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The Orbiting Carbon Observatory (OCO-2): spectrometer performance evaluation using pre-launch direct sun measurements

C. Frankenberg¹, R. Pollock¹, R. A. M. Lee¹, R. Rosenberg¹, J.-F. Blavier¹, D. Crisp¹, C. W. O'Dell², G. B. Osterman¹, C. Roehl³, P. O. Wennberg³, and D. Wunch³

¹Jet Propulsion Laboratory, California Institute of Technology, Pasadena, USA

²Colorado State University, Fort Collins, USA

³California Institute of Technology, Pasadena, USA

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Correspondence to: C. Frankenberg (christian.frankenberg@jpl.nasa.gov)

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OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Abstract

The Orbiting Carbon Observatory-2 (OCO-2), launched on 2 July 2014, is a NASA mission designed to measure the column-averaged CO₂ dry air mole fraction, XCO₂. Towards that goal, it will collect spectra of reflected sun-light in narrow spectral ranges centered at 0.76, 1.6 and 2.0 μm with a resolving power ($\lambda/\Delta\lambda$) of 20 000. These spectra will be used in an optimal estimation framework to retrieve XCO₂. About 100 000 cloud free soundings of XCO₂ each day will allow estimates of net CO₂ fluxes on regional to continental scales to be determined. Here, we evaluate the OCO-2 spectrometer performance using pre-launch data acquired during instrument thermal vacuum tests in April 2012. A heliostat and a diffuser plate were used to feed direct sunlight into the OCO-2 instrument and spectra were recorded. These spectra were compared to those collected concurrently from a nearby high-resolution Fourier Transform Spectrometer that was part of the Total Carbon Column Observing Network (TCCON). Using the launch-ready OCO-2 calibration and spectroscopic parameters, we performed total column scaling fits to all spectral bands and compared these to TCCON results. On 20 April, we detected a CO₂ plume from the Los Angeles basin at the JPL site with strongly enhanced short-term variability on the order of 1 % (3–4 ppm). We also found good (<0.5 ppm) inter-footprint consistency in retrieved XCO₂. The variations in spectral fitting residuals are consistent with signal-to-noise estimates from instrument calibration, while average residuals are systematic and mostly attributable to remaining errors in our knowledge of the CO₂ and O₂ spectroscopic parameters. A few remaining inconsistencies observed during TVAC may be attributable to the specific instrument setup on the ground and will be re-evaluated with in-orbit data, when the instrument is expected to be in a much more stable environment.

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



1 Introduction

After the launch failure of OCO (Crisp et al., 2004) in 2009, the National Aeronautics and Space Administration (NASA) authorized the development of OCO-2, which was successfully launched on 2 July 2014 at 2.56 a.m. PDT from Vandenberg Air Force Base on the California coast. To achieve its mission goal, OCO-2 spectra must enable retrievals of column averaged atmospheric CO₂ (denoted as XCO₂) with an accuracy of 1 ppm or better. A typical algorithm for the retrieval of XCO₂ (Bösch et al., 2006; Butz et al., 2009; O'Dell et al., 2012) concurrently employs three spectral bands, centered around 0.76 μm (O₂ A-band), 1.61 μm (weak CO₂ band) and 2.06 μm (strong CO₂ band). By using this multi-channel approach, XCO₂, surface albedos, as well as aerosol properties can be retrieved concurrently.

This paper describes spectral fitting results from direct sun observations collected with the OCO-2 instrument during Thermal Vacuum Tests (TVAC) in April 2012 at the Jet Propulsion Laboratory (JPL) in Pasadena, California. We compare the instrument performance with measurements collected by a high resolution Fourier Transform Spectrometer from the Total Carbon Column Observing Network (TCCON) (Washenfelder et al., 2006; Wunch et al., 2011) that was located at the same altitude, ≈ 200 m from the TVAC facilities at JPL. By using actual retrievals from the OCO-2 data, we can evaluate retrieval-relevant instrument properties that will help guide users of the OCO-2 spectra. While direct sun retrievals are simpler than retrievals of reflected sunlight from space, these measurements from OCO-2 provide an end-to-end description of the combined impact of the OCO-2 instrument calibration as well as other aspects of the retrieval, such as the spectrally-dependent gas absorption cross-sections. Detailed descriptions of the instrument calibration can be found in Rosenberg et al. (2014); O'Dell et al. (2011) for radiometric calibration and Lee et al. (2014); Day et al. (2011) for instrument lineshape (ILS) and spectral characterization. This paper is organized as follows: Sect. 2 provides a succinct overview of the OCO-2 instrument. Section 3 discusses the retrievals based on Thermal Vacuum Test (TVAC) data acquired by the

AMTD

7, 7641–7670, 2014

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



OCO-2. Section 4 focusses on the fortuitous detection of a CO₂-rich plume associated with the Los Angeles urban dome during the TVAC tests and Sect. 5 summarizes the overall work.

2 OCO-2 instrument overview

The OCO-2 instrument is a 3-channel grating spectrometer. It records spectra of the O₂ A-band (0.757–0.775 μm, FWHM = 0.042 nm, denoted as O2A), a weak CO₂ band (1.594–1.627 μm, FWHM = 0.076 nm, denoted as WCO2) and a strong CO₂ band (2.043–2.087 μm, FWHM = 0.097 nm, denoted as SCO2) with 8 independent along-slit focal plane array readouts, denoted as footprints (1–8). A common telescope feeds all 3 spectrometers through a series of beam splitters and re-imagers, with a linear polarizer selecting only the polarization vector parallel to the entrance slit. Each spectrometer works in first order of the holographic grating. The system is optically fast (F/1.8) and yields a high signal-to-noise ratio (SNR). See Crisp et al. (2004); Pollock et al. (2010); Haring et al. (2008) for a more general description of the OCO instrument and mission.

At each spectrometer's focus, a 1024 × 1024 imaging array collects the spectrum, such that one dimension measures field angles along the slit, while the other dimension measures wavelengths. The O₂ A-band detector is a silicon (HyViSi) Hawaii-1RG, and the two CO₂ detectors are HgCdTe Hawaii-1RG; all were manufactured by Teledyne Scientific and Imaging, LLC. Only 160 out of the 1024 pixels in the spatial dimension are used, and sets of ≈ 20 are averaged onboard to constitute the eight spatial along-slit footprints. In the spectral dimension, four reference pixels are blacked out on each end of the array, leaving 1016 pixels (or channels) per band and footprint. Figure 2 shows an example of the focal plane array (1016 × 160, with block averages of 20 rows in the spatial dimension) readout of all three bands using a direct sun measurement, corresponding to an airmass equivalent to a typical in-space viewing geometry (solar zenith angle of 67°). For retrieval purposes, it is important to note that each

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



spectral band (3), footprint (8) and spectral pixel (1016) has its own characterization; this is in contrast to the FTS instrument on-board GOSAT (Hamazaki et al., 2005; Kuze et al., 2009). The instrument line-shape (ILS), for instance, is given for each of the $3 \times 8 \times 1016$ detector elements independently, resulting in a 4-dimensional array with the dimensions $3 \times 8 \times 1016 \times 200$ (the ILS is defined for 200 spectral points around the center point).

The OCO-2 instrument was designed to have a spectral sampling of approximately 2.5 detector elements per full width at half maximum (FWHM), in each band, and a spectral resolving power $\lambda/\Delta\lambda$ of approximately 20 000 in the CO₂ channels and 17 000 in the O₂ A channel. The shape of the OCO-2 ILS function is determined by the slit width, pixel pitch, optical aberrations, diffraction, and detector crosstalk. Details of the ILS as well as spectral and radiometric characterization that are used in this manuscript can be found in Rosenberg et al. (2014); Lee et al. (2014); Day et al. (2011); O'Dell et al. (2011). Figure 1 shows an example of the OCO-2 ILS for footprint 4 in each band at four different detector array spectral positions. The ILS variation in the spatial direction (footprint) is not as strong as in the spectral direction but still needs to be taken into account for accurate retrievals. As mentioned above, ILS functions are provided in tables that can be interpolated because conventional line-shape functions (e.g. Gaussian, Voigt) can't fit the shapes well enough for accurate XCO₂ retrievals.

3 Thermal Vacuum Test (TVAC) results

3.1 Experimental setup

During the characterization and calibration of the OCO-2 flight instrument at the Jet Propulsion Laboratory in Pasadena, CA, a heliostat was used to direct sunlight onto a diffuser that was viewed by the flight instrument through a window in the TVAC chamber. Simultaneously, a TCCON instrument located at the same altitude about 200 m away recorded high resolution solar spectra through essentially the same atmospheric

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

I◀

▶I

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



column. The FTS spectra were acquired at an un-apodized resolution of 0.013 cm^{-1} (45 cm optical path difference), which is approximately 20 times higher than the spectral resolution of the OCO-2 instrument. In addition, the FTS spectra had high SNR and were characterized by a single, well-determined ILS for the entire spectral range (Wunch et al., 2011). Apart from the spectral resolution, there are a few other differences between both measurements. First, the integration time of the OCO-2 instrument was 0.333 s (a continuous 3 Hz measurement) while the FTS required 79 s to complete a scan. Second, the OCO-2 instrument had a larger field of view (FOV) than the FTS, observing the full solar disk and some of the surrounding sky, while the FTS observed the center of the solar disk. The solar lines observed by OCO-2 were therefore broadened relative to those observed by the FTS due to Doppler shifts caused by the sun's rotation. For this reason, solar line pixels in the FTS fits were excluded in the following analysis, as we aimed at using exactly the same retrieval setup as for the OCO-2 fits.

3.2 Spectral fits

For the spectral fits, we performed total column retrievals of trace gases for each band (O₂A, WCO₂, SCO₂) independently. For this purpose, we used the fast IMAP-DOAS preprocessor in a special up-looking retrieval mode (Frankenberg et al., 2005) using a single temperature and pressure a priori profile for the entire day, that was extracted from the National Centers for Environmental Prediction (NCEP) data, but replacing the surface pressure with the one locally measured at JPL. In this up-looking mode, the IMAP-DOAS algorithm merely fits spectral dispersion, solar shift and continuum baseline as a 3rd order polynomial, as well as a total column scaling of a pre-defined O₂, H₂O and CO₂ profile shape. Neither surface pressure nor a temperature scaling was retrieved.

In the following, we will show spectral fits for a low (1.1), typical (2.5) and high (5.1) air mass factor (AMF) in blue, red and black, respectively. The “typical” AMF will be very close to the lower end of AMFs encountered during flight, as the AMF in perfect

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



likely caused by the choice of the a priori temperature and pressure profile. Given that both instruments show very similar residuals, we conclude that they are still dominated by errors in the ABSCO tables of the O₂ A-band absorption cross sections, specifically relating to line mixing, line-shape and collision induced absorption. This can also be seen in the air mass dependence of systematic residuals. At higher air masses, the far line-wing shape becomes more important as line centers are entirely saturated. The periodic structures in the P-branch residuals point to the importance of the line-shape as residuals are smaller at lower air mass, at which the far wing line-shape becomes less important.

Weak CO₂ band (WCO₂) fits are shown in Fig. 4. At OCO-2 spectral resolution this band is not saturated even at a high airmasses. Most of the information content is thus in the line centers as measured by OCO-2, and the continuum can be isolated much easier from well separated CO₂ lines. Again, we observe residuals that strongly resemble those from TCCON fits (with many solar features masked), pointing to small remaining errors in CO₂ absorption cross sections. There is a strong similarity between footprints but, in general, most features are below 0.5 %, similar to measurements from the GOSAT satellite as well as TCCON. Some residual features, such as the slope at the short-wavelength edge of the fit window, can also be caused by broad spectral variations in the solar spectrum that are not accurately represented here.

Strong CO₂ band (SCO₂) fits are shown in Fig. 5. Line centers at OCO-2 resolution now gradually saturate at higher air mass and the continuum level radiances are harder to define than in the other bands because of overlapping line-wings. In general, conclusions for this band are similar to the other bands, with spectral residuals roughly resembling the TCCON residuals. The only exception may be at the short-wavelength edge where OCO-2 residuals are somewhat higher, potentially caused by either dispersion or ILS changes at the band edge, where the instrument calibration is less well characterized.

In summary, spectral fits to OCO-2 data yield expected fitting results but, as with all other current instruments measuring this spectral range, systematic residuals still

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



appear. Given the high SNR of the OCO-2 instrument, this would result in unrealistically high χ^2 values for spectral fits, which is why we will implement a residual fitting technique based on empirical orthogonal functions derived from systematic features as is currently done for data from GOSAT. Using these basis functions, as additional fit parameters, we expect fit residuals to be mostly noise driven resulting in realistic χ^2 values that will help the fit converge, allow us to use χ^2 as quality criteria, and result in more realistic a posteriori uncertainty estimates, as the true OCO-2 measurement noise can be provided to the weighted least squares fit.

To determine whether the variability in spectral residuals can be explained by OCO-2 detector noise, we computed the standard deviation of spectral residuals using 2000 OCO-2 fits, to spectra collected within a short time period, and define this as a surrogate for empirically derived noise in the OCO-2 data. Using the mean of the 2000 radiance spectra as signal level surrogate, we then derive a set of SNR points at varying signal levels within a spectrum, mostly caused by CO₂ and O₂ absorptions. Figure 6 shows the empirically derived SNR points as well as a typical SNR relationship based on the launch-ready OCO-2 calibration (Rosenberg et al., 2014). There is an excellent agreement between calibration curves and empirically derived SNR based on spectral fits. The spread of the individual data points for each band is due to the behavior of the detector, with varying gain coefficients and shot noise in each pixel. This is reflected in the OCO-2 launch-ready calibration curves as well but, for clarity, the theoretical SNR curve of just a typical pixel is shown in transparent lines, and corresponds very well with the empirical noise estimates. After accounting for the systematic residual features, expected χ^2 distributions of the residuals are thus attained.

Two advantages of the OCO-2 grating spectrometer are its high dynamic range and relatively low noise at low signal levels. This results in much lower noise levels within deep absorption lines of the O2A and the SCO2 bands than would be possible with space-based Fourier Transform Spectrometer data, which exhibit a constant absolute noise level across the entire spectrum (which would yield a straight line through the

origin in the SNR vs. signal level curve when plotting individual spectral points within a single acquisition).

3.3 Footprint dependencies

The OCO-2 spectrometer records spectra separately in the spatial domain of the focal plane array. To increase SNR, photons collected in 20 adjacent spatial rows are averaged to define a footprint, hence the 160 spatial detector rows produce 8 spatial footprints (see Fig. 2). Given that each footprint spans a considerable fraction of the focal plane array, instrument calibration (most importantly dispersion and ILS) vary from footprint to footprint. With the unprecedented accuracy requirements in the sub 1% range for atmospheric CO₂ measurements, inter-footprint difference in retrieved quantities for imaging spectrometers will cause challenges in using this data.

In Fig. 7, we show differences between the total column CO₂ estimates from each footprint and the all-footprint average, for the weak and strong CO₂ bands. Two main features are apparent: (1) there is a very small inter-footprint variation but at least 6 to 7 of the footprints agree to within ± 0.5 ppm; (2) the inter-footprint variations are somewhat unstable and appear to be influenced by how the heliostat is illuminating the OCO-2 slit, as can be seen by some jumps around 16:00 UTC, when the heliostat was realigned. Hence, it is not yet clear, what fraction of the variation is merely caused by the experimental setup and what is caused by potential small calibration inconsistencies. In orbit, the inter-footprint differences are not expected to vary in time and the variability of ± 0.5 ppm can be characterized using in-orbit data and corrected for using constant correction factors.

For the pre-launch OCO-2 calibration, two sets of instrument line-shapes were tested (Lee et al., 2014), of which one will eventually be implemented for official OCO-2 retrievals. One set is derived from tunable diode laser measurements and a second, enhanced set (see Lee et al., 2014), using direct comparisons against the high-resolution FTS spectra, where the ILS was optimized to best match the FTS data. In Fig. 8, the inter-footprint variations averaged over a longer time-period are shown independently

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



for the FTS enhanced ILS (solid lines) and the original laser-based ILS (dashed lines). The impact is not large and mostly affects footprint #1, at the edge of the field of view. For this particular footprint, both CO₂ bands would behave somewhat abnormally but deviate in opposing directions (positive or negative). This could have implications for a joint three-band retrieval. The enhanced ILS largely mitigates this problem to the best possible degree. Once OCO-2 is collecting in-orbit data, both ILS models will be evaluated in order to decide whether the FTS-enhanced ILS only mitigated an issue associated with the TVAC test setup or a real calibration issue. In any case, most footprints agree very well with each other for both choices of ILS and provide confidence, especially given the small fitting residuals discussed earlier. However, even this small footprint dependence will have to be scrutinized since the accuracy requirements of better than 1 ppm are very stringent. This also has implications for future imaging spectrometers with hundreds of spatial pixels, for which a careful per-footprint calibration as for OCO-2 might not be feasible.

3.4 The Matador test

The OCO-2 retrieval strategy imposes stringent requirements on knowledge about the detector linearity and dark current. Any imperfections can slightly modify the fractional depths of absorption lines depending on the detector fill level, which can create regional and/or seasonal biases in retrieved XCO₂. For OCO, the so-called “Matador test” was devised to detect any apparent non-linearities or dark offsets (O’Dell et al., 2011).

The general TVAC setup was used with direct sunlight being illuminating a diffuser viewed by the OCO-2 spectrometers via a series of mirrors. For the Matador test, special sheets with varying transmission were optionally inserted into the optical path. To create a spectrally uniform reduction in the solar intensity, aluminum sheets with small holes were fabricated with different hole-densities in a hexagonally packed pattern to create effective transmissions of approximately 75, 50, 25 and 10%. For the test, the sheets were rapidly (and manually) inserted into and removed from the heliostat optical

path, hence the name “Matador test” due to the visual analogy to a matador’s cape. The test was performed near local noon on 20 April 2012.

Figure 9 shows results for both CO₂ bands and the O₂ A-band independently, for each footprint, expressed as deviations in XCO₂ or O₂ column from the average across all footprints at the initial 100 % transmission level (i.e. before the sheets were inserted). The periods with reduced transmission are indicated with a gray background, as each sheet, in descending order of transmission, was inserted and removed three times.

For the 75, 50 and 25 % transmission levels, changes on the order of much less than 0.5 ppm occur. More importantly, deviations are neither consistent across footprints nor are they consistent across all three independent Matador tests per transmission level. This behavior hints at subtle changes in how the OCO-2 spectrometer slits were illuminated for each of these tests, which can also cause minor variations in the ILS and, hence, XCO₂. This is even more obvious at the <10 % transmission levels in which jumps in retrieved CO₂ are considerably higher but without any consistent pattern across all three independent tests.

For O₂ columns, all footprints agree to about 0.050 %, which would translate to better than 0.5 hPa or 0.2 ppm if errors propagate linearly into XCO₂. In general, the O₂ A-band appeared to be somewhat more consistent than the CO₂ bands and is excluded in the other plots, for the sake of clarity. It is, however, not clear whether calibration parameters are the cause for the very small variations in the O₂ fits or whether it can be merely attributed to the shape of the oxygen absorptions itself, where a lot of information is also coming from the deep Q-R-branch, which is less sensitive to ILS variations as no individual lines are sampled.

In summary, all of the 3 spectral bands show very good consistency, but achieving a near perfect consistency is currently not possible and small calibration factors for each footprint may need to be included for OCO-2. At this stage, we have gained confidence that the OCO-2 behaves well at variable illumination levels but also that we have obtained the maximum amount of information possible from the TVAC setup.

AMTD

7, 7641–7670, 2014

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



For these pre-launch tests, it is hard to reach the stability in terms of instrument thermal control or illumination that will be achieved in orbit.

4 Observing the Los Angeles urban dome with the OCO-2 instrument during the TVAC tests

5 Apart from the primary mission objective to better understand regional-scale biospheric fluxes, the OCO-2 data can also help us to estimate carbon fluxes from megacities (Kort et al., 2012). Even though its narrow swath will only rarely cross these localized sources, it will have unprecedented spatial resolution and XCO₂ sensitivity along the orbit track. Also, the glint repeat cycles will have shifted orbit paths, increasing the chances of obtaining direct cross-sections of the urban dome at least once in the OCO-2 mission.

10 In the Los Angeles basin, ground-based studies observed elevated amounts of CO₂ and CH₄ from total column measurements (Wunch et al., 2009). For CO₂, column enhancements measured with a TCCON FTS in 2007/08 at JPL revealed diurnal changes in XCO₂ on the order of up to 4 ppm, caused by the Los Angeles urban dome approaching the foothills at JPL around mid-day.

15 Here, we report on the first urban plume observed from the OCO-2 even before launch. Figure 11 shows time-series for 20/21 April of CO₂ column retrievals using the WCO2 and SCO2 band of OCO-2 as well as the TCCON FTS retrievals located nearby (Wunch et al., 2011). For TCCON retrievals, we extracted the official XCO₂ values retrieved using the latest GGG2012 data version (https://tcon-wiki.caltech.edu/Network_Policy/Data_Use_Policy/Data_Description). These data are based on two weak CO₂ bands, one of which is identical with the OCO-2 WCO2 band. CO₂ columns are scaled by retrievals of dry air, obtained from fits to the oxygen band at 1.27 μm. The official TCCON data applies both an airmass factor correction and a scaling factor to bring the measured XCO₂ onto the NOAA standard in situ CO₂ networks.

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



For OCO-2, all 8 footprints have been averaged but no further smoothing has been applied. Owing to the strengths of absorption features in the SCO₂ band, retrieval scatter in the strong band is substantially lower than for the weak band. TCCON data was divided by 1.005, indicating a small but consistent offset between the official TCCON results and OCO-2. The XCO₂ observed on these two TVAC days exhibits substantially different behavior. On 21 April, a relatively smooth diurnal cycle, with maxima around solar noon, were observed with an overall amplitude of about 2 ppm in XCO₂. Some short term features, such as the increase at 16:00 UTC or the swings around 18:30 and 19:30 UTC can be clearly observed in both TCCON and OCO-2, while the latter allows for a much faster sampling in time due to the 3 Hz readout rate. This also allows us to sample the very erratic behavior seen on 20 April. This day, interrupted by Matador tests around local noon, shows very high-frequency variations of observed CO₂ starting around 20:00 UTC. Figure 12 shows a zoom plot into this specific anomalous time-period with high-frequency variability unresolved by the TCCON sampling, but similar general features observed by both instruments. This behavior suggests the edge of the urban dome moving in and out of the observation site. Swings can be up 3–4 ppm within a 5–10 min period, as is observed at about 21:40 UTC. Assuming typical wind-speeds of 5–10 m s⁻¹, these temporal scales correspond to spatial scales of 1.5–3 km, which will be matched by the in-orbit OCO-2. In other words, spatial gradients of about 2–4 ppm within a few kilometers can be expected in strongly emitting areas such as Los Angeles. In addition, cross sections of the urban dome can be mapped with the OCO-2 with the caveat of a more complex retrieval from space, where retrievals are based on backscattered radiation and have to take aerosols into account. However, 1 % changes in XCO₂ as observed here are substantial and should be differentiable from light-path modifications within the urban dome.

Another feature is apparent in the time-series: the amplitude of the variation in the weak band is smaller than in the strong band. This could be caused by differences in the averaging kernels (AK) of both channels in total column scaling mode. The AK of the strong band, with more saturation, typically peaks at much higher values near

the surface if a simple total column is applied. Thus, CO₂ variability at the surface will be more amplified in the SCO₂ band retrievals, which is well reflected in the OCO-2 measurements and is another hint at boundary-layer CO₂ enhancements.

5 Conclusions

In this work, we performed spectral fits on calibrated OCO-2 data obtained during Thermal Vacuum Tests at the Jet Propulsion Laboratory in Pasadena, California, in April 2012. Direct sun-light was fed to the OCO-2 instrument via a heliostat and data could be compared against a high quality reference taken from a Fourier Transform Spectrometer located in close proximity. OCO-2 spectra were calibrated with the launch-ready calibration version from the OCO-2 instrument team. We performed the first CO₂ and O₂ total column retrievals from the OCO-2, providing a high level evaluation and reference for the expected OCO-2 instrument performance. We find that spectral residuals are still mostly dominated by systematic features but that these are also apparent in other measurements, including high-resolution FTS data. Hence, most of the systematic features can be attributed to uncertainties in the spectroscopy. We also find that OCO-2 footprints agree very well with each other, namely better than 1 ppm. There is a small time-dependence on the footprint dependency, which is probably related to the specific test setup and would thus not occur in space.

A strong CO₂ plume from the Los Angeles metropolitan area (3–4 ppm enhancement in the total column average) was observed on 20 April 2012, when the plume caused very high-frequency (5–10 min) swings of CO₂ at the JPL site. These are indicative of strong spatial gradients in XCO₂, which will be very helpful in characterizing localized carbon emissions when OCO-2 orbit tracks passes over such regions.

In summary, the OCO-2 instrument performed very well on the ground and none of the spectral fits calculated in this work showed anything unexpected. The team is now eagerly waiting for in-space OCO-2 data.

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



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AMTD

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OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



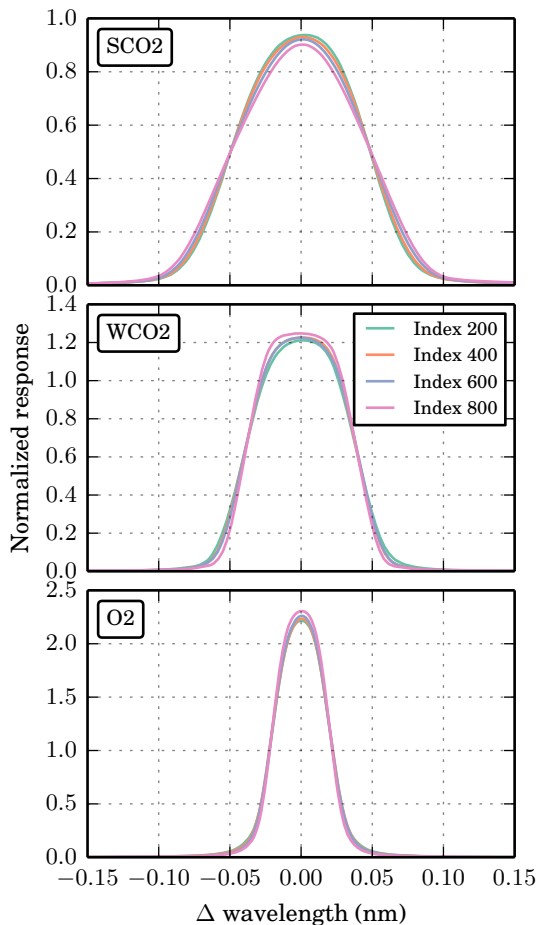


Figure 1. OCO-2 instrument line-shape functions (ILS) for a center footprint (#4) at 4 different spectral positions for each band independently. Full width at Half Maximum (FWHM) are about 0.04, 0.075 and 0.01 nm for the O₂, WCO₂ and SCO₂ band, respectively.

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



OCO-2 pre-flight performance

C. Frankenberg et al.

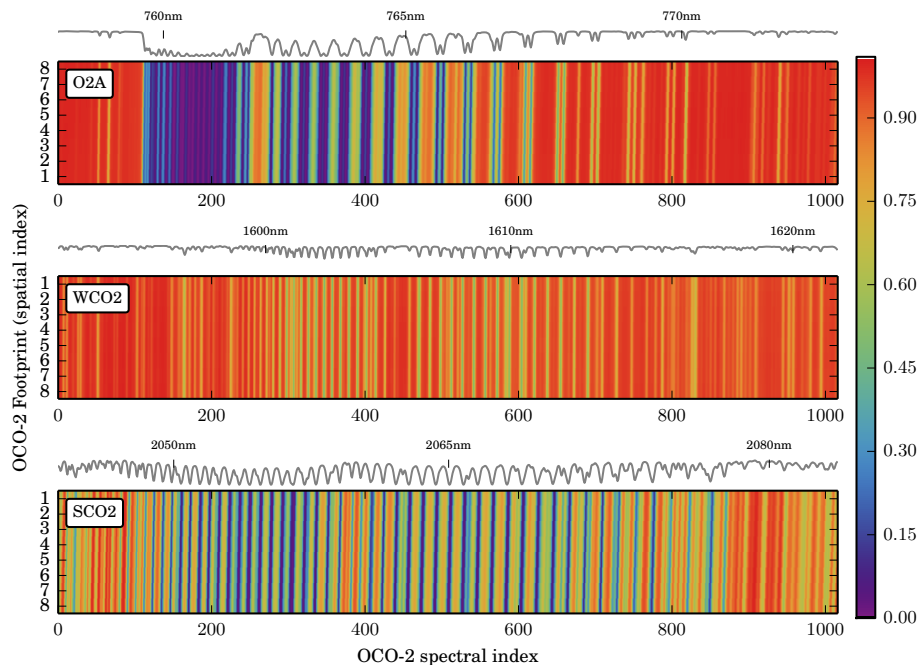


Figure 2. OCO-2 transmission spectra from TVAC direct-sun measurements at an airmass factor of ≈ 2.5 . Transmission levels from 0–1 are color-coded for the O₂ A-band (O2A), the weak (WCO2) and the strong (SCO2) CO₂ band, from top to bottom. The x axis represents the spectra dimension (1016 elements) of the OCO-2 focal plane and the y axis the spatial (160 elements co-added to 8 individual footprints). As a guide for the eye, a spectral cross-section of the center footprint is shown on top of each image, including wavelength indicators at the top. The detectors are slightly tilted with respect to the slit orientation for the O₂ A-band in particular, but also for the strong CO₂ band, causing a stronger footprint dependence of the dispersion coefficients.

OCO-2 pre-flight
performance

C. Frankenberg et al.

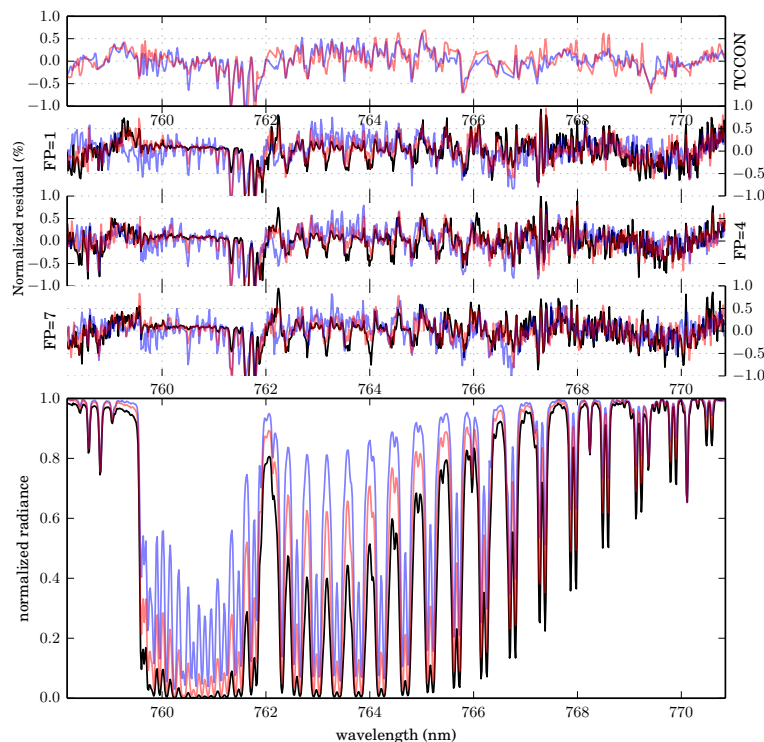


Figure 3. Spectral fits to direct-sun measurements in the O₂A band of OCO-2 and TCCON. Normalized radiances from OCO-2 are shown in the bottom panel for three different air masses (low (1.1), typical (2.5) and high (5.1) in blue, red and black, respectively). The middle panels show residuals of fits to normalized spectra, for three OCO-2 footprints, and the upper panel shows fit residuals to a high resolution FTS spectrum of identical airmass, that was convolved and resampled to OCO-2 spectral resolution prior to fitting.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

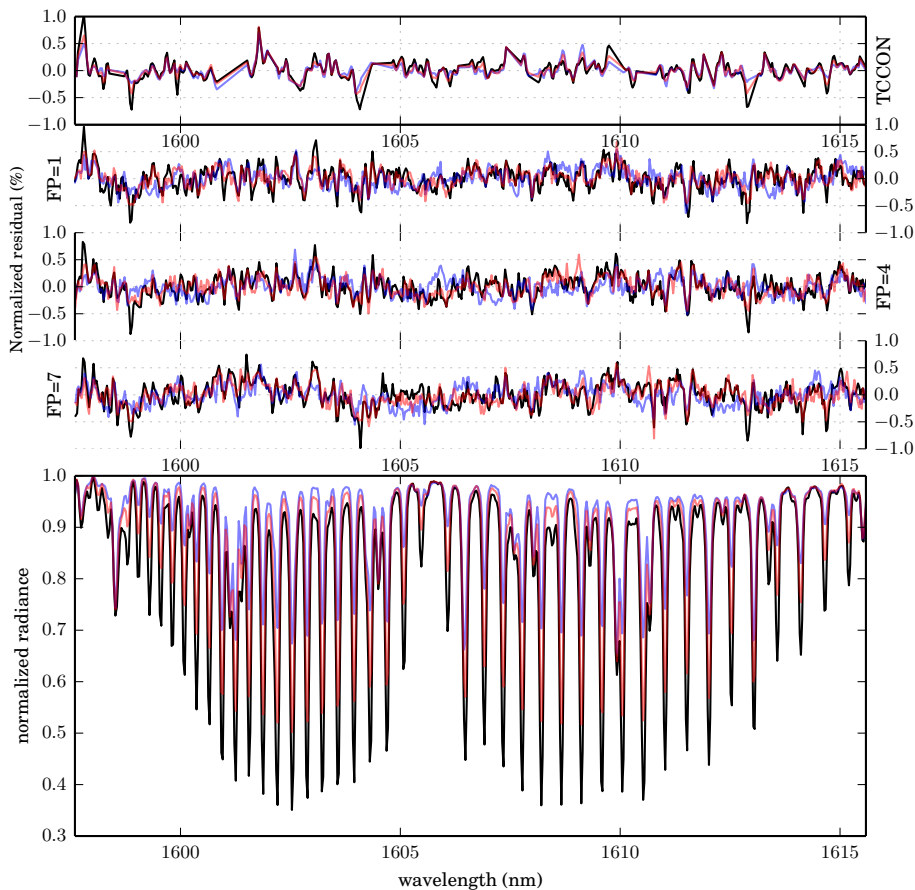


Figure 4. Spectral fits to direct-sun measurements in the WCO₂ band of OCO-2 and TCCON (analogous to Fig. 3).

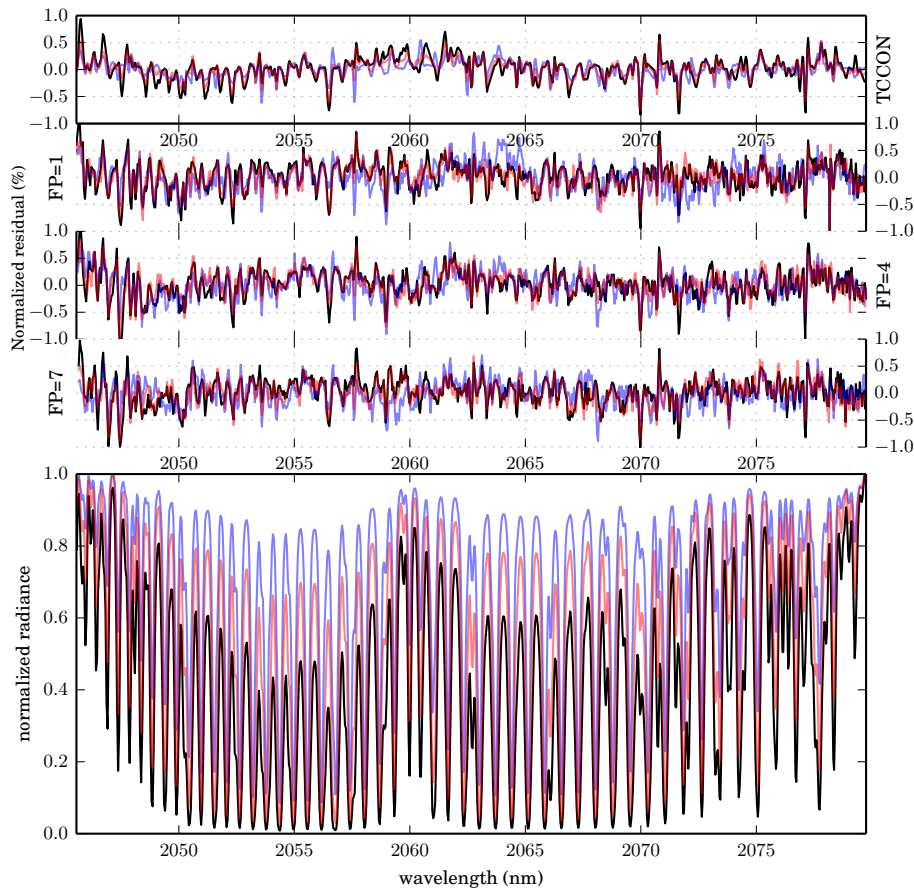


Figure 5. Spectral fits to direct-sun measurements in the SCO₂ band of OCO-2 and TCCON (analogous to Fig. 3).

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



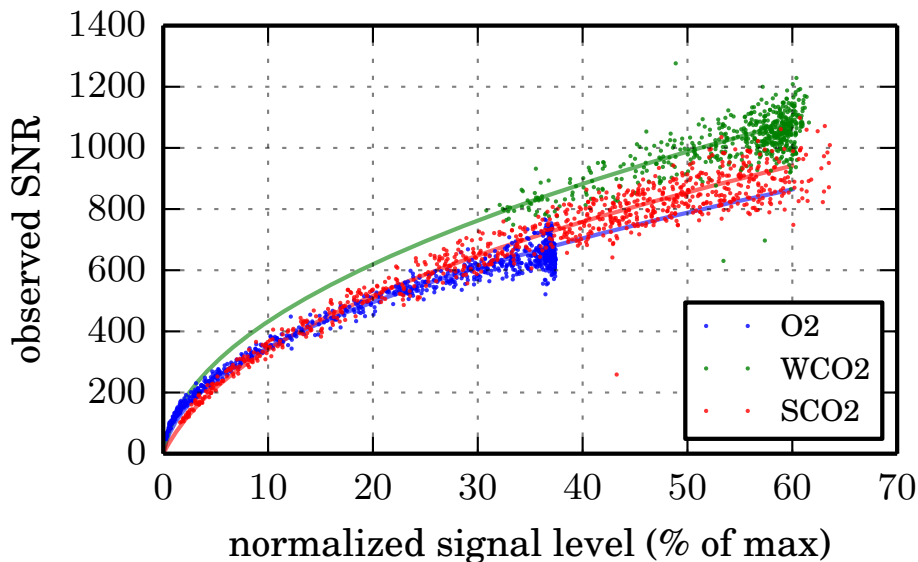


Figure 6. Empirical signal-to-noise ratios of each OCO-2 band derived from the standard deviation of spectral residuals from 2000 separate fits. Solid transparent lines represent the expected OCO-2 SNR for a typical detector pixel, which agrees well with the observed SNR. For each band (O2A, WCO2, SCO2), the max observable signal is defined as 7.0×10^{20} , 2.45×10^{20} and 1.25×10^{20} photons $\text{s}^{-1} \text{m}^{-2} \text{sr}^{-1} \mu\text{m}^{-1}$, respectively, where it is taken into account that OCO-2 will only measure one polarization direction (i.e. the maximum signal level in terms of full intensity would be twice as high as provided here).

OCO-2 pre-flight performance

C. Frankenberg et al.

Title Page	
Abstract	Introduction
Conclusions	References
Tables	Figures
◀	▶
◀	▶
Back	Close
Full Screen / Esc	
Printer-friendly Version	
Interactive Discussion	



OCO-2 pre-flight
performance

C. Frankenberg et al.

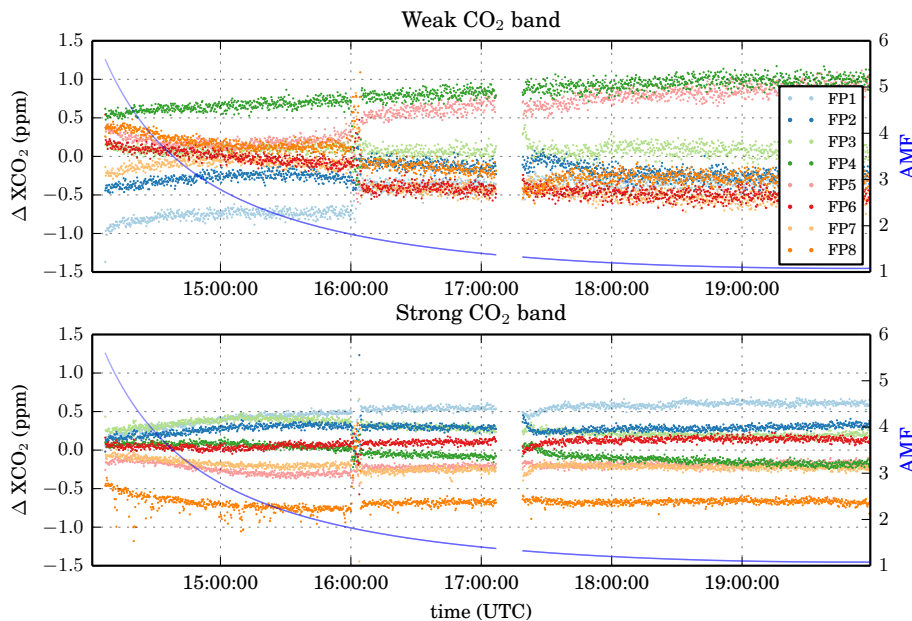


Figure 7. Footprint differences in retrieved CO₂ columns using the strong and weak band from OCO-2 TVAC data acquired on 21 April 2012.

OCO-2 pre-flight performance

C. Frankenberg et al.

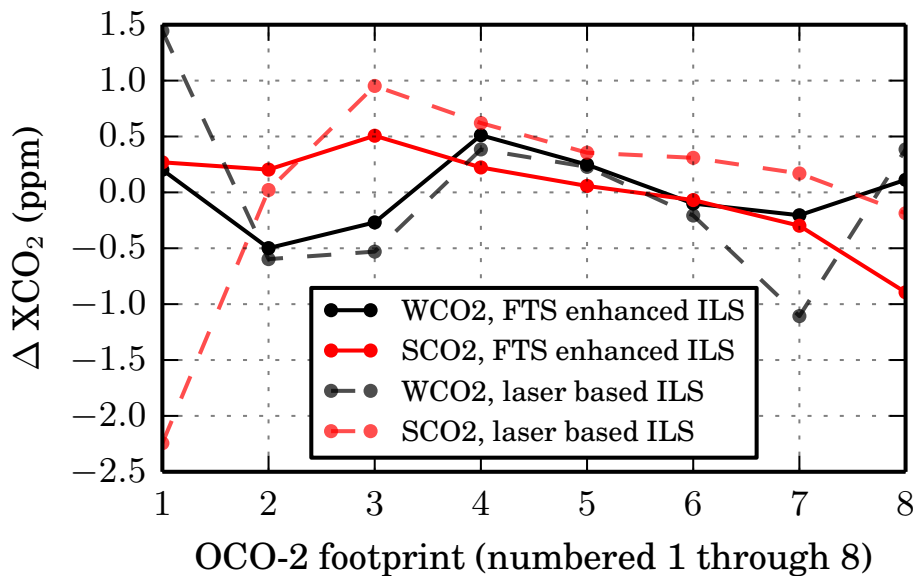


Figure 8. Average of inter-footprint differences in retrieved CO₂ columns from OCO-2 using two different sets of instrument line-shapes, namely the laser based ILS and the FTS enhanced ILS estimate.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

OCO-2 pre-flight performance

C. Frankenberg et al.

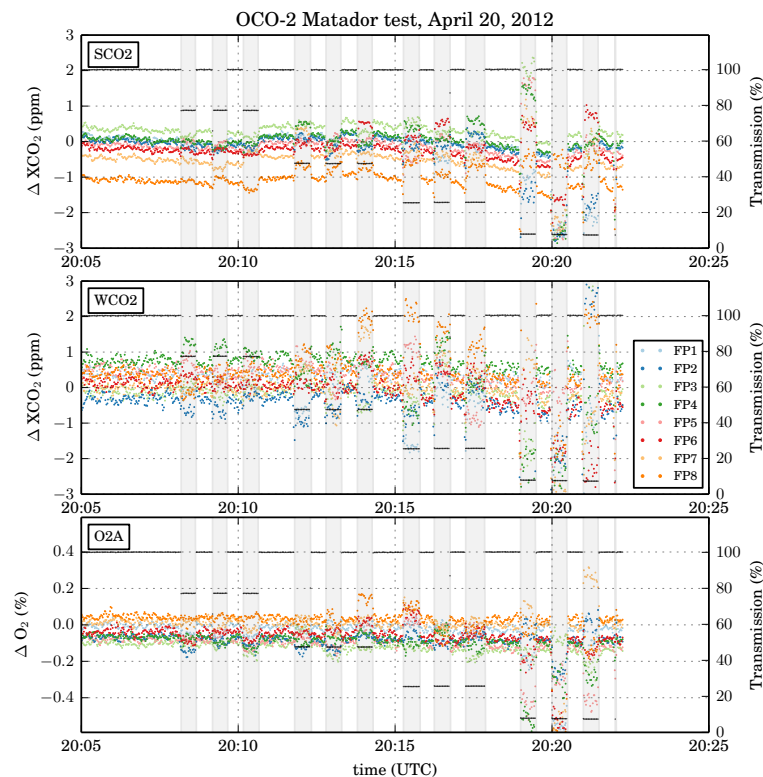


Figure 9. XCO₂ (in ppm) and O₂ (in %) differences from a common average in the OCO-2 Matador test, which changed signal levels through inserting aluminum sheets with variable hole density into the light beam. Gray shades indicate the time periods that the aluminum sheets were in the beam and the black horizontal lines indicate the transmission level for each of these tests. Each footprint for both CO₂ bands and the O₂ band is displayed separately.

OCO-2 pre-flight performance

C. Frankenberg et al.

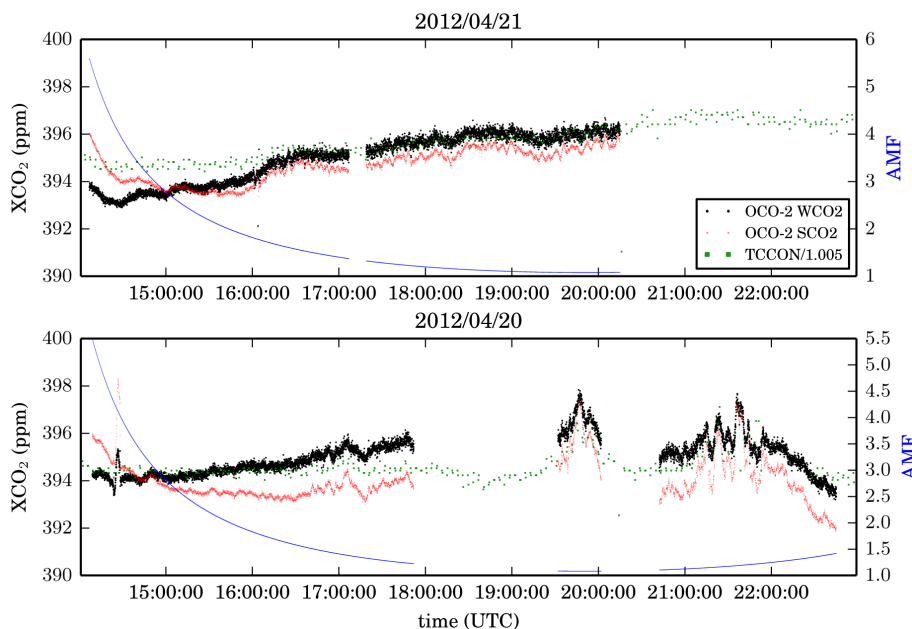


Figure 11. XCO₂ time-series on 20/21 April 2012 acquired at JPL near Pasadena, CA. Air mass is depicted in blue, SCO₂ and WCO₂ results in red and black, respectively, and scaled official TCCON data in green. Only every 10th OCO-2 data point is plotted but all 8 footprints are averaged.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**OCO-2 pre-flight
performance**

C. Frankenberg et al.

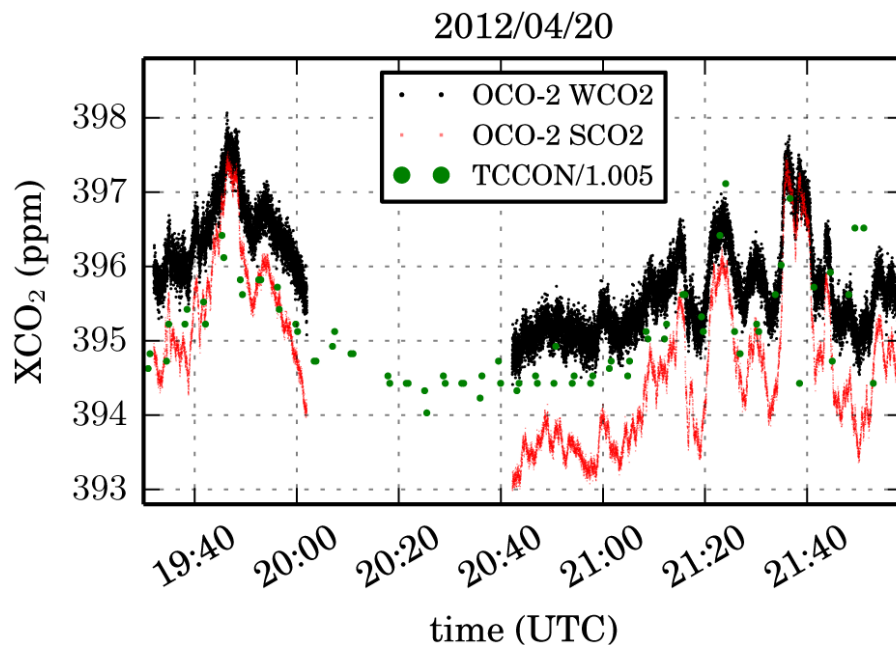


Figure 12. Zoom on the CO₂ column time-series on 20 April 2012 acquired at JPL near Pasadena, CA. Every OCO-2 data point is plotted and all 8 footprints are averaged.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)