



Remote sensing of  
CO<sub>2</sub>, CH<sub>4</sub>, CO, and  
H<sub>2</sub>O from  
geostationary orbit

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# Simulated retrievals for the remote sensing of CO<sub>2</sub>, CH<sub>4</sub>, CO, and H<sub>2</sub>O from geostationary orbit

X. Xi<sup>1</sup>, V. Natraj<sup>2</sup>, R. L. Shia<sup>1</sup>, M. Luo<sup>2</sup>, Q. Zhang<sup>1</sup>, S. Newman<sup>1</sup>, S. P. Sander<sup>2</sup>, and Y. L. Yung<sup>1,2</sup>

<sup>1</sup>Environmental Science and Engineering, California Institute of Technology, Pasadena, California, USA

<sup>2</sup>Jet Propulsion Laboratory, California Institute of Technology, Pasadena, California, USA

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Correspondence to: X. Xi (xixi@caltech.edu)

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## Abstract

The Geostationary Fourier Transform Spectrometer (GeoFTS) is designed to measure high-resolution spectra of reflected sunlight in three near-infrared bands centered around 0.76, 1.6, and 2.3  $\mu\text{m}$  and to deliver simultaneous retrievals of column-averaged dry air mole fractions of  $\text{CO}_2$ ,  $\text{CH}_4$ ,  $\text{CO}$ , and  $\text{H}_2\text{O}$  (denoted  $\text{XCO}_2$ ,  $\text{XCH}_4$ ,  $\text{XCO}$ , and  $\text{XH}_2\text{O}$ , respectively) at different times of day over North America. In this study, we perform radiative transfer simulations over both clear-sky and all-sky scenes expected to be observed by GeoFTS and estimate the prospective performance of retrievals based on results from Bayesian error analysis and characterization.

We find that, for simulated clear-sky retrievals, the average retrieval errors and single-measurement precisions are  $<0.2\%$  for  $\text{XCO}_2$ ,  $\text{XCH}_4$ , and  $\text{XH}_2\text{O}$ , and  $<2\%$  for  $\text{XCO}$ , when the a priori values have a bias of 3% and an uncertainty of 3%. In addition, an increase in the amount of aerosols and ice clouds leads to a notable increase in the retrieval errors and slight worsening of the retrieval precisions. Furthermore, retrieval precision is a strong function of signal-to-noise ratio and spectral resolution. This simulation study can help guide decisions on the design of the GeoFTS observing system, which can result in cost-effective measurement strategies while achieving satisfactory levels of retrieval precisions. The simultaneous retrievals at different times of day will be important for more accurate estimation of carbon sources and sinks on fine spatiotemporal scales and for studies to better understand the close coupling between the carbon and water cycles.

## 1 Introduction

The global carbon cycle has been significantly perturbed since the start of industrialization. The atmospheric concentrations of the greenhouse gases (GHG) carbon dioxide ( $\text{CO}_2$ ), methane ( $\text{CH}_4$ ), and the pollutant carbon monoxide ( $\text{CO}$ ) have all increased substantially due to anthropogenic activities (IPCC, 2013). These potent greenhouse

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gases trap outgoing long-wave radiation and warm the Earth's atmosphere and surface. As surface temperature continues to rise, the global water cycle is also notably changed (IPCC, 2013). Being the most potent greenhouse gas, water vapor (H<sub>2</sub>O) is of crucial importance to climate studies. An accurate understanding of carbon storage in the atmosphere, oceans, and terrestrial biosphere, as well as the exchanges between these reservoirs and the coupling with the water cycle, is critical to confidently project the future evolution of climate and to support societal efforts to mitigate climate change. However, despite decades of research, the Earth's carbon budget, especially the changes in the carbon sinks, still remains uncertain on regional and global scales (Le Quéré et al., 2009, 2014). In particular, two major challenges remain in balancing the carbon budget: quantifying uncertainty in fossil fuel CO<sub>2</sub> emissions, and constraining the variability and uncertainty of land fluxes (Schimel et al., 2001; Houghton et al., 2007).

To tackle these challenges in carbon cycle studies, several satellite missions have been measuring the atmospheric concentrations of greenhouse gases such as CO<sub>2</sub> and CH<sub>4</sub>. For example, the Atmospheric Infrared Sounder (AIRS), whose main scientific goal is to improve weather forecasting, has been able to measure mid-tropospheric CO<sub>2</sub> and CH<sub>4</sub> since 2002 (Aumann et al., 2003; Chahine et al., 2005, 2008; Xiong et al., 2008). The SCanning Imaging Absorption spectroMeter for Atmospheric CartographY (SCIAMACHY) instrument was operational for measurements of atmospheric compositions from 2002 to 2012, with CO<sub>2</sub>, CH<sub>4</sub>, CO, and H<sub>2</sub>O as four of its target species (Buchwitz et al., 2004, 2007; Schrijver et al., 2009; Schneising et al., 2011). The Greenhouse gases Observing SATellite (GOSAT), developed by the Japan Aerospace Exploration Agency and launched in January 2009, is the world's first mission dedicated to the monitoring of CO<sub>2</sub> and CH<sub>4</sub> (Yokota et al., 2009; Butz et al., 2011; Yoshida et al., 2011). In addition, the Orbiting Carbon Observatory (OCO)-2 mission, successfully launched in July 2014, now provides high-quality and worldwide column abundances of CO<sub>2</sub> (Crisp et al., 2004, 2012).

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Flying in a Low Earth Orbit (LEO), these missions can achieve nearly global coverage in about sixteen days and offer monthly averaged retrievals of greenhouse gases. The sources and sinks of greenhouse gases are then inferred through flux inversions (e.g., Baker et al., 2010; Nassar et al., 2011). However, the low spatiotemporal measurement densities of the current Earth observing system in LEO result in a lack of information about emissions on smaller spatiotemporal scales (Chevallier et al., 2005; Hungershofer et al., 2010; Wecht et al., 2014). For example, localized emissions from forest fires and megacities vary over days or even hours. Furthermore, as the carbon and water cycles are closely coupled in the terrestrial biosphere, there is a significant scientific need to simultaneously study both on a daily basis.

The Geostationary Fourier Transform Spectrometer (GeoFTS) combines an imaging Fourier Transform Spectrometer instrument with a geostationary Earth orbit vantage point to realize a transformational advance in carbon and water monitoring beyond the synoptic capability of the current LEO instruments. It is designed to measure high-resolution spectra of reflected sunlight in near-infrared bands and to deliver simultaneous retrievals of column-averaged dry air mole fractions of CO<sub>2</sub>, CH<sub>4</sub>, CO, and H<sub>2</sub>O (denoted XCO<sub>2</sub>, XCH<sub>4</sub>, XCO, and XH<sub>2</sub>O, respectively) at different times of day over North America. A prototype of the instrument on top of Mount Wilson, in the San Gabriel Mountains just north of Los Angeles, California, has been collecting high-quality data over that megacity since August 2011 (Key et al., 2012; Fu et al., 2014; Wong et al., 2015). The main scientific objectives of GeoFTS are to:

1. Provide observations with the spatial and temporal density required to enable reliable flux inversions, which allow assessments of surface-atmosphere carbon exchange on scales that mirror fundamental carbon cycle processes. To better constrain the flux inversions, the requirements for single-sounding XCO<sub>2</sub>, XCH<sub>4</sub>, and XCO retrieval precisions are less than 0.5 % (~ 2 ppmv), 1 % (~ 18 ppbv), and 10 % (~ 12 ppbv), respectively (Rayner et al., 2001, 2014; Miller et al., 2007).

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2. Quantify the magnitude and spatiotemporal patterns of atmospheric signatures of anthropogenic CO<sub>2</sub>, CH<sub>4</sub>, and CO emissions (e.g., Newman et al., 2013).
3. Examine the close coupling between the carbon and water cycles with precise retrievals of XCO<sub>2</sub>, XCH<sub>4</sub>, XCO, and XH<sub>2</sub>O.

5 Simulation studies have been done in the past to evaluate the design of proposed observing systems and have proven to be informative for the design of new missions. For example, Kuang et al. (2002) conducted an introductory study on the potential of using a 3-band spectrometric approach to accurately measure atmospheric CO<sub>2</sub> from space. The OCO-2 mission has adopted the measurement strategy pioneered  
10 by their simulation study. Buchwitz et al. (2013) developed an error parameterization scheme for the proposed CarbonSat mission and assessed atmospheric CO<sub>2</sub> and CH<sub>4</sub> retrieval errors in a simulated environment. Polonsky et al. (2014) conducted end-to-end simulated retrievals for the proposed geoCARB mission and found that accurate measurements of trace gases make it possible to monitor localized emission sources  
15 such as power plants.

In this study, we perform radiative transfer simulations over both clear-sky and all-sky scenes expected to be observed by GeoFTS and estimate the prospective performance of retrievals based on results from Bayesian error analysis and characterization. This paper is structured as follows. Section 2 explains the methodology of simulated  
20 retrievals. Section 3 presents and discusses the results of column-averaged retrievals, their sensitivity to signal-to-noise ratio (SNR) and spectral resolution, and the influence of aerosols and clouds. Section 4 summarizes the major scientific contributions and some limitations of this study.

## 2 Methodology

This section presents the design of the observing system for GeoFTS, the set-up for the simulated retrievals, the forward model, the retrieval algorithm, and some diagnostics of the observing system.

### 2.1 Design of the GeoFTS observing system

The science objectives of GeoFTS require continuous coverage of bioactive land surfaces, as well as major population centers. It will cover the latitude range of 60° N to 60° S based on an assumed 95° W geostationary longitude. For the purposes of this study, we focus only on North America from 60 to 30° N. Figure 1 shows a schematic diagram of the GeoFTS observing system. At a fixed position relative to Earth, an FTS with agile pointing optics can scan all of North America in any pattern. The spatial resolution is designed to match modeling scales: 1–10 km for urban models and 10–50 km for regional flux inversion models. In addition, to maximize the fraction of clear-sky scenes, small footprints are needed. In GeoFTS, each 0.28 × 0.28 degree Instantaneous Field of View (IFoV) scene is imaged onto 64 × 64 pixels of a focal plane array, which provides a 2.7 × 2.7 km<sup>2</sup> size pixel footprint at nadir. The evolving atmospheric compositions are then inferred from the high-resolution spectra measured by the instrument. Fu et al. (2014) provide more details about a prototype of the instrument on Mount Wilson. The real instrument in space will be similar to the prototype, although the viewing geometry and the spectral resolution will be different from those for the ground-based instrument.

For flux modeling, the temporal resolution needs to be paired to regional flux models, which use 1–6 h time steps and estimate daily fluxes (Wu et al., 2011). Considering that about 20–30 % of the scenes over North America are without excessive aerosol and cloud loadings (Bréon et al., 2005), and taking into account the time steps used in flux models, at least four revisits per day are required. This revisit frequency requires GeoFTS to scan all of North America in about three hours when there is sufficient

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reflected sunlight to be measured. This will result in four measurements each day, thus capturing the diurnal variations of trace gases. An integration time of about one minute for each footprint should give a SNR of about 300. The scanning pattern shown in Fig. 1 is one of the many possible options available for GeoFTS. Faster revisit intervals are possible but involve trade-offs between SNR, spatial resolution, geographic coverage, and data quality.

Three spectral bands are chosen to meet the GeoFTS science objectives. Table 1 shows the spectral bands used and the target species in each band. Band 1A is used for retrievals of CH<sub>4</sub>, CO, and H<sub>2</sub>O and Band 1B for CH<sub>4</sub>, CO<sub>2</sub>, and H<sub>2</sub>O. Band 2 is the O<sub>2</sub> A-band, which provides information about surface pressure, aerosols, and clouds (Crisp et al., 2004).

### 2.2 Set-up for simulated retrievals

After designing the observing system, studies need to be performed to evaluate whether this specific design will meet the requirements for single-sounding XCO<sub>2</sub>, XCH<sub>4</sub>, and XCO retrieval precisions. This is done through simulated retrievals, as illustrated in Fig. 2. We use realistic atmospheric profiles from the GEOS-CHEM global 3-D chemical transport model (Bey et al., 2001), with line positions and strengths obtained from the HITRAN 2008 molecular spectroscopic database. These are then input to a Line-By-Line Radiative Transfer Model (LBLRTM) (Clough et al., 2005) to generate the optical depths (ODs). The state-of-the-art Vector Linearized Discrete Ordinate Radiative Transfer (VLIDORT) model (Spurr, 2006) generates the synthetic radiance data after its model output is convolved with the instrument line shape (ILS) and white Gaussian measurement noise is added. A combination of a two-stream (2S) multiple scattering and an exact single scattering (ESS) model, henceforth referred to as the 2S-ESS model, generates the modeled radiances for the retrieval. The 2S-ESS model also accounts for scattering effects of aerosols and clouds (Spurr and Natraj, 2011; Natraj, 2013), albeit not as accurately as VLIDORT. The differences between the VLI-

DORT and 2S-ESS models give a conservative representation of the forward model error in real retrievals.

The atmospheric profile used to generate the synthetic data is regarded as the “truth”. In retrievals where we do not know the “truth”, we first have some prior knowledge about the concentration of gases in the air column,  $x_a$ . This a priori is used as input to the LBLRTM and 2S-ESS models to generate the modeled radiances. The state vector  $x$  is updated iteratively until the modeled radiances match the synthetic data adequately. If and when the two spectra do match, the state vector  $x$  is then called the “retrieved” state.

Ten test locations in North America are chosen over a wide range of latitudes and longitudes on 26 and 30 July 2006. These two days are a Wednesday and a Sunday, respectively, thus accounting for the weekday-weekend variations in anthropogenic emissions. We will show specific results from Location 1 in Central California and retrieval statistics based on all ten locations. Due to computational resource constraints, the original 116-level atmospheric profiles from the GEOS-CHEM model are interpolated to 20 carefully chosen levels. The Solar Irradiance Reference Spectra (SIRS) (Woods et al., 2009) is linearly interpolated to the wavenumbers in the GeoFTS spectral bands. For simplicity, surface reflection is assumed to be Lambertian. Broadband surface emissivity for each location is obtained from spaceborne measurements (Jacob et al., 2004).

### 2.2.1 Forward model

A forward model is a function that translates the concentration of gases in the atmosphere to the upwelling radiances measured by an instrument in space. In this study, there are two components to the forward model. The first component is the LBLRTM model that takes in an atmospheric profile and computes the ODs of each gaseous species based on spectroscopic information in the HITRAN 2008 database (Rothman et al., 2009). The absorption cross section ( $\sigma$ ) from the LBLRTM, the concentration of the gaseous species ( $[C]$ ) at various atmospheric heights ( $z$ ), and the total column-

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integrated OD of the gas ( $\tau$ ) are related as follows:

$$\tau = \int_0^{\infty} \sigma \cdot [C] dz. \quad (1)$$

Figure 3 shows the ODs of the four target species. CO<sub>2</sub>, CH<sub>4</sub>, and H<sub>2</sub>O absorb strongly in these spectral bands. Although CO only absorbs weakly, its absorption strength is sufficient for good retrievals in clear-sky conditions, as will be shown in Sect. 3.1. For all-sky retrievals, four types of aerosols (black carbon, organic carbon, seasalt, and sulphate) with realistic vertical profiles and diurnal variations are introduced into the model atmosphere. Based on cloud observations from the CALIPSO mission (Winker et al., 2010), we also introduce a small amount of ice clouds to the lower stratosphere. The optical properties of both aerosols and ice clouds are calculated using Mie theory with parameterized Gamma particle size distributions (Heymsfield et al., 2002).

The second component of the forward model is the radiative transfer model (VLIDORT for generating synthetic data and 2S-ESS for retrieval). Both models take in the same optical depths of all the absorbing and scattering species (GHG, aerosols, and ice clouds) and compute the upwelling radiances based on the viewing geometry, the solar irradiance, and other auxiliary parameters. The zenith and azimuth angles of the Sun and the satellite are computed for each location at different times of day. The main difference between the two models is the better characterization of the scattering effects of aerosols and clouds in VLIDORT. This comprehensive radiative transfer model can use any number of computational quadrature angles (“streams”) for the discrete ordinates calculation and we employ 16 streams in our calculations. Both models are run with a spectral resolution of 0.05 cm<sup>-1</sup>. Then, these high-resolution spectra are convolved with the GeoFTS ILS function to produce low-resolution spectra (0.25 cm<sup>-1</sup>) (see Fig. 2 of Fu et al., 2014 for the ILS). The radiances generated from VLIDORT, with measurement noise added in, mimic the measurements that GeoFTS would make from

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the geostationary orbit. The outputs from 2S-ESS are the modeled radiances used in retrievals. Initially, the SNR is set at 300 and the spectral resolution at 0.25 cm<sup>-1</sup>. These two parameters can be changed in the actual design of the observing system. The choices of SNR and spectral resolution and their impact on the retrieval results are discussed in Sect. 3.3.

Several assumptions have been made in the forward model. It has been assumed that the spectroscopic parameters and the ILS are known with certainty and the surface is Lambertian. In reality, uncertainties in these variables would lead to retrieval errors (e.g., Wunch et al., 2011; Fu et al., 2014). In addition, no micro-windows are used in this study. They could be selected in the future to alleviate the problem of interference among the retrieved parameters and to speed up the numerical computations (e.g., Rodgers, 1998; Vidot et al., 2009; Kuai et al., 2010).

### 2.2.2 Retrieval algorithm

We use a non-linear Maximum A Posteriori (MAP) method to iteratively retrieve the state vector  $\mathbf{x}$  (Rodgers, 2000). This method for retrieval has been widely used for GOSAT (Yoshida et al., 2011), Tropospheric Emission Spectrometer (Bowman et al., 2006), and OCO-2 (Kuang et al., 2002; Connor et al., 2008; O'Dell et al., 2012; Crisp et al., 2012). The state vector  $\mathbf{x}$ , model parameters  $\mathbf{b}$ , and error  $\mathbf{e}$  are related to the measurement vector  $\mathbf{y}$  by a forward model  $\mathbf{F}$ :

$$\mathbf{y} = \mathbf{F}(\mathbf{x}, \mathbf{b}) + \mathbf{e}. \quad (2)$$

The model parameters include trace gas spectroscopy and geometric and optical properties of aerosols and clouds. The errors include both the forward model error and the instrument noise. The goal of the non-linear MAP method is to minimize the Bayesian least-squares cost function, which is defined as follows:

$$J(\mathbf{x}) = [\mathbf{y} - \mathbf{F}(\mathbf{x}, \mathbf{b})]^T \mathbf{S}_e^{-1} [\mathbf{y} - \mathbf{F}(\mathbf{x}, \mathbf{b})] + (\mathbf{x} - \mathbf{x}_a)^T \mathbf{S}_a^{-1} (\mathbf{x} - \mathbf{x}_a), \quad (3)$$

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$\mathbf{S}_a$  and  $\mathbf{S}_e$  are the error covariance matrices for the a priori and the measurement, respectively.  $\mathbf{S}_e$  is assumed to be a diagonal matrix with values of the square of the noise.  $\mathbf{x}$  and  $\mathbf{x}_a$  are the true state and the a priori, respectively. The Levenberg-Marquardt method (Levenberg, 1944; Marquardt, 1963) is used to obtain stable retrieval results within a trust region for nonlinear least-squares problems. In each iteration, the  $(i + 1)$ th state vector,  $\mathbf{x}_{i+1}$ , is related to the  $i$ th state vector,  $\mathbf{x}_i$ , as follows:

$$\mathbf{x}_{i+1} = \mathbf{x}_i + [(1 + \gamma)\mathbf{S}_a^{-1} + \mathbf{K}_i^T \mathbf{S}_e^{-1} \mathbf{K}_i]^{-1} \{ \mathbf{K}_i^T \mathbf{S}_e^{-1} [\mathbf{y} - \mathbf{F}(\mathbf{x}_i)] - \mathbf{S}_a^{-1} [\mathbf{x}_i - \mathbf{x}_a] \}. \quad (4)$$

The parameter  $\gamma$  is set at 10 initially and updated after each iteration.  $\mathbf{F}(\mathbf{x}_i)$  is the forward model output based on the state vector  $\mathbf{x}_i$ .  $\mathbf{K}_i$  ( $\delta \mathbf{F}(\mathbf{x}_i) / \delta \mathbf{x}_i$ ) is a Jacobian matrix that describes how the modeled radiances change given an infinitesimal change in  $\mathbf{x}_i$ . It is computed using the finite difference method. The iterations stop when each element of  $\mathbf{x}_{i+1}$  differs from the corresponding one in  $\mathbf{x}_i$  by less than 0.001. Typically, it takes five to ten iterations for retrievals to converge to stable results. Since this study is meant to serve as a proof of concept, the state vector is slightly simpler than that anticipated to be used in the real retrievals. The state vector is defined in the following way:

$$\mathbf{x} = \begin{bmatrix} \eta(\text{CO}_2) \\ \eta(\text{CH}_4) \\ \eta(\text{CO}) \\ \eta(\text{H}_2\text{O}) \\ \eta(\text{T}) \\ \eta(\text{p}) \\ \eta(\varepsilon) \\ \eta(\text{fs}) \\ \eta(\text{zo}) \end{bmatrix} \text{ for total column retrieval in clear sky conditions.} \quad (5)$$

Here  $\eta$  stands for the multiplicative scaling factor retrieved for the atmospheric profile and other parameters (T: temperature profile; p: pressure profile;  $\varepsilon$ : broadband surface emissivity; fs: frequency shift; zo: zero-level offset). For each element in the

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state vector, the a priori is biased 3% with respect to the true state, so as to allow easy comparison among retrieved values. The a priori one-sigma ( $1\sigma$ ) error is also set at 3%, thus providing a fairly loose constraint on the a priori and making the retrieval dependent on the measurement rather than the prior knowledge. Note that the a priori constraints are small compared to the true uncertainties of XH<sub>2</sub>O and XCO. More realistic a priori values will be used in operational retrievals. For this simulation study, we choose to set all of the a priori biases and errors to 3% so as to compare the retrieval results for all four target species. All the elements are assumed to be independent of each other, i.e.,  $\mathbf{S}_a$  is diagonal. The retrieved state vector can be expressed as:

$$\hat{\mathbf{x}} = \mathbf{A}\mathbf{x} + (\mathbf{I}_n - \mathbf{A})\mathbf{x}_a + \mathbf{G}\mathbf{e}. \quad (6)$$

The first term on the right hand side,  $\mathbf{A}\mathbf{x}$ , represents the contribution of the true state to the retrieved state. The averaging kernel  $\mathbf{A}$  is defined as:

$$\mathbf{A} = \frac{\partial \hat{\mathbf{x}}}{\partial \mathbf{x}} = (\mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1} \mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K} = \mathbf{G}\mathbf{K}. \quad (7)$$

It quantifies the relative contribution to the retrieval from the true state, compared to that from the a priori. Correspondingly,  $(\mathbf{I}_n - \mathbf{A})$  represents the relative contribution to the retrieval from the a priori. The second term,  $(\mathbf{I}_n - \mathbf{A})\mathbf{x}_a$ , is also known as the smoothing error because it smoothes the solution towards the a priori. In the third term,  $\mathbf{G}$  stands for the gain matrix and  $\mathbf{G}\mathbf{e}$  represents the retrieval error due to the random instrument noise and the forward model error. The error covariance matrix for the retrieved state,  $\hat{\mathbf{S}}$ , quantifies the uncertainties in the retrieval based on  $\mathbf{S}_a$ ,  $\mathbf{S}_e$ , and  $\mathbf{K}$ :

$$\hat{\mathbf{S}} = (\mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1}. \quad (8)$$

The square roots of its diagonal values give the  $1\sigma$  uncertainty of the retrievals.

### 2.3 Diagnostics of the observing system

In order to examine whether the design of the GeoFTS observing system would result in precise trace gas retrievals, we implement the following diagnostics (Rodgers, 2000):

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- Jacobian (**K**): it is the sensitivity of the modeled radiances to the state vector. A large Jacobian indicates that a gaseous species absorbs strongly in the spectral band.
- Averaging Kernel (**A**): it quantifies the relative contribution to the retrieval from the true state, compared to that from the a priori. Ideally, it should be close to unity.
- Degrees of Freedom for signal (DoF): it describes the number of useful independent quantities there are in a measurement. DoF = trace(**A**).
- Information Content (*H*): it quantitatively describes how well the measurements increase our confidence in the estimation of the atmospheric state relative to the prior knowledge. As a scalar quantity, *H* is useful for characterizing and comparing observing systems. Mathematically, it is defined as:

$$H = \frac{1}{2} \ln(|\hat{\mathbf{S}}^{-1} \mathbf{S}_a|). \quad (9)$$

- Retrieval error: it is defined as the absolute percentage difference between the retrieval and the truth.
- Single-measurement precision: it is the  $1\sigma$  uncertainty of the retrieval  $\sigma_p$ , where “p” stands for the posterior state.
- Uncertainty Reduction (UR): it is defined as  $(\sigma_a - \sigma_p)/\sigma_a \times 100\%$ , where  $\sigma_a$  and  $\sigma_p$  stand for uncertainty of the a priori and the a posteriori, respectively.

### 3 Results and discussion

#### 3.1 Simulated retrievals in clear sky conditions

For clear-sky retrievals, the forward model generates the upwelling radiances based on the trace gas atmospheric profiles and other auxiliary parameters. Figure 4 illustrates the modeled radiances at four times of day at Location 1, a test location in Central California. As expected, the reflected sunlight captured by GeoFTS will be more intense at 12 p.m. and 3 p.m. and less so at 9 a.m. and 6 p.m. This demonstrates that radiative transfer models dedicated to geostationary measurements can realistically simulate the diurnal variations of radiances.

Figure 5 shows the Jacobians for the four target species.  $\text{CO}_2$ ,  $\text{CH}_4$ , and  $\text{H}_2\text{O}$  absorb strongly in the two bands, while the  $\text{CO}$  absorptions are much weaker. The relative absorption strengths suggest that the retrieval results are likely to be better for  $\text{CO}_2$ ,  $\text{CH}_4$ , and  $\text{H}_2\text{O}$ . Through the iterative retrieval algorithm, a fit is obtained between the modeled radiances and the synthetic data. Figure 6 shows a sample spectral fit and residuals in all spectral bands. The fact that the residuals fluctuate randomly around zero and follow an approximate Gaussian distribution (not shown) indicates that most of the information from the spectra has been extracted. Minor inaccuracies in the retrieved state result in some small spikes in the residuals. If the forward model is perfect and the SNR is 300, the theoretical lower limit for the root-mean-square (rms) value of the residuals is approximately 0.33%. The real rms values of the residuals from this simulated retrieval are close to the theoretical value, implying that most of the residuals are due to the random instrument noise, and not the forward model error.

The averaging kernel approaches a unitary matrix (not shown). This indicates that most of the elements in the retrieved state are close to the true state. Since the state vector for clear-sky retrievals contains 9 elements, as shown in Eq. (5), the theoretical upper limit for DoF is 9. Through simulated retrievals, the real DoF is calculated to be around 8.7. It falls slightly short of 9 because of the relatively weak absorption of  $\text{CO}$ .

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findings that, due to interferences, the quality of the XCO retrieval is systematically influenced by the shortcomings in CH<sub>4</sub> and H<sub>2</sub>O spectroscopy. Nonetheless, the precisions for the simulated retrievals far exceed the measurement requirements discussed in Sect. 1, and thus leave a large margin for other sources of error such as uncertainty in the spectroscopic parameters.

We also notice that the precision and the UR of XCO retrievals tend to follow a diurnal pattern: precision is worse and UR lower at 12 and 3 p.m. than at 9 a.m. and 6 p.m. This might be due to the different light paths at different times of day. At 12 and 3 p.m., the solar zenith angles are between 35 and 50°, so incoming sunlight travels through a relatively short path. In contrast, at 9 a.m. and 6 p.m., the solar zenith angles are between 65 and 85°, so the sunlight traverses a longer light path in the atmosphere, resulting in stronger signals and better precision. In other words, this diurnal pattern in precision might be a feature associated with the viewing geometry from geostationary orbit. Whether this diurnal pattern will persist in more simulated or real retrievals warrants more studies in the future.

### 3.2 Influence of aerosols and clouds

Aerosol and cloud contamination are substantial sources of systematic errors in trace gas retrievals (e.g., Dufour and Bréon, 2003; Houweling et al., 2005; Butz et al., 2011; O'Dell et al., 2012; Merrelli et al., 2015; Zhang et al., 2015). Novel cloud screening algorithms, such as the one designed for the OCO-2 mission, are able to screen out most of the scenes with thick clouds and aerosols. However, even after the pre-screening, some scenes are still contaminated with a small amount of aerosols and ice clouds that have a column-integrated OD less than ~ 0.30 (O'Dell et al., 2012). In these cases, inaccurate retrievals of the optical properties of aerosols and ice clouds will negatively affect the quality of trace gas retrievals.

Here we briefly explore the influence of aerosols and ice clouds on trace gas retrievals and defer detailed investigations to later studies. We add realistic profiles of aerosols and ice clouds to the model atmosphere. It is assumed that the geometric



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and optical properties of aerosols and ice clouds are known with certainty. Scaling factors for aerosol OD and ice cloud OD are added to the state vector in Eq. (5) and retrievals are repeated for the test cases at 12 p.m., a time with larger radiance values and higher instrument noise than those at other times of day. Again, the a priori is 3 % biased from the truth and has a 3 % uncertainty. Compared to the clear-sky retrievals in Sect. 3.1, these retrievals are referred to as the all-sky retrievals. The sum of the aerosol and ice cloud ODs is the total scattering OD.

Retrieving the ODs of aerosols and clouds from reflected sunlight in near-infrared bands, has been a significant challenge and an active area of scientific research (e.g., Reuter et al., 2010; Herbin et al., 2013). As there are now 11 elements in the state vector (9 elements in Eq. 5 and 2 new elements added), the theoretical upper limit for DoF is 11. However, the real DoF is only about 10.2, mainly due to the lack of information to accurately retrieve the ODs of aerosols and clouds.

Figure 8 shows the retrieval errors, precisions, and URs of the four target species as a function of the total scattering OD. Increases in the total scattering OD lead to notable increases in the retrieval errors of all four gases. For XCH<sub>4</sub> and XH<sub>2</sub>O, the average retrieval errors increase from < 0.4 % for clear-sky retrievals to 1–3 % for all-sky retrievals. For XCO<sub>2</sub>, the errors increase from ~ 0.2 to ~ 0.5 %. For XCO, the errors increase from ~ 1.8 to 2–6 %. Previous studies have obtained similar results. For example, Vidot et al. (2009) conducted simulated retrievals of XCO<sub>2</sub> over liquid water clouds with OCO-2's three spectral bands. They found that the presence of undetected thin cirrus clouds affect all retrieval parameters significantly. These results corroborate Vidot et al.'s findings that inaccurate retrievals of aerosol and cloud ODs cause substantial biases in XCO<sub>2</sub> retrievals. Through simultaneous retrievals of all four gases, we find that aerosols and clouds cause biases not only in XCO<sub>2</sub> retrievals, but also in all the trace gas retrievals studied here. Generally, the higher the total scattering OD, the worse the retrieval errors are. This suggests that more effective pre-screening algorithms and better prior knowledge of the aerosol and cloud properties and/or vertical distributions are needed in order to mitigate errors caused by aerosols and clouds.

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There are two main causes behind the increase in the retrieval error. One is the interference between absorption by gases and scattering by aerosols and clouds. The other one is the increase in the forward model error when the total scattering OD increases. VLIDORT-16 stream is better at characterizing the multiple light scattering by aerosols and clouds, while 2S-ESS treats light scattering in a more approximate manner. This leads to an increase in the forward model error, and therefore, a corresponding increase in the retrieval error. These results highlight the need for sophisticated radiative transfer models such as VLIDORT, along with better models to characterize the properties of aerosols and clouds and/or their vertical distributions, so that the forward model error can be minimized for real operational retrievals.

In contrast, the single-measurement precision does not change much when the total scattering OD changes. This is expected because, according to Eq. (8), the a posteriori uncertainty and the precision depend on  $\mathbf{S}_a$ ,  $\mathbf{S}_e$ , and  $\mathbf{K}$  only. Changes in total scattering OD only affect  $\mathbf{S}_e$  and  $\mathbf{K}$  slightly, resulting in some minor worsening of the retrieval precision, as illustrated in Fig. 8. These results corroborate Kuang et al. (2002)'s findings from a simulated study. They also investigated the achievable XCO<sub>2</sub> precisions as a function of the total scattering OD and found that, as OD increases from 0.05 to 0.25, the retrieval precision of XCO<sub>2</sub> over land only changes from  $\sim 0.5$  ppm ( $\sim 0.15\%$ ) to  $\sim 1.0$  ppm ( $\sim 0.30\%$ ). Our results show that the retrieval precisions of all four gases are only slightly affected by the small amount of aerosol and cloud contamination. Thus, this study provides evidence that the single-measurement precisions for both clear-sky and all-sky retrievals meet measurement requirements for accurate flux inversions (Sect. 1).

### 3.3 Choices of SNR and spectral resolution

The previous sections present retrieval results when SNR and spectral resolution are set at 300 and  $0.25\text{ cm}^{-1}$ , respectively. Optimal choices of SNR and spectral resolution can result in better cost-benefit ratio for a geostationary mission. Here we study the effect of varying SNR and spectral resolution on both clear-sky and all-sky retrievals.

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For simplicity, the two variables are assumed to be independent of each other. To allow easy comparison with retrieval results in Sect. 3.1 and 3.2, the a priori values are same as before. Figure 9 shows the retrieval precision as a function of SNR and spectral resolution, which are set to vary from 50 to 250 and from 0.10 to 0.30 cm<sup>-1</sup>, respectively.

These ranges are chosen so as to cover some possible scenarios for real spaceborne measurements. The total scattering OD is zero for clear-sky retrievals and 0.20 for all-sky retrievals. The results shown here are based on one test case at 12 p.m. at Location 1. Other test cases exhibit similar patterns (not shown).

For both clear-sky and all-sky retrievals, the retrieval precision is a strong function of both SNR and spectral resolution. Generally, it increases with higher SNR and finer spectral resolution. Comparing clear-sky and all-sky retrievals, we notice that the precisions for all trace gas retrievals are slightly worse in all-sky retrievals. This is a manifestation of the influence of aerosols and ice clouds, as discussed in Sect. 3.2. In particular, the XCO retrievals are more sensitive to changes in SNR and spectral resolution because of its weak absorption. It is clear that the quality of XCO retrievals is a significant function of both SNR and spectral resolution. The importance of XCO retrievals for surface flux inversions has been emphasized in Rayner et al. (2014). They investigated the ability of trace gas retrievals to constrain regional GHG emissions and found that XCO retrievals play the most important role in constraining urban emissions at 3 km resolution. Our results show that the choices of SNR and spectral resolution will be of paramount importance for XCO retrievals. This exploratory study could help guide decisions on GeoFTS SNR and spectral resolution, which can result in cost-effective measurement strategies while achieving satisfactory levels of retrieval precisions.

## 4 Conclusion

This study explores the prospective performance of trace gas retrievals using simulated spectra measured by a Geostationary Fourier Transform Spectrometer (GeoFTS). The major scientific contributions of this study are as follows:

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1. We have performed simulated retrievals over a wide range of atmospheric and surface scenarios in clear sky conditions over North America, one potential target for the GeoFTS. We find that, with no uncertainty in spectroscopy, the average retrieval errors and single-measurement precisions are  $<0.2\%$  for XCO<sub>2</sub>, XCH<sub>4</sub>, and XH<sub>2</sub>O, and  $<2\%$  for XCO. These results are based on a priori values with a bias of 3% and an uncertainty of 3%.
2. We have briefly explored the influence of aerosols and ice clouds on trace gas retrievals. An increase in the amount of aerosols and ice clouds leads to a notable increase in the retrieval errors and slight worsening of the retrieval precisions. This study provides evidence that the single-measurement precisions for simulated clear-sky and all-sky retrievals meet the requirements for GeoFTS.
3. We have studied the effects of varying signal-to-noise ratio (SNR) and spectral resolution on both clear-sky and all-sky retrievals. We find that retrieval precision is a strong function of the two variables. This exploratory study can help guide decisions on GeoFTS SNR and spectral resolution, which can result in cost-effective measurement strategies while achieving satisfactory levels of retrieval precisions.

While this study offers new scientific insights through simulated retrievals, it is worth pointing out that these numerical simulations often provide overoptimistic assessments of the capability of the observing system. This is partly because the instrument noise added to the synthetic data is random, which is unlikely to be the case for real measurements. Furthermore, uncertainty in the spectroscopic parameters, imperfect aerosol/cloud filtering, and various instrument issues will complicate real retrievals. Future studies will focus on how uncertainties in these factors affect the retrieval results.

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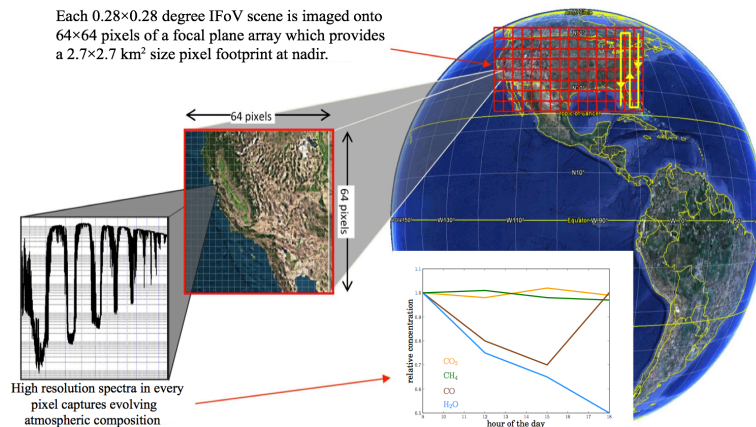


**Table 1.** The GeoFTS spectral bands and the target species.

Band Number	Spectral Band	Target Species
1A	2.3 $\mu\text{m}$ (4210–4320 $\text{cm}^{-1}$ )	CH <sub>4</sub> , CO, H <sub>2</sub> O
1B	1.6 $\mu\text{m}$ (5950–6100 and 6190–6260 $\text{cm}^{-1}$ )	CH <sub>4</sub> , CO <sub>2</sub> , H <sub>2</sub> O
2	0.76 $\mu\text{m}$ (13 000–13 170 $\text{cm}^{-1}$ )	O <sub>2</sub>

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**Figure 1.** A schematic illustrating the GeoFTS observing system. Each  $0.28 \times 0.28^\circ$  Instantaneous Field of View (IFOV) scene is imaged onto  $64 \times 64$  pixels of a focal plane array, which provides a  $2.7 \times 2.7 \text{ km}^2$  size pixel footprint at nadir. The evolving atmospheric compositions at different times of day are then inferred from the high-resolution spectra measured by the instrument. The relative concentrations of four target species with respect to those at 9 a.m. are shown for illustration purposes only.

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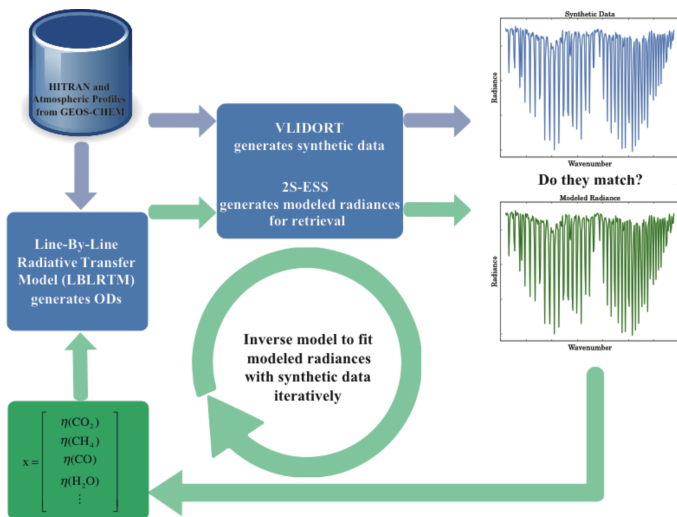
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**Figure 2.** The set-up for simulated retrievals. Realistic atmospheric profiles are obtained from GEOS-CHEM and line positions and strengths from the HITRAN 2008 molecular spectroscopic database. These are then input to a Line-By-Line Radiative Transfer Model (LBLRTM) to generate optical depths. VLIDORT, a state-of-the-art radiative transfer model, generates the synthetic data after its model output is convolved with the instrument line shape (ILS) and random measurement noise is added. A fast two-stream (2S), enhanced single scattering (ESS) radiative transfer model, 2S-ESS, generates the modeled radiances for retrievals. In simulated retrievals, we first have some prior knowledge about the concentration of gases and other parameters. This a priori is used as input to 2S-ESS to generate the modeled radiances. The state vector  $x$  is updated iteratively in an inverse model until the modeled radiances match the synthetic data adequately.

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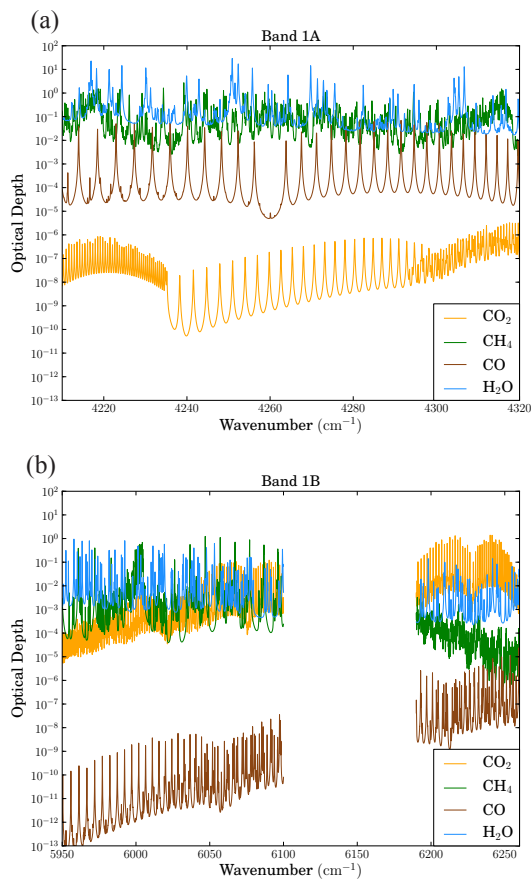
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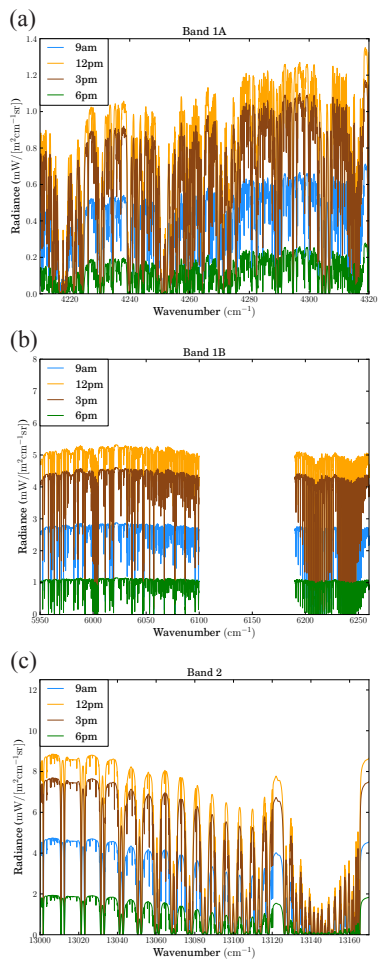


**Figure 3.** The optical depths of four target species in the GeoFTS spectral bands: **(a)** for Band 1A and **(b)** for Band 1B. Band 1A is for retrievals of CH<sub>4</sub>, CO, and H<sub>2</sub>O. Band 1B is for retrievals of CH<sub>4</sub>, CO<sub>2</sub>, and H<sub>2</sub>O.



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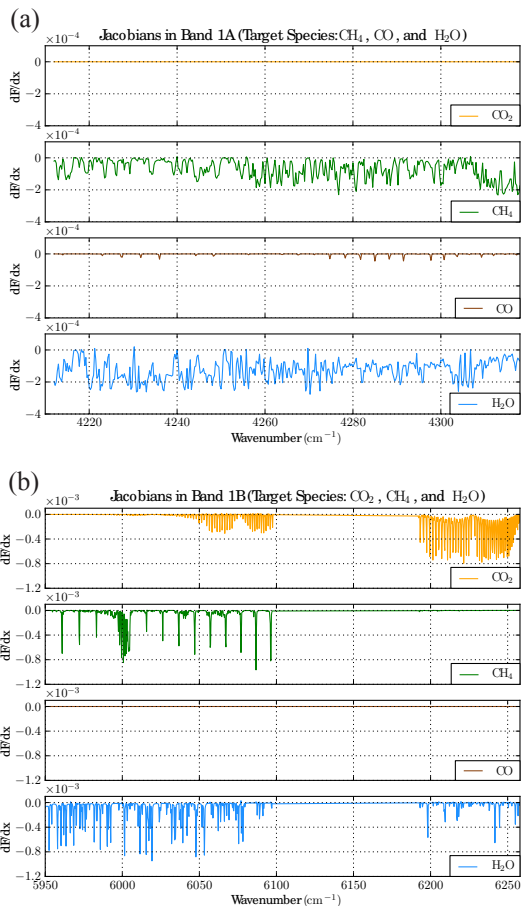
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**Figure 4.** Modeled radiances in three spectral bands at four times of day at Location 1 on 26 July 2006.

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**Figure 5.** The Jacobians of four target species in (a) Band 1A and (b) Band 1B. Since  $x$  is a unitless scaling factor, the Jacobian ( $dF/dx$ ) has the same unit as the modeled radiances in Fig. 4.

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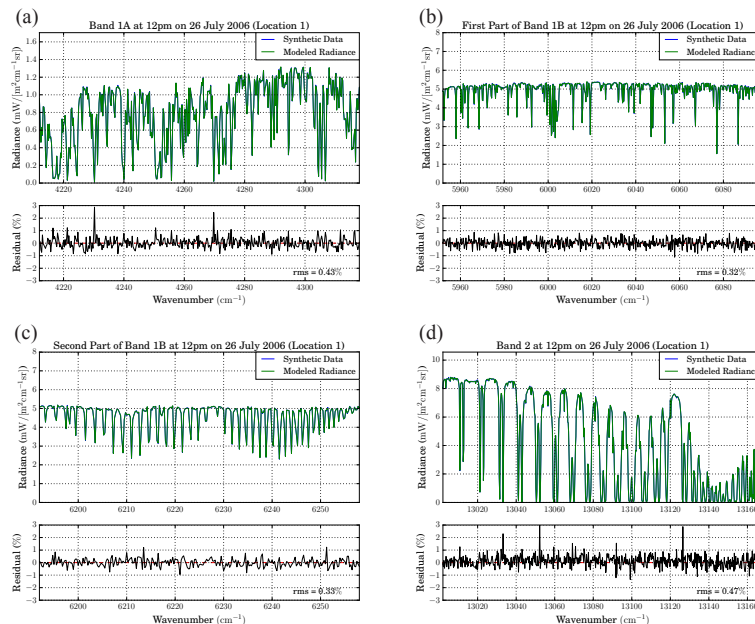
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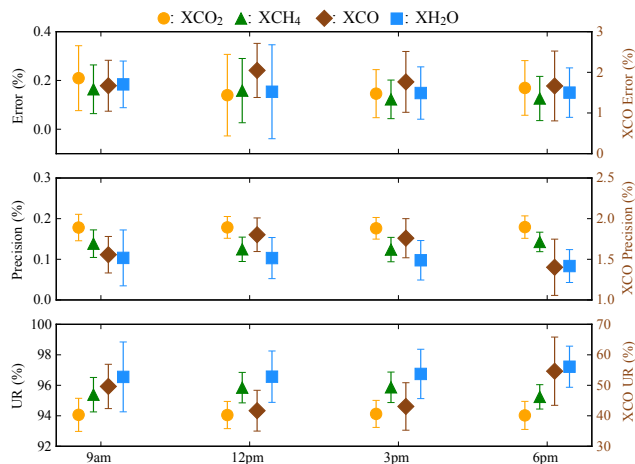
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**Figure 6.** A sample spectral fit and residuals from a simulated clear-sky retrieval on 26 July 2006 at Location 1. The root-mean-square (rms) values of the residuals are between 0.32 and 0.47 %.

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**Figure 7.** Average retrieval errors, precisions, and uncertainty reductions (UR) for simulated clear-sky retrievals at different times of day. SNR is 300 and spectral resolution is  $0.25\text{ cm}^{-1}$ . The a priori is biased 3% from the truth and has a 3% uncertainty. The error bars are the standard deviations of retrieval results from all the test cases. Note that the scales for XCO<sub>2</sub>, XCH<sub>4</sub>, and XH<sub>2</sub>O are on the left-hand side whereas those for XCO are on the right-hand side.

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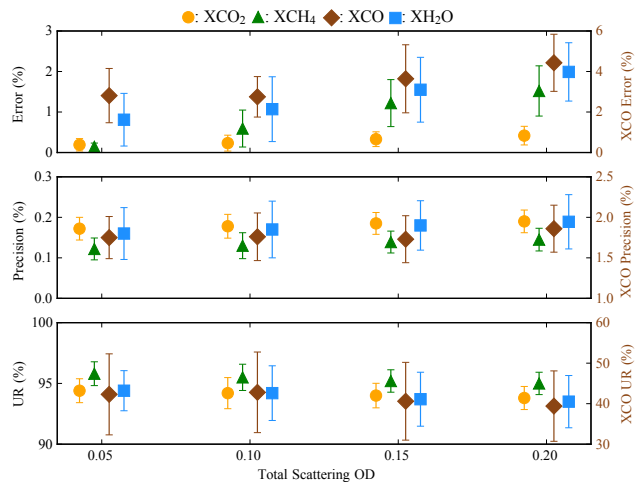
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**Figure 8.** Average retrieval errors, precisions, and uncertainty reductions (UR) for simulated all-sky retrievals as a function of the total scattering optical depth (OD). SNR is 300 and spectral resolution is  $0.25\text{ cm}^{-1}$ . The a priori is biased 3% from the truth and has a 3% uncertainty. The error bars are the standard deviations of retrieval results from the test cases.

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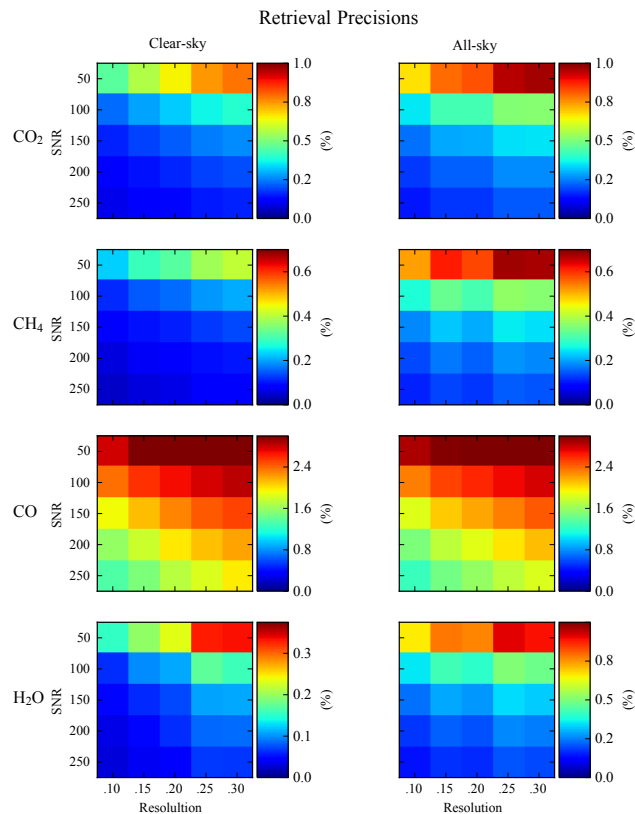
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**Figure 9.** Retrieval precision as a function of SNR and spectral resolution for simulated clear-sky and all-sky retrievals. SNR and spectral resolution are set to vary from 50 to 250 and from 0.10 to 0.30 cm<sup>-1</sup>, respectively. The a priori is biased 3% from the truth and has a 3% uncertainty for all species.

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