



# Development of automated preparation system for isotopocule analysis of N<sub>2</sub>O in various air samples

Sakae Toyoda<sup>1</sup> and Naohiro Yoshida<sup>1,2,3</sup>

<sup>1</sup> Department of Environmental Science and Technology, Tokyo Institute of Technology, Yokohama, 226-8502, Japan

5 <sup>2</sup> Department of Environmental Chemistry and Engineering, Tokyo Institute of Technology, Yokohama, 226-8502, Japan

<sup>3</sup> Earth-Life Science Institute, Tokyo Institute of Technology, Tokyo, 152-8550, Japan

*Correspondence to:* Sakae Toyoda (toyoda.s.aa@m.titech.ac.jp)

**Abstract.** Nitrous oxide (N<sub>2</sub>O), an increasingly abundant greenhouse gas in the atmosphere, is the most important  
10 stratospheric ozone-depleting gas of this century. Natural abundance ratios of isotopocules of N<sub>2</sub>O, NNO molecules  
substituted with stable isotopes of nitrogen and oxygen, are a promising index of various sources or production pathways of  
N<sub>2</sub>O and of its sink or decomposition pathways. Several automated methods have been reported to improve the analytical  
precision for the isotopocule ratio of atmospheric N<sub>2</sub>O and to reduce the labor necessary for complicated sample-preparation  
procedures related to mass spectrometric analysis. However, no method accommodates flask samples with limited volume or  
15 pressure. Here we present an automated pre-concentration system for gas samples of various amounts and various N<sub>2</sub>O  
concentrations. The shortest processing time for a single analysis of typical atmospheric sample is 40 min. Precision values  
of isotopocule ratio analysis are < 0.1‰ for δ<sup>15</sup>N<sup>bulk</sup> (average abundances of <sup>14</sup>N<sup>15</sup>N<sup>16</sup>O and <sup>15</sup>N<sup>14</sup>N<sup>16</sup>O relative to <sup>14</sup>N<sup>14</sup>N<sup>16</sup>O),  
< 0.2‰ for δ<sup>18</sup>O (relative abundance of <sup>14</sup>N<sup>14</sup>N<sup>18</sup>O), and < 0.5‰ for SP (difference between relative abundance of <sup>14</sup>N<sup>15</sup>N<sup>16</sup>O  
and <sup>15</sup>N<sup>14</sup>N<sup>16</sup>O). That precision is comparable to those of other automated systems, but better than that of our previously  
20 reported manual measurement system.

## 1 Introduction

Long-term monitoring of trace gases that are increasingly abundant in the atmosphere is fundamental for the analysis of the  
imbalance of their sources and sinks and for prediction of future environmental change on Earth. Nitrous oxide (N<sub>2</sub>O) is one  
such trace gas, with global warming potential that is 220 times as great as that of carbon dioxide (CO<sub>2</sub>) and being the most  
25 important stratospheric ozone-depleting gas of this century (Myhre et al., 2013; Ravishankara et al., 2009). Its globally  
averaged concentration is about 324 nmol mol<sup>-1</sup> in 2011 (Hartmann et al., 2013). Its rate of increase is about 0.73 nmol mol<sup>-1</sup>  
a<sup>-1</sup> (Ciais et al., 2013). Sources of N<sub>2</sub>O include natural and agricultural soils, aqueous environments such as oceans, rivers,  
and lakes, industrial processes such as fossil fuel combustion, biomass burning, and animal and human wastes (Ciais et al.,  
2013), but its major sink is photochemical decomposition in the stratosphere.



Although concentration analyses yield quantitative information related to trace gases straightforwardly, it is often difficult to differentiate the sources contributing to the increase of such gases in the atmosphere, especially for N<sub>2</sub>O. Natural abundance ratios of stable isotopes of the elements that compose trace gas molecules have qualitative information related to the origin and production–decomposition processes of the gases because isotope ratios are generally different among different compounds. Moreover, they can change during physical, chemical, and biological processes. Regarding N<sub>2</sub>O, measurements of the nitrogen isotope ratio (<sup>15</sup>N/<sup>14</sup>N) for the atmosphere and various sources since the 1980s have revealed that the imbalance of isotopically light N<sub>2</sub>O from surface sources and isotopically heavy N<sub>2</sub>O refluxed from the stratosphere after its partial decomposition causes a decrease in <sup>15</sup>N/<sup>14</sup>N (Ishijima et al., 2007; Röckmann et al., 2003a; Sowers et al., 2002). Furthermore, a technique developed for measuring isotopomers of N<sub>2</sub>O (<sup>14</sup>N<sup>15</sup>N<sup>16</sup>O and <sup>15</sup>N<sup>14</sup>N<sup>16</sup>O) expanded conventional isotopic analysis to isotopocule analysis by which ratios of NNO molecules substituted with stable isotopes of nitrogen or oxygen at any site relative to <sup>14</sup>N<sup>14</sup>N<sup>16</sup>O are obtained and by which production and decomposition pathways can be differentiated in greater detail (Toyoda et al., 2015a and references therein).

Compared to concentration analysis, stable isotope and isotopocule analyses require (1) larger sample amounts, (2) more time and labor to extract and purify the target compound from the sample, and (3) larger and more expensive apparatus. Although recently developed tunable diode laser absorption spectroscopy (TDLAS) relaxes some of the requirements above and although it has some potential for onsite monitoring of stable isotope – isotopocule ratios of trace gases (Harris et al., 2014; Mohn et al., 2012; Tuzson et al., 2008), mass spectrometry combined with flask sampling still holds advantages for high-precision isotopic monitoring at polar regions or remote areas and flight observation using a balloon or an airplane.

In most currently used mass spectrometric analytical methods for N<sub>2</sub>O isotopocules, air samples are first passed through chemical adsorbents to remove CO<sub>2</sub> and water vapor. Then, N<sub>2</sub>O is concentrated on chemically inert adsorbents or inner walls of narrow tubes at liquid nitrogen temperatures. It is further purified on a capillary column of GC and is introduced directly into IRMS. The analysis of a single sample takes 30–60 min. The precision reported in earlier studies is typically 0.1–0.5‰ for 1 nmol of N<sub>2</sub>O (e.g., Toyoda et al., 2001) (See Sect. 2.4 for notation of isotopocule ratios), which is worse than the ultimate precision expected from the “shot-noise limit” of the IRMS (Potter et al., 2013) and which is insufficient to resolve the secular trend of atmospheric N<sub>2</sub>O isotopocule ratios. This low precision is partly caused by incomplete separation or purification, or by imprecise manual handling during sample preparation.

To improve the precision of the isotopocule ratio analysis of atmospheric N<sub>2</sub>O and to reduce the labor for complicated sample-preparation procedures for mass spectrometric analysis, several automated methods have been reported. Röckmann et al. (2003b) improved the precision of fragment ion (NO<sup>+</sup>) analysis by modifying the gas chromatographic purification of N<sub>2</sub>O from interfering species such as halocarbons and less-volatile compounds (Röckmann et al., 2003b). Röckmann and Levin (2005) and Potter et al. (2013) reported further improvement in the precision by partially or fully automating sample preparation steps and by slightly increasing the sample size.

In addition to mass spectrometric method, an automated sample preparation-mid-infrared quantum cascade laser spectroscopy system has been reported recently for the monitoring of atmospheric isotopocules of N<sub>2</sub>O (Harris et al., 2014;



Mohn et al., 2010). However, previously reported automated methods entail several shortcomings. First, they are designed to measure pressurized air samples such as ambient air drawn by pumps or air collected into glass bottles or metal cylinders using pumps. For that reason, they are not applicable to samples at subatmospheric pressure. Second, they are not readily applicable to high-concentration samples such as source gases because the amount of injected sample is adjusted by changing the time for flowing the sample using a mass flow controller.

Here we present an automated pre-concentration system that is flexible in terms of the type of gas samples. It can accommodate samples of various pressures and various  $N_2O$  concentrations using a vacuum line and a computer program that controls valves to inject samples of a designated amount.

## 2 Preparation system

The preparation system developed in this study consists of a sample injection unit, cryogenic concentration unit, purification unit, and cryofocusing unit (Fig. 1). It is placed in a steel rack (60 cm width, 80 cm depth, 150 cm height) with wheels attached, and has a connected gas chromatograph–isotope ratio monitoring mass spectrometer (GC-IRMS). Details of each unit are presented below.

### 2.1 Sample injection unit

This unit consists of a multi-position six-port switching valve (E4SD6MWE; Valco Instruments Co. Inc., Houston, TX) equipped with an electric actuator, air-actuated diaphragm shut-off valves (FPR-ND-71-6.35-2; Fujikin, Osaka, Japan), a pressure gauge (VSHT21; Valcom Co. Ltd., Toyonaka, Japan), a capacitance manometer (Barocel Model 600; BOC Edwards, Wilmington, MA), a Pirani vacuum gauge (GP-2A; ULVAC, Inc., Chigasaki, Japan), a vacuum pump system (turbo drag pump TMH 071 P and diaphragm pump MVP 015-2 with a controller; Pfeiffer Vacuum GmbH, Asslar, Germany), three custom-made glass bottles, and stainless steel (SUS) tubing.

A sample flask (made of either glass or SUS) or a gas cylinder is manually connected to one of the switching valve ports with a SUS connector (Cajon Ultra-Torr or Swagelok; Swagelok Company, Solon, OH). The tubing between the diaphragm valve (V1, V2, or V3) and the flask/cylinder valve is evacuated by manually operating the valves and vacuum pump via the PC control panel (see below). Then, all the diaphragm valves are closed, the flask valve is opened by hand, and a PC program for sample preparation (see Sect. 2.5) is started.

First, the flask inner pressure is measured using the pressure gauge by expanding the sample gas into the vacuum line until V4. Based on the pressure and the volume of the flask and the sample size to be injected, the “sample expanding option” and final pressure of the sample injected into the vacuum line is calculated. Seven options exist for sample expansion into the calibrated volume in the vacuum line from  $100\text{ cm}^3$  (option #1) to  $510\text{ cm}^3$  (option #7). This expansion is realized by a combination of the three glass bottles (C, D, and E in Fig. 1) with different volumes.



Next, an aliquot of the sample in the flask is expanded by sequential open–close operation of diaphragm valves. The pressure is monitored using the manometer. When the pressure agrees with the pre-calculated value within  $\pm 5\%$ , valves V1–V4 are closed, the pressure is recorded, and the injected sample amount is calculated.

## 2.2 Concentration unit

5 This unit consists of a chemical trap (CT in Fig. 1), an electrically actuated two-position six-port switching valve (E4C6UWE; Valco Instruments Co. Inc., Houston, TX, SV2 in Fig. 1), a U-shaped concentration trap, and a mass flow controller (SEC-E40; Horiba Stec Co. Ltd., Kyoto, Japan) (Fig. 1). The chemical trap is a glass tube (9 mm inner diameter, 20 cm long) packed with  $\text{Mg}(\text{ClO}_4)_2$  (8–24 mesh; Wako Pure Chemical Industries Ltd., Osaka, Japan), NaOH on support (Ascarite, 20–30 mesh; Thomas Scientific), and  $\text{Mg}(\text{ClO}_4)_2$  (20–48 mesh) in series. The concentration trap is a SUS tube (1/4  
10 inch outer diameter, 30 cm long) packed with glass beads (Flusin GH 60–80 mesh; GL Sciences Inc., Tokyo, Japan). First, the concentration trap is purged with ultra-pure He ( $> 99.9999\%$ , Japan Air Gases Ltd., Tokyo, Japan) at  $100 \text{ cm}^3 \text{ min}^{-1}$  for 10 s by switching SV2 to the “inject” position. The He is purified in advance through a column packed with molecular sieves 5A, active charcoal, and molecular sieves 13X in series. Next, SV2 is switched to the “load” position and the trap is cooled with liquid nitrogen in an SUS dewar which is driven up and down by a custom-made air-actuated stage and which is  
15 filled with liquid nitrogen from an automatic liquid nitrogen supply system (Koshin Ltd., Tokyo, Japan). Then V5, V15, and valves relevant to the sample injection option (V9–V14) are opened. The sample gas in the calibrated volume is transferred to the concentration trap through the chemical trap by He carrier gas at  $30 \text{ cm}^3 \text{ min}^{-1}$ . The He is purified in a similar manner to that described above. When more than two glass bottles are filled with the sample gas, the transfer is conducted sequentially. The transfer time is set so that the total volume of He which flows through the bottle is twice the bottle volume.

## 20 2.3 Purification and cryofocusing unit

This unit consists of two electrically actuated two-position six-port switching valves (SV3 and SV4 in Fig. 1, E4C6UWE; Valco Instruments Co. Inc., Houston, TX), a gas chromatograph (GC-8AIT; Shimadzu Corp., Kyoto, Japan) equipped with a thermal conductivity detector (TCD), and a U-shaped cryofocusing trap (Fig. 1). The GC column is an SUS tube (4 mm o.d., 3 m length) packed with Porapak Q (80–100 mesh; Waters Corp., MA). It is kept at  $60^\circ\text{C}$ . The cryofocusing trap is an SUS  
25 tube (1/16 inch o.d., 70 cm long) with no packing material.

Initially, SV3 and SV4 are set to the "Load" position. After the sample concentration step is completed, SV2 is switched to the "Inject" position. The concentration trap is heated to  $70^\circ\text{C}$  by lowering the liquid nitrogen dewar and turning on an electric sheathed heater attached to the trap. The concentrated trace gases are transferred to the GC column with purified He at  $20 \text{ cm}^3 \text{ min}^{-1}$ . When 2 min have passed after the GC injection, the cryofocusing trap is cooled with liquid nitrogen by  
30 moving up another SUS dewar. Three minutes later, SV4 is switched to the "Inject" position. Purified  $\text{N}_2\text{O}$  from the GC is focused on the trap for 2 min.



## 2.4 Injection to GC-IRMS

After the cryofocusing step is completed, SV4 is switched to the "Load" position, the liquid nitrogen dewar is moved down, and the cryofocus trap is heated to 70°C similarly, as in the case of the concentration trap. The N<sub>2</sub>O is injected into another GC (GC6890; Agilent Technologies Inc., Santa Clara, CA) with He (2 cm<sup>3</sup> min<sup>-1</sup>). It is further purified with the GS Carbon  
 5 PLOT column (0.32 mm i.d., 3 μm film thickness, 30 m; Agilent Technologies Inc.) maintained at 35°C. The purified N<sub>2</sub>O is finally injected into an IRMS (MAT252; Thermo Fisher Scientific K.K., Yokohama, Japan) via an interface that includes a gas drier with a permeation tube and two open split interfaces for the sample and reference gas (GC-Combustion Interface; Thermo Fisher Scientific K.K., slightly modified).

Mass spectrometric analysis of N<sub>2</sub>O isotopocules was conducted as described elsewhere (Toyoda and Yoshida, 1999;  
 10 Toyoda et al., 2015b). Briefly, molecular (N<sub>2</sub>O<sup>+</sup>) and fragment (NO<sup>+</sup>) ions of N<sub>2</sub>O are analyzed in independent runs. Solving the following equations and applying correction for the rearrangement or scrambling reactions during fragmentation, the isotopocule ratios are obtained as delta values.

$${}^{45}R = {}^{15}R^{\alpha} + {}^{15}R^{\beta} + {}^{17}R \quad (1)$$

$${}^{46}R = {}^{18}R + ({}^{15}R^{\alpha} + {}^{15}R^{\beta}) {}^{17}R + {}^{15}R^{\alpha} {}^{15}R^{\beta} \quad (2)$$

15  ${}^{31}R = {}^{15}R^{\alpha} + {}^{17}R \quad (3)$

$${}^{17}R = A({}^{18}R)^{\gamma} \quad (4)$$

$$\delta^{15}N^i = {}^{15}R_{\text{sample}}^i / {}^{15}R_{\text{std}}^i - 1 \quad (i = \alpha, \beta, \text{ or bulk}) \quad (5)$$

$$\delta^{18}O = {}^{18}R_{\text{sample}} / {}^{18}R_{\text{std}} - 1 \quad (6)$$

$$SP = \delta^{15}N^{\alpha} - \delta^{15}N^{\beta} \quad (7)$$

20 In Eqs. (1)–(6), <sup>45</sup>R and <sup>46</sup>R respectively denote the measured ion-beam intensity ratios of *m/z* 45/44 and 46/44 in molecular ion analysis; <sup>31</sup>R shows a 31/30 ratio by fragment ion analysis; <sup>15</sup>R<sup>α</sup>, <sup>15</sup>R<sup>β</sup>, <sup>17</sup>R, and <sup>18</sup>R respectively denote the abundance of ions <sup>14</sup>N<sup>15</sup>N<sup>16</sup>O<sup>+</sup>, <sup>15</sup>N<sup>14</sup>N<sup>16</sup>O<sup>+</sup>, <sup>14</sup>N<sup>14</sup>N<sup>17</sup>O<sup>+</sup>, and <sup>14</sup>N<sup>14</sup>N<sup>18</sup>O<sup>+</sup> relative to <sup>14</sup>N<sup>14</sup>N<sup>16</sup>O<sup>+</sup>. In Eq. (4), A = 0.00937035 and γ = 0.516 (Kaiser et al., 2003). In Eq. (5), δ<sup>15</sup>N<sup>bulk</sup> denotes the average isotope ratios for <sup>15</sup>N/<sup>14</sup>N. The subscripts “sample” and “std” respectively denote the isotope ratios for the sample and the standard. International standards for N and O isotope ratios are,  
 25 respectively, atmospheric N<sub>2</sub> and standard mean ocean water (SMOW).

## 2.5 Operation by PC software

A personal computer (NI PXI-1042Q with a controller NI PXI-8196 and I/O boards NI PXI-6221, NI PXI-4351, NI PXI-8421, and NI PXI-6514; National Instruments Corp., Austin, TX) and programming software (LabVIEW Ver. 8.2; National Instruments Corp., Austin, TX) were used to activate the solenoid valves that regulate compressed air for air-actuated shut-



off valves and the air-actuated up–down stages, the multi/two-position switching valves, the vacuum pump system, the automatic liquid nitrogen supply system, and the temperature controller for the heaters. The PC also received analogue data from the pressure and vacuum gauges and from the manometer, received the TCD signal, and synchronized the GC-IRMS data acquisition with the end of the sample preparation procedure. The timing of each regulation function is presented in Fig.

5 2.  
The program developed in this study includes a special algorithm to adapt to samples of various types. As described briefly in Sect. 2.1, it has a user interface to obtain information related to the sample: the flask volume and the sample size to be injected. When the actual inner pressure is measured and obtained, it automatically determines the optimal procedure for sample injection. Figure 3 shows a flow chart for the algorithm.

## 10 3 Results and discussion

### 3.1 Sample injection to the system

The time required for sample injection depends on the “sample expanding option” (see section 2.1). It takes about five minutes when a 300 cm<sup>3</sup> aliquot of air, which contains ca. 4 nmol of N<sub>2</sub>O in the case of ambient air (ca. 320 nmol mol<sup>-1</sup>), is injected from a 1-L flask pressurized to about 2.5 atm (option #7). When the flask volume or inner pressure is lower, more  
15 time is needed because the number of repetitive sample diffusion steps increases and one valve was operated 5 s after actuating another valve to equilibrate the pressure in the inlet line and to avoid potential fractionation of isotopocules.

When a smaller sample with high N<sub>2</sub>O concentration is measured, sample injection is completed in a minute or less. However, the performance of quantitative sample injection becomes poor for samples with more than 10 μmol mol<sup>-1</sup> N<sub>2</sub>O because the system cannot fully adjust the introduction of a small amount of sample (< 10 cm<sup>3</sup>). Moreover, the relative error  
20 of the pressure measurement becomes greater for low pressure. Such highly concentrated samples are better introduced after dilution with N<sub>2</sub> or He using another flask. Modification of one glass bottle (e.g., bottle C) to enable manual injection with a microsyringe is also possible.

### 3.2 Concentration, purification, and cryofocusing of N<sub>2</sub>O

During cryogenic concentration of N<sub>2</sub>O, the flow rate and flowing time of the He carrier gas should be optimized carefully to  
25 ensure the perfect recovery of N<sub>2</sub>O and to minimize the accumulation of blank or contamination from the system. Our preliminary tests showed that each glass bottle (C, D, or E) is purged completely when the total volume of He is more than twice the bottle volume. This result indicates that laminar flow is predominant in the bottle. Turbular flow, which is expected to cause exponential dilution and to result in the consumption of a larger amount of He to sweep out the initial sample gas, is negligible.

30 The main purpose of the purification step is separation of N<sub>2</sub>O from CO<sub>2</sub> and compounds that are less volatile than N<sub>2</sub>O. CO<sub>2</sub> is a thousand times more abundant than N<sub>2</sub>O in ambient air samples. Its isotopocules have the same mass as those of N<sub>2</sub>O.



Therefore, it often interferes with mass spectrometric analysis of N<sub>2</sub>O molecular ions. We tested two column packing materials for this purpose: Porapak Q and silica gel (dimension of the column was identical to that of Porapak Q, 60–80 mesh; GL Sciences Inc., Tokyo, Japan.). Although silica gel has the unique property of eluting N<sub>2</sub>O before CO<sub>2</sub>, their separation took longer than in Porapak Q. It was not complete even at 50°C with the flow rate of 15–50 cm<sup>3</sup> min<sup>-1</sup>. We also strove to separate CO<sub>2</sub> and N<sub>2</sub>O without using chemical adsorbents, which revealed a condition under which CO<sub>2</sub> and N<sub>2</sub>O are separated almost completely in preliminary experiments using a thermal conductivity detector and a mixture of CO<sub>2</sub> and N<sub>2</sub>O in N<sub>2</sub> bath gas (mixing ratios of CO<sub>2</sub> and N<sub>2</sub>O were ca. 250 μmol mol<sup>-1</sup>). However, a small CO<sub>2</sub> peak was observed on the m/z 44 chromatogram after separation by the second GC. Mass ratios <sup>45</sup>R and <sup>46</sup>R showed dependence on the area of CO<sub>2</sub> peak, which indicates that separation on the first column was insufficient for precise isotopocule ratio analysis of N<sub>2</sub>O. Therefore, we inserted the chemical CO<sub>2</sub> trap before cryogenic concentration. Volatile compounds such as halocarbons and hydrocarbons have longer retention time than that of N<sub>2</sub>O on columns typically used for N<sub>2</sub>O analysis. Some of them are known to hamper the chromatography of successive runs caused by their very slow elution (Roeckmann et al., 2002). Similar to previous studies, such compounds were prevented from being transferred to the next step and were backflushed to vent by switching the flow path in the present system. The cryofocusing step was necessary to inject the N<sub>2</sub>O purified in the high-flow system to the low-flow capillary GC-IRMS system. It was optimized so that eluent from the first columns were trapped only in a short time band of N<sub>2</sub>O peak and so that N<sub>2</sub>O is recovered completely.

### 3.3 Optimization of GC-IRMS analysis and measurement precision

We tested two fused-silica capillary columns for the separation of N<sub>2</sub>O from other constituents in the second GC, a porous polymer PLOT column (HP PLOT Q, 0.32 mm i.d., 20 μm film thickness, 30 m; Agilent Technologies Inc.) and a PLOT column with a monolithic carbon layer (GS Carbon PLOT). The latter column was found to have benefits for the separation of CO<sub>2</sub>, N<sub>2</sub>O, and other interfering compounds such as fluorinated hydrocarbons (Fig. 4). A shortcoming was that the retention time of N<sub>2</sub>O at the optimized condition became longer than that obtained with the porous polymer column, which was used in previous studies (Potter et al., 2013; Röckmann et al., 2003b). The degrees for precision of the measurements was evaluated with the standard deviation of repeated analyses ( $n = 3$ ) of synthetic air (349 nmol mol<sup>-1</sup> N<sub>2</sub>O) pressurized in an aluminum cylinder that had been calibrated against the international isotopic standard and which was used as a working standard (Toyoda et al., 2013). As presented in Fig. 5, precision of δ<sup>15</sup>N<sup>bulk</sup>, δ<sup>18</sup>O, and SP values measured on a single day is typically better than 0.1‰, 0.2‰, and 0.5‰, respectively, when more than 4 nmol (which corresponds to about 300 cm<sup>3</sup> of the synthetic air) of N<sub>2</sub>O is injected. The peak area of major ions m/z 44 and 30 showed good linearity with respect to the sample size (data not shown). The N<sub>2</sub>O concentration was obtained by comparison of the peak area normalized to the specific sample size between the sample and the laboratory standard. The resulting precision of the concentration measurement is better than 0.5% (coefficient of variation,  $n = 3$ ). In addition,



measured isotopocule ratios are independent of the sample size of 4–8 nmol. Results show that routine analyses of atmospheric air samples can be conducted with samples of 320 cm<sup>3</sup> so that measurements of each sample are sandwiched by those of the working standard (Table 2).

The performance of the developed system is presented along with that of previous works in Table 1. The precision and required sample size of this work is comparable to that of similar automated GC-IRMS systems. It takes 40 min for a single run, which means that a total of 80 min is necessary to obtain a single set of  $\delta^{15}\text{N}^{\text{bulk}}$ ,  $\delta^{18}\text{O}$ , and SP on some mass spectrometers that are incapable of simultaneous monitoring of five ions ( $m/z$  44, 45, 46, 30, and 31). This might be a shortcoming of the present system, but it presents advantages in terms of flexibility of the sample pressure and sample size.

#### 4 Conclusions

A fully automated sample preparation system was developed for measurement of concentrations and isotopocule ratios of N<sub>2</sub>O in both pressurized and sub-atmospheric pressure samples. An ambient atmospheric sample of 320 cm<sup>3</sup> can be analyzed in 40 min with precision of < 0.5% (coefficient of variation) for concentration, < 0.1‰ (1 standard deviation) for  $\delta^{15}\text{N}^{\text{bulk}}$ , < 0.2‰ for  $\delta^{18}\text{O}$ , and < 0.5‰ for <sup>15</sup>N-site preference (SP). The system, not being limited to use for mass spectrometric analysis, can also be applied to concentration or isotopic analyses of other trace gases such as CO<sub>2</sub> and CH<sub>4</sub> by replacing the chemical trap, GC columns, and cryogenic concentration/focusing traps and by re-optimizing the temperature, flow rate, and flow switch conditions.

Unlike previously reported systems, this system enables analysis of grab-sampled air samples that are collected into a pre-evacuated container at atmospheric pressure. This capability is particularly valuable when compressors or pumps cannot be used for sampling because of logistic reasons such as electric power or weight.

#### 20 Acknowledgements

The authors thank N. Kuroki and Y. Watanabe for their assistance in the developing and optimizing the system. This work was conducted as part of the “Studies on greenhouse gas cycles in the Arctic and their responses to climate change” under the GRENE Arctic Climate Change Research Project, and also financially supported by JSPS KAKENHI Grant Numbers 17GS0203 and 23224013.

#### 25 References

Ciais, P., Sabine, C., Bala, G., Bopp, L., Brovkin, V., Canadell, J., Chhabra, A., DeFries, R., Galloway, J., Heimann, M., Jones, C., Quéré, C. L., Myneni, R. B., Piao, S., and Thornton, P.: Carbon and other biogeochemical cycles. In: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the





- Intergovernmental Panel on Climate Change, Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M. (Eds.), Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Harris, E., Nelson, D. D., Olszewski, W., Zahniser, M., Potter, K. E., McManus, B. J., Whitehill, A., Prinn, R. G., and Ono, S.: Development of a Spectroscopic Technique for Continuous Online Monitoring of Oxygen and Site-Specific Nitrogen Isotopic Composition of Atmospheric Nitrous Oxide, *Anal. Chem.*, 86, 1726-1734, 2014.
- Hartmann, D. L., Tank, A. M. G. K., Rusticucci, M., Alexander, L. V., Brönnimann, S., Charabi, Y., Dentener, F. J., Dlugokencky, E. J., Easterling, D. R., Kaplan, A., Soden, B. J., Thorne, P. W., Wild, M., and Zhai, P. M.: Observations: Atmosphere and Surface. In: *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M. (Eds.), Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Ishijima, K., Sugawara, S., Kawamura, K., Hashida, G., Morimoto, S., Murayama, S., Aoki, S., and Nakazawa, T.: Temporal variations of the atmospheric nitrous oxide concentration and its  $\delta^{15}\text{N}$  and  $\delta^{18}\text{O}$  for the latter half of the 20th century reconstructed from firn air analyses, *J. Geophys. Res.: Atmos.*, 112, D03305, 2007.
- Kaiser, J., Röckmann, T., and Brenninkmeijer, C. A. M.: Complete and accurate mass spectrometric isotope analysis of tropospheric nitrous oxide, *J. Geophys. Res.*, 108(D15), 4476, 2003.
- Mohn, J., Guggenheim, C., Tuzson, B., Vollmer, M. K., Toyoda, S., Yoshida, N., and Emmenegger, L.: A liquid nitrogen-free preconcentration unit for measurements of ambient N<sub>2</sub>O isotopomers by QCLAS, *Atmospheric Measurement Techniques*, 3, 609-618, 2010.
- Mohn, J., Tuzson, B., Manninen, A., Yoshida, N., Toyoda, S., Brand, W. A., and Emmenegger, L.: Site selective real-time measurements of atmospheric N<sub>2</sub>O isotopomers by laser spectroscopy, *Atmospheric Measurement Techniques*, 5, 1601-1609, 2012.
- Myhre, G., Shindell, D., Bréon, F.-M., Collins, W., Fuglestedt, J., Huang, J., Koch, D., Lamarque, J.-F., Lee, D., Mendoza, B., Nakajima, T., Robock, A., Stephens, G., Takemura, T., and Zhang, H.: Anthropogenic and Natural Radiative Forcing. In: *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M. (Eds.), Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Potter, K. E., Ono, S., and Prinn, R. G.: Fully automated, high-precision instrumentation for the isotopic analysis of tropospheric N<sub>2</sub>O using continuous flow isotope ratio mass spectrometry, *Rapid Commun. Mass Spectrom.*, 27, 1723-1738, 2013.
- Röckmann, T., Kaiser, J., and Brenninkmeijer, C. A. M.: The isotopic fingerprint of the pre-industrial and the anthropogenic N<sub>2</sub>O source, *Atmos. Chem. Phys.*, 3, 315-323, 2003a.



- Röckmann, T., Kaiser, J., Brenninkmeijer, C. A. M., and Brand, W. A.: Gas chromatography/isotope-ratio mass spectrometry method for high-precision position-dependent  $^{15}\text{N}$  and  $^{18}\text{O}$  measurements of atmospheric nitrous oxide, *Rapid Commun. Mass Spectrom.*, 17, 1897-1908, 2003b.
- Röckmann, T. and Levin, I.: High-precision determination of the changing isotopic composition of atmospheric  $\text{N}_2\text{O}$  from 1990 to 2002, *J. Geophys. Res.*, 110, D21304, 2005.
- Ravishankara, A. R., Daniel, J. S., and Portmann, R. W.: Nitrous oxide ( $\text{N}_2\text{O}$ ): The dominant ozone-depleting substance emitted in the 21st century, *Science*, 326, 123, 2009.
- Sowers, T., Rodebaugh, A., Yoshida, N., and Toyoda, S.: Extending records of the isotopic composition of the atmospheric  $\text{N}_2\text{O}$  back to 1800 A.D. from air trapped in snow at the South Pole and the Greenland Ice Sheet Project II ice core, *Global Biogeochem. Cycles*, 16(4), 1129, 2002.
- Toyoda, S., Koba, K., and Yoshida, N.: Isotopocule analysis of biologically produced nitrous oxide in various environments, *Mass Spectrometry Reviews*, doi: 10.1002/mas.21459, 2015a. 2015a.
- Toyoda, S., Kuroki, N., Yoshida, N., Ishijima, K., Tohjima, Y., and Machida, T.: Decadal time series of tropospheric abundance of  $\text{N}_2\text{O}$  isotopomers and isotopologues in the Northern Hemisphere obtained by using the long-term observation at Hateruma Island, Japan, *J. Geophys. Res.*, 118, 3369-3381, 2013.
- Toyoda, S. and Yoshida, N.: Determination of nitrogen isotopomers of nitrous oxide on a modified isotope ratio mass spectrometer, *Anal. Chem.*, 71, 4711-4718, 1999.
- Toyoda, S., Yoshida, N., and Koba, K.: Isotopocule analysis of biologically produced nitrous oxide in various environments, *Mass Spectrom. Rev.*, doi: 10.1002/mas.21459, 2015b.
- Toyoda, S., Yoshida, N., Urabe, T., Aoki, S., Nakazawa, T., Sugawara, S., and Honda, H.: Fractionation of  $\text{N}_2\text{O}$  isotopomers in the stratosphere, *J. Geophys. Res.*, 106, 7515-7522, 2001.
- Tuzson, B., Mohn, J., Zeeman, M. J., Werner, R. A., Eugster, W., Zahniser, M. S., Nelson, D. D., McManus, J. B., and Emmenegger, L.: High precision and continuous field measurements of  $\delta^{13}\text{C}$  and  $\delta^{18}\text{O}$  in carbon dioxide with a cryogen-free QCLAS, *Applied Physics B* 92, 451-458, 2008.



**Table 1: Comparison of the analytical precision obtained in this study and those reported in the literature.**

Reference	Sample size (mL of ambient air)	Precision (1s) (‰)					Analytic al time (min)	Notes
		$\delta^{15}\text{N}^{\text{bulk}}$	$\delta^{15}\text{N}^{\alpha}$	$\delta^{15}\text{N}^{\beta}$	SP	$\delta^{18}\text{O}$		
Toyoda et al. (2001)	100	0.1–0.5	0.5–1	0.5–1	1–2	0.1–0.5	25	Manual system with MAT252 ( $n = 3$ )
Toyoda et al. (2013)	300	0.1	0.3	0.4	0.6	0.3	35	Manual system with MAT252 ( $n = 3$ )
Roeckmann et al. (2003)	125–167	0.1	0.3*	0.4 <sup>#</sup>	0.6 <sup>#</sup>	0.2	n.a. <sup>†</sup>	Automated system with Delta Plus XL ( $n = 5–20$ )
Roeckmann and Levin (2005)	333	0.06	n.a.	n.a.	n.a.	0.09	20	Automated system with Delta Plus XP
Mohn et al. (2010)	10,000	n.a.	0.24	0.17	0.29 <sup>#</sup>	n.a.	ca. 30	Automated system with quantum cascade laser ( $n = 136$ )
Potter et al. (2013)	420	0.05	0.11	0.14	0.21	0.10	n.a.	Fully automated system with MAT253 ( $n = 3–5$ )
This work	320	0.09	0.19	0.30	0.45	0.23	40	Fully automated system with MAT252 ( $n = 3$ )

\*Obtained with 420 cm<sup>3</sup> air. <sup>#</sup>Estimated from the reported precision for  $\delta^{15}\text{N}^{\alpha}$ ,  $\delta^{15}\text{N}^{\beta}$ , or  $\delta^{15}\text{N}^{\text{bulk}}$ . <sup>†</sup>Not available or not described.



**Table 2: Example of measurement results conducted on a single day.**

Measurement #	Sample	$\delta^{15}\text{N}^{\text{bulk}}$	$\delta^{15}\text{N}^{\alpha}$	$\delta^{15}\text{N}^{\beta}$	$\delta^{18}\text{O}$	SP
1,2	S	-2.69	-4.17	-1.20	21.22	-2.97
3,4	X1	6.58	17.18	-4.02	43.05	21.20
5,6	X1	6.80	16.63	-3.04	43.86	19.68
7,8	S	-2.62	-4.16	-1.08	21.34	-3.08
9,10	X2	7.71	18.51	-3.09	44.54	21.61
11,12	X2	7.96	18.50	-2.59	44.45	21.09
13,14	S	-2.58	-4.43	-0.73	21.28	-3.70
15,16	X3	6.02	15.22	-3.17	43.15	18.39
17,18	X3	6.18	16.20	-3.85	43.68	20.05
Average	S	-2.63	-4.25	-1.00	21.28	-3.25
SD	( $n = 3$ )	0.05	0.15	0.25	0.06	0.39
	X1	6.69	16.91	-3.53	43.46	20.44
	( $n = 2$ )*	0.11	0.28	0.49	0.41	0.76
	X2	7.83	18.51	-2.84	44.49	21.35
	( $n = 2$ )*	0.12	0.01	0.25	0.05	0.26
	X3	6.10	15.71	-3.51	43.41	19.22
	( $n = 2$ )*	0.08	0.49	0.34	0.27	0.83

\* Difference/2 is shown

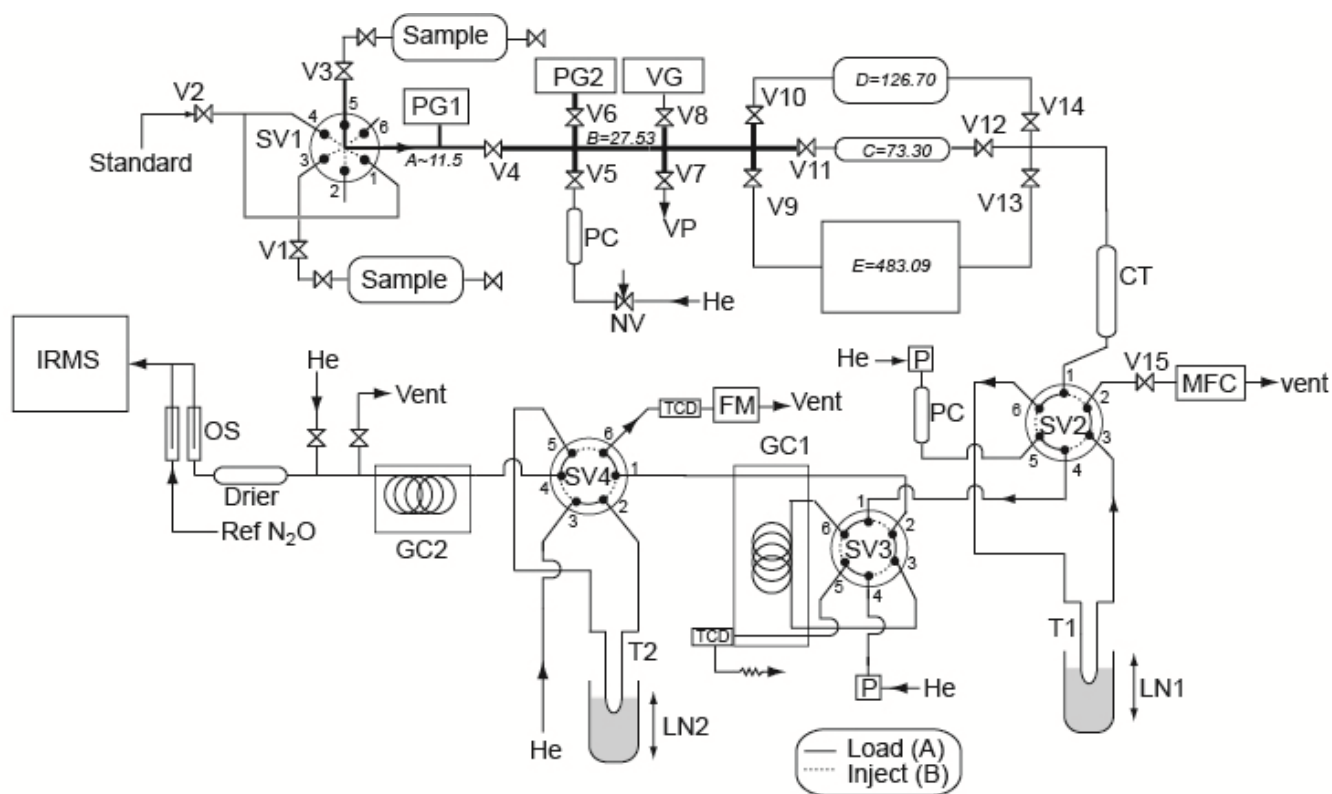


Figure 1: Schematic portraying the sample preparation system developed in this study: CT, chemical trap; FM, flow monitor; GC, gas chromatograph; IRMS, isotope ratio mass spectrometer; LN, liquid nitrogen; MFC, mass flow controller; NV, needle valve; OS, open split interface; P, pressure regulator; PC, purification column; PG, pressure gauge; SV, electrically actuated switching valve; T, trap; V, air-actuated diaphragm valve; VG, vacuum gauge; VP, vacuum pump. A–E denote parts of the vacuum line or glass bottles that are used to expand the sample, the volume ( $\text{cm}^3$ ) of which is also shown.

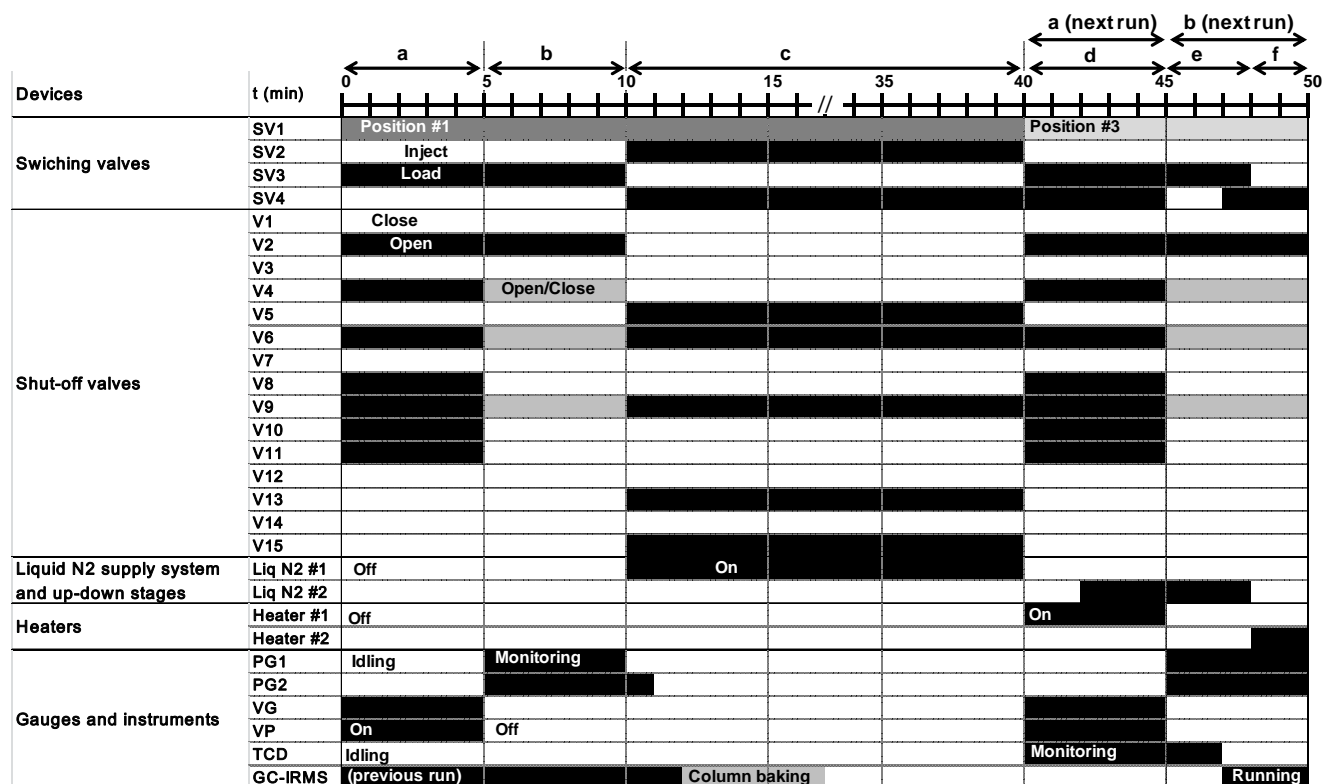


Figure 2: Time sequence of the sample preparation procedure. Horizontal arrows on the top indicate the periods for evacuation of the inlet line (a), sample injection (b), cryogenic concentration of N<sub>2</sub>O on T1 (c), purification of N<sub>2</sub>O by GC1 (d), cryofocusing of N<sub>2</sub>O on T2 (e), and injection of N<sub>2</sub>O into GC2 (f). See also Fig. 1.

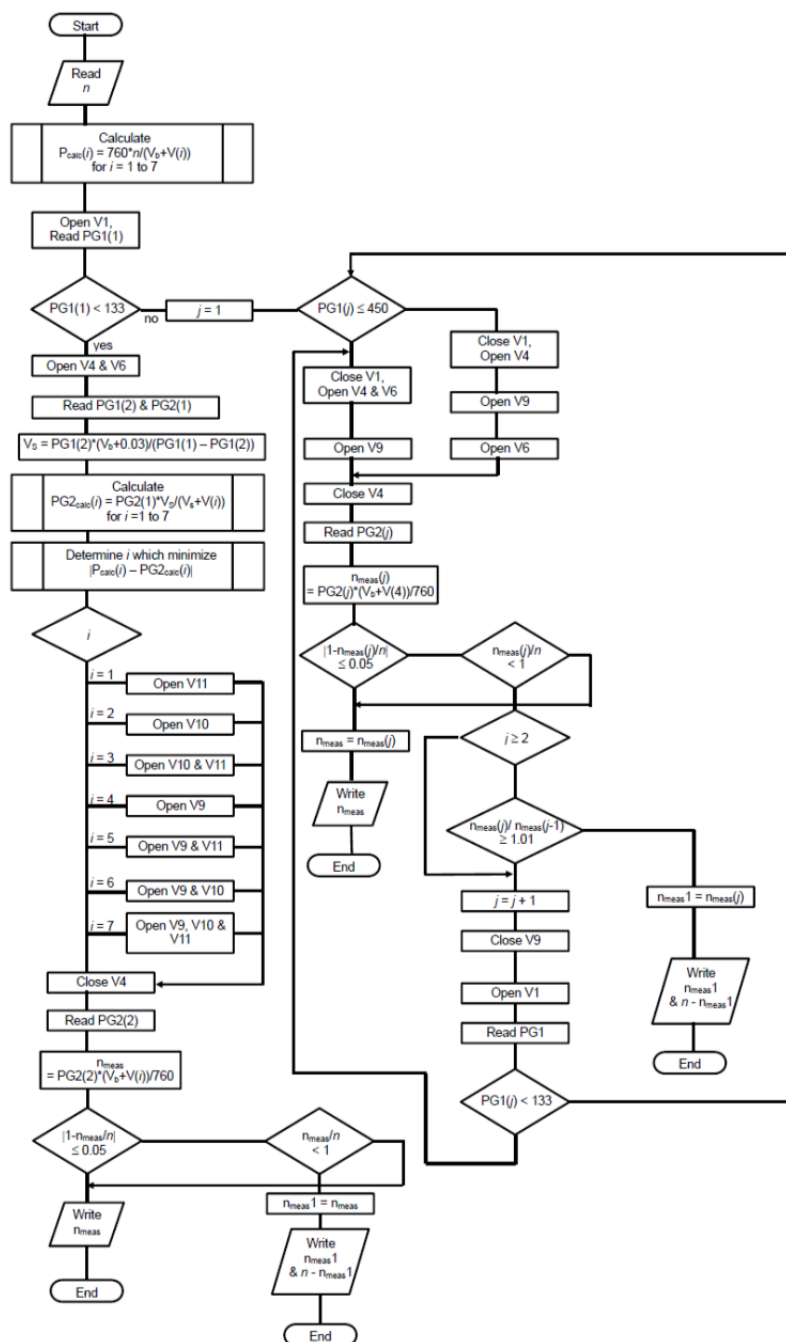


Figure 3: Flow chart of the algorithm for sample injection. “ $n$ ” denotes the sample size in cubic centimeters at 25°C and 1 atm. “ $V_b$ ” is partial volume ( $\text{cm}^3$ ) of the vacuum line indicated by “B” in Fig. 1 and “ $V(i)$ ” is calibrated volume that corresponds to the sample expanding option  $i$ . “ $PG1(j)$ ” and “ $PG2(j)$ ” are output of pressure gauges 1 (in kPa) and 2 (in Torr) (Fig. 1).

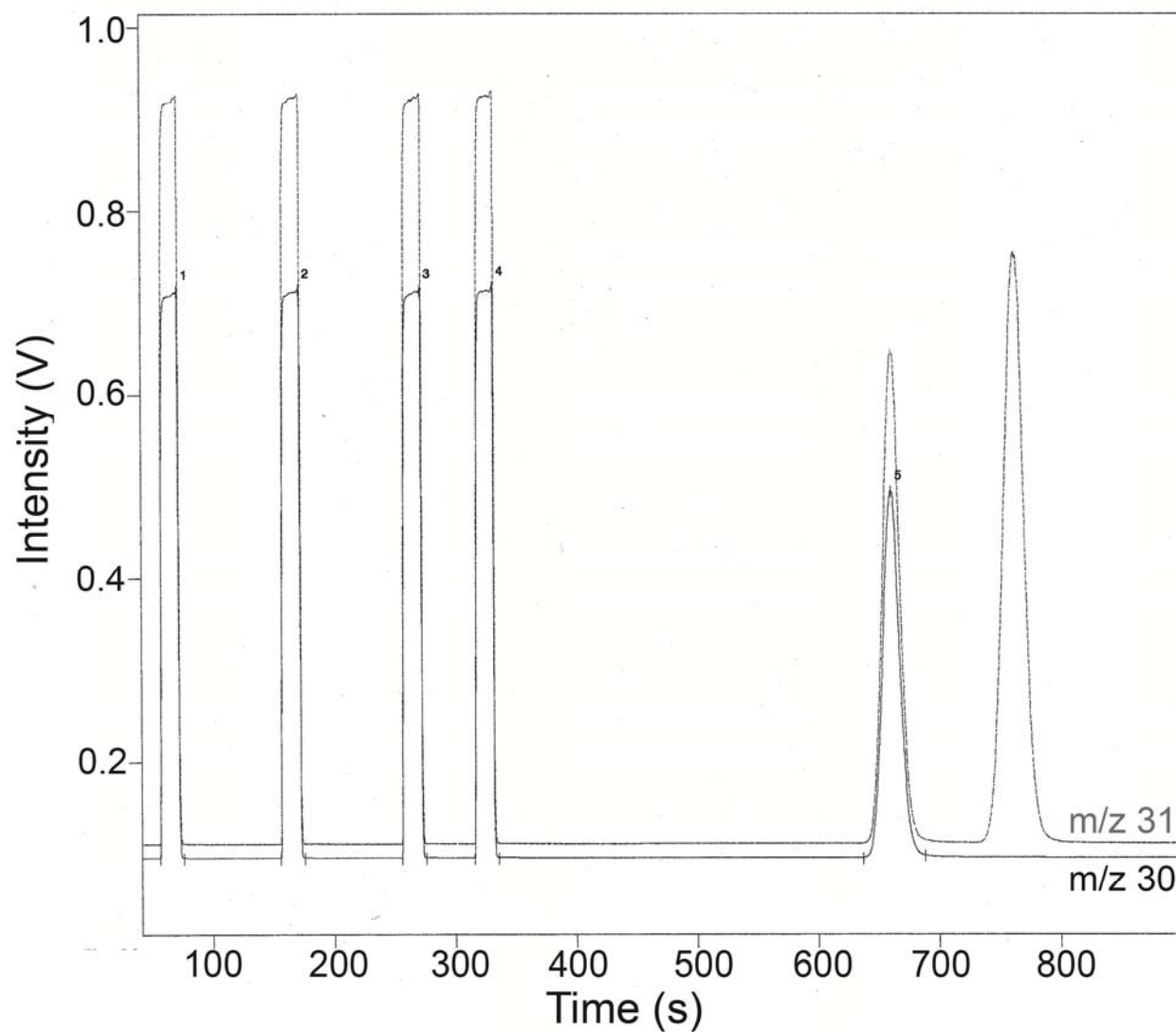


Figure 4: Typical chromatogram of background air sample obtained in a fragment-ion ( $\text{NO}^+$ ) monitoring run. Time 0 corresponds to the start of heating of the cryofocusing trap. After the peaks of reference  $\text{N}_2\text{O}$  injected from GC-IRMS interface (#1–#4), sample  $\text{N}_2\text{O}$  peak appears (#5). The peak elution about 100 s later is detected only on m/z 31 trace and is  $\text{CF}^+$  derived from a fluorinated carbon species.

5



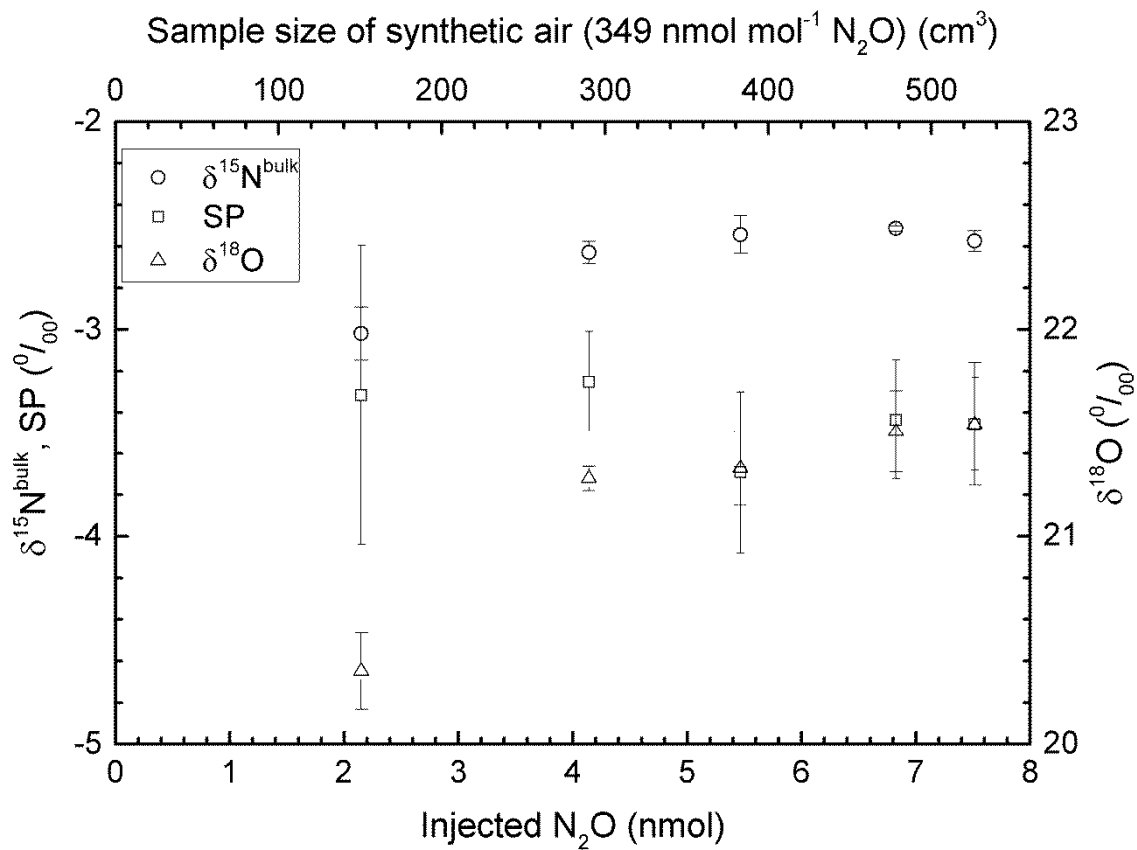


Figure 5: Precision of isotopic ratio measurements as a function of sample size.