

Intercomparison of total column ozone data from the Pandora spectrophotometer with Dobson, Brewer, and OMI measurements over Seoul, Korea

Jiyoung Kim¹, Jhoon Kim¹, Hi-Ku Cho¹, Jay Herman², Sang Seo Park^{1,3}, HyunKwang Lim¹, Jae-hwan Kim⁴ and Koji.Miyagawa^{5,6}

¹ Department of Atmospheric Sciences, Yonsei University, Seoul South Korea

² Joint Center for Earth Systems and Technology, University of Maryland, Baltimore County, UMBC-JCET and NASA Goddard Space Flight Center, Greenbelt, MD 20771, USA

³ Research Institute for Applied Mechanics, Kyushu University, Fukuoka, Japan

⁴ Department of Atmospheric Science, Pusan National University, Busan, Korea

⁵ Science & Technology Corp, Hampton, VA 23666-1393, United States

⁶ NOAA ESRL Global Monitoring Division

Correspondence to: Jiyoung Kim (jigang615@yonsei.ac.kr)

1 **Abstract**

2 Daily total column ozone (TCO) measured using the Pandora spectrophotometer (#19) was
3 intercompared with data from the Dobson (#124) and Brewer (#148) spectrophotometers, as
4 well as from the Ozone Monitoring Instrument (OMI), over the 2-year period between March
5 2012 and March 2014 at Yonsei University, Seoul, Korea. The Pandora TCO measurements
6 are closely correlated with those from the Dobson, Brewer, and OMI instruments with
7 regression coefficients (slopes) of 0.95, 1.00, 0.98 (OMI-TOMS), and 0.97 (OMI-DOAS),
8 respectively, and determination coefficients (R^2) of 0.95, 0.97, 0.96 (OMI-TOMS), and 0.95
9 (OMI-DOAS), respectively. In particular, they show a close agreement with the Brewer TCO
10 measurements. The difference between the Pandora and Dobson data can be explained by
11 smaller amount of Dobson data available to calculate the daily averages, observation times,
12 solar zenith angles, SO_2 effect, temperature, and humidity. The difference in the results
13 obtained from the Pandora instrument and Ozone Monitoring Instrument-Differential Optical
14 Absorption Spectroscopy (OMI-DOAS algorithm) can be explained by the dependence on
15 seasonal variations of about $\pm 2\%$ and solar zenith angle leading to overestimation by 5% of
16 OMI-DOAS measurements. For the Dobson measurements in particular, the difference
17 caused by the inconsistency in observation times when compared with the Pandora
18 measurements was up to 12.5% on 22 June 2013 because of diurnal variations in the TCO
19 values. However, despite these various differences and discrepancies, the daily TCO values
20 measured by the Pandora during the 2-year study period can be considered to be accurate
21 since they are closely correlated with measurements from other instruments as mentioned
22 above.

23

24 **1. Introduction**

25 Approximately 90% of total column ozone (TCO) is found in the stratosphere, and its density
26 peak occurs at altitudes between 20 and 30 km (Liou, 2002; Schott, 2007). This layer is
27 essential to the biosphere as it absorbs harmful solar ultraviolet (UV) radiation. In addition,
28 UV absorption by ozone influences global radiative forcing and climate change over long
29 timescales (e.g., Cho et al., 2003; Martens, 1998; WMO, 2014). After significant depletion of
30 the ozone layer was detected in the 1980s (Farman et al., 1985; Chubachi, 1985), many
31 related studies have conducted (e.g., Harris et al., 1997; Solomon, 1999; Newchurch et al.,
32 2003; Weatherhead et al., 2000). These studies found that the concentration of anthropogenic
33 ozone-depleting substances (ODSs) had decreased and that, consequently, global ozone
34 amounts would return to 1980 levels during the 21st century (WMO, 2014).

35 Over recent decades, ground-based instruments such as Dobson or Brewer
36 spectrophotometers have been used (and improved) to obtain stable and highly accurate
37 measurements of global ozone amounts. The Dobson spectrophotometer was developed in
38 1928 by G. M. B. Dobson to measure TCO levels under clear-sky conditions (Dobson, 1968).
39 TCO values are retrieved by measuring direct or scattered intensity ratios at two wavelength
40 pairs that have different absorption features in the UV band (A-pair: 305.5 and 325.4 nm; D-
41 pair: 317.6 and 339.8 nm, recommended by WMO; Komhyr et al., 1980; LEONARD, 1989).
42 Since the 1970s, many instruments have been installed and a global network established to
43 monitor TCO amounts and its vertical profile using Umkehr measurements.

44 The Brewer spectrophotometer was developed in the early 1980s and since commencing
45 operational use has supplemented measurements made by Dobson spectrophotometers

46 (Brewer, 1973; Kerr et al., 1985; Kerr, 2010). The measurement principle is similar to that of
47 the Dobson instrument; however, the single monochromator Brewer spectrophotometer
48 retrieves data on total UV (TUV), erythemal UV (EUV), TCO, and aerosols, as well as trace
49 gases such as NO₂ and SO₂, by measuring solar irradiance and zenith sky radiances. The
50 accuracy of well calibrated Brewer measurements is estimated to be about 1% for direct-sun
51 observations. Nearly 200 Brewer instruments are now operating in about 40 countries (Kerr,
52 2010), and the MK-IV version has been operating at Yonsei University, Seoul, Korea since
53 1997 (Kim et al., 2014). This instrument enables measurement of global UV spectral
54 irradiance, and this is used for retrieval of TUV and EUV from 290 to 363 nm with a spectral
55 resolution of 0.5 nm on a horizontal surface (cf. Sabburg et al., 2002). It also measures
56 normal direct UV radiation, which can be used to retrieve gas concentrations at five
57 wavelengths in the UV region (306.3, 310.1, 313.5, 316.7, and 320.0 nm; e.g., Kerr et al.,
58 1985; Kerr, 2002; Kim et al., 2011). In addition, satellite-based observations, such as Total
59 Ozone Mapping Spectrometer (TOMS) and OMI, have also been conducted since 1979
60 (Bhartia and Wellemeyer, 2002; Levelt et al., 2006) and have generated an extensive and
61 highly accurate global dataset. These data have been validated globally and over long periods,
62 and have been widely used in numerous studies (Balis et al., 2007; McPeters et al., 2008;
63 WMO, 2014).

64 The Pandora spectrophotometer was developed at NASA's Goddard Space Flight Center in
65 2006 to measure the concentrations of trace gases including ozone (Herman et al., 2009; Cede,
66 2011). It consists of a head sensor with fore-optics, mounted on a high-precision sun-tracker
67 and sky-scanner (ca. 1.6° field of view and ca. 0.01° pointing precision), and it measures the
68 direct solar beam in the spectral range between 280 and 500 nm using the Sun-only CMOS

69 detector, and 280–525 nm using the Sun-and-Sky CCD detector with absolute O₃ retrieval
70 errors of about 1% (± 3 DU) and a high precision of ± 0.1 DU (Herman et al., 2015; Reed et al.,
71 2015; Tzortziou et al., 2012). Absolute NO₂ retrieval errors are about ± 0.1 DU (Herman et al.,
72 2009). From the measured radiance, TCO levels, together with the total column of trace gases
73 (including NO₂, SO₂, BrO, water vapor, and formaldehyde), are retrieved using the
74 differential optical absorption spectroscopy (DOAS) technique (Wang et al., 2010; Yun et al.,
75 2013).

76 In this study, we intercompare the Pandora measurements from Seoul with two ground-based
77 and two satellite datasets over a 2-year period. Furthermore, the difference between Pandora
78 and the other measurements, and the causes of these differences, are discussed. The
79 remainder of this paper is organized as follows. Section 2 describes the ground-based and
80 satellite datasets used in this study. Section 3 describes the methodology and results of the
81 intercomparison together with our analysis and discussion. In addition, high-resolution
82 diurnal variations in the Pandora TCO data are compared with Dobson measurements. Finally,
83 our conclusions are summarized in Sect. 4.

84

85

86

87

88

89

90 **2. Data and Analysis**

91 In this study, the TCO data used for intercomparisons were measured using Pandora, Dobson,
92 and Brewer spectrophotometers from March 2012 to March 2014 at Yonsei University
93 (37.57°N, 126.95°E; 84 m above sea level) in Seoul, Korea. The university is one of the
94 WMO Global Ozone Observing System (GO3OS) stations (Station No. 252). OMI has also
95 recorded TCO data over this site since 2004. As part of the ongoing national monitoring
96 program of the Korea Meteorological Administration (KMA), TCO measurements have been
97 made at this station since 1984. The calibration history and characteristics of Dobson (Beck
98 #124), Brewer (SCI-TEC #148), and OMI instruments are described in Sect. 2.1 to 2.4.

99

100 **2.1. Dobson Spectrophotometer (Beck #124)**

101 The Dobson spectrophotometer (Beck #124) is located on the rooftop of the Science Hall of
102 Yonsei University and has been in operation since 1984, with regular calibration as a standard
103 for total ozone measurements (Cho, 1989, 1996; Cho et al., 2003; Kim et al., 2005). The
104 instrument retrieves TCO from the observed UV radiance in direct-sun and zenith-sky modes
105 three times a day. A direct-sun TCO value measured at noon under clear skies is generally
106 selected as a representative value; however, a value close to noon or the zenith-sky
107 measurement can be used instead if data from noon are unavailable. After the automation of
108 the Dobson instrument (in particular, Q-levers, Attenuator, R-dial, observation, and data
109 processing with test) in 2006, accuracy was improved such that the proportion of data points
110 within a $\pm 3\%$ error range increased from 92% to 98% (Kim et al., 2007; Miyagawa et al.,
111 2005). The calibration history of this instrument has been summarized by Kim et al. (2007)

112 and Hong et al. (2014). The Dobson instrument has provided a high-quality, objective, and
113 reliable dataset that can be used to monitor the variations and trends in ozone levels over the
114 Korean Peninsula. According to previous studies that have used this dataset, the annual mean
115 ozone level increased by $7.2\% \text{ decade}^{-1}$ from 2004 to 2010 (Kim et al., 2014). This recent
116 increase is contrary to the result reported in Kim et al. (2005) showing past decreasing trend
117 of $-2.39\% \text{ decade}^{-1}$ from 1979 to 1991 measured by satellite TOMS, in spite of slight
118 increasing trend of $0.75\% \text{ decade}^{-1}$ from 1992 to 2004. As for the Dobson instrument, the
119 calibrations are relevant for the comparison period in this result.

120

121 **2.2. Brewer Spectrophotometer (SCI-TEC #148)**

122 The Brewer MK-IV spectrophotometer (SCI-TEC #148) at the Dobson measurement site,
123 which has been in operation since 1997 (Kim et al., 2011), automatically measures TCO,
124 trace gases, and UV irradiance, and is regularly calibrated (Kim et al., 2014). Previous studies
125 based on Brewer spectrophotometer data have shown that annual EUV and TUV from 2004
126 to 2010 decreased by $0.83\% \text{ decade}^{-1}$ and $0.90\% \text{ decade}^{-1}$, respectively (Kim et al., 2011),
127 whereas the Aerosol Optical Depth (AOD) at 320 nm increased by $22.4\% \text{ decade}^{-1}$ (Kim et
128 al., 2014). And Hong and Cho (2007) showed the annual mean variation of the daily total
129 ozone amount showing a clear seasonal variation in Pohang from 1994 to 2005 using Brewer
130 spectrophotometer.

131

132

133 **2.3. Ozone Monitoring Instrument (OMI)**

134 The OMI onboard the Aura satellite has been dedicated to monitoring ozone and trace gases
135 for air quality and climate studies since 2004. This instrument detects the molecular
136 absorption of backscattered solar light in the visible and UV wavelengths (270–500 nm) with
137 a spatial resolution of 13×24 km at nadir (Buchard et al., 2008; Levelt et al., 2006). The total
138 ozone product from the OMI is calculated using two different algorithms: the TOMS
139 algorithm and the DOAS algorithm, which show fairly good agreement, with correlation
140 coefficients ranging from 0.89 to 0.99 (Antón et al., 2009; Kroon et al., 2008; McPeters et al.,
141 2008). The TOMS algorithm uses two wavelengths: a weak absorption wavelength (331.2 nm)
142 and a strong absorption wavelength (317.5 nm) to derive TCO. The derivation applies tables
143 calculated by the TOMS forward model (TOMRAD), which is based on successive iterations
144 of the characteristic equation in the theory of radiative transfer developed by Dave (1964)
145 (Bhartia and Wellemeyer, 2002; McPeters et al., 2008). The DOAS algorithm derives TCO
146 using the DOAS method, and the derivation follows three steps. First, spectral fitting is
147 performed (Platt and Stutz et al., 2008; Veefkind et al., 2006) to obtain the slant column
148 density (SCD); second, the SCD is converted to the vertical column density (VCD) using the
149 air mass factor (AMF). The final step in the derivation procedure is a correction for the cloud
150 effect (e.g., Hong et al., 2014; McPeters et al., 2008). For the level-3 daily product used in
151 this study, this step consists of gridding and averaging the level-2 TCO orbital swath data
152 onto the $0.25^\circ \times 0.25^\circ$ global grids (after a quality check).

153

154

155 **2.4. Pandora Spectrophotometer (#19)**

156 The Pandora spectrophotometer (#19) was installed at Yonsei University as part of the
157 Distributed Regional Aerosol Gridded Observation Networks (DRAGON)-NE Asia
158 Campaign (http://aeronet.gsfc.nasa.gov/new_web/DRAGON-Asia_2012_Japan_South_Korea.html) and has been used operationally since March 2012 to
159 monitor NO₂ and O₃ concentrations. Pandora is a passive system that measures direct sunlight
160 from 280 to 525 nm with a spectral resolution of 0.6 nm, and uses a UV sensitive CCD
161 detector of 2048 × 16 pixels. Two UV band-pass filters, BP300 (280–320 nm) and U340
162 (280–380 nm), are used to correct for the stray light effect. The temporal resolution of the
163 Pandora measurement is about 2 minutes (Tzortziou et al., 2012; Yun et al., 2013), and it was
164 installed at the same geolocation as the Dobson and Brewer instruments.
165

166

167

168

169

170

171

172

173

174

175 **3. Results**

176 The daily TCO datasets were calculated using the following methods. Real-time Pandora and
177 Brewer data obtained from the direct-sun measurements were averaged to obtain a single
178 representative daily value. For the Pandora data in particular, to avoid errors associated with
179 cloud contamination and stray light effects, the data were selected using the following criteria:
180 root mean square (RMS) of weighted spectral fitting residuals < 0.05 , solar zenith angle $< 75^\circ$,
181 and bias of ozone amount < 2 DU, as suggested in previous studies (Herman et al., 2015;
182 Reed et al., 2015; Tzortziou et al., 2012). For the Dobson spectrophotometer, daily values
183 measured in direct-sun mode under a clear sky were averaged for the comparison. Finally, for
184 the OMI instruments, the daily level-3 gridded data, the OMTO3e from the TOMS-like
185 algorithm, and the OMDOAO3e from the DOAS technique, together with site information
186 from Yonsei University, were spatially interpolated.

187

188 **3.1. TCO measured by the Pandora, Dobson, Brewer, and OMI instruments**

189 The time series of measurements from the four instruments are shown in Fig. 1 for
190 comparison, and all fall within the range of 230–500 DU and show obvious and similar
191 seasonal variations. These seasonal variations are caused by changes to the Brewer–Dobson
192 circulation in the Northern Hemisphere (Brewer, 1949; Wang et al., 2015; Weber et al., 2003).
193 Additionally, Fig. 2 shows similar annual cycles with an amplitude of about 0.15% of the
194 average values for the four different instruments. Maximum and minimum values of 2-year
195 averaged monthly TCO and annual ranges are also recorded in this figure. All graphs and
196 statistics were derived from Fig. 1 under the condition that the valid number of daily

197 observations was greater than 10 days per month. In this figure, the largest maximum
198 monthly mean TCO values are from the Dobson measurements (i.e., 371.5 DU in April) and
199 the smallest minimum monthly mean TCO values were from the Pandora measurements (i.e.,
200 268.9 DU in October). In addition, the largest annual range is seen in the Dobson
201 measurements (101.7 DU), whereas the smallest range is seen in the Brewer measurements
202 (81.3 DU). In addition, the maximum monthly mean TCO value of the Pandora
203 measurements shows the smallest relative difference with that of the Brewer measurements of
204 1.54%. And the minimum monthly mean TCO value of the Pandora measurements shows the
205 smallest relative difference with that of the Dobson measurements of 0.37%. The OMI-
206 DOAS measurements also showed the smallest difference in minimum value from that of the
207 Pandora of 0.37%. Next, we consider the daily data (Table 1).

208 Table 1 shows the average, standard deviation, and maximum and minimum values of the
209 daily TCO data measured by the four instruments, together with the relative differences
210 between the Pandora data and the Dobson, Brewer, and OMI data. The largest maximum and
211 smallest minimum TCO values were 467.1 DU on 10 April 2013 and 238.3 DU on 8 October
212 2013, respectively, measured by OMI-TOMS. For the 2-year average TCO value, the Dobson
213 value was the largest at 331.9 DU and with a standard deviation of 38.6. In contrast, the
214 Pandora measurement showed the smallest two-year average value of 317.2 DU, with a
215 standard deviation of 36.8 and a maximum of 436.7 on 6 April 2012, and the smallest
216 minimum value of 249.2 DU on 7 October 2013. The histograms of all daily TCO data (Fig.
217 3) show a bell-shaped curve, similar to a Gaussian distribution, and suggest that the 2-year
218 average values of the Pandora, Dobson, Brewer, and OMI instruments in Table 1 are a
219 reliable representation of the annual mean TCO value over the 2-year period for each

220 instrument. The annual mean TCO values from 2012 to 2014 are largest for the Dobson
221 instrument (331.9 ± 38.6 DU) and smallest for the Pandora instrument (317.2 ± 36.8 DU). The
222 annual and monthly mean TCO values over this period are similar to the result of the past
223 period, which the annual and monthly TCO values are measured by Dobson instrument from
224 1985 to 2000 (within $\sim 2\%$), as presented by Cho (2003). However, slight decreases in the
225 annual and monthly means of the TCO values are seen over our study period (2012–2014)
226 when compared with the earlier period (1985–2000) for Pandora (decreases of -1.49% and $-$
227 0.54% , respectively); the Dobson, Brewer, and OMI show slight increases of $\sim 1\%$. Statistical
228 details are presented below.

229

230 **3.2. Intercomparisons of Pandora TCO measurements**

231 In this study, the linear-least-squares regression method was used for all intercomparisons. To
232 ensure high reliability of intercomparison results, only datasets valid for all instruments were
233 selected. To this end, prior to making the comparisons it was necessary to verify the accuracy
234 of the datasets. Thus, intercomparisons of all available TCO values obtained from each
235 instrument (except for Pandora) were performed for the study period. As illustrated in Fig. 4,
236 all of the regressions show excellent agreement, with slopes close to 1:1 and coefficient of
237 determination (R^2) greater than 0.90. In particular, the Brewer data show results close to a
238 perfect correlation with those of the Dobson and OMI instruments, with slopes of 1.01 and
239 0.95, and R^2 values of 1.00 and 0.97 (OMI-TOMS), respectively. These strong correlations
240 among the datasets indicate reliable measurements with high accuracies, thereby enabling a
241 thorough regression analysis. Having established the reliability of the datasets, we next

242 carried out the intercomparisons of the Pandora TCOs.

243 Figure 5 shows scatter plots of the daily Pandora TCOs and daily Dobson, Brewer and OMI
244 TCOs, respectively, together with regression lines within an error range of $\pm 3\%$ (Basher,
245 1985; Tzortziou et al., 2012) calculated by linear least-squares regression. The slope of the
246 regression line and the coefficient of determination (R^2) from the intercomparison of the
247 Pandora data with the other instruments are 0.95 and 0.95 for the Dobson data, 1.00 and 0.97
248 for the Brewer data, 0.98 and 0.96 for the OMI-TOMS data, and 0.97 and 0.95 for the OMI-
249 DOAS data, respectively. That is, all linear regression lines between Pandora and the others
250 are close to 1 to 1 line. Furthermore, the Pandora data show the highest mean ratio value of
251 0.98 ± 0.001 ($\pm 1\sigma$) with the Brewer data, which is slightly higher than the others. According to
252 Park et al. (2012), the mean ratio value shows intercomparison accuracy. These high
253 correlation results are comparable with previous validation studies undertaken in Boulder,
254 Colorado (Herman et al., 2015) and in Greenbelt, Maryland (Tzortziou et al., 2012). Table 1
255 lists the mean relative differences, which are defined as the percentage differences between
256 the observation data. All of these values show that the measured TCO values from the
257 Dobson, Brewer, and OMI instruments are generally higher than those from Pandora. Figure
258 6 shows time series of the relative differences between the daily TCO values from Pandora
259 and the other instruments, which is smallest between the OMI-DOAS and Pandora data (2.01%
260 on average), but increases to 3.64%, 2.31%, and 2.55% for the Dobson, Brewer, and OMI-
261 TOMS data, respectively. Based on these results, we conclude that the Dobson, Brewer, and
262 OMI TCO data show good agreement with the Pandora measurements.

263 We used a generic analysis of variance table for simple linear regression (ANOVA) to
264 perform a more detailed analysis of these relationships. ANOVA tables for the Pandora

265 intercomparisons are presented in Table 2 and follow the procedure of Wilks (2006). Table 2
266 ((a)-(d)) shows the mean squared error (MSE), standard error (s.e.), and F-ratio of the
267 Pandora intercomparison results. The MSE indicates the variability of the data, with a large
268 MSE indicating a greater degree of scatter around the 1:1 line, and a small MSE the opposite.
269 The MSE value for the comparison of Pandora with Brewer was the smallest, at 42.8, and
270 was the largest (73.7) with OMI-DOAS. That is, in the case of comparison between the
271 Pandora and the Brewer, most of data are located close to the regression line (Fig. 5). The s.e.
272 of the slope and the intercept represent the uncertainty of the regression line. The s.e. value of
273 the slope was 0.02 for all Pandora intercomparisons, and that of the intercept was the smallest
274 (5.42) and largest (6.95) for the comparisons with the Brewer and OMI-DOAS instruments,
275 respectively. Finally, the F-ratio (mean square regression (MSR)/MSE) increases with the
276 strength of the regression (Draper et al., 1966; Neter et al., 1996). In Table 2, the F-ratio value
277 calculated from the regression analysis of Pandora with the Brewer is 3477.9, which is much
278 greater than the others (2351.5, 2607.4, and 1974.8 for Dobson, OMI-TOMS, and OMI-
279 DOAS, respectively). Taking all of these results into consideration, the TCO data measured
280 by Pandora are in closest agreement with the Brewer data, similar to the validation results of
281 Tzortziou et al. (2012).

282 The relatively small slopes, R^2 , and F-ratios, and large MSE show that the Pandora data have
283 a slightly weaker linear relationship with the Dobson and OMI-DOAS data than with the
284 Brewer and OMI-TOMS data (Fig. 5; Table 2). In particular, in the case of OMI-DOAS, the
285 regression result shows the smallest R^2 and F-ratio values of 0.95 and 1974.8, respectively,
286 and the largest MSE of 73.7, even though it has the smallest mean relative difference of
287 2.01%. However, the time series of relative differences between the Pandora and OMI-DOAS

288 TCO data in Fig. 6(d) shows more negative relative difference values than for the other
289 relationships, and these compensate for the positive values. That is, for OMI-DOAS, there are
290 more and larger underestimated TCO values when compared with the Pandora data than for
291 the Dobson, Brewer, and OMI-TOMS data and these underestimated TCO values lead to the
292 small mean relative difference. As a result, it is difficult to conclude that the Pandora and
293 OMI-DOAS TCO values are in good agreement only with small mean relative difference
294 value between two data. Moreover, the largest MSE and smallest F-ratio values, which are
295 used to assess the correlation between the Pandora and OMI-DOAS data, represent a poorer
296 agreement among all intercomparison results with an MSE of 73.7 and F-ratio of 1974.8
297 (Table 2(d)). Thus, in summary, the Pandora TCO data show a better correlation with the
298 Brewer or OMI-TOMS data than do the OMI-DOAS data. This result can be explained by the
299 dependence of the OMI-DOAS measurements on seasonal variations and solar zenith angle.
300 According to previous studies, for a comparison between ground-based and OMI instruments,
301 OMI-DOAS TCO data have a seasonal variation of about $\pm 2\%$ and can be overestimated by 5%
302 depending on solar zenith angle (Balis et al., 2007; Kroon et al., 2008; McPeters et al., 2008).
303 The Pandora TCO values show very good agreement with the Dobson values, with a slope of
304 0.95 and R^2 of 0.95. This result is similar to the findings of Herman et al. (2015), despite the
305 following error sources in Dobson and Pandora measurements:

- 306 • the limited amount of data used to calculate the single representative daily
307 average;
- 308 • the dependence on solar zenith angle, meaning that measurements are
309 underestimated by 6% or overestimated by 20%–30% when solar zenith angles are
310 less than 57° and greater than 60° , respectively (Bojkov, 1969; Komhyr, 1980;

311 Miyagawa et al., 2005; Nichol and Valenti, 1993);

312 • the SO₂ absorption effect (De Backer and De Muer, 1991; Komhyr, 1980;

313 Miyagawa et al., 2005); and

314 • fixed temperature and high humidity lead to a bias in TCO retrievals

315 (Herman et al., 2015; Komhyr, 1980).

316 The SO₂ absorption effect occurs when the Dobson measures TCO considering the

317 attenuation of irradiance by SO₂ in polluted air. According to Herman et al. (2015), both the

318 standard Dobson and Pandora TCO retrievals required a correction using a monthly varying

319 effective ozone temperature for removing seasonal bias.

320

321 **3.3. Diurnal variations in Pandora TCO**

322 As mentioned above, the temporal resolution of the Pandora measurements is about 2 minutes,

323 and this allows us to detect diurnal variations in ozone using the Pandora data. Figure 7

324 shows six cases of diurnal variation for the TCO values measured by the Pandora instrument

325 with average, minimum, and maximum values under clear-sky condition when the cloud

326 amount is less than 3 tenths during the study period. In this figure, TCO data measured at

327 solar zenith angles greater than 75° are shaded and excluded from the statistical calculations.

328 According to Herman et al., (2015), the Pandora (#34) TCO data measured at Boulder,

329 Colorado over 13 consecutive days in December 2013 showed considerable variations.

330 Similarly, in Fig. 7 there are substantial daytime variations for all six cases, especially on 5

331 March 2013 (Fig. 7 (c)), which shows the largest standard deviation of 15.4 DU. Moreover,

332 the range of TCO values on a given day shows a largest value of 53.4 DU, about 15.3% of the
333 daily average value. Because of these variations, the inconsistency of time intervals between
334 measurements selected for the daily averaging in the intercomparison can result in a sampling
335 bias. In particular, direct-sun observations by the Dobson instrument were performed, at most,
336 three times a day in this study. Observation times and real-time TCO values, as well as the
337 daily average values of the Dobson measurements, are shown for each diurnal cycle in Fig. 7.
338 In the six cases, the daily TCO values from the Pandora instrument were underestimated by
339 about 5% compared with those of the Dobson. For the entire period, the maximum difference
340 between the daily TCO values of the Pandora and Dobson was ~12.5% on 22 June 2013.
341 Thus, Herman et al. (2015) suggested that the Pandora time interval for intercomparison with
342 Dobson should be kept fairly short (e.g., 8 minutes) to avoid under-sampling of the
343 coincident time series. More reliable characteristics of diurnal variability of TCO will be
344 found using the long-term Pandora measurement data in the future.

345

346

347

348

349

350

351

352

353

354 **4. Summary and Conclusions**

355 In this study, daily total ozone measured by the Pandora spectrophotometer were
356 intercompared using ground-based and satellite measurements (Dobson and Brewer
357 spectrophotometers, and OMI) over a 2-year period at Yonsei University, Seoul, Korea. A
358 linear least-squares regression analysis revealed that the Pandora TCO data show excellent
359 agreement with the other instruments, with slopes close to 1 and R^2 values greater than 0.95,
360 which are within $\pm 5\%$ of perfect regression. In addition, comparison of the mean relative
361 differences shows that the Pandora TCO data were underestimated when compared with the
362 Dobson, Brewer, and OMI data. Through detailed comparison using the ANOVA approach,
363 we found that the regression of the Pandora intercomparison with the Brewer data shows the
364 smallest MSE value of 42.8 and the largest F-ratio of 3477.9, indicating a close relationship.
365 Several internal and external factors may result in slight differences between the Pandora
366 measurements and other data; i.e., the time interval difference for daily averaging,
367 dependence on solar zenith angle, SO_2 effect, temperature, and humidity for the Dobson data,
368 and dependence on seasonal variations and solar zenith angle for the OMI-DOAS data. In
369 particular, the Pandora measurements were underestimated by up to about 12.5% compared
370 with the TCO obtained from the Dobson instrument on 22 June 2013. Despite these factors,
371 daily TCO values retrieved from Pandora showed very good agreement with the Dobson,
372 OMI-DOAS, Brewer, and OMI-TOMS data. Consequently, daily total ozone data measured
373 by the Pandora spectrophotometer show high reliability, and are expected to improve
374 substantially with the regular and accurate calibration and validation associated with the
375 operational monitoring of trace gases and pollutants.

376 *Acknowledgements.* This research was supported by the GEMS program of the Ministry of
377 Environment, Korea and the Eco Innovation Program of KEITI (2012000160002).

378 **References**

- 379 Antón, M., López, M., Vilaplana, J. M., Kroon, M., McPeters, R., Bañón, M., and Serrano, A.: Validation of
380 OMI-TOMS and OMI-DOAS total ozone column using five Brewer spectroradiometers at the Iberian peninsula,
381 *Journal of Geophysical Research*, 114, 10.1029/2009jd012003, 2009.
- 382 Balis, D., Kroon, M., Koukouli, M. E., Brinksma, E. J., Labow, G., Veefkind, J. P., and McPeters, R. D.:
383 Validation of Ozone Monitoring Instrument total ozone column measurements using Brewer and Dobson
384 spectrophotometer ground-based observations, *Journal of Geophysical Research*, 112, 10.1029/2007jd008796,
385 2007.
- 386 Basher, R. E.: *Review of the Dobson spectrophotometer and its accuracy*, Springer, 1985.
- 387 Bhartia, P., and Wellemeyer, C.: OMI TOMS-V8 Total O3 algorithm, algorithm theoretical baseline document:
388 OMI ozone products. PK Bhartia (ed.), vol. II, ATBD-OMI-02, version 2.0, 2002.
- 389 Bojkov, R. D.: Differences in Dobson spectrophotometer and filter ozonometer measurements of total ozone,
390 *Journal of Applied Meteorology*, 8, 362-368, 1969.
- 391 Brewer, A.: Evidence for a world circulation provided by the measurements of helium and water vapour
392 distribution in the stratosphere, *Quarterly Journal of the Royal Meteorological Society*, 75, 351-363, 1949.
- 393 Brewer, A.: A replacement for the Dobson spectrophotometer?, *Pure and Applied Geophysics*, 106, 919-927,
394 1973.
- 395 Buchard, V., Brogniez, C., Auriol, F., Bonnel, B., Lenoble, J., Tanskanen, A., Bojkov, B., and Veefkind, P.:
396 Comparison of OMI ozone and UV irradiance data with ground-based measurements at two French sites,
397 *Atmospheric Chemistry and Physics*, 8, 4517-4528, 2008.
- 398 Cede, A.: *Manual for Pandora Software Suite Version 1.3*, 2011.
- 399 Cho, H.: *Ozone layer monitoring over Korea, 1985-1994*, Global Environment Laboratory, Yonsei University,
400 224pp, 1996.
- 401 Cho, H., Lee, S., and Choi, C.: The seasonal variations of total ozone at Seoul, *Journal of Korean Meteorology*,
402 25, 21-29, 1989.
- 403 Cho, H.-K., Kim, J., Oh, S. N., Kim, S.-K., Baek, S.-K., and Lee, Y. G.: A climatology of stratospheric ozone
404 over Korea, *Korean Journal of the Atmospheric Sciences*, 6, 97-112, 2003.
- 405 Chubachi, S.: A special ozone observation at Syowa Station, Antarctica from February 1982 to January 1983, in:
406 *Atmospheric ozone*, Springer, 285-289, 1985.
- 407 Dave, J.: Meaning of Successive Iteration of the Auxiliary Equation in the Theory of Radiative Transfer, *The*
408 *Astrophysical Journal*, 140, 1292, 1964.
- 409 De Backer, H., and De Muer, D.: Intercomparison of total ozone data measured with Dobson and Brewer ozone
410 spectrophotometers at Uccle (Belgium) from January 1984 to March 1991, including zenith sky observations,
411 *Journal of Geophysical Research: Atmospheres*, 96, 20711-20719, 1991.
- 412 Dobson, G. M. B.: Forty years' research on atmospheric ozone at Oxford: a history, *Applied Optics*, 7, 387-405,
413 1968.
- 414 Draper, N. R., Smith, H., and Pownell, E.: *Applied regression analysis*, Wiley New York, 1966.
- 415 Farman, J., Gardiner, B., and Shanklin, J.: Large losses of total ozone in Antarctica reveal seasonal ClO_x/NO_x
416 interaction, 1985.

417 Harris, N., Ancellet, G., Bishop, L., Hofmann, D., Kerr, J., McPeters, R., Prendez, M., Randel, W., Staehelin, J.,
418 and Subbaraya, B.: Trends in stratospheric and free tropospheric ozone, *Journal of Geophysical Research:*
419 *Atmospheres*, 102, 1571-1590, 1997.

420 Herman, J., Cede, A., Spinei, E., Mount, G., Tzortziou, M., and Abuhassan, N.: NO₂ column amounts from
421 ground-based Pandora and MFDOAS spectrometers using the direct-sun DOAS technique: Intercomparisons
422 and application to OMI validation, *Journal of Geophysical Research*, 114, 10.1029/2009jd011848, 2009.

423 Herman, J., Evans, R., Cede, A., Abuhassan, N., Petropavlovskikh, I., and McConville, G.: Comparison of
424 ozone retrievals from the Pandora spectrometer system and Dobson spectrophotometer in Boulder, Colorado,
425 *Atmospheric Measurement Techniques*, 8, 3407-3418, 10.5194/amt-8-3407-2015, 2015.

426 Hong, G.-M., and Cho, C.-H.: A Variability of Total Ozone Amount over Pohang using the Brewer Ozone
427 Spectrophotometer and Ozonesonde, *Asia-Pacific Journal of Atmospheric Sciences*, 43, 31-39, 2007.

428 Hong, H., Lee, H., Kim, J., and Lee, Y.-G.: First comparison of OMI-DOAS total ozone using ground-based
429 observations at a megacity site in East Asia: Causes of discrepancy and improvement in OMI-DOAS total ozone
430 during summer, *Journal of Geophysical Research: Atmospheres*, 119, 10058-10067, 10.1002/2014jd021829,
431 2014.

432 Kerr, J.: New methodology for deriving total ozone and other atmospheric variables from Brewer
433 spectrophotometer direct sun spectra, *Journal of Geophysical Research: Atmospheres*, 107, 2002.

434 Kerr, J.: The Brewer Spectrophotometer, in: *UV Radiation in Global Climate Change*, Springer, 160-191, 2010.

435 Kerr, J., McElroy, C., Wardle, D., Olafson, R., and Evans, W.: The automated Brewer spectrophotometer, in:
436 *Atmospheric Ozone*, Springer, 396-401, 1985.

437 Kim, J., Cho, H., Lee, Y., Oh, S., and Baek, S.: Updated trends of stratospheric ozone over Seoul, *Atmosphere*,
438 15, 101-118, 2005.

439 Kim, J., Park, S., Moon, K., Koo, J., Lee, Y., Miyagawa, K., and Cho, H.: Automation of Dobson
440 spectrophotometer (No. 124) for ozone measurements, *Atmos. Korean Meteor. Soc*, 17, 339-348, 2007.

441 Kim, J., Park, S., Cho, N., Kim, W., and Cho, H. K.: Recent Variations of UV Irradiance at Seoul 2004~ 2010,
442 *Atmosphere*, 21, 429-438, 2011.

443 Kim, W., Kim, J., Park, S. S., and Cho, H.-K.: UV Sensitivity to Changes in Ozone, Aerosols, and Clouds in
444 Seoul, South Korea, *Journal of Applied Meteorology and Climatology*, 53, 310-322, 10.1175/jamc-d-13-052.1,
445 2014.

446 Komhyr, W.: *Operations handbook-Ozone observations with a Dobson spectrophotometer*, NOAA
447 Environmental Research Laboratories, Air Resources Laboratory, 1980.

448 Kroon, M., Veefkind, J. P., Sneep, M., McPeters, R., Bhartia, P., and Levelt, P.: Comparing OMI-TOMS and
449 OMI-DOAS total ozone column data, *Journal of Geophysical Research: Atmospheres*, 113, 2008.

450 LEONARD, R.: Dobson spectrophotometer 83: A standard for total ozone measurements, 1962-1987, *Journal of*
451 *Geophysical Research*, 94, 9847-9861, 1989.

452 Levelt, P. F., Van den Oord, G. H., Dobber, M. R., Malkki, A., Visser, H., De Vries, J., Stammes, P., Lundell, J.
453 O., and Saari, H.: The ozone monitoring instrument, *IEEE Transactions on geoscience and remote sensing*, 44,
454 1093-1101, 2006.

455 Liou, K.N.: *An introduction to to atmospheric radiation (Vol. 84)*, Academic press, 2002.

456 Martens, W.: Health impacts of climate change and ozone depletion: an ecoepidemiologic modeling approach,
457 *Environmental Health Perspectives*, 106, 241, 1998.

- 458 McPeters, R., Kroon, M., Labow, G., Brinksma, E., Balis, D., Petropavlovskikh, I., Veefkind, J. P., Bhartia, P. K.,
459 and Levelt, P. F.: Validation of the Aura Ozone Monitoring Instrument total column ozone product, *Journal of*
460 *Geophysical Research*, 113, 10.1029/2007jd008802, 2008.
- 461 Miyagawa, K., Kim, J., and Cho, H.: Intercomparison of Dobson Spectrophotometer in Yonsei Univ., Korea,
462 *Journal of Aerological Observatory*, 65, 93-98, 2005.
- 463 Neter, J., Kutner, M. H., Nachtsheim, C. J., and Wasserman, W.: *Applied linear statistical models*, Irwin Chicago,
464 1996.
- 465 Newchurch, M., Yang, E. S., Cunnold, D., Reinsel, G. C., Zawodny, J., and Russell, J. M.: Evidence for
466 slowdown in stratospheric ozone loss: First stage of ozone recovery, *Journal of Geophysical Research:*
467 *Atmospheres*, 108, 2003.
- 468 Nichol, S., and Valenti, C.: Intercomparison of total ozone measured at low sun angles by the Brewer and
469 Dobson spectrophotometers at Scott Base, Antarctica, *Geophysical research letters*, 20, 2051-2054, 1993.
- 470 Park, S. S., Kim, J., Cho, H. K., Lee, H., Lee, Y., and Miyagawa, K.: Sudden increase in the total ozone density
471 due to secondary ozone peaks and its effect on total ozone trends over Korea, *Atmospheric Environment*, 47,
472 226-235, 10.1016/j.atmosenv.2011.11.011, 2012.
- 473 Platt, U., and Stutz, J.: *Differential absorption spectroscopy*, Springer, 2008.
- 474 Reed, A. J., Thompson, A. M., Kollonige, D. E., Martins, D. K., Tzortziou, M. A., Herman, J. R., Berkoff, T. A.,
475 Abuhassan, N. K., and Cede, A.: Effects of local meteorology and aerosols on ozone and nitrogen dioxide
476 retrievals from OMI and Pandora spectrometers in Maryland, USA during DISCOVER-AQ 2011, *Journal of*
477 *atmospheric chemistry*, 72, 455-482, 2015.
- 478
479 Sabburg, J., Rives, J. E., Meltzer, R. S., Taylor, T., Schmalzle, G., Zheng, S., Huang, N., Wilson, A., and
480 Udelhofen, P. M.: Comparisons of corrected daily integrated erythemal UVR data from the U.S. EPA/UGA
481 network of Brewer spectroradiometers with model and TOMS-inferred data, *Journal of Geophysical Research:*
482 *Atmospheres*, 107, ACL 5-1-ACL 5-10, 10.1029/2001jd001565, 2002.
- 483 Schott, J. R.: *Remote Sensing*, Oxford University Press, 2007.
- 484 Solomon, S.: Stratospheric ozone depletion: A review of concepts and history, *Reviews of Geophysics*, 37, 275-
485 316, 1999.
- 486 Tzortziou, M., Herman, J. R., Cede, A., and Abuhassan, N.: High precision, absolute total column ozone
487 measurements from the Pandora spectrometer system: Comparisons with data from a Brewer double
488 monochromator and Aura OMI, *Journal of Geophysical Research: Atmospheres*, 117, n/a-n/a,
489 10.1029/2012jd017814, 2012.
- 490 Veefkind, J. P., De Haan, J. F., Brinksma, E. J., Kroon, M., and Levelt, P. F.: Total ozone from the Ozone
491 Monitoring Instrument (OMI) using the DOAS technique, *IEEE Transactions on Geoscience and Remote*
492 *Sensing*, 44, 1239-1244, 2006.
- 493 Wang, S., Pongetti, T. J., Sander, S. P., Spinei, E., Mount, G. H., Cede, A., and Herman, J.: Direct Sun
494 measurements of NO₂ column abundances from Table Mountain, California: Intercomparison of low- and high-
495 resolution spectrometers, *Journal of Geophysical Research*, 115, 10.1029/2009jd013503, 2010.
- 496 Wang, Y., Zhao, P., Xu, H., and Liu, G.: Anomalies of Northern Hemisphere ozone associated with a
497 tropopause-lower stratosphere teleconnection during summer, *International Journal of Climatology*, 2015.
- 498 Weatherhead, E. C., Reinsel, G. C., Tiao, G. C., Jackman, C. H., Bishop, L., Frith, S. M. H., DeLuisi, J., Keller,
499 T., Oltmans, S. J., and Fleming, E. L.: Detecting the recovery of total column ozone, *Journal of Geophysical*
500 *Research: Atmospheres*, 105, 22201-22210, 2000.

501 Weber, M., Dhomse, S., Wittrock, F., Richter, A., Sinnhuber, B. M., and Burrows, J.: Dynamical control of NH
502 and SH winter/spring total ozone from GOME observations in 1995–2002, *Geophysical Research Letters*, 30,
503 2003.

504 Wilks, D. S.: *Statistical methods in the atmospheric sciences*, Academic press, 2006.

505 WMO (World Meteorological Organization): *Assessment for Decision-Makers: Scientific Assessment of*
506 *Ozone Depletion: 2014*, 88 pp., Global Ozone Research and Monitoring Project—Report No. 56, Geneva,
507 Switzerland, 2014.

508 Yun, S., Lee, H., Kim, J., Jeong, U., Park, S. S., and Herman, J.: Inter-comparison of NO₂ column densities
509 measured by Pandora and OMI over Seoul, Korea, *Korean Journal of Remote Sensing*, 29, 663-670,
510 10.7780/kjrs.2013.29.6.9, 2013.

511

512

513

514

515

516

517

518

519

520

521

522

523

524

525

526

527 **Table 1.** Summary of intercomparison statistics for the 2 years from March 2012 to March 2014.

	Pandora	Dobson	Brewer	OMI-TOMS	OMI-DOAS
Average [DU]	317.2	331.9	325.1	324.1	322.0
Standard deviation	36.8	38.6	36.2	38.0	38.6
Max (date)	436.7 (2012/Apr/6)	463.0 (2013/Apr/26)	449.3 (2013/Apr/26)	467.1 (2013/Apr/10)	465.1 (2013/Apr/10)
Min (date)	249.2 (2013/Oct/7)	239.0 (2013/Oct/7)	246.5 (2013/Oct/7)	238.3 (2013/Oct/8)	241.8 (2013/Oct/8)
Mean relative difference [%]					
	Dobson–Pandora	Brewer–Pandora	OMI-TOMS–Pandora	OMI-DOAS–Pandora	
	3.64	2.31	2.55	2.01	

528

529

530 **Table 2(a).** ANOVA table for simple linear regression between the Pandora and Dobson data.¹ .

Source	df	SS	MS	F	P
Total	114	153,818			
Regression	1	146,765	146,765	2351.5	< 0.0001
Residual (error)	113	7053	62.4		
Variable	<i>Coefficient</i>	<i>s.e.</i>	<i>t ratio</i>		
Intercept	5.21	6.35	0.82		
Slope	0.95	0.02			

531

532 **Table 2(b).** As for Table 2(a) but for comparison of the Pandora and Brewer data.

Source	Df	SS	MS	F	P
Total	114	153,818			
Regression	1	148,978	148,978	3477.9	< 0.0001
Residual (error)	113	4840	42.8		
Variable	<i>Coefficient</i>	<i>s.e.</i>	<i>t ratio</i>		
Intercept	-6.15	5.42	-1.14		
Slope	1.00	0.02			

533

¹ The column headings df, SS, MS, F, and P stand for degrees of freedom, sum of squares, mean square, F-ratio, and P-value, respectively.

534 **Table 2(c).** As for Table 2(a) but for comparison of the Pandora and OMI-TOMS data.

Source	df	SS	MS	F	P
Total	114	153,818			
Regression	1	147,429	147,429	2607.4	< 0.0001
Residual (error)	113	6389	56.5		
Variable	<i>Coefficient</i>	<i>s.e.</i>	<i>t ratio</i>		
Intercept	-1.66	6.17	-0.27		
Slope	0.98	0.02			

535

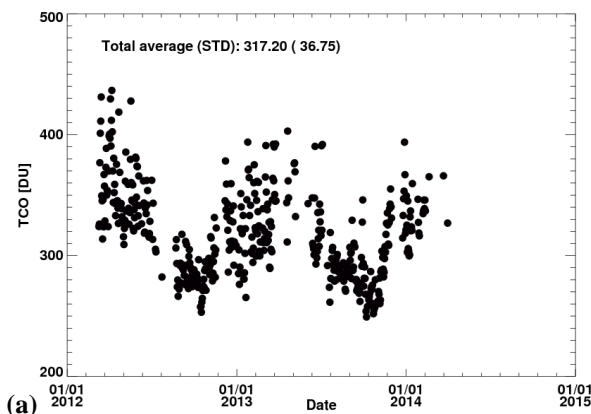
536

537 **Table 2(d).** As for Table 2(a) but for comparison of the Pandora and OMI-DOAS data.

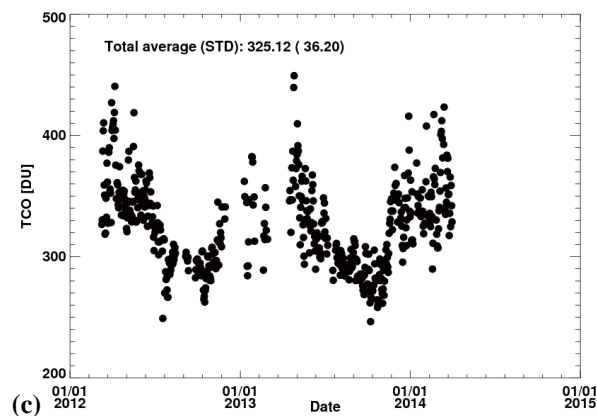
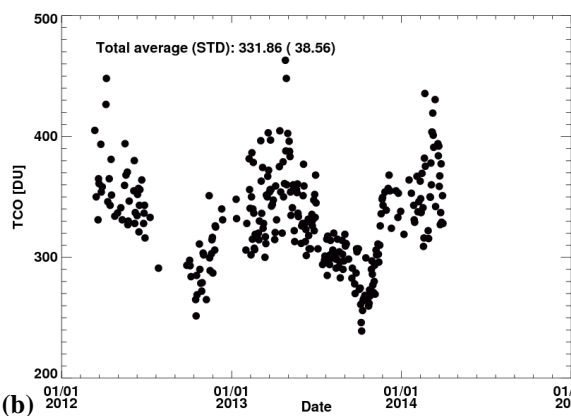
Source	df	SS	MS	F	P
Total	114	153,818			
Regression	1	145,493	145,493	1974.8	< 0.0001
Residual (error)	113	8325	73.7		
Variable	<i>Coefficient</i>	<i>s.e.</i>	<i>t ratio</i>		
Intercept	4.46	6.95	0.64		
Slope	0.97	0.02			

538

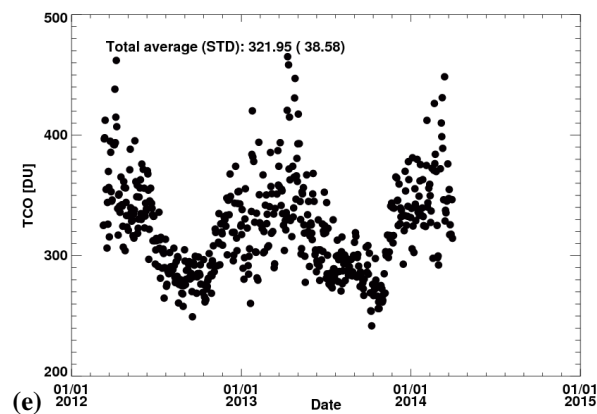
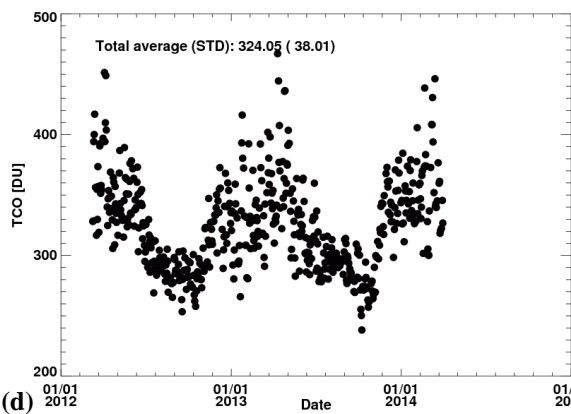
540



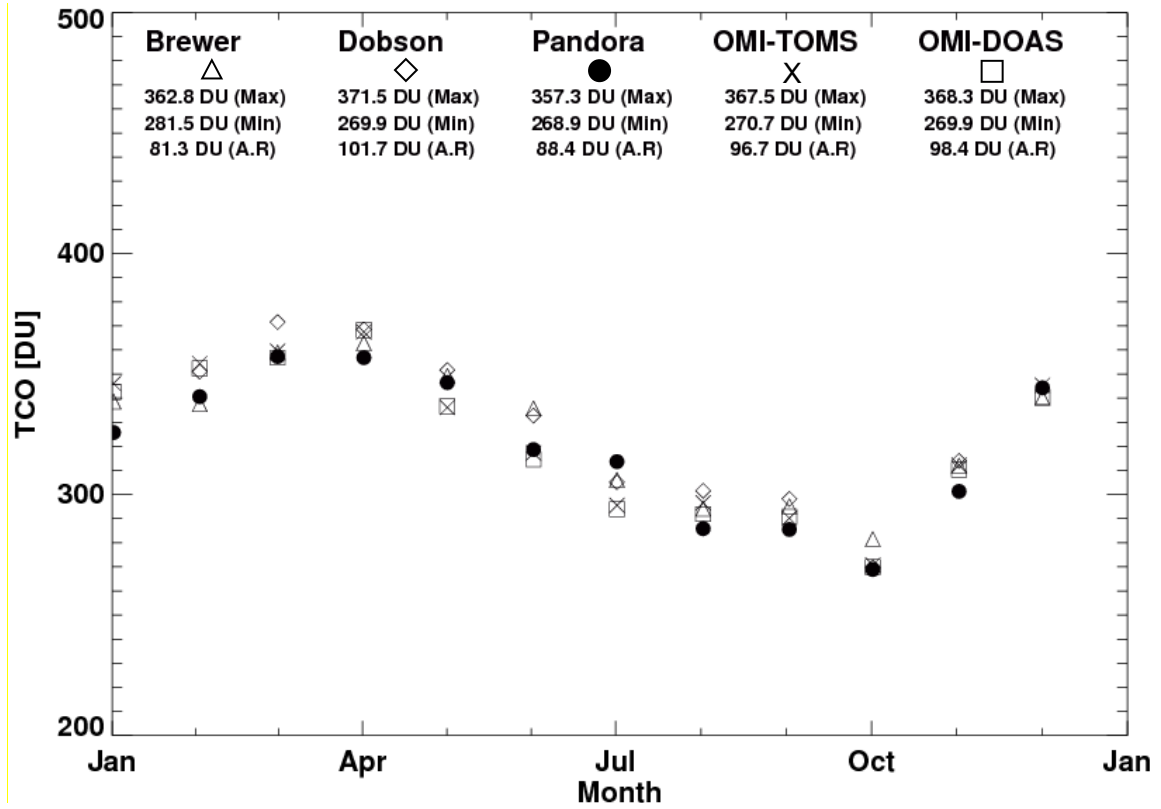
541



542



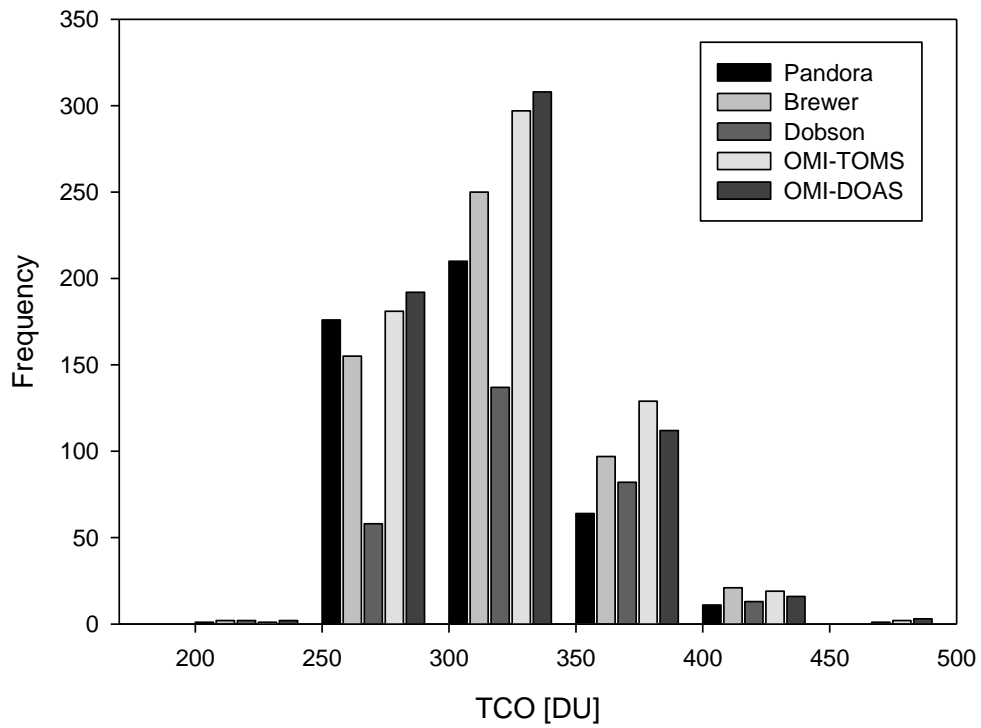
543 **Figure 1.** Daily TCO values from the following instruments: (a) Pandora, (b) Dobson, (c) Brewer, (d) OMI-
544 TOMS, and (e) OMI-DOAS, for the 2 years from March 2012 to March 2014.



545

546 **Figure 2.** 2-year averaged monthly TCO values, together with the maximum, minimum values and annual
 547 ranges (A.R) from the four instruments over the study period.

548



549

550 **Figure 3.** Histogram of daily TCO values from the four instruments (Pandora, Brewer, Dobson, and OMI
 551 (TOMS and DOAS). Vertical axis (frequency) stands for the number of data in each TCO interval.

552

553

554

555

556

557

558

559

560

561

562

563

564

565

566

567

568

569

570

571

572

573

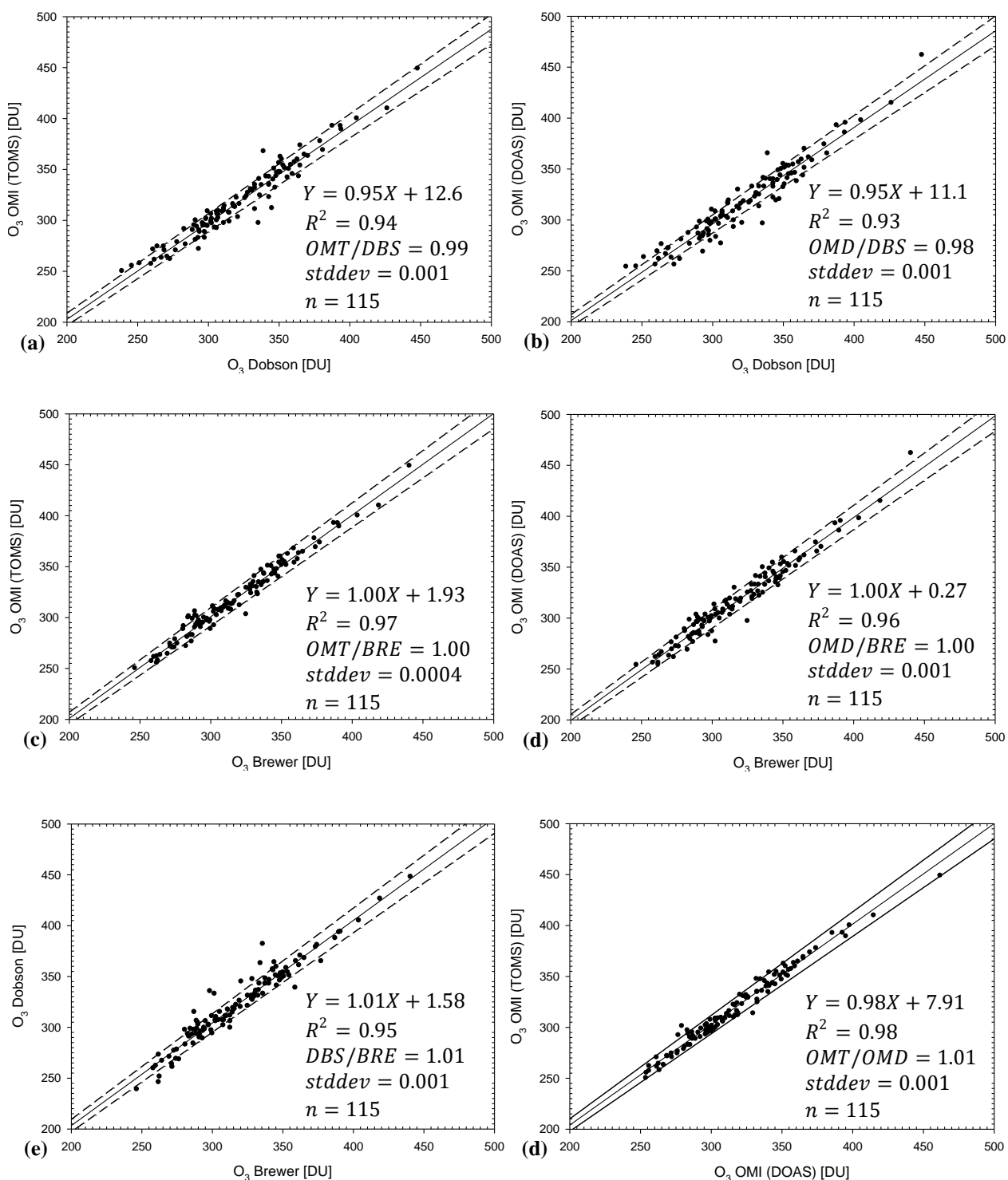


Figure 4. Intercomparison of daily TCO values between (a) Dobson and OMI-TOMS, (b) Dobson and OMI-DOAS, (c) Brewer and OMI-DOAS, (e) Brewer and Dobson, and (f) OMI-DOAS and OMI-TOMS.

574

575

576

577

578

579

580

581

582

583

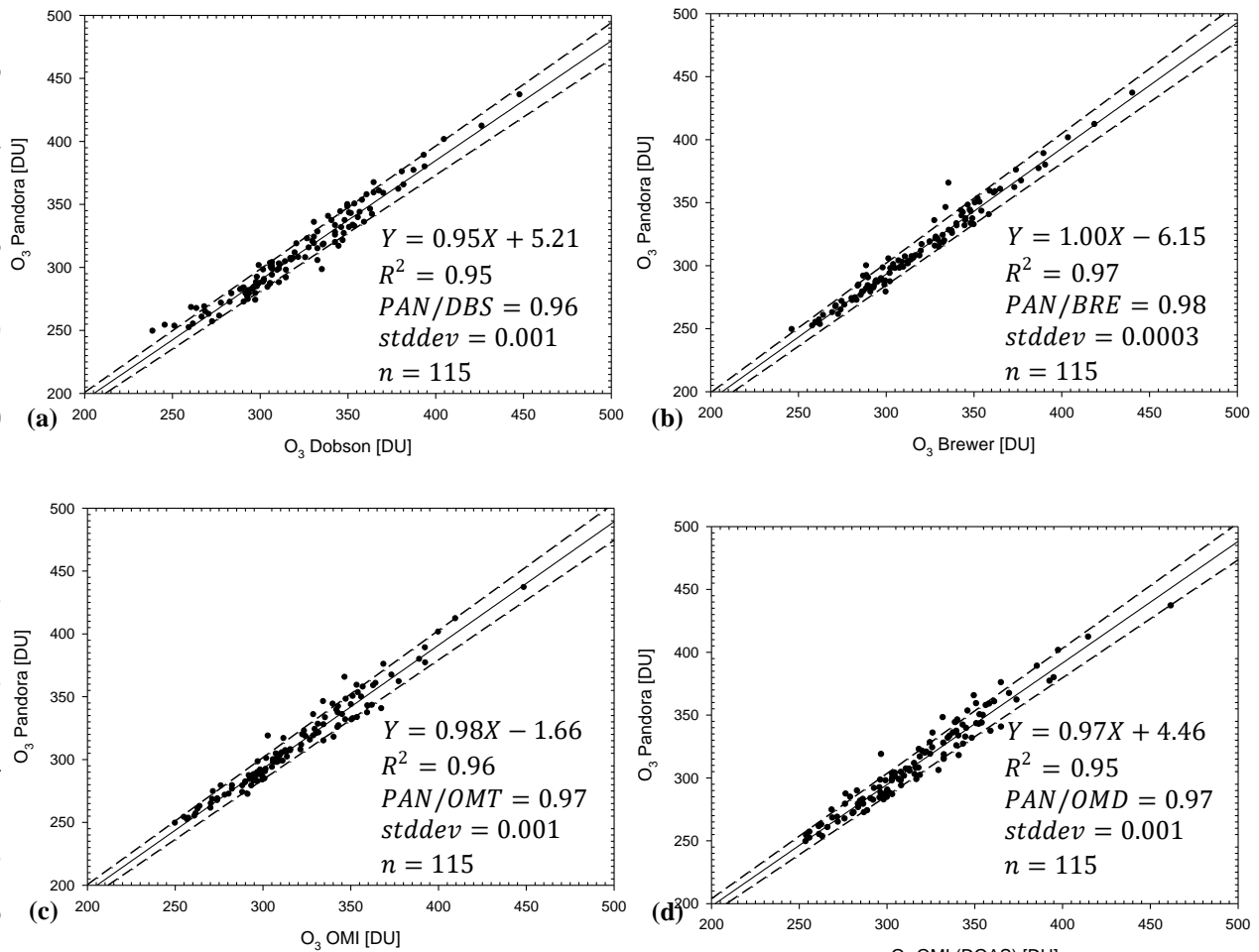
584

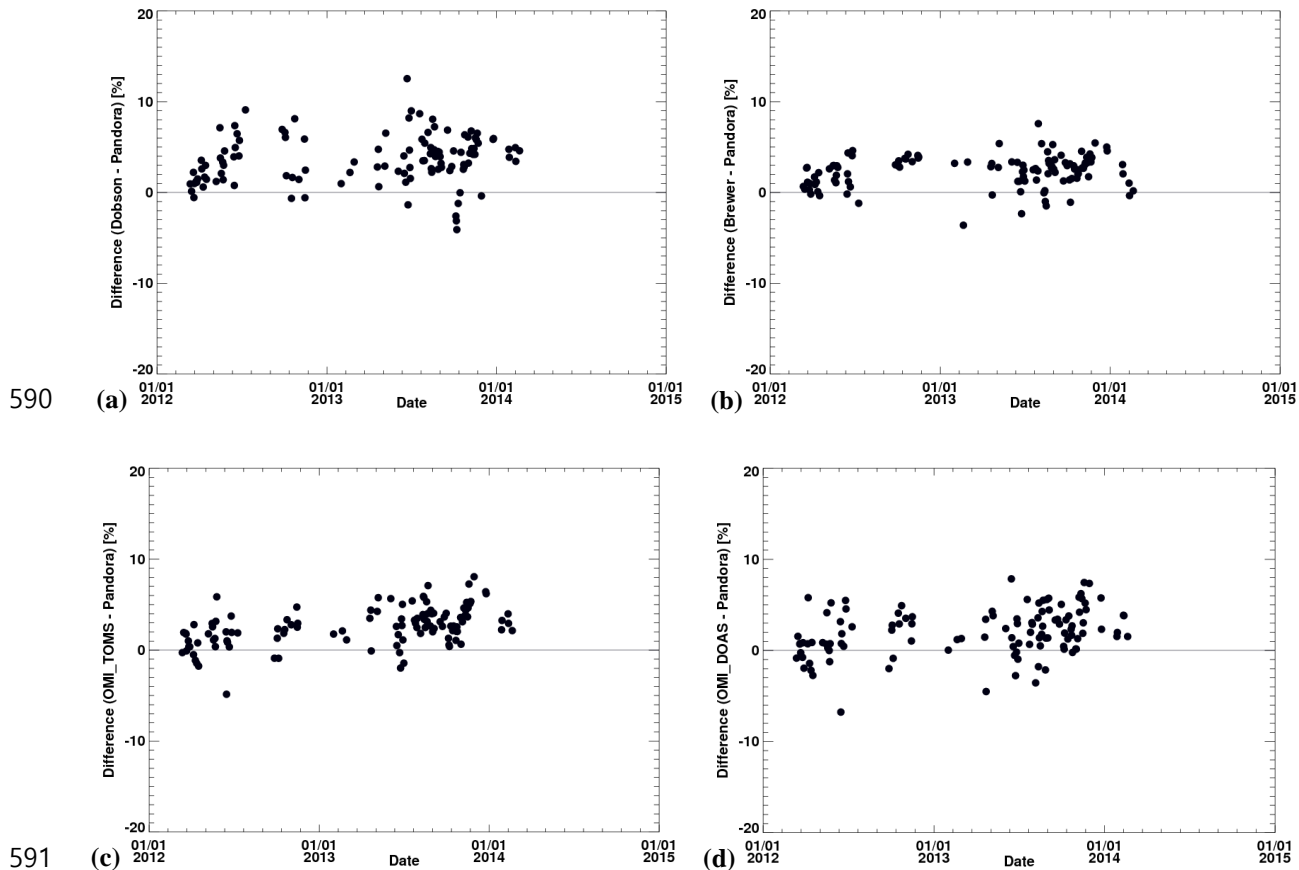
585

586

587 **Figure 5.** Intercomparison of daily TCO values from Pandora with Dobson (a), Brewer (b), and OMI (OMI (c)
588 and DOAS (d)). Solid lines represent regression lines and dashed lines indicate an error range of $\pm 3\%$.

589





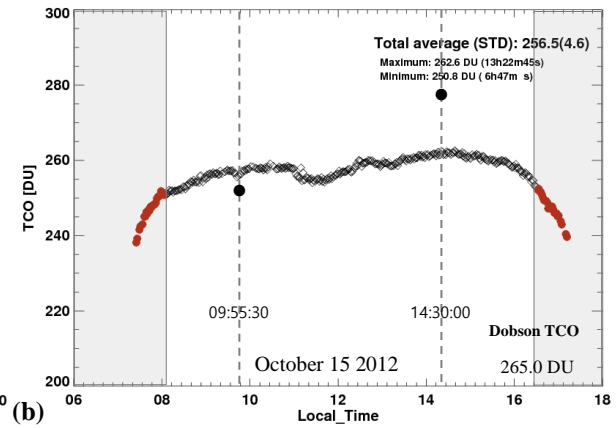
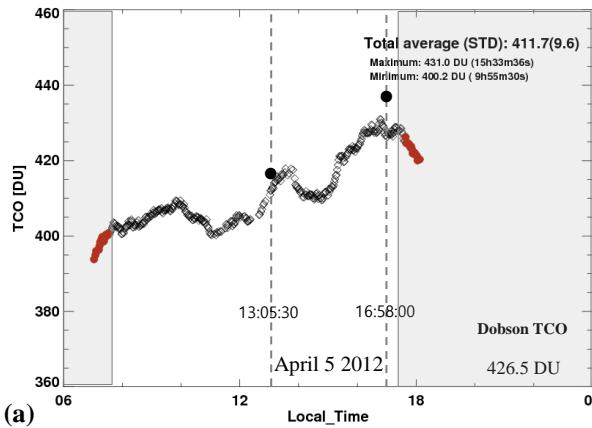
590

591

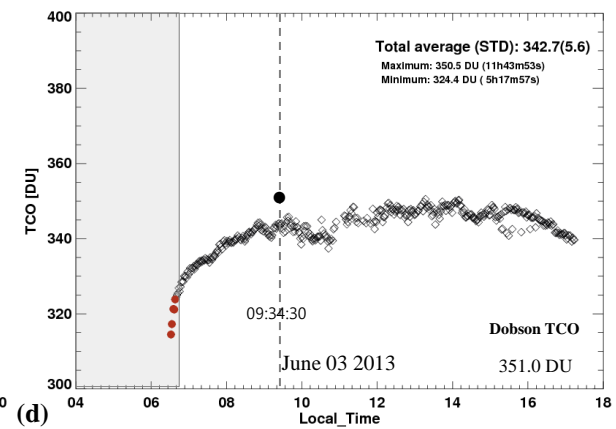
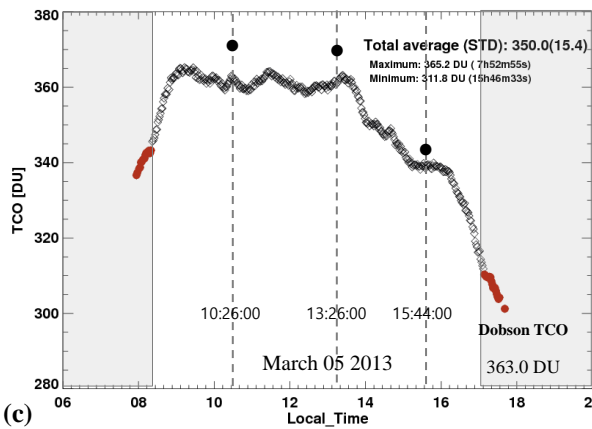
592 **Figure 6.** Time series of relative differences in daily TCO values from Pandora and those from (a) Dobson, (b)
 593 Brewer, (c) OMI-TOMS, and (d) OMI-DOAS ($\frac{TCO_{inst} - TCO_{Pan}}{TCO_{Pan}}$ [%]). Gaps in the time series indicate at least one
 594 missing value in the observations from the four instruments.

595

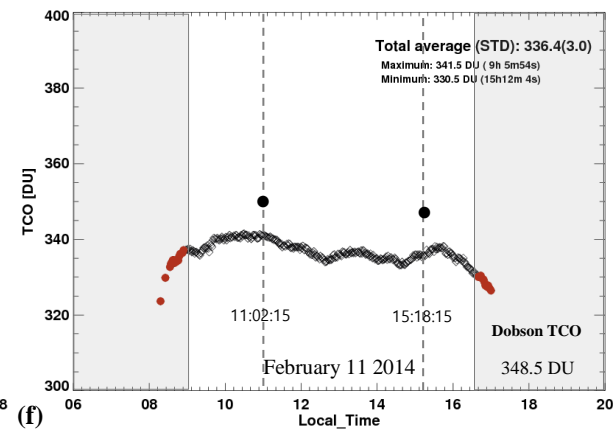
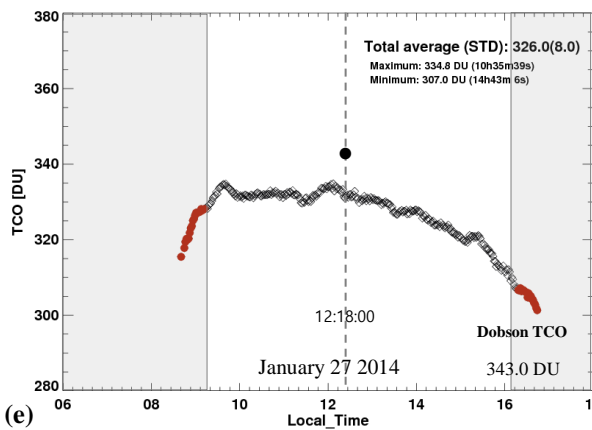
596



597



598



599 **Figure 7.** Diurnal variations in TCO values retrieved from Pandora for six randomly selected clear-sky days
 600 (cloud amount < 3) during the study period. TCO values measured at solar zenith angle > 75° and
 601 were removed from the calculations. Filled circles and dashed lines represent direct-sun TCO values measured
 602 by the Dobson instrument and observation times, respectively. All vertical axes have the same scale-range of
 603 100 DU.

605