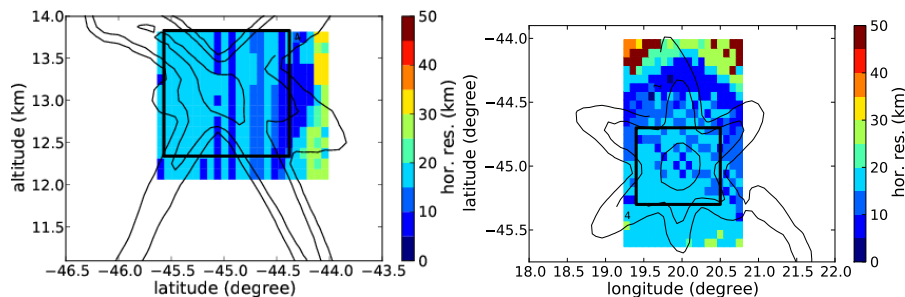
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**Fig. 16.** Retrieved  $\text{HNO}_3$ ,  $\text{O}_3$  and CFC-12 mixing ratios at  $20^\circ \text{E}$  (left) and 13 km altitude (right) for 13:30–14:30 UTC, 13 September 2012. Contour lines (at 2, 4, 16) indicate the number of tangent points per grid cell; highest numbers are at the center of the plot.

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**Fig. 17.**  $\text{HNO}_3$  horizontal resolution at  $20^\circ \text{E}$  (left) and  $13 \text{ km}$  altitude (right). Contour lines (at 2, 4, 16) indicate the number of tangent points per grid cell; highest numbers are at the center of the plot.

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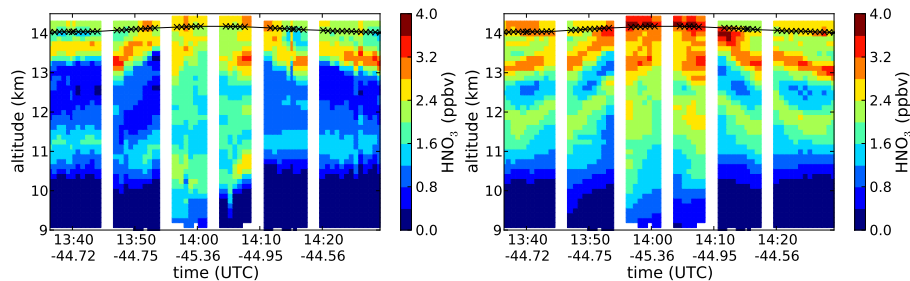
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**Fig. 18.** Comparison between a one-dimensional (left) and three-dimensional (right) reconstruction of  $\text{HNO}_3$  during the GLORIA hexagonal flight. The segments of the hexagon are separated by white stripes. Measurement time (UTC) and latitude (deg) for the beginning of each segment are given on the x axis.

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