Uncertainties of satellite-derived surface skin temperatures in the polar oceans: MODIS, AIRS/AMSU, and AIRS only

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Abstract

Uncertainties in the satellite-derived Surface Skin Temperature (SST) data in the polar oceans during two periods (April 16-24 and September 15-23) of 2003-2014 were investigated and the three datasets were intercompared as follows: MODerate Resolution Imaging Spectroradiometer Ice Surface Temperature (MODIS IST), the SST of the Atmospheric Infrared Sounder/Advanced Microwave Sounding Unit-A (AIRS/AMSU), and AIRS only. AIRS only algorithm was developed in preparation for the degradation of the AMSU-A. MODIS IST was systematically warmer up to 1.65 K at the sea ice boundary and colder down to -2.04 K in the polar sea ice regions of both the Arctic and Antarctic than that of the AIRS/AMSU. This difference in the results could have been caused by the surface classification method. The spatial correlation coefficient of the AIRS only to the AIRS/AMSU (0.992-0.999) method was greater than that of the MODIS IST to the AIRS/AMSU (0.968-0.994). The SST of the AIRS only compared to that of the AIRS/AMSU had a bias of 0.168 K with a RMSE of 0.590 K over the northern hemisphere high latitudes and a bias of -0.109 K with a RMSE of 0.852 K over the southern hemisphere high latitudes. There was a systematic disagreement between the AIRS retrievals at the boundary of the sea ice, because the AIRS only algorithm utilized a less accurate GCM forecast over the seasonally-varying frozen oceans than the microwave data. The three datasets (MODIS, AIRS/AMSU and AIRS only) showed significant warming rates (2.3±1.7 ~ 2.8±1.9 K/decade) in the northern high regions (70-80 N) as expected from the ice-albedo feedback. The systematic temperature disagreement associated with surface type classification had an impact on the resulting temperature trends.

1. Introduction

The satellite observations of the polar oceans have been more challenging than those of non-frozen ocean and land, because it is more difficult to identify clouds over the various surfaces (Tobin et al., 2006). The surface skin temperature (SST) is one of the most important climate variables that is related to the surface energy balance and the thermal state of the atmosphere (Jin et al., 1997). Compared to ground-based observations, satellite-observed SST data play a crucial role in climate study and model development by providing a uniform resolution data over the globe. The retrievals of AIRS data over the last decade have a significant contribution to various climate studies and model evaluations (Aumann et al., 2003; Tian et al., 2013; Yoo et al., 2013). AIRS retrievals have produced atmospheric temperature, moisture, and ozone profiles on a global scale by the AIRS method itself or together with other instruments (Liu at al., 2008). A lot of comparisons of the AIRS/AMSU data against data from numerical forecast model analysis fields, radiosondes, lidar, and retrievals from high-altitude aircraft have been used to assess the accuracy of the retrievals (Tobin et al., 2006; Susskind et al., 2014). The AIRS retrieval algorithm has been developed and validated gradually with clear sky and clear/cloudy conditions over a non-frozen ocean and then the non-polar land and polar cases (Tobin et al., 2006).

The AIRS/AMSU has an advantage of measuring the radiation penetrating through clouds and polar darkness, and has high spectral resolution and coarse spatial resolution (Dong et al., 2006). However, the AIRS only algorithm using only AIRS observations has been developed due to the degradation of the AMSU-A. Microwave and multispectral radiometers were used for global mapping of the sea ice extent and dynamics, while the visible, near-infrared, and infrared sensors could obtain details on the ice concentration, snow/ice albedo, thickness, and IST during clear-sky conditions (Hall et al., 2004; Scott et al., 2014). The MODIS on Earth Observing System (EOS) Aqua, used as infrared measurements, was influenced by water and cloud contamination, but had a higher spatial resolution

(Dong et al., 2006). In order to remove the cloud effects in the MODIS IST algorithm, MODIS cloud mask products were used (Hall et al., 2004).

Since AIRS and MODIS were co-located on Aqua, they have often been used to make a synergistic algorithm and they have been compared to each other frequently (Li et al., 2005; Molnar and Susskind, 2005). Molnar and Susskind (2005) validated the accuracy of the AIRS/AMSU cloud products using MODIS cloud analyses, which have a higher spatial resolution than that of AIRS. Knuteson et al. (2006) compared the MODIS Collection 4 (C4) with the AIRS Version 3 (V3) on the land surface temperature (LST) for the eastern half of the U.S., showing that the monthly differences were approximately 3 K. Lee et al. (2013) investigated the characteristics of the differences between the MODIS land surface skin temperature/sea surface temperature and the AIRS/AMSU surface skin temperature across the globe, and found that the MODIS C5 product was systematically lower by 1.7 K than the AIRS/AMSU V5 product over land in the 50 N-50 S regions, but it was higher by 0.5 K than the AIRS/AMSU product over ocean. Particularly in the sea ice regions, the MODIS annual averages were larger than the AIRS/AMSU values, due to the differential errors in ice/snow emissivity between the retrieval methods (or channels) for the two data products. The differences between the MODIS and AIRS methods were reduced when the MODIS IST and AIRS/AMSU surface skin temperatures were compared for 9-days. The possible reasons for this include the satellite local crossing time (LCT) difference between them due to the different swath width in the high latitude regions, and the emissivity difference between microwave and infrared channels, but more comparison studies are necessary for a longer period to pin down the reasons of such skin temperature discrepancies between MODIS and AIRS/AMSU.

The primary purpose of this study was to investigate a relative degree of agreement (or disagreement) among different SST datasets using the MODIS IST C5, the SST of the AIRS/AMSU, and AIRS only V6. The second purpose of this paper was to analyze the temperature trend differences affected by the temperature differences among different data products. The datasets used in this study were described in section 2. In section 3, we compared the MODIS and AIRS only data with the AIRS/AMSU values. We also analyzed the temperature trends from the three satellite-based datasets in section 4, and in the conclusion we summarized our study.

2. Data and methods

The Aqua satellite carrying the AIRS, AMSU and MODIS instruments was launched on May 4, 2002 with the Earth's Radiant Energy System (CERES), Humidity Sounder for Brazil (HSB) and the Advanced Microwave Scanning Radiometer-EOS (AMSR-E). It has far exceeded its designed life span of 6 years and has a chance of operating into the 2020's (http://aqua.nasa.gov/). The Aqua satellite orbits the earth every 98.8 minutes with an equatorial crossing time going north (ascending) at 1:30 p.m. local time (daytime) and going south (descending) at 1:30 a.m. (nighttime) in a sun-synchronous, near polar orbit with an inclination of 98.2° and an operational altitude of 705 km (Tian et al., 2013).

As shown in Table 1, we used the datasets of MODIS IST (e.g., Hall et al., 2006) and SSTs of AIRS/AMSU and AIRS only over the northern hemisphere during April 16-24 and over the southern hemisphere on September 15-23 from 2003 to 2014 in order to avoid the polar night when the visible channels of the MODIS did not operate (Hall et al., 2004). The sea surface temperature observed from infrared channels of satellites indicates the values at the skin of sea water, in contrast with the sea surface temperature measured from buoys, of which values represent the temperature of bulk water near the sea surfaces. The infrared sea SST was measured at depths of approximately 10 μm within the oceanic skin layer (~500 μm) at the water side of the air-sea interface where the conductive and diffusive heat transfer processes dominated (Emery et al., 2001; Donlon et al., 2002; Liou, 2002).

As an imaging spectroradiometer, the MODIS with 36 bands has retrieved various physical parameters such as aerosol optical thickness, land and water surface temperature, leaf area index, and snow cover, etc. (Barnes et al., 1998; Hall and Riggs, 2007). MODIS produced the 'sea ice by reflectance' and 'IST' in order to identify sea ice (Riggs et al., 1999). The 'sea ice by reflectance' was determined by the Normalized Difference Snow Index (NDSI), and the reflectance of Band 1 (0.645 μm) and Band 2 (0.858 μm). The NDSI was calculated using Band 4 (0.555 μm) and Band 7 (2.130 μm). IST is used as another method for identifying sea ice. The IST derived from the "split-window method" in Eq. (1), where bands 31 and 32 are centered at approximately 11 μm and 12 μm, respectively. The method was applied in order to identify the ice when the IST was less than 271.5 K. The cutoff temperature between water and ice (271.5 K) may vary depending on the region and season. The IST is calculated as follows:

$$IST = a + b\,T_{11} + c\,(T_{11}-T_{12}) + d\,[(T_{11}-T_{12})(\sec(\theta)-1)] \qquad (1),$$

where T_{11} is the brightness temperature (K) in 11 μm, T_{12} is the brightness temperature (K) in 12 μm, and θ is the scan angle from nadir. The difference between the T_{11} and the skin temperature from the LOWTRAN can be less than 3 K for a skin temperature between 230 K and 260 K (Key et al., 1997). Since the value of T_{11} itself was a good estimate, coefficients a-d were defined for the following temperature ranges: $T_{11} < 240$ K, 240 K $\leq T_{11} \leq$ 260 K, and 260 K $< T_{11}$ (Riggs et al., 2006). The IST algorithm was only applied to the polar ocean pixels that were determined to be clear by the MODS cloud mask using visible reflectance (Hall et al., 2004). The surface in the IST algorithm was assumed to be snow (Key et al., 1997). MODIS ISTs were provided as daily polar fields with a 4 km by 4 Km resolution.

The AIRS spectrometer is a high spectral resolution spectrometer with 2,378 channels in the thermal infrared spectrum and 4 bands in the visible spectrum (Won, 2008). The AIRS and AMSU were coupled in order to play a role as an advanced sounding system under clear and cloudy conditions (Aumann et al., 2003). The AIRS/AMSU algorithm is independent of the GCM, except for the use of GCM surface pressure to determine the bottom boundary conditions (Molnar and Susskind, 2005). V6 is the most current retrieval algorithm since the launch of AIRS instrument, and detailed descriptions are given in Olsen (2013b). The primary products from AIRS suite include the atmospheric temperature-humidity profiles, ozone profiles, sea/land surface skin temperature (SST), and cloud related parameters such as the outgoing longwave radiation (OLR) (Susskind et al., 2011). In the AIRS/AMSU algorithm, the surface classification was conducted using the brightness temperature difference in 23 GHz (AMSU ch1) and 50 GHz (AMSU ch3). The difference (brightness temperature at 23 GHz minus brightness temperature at 50 GHz) had a negative value on the sea ice and a positive value on the water (Grody et al., 1999; Hewison and English, 1999). Also, the brightness temperature difference between 23 GHz (AMSU ch1) and 31 GHz (AMSU ch2) could distinguish the age of the sea ice (Kongoli et al., 2008). The accuracy of AIRS/AMSU SST can be affected by surface misclassification, which is caused by the surface emissivity changes, the pixel mixed with the various surface types, and the ice pixel pooled with water.

After the surface type classification from the AMSU retrieval, the initial state for atmospheric and surface parameters, cloud parameters and OLR was generated using the Neural Network methodology (Susskind et al., 2011, 2014). The methodology was used to approximate some functions between the input and output vectors by training (Gardner and Dorling, 1998). Next, the initial clear column radiances were generated, which were based on the initial state and the observed infrared radiances. The surface and atmospheric variables, including the surface skin temperature, were retrieved by updating the cloud cleared infrared radiance, iteratively. The cloud properties and outgoing longwave radiation were then retrieved, followed by the error estimates and quality control. In the AIRS only V6, shortwave window region 3.76–4.0 μm was used in order to derive the surface skin temperature and surface spectral emissivity (ε).

AMSU channels 4-5 had not been available since 2007 and 2010, respectively, due to radiometric noise. In preparation for the degradation of the other AMSU channels, the AIRS only algorithm was developed excluding the AMSU observations. The algorithm was similar to that of AIRS/AMSU, but it did not use the AMSU-A observations in any step of the physical retrieval process and the quality control methodology. Instead, the AIRS only algorithm used the NOAA Global Forecast System (GFS) for surface classification purposes (Olsen, 2013a).

We calculated the climatology and anomaly values from the yearly 9-day mean temperatures in a 1° by 1° grid for the 12-year period of Aqua satellite observations in order to estimate the temperature anomaly trends. The trends of the MODIS IST were derived only when the number of yearly data was at least 10 out of 12 entire years at each grid point. The trends of the AIRS/AMSU and AIRS only were derived only when the number of yearly datasets was 12, covering the entire years of the analysis at each grid. The bootstrap method (Wilks, 1995) was used to calculate at a 95% confidence interval. In the method, 10,000 linear temperature trends were generated by random sampling, allowing repetition of 10,000 yearly anomaly temperature datasets. Then, we estimated the 95% confidence interval of 10,000 temperature trends.

3. Comparison of the satellite-derived surface skin temperatures

Figure 1a shows the spatial coverage and the averaged value of the MODIS IST over the southern hemisphere during September 15-23 of 2003-2014. In order to solve the spatial resolution mismatches, the original resolution of the MODIS data with a 4 km by 4 km grid (Fig. 1a) was re-gridded to a 1° by 1° grid in the case of MODIS data present over 50 % (Fig. 1b). A grid spacing of 1° corresponds to approximately 111 km on the equator, and it becomes reduced poleward. In the zonal averaged SST analysis, this 50% criterion was used. During the same period, the spatial distributions of the climatological T_{skin} (AA_V6) and T_{skin} (AO_V6) were also shown in Fig. 1c-d, respectively. As expected, the MODIS and AIRS showed the spatial distribution of the climatological SST, warmer at the lower latitudes than the higher latitudes. The SST distributions over the northern hemisphere during April 16-24 of 2003-2014 have been shown in Kang and Yoo (2015).

Figure 2a displays the number of years when both T_{skin} (MODIS) and T_{skin} (AA_V6) are available at each grid point over the southern hemisphere. The number near 60 S was smaller than that of the other regions because the MODIS IST algorithm only produced its data in the cloud-free pixels. Similar distributions by the clouds were shown in Fig. 2b and 2d for the same reason. The reduced number of observations near 60 S had a spatial distribution similar to that of the frontal cloud bands that were likely associated with the mid-/high-latitude depressions encircling the Antarctica (e.g., Jakob, 1999; Comiso and Stock, 2001; Lachlan-Cope, 2010; Boucher et al., 2013). Figure 2c shows the number of years when both T_{skin} (AO_V6) and T_{skin} (AA_V6) were available at each grid. Most of the grids had both T_{skin} (AO_V6) and T_{skin} (AA_V6) for a period of more than 10 years.

Figure 3a presents the spatial distribution of the temporal difference in a 1° x 1° grid between the climatological T_{skin} (MODIS) and T_{skin} (AA_V6) during April 16-24 of 2003-2014 over the northern hemisphere. In general, T_{skin} (MODIS) at 60-70 N was higher than the T_{skin} (AA_V6). T_{skin} (MODIS) was about 3 K higher than the T_{skin} (AA_V6) on the Hudson Bay and near Greenland, whereas it was about -2 K lower near the center of the Arctic Ocean. The relationship between the climatological T_{skin} (MODIS) and T_{skin} (AA_V6) was presented in the scatter diagrams (Fig. 3b). The scatter plot revealed a temperature interval which deviated from the simple linear line. The discontinuous shape appeared at the freezing point (~273 K) and the turning point (~260 K) in terms of T_{skin} (MODIS), changing the coefficient of the MODIS IST algorithm. In the interval, T_{skin} (MODIS) was systematically higher than the T_{skin} (AA_V6) in the 260 – 273 K range of T_{skin} (MODIS). The slope in the range was 0.85, lower than the slope for the whole regression line (0.97). There was a better agreement in the 240-260 K range, where the difference between the T_{11} and the SST in the LOWTRAN was less than 3 K (Key et al., 1997). The better agreement in the range greater than 280K was also shown.

Figure 3b was the same as Fig. 3a except for T_{skin}(MODIS) vs. T_{skin}(AO_V6). The differences between the two datasets were very similar to those in Fig. 3a. However, T_{skin} (MODIS) was more than 4 K higher than T_{skin} (AO_V6) in some regions near the Greenland and the Barents Sea. The slope (0.93) in the 260 – 273 K range of T_{skin}(MODIS) also indicated a deviation from the total slope (0.98) in the scatter plot (Fig. 3d), similar to that in Fig. 3b. Figure 3e showed the difference between T_{skin} (AO_V6) to T_{skin} (AA_V6). Overall, the agreement was much better than the previous two cases (Fig. 3a and c), except for in the Greenland Sea, the Barents Sea, and the Okhotsk Sea. Both T_{skin} (AO_V6) and T_{skin} (AA_V6) agreed with each other (r=0.999) well except for near the freezing point.

Figure 4 showed discrepancies among the three types of SST datasets over the southern hemisphere during September 15-23 of 2003-2014. It has been noted that there was a latitudinal band encircling Antarctica at 60-70 S, where T_{skin} (MODIS) was higher than both the T_{skin} (AA_V6) and T_{skin} (AO_V6) (Fig. 4a and c). The circular region corresponded to the sea ice/water boundary which was expected to move seasonally. This implies a systematic difference between T_{skin} (MODIS) and T_{skin} (AA_V6) in the sea ice classification. The corresponding scatter plots also revealed a discontinuous (i.e., not linear) shape in the 260 – 273 K range of T_{skin}(MODIS) (Fig. 4b and d). The slopes in that range were 0.84-0.94, which were smaller than the slope (0.98) in the whole range. In addition, T_{skin} (MODIS) showed lower temperature values than T_{skin} (AA_V6) and T_{skin} (AO_V6) near the Antarctic peninsula, in the region from Weddell Sea to Ross Sea (Fig. 4a and c).

The comparison between two types of AIRS datasets also showed the circular pattern around Antarctica where T_{skin} (AO_V6) was lower by 1.5-5.6 K than T_{skin} (AA_V6) (Fig. 4e). The discrepancy near the sea ice/water boundary was clear, possibly due to the difference in the sea ice detection method between the two datasets. The uncertainty of the SST at the sea ice boundary was distinguished from the other regions. Both T_{skin} (AO_V6) and T_{skin} (AA_V6) were in good agreement, other than the sea ice/water boundary regions. The scatter pattern of the T_{skin} (AA_V6) versus that of the T_{skin} (AO_V6) showed that the two datasets generally agreed with each other, but the disagreement near the freezing point again occurred indicating a cold bias of AIRS only with respect to AIRS/AMSU (Fig. 4f).

Figure 5 showed the annual-average spatial distributions for T_{skin} (MODIS) minus T_{skin} (AA_V6) in the southern hemisphere from September 15-23 of 2003-2014. Although the 9-day composite values were used in each year, T_{skin} (MODIS) data did not exist in some areas. It was because the MODIS IST algorithm was valid only for cloud-free pixels. The systematic positive values at the boundary of the sea ice consistently occurred, while the negative ones occurred on some areas of the sea ice near Antarctica every year.

Figure 6 presented the interannual variation of the spatial distribution of T_{skin} (AO_V6) minus T_{skin} (AA_V6) for the study period. As already seen in Fig. 4e, the values of T_{skin} (AO_V6) compared to T_{skin} (AA_V6) show systematic negative values encircling Antarctica during the period. In addition, there were positive values over the sea-ice prevailing areas inside the circle, with the location varying from year to year, which must be related to the difference in the surface type characterization.

Table 2 showed the statistics of bias, spatial correlation coefficient (r), and root mean square error (RMSE) obtained from the 12-year climatologies of 2003-2014 in order to analyze the systematic error among the three types of satellite-observed temperatures quantitatively. This analysis for each hemispheric vernal period has been performed over the two regions (35-90 N, 60-90 N) of the northern hemisphere during April 16-24, and over the regions (40-90 S, 60-90 S) of the southern hemisphere during September 15-23. The spatial correlation coefficient between the two satellite data sets was computed in this study as follows; ⅰ) The climatological 9-day composite data of SSTs during 2003-2014 were computed in a 1° × 1° grid of the two data sets, respectively. ⅱ) We computed the spatial correlation coefficient between the two datasets, using their climatological values in a 1° × 1° grid within a given latitude band. The values in parentheses indicated the average obtained from the statistics for each year and their corresponding standard deviations. Based on the climatology values, the SST of the AIRS retrievals were comparable with respect to the T_{skin} (MODIS) (r=0.959-0.994). T_{skin} (MODIS) tended to systematically exceed the AIRS retrievals over the polar oceans (bias = 0.198-0.597 K). Hall et al. (2004) reported the accuracy of T_{skin} (MODIS) with the bias values of 1.2-1.3 K near the South Pole and the Arctic Ocean. The RMSE of 1.847 K for T_{skin} (MODIS) vs. the T_{skin}(AA_V6) over 60-90 S in our study was slightly higher than that in the study of Hall et al. (2004).

From the intercomparison of the three datasets, the bias (-0.109-0.597) and RMSE (0.590-2.173) over the high latitude belt (60-90 N and S) tended to be larger, and the correlation coefficients (r=0.959-0.986) was smaller than those over 35-90 N and 40-90 S among the three comparisons (Table 2). This result indicated that there was more disagreement over the high latitudes than over other regions. The spatial correlation coefficient (0.992-0.999) between T_{skin} (AO_V6) and T_{skin} (AA_V6) was greater than those (0.968-0.994) between T_{skin} (MODIS) and T_{skin} (AA_V6). In the high latitudes T_{skin} (AO_V6) with respect to T_{skin} (AA_V6) had a positive bias of 0.168 K with a RMSE of 0.590 K in the northern hemisphere, but a bias of -0.109 K with a RMSE of 0.852 K in the southern hemisphere. The high correlations (r = 0.998-0.999) between the AIRS/AMSU and AIRS only (i.e., AIRS retrievals) over the 35-90 N and 40-90 S areas showed that the AIRS only can be a good alternative for the AIRS/AMSU, except for at the region of the sea ice boundary (r = 0.992 over the 60-90 S). The disagreement between T_{skin} (AA_V6) and T_{skin} (AO_V6) at the region where the sea ice and water mixed appeared, because the AIRS only used less accurate GCM forecast data for surface classification over the potentially frozen oceans.

Figure 7 presents the zonal mean temperature difference among the three satellite-observed datasets in a 1° × 1° grid over the northern hemisphere during April 16–-24 of 2003-2014 and over the southern hemisphere during September 15-23, 2003-2014. The red, blue and green lines represent the zonally averaged annual values of T_{skin} (MODIS) minus T_{skin} (AA_V6), T_{skin} (MODIS) minus T_{skin} (AO_V6), and T_{skin} (AO_V6) minus T_{skin} (AA_V6), respectively. The climatological annual values have been calculated from the interannually-varying yearly data, shown in Fig. 8. The black dashed line, the difference between the original MODIS IST data (4km x 4km) and converted T_{skin}(MODIS) (1°×1°) indicated the possible error from the conversion of spatial resolution. The differences by the conversion over both hemispheres were within 0.3 K and 0.5 K, respectively. The original T_{skin} (MODIS), converted T_{skin} (MODIS), T_{skin} (AA_V6), and T_{skin} (AO_V6) were chosen under the same condition in space and time, and each grid (1°×1°) of a degree latitudinal band.

It is hard to see in Fig. 3a the systematic difference due to the sea ice detection over the northern hemisphere because of the continental distribution. However, Fig. 7 clearly showed that the difference among the T_{skin} (MODIS), T_{skin} (AA_V6), and T_{skin} (AO_V6) existed over the northern hemisphere. T_{skin}(MODIS) was warmer than T_{skin}(AA_V6) in 56-81 N and 54-69 S, while cooler than T_{skin}(AA_V6) in the other latitudinal zone. It has been noted that the peak of the difference between T_{skin} (MODIS) and two AIRS datasets in the northern hemisphere high-latitude region took place in a broader region than in the southern hemisphere. T_{skin} (MODIS) was up to 1.65 K higher than the AIRS datasets at the boundaries of the sea ice/water, whereas it was lower by up to -2.04 K over the sea ice region. The

MODIS IST algorithm was the optimized on the snow/ice surface type, and thus the underestimation of T_{skin} (MODIS) in the 35-54 N and 40-55 S may not be unexpected. In general, the overestimation of T_{skin} (MODIS) to the AIRS retrievals occurred at the sea ice boundary and the underestimation occurred in the sea ice region that can be covered with snow/ice.

The grey solid lines in Fig. A1a-b mean the 5% significance level of the differences between T_{skin} (MODIS) and T_{skin} (AA_V6), and between T_{skin} (AO_V6) and T_{skin} (AA_V6) over a possibly frozen region (poleward from 50 N and 50 S, respectively). Based on the *t*-test (von Storch and Zwiers, 1999) at significance level of $p<0.05$, the temperature disagreement between T_{skin} (MODIS) and T_{skin} (AA_V6) (red solid line) is significant in 50-55 N, 58-70 N, 89-90 N, 50-53 S, and 57-62 S (Fig. A1a). Considering the uncertainty of MODIS due to the conversion of spatial resolution (black dashed line), the temperature disagreement in 57-62 S can become insignificant. However, the discrepancy in 58-70 N is significant even if the uncertainty of MODIS is considered. The difference between T_{skin} (AO_V6) and T_{skin} (AA_V6) in 53-60 S is significant (Fig. A1b).

The color-coded lines in Fig. 8 interannually represent the differences in temperature among the three datasets for individual years. The thick black lines indicated the yearly difference averages. There existed a significant degree of interannual variation in the difference between T_{skin} (MODIS) and the two AIRS datasets (Fig. 8a-b). The variation was larger in 2009, 2010 and 2011 over the regions northward of 60 N and southward of 55 S where sea ice existed. Figure 8b shows a value of T_{skin} (MODIS) minus T_{skin} (AO_V6) that was similar to that in Fig. 8a. T_{skin} (MODIS) was lower than T_{skin} (AO_V6) at the ice surface, but higher than T_{skin} (AO_V6) at the boundary of the sea ice. Figure 8c showed the interannual variation of T_{skin} (AO_V6) minus T_{skin} (AA_V6). The interannual variation of the difference between the AIRS retrievals was much larger in the high latitude than in the mid-latitudes. The maximum difference of 1.56 K between the AIRS retrievals was found at 87-88 N in 2011.

There could be several reasons for the observed differences between T_{skin} (MODIS) and T_{skin} (AA_V6). The main one can be attributed to the difference in the channel used for the retrievals of the skin temperature. The AIRS/AMSU V6 only utilized shortwave window channels for the surface skin temperature, while the MODIS IST algorithm used the longwave window regions. The shortwave window could be mixed with the solar radiation during the daytime, but it was suitable for temperature sounding (Chahine, 1975, 1977; Susskind et al., 2014). The advantage of the longwave window was that its range corresponded to the peak of the infrared radiation emitted from the earth (Prakash, 2000). On the other hand, the longwave window radiation could be affected more by clouds. In order to avoid cloud contamination, the MODIS IST algorithm analyzed the pixel when the MODIS cloud mask was reported as clear sky (Hall et al., 2004). The MODIS cloud mask using visible reflectance had a high accuracy during the daytime, but a lower accuracy during the nighttime due to low illumination. As another reason for the temperature difference, Lee et al. (2013) suggested that there were substantial differences in LCT between MODIS and AIRS in the high latitude regions, since the different scan angles of the two instruments resulted in different footprints, which could lead to the observed difference in temperature. However, we suggested that the surface type classification method could be the primary reason for the temperature difference between the MODIS-based and AIRS-based datasets. AIRS/AMSU SST was retrieved after the surface type was classified. On the other hand, the MODIS IST was calculated without the surface type classification step. Then, the MODIS algorithm categorized pixels being ice if IST was less than the cutoff temperature. MODIS IST was calculated on the snow, sea ice, and ocean, assuming the surface was snow-covered (sea ice). The IST was utilized as a criterion for identifying the ice/water which might cause significant disagreement between the T_{skin} (MODIS) and T_{skin} (AA_V6) in the range of 260-273 K.

4. Comparison of the surface skin temperature trends from the MODIS and AIRS/AMSU

In order to further investigate the effects of the difference among the satellite-observed temperatures from different measurement techniques or algorithms on the temperature anomaly trend, we calculated the trend in some latitude belts, using the three satellite-observed temperature datasets at each grid during April 16-24 of 2003-2014 (in the northern hemisphere) and September 15-23 of 2003-2014 in the southern hemisphere. During this period, an unusually extensive surface melting event was observed in 2012 (Nghiem et al., 2012; Hall et al., 2013; Comiso and Hall, 2014).

Table 3 shows the temperature anomaly trend with a 95% confidence level on the 10° latitude belt. We arranged the data of MODIS IST, AIRS/AMSU, and AIRS only under the same condition in space and time. The significant warming trend in 70-80 N was estimated in the following order: AIRS/AMSU (2.83 K/decade) > AIRS only (2.71 K/decade) > T_{skin}(MODIS) (2.30 K/decade). The warming (0.10 to 0.38 K/decade) at 40-50 N and 50-60 S, and the cooling (-0.08 to -1.94 K/decade) at 80-90 N, 60-70 N, 50-60 N, 60-70 S and 70-80 S of the three datasets occurred, but the trends were not significant. Comiso and Hall (2014) reported the SST trend using the Goddard Institute for Space Studies (GISS) dataset as 0.60 K/decade and the trend using the Advanced Very High Resolution Radiometer (AVHRR) dataset as 0.69 K/decade in the Arctic (> 64 N) during 1981-2012. Our result in 70-80 N, compared with the above studies, seems to indicate an acceleration in the Arctic warming.

The warming trend in the northern hemispheric high latitudes had been known to be caused in part by the well-known positive feedback among snow/ice, surface albedo and temperature (Curry et al., 1995; Comiso and Hall, 2014). T_{skin} (MODIS) had a greater cooling tendency compared to T_{skin} (AA_V6) in the higher latitude regions (70-90 N and 60-80 S) (Table 3). The trend difference between the two temperatures was -0.69 K/decade at 70-80 S. The trend difference of the T_{skin} (AA_V6) and T_{skin} (AO_V6) (i.e., AIRS only minus AIRS/AMSU) was the largest (-0.26 K/decade) at 60-70 N. The cooling trend (-0.90 K/decade) of the T_{skin} (AO_V6) was greater than that (-0.65 K/decade) of T_{skin} (AA_V6) at the latitude band.

Figure 9a-b showed the SST anomaly trends from the T_{skin} (MODIS) in a 1° by 1° grid over the northern hemisphere during April 16-24 of 2003-2014 and over the southern hemisphere during September 15-23 of 2003-2014. The T_{skin} (MODIS) trend was calculated on the grid, which had available data that existed for over 10 years. Figure 9c-d and Fig. 9e-f showed the trend data for T_{skin} (AA_V6) and T_{skin} (AO_V6), respectively, which all had 12-year data, individually. The trend distributions in all three of the datasets were similar over the northern hemisphere. Warming trend in the Beaufort Sea, East Siberian Sea and Kara Sea was detected, while cooling was observed in the Hudson Bay and near Greenland. The significant warming trend appeared at 70-80 N as shown in Table 3, and the trend based on the spatial distribution varied depending on the regions (Fig. 9a, c and e). According to Comiso and Hall (2014), a strong warming trend (> 1.5 K/decade) existed near the Kara Sea and Baffin Bay among the entire Arctic, consistent with the noticeable trend revealed near the Kara Sea in our study. Over the southern hemisphere, there were not enough data to derive a trend for T_{skin} (MODIS) mostly due to clouds. The trend analysis over the sea ice regions from T_{skin} (AA_V6) and T_{skin} (AO_V6) showed a strong cooling trend, especially near the Antarctic peninsula between the Weddell and Ross Seas (Fig. 9d and f). The cooling trend was generally dominant over the southern hemisphere. Marshall et al. (2015) suggested that based on the model experiments, the cooling trend around Antarctica as opposed to the warming trend around the Arctic Ocean was the result of the offset between the greenhouse gas and ozone hole responses, emphasizing the larger cooling effects associated with the Antarctic ozone hole.

The 12-year mean of the T_{skin} (MODIS) minus T_{skin} (AA_V6) (Fig. 10a and c) and of the trend difference between T_{skin} (MODIS) and T_{skin} (AA_V6) (Fig. 10b and d) were compared in order to reveal the relationship between the temperature difference and the corresponding trend difference over the northern hemisphere during April 16-24 of 2003-2014 and over the southern hemisphere during September 15-23 of 2003-2014. T_{skin} (MODIS) was higher than T_{skin} (AA_V6) over the bays of

Hudson and Baffin, and Bering Sea (Fig. 10a). The warming trend of the T_{skin} (MODIS) was also greater than that of the T_{skin} (AA_V6) over the Hudson Bay and near the Kara Sea (Fig. 10b). The data for the trend difference in the southern hemisphere was not sufficient due to the missing data of T_{skin}(MODIS) in the cloudy condition (Fig. 10d).

Figure 11 showed over both hemispheres the 12-year mean of the T_{skin} (AO_V6) minus T_{skin} (AA_V6) (Fig. 11a and c) and the corresponding trend difference of the T_{skin} (AO_V6) and T_{skin} (AA_V6) (Fig. 11b and d). The relationship of the temperature difference and trend difference over the southern hemisphere in Fig. 10 was hard to analyze due to the absence of a T_{skin} (MODIS) trend (Fig. 10c-d). However, Fig. 11c-d clearly showed that the temperature difference had a significant impact on the trend difference over the southern hemisphere. The trend of the T_{skin} (AA_V6) and T_{skin} (AO_V6) agreed well except for at the region of the sea ice boundary, implying that the algorithm identifying the sea ice affected the SST trend.

Uncertainties among satellite observations (T_{skin} (MODIS), T_{skin} (AA_V6), and T_{skin} (AO_V6)) in the sea ice region of the northern hemisphere are generally similar to those of the southern hemisphere in terms of zonal averages. However, the systematic difference between the observations can be more clearly seen in the latter region than in the former region due to more oceanic regions in the southern hemisphere (Figs. 10-11, and see also Fig. 7).

Table 4 quantitatively showed how the temperature differences among the three types of SST affected each trend difference over the hemispheric regions poleward either from 50 N (shown in the left side of the table) during April 16-24 of 2003-2014 or from 50 S during September 15-23 of 2003-2014. In the upper portion, the average of the temperature difference and the trend difference in the grid corresponding to the temperature difference condition was used, whereas the average values on the grids that had the same signs for the temperature difference and the trend difference were used in the lower portion. Only the cases where grid number was greater than 100 were considered. The warmer temperature led to relatively warming trend, the cooler temperature led to relatively cooling trend. When the T_{skin} (MODIS) was greater than T_{skin} (AA_V6) in the regions poleward from 50 S, the trend difference was in the reduced cooling trend (i.e., warmer direction) as -0.96, -0.66, and -0.21 K/decade with the conditions of T_{skin} (MODIS) minus T_{skin} (AA_V6) rising as more than 1 K, 1.5 K, and 2 K, respectively. The uncertainty of the satellite-derived temperatures had a substantial effect on the uncertainty of the temperature trends. The data set has been reduced in the lower section of Table 4. The sample size can affect the estimated impact of ΔT on ΔTrend, but it looks like that the impact on the trends in the lower section is almost consistent with that in the upper section despite the reduced sample sizes.

5. Conclusions

The satellite-derived L3 products of MODIS IST and two SSTs from AIRS/AMSU and AIRS only were investigated with a comparative analysis during the vernal periods of 2003-2014: April 16-24 over the northern hemisphere and September 15-23 over the southern hemisphere. The original MODIS IST data were regridded onto a 1° × 1° grid box for comparison with the AIRS retrievals. The difference between the original MODIS IST and the converted one was within 0.5 K in a latitudinal belt.

The differences among the three types of satellite derived SST data were most prominent over the sea ice regions. T_{skin} (MODIS) and T_{skin} (AA_V6) were comparable (r=0.97-0.99), but there existed systematic disagreement occurred in the T_{skin} (MODIS) range of 260-273 K. The southern hemispheric high latitude (60 S-90 S) was the primary contributor to the disagreement between them. In comparison with the T_{skin} (AA_V6) in a latitudinal belt, the T_{skin} (MODIS) was higher by up to 1.65 K than T_{skin} (AA_V6) on the boundary of the sea ice/water, whereas it was lower by up to -2.04 K in the sea ice region.

The spatial correlation coefficients (0.992-0.999) of the T_{skin} (AO_V6) and T_{skin} (AA_V6) over both hemispheres were greater than those (0.968-0.994) between T_{skin} (MODIS) and T_{skin} (AA_V6). The T_{skin} (AO_V6) compared to the T_{skin}(AA_V6) had a bias of 0.168 K with a RMSE of 0.590 K over the northern hemisphere high latitudes and a bias of -0.109 K with a RMSE of 0.852 K over the southern hemispheric high latitudes. There was a systematic disagreement between the T_{skin} (AA_V6) and T_{skin} (AO_V6) at the sea ice boundary. It is likely due to the fact that the AIRS only algorithm utilized a less accurate GCM forecast than the microwave data over the seasonally-varying frozen oceans.

The temperature differences among the three types of datasets showed a high degree of interannual variations over the latitudinal belts where sea ice existed. The significant warming rates (2.3±1.7 ~ 2.8±1.9 K/decade) were revealed by all three datasets in the northern hemispheric high-latitude regions (70-80 N) could be interpreted as the ice-albedo feedback. The discrepancies between the trends of the T_{skin} (AA_V6) and T_{skin} (AO_V6) occurred at the sea ice boundary. When the T_{skin} (AA_V6) trends were compared to those of the T_{skin} (MODIS) or T_{skin} (AO_V6) in a 1° × 1° grid, the warmer temperature difference tended to lead to a relative warming trend, whereas the cooler temperature difference tended to lead to a relative cooling trend.

The systematic disagreement between the T_{skin} (MODIS) and T_{skin} (AA_V6) could be caused by (1) the channels used for the surface skin temperature, (2) the cloud contamination, (3) the LCT difference between the MODIS and AIRS, and (4) the surface type classification method. Whereas the AIRS/AMSU V6 used only the shortwave window channels for the surface skin temperature, MODIS IST used the longwave window regions. The MODIS IST product utilized the MODIS cloud mask with visible reflectance, which had lower accuracy during the night (Hall et al., 2004). Lee et al.,

(2013) reported that the LCTs between the MODIS and AIRS were almost the same from 60 N-60 S, but they were quite different in the high latitude regions. It is likely that the main cause to the observed SST differences near the sea ice boundary was in the way the surface type was classified. The AIRS/AMSU algorithm conjugated the emissivity difference in the low and high frequency microwave bands (23 and 50 GHz) in order to identify sea ice. However, MODIS IST was calculated without the surface type classification.

The AIRS/AMSU L2 data offer the surface type (coastline, land, ocean, two types of sea ice, two types of snow, and glacier/snow), and the AIRS/AMSU L3 data offer the number of these various surface types in a grid. The AIRS only L2 also offer the surface type (coastline, land, ocean, two types of sea ice, and snow), and its L3 data offer the number of these various surface types in a grid. Under the condition without ground truth, the direct validation has a limit because the surface classifications of AIRS/AMSU and AIRS only have some difference. Although the AIRS only has utilized the GCM forecast, there is a good agreement in SST between AIRS/AMSU and AIRS only in most regions. However, the disagreement between them over the land regions of the Sahara desert, parts of Spain and in the US with snow cover at night has been reported (Dang et al., 2012).

The SST in the polar region is a useful parameter being used to derive the climate change signal, although it has been challenging to measure an accurate SST. Based on our results from detailed comparative investigation, we cautiously suggested that the observed difference and uncertainty among the satellite-derived SSTs were likely caused by the different sea ice detecting methods used in each algorithm. In addition, the methods also affected the temperature trend. In this study, we aimed to help in understanding characteristics of the infrared and microwave measurements for the surface skin temperature, and the method for identifying sea ice. We believe the results of this study can be useful for the interpretation and the modeling of the climate change associated with the temperature trends.

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Table 1. The information on the satellite-observed surface skin temperature (T_{skin}) Level 3 (L3) data used in this study. Three datasets of T_{skin} were compared over the northern hemisphere during April 16-24 of 2003-2014, and over the southern hemisphere during September 15-23 of 2003-2014. The abbreviations used in this table are as follows: temp (temperature), IST (ice surface skin temperature), OBS (observation), and SFC (surface).

Satellite-observed dataset	Version (Collection)	Temp type	Area	Spatial resolution	Number of OBS	Satellite sensor	Abbreviation	Reference
MODIS IST	MYD29E1D/5	Skin	Polar ocean	4km×4km	1/day	Aqua MODIS	T_{skin}(MODIS)	Hall et al.(2004)
AIRS/AMSU SFC skin temp	AIRX3STD/6	Skin	Globe	1°×1°	2/day	Aqua AIRS/AMSU-A	T_{skin}(AA_V6)	Susskind et al. (2014)
AIRS only SFC skin temp	AIRS3STD/6	Skin	Globe	1°×1°	2/day	Aqua AIRS	T_{skin}(AO_V6)	Susskind et al. (2014)

Table 2. Statistical comparisons of the climatological 9-day composite data during 2003-2014 over both hemispheres; T_{skin} (MODIS) vs. T_{skin} (AA_V6), T_{skin} (MODIS) vs. T_{skin} (AO_V6), and T_{skin} (AO_V6) vs. T_{skin} (AA_V6). The values in this table were calculated based on the 12-year composite mean values. The values in parentheses indicate the 12-year mean values and their standard deviations during 2003-2014. Bias: T_{skin} (MODIS) minus T_{skin} (AA_V6), T_{skin} (MODIS) minus T_{skin} (AO_V6), and T_{skin} (AO_V6) minus T_{skin} (AA_V6), r: correlation coefficient, RMSE: root mean square error.

Region	T_{skin} (MODIS) vs. T_{skin} (AA_V6)			T_{skin} (MODIS) vs. T_{skin} (AO_V6)			T_{skin} (AO_V6) vs. T_{skin} (AA_V6)		
	Bias (K)	r	RMSE (K)	Bias (K)	r	RMSE (K)	Bias (K)	R	RMSE (K)
35-90N	-0.169 (-0.161±0.231)	0.994 (0.990±0.002)	1.491 (1.909±0.156)	-0.289 (-0.324±0.308)	0.993 (0.990±0.003)	1.563 (1.963±0.260)	0.133 (0.137±0.130)	0.999 (0.997±0.001)	0.574 (1.018±0.131)
40-90S	0.026 (-0.010±0.218)	0.989 (0.982±0.003)	1.480 (2.082±0.144)	0.203 (0.035±0.282)	0.985 (0.980±0.003)	1.756 (2.184±0.119)	-0.141 (-0.139±0.079)	0.998 (0.994±0.001)	0.750 (1.272±0.092)
60-90N	0.223 (0.194±0.357)	0.986 (0.973±0.009)	1.501 (1.986±0.227)	0.597 (-0.013±0.475)	0.986 (0.972±0.011)	1.591 (2.033±0.370)	0.168 (0.170±0.214)	0.998 (0.992±0.003)	0.590 (1.027±0.216)
60-90S	0.198 (0.368±0.537)	0.968 (0.906±0.023)	1.847 (2.871±0.276)	0.306 (0.295±0.620)	0.959 (0.898±0.021)	2.173 (2.987±0.271)	-0.109 (-0.108±0.142)	0.992 (0.976±0.005)	0.852 (1.498±0.112)

Table 3. The rate of the surface skin temperature change (K/decade) of the MODIS, AIRS/AMSU, and AIRS only in each 10° latitudinal belt over the northern hemisphere (NH) during April 16-24 and over the southern hemisphere (SH) during September 15-23 of 2003-2014, using their collocated data in a $1^\circ \times 1^\circ$ grid. The ± values define the 95% confidence intervals for the trends. The symbol '*' means the significant value at a 95% confidence interval. Note that the rates are subject to large uncertainty due to the short periods of the satellite-based temperature records.

Latitudinal belt	MODIS	AIRS/AMSU	AIRS only	MODIS minus AIRS/AMSU	MODIS minus AIRS only	AIRS only minus AIRS/AMSU
<NH>						
80-90 N	-0.558±3.101	-0.100±3.673	-0.093±3.736	-0.458	-0.465	0.007
70-80 N	2.302±1.701*	2.826±1.878*	2.711±1.788*	-0.524	-0.409	-0.115
60-70 N	-0.506±1.173	-0.646±1.294	-0.902±1.050	0.140	0.396	-0.256
50-60 N	-0.345±0.539	-0.522±0.628	-0.466±0.550	0.177	0.121	0.056
40-50 N	0.292±0.402	0.103±0.576	0.191±0.565	0.189	0.101	0.088
<SH>						
50-60 S	0.375±0.400	0.315±0.466	0.316±0.600	0.060	0.059	0.001
60-70 S	-1.944±2.271	-1.304±1.890	-1.300±1.918	-0.640	-0.644	0.004
70-80 S	-0.769±2.687	-0.081±2.586	-0.135±2.633	-0.688	-0.634	-0.054

Table 4. Uncertainties of the satellite-derived surface skin temperature rate (or trend; ΔTrend) due to the temperature difference (ΔT_{skin}) for the cases of T_{skin} (MODIS) minus T_{skin} (AA_V6) and T_{skin} (AO_V6) minus T_{skin} (AA_V6) in the upper portion of the table. Also, the values of uncertainties provided in the lower portion of the table indicate the cases of ±ΔTrend with respect to $\pm\Delta T_{skin}$ (double signs in the same order). The uncertainties are not shown when the number of the grid (1°× 1°) points (i.e., No. of grids in the table) is less than 100.

ΔT_{skin} (K)	Poleward from 50N						Poleward from 50S					
	T_{skin} (MODIS) vs. T_{skin} (AA_V6)			T_{skin} (AO_V6) vs. T_{skin} (AA_V6)			T_{skin} (MODIS) vs. T_{skin} (AA_V6)			T_{skin} (AO_V6) vs. T_{skin} (AA_V6)		
	No. of grids	ΔT_{skin}	ΔTrend (K/decade)	No. of grids	ΔT_{skin}	ΔTrend (K/decade)	No. of grids	ΔT_{skin}	ΔTrend (K/decade)	No. of grids	ΔT_{skin}	ΔTrend (K/decade)
≥ 1.0	2155	2.01	-0.10	95	-	-	425	2.01	-0.96	378	1.37	0.03
≥ 1.5	1506	2.34	0.04	19	-	-	253	2.25	-0.66	104	1.80	0.24
≥ 2.0	940	2.71	0.19	5	-	-	134	2.69	-0.21	22	-	-
≤ -1.0	1839	-1.59	-0.45	236	-2.25	-0.34	224	-1.71	-0.43	877	-2.19	-0.37
≤ -1.5	921	-1.94	-0.45	162	-2.72	-0.47	115	-2.18	-0.16	654	-2.52	-0.55
≤ -2.0	367	-2.27	-0.60	109	-3.20	-0.78	55	-	-	472	-2.82	-0.69
≥ 1.0	912	2.15	1.21	40	-	-	139	2.09	1.36	179	1.40	1.01
≥ 1.5	707	2.41	1.22	8	-	-	94	-	-	51	-	-
≥ 2.0	499	2.70	1.22	1	-	-	64	-	-	15	-	-
≤ -1.0	1309	-1.59	-0.90	126	-2.51	-2.02	122	-1.69	-2.42	500	-2.26	-1.96
≤ -1.5	643	-1.96	-0.92	89	-	-	61	-	-	387	-2.55	-2.06
≤ -2.0	272	-2.28	-1.02	69	-	-	27	-	-	293	-2.81	-2.06

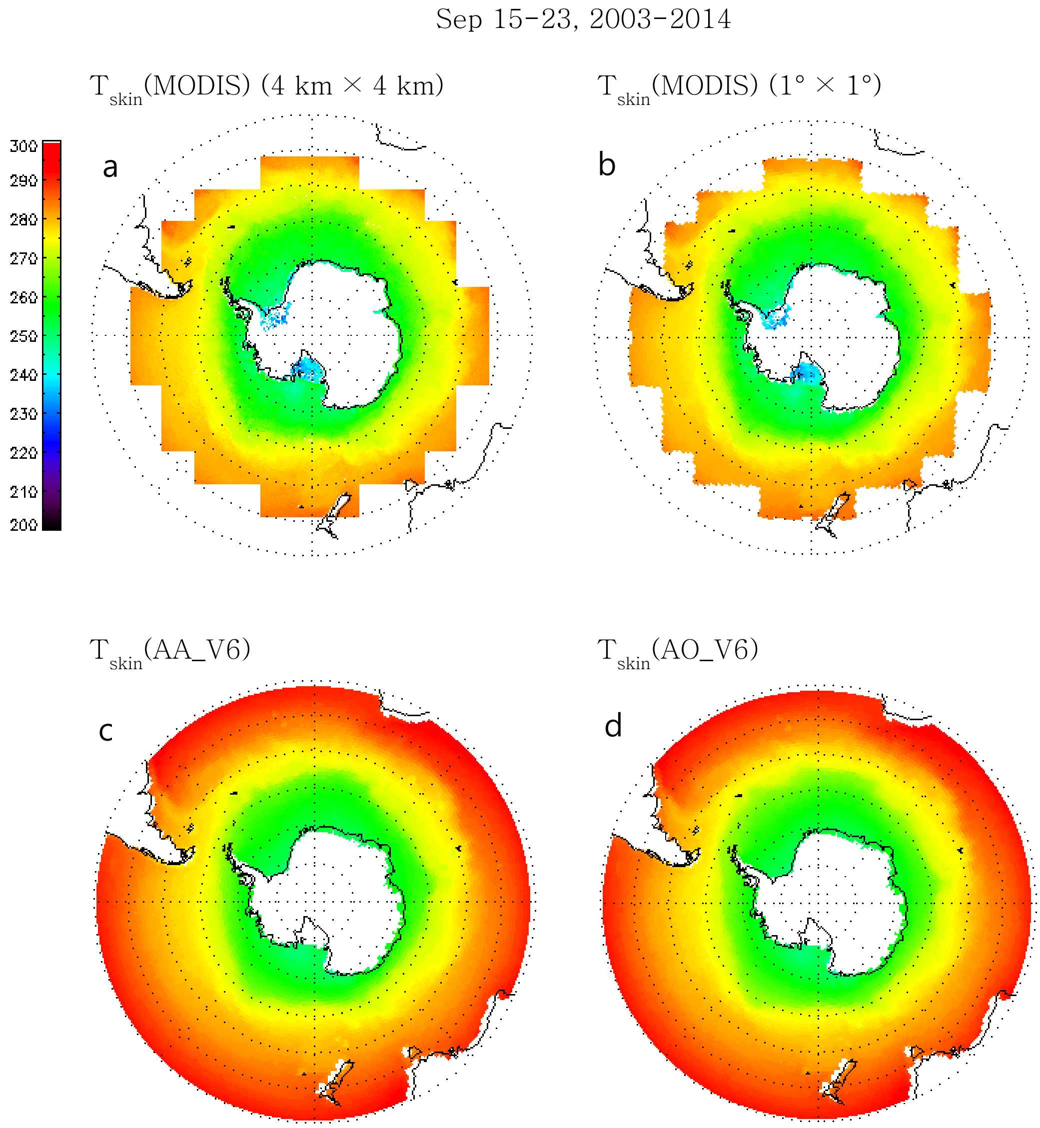

Fig. 1. (a) 12-year composite skin temperatures (K) of the MODIS IST over the southern hemisphere during September 15-23 of 2003-2014. The original MODIS data (MYD29E1D) have a 4 km × 4 km spatial resolution. Their spatial resolution has been reconstructed to 1° × 1° in Fig. 1b in order to compare this data with the AIRS/AMSU data. Figure 1c-d is the surface skin temperatures of the AIRS/AMSU and AIRS only over the southern hemisphere ocean during September 15-23 of 2003-2014, respectively.

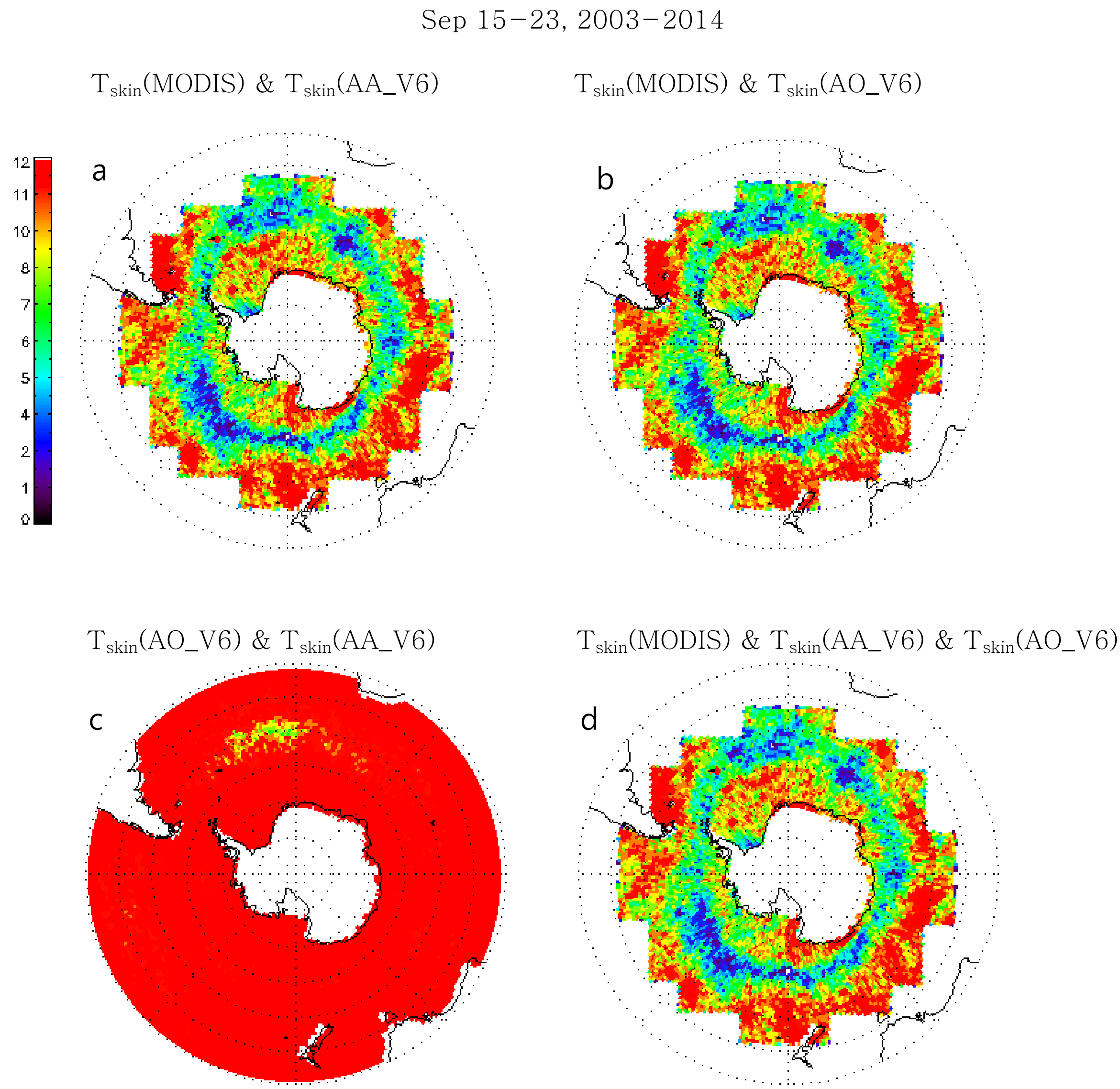

Fig. 2. The number of co-located observations of (a) T_{skin} (MODIS) and T_{skin} (AA_V6), (b) T_{skin} (MODIS) and T_{skin} (AO_V6), and (c) T_{skin} (AO_V6) and T_{skin} (AA_V6) over the southern hemisphere during September 15-23 of 2003-2014. (d) Same as in Fig. 2c except for three different datasets (T_{skin} (MODIS), T_{skin} (AA_V6), and T_{skin} (AO_V6)).

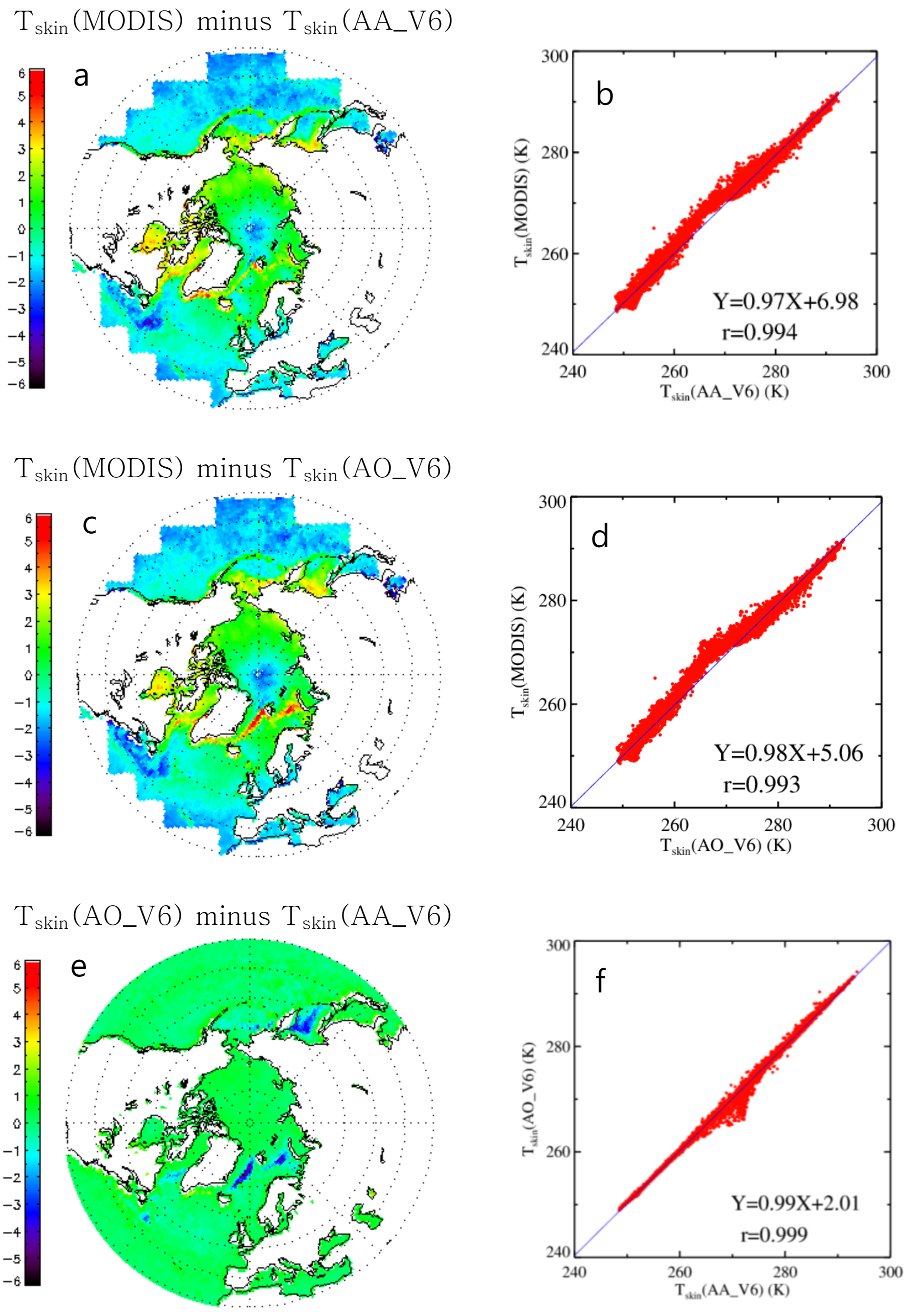

Fig. 3. The distributions of (a) T_{skin} (MODIS) minus T_{skin} (AA_V6), (c) T_{skin} (MODIS) minus T_{skin} (AO_V6), and (e) T_{skin} (AO_V6) minus T_{skin} (AA_V6) over the northern hemisphere during April 16-24 of 2003-2014. The scatter plots of (b) T_{skin} (MODIS) versus T_{skin} (AA_V6), (d) T_{skin} (MODIS) versus T_{skin} (AO_V6), and (f) T_{skin} (AO_V6) versus T_{skin} (AA_V6).

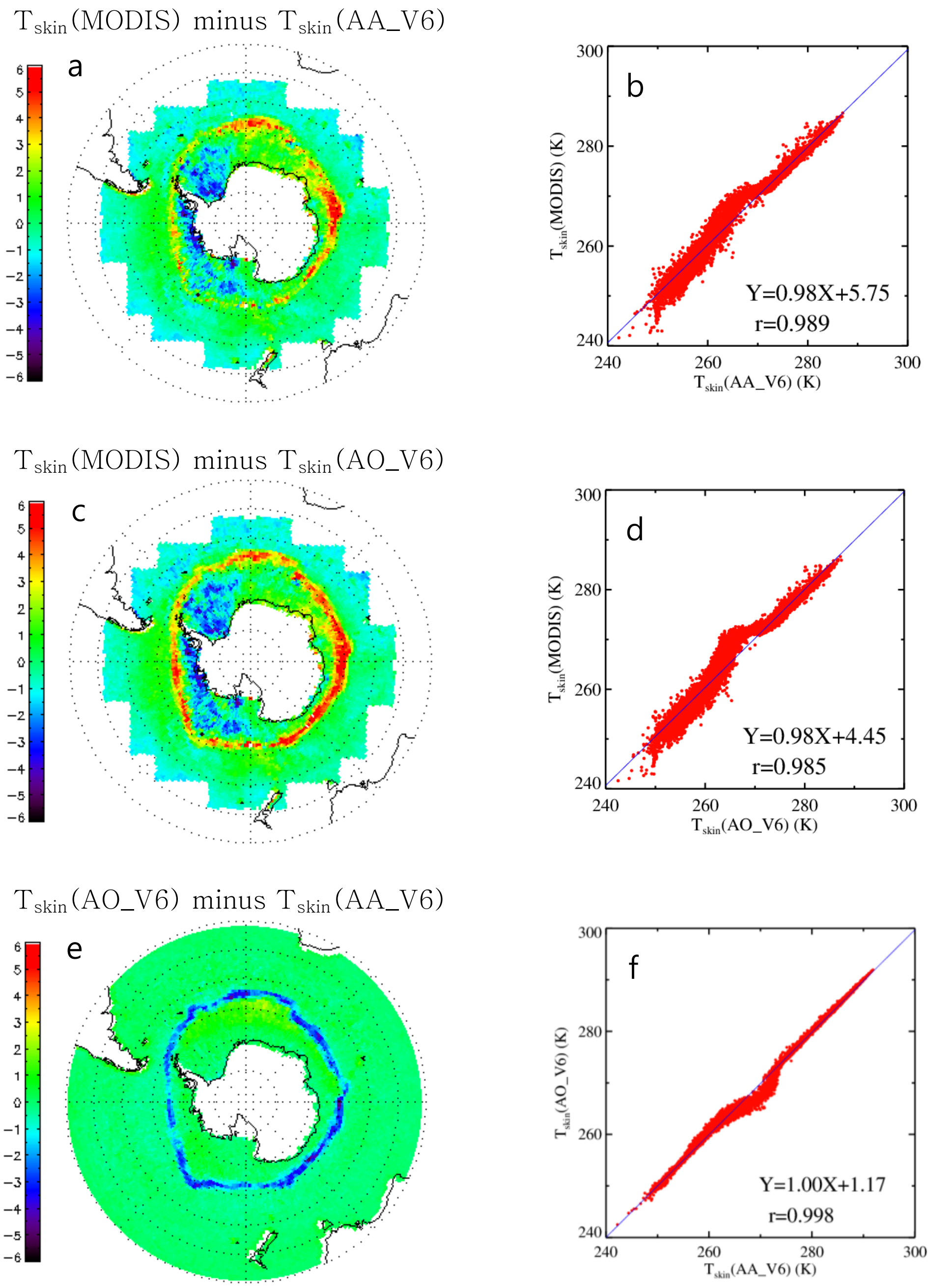

Fig. 4. Same as in Fig. 3 except for the data taken during September 15-23 of 2003-2014, over the southern hemisphere.

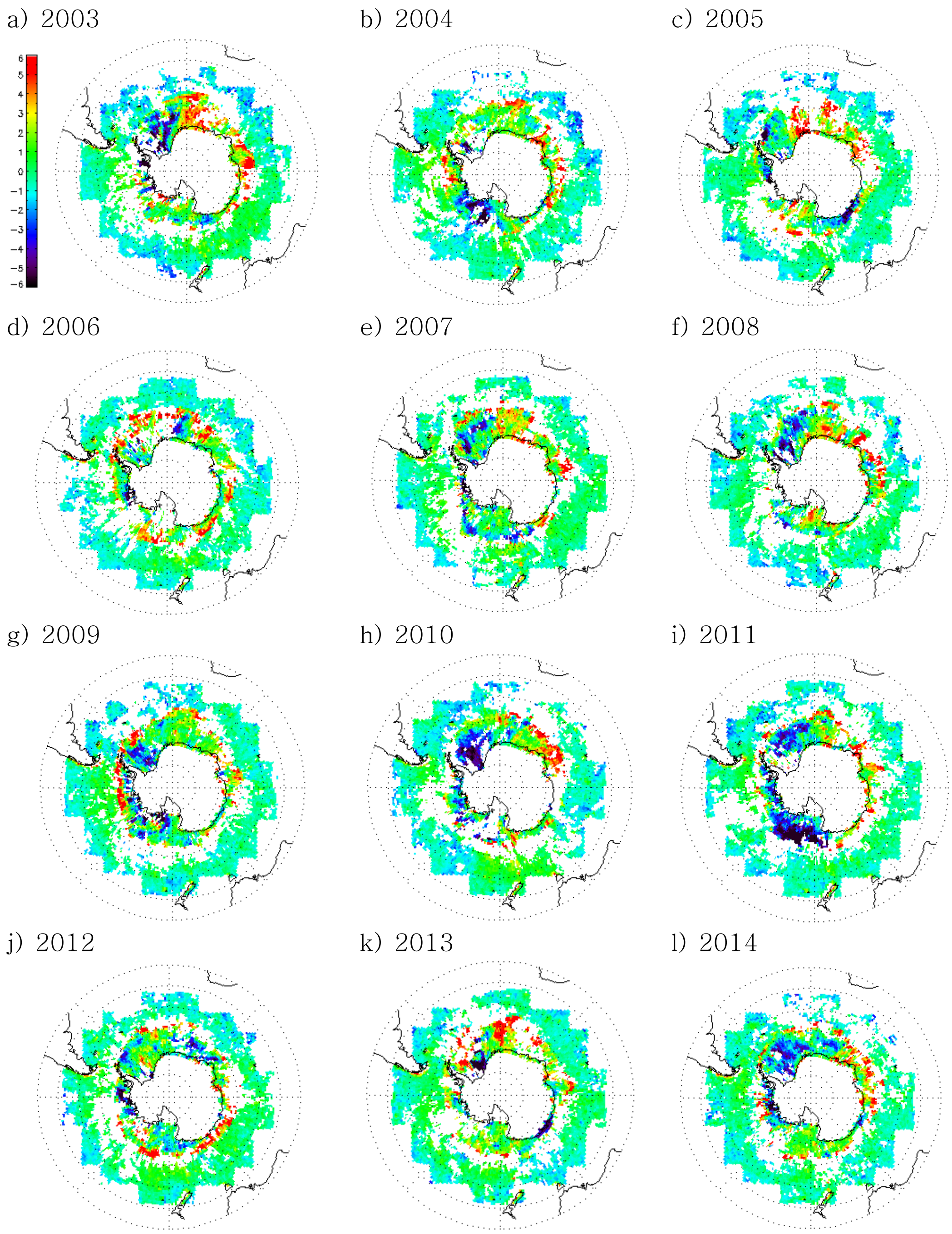

Fig. 5. Annual-average spatial distributions of the T_{skin} (MODIS) minus T_{skin} (AA_V6) over the southern hemisphere during September 15-23.

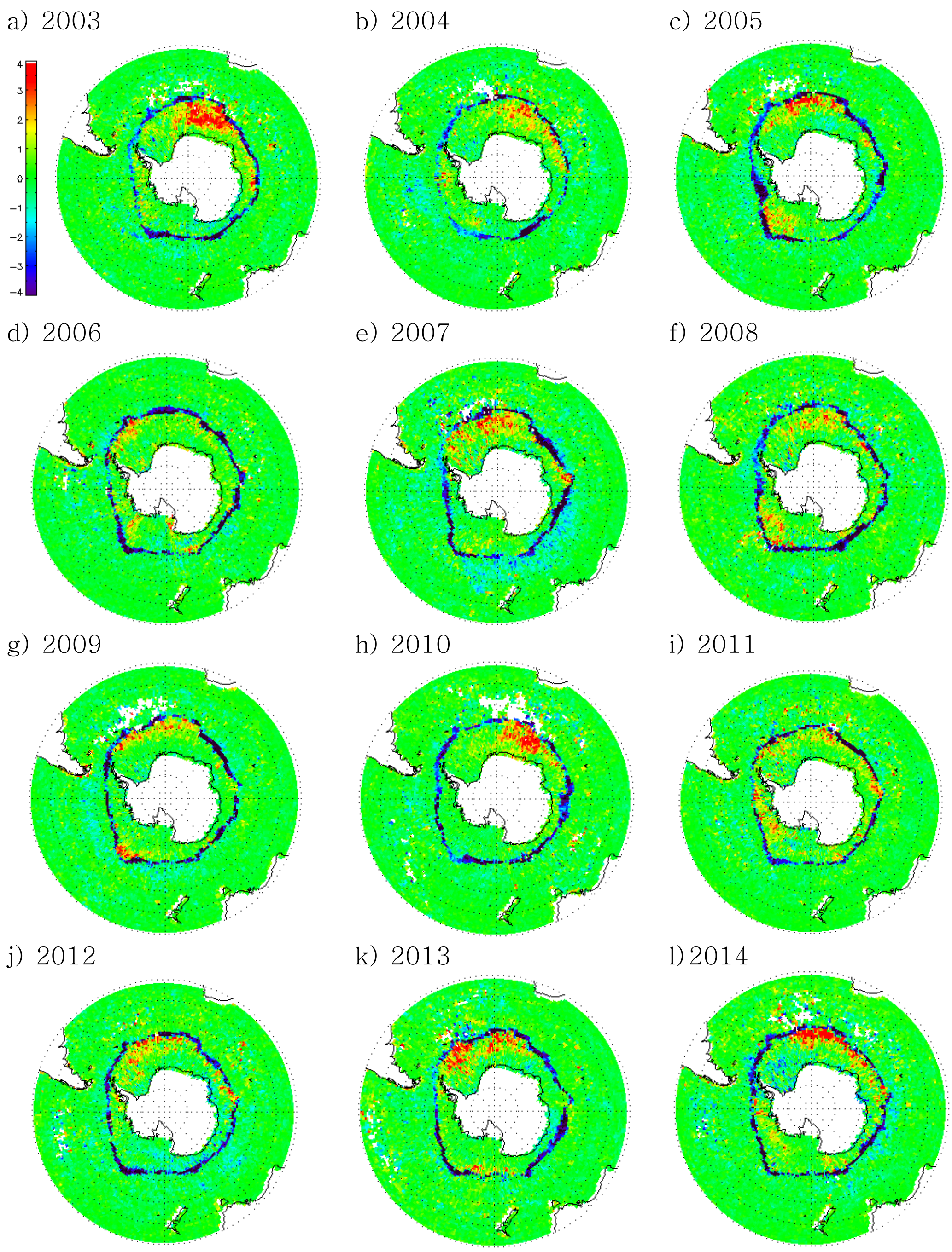

Fig. 6. Same as Fig. 5 except for T_{skin} (AO_V6) minus T_{skin} (AA_V6).

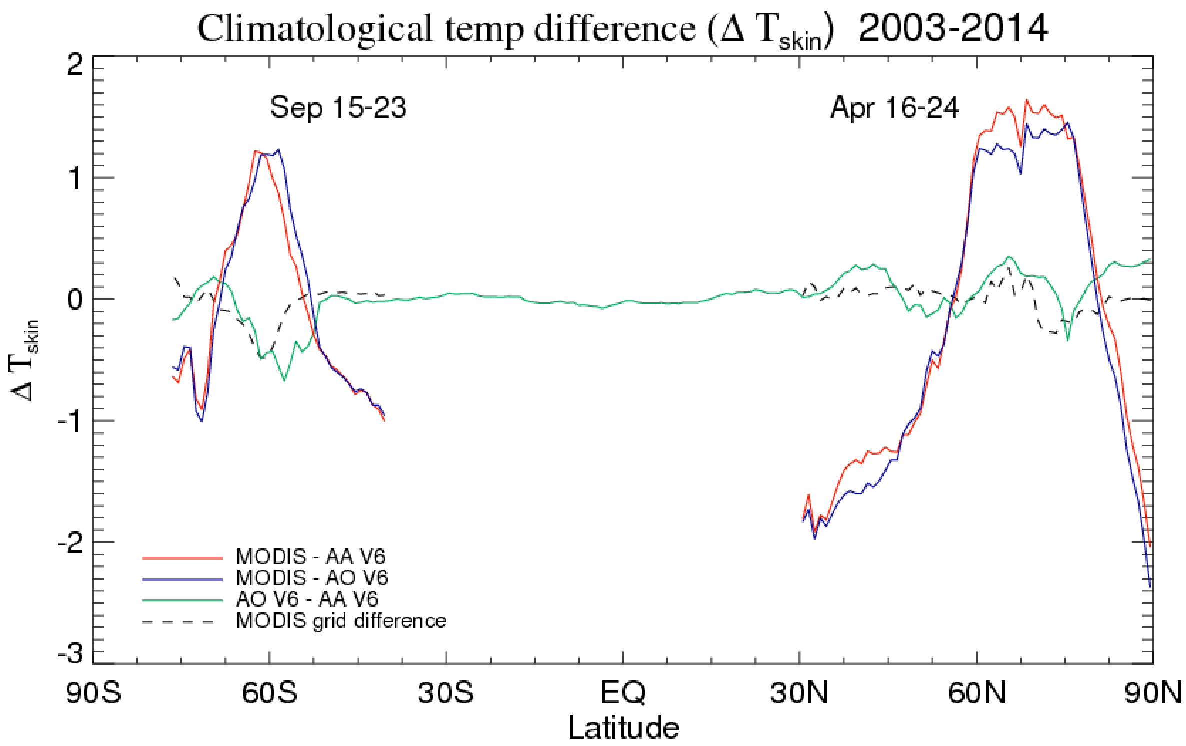

Fig. 7. Zonal averaged values of T_{skin} (MODIS) minus T_{skin} (AA_V6) (red solid line), T_{skin} (MODIS) minus T_{skin} (AO_V6) (blue solid line), and T_{skin} (AO_V6) minus T_{skin} (AA_V6) (green solid line). The difference in spatial grid averages of the MODIS data between 4 km by 4 km and 1° by 1° is shown by the black dashed line. The difference values are calculated at one degree interval along each latitudinal belt. The climatological data periods are April 16-24, 2003-2014 over the northern hemisphere, and September 15-23, 2003-2014 over the southern hemisphere.

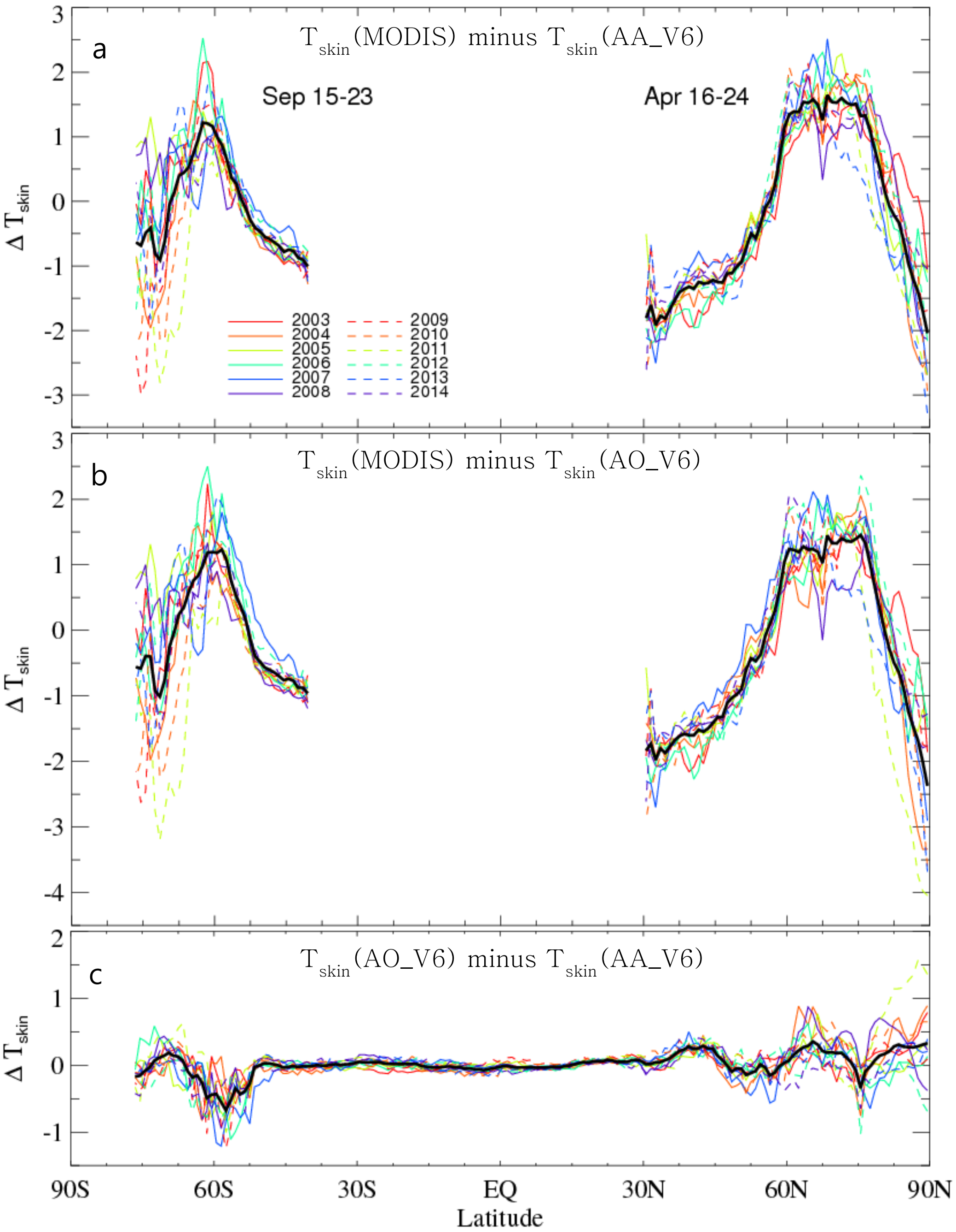

Fig. 8. Zonal averaged values of (a) T_{skin} (MODIS) minus T_{skin} (AA_V6), (b) T_{skin} (MODIS) minus T_{skin} (AO_V6), (c) T_{skin} (AO_V6) minus T_{skin} (AA_V6) over the northern hemisphere from April 16 to 24, 2003-2014, and over the southern hemisphere from September 15 to 23, 2003-2014. The values in each year represent the corresponding color lines. The thick black line indicates the mean difference values.

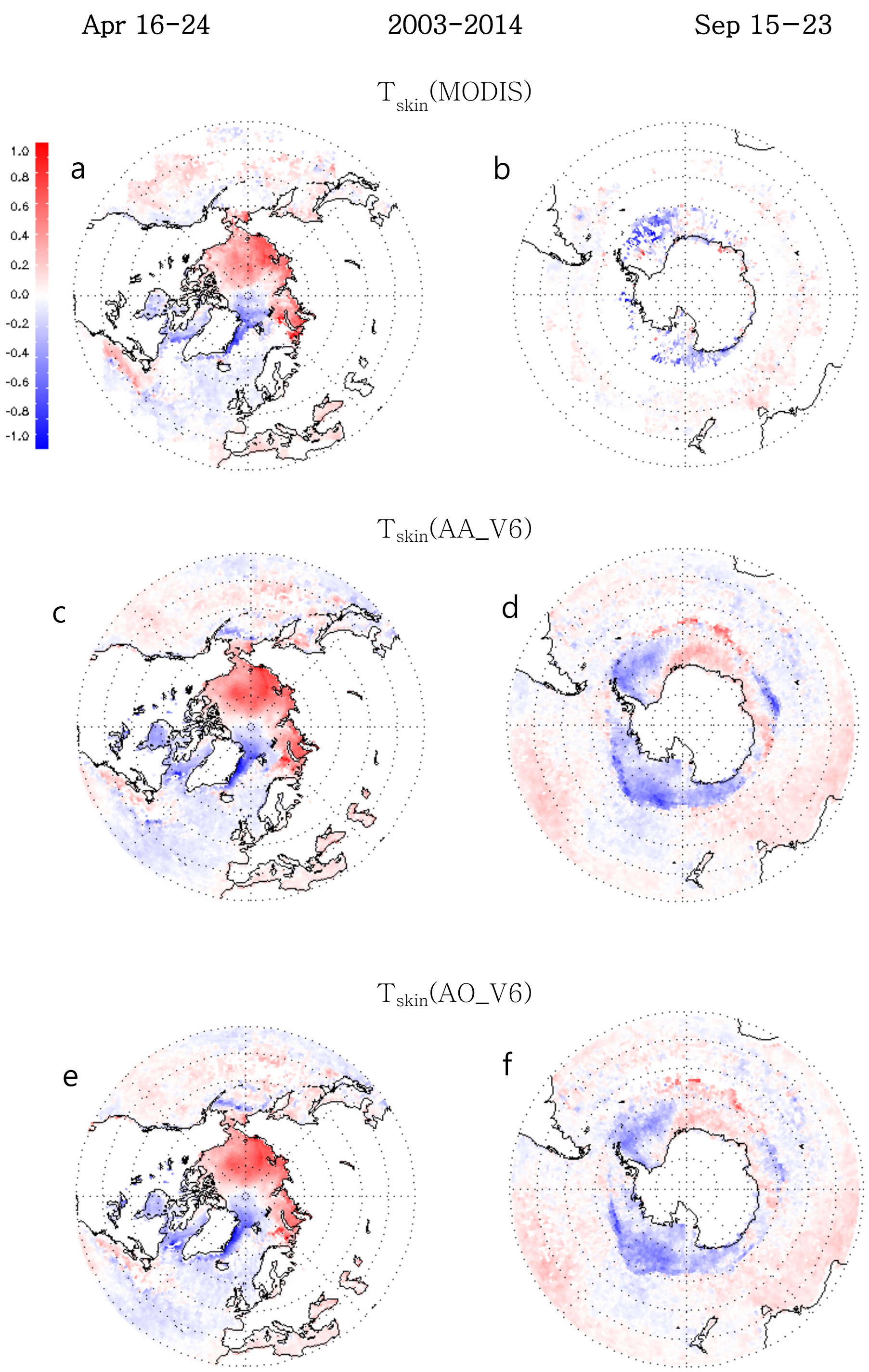

Fig. 9. Satellite-derived 9-day anomaly trends (K yr^{-1}) in a grid box of 1°× 1° over the northern hemisphere during April 16-24 of 2003-2014, for the (a) T_{skin} (MODIS), (c) T_{skin} (AA_V6), and (e) T_{skin} (AO_V6), and over the southern hemisphere during September 16-24 of 2003-2014, for the (b) T_{skin} (MODIS), (d) T_{skin} (AA_V6), and (f) T_{skin} (AO_V6).

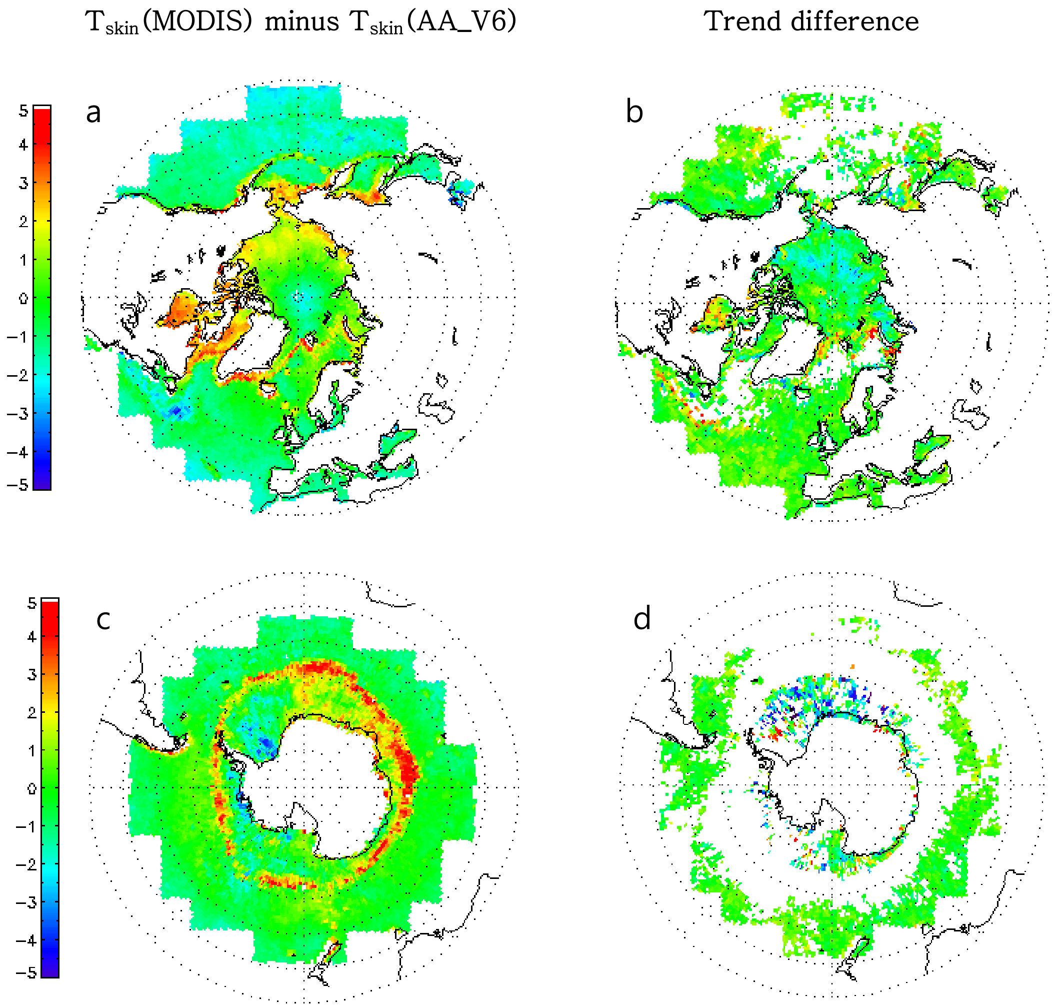

Fig. 10. (a) 12-year mean of T_{skin} (MODIS) minus T_{skin} (AA_V6) (K) over the northern hemisphere during April 16-24 of 2003-2014, and (b) difference in the thermal trend (K/decade) between T_{skin} (MODIS) and T_{skin} (AA_V6). Figures 10c-d are the same as Figs.10a-b except for over the southern hemisphere during September16-24 of 2003-2014, respectively. Figure 10a is the same as Fig. 3a in Kang and Yoo (2015).

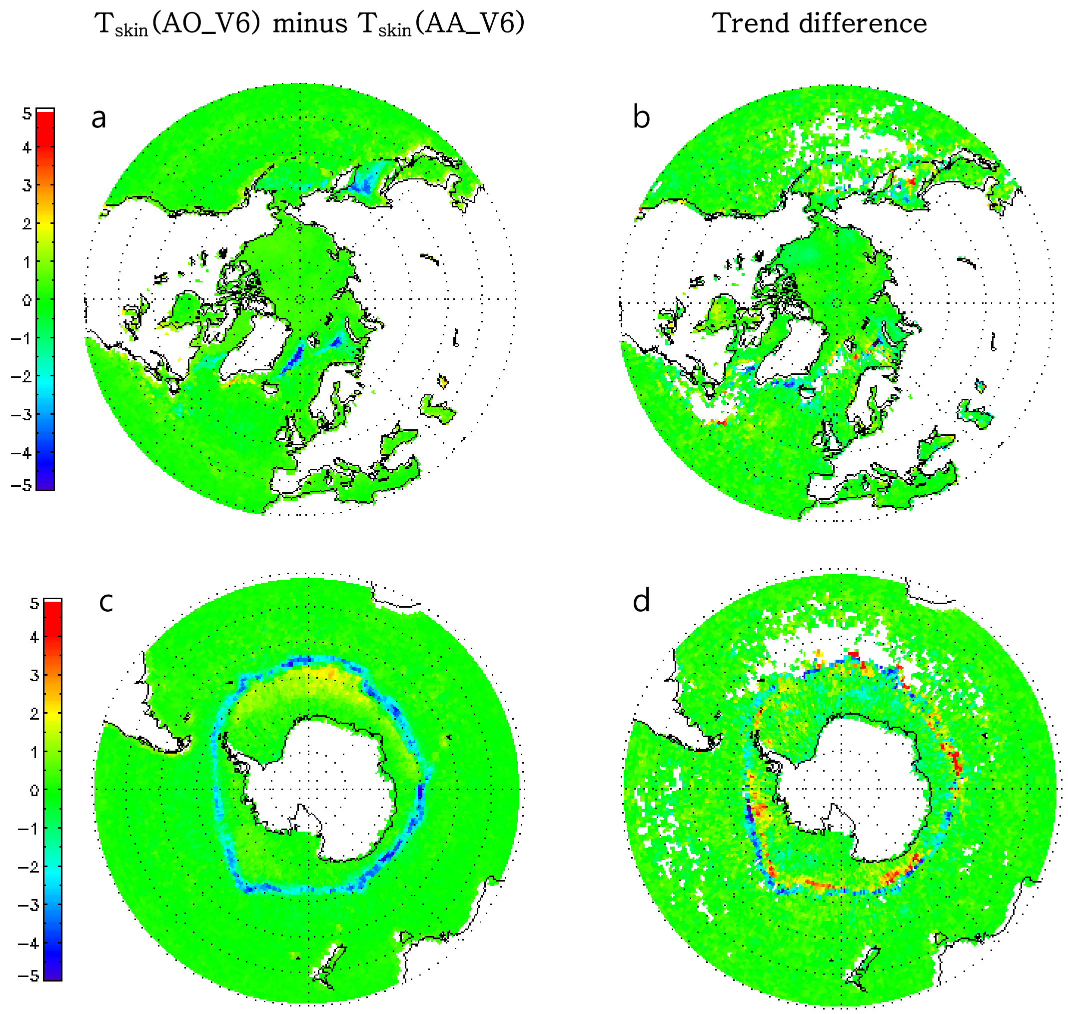

Fig. 11. Same as Fig. 10 except for T_{skin}(AO_V6) minus T_{skin}(AA_V6). Figure 11a is the same as Fig. 3c in Kang and Yoo (2015).

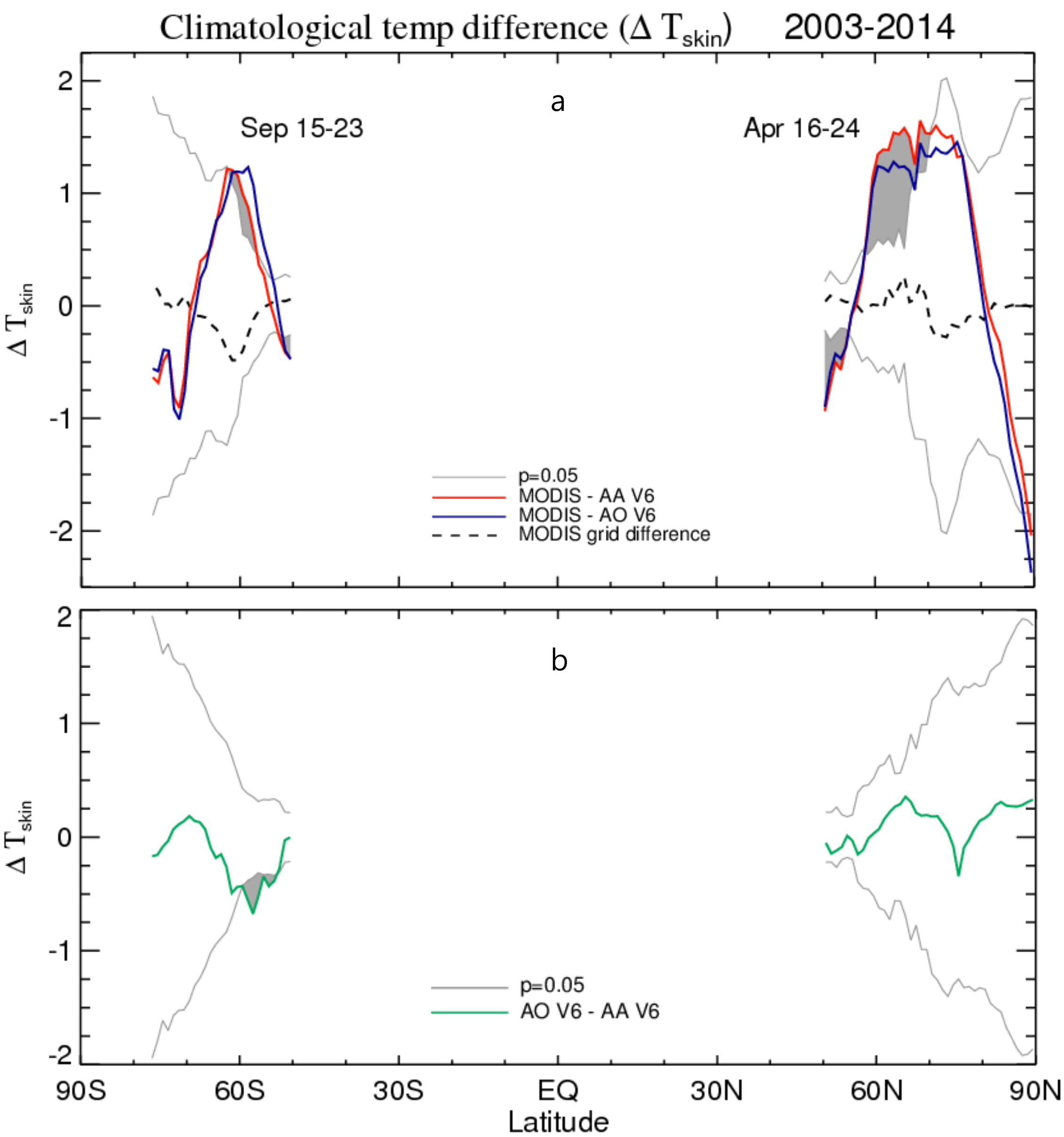

Fig. A1. The difference values (a) between T_{skin} (MODIS) and T_{skin} (AA_V6), and (b) T_{skin} (AA_V6) and T_{skin} (AO_V6) over a possibly frozen region; shown in Fig. 7. The 5% significance level is presented as grey solid lines, and the shaded areas are statistically significant at the 0.05 level.