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# The white-light humidified optical particle spectrometer (WHOPS) – a novel airborne system to characterize aerosol hygroscopicity

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Aerosol particles experience hygroscopic growth at enhanced relative humidity (RH) which leads to changes in their optical properties. We developed the white-light humidified optical particle spectrometer (WHOPS), a new instrument to investigate the particles' hygroscopic growth. Here we present a detailed technical description and characterization of the WHOPS in laboratory and field experiments. The WHOPS consists of a differential mobility analyzer, a humidifier/bypass and a WELAS (white-light aerosol spectrometer) connected in series to provide fast measurements of particle hygroscopicity at sub-saturated RH and optical properties on airborne platforms. The WELAS employs a white-light source to minimize ambiguities in the optical particle sizing. In contrast to other hygroscopicity instruments, the WHOPS retrieves information of relatively large particles (i.e. diameter  $D > 280 \, \text{nm}$ ), therefore investigating the more optically relevant size ranges.

The effective index of refraction of the dry particles is retrieved from the optical diameter measured for size-selected aerosol samples with a well-defined dry mobility diameter. The data analysis approach for the optical sizing and retrieval of the index of refraction was extensively tested in laboratory experiments with polystyrene latex size standards and ammonium sulfate particles of different diameters. The hygroscopic growth factor (GF) distribution and aerosol mixing state is inferred from the optical size distribution measured for the size-selected and humidified aerosol sample. Laboratory experiments with pure ammonium sulfate particles revealed good agreement with Köhler theory (mean bias of  $\sim 3\%$  and maximal deviation of 9% for GFs at RH = 95%).

First airborne measurements in the Netherlands observed GFs (mean value of the GF distribution) at RH = 95 % between 1.74 and 2.67 with a median of 1.94 for particles with a dry diameter of 500 nm. This corresponds to hygroscopicity parameters ( $\kappa$ ) between 0.21 and 0.93 with a median of 0.33. The GF distributions indicate externally mixed particles covering the whole range of GFs between  $\sim$  1.0–3.0. On average  $\sim$  74 % of the particles were "more hygroscopic" with GFs > 1.5,  $\sim$  15 % were non- or

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slightly hygroscopic with GF < 1.1 and the remaining ~ 11 % were "less hygroscopic" with 1.1 < GF < 1.5. The more hygroscopic mode sometimes peaked at GF > 2, indicating influence of sea salt particles, consistent with previous ground-based particle hygroscopicity measurements in this area. The mean dry effective index of refraction for 500 nm particles was found to be rather constant with a value of  $1.42 \pm 0.04$ .

#### Introduction

Atmospheric aerosols affect the Earth's radiative budget directly and indirectly. The scattering or absorption of incident sunlight by aerosol particles is called the direct aerosol-radiation effect. The aerosol optical properties depend on various factors like chemical composition, size and the ability of the particles to take up water. If particles are hygroscopic, their size and index of refraction and hence their optical properties change with changing relative humidity (e.g. Zieger et al., 2013). The knowledge of their hygroscopic properties and mixing state is central for climate forcing calculations and for the validation of remote sensing measurements with in-situ ones. This is of particular importance if the chemical composition and thus the hygroscopicity changes with altitude e.g. due to temperature-dependent ammonium nitrate partitioning (Morgan et al., 2010). Therefore, vertical profiles of aerosol properties are essential to enable more accurate model calculations and improve remote sensing retrieval validation.

The hygroscopic growth of the particles can be described by the growth factor GF which is defined as the particle diameter  $D_{\text{wet}}$  at a certain relative humidity (RH) divided by the dry diameter  $D_{dry}$ :

$$GF(RH) = \frac{D_{wet}(RH)}{D_{dry}}$$
 (1)

For a given chemical composition, the GF is dependent on RH and on the particle size (Kelvin effect). Therefore, to compare to other hygroscopicity measurements and

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$$\kappa = \frac{(\mathsf{GF}(\mathsf{RH})^3 - 1) \times (1 - a_{\mathsf{w}})}{a_{\mathsf{w}}} \tag{2}$$

The water activity can be inferred from the RH and equilibrium droplet diameter ( $D_{wet}$ ):

$$a_{\rm W} = \frac{\rm RH}{\exp\left(\frac{4\sigma_{\rm s}M_{\rm w}}{RT\rho_{\rm w}D_{\rm wet}}\right)} \tag{3}$$

Where  $\sigma_{\rm s}$  is the surface tension of the solution, R the ideal gas constant, T the absolute temperature,  $\rho_{\rm w}$  the density of water and  $M_{\rm w}$  the molecular weight of water. The  $\kappa$  parameter has the advantage that it is for the most part determined by the chemical composition of a particle, while it is independent of particle size and only weakly dependent on RH (at least over limited RH ranges). It allows to directly compare measurements made at different dry sizes and RH.

The need to quantify the hygroscopicity of aerosol particles led to the development of several instruments including the hygroscopicity tandem differential mobility analyzer (HTDMA; e.g. Liu et al., 1978; Swietlicki et al., 2008). It measures dry and humidified mobility diameters, whereby the RH of the humidified DMA (differential mobility analyzer) can be varied, making it possible to record humidograms. Even if this instrument provides high precision measurements of GFs and hygroscopic mixing state of the aerosol, the measurement duration for one measurement amounts to several minutes which makes it rather unsuited for airborne studies. Besides the maximal dry mobility diameter that can be detected is approximately 250 nm. The diameter size range of the standard HTDMA is therefore limited to small particles, and might miss hygroscopic features of larger particles (e.g. sea salt, Zieger et al., 2011) which are also important for aerosol light scattering.

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In contrast, the differential aerosol sizing and hygroscopicity spectrometer probe (DASH-SP; Sorooshian et al., 2008) was developed specifically to provide a fast measurement of the GF. It can measure at different RHs simultaneously and the maximum dry diameter it can detect accurately is 450 nm. The measurement principle is based on a combination of a DMA and an OPSS (optical particle size spectrometer) which has been utilized also in previous techniques like for the LACIS (Leipzig Aerosol Cloud Interaction Simulator; Stratmann et al., 2004). In the DASH-SP a specific monodisperse aerosol is selected in the DMA, then exposed to different RHs and finally measured in the OPSS. However, the OPSS uses monochromatic light which can lead to ambiguities for the diameter retrieval.

Another instrument that can be compared to the ones above is the humidified nephelometer (Wet-Neph; e.g. Fierz-Schmidhauser et al., 2010) which measures the aerosol light scattering coefficient at dry and elevated RH and allows to retrieve the light scattering enhancement factor. However, in the usual setup, no size specific or mixing state information can be retrieved.

In this paper, we describe a novel instrument that has specifically been developed for airborne measurements of aerosol particles in the optically relevant size range. The white-light humidified optical particle spectrometer (WHOPS) uses a DMA – OPSS setup to allow the detection of larger aerosol sizes. A WELAS 2300 sensor (white-light aerosol spectrometer; Palas GmbH, Karlsruhe) was chosen for the optical measurements since it uses a white-light source to illuminate the particles which in turn allows for an unambiguous relation between the light intensity scattered by a particle and its size. Another advantage is the simultaneous measurement of all sizes in the detectable range of approximately 280 nm to  $10\,\mu m$  which therefore allows a faster measurement.

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## Setup of the PSI-Zeppelin rack

The instrument was specifically developed for airborne measurements onboard an airship (Zeppelin NT). The compact rack contains the WHOPS, another WELAS and an 5 aethalometer, see Fig. 1.

To collect the aerosol an inlet was built for isokinetic sampling at an average flight velocity of 50 km h<sup>-1</sup>. Then, the flow is split in three ways:

5 liters per minute (Lmin<sup>-1</sup>) are directed to a WELAS 2300 sensor to measure the polydisperse size distribution of the ambient aerosol (labelled as WELAS1 in Fig. 1). This sensor uses an intense source of white light (Osram XBO-75 xenon short arc lamp) to illuminate a certain sensing volume (approximately 2.24 cm<sup>3</sup> × 2.24 cm<sup>3</sup> × 1.4 cm<sup>3</sup>), which has a 3-D T-shape to minimize "border-zone" errors, i.e. to ensure a uniform light-intensity across the whole sensing volume (see Heim et al., 2008 for more details). Optical lenses collect the scattered light between 78° and 102° with respect to the incident beam and direct it to a photomultiplier tube (PMT) which detects the partial scattering cross section of the particles for this solid angle. The so called "PROMO 3000 box" contains the PMT and lamp which are connected to the sensor via optical fibers to minimize heat input from the lamp and thus a temperature increase. Since each signal pulse height recorded by the PMT can be related to a corresponding optical particle size (see Sect. 2.2.1), the particle number size distribution can be retrieved. The diameters that can be detected range from approximately 280 nm to 10 µm in PSL (polystyrene latex) equivalent particle diameters, though with decreasing detection efficiency for the smallest sizes. The lower limit is given by the minimum detection threshold of the PMT while the upper limit is reached above saturation of the detection electronics.

The WHOPS uses a combined setup of a DMA and another WELAS 2300 sensor (labelled as WELAS2 in Fig. 1) connected in series. This WELAS2 was modified to operate at a reduced sample flow rate of 0.5 L min<sup>-1</sup> in order to be able to operate the

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DMA with sample and sheath air flow rates of 0.5 L min<sup>-1</sup> and 5.0 L min<sup>-1</sup>, respectively, without adding dilution air between DMA and WELAS2. In order to increase the detection probability of the particles, a nozzle was implemented by the manufacturer in the WELAS2 inlet, guiding the flow directly towards the measurement volume.

In the WHOPS the air passes first a silica gel dryer and reaches a DMA where particles with a certain mobility diameter are selected. In the next step these monodisperse particles can be measured either under dry or wet conditions. The WHOPS is operated in the wet mode to study the hygroscopic behavior: the dried and size-selected particles pass through a custom-built humidifier and finally enter a growth chamber with a residence time of about 10 s. The WELAS2 is then used to determine the resulting equilibrium diameters of the grown particles at elevated RH, typically around 95%. To maintain a constant RH, the humidifier and the optical sensor are situated in an insulated box with a fan to ensure a well-stirred air circulation and therefore a homogeneous temperature. The measurement of the RH is crucial for a correct interpretation of the observed GF. Figure 1 shows that 3 RH and temperature (T) measurements are included in the setup. We use HC-S probes (Rotronic Hygroclip Sensors) which monitor RH changes with a precision of ±0.8 %. However, extensive laboratory tests showed that at changing conditions at high RH (above 90 % RH) the sensors and the tubes need an equilibrium time of approximately 3-5 min. Therefore, all RH measurements under transient conditions were corrected for equilibration time effects. The accuracy of the corrected RH measurements is estimated to be ~ ±2%. The good agreement between theoretical and measured hygroscopic growth factors of ammonium sulfate results presented in Sect. 3.1.4, supports this assumption. To ensure this accuracy and precision, the sensors were regularly calibrated using saturated salt solutions with different deliquescence RH.

The WHOPS is operated in the dry mode to investigate the optical properties of the dry aerosol: the particles selected by the DMA are led through a bypass directly into the WELAS2. In this way it is possible to compare the dry mobility diameter of the size-selected particles with their partial scattering cross section. This relationship is

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then used to retrieve optical properties of the dry aerosol (i.e. the dry effective index of refraction; see Sect. 2.2.2). However, the WELAS 1 and 2 are connected to the same PMT and lamp in the PROMO box, which further means that the measurements can only be performed alternately and not simultaneously. Additionally the pressure change at different altitudes has to be taken into account for the actual selected diameters in the DMA.

Another  $4 \, \text{L} \, \text{min}^{-1}$  reach the portable aethalometer (AE42, MAGEE Scientific; Berkeley, USA) which contains LEDs of 7 different wavelengths between  $\lambda = 370$  and 950 nm, to measure the light absorption coefficient and retrieve the equivalent black carbon (BC) mass concentration. The flow rate is regulated by a mass flow controller; therefore corrections for pressure changes were performed.

#### 2.2 WHOPS characterization

## 2.2.1 Conversion of light scattering pulse height to optical diameters

The state-of-the-art approach for particle sizing with an optical particle size spectrometer consists of two steps: First, the instrument calibration factor relating the pulse height (voltage) to the scattering cross section (particle property) is determined with an empirical calibration using particles of known scattering cross section. Second, the relationship between scattering cross section and particle size is obtained from Mie theory (Mie, 1909; Bohren and Huffman, 2007). The combination of the instrument calibration factor with the Mie curves provides the relationship between pulse height and corresponding optical particle size. In order to obtain a meaningful optical diameter, e.g. equal to the geometric diameter for homogeneous spherical particles, it is crucial to consider all relevant properties of the light source, the detector geometry and the observed particles in the Mie calculations.

The pulse height recorded as a raw voltage ( $V_{raw}$ ) by the PROMO, which is a measure for the scattered light intensity, is directly proportional to the light intensity of the xenon lamp ( $I_{xenon}$ ) and to the particles' partial scattering cross section ( $\overline{\sigma}$ ; from now

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$$V_{\text{raw}} \propto \overline{\sigma} I_{\text{xenon}}$$
 (4)

In order to calculate instrument-specific Mie curves (i.e. scattering cross section  $\overline{\sigma}$  as a function of the particle diameter), the theoretical scattering cross section  $(\sigma(\lambda, D, m))$ , which depends on the wavelength  $(\lambda)$ , diameter (D), refractive index (m), and the solid angle  $(\Omega)$  covered by the light collection optics of the WELAS, has to be averaged over all wavelengths weighted with the relevant light intensity spectrum,  $\frac{dI_{Mie}(\lambda)}{d\lambda}$ :

$$\overline{\sigma}(D, m, \Omega) = \frac{\int \sigma(\lambda, D, m, \Omega) \frac{d'_{\text{Mie}}(\lambda)}{d\lambda} d\lambda}{\int \frac{d'_{\text{Mie}}(\lambda)}{d\lambda} d\lambda}.$$
(5)

The relevant light intensity spectrum differs from the spectrum of the xenon lamp,  $\frac{dl_{Xenon}(\lambda)}{dl}$ , as the light transmission and detection efficiency of the optical components and detector depend on the wavelength. They are related through:

$$\frac{dI_{Mie}(\lambda)}{d\lambda} = \frac{dI_{Xenon}(\lambda)}{d\lambda} f_{fibre, lamp}(\lambda) f_{lens}(\lambda) f_{fibre, PMT}(\lambda) f_{PMT}(\lambda)$$

$$\approx \frac{dI_{Xenon}(\lambda)}{d\lambda} f_{lens}(\lambda) f_{PMT}(\lambda) \tag{6}$$

where the factors  $f_{x}(\lambda)$  describe the relative wavelength dependence for the key components.  $f_{\rm fibre.\ lamb}(\lambda)$  is the light transmission efficiency between the xenon lamp and the detection volume (assumed to be constant, i.e.  $f_{\text{fibre, lamp}}(\lambda) \equiv 1$ ),  $f_{\text{lens}}(\lambda)$  is the light transmission efficiency of the lens collecting the scattered light,  $f_{\text{fibre.PMT}}(\lambda)$  is the light transmission efficiency between the lens and the PMT (assumed to be constant, i.e.  $f_{\text{fibre.PMT}}(\lambda) \equiv 1$ ), and  $f_{\text{PMT}}(\lambda)$  is light detection efficiency of the bialkali PMT. Note that the normalization of any  $f_{\nu}(\lambda)$  can randomly be chosen, as any changes of the normalization factors are cancelled out in Eq. (5). Figure 2 illustrates all important spectra and

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efficiency curves for the WELAS. The PMT efficiency ( $f_{PMT}$ ; turquoise curve in Fig. 2), which gradually decreases from its maximum at wavelengths below 400 nm down to an efficiency of 0 % at ~ 600 nm, has a strongly distorting effect on the relevant spectrum  $\frac{df_{Mie}(\lambda)}{d\lambda}$  by reducing the contribution by larger wavelengths. The lens transmission ( $f_{lens}$ ; purple curve in Fig. 2), which is unity for the dominant fraction of the lamp spectrum while it only drops at shorter wavelengths (< 400 nm), has a minor influence on  $\frac{df_{Mie}(\lambda)}{d\lambda}$ .

The raw voltage recorded for particles with a certain scattering cross section can potentially drift over time due to drifts of the light intensity of the xenon lamp ( $I_{xenon}$ ), PMT sensitivity, detector electronics, optics contamination, etc. Such temporal drifts of the WELAS sensitivity are accounted for by correcting the measured raw voltages with a time-dependent voltage calibration factor  $C_{V}$ :

$$V_{\rm cal} = C_{\rm V}(t)V_{\rm raw} \tag{7}$$

This approach, which follows the manufacturer's instructions, ensures that a certain calibrated voltage signal ( $V_{cal}$ ) always corresponds to the same scattering cross section. In practice, the voltage calibration factor is obtained by relating the raw voltage signal measured for quasi-monodisperse Caldust (calibration dust with a diameter of 850 nm and an index of refraction similar to PSL; provided by the manufacturer) at a certain time t, to the raw signal measured at a reference time  $t_0$ :

$$C_{V}(t) = \frac{V_{\text{raw}}(\text{Caldust}, t_0)}{V_{\text{raw}}(\text{Caldust}, t)}$$
(8)

An absolute calibration factor  $(C_{\sigma})$  for scattering cross section measurements is obtained by recording the calibrated voltage signal for PSL spheres with a certified diameter and relating it to results from Mie calculations:

$$C_{\sigma} = \frac{\overline{\sigma}(PSL)}{V_{cal}(PSL)}$$
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 $\overline{\sigma}(\text{PSL})$  of the PSL spheres (m=1.59+0i) is calculated using Eq. (5) to account for the instrument specific intensity spectrum. In principle, the calibration factor  $C_{\sigma}$  is expected to be independent of the PSL diameter used to determine it. However, tests with different PSL sizes revealed a weak dependence of  $C_{\sigma}$  on PSL size (see Sect. 3.1.2). Thus, a calibration function  $C_{\sigma}(V_{\text{cal}})$ , rather than a simple calibration constant, is used to infer absolute scattering cross section values from the calibrated voltages measured for any sample aerosol according to the following equation:

$$\overline{\sigma}(V_{\text{cal}}) = C_{\sigma}(V_{\text{cal}})V_{\text{cal}} \tag{10}$$

Absolute scattering cross section values, determined with Eq. (10) for any kind of aerosol sample, form the basis for meaningful optical sizing. The relationship between scattering cross section and corresponding geometric particle diameters can be calculated with Mie theory (and Eq. 5 to account for the instrument specific intensity spectrum), if the particles are spherical with a known refractive index. If the particles are non-spherical or if the refractive index is unknown, then it is only possible to infer an "equivalent optical diameter" representing the size of a sphere with a certain refractive index that has the same scattering cross section as the particle under investigation. It is further detailed in Sect. 3.1.2, how meaningful optical sizes are inferred from WELAS measurements of the dried and humidified aerosol samples probed in the WHOPS setup.

The real part of the index of refraction m of PSL shows a weak wavelength dependence, decreasing from 1.617 at  $\lambda = 400$  nm to 1.572 at 1000 nm (Kasarova et al., 2007; Ma et al., 2003). Furthermore, a non-zero imaginary part was reported from measurements at a shorter wavelength of about 260 nm (Gaigalas et al., 2013). In this study, a constant index of refraction of 1.59 + 0i across the whole relevant spectrum is used in the Mie calculations for PSL spheres. The weak wavelength dependence and the marginal imaginary part are ignored, as they only have a negligible effect.

A mean index of refraction m of dry ammonium sulfate crystals is reported in the literature as  $1.53 \pm 0.01$  (real part; Toon et al., 1976) at wavelengths between 300 and

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800 nm. This value was used in the Mie calculations for the optical sizing of pure ammonium sulfate aerosols in laboratory experiments. The assumptions made for m in the analysis of ambient WHOPS measurements are discussed in detail in Sects. 2.2.2 and 3.1.3.

Ideally, the firmware of the WELAS instrument would also report the raw voltages and calibrated voltages corresponding to the reported optical diameter values. Unfortunately, this is not the case. The WELAS only reports PSL equivalent optical diameters, which are internally calculated by the instruments' firmware using a preset empirical calibration curve ( $D_{opt}$  vs.  $V_{cal}$ ) based on PSL measurements done by the manufacturer. Thus, in this study the calibrated voltages were reconstructed from the reported diameter values using this preset calibration curve. The raw voltages corresponding to the calibrated voltages were obtained with Eq. (7).

## 2.2.2 Retrieval of index of refraction and hygroscopic growth factor

Using a lookup table is useful to avoid iterative approaches and repeated computationally expensive Mie calculations for the retrieval of the index of refraction  $(m_{drv})$  of the dry particles from the dry mode WHOPS measurement and of the hygroscopic growth factor (GF) from the wet mode WHOPS measurement. Figure 3 presents the algorithms used to infer these quantities. A similar approach was described in previous articles like in Kiselev et al. (2005) for the LACIS analysis or in Flores et al. (2009) for WELAS measurements.

Box 1 in Fig. 3 (marked by the blue dotted frame) describes how the lookup table is generated. Assuming a dry particle size  $(D_{dry})$  and a hygroscopic growth factor (GF) defined through Eq. (1), the corresponding solution droplet diameter  $(D_{wet})$  and the amount of absorbed water. The GF and the indices of refraction of the dry particle  $(m_{dry})$  and pure water  $(m_{H_2O} = 1.333)$  determine the index of refraction of the solution droplet  $(m_{\text{mix}})$ , which is calculated using volume weighted mixing rule of  $m_{\text{dry}}$  and  $m_{\text{HaO}}$ (Hale and Querry, 1973). Mie theory for homogeneous spherical particles (see also Sect. 2.2.1) is then used to calculate the scattering cross section  $(\sigma_{wet}(D_{wet}, m_{mix}))$ 

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of this humidified particle. The complete lookup table is obtained by calculating  $\sigma_{\rm wet}$  with the above approach for the relevant ranges of the parameters GF,  $D_{\rm dry}$ , and  $m_{\rm dry}$ . The lookup table is calculated only once in a preparation step, before it is repeatedly accessed in steps 2 and 3 of the WHOPS data evaluation.

The retrieval of the effective index of refraction of the dry particles from the dry mode WHOPS measurements provide the second data analysis step (red dotted frame in Fig. 3): Operating the WHOPS in the dry bypass mode (see Fig. 1) yields a measurement of the scattering cross section ( $\sigma_{\rm dry}$ ) of dry particles with a well-defined mobility diameter ( $D_{\rm dry,\ mob}$ ). For a certain  $D_{\rm dry,\ mob}$  and GF = 1 (no hygroscopic growth in the dry mode), the corresponding  $\sigma$ -m-relationship can be extracted from the lookup table. Finally the retrieved effective index of refraction ( $m_{\rm dry,\ meas}$ ) of the dry particle is obtained by selecting m corresponding to  $\sigma_{\rm dry,\ meas}$  in the extracted  $\sigma$ -m-relationship. The qualifier "effective" is used because the true index of refraction might be different due to the applied simplifications: The imaginary part of the index of refraction is assumed to be zero and all particles are assumed to be spherical and perfectly internally mixed (equal composition of all individual particles of equal size).

After the dry measurement described above, the WHOPS is switched to the wet mode (Fig. 1) such that the WELAS2 measures the scattering cross section distribution of the selected particles at high RH. Converting the  $\sigma_{\rm wet,\,meas}$  distribution to a meaningful size (GF) distribution is not straightforward since the hygroscopic water uptake gradually reduces the index of refraction of the solution droplets with increasing size (GF). Nevertheless, iterative processes can again be avoided by using the lookup table described above and the approach described in Fig. 3 (box 3, green dotted line). First the  $\sigma$ -GF-relationship is extracted for the dry size selected in the DMA ( $D_{\rm dry,\,mob}$ ) and the corresponding index of refraction of the dry particles ( $m_{\rm dry,\,meas}$ ), which was previously retrieved from the nearest dry mode WHOPS measurement as described above. Finally, the measured GF (GF<sub>meas</sub>) corresponding to the measured  $\sigma_{\rm wet,\,meas}$  from the lookup table. This allows to convert the scattering cross section axis to the corresponding

GF axis and thus to transform the scattering cross section distribution measured by the

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The scattering cross section as a function of GF (wet diameter) obtained with the

above approach is shown in Fig. 4 (green lines) for dry particle diameters of (a) 500 nm

and (b) 300 nm. The Mie curves for pure water droplets ( $m_{\rm H_2O}$  = 1.333; blue lines) and

particles with a constant index of refraction equal to that of the dry particle ( $m_{dry} = 1.45$ ;

red lines) are also shown as a reference. The red and blue lines illustrate that the

scattering cross section increases with increasing particle size at a constant index of

refraction. The blue line is below the red line for the whole size range shown in Fig. 4a,

WELAS2 to a meaningful GF distribution.

because the scattering cross section decreases with decreasing index of refraction

at constant particle size. Hygroscopic growth results in a concurrent increase of the

particle size and decrease of the index of refraction. Thus, the green curve, which starts off at the level of the red curve for the initially dry particle, gradually approaches the

blue curve with increasing growth factor. First focusing on particles with a dry diameter of  $D_{\text{drv. opt}} = 500 \,\text{nm}$  (panel a), it can be seen that the size effect dominates over the

counteracting effect from the change of the index of refraction, thus resulting in a strictly monotonically increasing relationship between the scattering cross section and the GF

of the solution droplet (green line). This makes unambiguous assignment of a GF to a measured (wet WHOPS mode) scattering cross section of a solution droplet possible.

However, one limitation is given by the Mie wiggles that appear in the green line at sizes above 1400 nm, which corresponds to GF = 2.8. Ambiguities have therefore to be considered above this GF for particles with  $D_{\text{dry, opt}} = 500 \text{ nm}$ .

On the other hand the scattering behavior of humidified particles with a dry diameter of  $D_{drv, opt} = 300 \, \text{nm}$  (panel b) is quite different: here, the scattering cross section of the growing particle (green line) decreases with initial water uptake, as the effect of changing index of refraction dominates over the size effect, and reaches a minimum at a droplet size of 360 nm, before it increases strictly monotonically at further water uptake, where the size effect dominates. This results in an ambiguous relationship between the scattering cross section and the growth factor in the range 1.00 < GF < 1.31

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for 300 nm particles, thus making it impossible to distinguish between growth factors in this range, when using the WELAS as an optical sizer. Growth factors larger than 1.31 can be unambiguously determined. The exact threshold GF depends weakly on the index of refraction of the dry particles.

The measurement of the GF-PDF for dry diameters below ~ 400 nm using the WHOPS is limited by the ambiguity in the optical sizing, described above, and the sharp drop in counting efficiency of the WELAS (see Sect. 3.1.1). While the counting efficiency could potentially be improved, resolving the ambiguity issue would require a different light spectrum.

Figure 4 also demonstrates the crucial importance of using an appropriate index of refraction for the optical sizing of grown particles, in particular the need to account for the effect of absorbed water. For instance a particle with  $D_{\rm dry,\ opt} = 500\ \rm nm$  and GF = 1.4 (i.e.,  $D_{\rm wet,\ opt} = 700\ \rm nm$ ) has a scattering cross section of  $9.91\times 10^{-14}\ \rm m^2$  according to theory (see green line in Fig. 4a). If a fixed index of refraction of m=1.45 (dry particle) or m=1.333 (pure water) were used to infer the size of this solution droplet, the resulting diameters would be 580 nm or 765 nm, respectively, according to the red or blue lines, and the corresponding inferred GFs would be 1.16 or 1.53. This would be substantially different (-17% or +9%) from the true value of 1.4.

## 2.3 Laboratory measurements

Aqueous suspensions of PSL size standards (Duke Standards, Thermo Scientific) were diluted with MilliQ water, nebulized, led through a drier and subsequently measured in the WELAS2. PSL experiments were used to infer the scattering cross section calibration function (Eqs. 9 and 10 and Sect. 3.1.2) as well as for comparison with the manufacturer's calibration.

On the other hand ammonium sulfate (AS; Fluka Biochemica; purity > 99.5%) particles were used to test the optical sizing for atmospherically relevant particles with an index of refraction that differs from that of PSL, to test the method to retrieve the

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effective index of refraction of a dried monodisperse particle sample, and to determine the counting efficiency of the WELAS. For this purpose, an ammonium sulfate and MilliQ water solution was prepared, nebulized, dried and passed through a neutralizer before being fed to a DMA, in order to select particles with a well-defined mobility diameter for testing the WELAS.

For monodisperse samples such as PSL spheres or size-selected AS, the mean value of the measured signal distribution was used as a representative value (ignoring the signals from multiply charged particles in the case of size-selected AS samples).

#### 2.4 Field measurements

Within the Pan-European Gas-Aerosols-climate interaction Study (PEGASOS; pegasos.iceht.forth.gr) a Zeppelin NT was used to explore the planetary boundary layer (for more information see Stull, 1988). The PSI-Zeppelin rack containing the WHOPS, the aethalometer and an extra WELAS was mounted in the airship to investigate properties of secondary organic aerosol (SOA). The first campaign took place in the Netherlands, in the region of Rotterdam, in May 2012. Simultaneously, chemical composition and size distributions were recorded by an Aerodyne aerosol mass spectrometer (AMS; Rubach, 2013) and a scanning mobility particle sizer (SMPS; e.g. Wiedensohler et al., 2012), respectively. Besides, various gas phase measurements (e.g. NO<sub>x</sub>, CO, OH, O<sub>3</sub>) together with GPS (global positioning system) and meteorological data were logged.

During the Zeppelin NT flights the measurement cycle per probed particle size was chosen to be 250 s, where 150 s are used for the wet mode and 100 s for the dry mode WHOPS measurement. The dry diameters selected were 300 and 500 nm. The polydisperse size distribution was recorded for 50 s. Equivalent BC concentrations were logged with a resolution of 2 min.

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#### Characterization of the WELAS and WHOPS

#### 3.1.1 Counting efficiency

The size-resolved counting efficiency (CE) of the WELAS was tested to verify the number concentration measurements and to determine the lower cut-off diameter (scattering cross section). For this purpose, a comparison with a condensation particle counter (CPC 3025a, TSI Inc.) using quasi-monodisperse ammonium sulfate particles sizeselected by a DMA was performed. The measured counting efficiency expressed as a function of the raw voltage signal,  $\varepsilon_{\rm CE, meas}(V_{\rm raw})$ , is defined as:

$$\varepsilon_{\text{CE, meas}}(V_{\text{raw}}) = \frac{n_{\text{Welas, raw}}(V_{\text{raw}})}{n_{\text{CPC}}},$$
 (11)

where  $n_{\text{Welas. raw}}(V_{\text{raw}})$  is the particle number concentration measured by the WELAS for a particle size that produces a raw voltage signal of  $V_{\rm raw}$ , and  $n_{\rm CPC}$  is the corresponding value reported by the CPC. The "raw pulse height" ( $V_{\text{raw}}$ ) is chosen as a reference axis for the counting efficiency, as  $V_{raw}$ , rather than particle size or scattering cross section, determines the detection probability. Note that  $V_{raw}$  is inferred from the instrument output as described in Sect. 2.2.1. This calculation involves the WELAS-internal voltage calibration factor  $C_V$  (Eq. 8), which underlines that the knowledge of this factor is essential.

Figure 5 shows the measured WELAS counting efficiency as a function of the raw pulse height. The plateau value, reached for  $V_{\text{raw}} > 8 \,\text{mV}$ , is at  $\sim 132 \,\%$ . This means that the WELAS reports higher number concentration readings than the CPC. Such a bias was already observed by Heim et al. (2008), who explained it by the fact that the size of the measuring volume in the WELAS used in the firmware calculations is incorrect, thus leading to a bias (fixed factor) in the concentration readings. As a consequence, the number concentration readings of this particular WELAS have to be

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corrected whenever absolute number concentration values are important, given the fact that the disagreement between the instruments is larger than the accuracy of the CPC ( $\pm 10\,\%$  for particle number concentrations  $< 10^5\,\mathrm{cm}^{-3}$  and in the particle size range, which is of interest in this study). The measurements shown in Fig. 5 were performed with several Xe-lamps of the same type but different amount of hours of operation. The counting efficiency as a function of  $V_{\mathrm{raw}}$  was equal within experimental uncertainty for all lamps, as expected. Consequently the efficiency curve presented in Fig. 5 is valid for all measurements carried out with this lamp type, as long as the fiber transmission remains constant.

In the range of the lower detection limit, the counting efficiency for particles with a certain scattering cross section is expected to depend on the incident light intensity, as  $V_{\text{raw}}$  is proportional to it (Eq. 4). However, the voltage calibration factor  $C_{\text{V}}$  varied less than 10%, thus indicating a small variability in lamp intensity. This makes an unambiguous assignment of optical diameters to the primary  $V_{raw}$  axis possible (bottom axes in Fig. 5). A decrease of the efficiency for small particles can be explained as follows: the intensity of the light scattered by a particle has to be above a certain threshold to be detected by the scattering detector. Fluctuations of the ratio between scattering cross section and the raw voltage signal due to random noise are responsible for the fact that only a certain percentage of particles with a scattering cross section slightly above/below the lower detection limit will be measured/not be recorded. The finite width of the DMA transfer function has little influence here, as it is much smaller than the observed width of the lower cut-off of the WELAS. The raw voltages at 50%, 90 % and 99 % of the plateau value of the counting efficiency are 3.7, 5.2, and 7.7 mV, respectively (Fig. 5). In order to relate these  $V_{raw}$  to actual optical diameters a mean  $C_{\rm M} = 2.23$  was used to calculate corresponding diameters for PSL (m = 1.59), ammonium sulfate (m = 1.53) and water (m = 1.333) seen in Fig. 5 as additional abscissa scales. The equivalent optical diameters for 50 %, 90 % and 99 % of the plateau value are  $\sim 280$  nm,  $\sim 325$  nm and  $\sim 395$  nm for PSLs and  $\sim 330$  nm,  $\sim 390$  nm and  $\sim 430$  nm for ammonium sulfate. This lower detection limit is significantly higher than specified by

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the manufacturer (200 nm for PSLs). The largest tested particle size corresponds to a PSL equivalent optical diameter of approximately 600 nm. The counting efficiency is expected to remain constant at the observed plateau value for larger diameters, at least in the submicron size range, where potential impaction losses are unimportant.

In the following all data was corrected with the counting efficiency curve shown in Fig. 5, in order to correct for the concentration bias as well as the decreasing detection efficiency in the range of the lower detection limit.

## 3.1.2 Calibration of the optical sizing

The data analysis and calibration approach to retrieve quantitative scattering cross section values and optical diameters from the raw pulse height signals of the WELAS is described in detail in Sect. 2.2.1. The calibration factor  $C_{\sigma}$  for the scattering cross section is obtained by measuring PSL spheres with a certified size (see Eq. 9).  $C_{\sigma}$  is expected to be independent of the raw pulse height, i.e. independent of the size of the PSL spheres used to determine it. In order to test this, measurements with PSL spheres of nine different diameters between 260 nm and 1  $\mu$ m were performed, thereby covering the range of interest for the WHOPS measurements.

Figure 6 shows the comparison of the scattering cross section measured for the PSL particles with the theoretical Mie curve calculated for the WELAS specifications as described in Sect. 2.2. The squares in different colors depict four different experiments (distributed over the course of five months) and the abscissa denotes the PSL diameters. For Fig. 6a, a constant scattering cross section calibration factor ( $C_{\sigma} = 3.71 \times 10^{-15} \,\mathrm{m^2\,mV^{-1}}$ ) was applied in Eq. (10) to all experimental data ( $V_{\mathrm{cal}}$ ).  $C_{\sigma}$  was chosen such to minimize the " $\chi$ -square value" between the single points and theoretical curve. Figure 6a reveals that the main features of the theoretical Mie curve, like the shoulder between 300–400 nm, are well reproduced by the WELAS measurement. However, PSL particles with diameters above 800 nm seem to scatter about 10 % less than predicted by Mie theory and small particles with diameters < 400 nm scatter approximately 10 % more than expected. Sensitivity studies with variations of the input

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parameters in the Mie calculations showed that this systematic size-dependent bias cannot be significantly reduced by changing the scattering angles, the light spectrum, the index of refraction or assuming a narrow width of the PSL size distribution rather than a perfectly monodisperse aerosol. The exact reason for this minor disagreement is currently unknown. A possible explanation could be a non-linear behavior of the detector or electronic parts amplifying and processing these signals.

The systematic deviations observed in Fig. 6a imply that  $C_{\sigma}$  is not a constant, contrary to expectations. Therefore,  $C_{\sigma}$  was separately determined for each PSL data point, using Eq. (9), and plotted against the calibrated voltage (Fig. 7). This analysis reveals a systematic increase of the calibration factor  $C_{\sigma}$  with increasing voltage signal. Therefore, we decided to use a voltage dependent calibration function,  $C_{\sigma}(V_{\rm cal})$ , for the quantitative analysis of the WELAS signals (see Eq. 10). A linear fit,  $C_{\sigma}(V_{\rm cal})\left[\frac{{\rm m}^2}{{\rm mV}}\right] = 1.25\times 10^{-17}\frac{{\rm m}^2}{({\rm mV})^2}V_{\rm cal}[{\rm mV}] + 3.25\times 10^{-15}\frac{{\rm m}^2}{{\rm mV}}$ , was selected for this calibration function. 85% of the measurement points lie within a 15% deviation band around the linear fit, indicating that the scattering cross section can be measured within a 2- $\sigma$  uncertainty of 15% when using this calibration function. Figure 6b shows the PSL measurements analyzed with this calibration function. The fact that the systematic deviations presented in Fig. 6a disappeared justifies the use of this calibration function for all further data analyses.

To retrieve the uncertainty in the optical diameter measurement, the 15% uncertainty of the scattering cross section measurement is used as input for error propagation calculations. Figure 8 depicts the absolute values of the resulting estimated relative uncertainty of the optical diameters for four different indices of refraction: 1.59, 1.50, 1.40 and 1.333, representing PSL, the upper and lower end of the range often found for atmospheric aerosol and pure water, respectively. The uncertainty for optical diameters, resulting from a fixed scattering cross section uncertainty, depends on the local gradient of the Mie curve and thus on the particle size and index of refraction. This explains why the relative sizing uncertainty shown in Fig. 8 strongly depends on diameter and index of refraction. The average relative sizing uncertainty for these indices of refraction,

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in the range between 300 nm and 1  $\mu$ m, is 9 %, with minimal and maximal values of 4 % and 22 %, respectively, for m = 1.50 at 0.469  $\mu$ m and m = 1.333 at 0.300  $\mu$ m. For the optical sizing of hygroscopic particles at high RH, which have an index of refraction close to the one of pure water (m = 1.333), follows that the relative sizing uncertainty remains below 8 % for solution droplets with diameters above 400 nm. Below approximately 200 nm the uncertainty would be smallest (< 4 %) and independent of the index of refraction, however, this is below the detection limit of the WELAS.

#### 3.1.3 Retrieval of the index of refraction of PSL and ammonium sulfate

The dry effective index of refraction  $(m_{\rm dry})$  of an aerosol sample can be retrieved from the dry mode WHOPS measurement, as described in Sect. 2.2.2. To verify this procedure, tests were performed in the laboratory with PSL spheres and size-selected (by the DMA in the WHOPS) ammonium sulfate  $((NH_4)_2SO_4)$  particles, which were prepared as described in Sect. 2.3. Note that the size-selected ammonium sulfate aerosol sample also contains multiply charged particles which are larger than the singly charged particles. However, they appear as a well-separated mode in the WELAS size distribution measurement and are neglected for this analysis.

Figure 9 illustrates the retrieved effective indices of refraction for particle diameters between 290 nm and 1000 nm. The four different experiments with certified PSL spheres are the same as those already shown in Figs. 6 and 7. The retrieved index of refraction for PSLs was found to be  $m=1.60\pm0.04$  (mean  $\pm1$  SD) which agrees well with the literature value of  $m\approx1.59$  (see Sect. 2.2.1). These four PSL experiments were used to determine the scattering cross section calibration function  $C_{\sigma}(V_{\rm cal})$  for the WELAS (see Sect. 2.2.1/Eq. 10 and Sect. 3.1.2/Fig. 7). Therefore the retrieval of the index of refraction for the very same PSL experiments is an internal consistency check of the instrument calibration and data analysis approaches. Besides this, the scatter of the PSL data points in Fig. 9 reveals the precision of the retrieved index of refraction (1 SD  $\approx\pm0.04$  absolute), which is given by the random noise of the individual PSL data points around the mean calibration function (see Fig. 7).

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The dry mode WHOPS measurements of the size-selected ammonium sulfate particles provide an independent test of the index of refraction retrieval for an aerosol sample that has different optical properties and that was not used for calibration purposes (except for testing the counting efficiency, which is independent of the optical sizing). The retrieved index of refraction for dry ammonium sulfate particles is found to be  $m = 1.49 \pm 0.02$  (mean  $\pm 1$  SD) which is  $\sim 2.3$ % lower than the mean literature value of m = 1.53 (Toon et al., 1976), which is found by weighing the stated vales for the wavelength range of interest. This systematic difference is small, given the facts that the selected mobility diameter is tainted with some uncertainty and that part of it can likely be explained with slight non-sphericity of dry ammonium sulfate particles produced by nebulizing a solution and subsequent drying. Zelenyuk et al. (2006) reported a decreasing effective density with increasing mobility diameter for ammonium sulfate particles. This is qualitatively consistent with the observed increasingly negative bias in the retrieved index of refraction, though non-sphericity may also have some influence on the light scattering cross sections.

In summary, the retrieval of the effective index of refraction of the dry particles gives a good approximation, given the fact that the main use of it is to ensure that meaningful growth factors are obtained by relating the inferred optical diameters to the selected mobility diameters, as described in Sect. 2.2.2.

## 3.1.4 Measured hygroscopicity of ammonium sulfate particles

The accuracy of hygroscopic growth factors measured by the WHOPS in the wet mode was verified against Köhler theory using pure ammonium sulfate particles which were generated as described in Sect. 2.3. The WHOPS measurements were analyzed using the approach presented in Sect. 2.2.2 and Fig. 3, i.e. with accounting for the effect of absorbed water on the index of refraction of the solution droplets. Figure 10a shows a typical example of the GFs measured by the WHOPS for different dry particles diameters (red markers), along with the relevant relative humidity (blue markers) simultaneously measured in the WHOPS. The error bars on the measured GFs reflect

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the uncertainty of the optical diameters (taken from Fig. 8) and mobility diameters (assumed to be ±5%). The corresponding theoretical growth factors, calculated for the measured RH using Köhler theory (ADDEM Model; Topping et al., 2005), are shown with black markers. As explained in Sect. 2.1, the RH sensors applied in the WHOPS <sub>5</sub> have an absolute accuracy of ~ ±2 % RH. The error bars around the theoretical GFs depict the propagated uncertainty corresponding to this RH uncertainty. The agreement between the measured and theoretical GF is very good for the example shown in Fig. 10a. In order to determine the long-term performance of the WHOPS, results of 8 independent ammonium sulfate experiments distributed over 7 months are shown in Fig. 10b. The crosses depict the mean relative deviation of the measured from the theoretical GF at the same RH (between 89–97%), while the error bars show the standard deviation. The right y axis in Fig. 10b explains what absolute RH bias would be necessary to explain the corresponding GF bias shown on the left y axis. The mean relative deviation of measured GFs from theory is almost negligible with values between -3% and +1% for different dry diameters, which corresponds to an absolute RH bias between -0.52 % and +0.15 %. The fact that also all error bars fall between -9% and +6% GF bias, which corresponds to -1.5% and +0.7% absolute RH bias, indicates that all individual data points fall within the expected uncertainty, this in view of the fact that the uncertainty of the RH measurement, which is likely the main limiting factor of growth factor measurements by the WHOPS, is estimated to be as much as ±2% (absolute).

Figure 10 shows that it is possible to accurately measure the GF of pure ammonium sulfate particles between 300 nm and 531 nm. The measurements for a dry diameter as small as 300 nm are only possible, since the ammonium sulfate solution droplets grow clearly above the lower WELAS detection limit, which is at ~ 390 nm for dry ammonium sulfate particles (Sect. 3.1.1 and Fig. 5). Furthermore, the GF of 300 nm ammonium sulfate particles is also above the threshold, which makes unambiguous optical sizing possible even at this dry diameter (Sect. 2.2.2 and Fig. 4). However, it is not possible to measure the complete growth factor distribution of aerosol samples that contain non- or

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only slightly hygroscopic particles for dry sizes below ~ 400 nm, as they would remain undetected by the WELAS, thereby biasing the measured GF distribution.

#### 3.2 First field measurements

Here we present first results from an airborne application of the WHOPS on board the Zeppelin NT during one long-distance flight in the Netherlands which took place on 22 May 2012 in the afternoon. The Zeppelin started at the international airport of Rotterdam – The Hague, then reached the Cabauw Experimental Site for Atmospheric Research (CESAR)  $\sim 40\,\mathrm{km}$  to the east of Rotterdam, where it turned back westwards, passed Rotterdam again, flew over the North Sea until  $\sim 20\,\mathrm{km}$  from the coast, before turning back eastwards and finally landed at the airport (see Fig. 11a). The average altitude flown was 200 m above ground, except for the part over the sea, marked with a red circle in Fig. 11a, where a short height profile was performed reaching up to 700 m above ground, before flying back to the airport at an altitude of 200 nm above ground.

Figure 11b displays the mean GFs at 95 % RH (color coded; averaged over 50 s) measured by the WHOPS along the flight track for 500 nm particles. The measured GFs were not recalculated for small deviations between the target RH of 95 % and the actual RH measured in the WHOPS (corrected for equilibration time effects), because the latter was always within 95±0.6 %. For the 500 nm particles mean GFs were found to be between 1.74 and 2.67, with a median value of 1.94. In order to compare the aerosol hygroscopicity measurements made in this study with literature data, the hygroscopicity parameter  $\kappa$  was calculated from the measured mean GF according to Eqs. (2) and (3) (see Sect. 1 and Petters and Kreidenweis, 2007). The WHOPS-derived  $\kappa$  values are listed in Table 1, along with literature data for similar aerosol types. The WHOPS values reach from  $\kappa$  = 0.21 to 0.93, with a median  $\kappa$  of 0.33 for  $D_{\rm dry}$  = 500 nm. This matches well with the range of kappa values of 0.3±0.1 (mean ± SD) and 0.7±0.2 that are considered to be representative for continental and marine aerosols, respectively, based on a review of cloud condensation nuclei measurements (Andreae and

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Rosenfeld, 2008). Pringle et al. (2010) performed global model simulations and reported a  $\kappa$  value of 0.27 ± 0.21 for continental aerosol. Higher  $\kappa$  values of 0.4–0.6 and 0.9–1.0 were reported for mixed marine/continental aerosols and remote marine aerosols respectively. These values cover well the range from in-situ measurements.

The time series of the measured mean GFs and  $\kappa$  values for 500 nm particles are shown in Fig. 11c with black symbols. Somewhat higher  $\kappa$  values (median = 0.43) were found during the first part of the flight, when the Zeppelin headed inland towards Cabauw and back to Rotterdam (from now on referred to as "Part 1"), while lower  $\kappa$  values (median = 0.31) were found during the second part of the flight which was directed towards the sea and back to Rotterdam (from now on referred to as "Part 2"). One difference that was found between these two regions was the prevailing local wind direction, which was measured on board of the Zeppelin NT (Fig. 11c; the change in local wind direction is marked by the dashed, grey line): during "Part 1", north-westerly winds were dominant, while during "Part 2" north-easterly winds prevailed. Obviously, the Zeppelin entered a different air mass during the second part of the flight. Back trajectory analysis (not shown) reveals that the probed air masses had maritime and/or continental influences but it is not possible to assign well defined air mass types to these two parts of the flights.

Figure 11d illustrates the averaged GF distribution (GF-PDF) for a dry diameter of 500 nm, where the dashed turquoise line and the solid blue line illustrate the results for "Part 1" and "Part 2", respectively. Particles covering the whole GF range between approximately 1 up to  $\sim$  3 were observed. This does not necessarily imply external mixture at a certain time. However, the fact that particles with GFs substantially smaller and larger than the observed minimal and maximal mean GF occurred, indicates that the aerosol was at least sometimes externally mixed (a time-resolved analysis of the mixing state is not possible due to the limited counting statistics). On average, about 15% of the particles were non- or slightly-hygroscopic with GF < 1.1, which could be explained with externally mixed dust (e.g. Herich et al., 2009), soot (e.g. Tritscher et al., 2011) or biological particles (e.g. Després et al., 2012). Around

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74% of the particles were "more hygroscopic" with GF > 1.5 at the dry size of 500 nm. These particles must have contained a high fraction of inorganic salts, which are the most hygroscopic aerosol components (e.g. Topping et al., 2005). The remaining 11% of the particles were "less hygroscopic" with GFs between 1.1 and 1.5, most likely an 5 internal mixture of organic/inorganic compounds (e.g. Gysel et al., 2004; Meyer et al., 2009). The dominance of more hygroscopic particles explains the relatively large mean and median GFs and  $\kappa$  values presented above. One clear distinction between the two air masses is the shift of the more hygroscopic fraction to smaller GFs during "Part 2" compared to "Part 1", while the non- and less-hygroscopic fractions remained essentially unchanged. During "Part 1", the main, more hygroscopic mode peaked at a GF of  $\sim 2.5$  ( $\kappa \sim 0.8$ ). This is more hygroscopic than e.g. pure ammonium sulfate or ammonium nitrate particles (e.g. Petters and Kreidenweis, 2007; Topping et al., 2005) and can only be explained with very hygroscopic sea salt particles (pure NaCl, NaNO<sub>2</sub> and Na<sub>2</sub>SO<sub>4</sub> particles have  $\kappa$  values at RH = 95% of 1.32, 0.86 and 0.67, respectively). Sea salt particles are frequently observed in the accumulation mode aerosol of marine environments (Swietlicki et al., 2008). During "Part 2", the main, more hygroscopic mode shifted towards smaller GFs peaking at a GF of  $\sim 2$  ( $\kappa \sim 0.4$ ), which is typical for a continental background aerosol (Swietlicki et al., 2008) although aged sea salt particles with a substantial organic fraction would also be possible.

For the 300 nm particles, shown in Fig. 11f, the GF-PDF could only be measured by the WHOPS for the range GF > 1.5, due to the lower detection limit of the WELAS (see Sect. 3.1.4). This leads to incomplete GF-PDFs and thus also to a high bias in the inferred mean GFs. Nevertheless, these measurements still provide some useful insight. The GF-PDF shown in Fig. 11f also reveals a similar trend with the air mass change as seen for the 500 nm particles diameter: The more hygroscopic mode shifts from very high GFs during "Part 1" towards slightly smaller GFs during "Part 2". Figure 11f further reveals that the averaged GF-PDF in the range GF > 1.5 is guite similar for the 300 nm and 500 nm particles, although the 300 nm particles seem to be shifted towards just slightly larger growth factors. The lower detection limit of the WELAS made

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it impossible to determine whether the number fraction of particles with GF < 1.5 is also similar for the two investigated dry sizes. For a spatially resolved comparison of the GFs between the two selected aerosol sizes, Fig. 11d additionally contains the mean GF of all particles with GF > 1.5 for the 300 nm (green symbols) and the 500 nm particles (red symbols). This comparison illustrates a similar spatial pattern for both sizes, while the mean GF values of the more hygroscopic particles are just slightly larger at 300 nm dry diameter, as already seen with the observed GF-PDFs. However, this subtle difference is within the experimental uncertainty of the WHOPS. The time series of the "mean GF" of the 300 nm particles shown in Fig. 11d further reveals little spatial (horizontal and vertical) variability, except for the air mass change between "Part 1" and "Part 2". This is not unexpected, as the Zeppelin always stayed within the well-mixed boundary layer during this flight. Somewhat larger variability in the mean GF of the 500 nm particles can likely be attributed to limited counting statistics at this size.

Zieger et al. (2011) reported aerosol hygroscopicity data from previous measurements at the CESAR site in Cabauw, at which the Zeppelin passed by in this study, determined with two different experimental techniques: monodisperse HTDMA measurements at RH = 90 % as well as polydisperse humidified nephelometer (Wet-Neph) measurements. These two approaches differ from the WHOPS technique with regard to the selected/representative dry diameters as well as the set RH (see Table 1 for more details).

The  $\kappa$  values determined by Zieger et al. (2011) using the Wet-Neph and size distribution measurements were between 0.14 and 0.53 (median = 0.29) and those derived from the HTDMA measurement lied between 0.11 and 0.44 (median = 0.26), which compares well to the WHOPS data of this study, even though these measurements were taken at different times and dry sizes. The lower  $\kappa$  values of the HTDMA were explained by the not fully detected contribution of sea salt above 165 nm (maximum dry diameter set in the HTDMA). Zieger et al. (2011) could further classify the observations into different air mass types with distinct aerosol hygroscopic properties, including marine, polluted marine and continental influence. Within this study, it was

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further observed that the particle hygroscopicity was increased during periods with marine air masses. This further supports the above interpretation of the two air mass types encountered in this study during the first and second parts of the flight.

Using the approach described in Sect. 2.2.2, it was also possible for the 500 nm particles to retrieve the effective index of refraction of the dry particles which was found to be 1.42 ± 0.01 (mean ± 1 SD; the estimated measurement uncertainty is ±0.04) for either part of the flight. This compares well with literature values for the index of refraction in the visible wavelength range of 1.44–1.52 for marine aerosol (Stock et al., 2011), while it is somewhat lower than values of 1.5–1.57 reported for continental aerosol (Stock et al., 2011; Ebert et al., 2002).

#### 4 Conclusions

We developed a new airborne instrument to quantify the hygroscopic growth of large aerosol particles above a mobility diameter of approximately 280 nm at defined high relative humidity, typically around 95 %. One main advantage is that the system is able to select dry particle sizes up to approximately 630 nm and therefore to investigate the optical and hygroscopic properties of particles in the optically relevant size range. Besides, information on the mixing state and the particles' refractive index can be inferred. The instrument was further employed in a Zeppelin NT to perform vertical and horizontal profiles of the hygroscopic growth factor in the lower part of the troposphere.

The calibration and data analysis procedures for the optical sizing were successfully verified in laboratory experiments using monodisperse PSL and ammonium sulfate particles. The uncertainty for the optical sizing was found to be  $\pm 9$ %. Measured hygroscopic growth factors of pure ammonium sulfate particles agreed on average within 3% with theory, with a maximal difference of 9%. The index of refraction retrieved for pure ammonium sulfate particles agreed within 2% with literature values, which is within experimental uncertainty.

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During first airborne measurements in the Netherlands, GFs (mean value of the GF distribution) at RH = 95 % between 1.74 and 2.67 with a median of 1.94 for particles with a dry diameter of 500 nm were observed. The measured growth factor distributions for these particles revealed an externally mixed aerosol containing non-, less- and more-hygroscopic particles, including clear evidence for the presence of sea salt particles at certain times. The hygroscopic properties observed in this study are well comparable to literature values from previous ground-based measurements in the area. The mean effective index of refraction for 500 nm particles was found to be  $1.42 \pm 0.04$  (real part).

Acknowledgements. This work was supported by the EC project PEGASOS, funded by the European Commission under the Framework Programme 7 (FP7-ENV-2010-265148). We thank all PEGASOS participants for the excellent team work during the campaigns. Additionally we thank Frank Holland for the analysis and maintenance of all meteorological and GPS data on the Zeppelin.

P. Zieger was supported by a fellowship of the Swiss National Science Foundation (grant no. P300P2\_147776). M. Gysel was supported by the ERC under grant 615922-BLACARAT.

#### References

Andreae, M. O. and Rosenfeld, D.: Aerosol-cloud-precipitation interactions, Part 1. The nature and sources of cloud-active aerosols, Earth-Sci. Rev., 89, 13–41, doi:10.1016/j.earscirev.2008.03.001, 2008.

Bohren, C. F. and Huffman, D. R.: Absorption and Scattering of Light by Small Particles, Wiley-VCH Verlag GmbH, 530, 2007.

Després, V. R., Huffman, J. A., Burrows, S. M., Hoose, C., Safatov, A. S., Buryak, G., Fröhlich-Nowoisky, J., Elbert, W., Andreae, M. O., Pöschl, U., and Jaenicke, R.: Primary biological aerosol particles in the atmosphere: a review, Tellus B, 64, 15598, doi:10.3402/tellusb.v64i0.15598, 2012.

Ebert, M., Weinbruch, S., Rausch, A., Gorzawski, G., Helas, G., Hoffmann, P., and Wex, H.: Complex refractive index of aerosols during LACE 2010; as derived from the analysis of individual particles, J. Geophys. Res.-Atmos., 107, 8121, doi:10.1029/2000JD000195, 2002.

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- Fierz-Schmidhauser, R., Zieger, P., Wehrle, G., Jefferson, A., Ogren, J. A., Baltensperger, U., and Weingartner, E.: Measurement of relative humidity dependent light scattering of aerosols, Atmos. Meas. Tech., 3, 39–50, doi:10.5194/amt-3-39-2010, 2010.
- Flores, J. M., Trainic, M., Borrmann, S., and Rudich, Y.: Effective broadband refractive index retrieval by a white light optical particle counter, Phys. Chem. Chem. Phys., 11, 7943–7950, 2009.
- Gaigalas, A. K., Wang, L., and Choquette, S.: Measurement of scattering and absorption cross sections of microspheres for wavelengths between 240 nm and 800 nm, J. Res. Natl. Inst. Stan., 118, 1–14, 2013.
- Gysel, M., Weingartner, E., Nyeki, S., Paulsen, D., Baltensperger, U., Galambos, I., and Kiss, G.: Hygroscopic properties of water-soluble matter and humic-like organics in atmospheric fine aerosol, Atmos. Chem. Phys., 4, 35–50, doi:10.5194/acp-4-35-2004, 2004.
  - Hale, G. M. and Querry, M. R.: Optical constants of water in the 200-nm to 200-μm wavelength region, Appl. Optics, 12, 555–563, doi:10.1364/AO.12.000555, 1973.
- Heim, M., Mullins, B. J., Umhauer, H., and Kasper, G.: Performance evaluation of three optical particle counters with an efficient "multimodal" calibration method, J. Aerosol Sci., 12, 1019– 1031, 2008.
  - Herich, H., Tritscher, T., Wiacek, A., Gysel, M., Weingartner, E., Lohmann, U., Baltensperger, U., and Cziczo, D. J.: Water uptake of clay and desert dust aerosol particles at sub- and supersaturated water vapor conditions, Phys. Chem. Chem. Phys., 11, 7804–7809, doi:10.1039/b901585j, 2009.
  - Kasarova, S. N., Sultanova, N. G., Ivanov, C. D., and Nikolov, I. D.: Analysis of the dispersion of optical plastic materials, Opt. Mater., 29, 1481–1490, doi:10.1016/j.optmat.2006.07.010, 2007.
- Kiselev, A., Wex, H., Stratmann, F., Nadeev, A., and Karpushenko, D.: White-light optical particle spectrometer for in situ measurements of condensational growth of aerosol particles, Appl. Optics, 44, 4693–4701, doi:10.1364/AO.44.004693, 2005.
  - Liu, B. Y. H., Pui, D. Y. H., Whitby, K. T., Kittelson, D. B., Kousaka, Y., and McKenzie, R. L.: Aerosol mobility chromatograph new detector for sulfuric acid aerosols, Atmos. Environ., 12, 99–104, doi:10.1016/0004-6981(78)90192-0, 1978.
  - Ma, X., Lu, J. Q., Brock, R. S., Jacobs, K. M., Yang, P., and Hu, X.-H.: Determination of complex refractive index of polystyrene microspheres from 370 to 1610 nm, Phys. Med. Biol., 48, 4165, doi:10.1088/0031-9155/48/24/013, 2003.

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Meyer, N. K., Duplissy, J., Gysel, M., Metzger, A., Dommen, J., Weingartner, E., Alfarra, M. R., Prevot, A. S. H., Fletcher, C., Good, N., McFiggans, G., Jonsson, Å. M., Hallquist, M., Baltensperger, U., and Ristovski, Z. D.: Analysis of the hygroscopic and volatile properties of ammonium sulphate seeded and unseeded SOA particles, Atmos. Chem. Phys., 9, 721-732, doi:10.5194/acp-9-721-2009, 2009.

Morgan, W. T., Allan, J. D., Bower, K. N., Esselborn, M., Harris, B., Henzing, J. S., Highwood, E. J., Kiendler-Scharr, A., McMeeking, G. R., Mensah, A. A., Northway, M. J., Osborne, S., Williams, P. I., Krejci, R., and Coe, H.: Enhancement of the aerosol direct radiative effect by semi-volatile aerosol components: airborne measurements in North-Western Europe, Atmos. Chem. Phys., 10, 8151-8171, doi:10.5194/acp-10-8151-2010, 2010.

Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud condensation nucleus activity, Atmos. Chem. Phys., 7, 1961-1971. doi:10.5194/acp-7-1961-2007, 2007.

Pringle, K. J., Tost, H., Pozzer, A., Pöschl, U., and Lelieveld, J.: Global distribution of the effective aerosol hygroscopicity parameter for CCN activation, Atmos. Chem. Phys., 10, 5241-5255. doi:10.5194/acp-10-5241-2010, 2010.

Rubach, F.: Ph.D. thesis: Aerosol Processes in the Planetary Boundary Layer: High Resolution Aerosol Mass Spectrometry on a Zeppelin NT Airship, Doktorgrad, Fachbereich Mathematik und Naturwissenschaften, Bergische Universität Wuppertal, 2013.

Sorooshian, A., Hersey, S. P., Brechtel, F. J., Corless, A., Flagan, R. C., and Seinfeld, J. H.: Rapid, size-resolved aerosol hygroscopic growth measurements: differential aerosol sizing and hygroscopicity spectrometer probe (DASH-SP), Aerosol Sci. Tech., 42, 445–464, 2008.

Stock, M., Cheng, Y. F., Birmili, W., Massling, A., Wehner, B., Müller, T., Leinert, S., Kalivitis, N., Mihalopoulos, N., and Wiedensohler, A.: Hygroscopic properties of atmospheric aerosol particles over the Eastern Mediterranean: implications for regional direct radiative forcing under clean and polluted conditions, Atmos. Chem. Phys., 11, 4251-4271, doi:10.5194/acp-11-4251-2011, 2011.

Stratmann, F., Kiselev, A., Wurzler, S., Wendisch, M., Heintzenberg, J., Charlson, R. J., Diehl, K., Wex, H., and Schmidt, S.: Laboratory studies and numerical simulations of cloud droplet formation under realistic supersaturation conditions, J. Atmos. Ocean. Tech., 21, 876-887, doi:10.1175/1520-0426(2004)021<0876:LSANSO>2.0.CO;2, 2004.

Stull, R. B.: An Introduction to Boundary Layer Meteorology, Kluwer Academic Publisher, Vancouver, Canada, 673 pp., 1988.

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Swietlicki, E., Hansson, H. C., Hameri, K., Svenningsson, B., Massling, A., McFiggans, G., McMurry, P. H., Petaja, T., Tunved, P., Gysel, M., Topping, D., Weingartner, E., Baltensperger, U., Rissler, J., Wiedensohler, A., and Kulmala, M.: Hygroscopic properties of submicrometer atmospheric aerosol particles measured with H-TDMA instruments in various environments – a review, Tellus B, 60, 432–469, doi:10.1111/j.1600-0889.2008.00350.x, 2008.

Toon, O. B., Pollack, J. B., and Khare, B. N.: Optical-constants of several atmospheric aerosol species – ammonium-sulfate, aluminum-oxide, and sodium-chloride, J. Geophys. Res.-Oceans, 81, 5733–5748, doi:10.1029/JC081i033p05733, 1976.

Topping, D. O., McFiggans, G. B., and Coe, H.: A curved multi-component aerosol hygroscopicity model framework: Part 1 – Inorganic compounds, Atmos. Chem. Phys., 5, 1205–1222, doi:10.5194/acp-5-1205-2005, 2005.

Tritscher, T., Juranyi, Z., Martin, M., Chirico, R., Gysel, M., Heringa, M. F., DeCarlo, P. F., Sierau, B., Prevot, A. S. H., Weingartner, E., and Baltensperger, U.: Changes of hygroscopicity and morphology during ageing of diesel soot, Environ. Res. Lett., 6, 034026, doi:10.1088/1748-9326/6/3/034026, 2011.

Wiedensohler, A., Birmili, W., Nowak, A., Sonntag, A., Weinhold, K., Merkel, M., Wehner, B., Tuch, T., Pfeifer, S., Fiebig, M., Fjäraa, A. M., Asmi, E., Sellegri, K., Depuy, R., Venzac, H., Villani, P., Laj, P., Aalto, P., Ogren, J. A., Swietlicki, E., Williams, P., Roldin, P., Quincey, P., Hüglin, C., Fierz-Schmidhauser, R., Gysel, M., Weingartner, E., Riccobono, F., Santos, S., Grüning, C., Faloon, K., Beddows, D., Harrison, R., Monahan, C., Jennings, S. G., O'Dowd, C. D., Marinoni, A., Horn, H.-G., Keck, L., Jiang, J., Scheckman, J., McMurry, P. H., Deng, Z., Zhao, C. S., Moerman, M., Henzing, B., de Leeuw, G., Löschau, G., and Bastian, S.: Mobility particle size spectrometers: harmonization of technical standards and data structure to facilitate high quality long-term observations of atmospheric particle number size distributions, Atmos. Meas. Tech., 5, 657–685, doi:10.5194/amt-5-657-2012, 2012.

Zelenyuk, A., Cai, Y., and Imre, D.: From agglomerates of spheres to irregularly shaped particles: Determination of dynamic shape factors from measurements of mobility and vacuum aerodynamic diameters, Aerosol Sci. Tech., 40, 197–217, doi:10.1080/02786820500529406, 2006.

Zieger, P., Weingartner, E., Henzing, J., Moerman, M., de Leeuw, G., Mikkilä, J., Ehn, M., Petäjä, T., Clémer, K., van Roozendael, M., Yilmaz, S., Frieß, U., Irie, H., Wagner, T., Shaiganfar, R., Beirle, S., Apituley, A., Wilson, K., and Baltensperger, U.: Comparison of ambient aerosol extinction coefficients obtained from in-situ, MAX-DOAS and LIDAR measure-

ments at Cabauw, Atmos. Chem. Phys., 11, 2603-2624, doi:10.5194/acp-11-2603-2011, 2011.

Zieger, P., Fierz-Schmidhauser, R., Weingartner, E., and Baltensperger, U.: Effects of relative humidity on aerosol light scattering: results from different European sites, Atmos. Chem. Phys., 13, 10609–10631, doi:10.5194/acp-13-10609-2013, 2013.

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**Table 1.**  $\kappa$  values derived from different measurements at the CESAR site in Cabauw in this and previous studies, plus simulated global values.

compound	$\kappa_{\text{low}}$	$\kappa_{\sf up}$	$\kappa_{ m median}$	instrument	dry diameter [nm]
Cabauw (ambient) Cabauw (ambient)	0.21 0.14 <sup>a</sup>	0.93 0.53 <sup>a</sup>	0.33 0.29 <sup>a</sup>	WHOPS (this study) Wet-Neph	500 mean of size distribution
Cabauw (ambient) global continental global remote marine global mixed marine/continental	0.11 <sup>b</sup> 0.06 <sup>c</sup> 0.90 <sup>c</sup> 0.40 <sup>c</sup>	0.44 <sup>b</sup> 0.48 <sup>c</sup> 1.00 <sup>c</sup> 0.60 <sup>c</sup>	0.26 <sup>b</sup>	HTDMA EMAC Model EMAC Model EMAC Model	165 - - -
	K <sub>mean</sub>	$\kappa_{std}$			
continental (ambient) marine (ambient)	0.3 <sup>d</sup> 0.7 <sup>d</sup>	0.1 <sup>d</sup> 0.2 <sup>d</sup>		Review – various methods Review – various methods	-

<sup>&</sup>lt;sup>a</sup> Zieger et al. (2011),  $\kappa$  calculated from GF retrieved from Wet-Neph data, from their Fig. 6.

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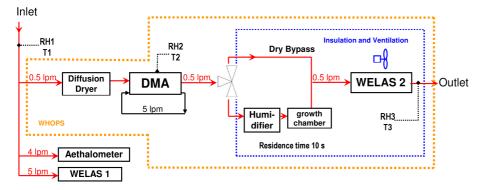
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 $<sup>^{\</sup>rm b}$  Zieger et al. (2011),  $\kappa$  calculated from HTDMA data from their Fig. 6.

<sup>&</sup>lt;sup>c</sup> Pringle et al. (2010), range of *κ* from Table 1 simulated at Earth's surface level from ECHAM/MESSy Atmospheric Chemistry (EMAC) model to simulate global fields of the effective hygroscopicity parameter *κ* which approximately describes the influence of chemical composition on the cloud condensation nucleus (CCN) activity of aerosol particles.

d Andreae and Rosenfeld (2008), summary for κ values from Sect. 2.1.3 from several HTDMA and CCNC (Cloud Condensation Nuclei Counter) measurements.



**Figure 1.** Setup of the PSI-Zeppelin Rack containing the WHOPS along with an Aethalometer and the WELAS1 sharing the sampling line in configuration for the airborne measurements. The positions of the RH and *T* sensors are shown (RH1, T1, etc).

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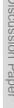




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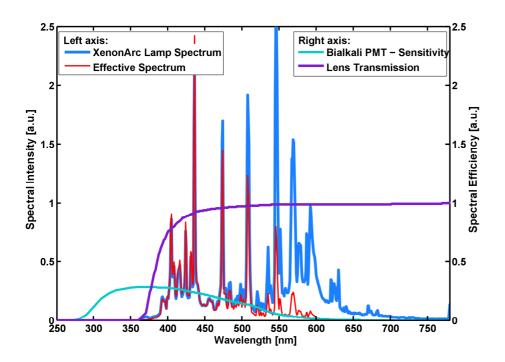


Figure 2. The effective light spectrum (red line; a.u.), which is relevant for scattering cross section calculations, is obtained by multiplication of the xenon arc lamp light spectrum (blue line; a.u.) with the wavelength dependent lens transmission (WELAS light collection optics; purple line: absolute values) and bialkali PMT sensitivity (turquoise line; a.u.) and an arbitrary constant scaling factor (chosen to get equal values of the red and blue lines at 450 nm wavelength).

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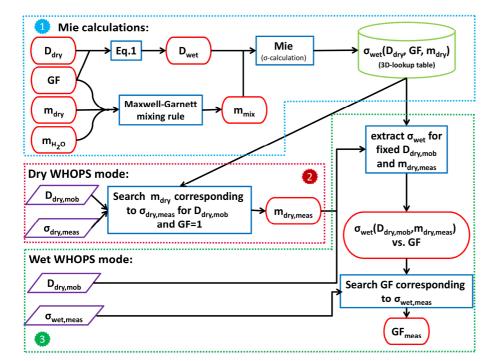
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**Figure 3.** Schematic of the approach applied to analyze the WHOPS measurements, indicated with the violet rhombi, which can be divided into three main steps: first (blue dashed box), a 3-D-lookup table containing  $\sigma_{\rm wet}$  as a function of  $D_{\rm dry}$ ,  $m_{\rm dry}$  and GF is prepared using Mie theory. Second (red dashed box), the results from the dry mode WHOPS measurements are used to infer the effective index of refraction ( $m_{\rm dry,\ meas}$ ) of the dry particles. Third (green dashed box), the scattering cross sections  $\sigma_{\rm wet,\ meas}$  of the humidified particles measured in the wet WHOPS mode are converted to corresponding growth factors (GF<sub>meas</sub>).

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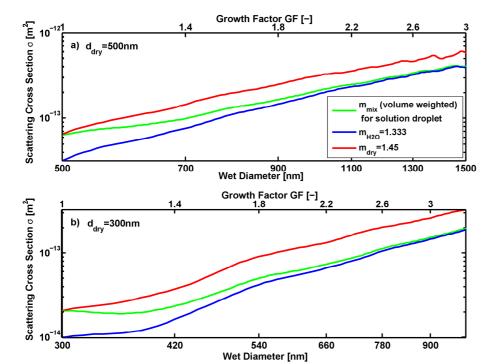


Figure 4. Calculated scattering cross section as a function of particle diameter (hygroscopic growth factor) for solution droplets (green lines) formed on particles with a dry diameter of (a) 500 nm and (b) 300 nm. A volume weighted mixing rule is used to determine the index of refraction of the solution droplet (see Sect. 2.2.2 and Fig. 3). The Mie curves for pure water droplets ( $m_{H_0O}$  = 1.333; blue lines) and particles with a constant index of refraction equal to that of the dry particle ( $m_{dry} = 1.45$ ; red lines) are also shown as a reference.

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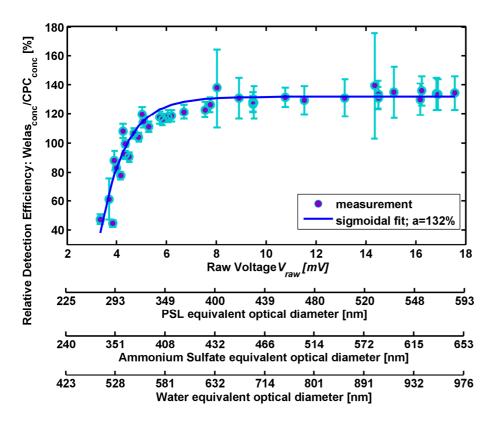
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**Figure 5.** Counting efficiency for size-selected ammonium sulfate particles retrieved by dividing the WELAS number concentration measurement by the corresponding value measured by the CPC. The primary abscissa is the measured raw pulse height ( $V_{raw}$ ). The additional abscissa scales representing the PSL equivalent, ammonium sulfate and water equivalent optical diameters are calculated from  $V_{raw}$  using the mean WELAS calibration value  $C_V = 2.23$ . The value a depicts the plateau value of the sigmoidal fit.

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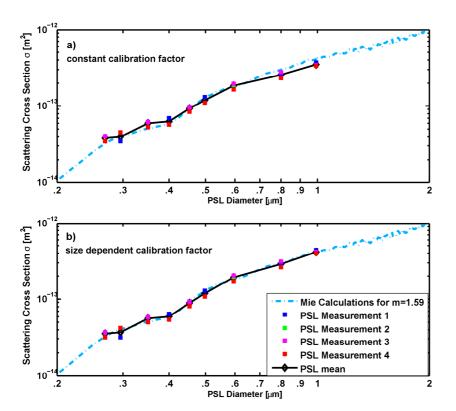






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**Figure 6.** Comparison of measured and theoretical Mie curve calculated for m=1.59. The square markers "PSL measurement 1–4" show results from 4 different experiments and the black line "PSL mean" shows the mean value over the 4 experiments. In **(a)** a constant scattering cross section calibration factor of  $C_{\sigma} = 3.71 \times 10^{-15} \,\mathrm{m^2\,mV^{-1}}$  was applied to analyze raw signals of the PSL measurements, while in **(b)** the voltage dependent calibration function  $C_{\sigma}(V_{\mathrm{cal}})$  described in Eq. (9) and Fig. 7 was applied.

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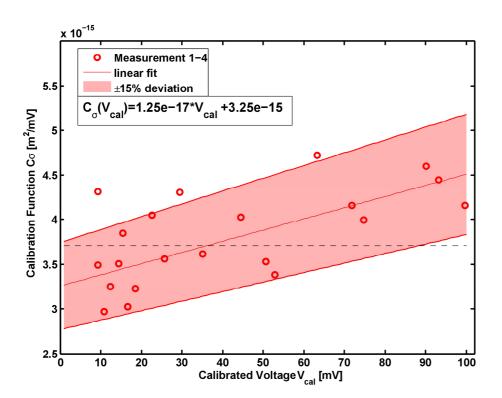
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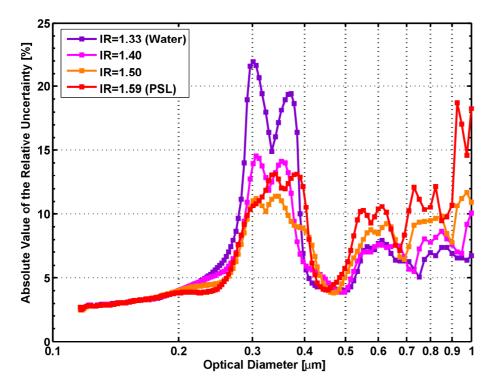
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**Figure 7.** Actual ratio ( $C_{\sigma}$ ) between the calculated scattering cross section of PSL spheres and the calibrated voltage  $(V_{cal})$  measured by the light scattering detector plotted against  $V_{cal}$ . A linear fit (red line) through the data points is used as scattering cross section calibration function,  $C_{\sigma}(V_{\rm cal})$ , for the WELAS data analysis. Additionally the shaded area represents  $\pm 15\%$ deviation around the fit. Approximately 85% of all data points fall within this uncertainty. The black dashed line depicts the constant calibration factor  $C_{\sigma} = 3.71 \times 10^{-15} \,\mathrm{m}^2 \,\mathrm{mV}^{-1}$  applied for Fig. 6a.



**Figure 8.** Absolute values of the relative uncertainty of the optical sizing by the WELAS as a function of optical diameter and for different indices of refraction. This uncertainty estimate is based on error propagation calculations for a fixed uncertainty of  $\pm 9\%$  for the measured scattering cross sections.

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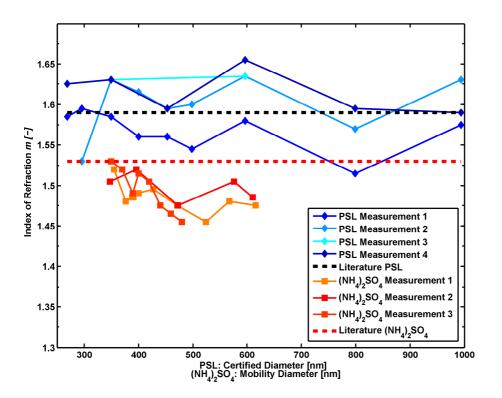
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**Figure 9.** Retrieved refractive indices for dry, nebulized ammonium sulfate ( $(NH_4)_2SO_4$ ; reddish colors; plotted against mobility diameter) and different polystyrene latex spheres (PSL; bluish colors; plotted against certified diameter) measurements.

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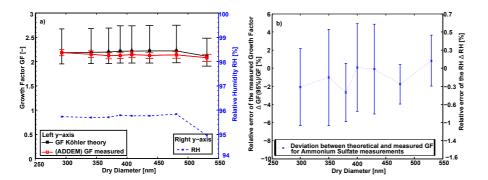




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**Figure 10. (a)** shows a typical example of the hygroscopic growth factors (red markers) measured by the WHOPS for pure ammonium sulfate particles of different dry size along with the relevant relative humidity (blue markers) simultaneously measured in the WHOPS. In **(a)** the error bars around the measured GFs indicate the uncertainty resulting from the optical and mobility sizing. The theoretical growth factors (ADDEM Model; Topping et al., 2005) corresponding to the measured RH are shown in black, where the error bars indicate the propagated uncertainty corresponding to an experimental uncertainty of the measured relative humidity of  $\sim \pm 2\%$  (absolute). The markers and error bars shown in **(b)** indicate the statistics (mean and standard deviation from 8 independent experiments) of the relative difference between measured and theoretical growth factor of ammonium sulfate particles at RHs between 89% and 97% for different dry sizes. The right y axis shows the hypothetical RH bias, which would explain the corresponding observed GF bias shown on the left y axis.

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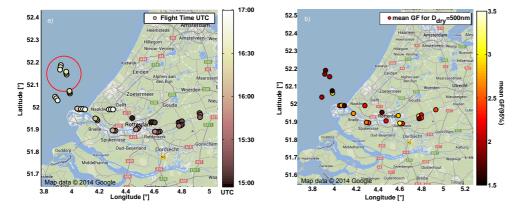






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**Figure 11.** Airborne WHOPS measurements onboard the Zeppelin NT from the 22 May 2012 in the Netherlands; **(a)** flight track color-coded by UTC time; **(b)** mean GFs (color-coded) at RH = 95% for  $D_{\text{dry}} = 500$  nm particles.

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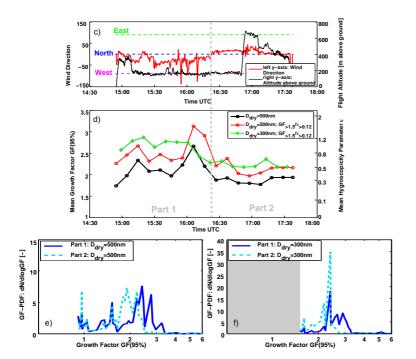




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**Figure 11.** Airborne WHOPS measurements onboard the Zeppelin NT from the 22 May 2012 in the Netherlands; **(c)** time series of the local wind direction measured on board of the Zeppelin NT; **(d)** time series of the measured mean GFs (left axis) and corresponding  $\kappa$  values (right axis) for the selected particle size of 500 nm; additionally the mean GF of all particles with GF > 1.5 ( $\kappa$  > 0.12; lower detection limit when selecting 300 nm particles) is shown for the 300 nm and 500 nm particles; the grey, dashed, vertical line across panels **(c)** and **(d)** marks the point in time when the local wind direction changed; the track before this wind direction change is referred to as "Part 1" whereas the remaining time is referred to as "Part 2"; **(e)** GF-PDF for humidified 500 nm particles separately averaged over "Part 1" and "Part 2" of the flight; **(f)** equivalent to **(e)** but showing results for 300 nm particles. The grey area covers the GFs that cannot be reliably detected for this particle size.

## **AMTD**

7, 7321–7366, 2014

A novel airborne system to characterize aerosol hygroscopicity

B. Rosati et al.

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