



1 **A high-resolution time-of-flight chemical ionization mass spectrometer**
2 **utilizing hydronium ions (H_3O^+ ToF-CIMS) for measurements of**
3 **volatile organic compounds in the atmosphere**

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16



17 **Abstract**

18 Proton transfer reactions between hydronium ions (H_3O^+) and volatile organic
19 compounds (VOCs) provide a fast and high sensitive measurement technique for VOCs,
20 leading to extensive use of proton-transfer-reaction mass spectrometry (PTR-MS) in
21 atmospheric research. Based on the same ionization approach, we describe the
22 development of a high-resolution (HR) time of flight chemical ionization mass
23 spectrometer (ToF-CIMS) utilizing H_3O^+ as the reagent ions. The new H_3O^+ ToF-CIMS
24 has sensitivities of 100-1000 cps/ppb (ion counts per second per part-per-billion mixing
25 ratio of VOC) and detection limits of 20-600 ppt at 3σ for a 1-second integration time for
26 simultaneous measurements of many VOC species of atmospheric relevance. Compared
27 with similar instruments with quadrupole mass spectrometer, e.g. proton-transfer-reaction
28 mass spectrometers, the ToF analyzer with mass resolution ($m/\Delta m$) of up to 6000 not
29 only increases measurement frequency of the instrument, but also expands the number of
30 measurable species. The humidity dependence of the instrument was characterized for
31 various VOC species and the behaviors for different species can be explained by
32 compound-specific properties that affect the ion chemistry. The new H_3O^+ ToF-CIMS
33 was successfully deployed on the NOAA WP-3D research aircraft for the SONGNEX
34 campaign in spring of 2015. The measured mixing ratios of several aromatics from the
35 H_3O^+ ToF-CIMS agreed within $\pm 10\%$ with independent gas chromatography (GC)
36 measurements from whole air samples. Initial results from the SONGNEX measurements
37 demonstrate that the H_3O^+ ToF-CIMS dataset will be valuable for the identification and
38 characterization of emissions from various sources, investigation of secondary formation
39 of many photochemical organic products and therefore the chemical evolution of gas-
40 phase organic carbon in the atmosphere.

41



42 **1. Introduction**

43 Volatile organic compounds (VOCs) are ubiquitous in the atmosphere. The
44 oxidation of VOCs contributes to formation of ozone (Atkinson, 2000) and secondary
45 organic aerosol (SOA) (Hallquist et al., 2009). Accurate measurements of VOCs in the
46 atmosphere are essential to understand their sources, chemical transformations and hence
47 their effects on ozone and SOA formation (Isaksen et al., 2009). Measurements of VOCs
48 require techniques with high time resolution to effectively capture their high variability in
49 the atmosphere, especially for measurements performed on mobile platforms.

50 Proton transfer reaction mass spectrometry (PTR-MS) has been an important
51 scientific tool for VOC measurements, associated with high sensitivity and fast time
52 response (de Gouw and Warneke, 2007;Lindinger et al., 1998;Blake et al., 2009). The
53 first generation of PTR-MS instruments used quadrupole mass spectrometers (QMS) for
54 the detection of reagent and product ions. For atmospheric measurements, PTR-QMS
55 instruments are usually used in selected-ion mode, in which the mass filter steps through
56 several pre-selected masses consecutively each with a dwell time of 0.1-1 s (Warneke et
57 al., 2015). The measurement frequency for each species is thus determined by the total
58 cycle length, which is usually on the order of several seconds to 1 minute dependent on
59 user settings. The datasets produced from PTR-QMS are usually disjunct and the
60 information for masses beyond the selected ones is not available. The time resolution of
61 PTR-QMS is generally sufficient for ground-based measurements targeting VOC mixing
62 ratios, but not ideal for aircraft measurements or eddy covariance flux measurements of
63 more than a few species. In recent years, other mass analyzer, e.g. ion traps (Warneke et
64 al., 2005;Mielke et al., 2008;Steeghs et al., 2007) and time of flight (ToF) mass
65 spectrometers (Blake et al., 2004;Ennis et al., 2005;Tanimoto et al., 2007), were used in
66 PTR-MS to overcome some of these inherent drawbacks in PTR-QMS.

67 PTR-MS instruments that use time-of flight mass spectrometers (PTR-TOF) were
68 made commercially available in 2009 by Ionicon Analytik (Jordan et al., 2009). The
69 Ionicon PTR-TOF, equipped with Toferk orthogonal acceleration reflectron TOF-MS
70 (HTOF), had both much higher sensitivities (10-50 cps/ppbv) and better mass resolution
71 ($m/\Delta m \sim 6000$) than earlier PTR-TOF instruments (< 5 cps/ppbv and $m/\Delta m = 100-1200$)
72 (Jordan et al., 2009). Recently, a newer version of the Ionicon PTR-TOF equipped with a



73 quadrupole ion guide (PTR-QiTOF) was developed and is more sensitive (by a factor of
74 ~25) than the previous version of Ionicon PTR-TOF that uses an electrostatic ion lens
75 system (Sulzer et al., 2014). The performance of these PTR-TOF instruments surpasses
76 those of PTR-QMS in several ways due to the ability of the ToF to obtain whole mass
77 spectra and their high mass resolution. The high mass resolution of the ToF analyzer
78 facilitates separation of isobaric isomers (Graus et al., 2010; Sulzer et al., 2014), which
79 increases the number of measurable VOC species and reduces the possible chemical
80 interferences (Warneke et al., 2015). PTR-TOF has been used successfully in several
81 field campaigns at ground sites to measure concentrations and fluxes of a large suite of
82 VOC species (Park et al., 2013; Müller et al., 2010; Kaser et al., 2013; Holzinger et al.,
83 2013). Recently, an aircraft-deployable PTR-TOF with mass resolution up to ~1000 was
84 developed and successfully deployed on the NASA P-3B research aircraft during the
85 DISCOVER-AQ campaign (Müller et al., 2014a). The deployment of this PTR-TOF has
86 demonstrated many advantages over PTR-QMS for aircraft measurements, such as much
87 higher time resolution and hence better spatial resolution, and separation of several
88 isobaric masses in the lower mass ranges (Müller et al., 2014a).

89 In this study, a high-resolution ToF-CIMS utilizing hydronium ions (H_3O^+) similarly
90 to PTR-MS was developed based on the commercial Aerodyne ToF-CIMS (Lee et al.,
91 2014). This instrument contains two quadrupole ion guides to transport ions from the drift
92 tube reaction region to the time-of-flight mass analyzer. The quadrupole ion guides, as
93 demonstrated in the PTR-QiTOF recently (Sulzer et al., 2014), provide better
94 transmission efficiency for the reagent and product ions than the conventional ion lens
95 system. The new instrument was deployed onboard the NOAA WP-3D research aircraft
96 during the SONGNEX campaign in spring of 2015. Here, we will present development
97 and characterization of the instrument and its performance during the SONGNEX
98 campaign.

99 **2. H_3O^+ TOF-CIMS instrument description**

100 A commercial Aerodyne chemical ionization mass spectrometer (CIMS) (Lee et al.,
101 2014) was used to construct the airborne H_3O^+ ToF-CIMS (Figure 1). Briefly, the
102 Aerodyne ToF-CIMS consists of (1) an ion-molecule reaction (IMR) chamber, (2) a
103 small segmented radio frequency (RF) only quadrupole ion guide that is used as



104 collisional dissociation chamber (SSQ), (3) a second big segmented RF-only quadrupole
105 (BSQ), (4) a series of DC optics that further focus and accelerate the primary beam (PB),
106 and (5) a high-resolution ToF detector (HTOF, ToFwerk AG, Switzerland) (Bertram et
107 al., 2011). In the new H_3O^+ ToF-CIMS, the IMR chamber was replaced by a drift tube
108 (Figure 1), which provides a homogenous and controllable electric field for ion-molecule
109 reactions. The drift tube is comprised of stainless steel rings separated by Teflon rings for
110 both vacuum sealing and electronic insulation (de Gouw and Warneke, 2007). A hollow
111 cathode discharge ion source was connected in front of the drift tube to produce high
112 purity hydronium ions by introducing a flow of 5 mL min^{-1} of water vapor. The detailed
113 description for both drift tube and ion source can be found in previous review papers (de
114 Gouw and Warneke, 2007; Blake et al., 2009). In addition to the 3-stage split-flow turbo
115 pump (Pfeiffer SplitFlow 310) used for the high vacuum of BSQ, PB and TOF, another
116 turbo pump (Pfeiffer TMH-071-P) was installed to draw air from the SSQ and water
117 vapor from the ion source. A butterfly exhaust throttle valve (MKS T3Bi) was used to
118 control the pumping rate of the newly installed turbo pump (Figure 1), and the pressure of
119 the SSQ can be adjusted by the opening percentage of the valve. In practice, the pressure
120 of drift tube is also actively controlled (see below) and SSQ pressure only needs to be
121 adjusted sporadically.

122 The newly built H_3O^+ ToF-CIMS was deployed on the NOAA WP-3D aircraft in
123 March-April, 2015 as part of the Shale Oil and Natural Gas Nexus (SONGNEX)
124 campaign. During SONGNEX, a total of 19 flights were carried out mainly over oil and
125 gas production basins in the western United States. For the detailed flight strategies, the
126 readers are referred to the SONGNEX website
127 (www.esrl.noaa.gov/csd/projects/songnex).

128 A heated Teflon tube ($40 \text{ }^\circ\text{C}$) mounted inside a winglet on a window plate was used
129 to transport air into the instrument (Figure S1). Pressure at the downstream end of the
130 inlet was regulated to 180 mbar using a pressure controller. A PEEK capillary was placed
131 between the pressure controller and drift tube to further reduce the pressure from 180
132 mbar to the maintained pressure in the drift tube ($2.40 \pm 0.01 \text{ mbar}$). During SONGNEX,
133 the 11 cm long drift tube had a drift voltage of 710 V and was heated to $50 \text{ }^\circ\text{C}$, which
134 resulted in an E/N ratio (electric field to number density) of 120 Td ($1 \text{ Td} = 10^{-17} \text{ V cm}^2$)



135 inside the drift tube. The flow of the inlet varied from $\sim 500 \text{ mL min}^{-1}$ at sea level to ~ 100
136 mL min^{-1} at 6.5 km of altitude. Background signals of the instrument were determined by
137 passing ambient air for 90 s every 30-40 min through a Platinum catalytic converter
138 heated to 350 °C. In-flight calibrations were performed by adding a small flow (0.5-3.0
139 mL min^{-1}) of a ten-component gas standard (see the list in Table 1) to the inlet flow
140 automatically for 90 s every 1-2 hours (Figure S1). Calibrations were conducted by
141 adding the gas standard flow to either clean air from the catalytic converter or ambient air
142 during SONGNEX. A diffusion cell introduced a small amount of 1,3,5-trichlorobenzene
143 ($\text{C}_6\text{H}_3\text{Cl}_3$) into the instrument continuously to facilitate ToF mass calibration at the high
144 ends of the m/z range. Laboratory experiments show that 1,3,5-trichlorobenzene mainly
145 generates $\text{C}_6\text{H}_3\text{Cl}_3\text{H}^+$ (m/z 180.9373, 98.5%) in the instrument, with small contributions
146 from $\text{C}_6\text{H}_3\text{Cl}_3^+$ (m/z 179.9295, 1.0%) and $\text{C}_6\text{H}_3\text{Cl}_2\text{H}^+$ (m/z 145.9685, 0.4%). The signals
147 of $\text{C}_6\text{H}_3\text{Cl}_3\text{H}^+$ ion were around 2000 cps during the SONGNEX campaign.

148 The axial voltage gradients for SSQ and BSQ were tuned using the Thuner
149 software (Tofwerk AG) by maximizing signals of the protonated product ions of several
150 VOC species (acetone, benzene and isoprene), minimizing VOC signals from charge
151 transfer reactions with O_2^+ , and maximizing mass resolution of the TOF analyzer before
152 the campaign. In addition, the effects of turning the RF amplitudes of the SSQ and BSQ
153 were explored manually (see below). The voltage of micro-channel plate (MCP) detector
154 for TOF are determined and set before takeoff for each individual flight by maintaining
155 the single ion signal (SIS) at around 1.8 mV-ns, which effectively prevents mass
156 discrimination as the result of MCP aging (Müller et al., 2014b). The extraction
157 frequency of the TOF was set at 25 kHz, which enables to measure masses up to m/z 500.
158 During SONGNEX, mass spectra were averaged and stored to 10 Hz and were further
159 averaged afterward to 1 Hz for the analysis shown in this paper.

160 ToF data were processed using the Tofware software package (v2.5.1)
161 (www.tofwerk.com/tofware) developed by Tofwerk and Aerodyne Research Inc.
162 (www.aerodyne.com). The detailed procedures of the ToF data processing have been
163 presented in a previous publication (Stark et al., 2015). Post-measurement mass
164 calibrations were conducted using 6-7 isolated masses, including m/z 19.0178 (H_3O^+),
165 m/z 29.9974 (NO^+), m/z 31.9893 (O_2^+), m/z 37.0284 ($\text{H}_3\text{O}^+(\text{H}_2\text{O})$), m/z 55.0390



166 $(\text{H}_3\text{O}^+(\text{H}_2\text{O})_2)$, m/z 125.9610 ($\text{FeH}_5\text{O}_4\text{H}^+$) and m/z 180.9373 ($\text{C}_6\text{H}_3\text{Cl}_3\text{H}^+$). Mass
167 calibration of the ToF was accurate within 5-10 ppm for various masses, similar to the
168 results reported for another Aerodyne ToF-CIMS instrument (Lee et al., 2014).
169 Instrument functions (peak shape, peak width and baseline) were derived for each
170 individual flight based on the algorithms shown in Stark et al. (2015). High-resolution
171 peak fitting to the mass spectra were performed using a user defined peak list with
172 masses up to m/z 181.

173 In the orthogonal extraction region of the TOF analyzer, ions of different masses
174 have the same energy but different velocities, and are therefore extracted at different duty
175 cycles (Chernushevich et al., 2001). The Tofware software corrects for this effect using
176 the equation of $I_{corr} = I_{raw} \times \sqrt{\frac{55}{m/z}}$ to a reference mass of m/z 55. In this study, all of the
177 signals are reported as the ToF duty cycle corrected signals (I_{corr}), unless otherwise
178 noted.

179 **3. H_3O^+ TOF-CIMS instrument performance**

180 **3.1 Quadrupole ion guides**

181 In addition to hydronium ions (H_3O^+), water clusters are also present in the
182 instrument, including protonated water dimers ($\text{H}_3\text{O}^+(\text{H}_2\text{O})$) and protonated water trimers
183 ($\text{H}_3\text{O}^+(\text{H}_2\text{O})_2$). In standard PTR-MS operation, the product ion signals are normalized to
184 the reagent ion count rates to account for drifts in the ion source. Because the reagent
185 ions in the H_3O^+ ToF-CIMS are so strongly dependent on humidity, the measured cluster
186 ion distribution needs to be understood in detail as it affects the normalization procedure
187 and therefore the measured mixing ratios. As the drift tube in H_3O^+ ToF-CIMS is
188 identical to those in PTR-MS, the cluster distribution and their humidity dependence in
189 the drift tube are expected to be similar as in PTR-MS, which has been understood well
190 and described in detail in de Gouw and Warneke (2007). However, the cluster
191 distribution can be altered in the quadrupole ion guides, because the electric field inside
192 the ion guides is not always the same as that in the drift tube. Thus, the reagent ions and
193 VOC signals as a function of various settings of the quadrupole ion guides are investigated
194 here.



195 The signals of the reagent ions as a function of the pressure inside the SSQ are
196 shown in Figure 2. As the SSQ pressure increases from 1.0 mbar to 1.5 mbar, the
197 intensities of all reagent ions increase, indicative of better ion transmission through the
198 SSQ at higher pressures. When the SSQ pressure is higher than 1.5 mbar, the H_3O^+ signal
199 starts to decline slightly and water clusters continue to increase, implying that the reagent
200 ions shift to larger clusters at these SSQ pressures. The ratios of $\text{H}_3\text{O}^+(\text{H}_2\text{O})/\text{H}_3\text{O}^+$
201 ($R_{37/19}$) increase with the SSQ pressure throughout the range studied here. At higher
202 SSQ pressure, the effective de-clustering ability or E/N ratio is reduced. The $\text{H}_3\text{O}^+(\text{H}_2\text{O})_2$
203 ions account for small a fraction of the reagent ions throughout the explored range of
204 SSQ pressures. The signals of protonated benzene and acetone at constant mixing ratios
205 (2.5 ppb) both increase with higher SSQ pressure, although the increase of benzene
206 signals is small when SSQ pressures go beyond 1.5 mbar. The protonated VOC signals
207 correlate better with H_3O^+ than with $\text{H}_3\text{O}^+(\text{H}_2\text{O})$, suggesting that H_3O^+ ions are the
208 dominant participant in proton-transfer reactions, even though H_3O^+ ions are measured at
209 comparable levels as $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ at high SSQ pressures. Along with this observation, the
210 observed slight reduction of H_3O^+ and disproportionally larger enhancement of
211 $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ with SSQ pressures above 1.5 mbar imply that the transmission of H_3O^+ in
212 the SSQ is significantly lower than that of $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ (and other higher m/z). Choosing
213 the optimum SSQ pressure therefore represents a trade-off between higher intensities of
214 the reagent (and VOC) ions and lower signals of water clusters. We have selected to run
215 the instrument with the SSQ pressure at 1.30 ± 0.01 mbar during the SONGNEX
216 campaign.

217 In addition to the SSQ pressure, the signals of reagent ions and VOC product ions
218 were also explored as a function of RF amplitudes for both SSQ and BSQ (Figure S2 and
219 Figure S3). From these experiments, the main findings are: (1) as expected, lower RF
220 amplitudes in the BSQ lead to better transmission for light ions but worse ion focusing
221 (Chernushevich et al., 2001). (2) Ion chemistry needs to be taken into account in the SSQ,
222 for example H_3O^+ declines and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ increases when RF amplitudes decrease from
223 80 V to 40 V. (3) Variations of VOC signals generally follow the reagent ions, but they
224 are influenced more by poorer focusing at low RF amplitudes, especially for heavier
225 VOC masses. (4) O_2^+ signals and charge transfer products of aromatics increase quickly



226 when RF amplitudes of the SSQ are higher than 80 V, possibly due to discharge in the
227 SSQ (TOF-CIMS manual indicates that discharge happens with >200 V). Based on these
228 results, the RF amplitudes of the SSQ and BSQ were set to 50 V and 350 V, respectively.

229 **3.2 Transmission of reagent ions and their humidity dependence**

230 The humidity of the sampled air affects the distribution of the reagent ions in PTR-
231 MS (de Gouw and Warneke, 2007). Figure 3A shows the reagent ion signals versus water
232 vapor mixing ratios (w) of the sampled air from a laboratory experiment. As expected, the
233 $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ signals increase as the air gets more humidified. The signals of $\text{H}_3\text{O}^+(\text{H}_2\text{O})_2$
234 were small, but increase quickly with humidity. However, there is also an increase of
235 H_3O^+ signals with humidity. Compared to dry air with $w=0$ g/kg, the H_3O^+ signals are
236 52% higher under the condition with $w=22.4$ g/kg equivalent to a relative humidity (RH)
237 of 90% at 25 °C. This behavior is in marked contrast to the reported dependence of the
238 reagent ions on humidity in a conventional PTR-MS using ion lenses to transfer ions (de
239 Gouw and Warneke, 2007): in these instruments, H_3O^+ decreases and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$
240 increases with rising humidity and the total intensities of the two reagent ions are
241 relatively stable with humidity with the remaining difference explainable from the
242 difference in detection efficiency between H_3O^+ and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$.

243 The distribution of H_3O^+ and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions exiting the SSQ is determined by
244 the distribution of these same ions entering the SSQ as well as the fragmentation and
245 clustering processes that modify the distribution in the SSQ. At the SSQ pressure used, an
246 individual ion entering the SSQ undergoes many collisions that can add or subtract a
247 ligand molecule. As a result, the transmission efficiency of an reagent ion in the SSQ
248 reflects the averaged transmission efficiencies of ions with m/z 19 and m/z 37 weighed
249 by the time the ion spends as H_3O^+ vs. $\text{H}_3\text{O}^+(\text{H}_2\text{O})$. This issue is qualitatively similar to
250 the measured mobilities of H_3O^+ and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ observed in the drift tube of PTR-MS
251 (Warneke et al., 2001). The larger H_3O^+ signals at higher humidity in the H_3O^+ ToF-
252 CIMS are the result of the low transmission efficiency of H_3O^+ ions compared to other
253 heavier masses, which has been inferred from the dependence of the reagent ions with
254 SSQ pressures above. The low transmission of H_3O^+ ions is related to the low-mass cut-
255 off of the RF-only quadrupoles (Chernushevich et al., 2001). The strong increase in H_3O^+
256 signal intensity with humidity reflects the fact that while ions may be detected as H_3O^+ ,



257 they spent a larger fraction of time as $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ in the SSQ at higher humidity and are
258 therefore transferred with a higher average transmission efficiency.

259 The transmission efficiency of H_3O^+ ions relative to $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ is quantified
260 using additional laboratory experiments. Methanol and acetonitrile were introduced into
261 the instrument at such high concentrations that significant fractions of the reagent ions
262 were depleted (Figure S4). Methanol and acetonitrile are used, because their product ion
263 masses (m/z 33 and m/z 42) bracket the mass of $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions. We observed more
264 product ions of both methanol (CH_4OH^+ , m/z 33.0335) and acetonitrile ($\text{C}_2\text{H}_3\text{NH}^+$, m/z
265 42.0338) than the summed depletion of H_3O^+ and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$. As the transmission factor
266 of $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions should be in between protonated methanol and acetonitrile ions, the
267 ratios between H_3O^+ changes and the changes of the sum of product ions and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$
268 ions in each depletion experiment reflect the ratios of the transmission efficiency between
269 H_3O^+ and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ($T_{\text{H}_3\text{O}^+}/T_{\text{H}_3\text{O}^+(\text{H}_2\text{O})}$). From the scatterplots in Figure S4, the
270 $T_{\text{H}_3\text{O}^+}/T_{\text{H}_3\text{O}^+(\text{H}_2\text{O})}$ ratios are determined to be in the range of 0.21-0.35 and 0.14-0.25
271 from the experiments with methanol and acetonitrile, respectively (Figure S5). The
272 discrepancies between the estimates from methanol and acetonitrile may come from the
273 differences between the transmission factors of m/z 33 and m/z 42 (Figure S6), and/or
274 small amount of methanol forming methanol-water clusters (see section 3.3). The
275 determined $T_{\text{H}_3\text{O}^+}/T_{\text{H}_3\text{O}^+(\text{H}_2\text{O})}$ ratios are larger at higher water vapor mixing ratios
276 (Figure 3B, also in Figure S5), in accordance with the expectation of a higher effective
277 transmission efficiency of H_3O^+ ions with increasing humidity discusses above.

278 The relationship of $T_{\text{H}_3\text{O}^+}/T_{\text{H}_3\text{O}^+(\text{H}_2\text{O})}$ versus water vapor mixing ratios shown in
279 Figure 3B can be used to correct the measured H_3O^+ signals to the same transmission
280 efficiency as $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ (Figure 3C). The corrected H_3O^+ signals exhibit slightly lower
281 signals at higher humidity, which is more similar to the relationship observed in the PTR-
282 MS drift tube (de Gouw and Warneke, 2007). The summed signals of the corrected H_3O^+
283 and measured $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ are also shown in Figure 3C. We see higher total reagent ion
284 signals at increasing humidity. It is clear that H_3O^+ ions dominate the reagent ions in the
285 drift tube throughout the explored humidity range, even though the measured H_3O^+
286 signals are lower than $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions at high humidity levels. At a high humidity level



287 (RH=90% at 25 °C, $w=22.4$ g/kg), $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions account for 24-30% of the total
288 reagent ions.

289 Taken together, both the signals of the reagent ions and their transmission
290 efficiency in the instrument exhibit non-linear relationships with humidity. The total
291 reagent ion signals can be derived from the determined $T_{\text{H}_3\text{O}^+}/T_{\text{H}_3\text{O}^+(\text{H}_2\text{O})}$, but it is
292 associated with significant uncertainties (Figure 3C). A simple equation to derive a
293 parameter with little humidity dependence from reagent ion signals seems to be difficult.
294 Thus, the product ion signals will be normalized to H_3O^+ signals of 10^6 cps to account for
295 drifts in the ion source. Normalization to H_3O^+ signals does not involve any mathematical
296 computation of the reagent ions, which may introduce extra uncertainty.

297 The ratios of $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions to H_3O^+ ions ($R_{37/19}$) can be used as a proxy for
298 humidity in PTR-MS (de Gouw et al., 2003a). The dependence of $R_{37/19}$ with water
299 vapor mixing ratios of the sampled air is shown in Figure 3D. We see higher $R_{37/19}$ ratios
300 with increasing humidity from both laboratory experiments and ambient measurements
301 during the SONGNEX, suggesting that $R_{37/19}$ is also a good internal humidity indicator
302 for the instrument. The agreement between laboratory experiments and ambient data is
303 good, considering that the humidity sensors used for laboratory experiments were not
304 cross-calibrated with the sensors on the NOAA WP-3D. The correlation coefficients from
305 laboratory experiments and SONGNEX ambient data are 0.98 and 0.99, respectively,
306 both indicative of tight linear relationship of the data points. The intercept in Figure 3D is
307 the result of excess water vapor entering the drift tube from the ion source. Based on the
308 linear fit to SONGNEX data points in Figure 3D, the $R_{37/19}$ ratio in the instrument is 1.4
309 at a humidity level of RH=90% at 25 °C (or $w=22.4$ g/kg), which will be used to
310 characterize the instrument response and compare with dry conditions in this study.

311 **3.3 The humidity dependence of VOC sensitivities**

312 *3.3.1 Species without significant dehydration and hydration*

313 A series of laboratory experiments were performed to describe instrument
314 sensitivities of various VOC species as a function of humidity. Figure 4 shows the results
315 for acetone and benzene from one of the experiments. Constant mixing ratios (8.0 ppb in
316 Figure 4) of acetone and benzene were introduced into the instruments and the



317 background subtracted raw signals of protonated product ions of the two compounds
318 were observed at various humidities. As described in the previous section, $R_{37/19}$ is used
319 as the indicator of humidity. Protonated acetone signals increase with rising humidity,
320 whereas protonated benzene signals decline under humidified conditions. The signals of
321 protonated acetone and protonated benzene at the humidity level of $R_{37/19}=1.4$,
322 equivalent to RH=90% at 25 °C, are 111% and 32% of those under dry condition,
323 respectively. The humidity dependencies of these two species are similar to that reported
324 for conventional PTR-MS in many previous studies (Warneke et al., 2001; de Gouw et al.,
325 2003a).

326 The normalized signals of acetone and benzene from the humidity experiments
327 are shown in Figure 4B. As illustrated in the previous section, H_3O^+ signals are 52%
328 higher at $R_{37/19}=1.4$ relative to dry condition. The increase of H_3O^+ signals with
329 humidity is apparently larger than the increase of protonated acetone, which in turn leads
330 to a reduction of the normalized protonated acetone signals with increasing humidity.
331 Because of the low transmission of H_3O^+ ions and clustering/de-clustering effects as
332 described earlier, the normalization to reagent ions leads to a different humidity
333 dependence of the normalized signals in the H_3O^+ ToF-CIMS compared to that in the
334 conventional PTR-MS.

335 The normalized signals at varying humidity levels relative to that at dry condition
336 for acetone, benzene and other VOC species are determined. After attempts using several
337 different fit functions to describe the data points, a double exponential function was
338 found to achieve the best representation of the data. The fitted humidity curves for
339 acetone, benzene and other VOC species are compiled in Figure 5. Generally, a stronger
340 humidity dependence was observed for hydrocarbons than for OVOCs. Comparing the
341 results of the aromatics (benzene, toluene, o-xylene, methylstyrene, 1,2,4-
342 trimethylbenzene and p-cymene), we see less humidity dependence for heavier aromatics.
343 The same trend is also observed for ketones (acetone, MEK, pentanone and hexanone).
344 The heavier species in the two compound series tend to have higher proton affinities (PA)
345 (Hunter and Lias, 1998), suggesting that proton affinities of the species play a role in the
346 humidity dependence.



347 The fractions of normalized signals at $R_{37/19}=1.4$ relative to dry condition
348 ($f_{R_{37/19}=1.4}$) were determined and plotted as a function of proton affinities of the species
349 in Figure 5C. A positive correlation between the determined fractions and proton
350 affinities of VOC species is generally observed. This is expected, as a higher proton
351 affinity for a VOC species has several implications: (1) the reaction with $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions
352 is exothermic when the proton affinity exceeds that of water dimer (801 kJ/mol); and (2)
353 the reaction more readily becomes exothermic through either direct proton transfer or
354 other routes (Midey et al., 2002). The data points in Figure 5C are color-coded using the
355 permanent dipole moment (μ_D) of the neutrals, which has been shown to affect the
356 efficiency of ligand switching reactions between VOC species and water clusters (Spanel
357 and Smith, 1995). It is clear that species with $\mu_D < 1$ D are more strongly humidity
358 dependent than the species with $\mu_D > 1$ D, which implies that ligand switching reactions
359 are important in the instrument as the result of substantial amounts of $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions.
360 Based on proton affinity and permanent dipole moment, the species shown in Figure 5C
361 can be divided into four groups: group I with $\text{PA} > 801$ kJ/mol and $\mu_D > 1$ D (e.g. acetone);
362 group II with $\text{PA} < 801$ kJ/mol and $\mu_D > 1$ D (e.g. acetaldehyde); group III with $\text{PA} > 801$
363 kJ/mol and $\mu_D < 1$ D (e.g. isoprene); group IV with $\text{PA} < 801$ kJ/mol and $\mu_D < 1$ D (e.g.
364 benzene). Species in group I (high PA and μ_D) can undergo proton transfer with
365 $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ and H_3O^+ at similar rate constants, whereas the reactions with $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ are
366 either not happening or inefficient for species in group IV (low PA and μ_D). For species
367 in group II (low PA, high μ_D) and III (high PA, low μ_D), ligand switching and direct
368 proton transfer reaction with $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ can occur in the instrument, respectively.

369 The inverse of H_3O^+ signal relative to dry conditions ($\frac{1}{m_{19}}$) was included in Figure
370 5 (A and B) as a reference. If the humidity dependence for a VOC species follows this
371 reference line, the raw signal of the VOC actually has no humidity dependence and the
372 lower normalized signals at higher humidity levels are solely the result of normalization
373 to higher H_3O^+ signals. Species above this reference line are associated with higher raw
374 signals at higher humidity, and vice versa. As discussed in section 3.2, the increase of
375 H_3O^+ signals with rising humidity is caused by a higher effective transmission of ions
376 detected as H_3O^+ (Figure 3B). Thus, the upper limit of the humidity dependence curve
377 should be a flat unity line in Figure 5 (A and B) (see exceptions in section 3.3.2). Many



378 heavier OVOC masses were observed during SONGNEX, but their humidity dependence
379 was not explored. As higher proton affinities are expected for these heavier OVOC
380 species, the humidity dependence curves for these heavier OVOCs should lie in the
381 shaded area filled by patterns in Figure 5 (A and B).

382 3.3.2 *Species with significant dehydration and hydration*

383 Product ions of the species explored for humidity dependence in the previous section
384 (3.3.1) are associated with minor dehydration following the proton transfer reaction.
385 Dehydration happens when the product ion of a VOC species fragments by losing one or
386 more water molecules. Propanols and higher alcohols are known to mainly yield
387 dehydration product ions in PTR-MS (Spanel and Smith, 1997; Blake et al.,
388 2009; Warneke et al., 1996). For example, isopropanol (C_3H_7OH , $PA=795.4$ kJ/mol) is
389 mainly detected as $C_3H_7^+$ (m/z 43.0542), rather than $C_3H_8OH^+$ (m/z 61.0648) in H_3O^+
390 ToF-CIMS (Figure 6A). We see higher $C_3H_8OH^+/C_3H_7^+$ ratios with increasing humidity
391 in the instrument. The measured humidity dependence of $C_3H_7^+$ ions is similar to that of
392 acetaldehyde (and other OVOCs), whereas more $C_3H_8OH^+$ ions are detected with
393 increasing humidity and the normalized signal of $C_3H_8OH^+$ ions at $R_{37/19}=1.4$ is 1.7
394 times of in dry air. Similar to i-propanol, a positive dependence of protonated ethanol on
395 humidity is also observed (Figure 6B), as ethanol fragments significantly by losing a
396 water molecule as well (Baasandorj et al., 2015).

397 In addition to dehydration, hydrated product ions in the form of $MH^+(H_2O)$ (where
398 M is the formula of VOC species) are observed, including acetaldehyde, acetone,
399 methanol and acetic acid (Figure 7A). The ratios of hydration ions to protonated
400 molecular ions ($MH^+(H_2O)/MH^+$) are observed to increase with humidity for these
401 compounds. Among the species investigated, formic acid ($HCOOH$, $PA=741.8$ kJ/mol)
402 and isocyanic acid ($HNCO$, $PA=753.1$ kJ/mol) are clustering the most, with
403 $MH^+(H_2O)/MH^+$ ratios of 0.34 and 1.2 at $R_{37/19}=1.4$, respectively. The protonated ions
404 of the two compounds both decrease quickly with humidity, and the normalized signals at
405 $R_{37/19}=1.4$ are both only approximate 2% of those at dry conditions. The signals of
406 hydration ions cannot account for the difference between the humidity dependent curves
407 of these two compounds and the region other OVOCs occupy in Figure 5. The reason for
408 the strong humidity dependence of the two compounds is not known, but it is possibly



409 due to enhanced loss of ions after hydration. A similar enhanced loss of ions at higher
410 humidity was observed for formic acid in a PTR-QMS at a low E/N ratio (85 Td), but not
411 at higher E/N ratios (Baasandorj et al., 2015). Baasandorj et al. (2015) (see Figure 5f in
412 their paper) showed that, at the low E/N ratio (85 Td), protonated ions of formic acid
413 decrease by 12 ncps/ppb and hydration ions only increase by 6 ncps/ppb going from
414 RH=18% to RH=88%. The enhanced loss of isocyanic acid at higher humidity might be
415 related to its hydrolysis, which is reported to be accelerated by water dimer and trimer
416 (Raspoet et al., 1998).

417 The examples of isopropanol, formic acid and isocyanic acid suggest that the
418 humidity dependence can be affected by dehydration and hydration processes of the
419 product ions. The abundance of water molecules affects the equilibrium ratios of the
420 respective ion pairs. In contrast, there are some species associated with significant
421 fragmentation without water molecule as a neutral product, e.g., monoterpenes and p-
422 cymene, which exhibit similar humidity dependence as other VOCs (Figure 5C).
423 Therefore, special attention should be paid to the humidity dependence of the species
424 associated with either significant fragmentation by losing water molecule or significant
425 molecule-water clusters.

426 **3.4 Mass resolving power and separation of isobaric masses**

427 The mass resolving power ($m/\Delta m$) was derived from the observed linear
428 relationship between the full-width half maximum (FWHM) of several isolated mass
429 peaks and their m/z using Tofware (Stark et al., 2015). The typical $m/\Delta m$ for the H_3O^+
430 ToF-CIMS during the SONGNEX campaign is shown in Figure S7. The $m/\Delta m$ in the
431 range of m/z 30 - 200 (where most VOCs were detected), are 3900-5900 with higher
432 resolution for heavier masses. These mass resolutions are sufficient to separate many
433 isobaric ions residing at the same nominal masses. Figure 8 shows the separation of C8
434 aromatics ($\text{C}_8\text{H}_{10}\text{H}^+$, m/z 107.0855) from benzaldehyde ($\text{C}_7\text{H}_6\text{OH}^+$, m/z 107.0491) for m/z
435 107 during a flight leg (18:15-18:45 UTC) over the Permian Basin in Texas, US. The two
436 mass spectra in A and B are 30 s averages centered around 18:25:00 and 18:35:10,
437 respectively. The two time windows had either benzaldehyde or C8 aromatics as the
438 higher peak between the two. The time series of benzaldehyde and C8 aromatics
439 determined from the mass spectral fits for this 30-min period correlated well with acetone



440 ($R=0.87$) and benzene ($R=0.94$), respectively. This is consistent with the expectation that
441 benzaldehyde was mainly from secondary formation and C8 aromatics were dominated
442 by primary emissions from oil and gas activities. The signals at m/z 107 are usually
443 assigned to C8 aromatics in PTR-QMS studies (de Gouw and Warneke, 2007). Although
444 the possible interference from benzaldehyde to C8 aromatics has been known, it usually
445 constitutes a small fraction of the total signal at nominal mass 107 (Warneke et al., 2003).
446 The example shown in Figure 8 indicates that benzaldehyde, in some environments, can
447 contribute significantly to signals at nominal mass 107.

448 Based on results in previous studies (Stark et al., 2015; Graus et al., 2010), the mass
449 resolving power of the ToF analyzer presented in this study can separate many of the
450 isobaric masses in the mass range of $m/z < 200$ by utilizing the high-resolution peak fitting
451 algorithms. The separation of isobaric ions has several advantages over the nominal mass
452 data of PTR-QMS: (1) reduce chemical interferences, e.g. the interference of
453 benzaldehyde to C8 aromatics at nominal m/z 107 as shown above; (2) decrease
454 background signals for several compounds of interest, e.g. for methanol (Müller et al.,
455 2014a) and acetaldehyde, the instrument backgrounds of which have interferences from
456 O_2H^+ and CO_2H^+ at the same nominal masses, respectively; (3) increase the number of
457 species that can be measured. In general, 10-20 compounds are usually measurable by the
458 PTR-QMS explicitly without significant interference depending on the origin of the air
459 masses (de Gouw and Warneke, 2007). During the SONGNEX campaign, a total of 1055
460 peaks between m/z 12 and m/z 181 were identified and assigned to signals in the mass
461 spectra of H_3O^+ ToF-CIMS. Not all of these mass signals can be used for VOC
462 quantification. Many masses (1) are associated with large errors from high-resolution
463 peak fittings as the result of a much larger peak nearby and/or poor separation from
464 another peak; or (2) have no significant enhancement over instrument background. Over
465 260 masses had periods (at least 1 min) with signal to background ratios larger than three
466 during the flight over the Permian Basin on April 23, 2015. Although not all of these
467 masses are attributable to specific compounds, chemical formulas of the masses give
468 more detailed information on the chemical composition than just the nominal m/z .
469 Detailed interpretation of the mass spectra in various air masses during the SONGNEX
470 campaign will be presented in a separate publication (Koss et al., in prep).



471 **3.5 Background correction, in-flight calibration and detection limits**

472 The humidity dependence curves illustrated in Figure 5 were used to correct the
473 normalized signals of various VOC species from measurements made during SONGNEX.
474 After that, background signals were averaged for each measurement cycle and
475 interpolated to the periods between background measurements. Two features in ambient
476 measurements were considered in the interpolation of background signals. First, we
477 observed continuously decrease of background signals for some ions during the flights
478 (Figure S8B). The decrease of background signals cannot be avoided during aircraft
479 measurements, as the instrument only had 2-3 hours start-up time in the morning of each
480 flight day and the instrument backgrounds usually become lower as the instrument runs.
481 Second, background signals for many ions are dependent on humidity in the air, even
482 after the correction for the humidity dependence of their sensitivities (Figure S8D). This
483 is also of particular importance for aircraft measurements, as a rapid change of humidity
484 is encountered during aircraft ascents and descents (de Gouw and Warneke, 2007).
485 Exponential decay function was used to describe the continuous decline of background
486 signals, whereas humidity dependence of background signals was described using linear
487 relationships with one of the humidity indicators ($R_{37/19}$, $H_3O^+(H_2O)_2/H_3O^+$, O_2^+ signals
488 and CO_2^+ signals). Both effects were important for some ions and they were taken into
489 account by consecutive implementation of the procedures (Figure S8C). For other ions, a
490 simple linear interpolation was adopted (Figure S8E). The variations of instrument
491 backgrounds with instrument running time and humidity may also be critical for other
492 PTR-MS and CIMS instruments, especially for aircraft deployments and in some
493 circumstances when meteorological conditions change quickly during ground
494 measurements. The procedures shown here should be easy to incorporate into these
495 measurements if necessary.

496 The measurements of each in-flight calibration are averaged in the same way as the
497 background measurements. The results for benzene, isoprene, acetaldehyde and acetone
498 during the SONGNEX campaign are shown in Figure 9. Each of the four species
499 represents one of the four groups of compounds described in section 3.3. The calibration
500 results for the four species show tight linear correlations between corrected normalized
501 signals and calculated mixing ratios from the gas standard ($R > 0.995$), indicating stable



502 instrument performance from flight to flight during the SONGNEX campaign. No clear
503 dependence of sensitivities on $R_{37/19}$ is observed, demonstrating that the effects of
504 humidity on the sensitivities are properly accounted for by the procedures described
505 above. The measured sensitivities for various VOC species are shown in Table 1, which
506 lists sensitivities under dry conditions. The sensitivities for most VOC species of interest
507 are better than 400 cps/ppb, with several species higher than 800 cps/ppb. The lower
508 sensitivity for methanol is consistent with previous studies (Warneke et al., 2015),
509 whereas lower sensitivities for isoprene and α -pinene are due to the fragmentation of their
510 product ions.

511 The counting statistics of the ions follow a Poisson distribution, i.e. the $1-\sigma$ error of
512 counting N ions is \sqrt{N} . Recent studies showed that the high-resolution peak fitting to the
513 TOF mass spectra can add significant additional noise to the fitted peak intensities of the
514 masses (Cubison and Jimenez, 2015; Corbin et al., 2015). Figure 10 shows standard
515 deviations of the background signals versus the background signals themselves from the
516 individual zeroing periods for the masses listed in Table 1 during one SONGNEX flight
517 (April 27, 2015). Note that the signals shown in Figure 10 were not corrected for the ToF
518 duty cycle, so they represent the actual ion signals detected by the MCP. Real variations
519 of background signals due to instrument drift may contribute some to the standard
520 deviations of background signals. Thus, data points in Figure 10 are best viewed as the
521 upper limits of the errors from counting statistics. Most of the data are observed in the
522 region between \sqrt{N} and $2\times\sqrt{N}$, suggesting that high-resolution peak fitting can increase
523 the errors in the ion signals by as much as a factor of 2 for the VOC masses in Table 1.

524 The signal to noise ratio (S/N) of the species X from a CIMS instrument can be
525 expressed by (Bertram et al., 2011):

$$\frac{S}{N} = \frac{C_f[X]t}{\alpha \times \sqrt{C_f[X]t + 2Bt}} \quad (1)$$

526 Here, $[X]$ is the mixing ratios of the species (ppb), C_f is the sensitivity of the species
527 (cps/ppb) and B is the background count (cps), both of which are values without
528 corrections for the ToF duty cycle. t is the sampling time (s). α is the scaling factor of the
529 errors in ion signals relative to Poissonian statistics, and used to account for the
530 additional errors from HR peak fitting. α for each species is determined from Figure 10



531 (see Table 1 for values). We define the detection limit as the mixing ratio with a S/N ratio
532 of 3. The calculated 1-s detection limits for various species are listed in Table 1. The 1-s
533 detection limits are better than 100 ppt for most species. Higher detection limits for
534 acetaldehyde, acetic acid, isoprene, methanol and ethanol are the result of higher
535 background counts for the former two species and lower sensitivities for the latter three
536 species, respectively.

537 Figure 11 compares sensitivities and detection limits of methanol, acetone and
538 benzene between the NOAA PTR-QMS from several previous campaigns (Warneke et
539 al., 2011) and the new H_3O^+ ToF-CIMS during SONGNEX. As discussed in Warneke et
540 al. (2011), the performance of the NOAA PTR-QMS improved gradually as a result of
541 many instrumental developments, which led to higher sensitivities and lower detection
542 limits. As shown in Figure 11, the sensitivity of acetone in the new H_3O^+ ToF-CIMS is
543 similar to that of PTR-QMS in the recent campaigns (CalNex and UBWOS 2013),
544 whereas sensitivities of methanol and acetone are somewhat lower than the PTR-QMS.
545 The performance of a commercial PTR-TOF (PTR-TOF 8000, Ionicon Analytik) during
546 the UBWOS 2013 (Warneke et al., 2015) is also included in Figure 11. We observed
547 much higher sensitivities for our H_3O^+ ToF-CIMS than the PTR-TOF used in UBWOS
548 2013. The large difference in the sensitivities between the PTR-TOF 8000 and H_3O^+
549 ToF-CIMS is mainly attributed to the difference between ion lenses used in the PTR-TOF
550 to transfer ions from the drift tube to the mass analyzer and the quadrupole ion guides
551 used in H_3O^+ TOF-CIMS. A large enhancement in sensitivities of a PTR-TOF by using
552 quadrupole ion guides was recently also demonstrated in a PTR-QiTOF instrument from
553 Ionicon Analytik (Sulzer et al., 2014), which achieved a benzene sensitivity up to 2900
554 cps/ppb.

555 As discussed in the introduction, the operation of a QMS and a TOF-MS is
556 different. The TOF instruments acquire full mass spectra every second (or faster),
557 whereas PTR-QMS is usually operated in selected-ion mode with 1-sec measurement for
558 each of the selected masses every ~15 seconds. So if the total counts measured in each
559 15-second period are calculated (the middle panel in Figure 11), the overall signal to
560 noise of the ToF instruments will compares very favorably to the PTR-QMS.



561 Consistent with the sensitivities, 1-s detection limits of H_3O^+ ToF-CIMS are close
562 to those in PTR-QMS and better than the PTR-TOF used during the UBWOS 2013. As
563 with sensitivities, the detection limits of TOF instruments compare more favorably to
564 QMS if measured data is averaged to the total cycle length of a PTR-QMS to scan the
565 selected masses for their dwell time. We note that the detection limits for methanol,
566 acetone and benzene in the PTR-QiTOF were not reported in Sulzer et al. (2014), and
567 hence a comparison with H_3O^+ ToF-CIMS is not possible at this point.

568 **3.6 Inter-comparisons with GC-MS**

569 In addition to H_3O^+ ToF-CIMS for VOC measurements, whole air samples were
570 collected into canisters in-flight and were analyzed post-flight by GC-MS (iWAS) during
571 the SONGNEX campaign. A total of 72 samples were obtained for most flights. The fill
572 time for each canister was 3-15 s dependent on aircraft altitude. A brief description of the
573 iWAS system was provided in recent publications (de Gouw et al., 2015;Warneke et al.,
574 2016). A few compounds were measured by both H_3O^+ ToF-CIMS and the iWAS
575 system, including benzene, toluene and C8 aromatics. For C8 aromatics, the H_3O^+ ToF-
576 CIMS measured the total mixing ratios of the isomers, whereas the iWAS system was
577 able to measure mixing ratios for the individual isomers (o-xylene, m/p-xylene and
578 ethylbenzene).

579 Figure 12 compares measured results for benzene, toluene and C8 aromatics
580 between H_3O^+ ToF-CIMS and iWAS. The agreement between the two instruments is
581 good for all three aromatic species, with slopes in the range of 0.89-1.1 and correlation
582 coefficients (R) larger than 0.9. We note that there are a few data points with larger
583 disagreement, which are the result of imperfect time alignment of the two measurements
584 during large transient concentration spikes downwind from point sources.

585 **3.7 Applications to ambient measurements**

586 As mentioned above, the H_3O^+ ToF-CIMS was deployed onboard the NOAA WP-
587 3D during the SONGNEX campaign. Figure 13 shows measurement results from a
588 portion of a flight on April 13, 2015 over the Denver-Julesburg Basin, which is an active
589 oil and gas extraction region. Mixing ratios of benzene and toluene were elevated over
590 and downwind from the oil/gas field. Several concentration spikes of the two aromatics
591 were observed and attributed to large point sources related to oil and gas activities. Urban



592 emissions may also contribute to the mixing ratios of benzene and toluene, when the
593 NOAA WP-3D flew over the cities (Loveland, Fort Collins and Greeley) in the area. As
594 shown in Figure 13, variations of benzene and toluene mixing ratios below 100 ppt still
595 tracked each other well, illustrating the low detection limits of the two compounds in the
596 H_3O^+ ToF-CIMS. As the emissions of aromatics and other VOCs from motor vehicles are
597 declining in both U.S. and European cities (Warneke et al., 2012; Derwent et al., 2014),
598 the need for techniques (e.g. our H_3O^+ ToF-CIMS) that can measure these compounds
599 rapidly and accurately with low detection limits is increasing.

600 Concentration peaks of acetonitrile were detected several times in this period, when
601 the aircraft sampled plumes from agricultural burns in the area (de Gouw et al., 2003b).
602 The concentration peaks of acetonitrile were pretty narrow, especially the one at 21:34
603 UTC (see the inserted graph in Figure 13B). These small fire plumes can be easily missed
604 if a PTR-QMS had been applied for the detection of acetonitrile, as PTR-QMS has a duty
605 cycle of only a few percent for each individual mass, as shown in the inserted plot in
606 13B. Along with acetonitrile, furan and furfural were also observed in these biomass
607 burning plumes. Furan is a known emission from biomass burning and has been detected
608 in the atmosphere mainly using PTR-MS (de Gouw and Warneke, 2007; Karl et al.,
609 2007). In PTR-QMS, furan cannot be distinguished from isoprene. Its aldehyde
610 derivatives, furfurals, were only measured in a recent laboratory experiment of biomass
611 burning emissions (Stockwell et al., 2015) and in a forest fire plume in the US (Müller et
612 al., 2015) both by PTR-TOF.

613 The third class of compounds shown in Figure 13 is acetaldehyde and acetone.
614 Time series of the two OVOCs were similar to those of benzene and toluene with
615 elevated mixing ratios over and downwind from the oil and gas wells, but they lacked the
616 concentration spikes observed for the aromatics, consistent with secondary formation as
617 their main sources. Acetaldehyde and acetone were also enhanced in the biomass burning
618 plumes.

619 As a final example, acetic acid and ethanol were observed in both oil/gas and
620 biomass burning plumes. Additional peaks of acetic acid and ethanol were detected
621 shortly before and after 21:00 UTC. These enhanced mixing ratios were believed to be
622 the result of emissions from agricultural facilities in this area. Weld County, the main



623 flight area shown in Figure 13, is home to over half million beef and dairy cattle in over
624 100 feedlots facilities. Emissions of acetic acid and ethanol from dairy operations have
625 been reported recently in the Central Valley of California (Gentner et al., 2014).

626 Figure 13 demonstrates that four different groups of VOC species exhibited with
627 distinctly different time series and spatial distributions, as the result of their different
628 sources. The examples in Figure 13 illustrate that the high time resolution dataset from
629 H_3O^+ ToF-CIMS provides important information on characterizations of various air
630 masses in the atmosphere. Along with unique tracers (e.g. acetonitrile for biomass
631 burning), rich information on the chemical signatures will help to identify and separate
632 the contributions of different sources to air mass components in the atmosphere.
633 Measurements of secondary products, in conjunction with primary emissions will be
634 valuable to constrain chemical evolution of gas-phase organic carbon in the atmosphere.

635 **4. Conclusions**

636 In this study, an aircraft-deployable H_3O^+ ToF-CIMS instrument was developed
637 based on a commercial Aerodyne ToF-CIMS. We characterize the humidity dependence
638 of reagent ions and instrument sensitivities for various VOC species. The new H_3O^+ ToF-
639 CIMS has sensitivities in the range of 100-1000 cps/ppbv for many VOCs of interest and
640 the 1-s detection limits are in the range of 20-500 ppt depending on product ion masses
641 and their instrument backgrounds. The instrument was deployed onboard the NOAA WP-
642 3D research aircraft as part of the SONGNEX campaign in March-April of 2015. The
643 measured mixing ratios for several aromatics from the H_3O^+ ToF-CIMS agreed very well
644 with independent GC measurements from whole air samples. Some initial results from
645 the instrument demonstrate that the H_3O^+ ToF-CIMS dataset will be extremely valuable
646 for the characterization of VOC emissions and photochemistry in the atmosphere.

647 We showed that the low transmission of H_3O^+ ions and secondary ion chemistry
648 inside the SSQ complicates signal normalization and interpretation of the humidity
649 dependence of VOC sensitivities. Further development by removal of the SSQ may
650 ameliorate this issue and is being considered. We also anticipate a further increase in
651 sensitivities from these modifications, as shown in a recent publication (Sulzer et al.,
652 2014). We show that many instrument settings, including SSQ pressure, RF voltages in
653 quadrupoles and the axial voltage gradients along the quadrupoles, affect both reagent



654 ions and VOC signals significantly in our instrument. Some of these effects were
655 explored or seen in other Aerodyne TOF-CIMS instruments as well (Zheng et al.,
656 2015;Lopez-Hilfiker et al., 2015). The next generation of PTR-MS (PTR-QiTOF) that
657 uses quadrupole ion guides for ion transmission may also be affected by some of
658 secondary ion chemistry in the quadrupole. Instruments using different ion chemistry are
659 expected to behave differently with these settings. Therefore, characterization of each
660 individual instrument for these variations and calibrations using authentic standards are
661 essential for generating high quality measurement data with these sophisticated CIMS
662 instruments.

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665

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673



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675 **References:**

- 676 Atkinson, R.: Atmospheric chemistry of VOCs and NO_x, *Atmospheric Environment*, 34,
677 2063-2101, 2000.
- 678 Baasandorj, M., Millet, D. B., Hu, L., Mitroo, D., and Williams, B. J.: Measuring acetic
679 and formic acid by proton-transfer-reaction mass spectrometry: sensitivity, humidity
680 dependence, and quantifying interferences, *Atmospheric Measurement Techniques*,
681 8, 1303-1321, 10.5194/amt-8-1303-2015, 2015.
- 682 Bertram, T. H., Kimmel, J. R., Crisp, T. A., Ryder, O. S., Yatavelli, R. L. N., Thornton, J.
683 A., Cubison, M. J., Gonin, M., and Worsnop, D. R.: A field-deployable, chemical
684 ionization time-of-flight mass spectrometer, *Atmos. Meas. Tech.*, 4, 1471-1479,
685 10.5194/amt-4-1471-2011, 2011.
- 686 Blake, R. S., Whyte, C., Hughes, C. O., Ellis, A. M., and Monks, P. S.: Demonstration of
687 proton-transfer reaction time-of-flight mass spectrometry for real-time analysis of
688 trace volatile organic compounds, *Analytical Chemistry*, 76, 3841-3845,
689 10.1021/ac0498260, 2004.
- 690 Blake, R. S., Monks, P. S., and Ellis, A. M.: Proton-Transfer Reaction Mass
691 Spectrometry, *Chemical Reviews*, 109, 861-896, Doi 10.1021/Cr800364q, 2009.
- 692 Chernushevich, I. V., Loboda, A. V., and Thomson, B. A.: An introduction to
693 quadrupole-time-of-flight mass spectrometry, *Journal of Mass Spectrometry*, 36,
694 849-865, 10.1002/jms.207, 2001.
- 695 Corbin, J., Othman, A., D. Allan, J., R. Worsnop, D., D. Haskins, J., Sierau, B.,
696 Lohmann, U., and A. Mensah, A.: Peak-fitting and integration imprecision in the
697 Aerodyne aerosol mass spectrometer: effects of mass accuracy on location-
698 constrained fits, *Atmospheric Measurement Techniques*, 8, 4615-4636, 10.5194/amt-
699 8-4615-2015, 2015.
- 700 Cubison, M. J., and Jimenez, J. L.: Statistical precision of the intensities retrieved from
701 constrained fitting of overlapping peaks in high-resolution mass spectra,
702 *Atmospheric Measurement Techniques*, 8, 2333-2345, 10.5194/amt-8-2333-2015,
703 2015.
- 704 de Gouw, J., and Warneke, C.: Measurements of volatile organic compounds in the
705 earth's atmosphere using proton-transfer-reaction mass spectrometry, *Mass
706 Spectrometry Reviews*, 26, 223-257, 2007.
- 707 de Gouw, J. A., Goldan, P. D., Warneke, C., Kuster, W. C., Roberts, J. M., Marchewka,
708 M., Bertman, S. B., Pszenny, A. A. P., and Keene, W. C.: Validation of proton
709 transfer reaction-mass spectrometry (PTR-MS) measurements of gas-phase organic
710 compounds in the atmosphere during the New England Air Quality Study (NEAQS)
711 in 2002, *Journal of Geophysical Research-Atmospheres*, 108,
712 doi:10.1029/2003JD003863, 10.1029/2003jd003863, 2003a.
- 713 de Gouw, J. A., Warneke, C., Parrish, D. D., Holloway, J. S., Trainer, M., and
714 Fehsenfeld, F. C.: Emission sources and ocean uptake of acetonitrile (CH₃CN) in the
715 atmosphere, *Journal of Geophysical Research-Atmospheres*, 108,
716 doi:10.1029/2002JD002897, 10.1029/2002jd002897, 2003b.
- 717 de Gouw, J. A., McKeen, S. A., Aikin, K. C., Brock, C. A., Brown, S. S., Gilman, J. B.,
718 Graus, M., Hanisco, T., Holloway, J. S., Kaiser, J., Keutsch, F. N., Lerner, B. M.,
719 Liao, J., Markovic, M. Z., Middlebrook, A. M., Min, K. E., Neuman, J. A., Nowak, J.



- 720 B., Peischl, J., Pollack, I. B., Roberts, J. M., Ryerson, T. B., Trainer, M., Veres, P.
721 R., Warneke, C., Welti, A., and Wolfe, G. M.: Airborne measurements of the
722 atmospheric emissions from a fuel ethanol refinery, *Journal of Geophysical*
723 *Research: Atmospheres*, 120, 4385-4397, 10.1002/2015jd023138, 2015.
- 724 Derwent, R. G., Dorn, J. I. R., Dollard, G. J., Dumitrean, P., Mitchell, R. F., Murrells,
725 T. P., Telling, S. P., and Field, R. A.: Twenty years of continuous high time
726 resolution volatile organic compound monitoring in the United Kingdom from 1993
727 to 2012, *Atmospheric Environment*, 99, 239-247, 10.1016/j.atmosenv.2014.10.001,
728 2014.
- 729 Ennis, C. J., Reynolds, J. C., Keely, B. J., and Carpenter, L. J.: A hollow cathode proton
730 transfer reaction time of flight mass spectrometer, *International Journal of Mass*
731 *Spectrometry*, 247, 72-80, 10.1016/j.ijms.2005.09.008, 2005.
- 732 Gentner, D. R., Ford, T. B., Guha, A., Boulanger, K., Brioude, J., Angevine, W. M., de
733 Gouw, J. A., Warneke, C., Gilman, J. B., Ryerson, T. B., Peischl, J., Meinardi, S.,
734 Blake, D. R., Atlas, E., Lonneman, W. A., Kleindienst, T. E., Beaver, M. R., Clair, J.
735 M. S., Wennberg, P. O., VandenBoer, T. C., Markovic, M. Z., Murphy, J. G., Harley,
736 R. A., and Goldstein, A. H.: Emissions of organic carbon and methane from
737 petroleum and dairy operations in California's San Joaquin Valley, *Atmos. Chem.*
738 *Phys.*, 14, 4955-4978, 10.5194/acp-14-4955-2014, 2014.
- 739 Graus, M., Müller, M., and Hansel, A.: High Resolution PTR-TOF: Quantification and
740 Formula Confirmation of VOC in Real Time, *Journal of the American Society for*
741 *Mass Spectrometry*, 21, 1037-1044, DOI 10.1016/j.jasms.2010.02.006, 2010.
- 742 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
743 Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F.,
744 Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L.,
745 Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A.,
746 Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The
747 formation, properties and impact of secondary organic aerosol: current and emerging
748 issues, *Atmospheric Chemistry and Physics*, 9, 5155-5236, 2009.
- 749 Holzinger, R., Goldstein, A. H., Hayes, P. L., Jimenez, J. L., and Timkovsky, J.:
750 Chemical evolution of organic aerosol in Los Angeles during the CalNex 2010 study,
751 *Atmospheric Chemistry and Physics*, 13, 10125-10141, 10.5194/acp-13-10125-2013,
752 2013.
- 753 Hunter, E. P. L., and Lias, S. G.: Evaluated Gas Phase Basicities and Proton Affinities of
754 Molecules: An Update, *Journal of Physical and Chemical Reference Data*, 27, 413-
755 656, doi:<http://dx.doi.org/10.1063/1.556018>, 1998.
- 756 Isaksen, I. S. A., Granier, C., Myhre, G., Berntsen, T. K., Dalsøren, S. B., Gauss, M.,
757 Klimont, Z., Benestad, R., Bousquet, P., Collins, W., Cox, T., Eyring, V., Fowler, D.,
758 Fuzzi, S., Jöckel, P., Laj, P., Lohmann, U., Maione, M., Monks, P., Prevot, A. S. H.,
759 Raes, F., Richter, A., Rognerud, B., Schulz, M., Shindell, D., Stevenson, D. S.,
760 Storelvmo, T., Wang, W. C., van Weele, M., Wild, M., and Wuebbles, D.:
761 Atmospheric composition change: Climate-Chemistry interactions, *Atmospheric*
762 *Environment*, 43, 5138-5192, 10.1016/j.atmosenv.2009.08.003, 2009.
- 763 Jordan, A., Haidacher, S., Hanel, G., Hartungen, E., Mark, L., Seehauser, H.,
764 Schottkowsky, R., Sulzer, P., and Mark, T. D.: A high resolution and high sensitivity
765 proton-transfer-reaction time-of-flight mass spectrometer (PTR-TOF-MS),



- 766 International Journal of Mass Spectrometry, 286, 122-128, DOI
767 10.1016/j.ijms.2009.07.005, 2009.
- 768 Karl, T. G., Christian, T. J., Yokelson, R. J., Artaxo, P., Hao, W. M., and Guenther, A.:
769 The Tropical Forest and Fire Emissions Experiment: method evaluation of volatile
770 organic compound emissions measured by PTR-MS, FTIR, and GC from tropical
771 biomass burning, *Atmospheric Chemistry and Physics*, 7, 5883-5897, 2007.
- 772 Kaser, L., Karl, T., Schnitzhofer, R., Graus, M., Herdinger-Blatt, I. S., DiGangi, J. P.,
773 Sive, B., Turnipseed, A., Hornbrook, R. S., Zheng, W., Flocke, F. M., Guenther, A.,
774 Keutsch, F. N., Apel, E., and Hansel, A.: Comparison of different real time VOC
775 measurement techniques in a ponderosa pine forest, *Atmos. Chem. Phys.*, 13, 2893-
776 2906, 10.5194/acp-13-2893-2013, 2013.
- 777 Lee, B. H., Lopez-Hilfiker, F., Mohr, C., Kurtén, T. C., Worsnop, D., and Thornton, J.
778 A.: An Iodide-Adduct High-Resolution Time-of-Flight Chemical-Ionization Mass
779 Spectrometer: Application to Atmospheric Inorganic and Organic Compounds,
780 *Environmental Science & Technology*, 10.1021/es500362a, 2014.
- 781 Lerner, B.: A new analysis system for whole air sampling: description and early results,
782 in prep.
- 783 Lindinger, W., Hansel, A., and Jordan, A.: On-line monitoring of volatile organic
784 compounds at pptv levels by means of proton-transfer-reaction mass spectrometry
785 (PTR-MS) - Medical applications, food control and environmental research,
786 *International Journal of Mass Spectrometry*, 173, 191-241, 1998.
- 787 Lopez-Hilfiker, F. D., Iyer, S., Mohr, C., Lee, B. H., D'Ambro, E. L., Kurtén, T., and
788 Thornton, J. A.: Constraining the sensitivity of iodide adduct chemical ionization
789 mass spectrometry to multifunctional organic molecules using the collision limit and
790 thermodynamic stability of iodide ion adducts, *Atmos. Meas. Tech. Discuss.*, 8,
791 10875-10896, 10.5194/amtd-8-10875-2015, 2015.
- 792 Midey, A. J., Williams, S., Arnold, S. T., and Viggiano, A. A.: Reactions of
793 $\text{H}_3\text{O}^+(\text{H}_2\text{O})_0,1$ with Alkylbenzenes from 298 to 1200 K \ddagger , *The Journal of Physical*
794 *Chemistry A*, 106, 11726-11738, 10.1021/jp014141e, 2002.
- 795 Mielke, L. H., Erickson, D. E., McLuckey, S. A., Muller, M., Wisthaler, A., Hansel, A.,
796 and Shepson, P. B.: Development of a Proton-Transfer Reaction-Linear Ion Trap
797 Mass Spectrometer for Quantitative Determination of Volatile Organic Compounds,
798 *Analytical Chemistry*, 80, 8171-8177, 10.1021/ac801328d, 2008.
- 799 Müller, M., Graus, M., Ruuskanen, T. M., Schnitzhofer, R., Bamberger, I., Kaser, L.,
800 Titzmann, T., Hörtnagl, L., Wohlfahrt, G., Karl, T., and Hansel, A.: First eddy
801 covariance flux measurements by PTR-TOF, *Atmos. Meas. Tech.*, 3, 387-395,
802 10.5194/amt-3-387-2010, 2010.
- 803 Müller, M., Mikoviny, T., Feil, S., Haidacher, S., Hanel, G., Hartungen, E., Jordan, A.,
804 Märk, L., Mutschlechner, P., Schottkowsky, R., Sulzer, P., Crawford, J. H., and
805 Wisthaler, A.: A compact PTR-ToF-MS instrument for airborne measurements of
806 volatile organic compounds at high spatiotemporal resolution, *Atmos. Meas. Tech.*,
807 7, 3763-3772, 10.5194/amt-7-3763-2014, 2014a.
- 808 Müller, M., Mikoviny, T., and Wisthaler, A.: Detector aging induced mass discrimination
809 and non-linearity effects in PTR-ToF-MS, *International Journal of Mass*
810 *Spectrometry*, 365-366, 93-97, 10.1016/j.ijms.2013.12.008, 2014b.



- 811 Müller, M., Anderson, B., Beyersdorf, A., Crawford, J. H., Diskin, G., Eichler, P., Fried,
812 A., Keutsch, F. N., Mikoviny, T., Thornhill, K. L., Walega, J. G., Weinheimer, A. J.,
813 Yang, M., Yokelson, R., and Wisthaler, A.: In situ measurements and modeling of
814 reactive trace gases in a small biomass burning plume, *Atmospheric Chemistry and*
815 *Physics Discussions*, 15, 31501-31536, 10.5194/acpd-15-31501-2015, 2015.
- 816 Park, J.-H., Goldstein, A. H., Timkovsky, J., Fares, S., Weber, R., Karlik, J., and
817 Holzinger, R.: Active Atmosphere-Ecosystem Exchange of the Vast Majority of
818 Detected Volatile Organic Compounds, *Science*, 341, 643-647,
819 10.1126/science.1235053, 2013.
- 820 Raspoet, G., Nguyen, M. T., McGarraghy, M., and Hegarty, A. F.: Experimental and
821 Theoretical Evidence for a Concerted Catalysis by Water Clusters in the Hydrolysis
822 of Isocyanates, *The Journal of Organic Chemistry*, 63, 6867-6877,
823 10.1021/jo980639+, 1998.
- 824 Spanel, P., and Smith, D.: Reactions of Hydrated Hydronium Ions and Hydrated
825 Hydroxide Ions with Some Hydrocarbons and Oxygen-Bearing Organic Molecules,
826 *The Journal of Physical Chemistry*, 99, 15551-15556, 10.1021/j100042a033, 1995.
- 827 Spanel, P., and Smith, D.: SIFT studies of the reactions of H₃O⁺, NO⁺ and O-2(+) with a
828 series of alcohols, *International Journal of Mass Spectrometry*, 167, 375-388, 1997.
- 829 Stark, H., Yatawelli, R. L. N., Thompson, S. L., Kimmel, J. R., Cubison, M. J., Chhabra,
830 P. S., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R., and Jimenez, J. L.: Methods
831 to extract molecular and bulk chemical information from series of complex mass
832 spectra with limited mass resolution, *International Journal of Mass Spectrometry*,
833 10.1016/j.ijms.2015.08.011, 2015.
- 834 Steeghs, M. M. L., Sikkens, C., Crespo, E., Cristescu, S. M., and Harren, F. J. M.:
835 Development of a proton-transfer reaction ion trap mass spectrometer: Online
836 detection and analysis of volatile organic compounds, *International Journal of Mass*
837 *Spectrometry*, 262, 16-24, 10.1016/j.ijms.2006.09.031, 2007.
- 838 Stockwell, C. E., Veres, P. R., Williams, J., and Yokelson, R. J.: Characterization of
839 biomass burning emissions from cooking fires, peat, crop residue, and other fuels
840 with high-resolution proton-transfer-reaction time-of-flight mass spectrometry,
841 *Atmos. Chem. Phys.*, 15, 845-865, 10.5194/acp-15-845-2015, 2015.
- 842 Sulzer, P., Hartungen, E., Hanel, G., Feil, S., Winkler, K., Mutschlechner, P., Haidacher,
843 S., Schotchkowsky, R., Gunsch, D., Seehauser, H., Striednig, M., Jürschik, S., Breiev,
844 K., Lanza, M., Herbig, J., Märk, L., Märk, T. D., and Jordan, A.: A Proton Transfer
845 Reaction-Quadrupole interface Time-Of-Flight Mass Spectrometer (PTR-QiTOF):
846 High speed due to extreme sensitivity, *International Journal of Mass Spectrometry*,
847 368, 1-5, 10.1016/j.ijms.2014.05.004, 2014.
- 848 Tanimoto, H., Aoki, N., Inomata, S., Hirokawa, J., and Sadanaga, Y.: Development of a
849 PTR-TOFMS instrument for real-time measurements of volatile organic compounds
850 in air, *International Journal of Mass Spectrometry*, 263, 1-11,
851 <http://dx.doi.org/10.1016/j.ijms.2007.01.009>, 2007.
- 852 Warneke, C., Kuczynski, J., Hansel, A., Jordan, A., Vogel, W., and Lindinger, W.: Proton
853 transfer reaction mass spectrometry (PTR-MS): propanol in human breath, *Int J Mass*
854 *Spectrom*, 154, 61-70, [http://dx.doi.org/10.1016/0168-1176\(96\)04369-8](http://dx.doi.org/10.1016/0168-1176(96)04369-8), 1996.
- 855 Warneke, C., van der Veen, C., Luxembourg, S., de Gouw, J. A., and Kok, A.:
856 Measurements of benzene and toluene in ambient air using proton-transfer-reaction



- 857 mass spectrometry: calibration, humidity dependence, and field intercomparison,
858 *International Journal of Mass Spectrometry*, 207, 167-182, 2001.
- 859 Warneke, C., De Gouw, J. A., Kuster, W. C., Goldan, P. D., and Fall, R.: Validation of
860 atmospheric VOC measurements by proton-transfer-reaction mass spectrometry
861 using a gas-chromatographic pre-separation method, *Environmental Science &*
862 *Technology*, 37, 2494-2501, 10.1021/es026266i, 2003.
- 863 Warneke, C., Kato, S., De Gouw, J. A., Goldan, P. D., Kuster, W. C., Shao, M., Lovejoy,
864 E. R., Fall, R., and Fehsenfeld, F. C.: Online volatile organic compound
865 measurements using a newly developed proton-transfer ion-trap mass spectrometry
866 instrument during New England Air Quality Study - Intercontinental Transport and
867 Chemical Transformation 2004: Performance, intercomparison, and compound
868 identification, *Environmental Science & Technology*, 39, 5390-5397,
869 10.1021/es050602o, 2005.
- 870 Warneke, C., Veres, P., Holloway, J. S., Stutz, J., Tsai, C., Alvarez, S., Rappenglueck, B.,
871 Fehsenfeld, F. C., Graus, M., Gilman, J. B., and de Gouw, J. A.: Airborne
872 formaldehyde measurements using PTR-MS: calibration, humidity dependence,
873 inter-comparison and initial results, *Atmos. Meas. Tech.*, 4, 2345-2358, 10.5194/amt-
874 4-2345-2011, 2011.
- 875 Warneke, C., de Gouw, J. A., Holloway, J. S., Peischl, J., Ryerson, T. B., Atlas, E.,
876 Blake, D., Trainer, M., and Parrish, D. D.: Multiyear trends in volatile organic
877 compounds in Los Angeles, California: Five decades of decreasing emissions, *J.*
878 *Geophys. Res.*, 117, D00V17, 10.1029/2012jd017899, 2012.
- 879 Warneke, C., Veres, P. R., Murphy, S. M., Soltis, J., Field, R. A., Graus, M. G., Koss, A.,
880 Li, S. M., Li, R., Yuan, B., Roberts, J. M., and de Gouw, J. A.: PTR-QMS vs. PTR-
881 TOF comparison in a region with oil and natural gas extraction industry in the Uintah
882 Basin in 2013, *Atmos. Meas. Tech.*, 8, 411-420, 10.5194/amt-7-6565-2014, 2015.
- 883 Warneke, C., Trainer, M., de Gouw, J. A., Parrish, D. D., Fahey, D. W., Ravishankara, A.
884 R., Middlebrook, A. M., Brock, C. A., Roberts, J. M., Brown, S. S., Neuman, J. A.,
885 Lerner, B. M., Lack, D., Law, D., Huebler, G., Pollack, I., Sjostedt, S., Ryerson, T.
886 B., Gilman, J. B., Liao, J., Holloway, J., Peischl, J., Nowak, J. B., Aikin, K., Min, K.
887 E., Washenfelder, R. A., Graus, M. G., Richardson, M., Markovic, M. Z., Wagner,
888 N. L., Welti, A., Veres, P. R., Edwards, P., Schwarz, J. P., Gordon, T., Dube, W. P.,
889 McKeen, S., Brioude, J., Ahmadov, R., Bougiatioti, A., Lin, J., Nenes, A., Wolfe, G.
890 M., Hanisco, T. F., Lee, B. H., Lopez-Hilfiker, F. D., Thornton, J. A., Keutsch, F. N.,
891 Kaiser, J., Mao, J., and Hatch, C.: Instrumentation and Measurement Strategy for the
892 NOAA SENEX Aircraft Campaign as Part of the Southeast Atmosphere Study 2013,
893 *Atmos. Meas. Tech. Discuss.*, 2016, 1-39, 10.5194/amt-2015-388, 2016.
- 894 Zheng, J., Ma, Y., Chen, M. D., Zhang, Q., Wang, L., Khalizov, A. F., Yao, L., Wang, Z.,
895 Wang, X., and Chen, L. X.: Measurement of atmospheric amines and ammonia using
896 the high resolution time-of-flight chemical ionization mass spectrometry,
897 *Atmospheric Environment*, 102, 249-259, 10.1016/j.atmosenv.2014.12.002, 2015.
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901 **Tables**

902 **Table 1. Sensitivities and detection limits of H₃O⁺ ToF-CIMS for various VOCs**
 903 **species during the SONGNEX campaign**

VOC species	Ion Formula	Sensitivity ¹		Background, cps ²	α value ³	1-s detection limit, ppt
		ncps/ppb	cps/ppb ²			
Methanol*	CH ₄ OH ⁺	81	158	128	1.34	397
Acetonitrile*	C ₂ H ₃ NH ⁺	376	822	26	1.33	45
Acetaldehyde*	C ₂ H ₄ OH ⁺	289	654	499	1.36	195
Acetone*	C ₃ H ₆ OH ⁺	354	916	261	1.28	97
Acetic acid	C ₂ H ₄ O ₂ H ⁺	209	551	660	1.40	283
Furan	C ₄ H ₄ OH ⁺	168	470	11	1.26	58
Isoprene*	C ₅ H ₈ H ⁺	73	206	26	1.21	162
MVK	C ₄ H ₆ OH ⁺	160	454	46	1.24	85
MEK*	C ₄ H ₈ OH ⁺	308	886	48	1.21	45
Benzene*	C ₆ H ₆ H ⁺	165	493	22	1.48	96
Toluene*	C ₇ H ₈ H ⁺	178	579	6	1.39	47
o-xylene*	C ₈ H ₁₀ H ⁺	193	673	4	1.49	40
1,2,4-TMB*	C ₉ H ₁₂ H ⁺	185	686	3	1.69	45
α -pinene	C ₁₀ H ₁₆ H ⁺	67	263	2	1.24	67

904 ¹ Sensitivity under dry condition;

905 ²Calculations using H₃O⁺ signal at the typical level of 2.5×10⁶ cps during the SONGNEX
 906 campaign. The signals or sensitivities are not corrected for the ToF duty cycle (see text).

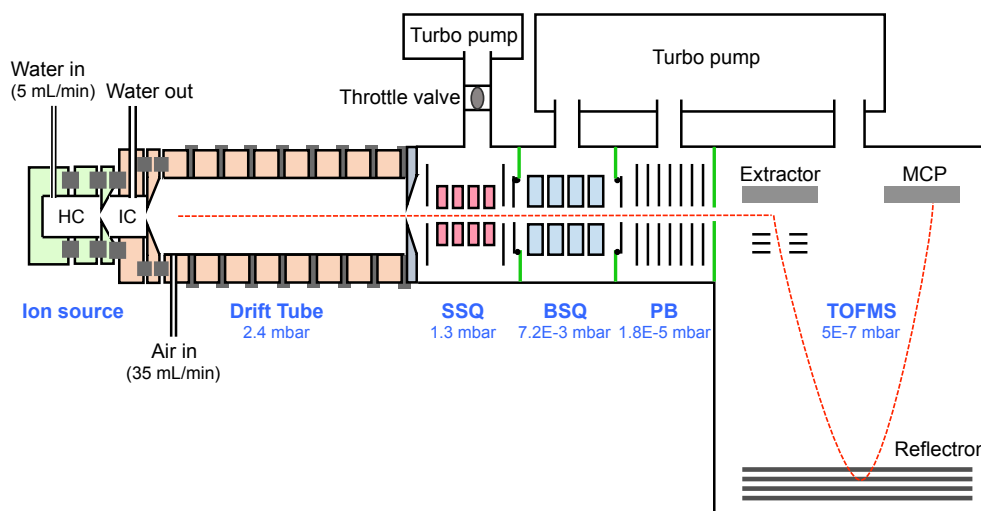
907 ³ α is the scaling factor of the errors in ion signals relative to Poissonian statistics
 908 (Equation 1).

909 *Stars indicate that the compounds were calibrated in-flight using the ten-component gas
 910 standard.

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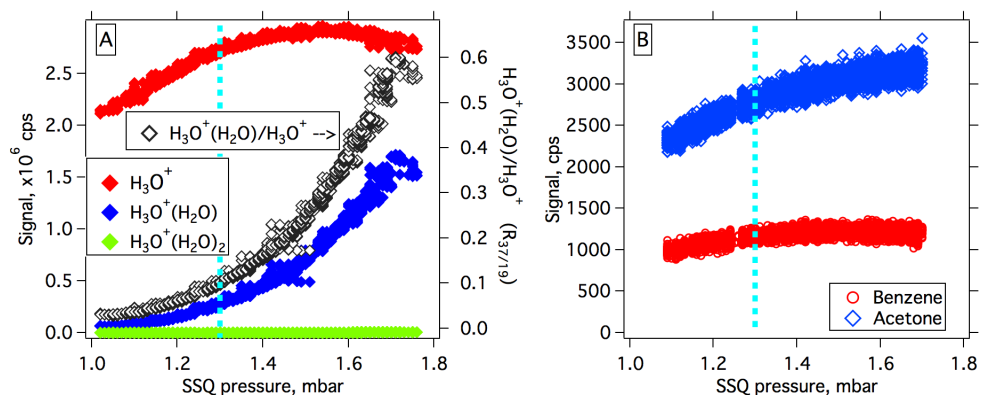
912 **Figures:**



913

914 Figure 1. Schematic drawing of the H_3O^+ ToF-CIMS. HC: hollow-cathode discharge; IC:
915 intermediate chamber; SSQ: small segmented quadrupole; BSQ: big segmented
916 quadrupole; PB: primary beam; MCP: microchannel plate detector.

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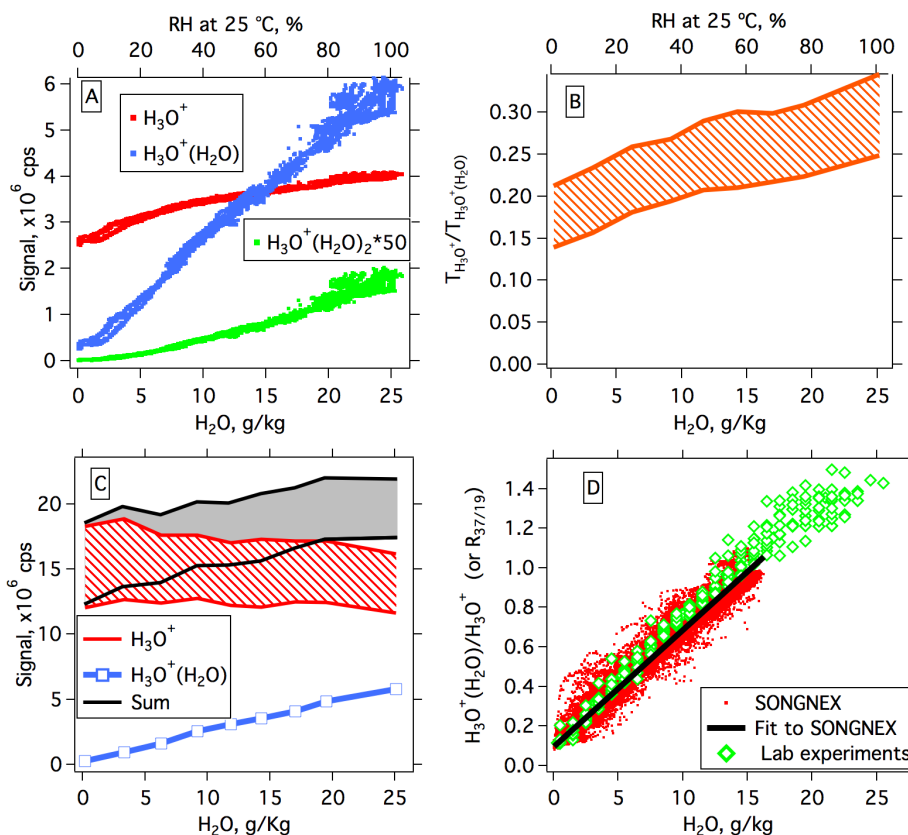


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919 Figure 2. (A) Signals of reagent ions and $\text{H}_3\text{O}^+(\text{H}_2\text{O})/\text{H}_3\text{O}^+$ ratio ($R_{37/19}$) as a function of
920 SSQ pressure. (B) Signals of protonated benzene and acetone at constant mixing ratios of
921 2.5 ppb as a function of SSQ pressure. The light blue vertical dashed lines indicate the
922 SSQ pressure (1.30 mbar) used during the SONGNEX campaign.

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926 Figure 3. (A) Reagent ion signals as a function of water vapor mixing ratios in the
 927 instrument. (B) Inferred transmission ratios between H_3O^+ and $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions
 928 ($T_{\text{H}_3\text{O}^+}/T_{\text{H}_3\text{O}^+(\text{H}_2\text{O})}$) as a function of water vapor mixing ratios. (C) The H_3O^+ signals and
 929 total signals of the two reagent ions that are corrected to the transmission factor of
 930 $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ ions as a function of water vapor mixing ratios. The measured $\text{H}_3\text{O}^+(\text{H}_2\text{O})$
 931 signals are also shown. The shaded areas in (B) and (C) indicate the possible ranges of
 932 the parameters, which are bounded by estimates from the depletion experiments of
 933 methanol and acetonitrile, respectively. (D) $\text{H}_3\text{O}^+(\text{H}_2\text{O})/\text{H}_3\text{O}^+$ ($R_{37/19}$) as a function of
 934 water vapor mixing ratios during SONGNEX and from laboratory experiments. The
 935 black line is the linear fit to all of the SONGNEX data.

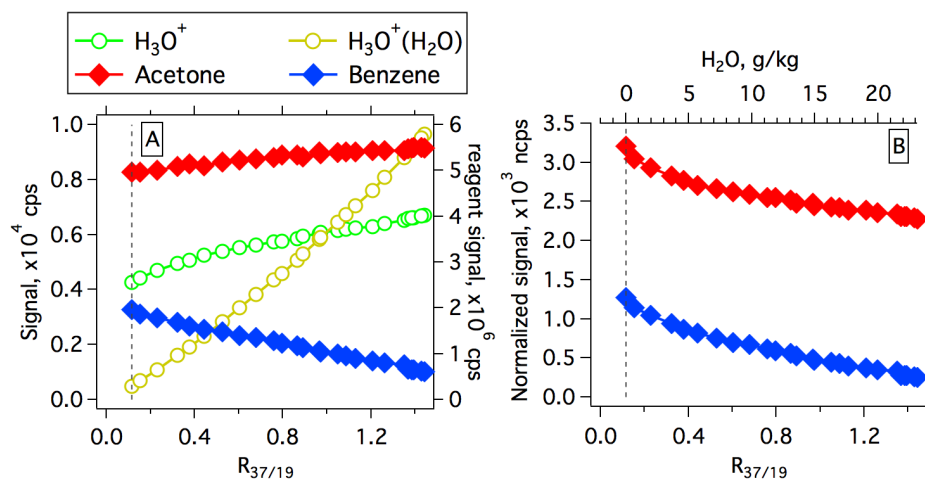
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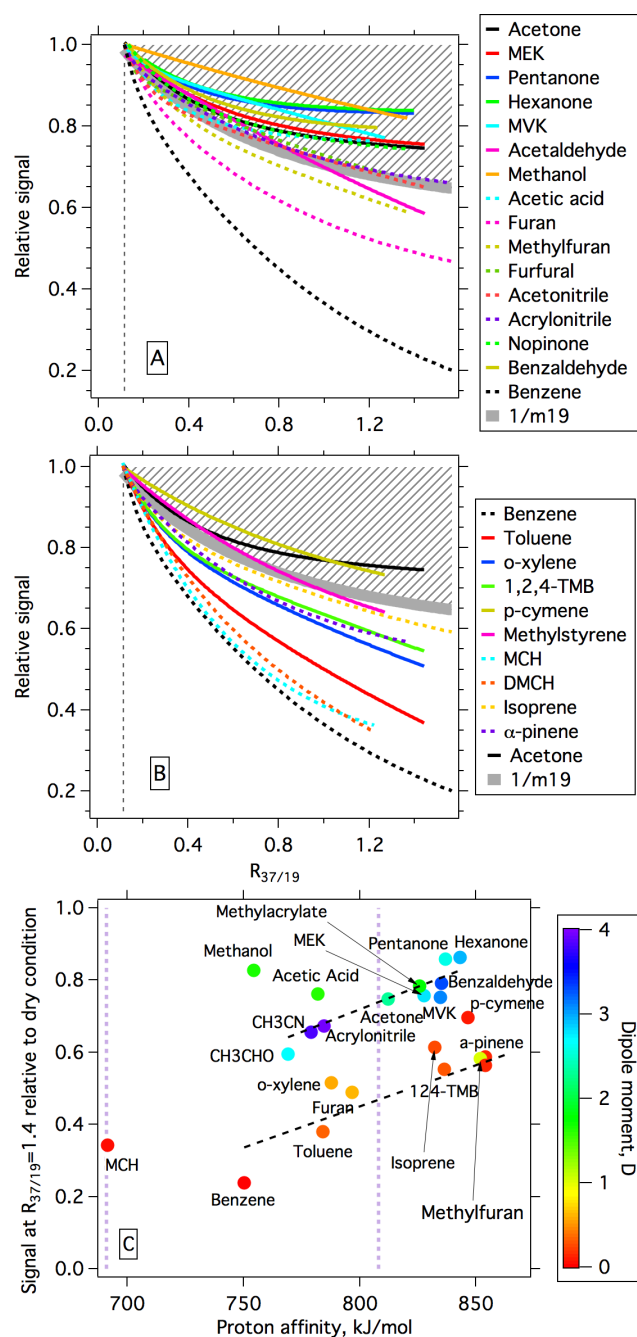
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941 Figure 4. Normalization and humidity dependence of acetone and benzene detection from
942 a laboratory experiment. (A) Background subtracted raw signals of protonated acetone
943 and benzene with mixing ratios of 8.0 ppb as the function of $R_{37/19}$. The signals of H_3O^+
944 and $H_3O^+(H_2O)$ are also shown for reference. (B) Signals of protonated acetone and
945 benzene normalized to the H_3O^+ signal as a function of $R_{37/19}$.

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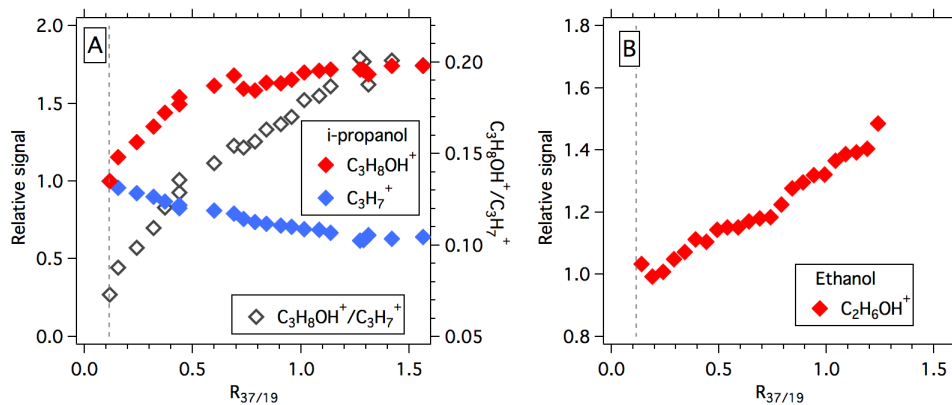


947

948 Figure 5. Derived humidity dependence curves for the normalized signals relative to dry
 949 condition as a function of $R_{37/19}$ ratios for oxygenates (A) and hydrocarbons (B),
 950 respectively. The humidity dependent curves of benzene and acetone are shown in both



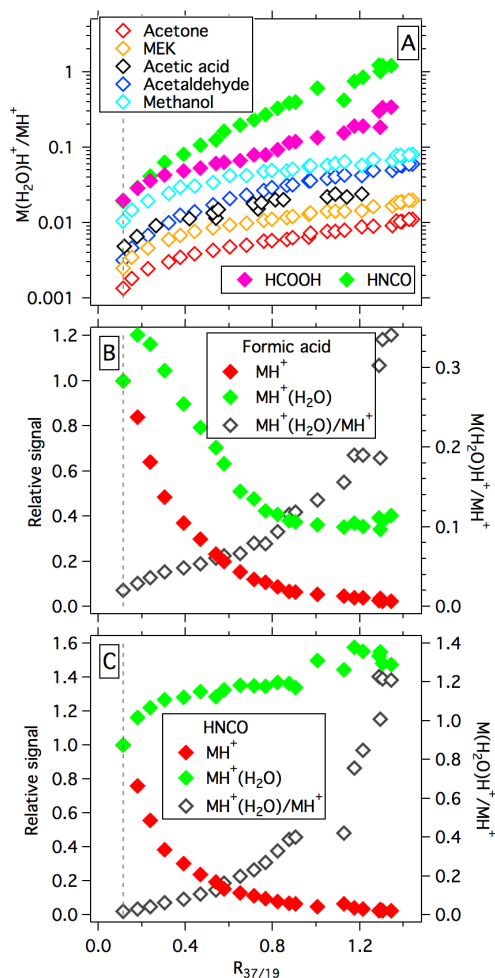
951 (A) and (B). The thick gray lines indicate the reciprocal of the H_3O^+ signal relative to dry
952 conditions ($\frac{1}{m_{19}}$). Vertical dashed lines in A and B indicate the level of $R_{37/19}$ under dry
953 condition. (C) Normalized signals at $R_{37/19}=1.4$ (equivalent to RH=90% at 25 °C)
954 relative to dry condition for different VOC species as a function of their proton affinities.
955 Data points are color-coded using the permanent dipole moment of the species. The two
956 purple vertical dashed lines are proton affinities of water and water dimer, respectively.
957 The two black dashed lines indicate the linear fits to the species with permanent dipole
958 moment larger than 1 ($y=-1.30+2.52\times 10^{-3}\times\text{PA}$) and lower than 1 ($y=-1.38+2.29\times 10^{-3}\times\text{PA}$),
959 respectively.
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962

963 Figure 6. (A) Normalized signals relative to dry condition for protonated and dehydrated
964 ions of isopropanol as a function of $R_{37/19}$ ratios. The vertical dashed line indicates the
965 level of $R_{37/19}$ under dry condition. (B) Normalized signals relative to dry condition for
966 protonated ion of ethanol as a function of $R_{37/19}$ ratios.

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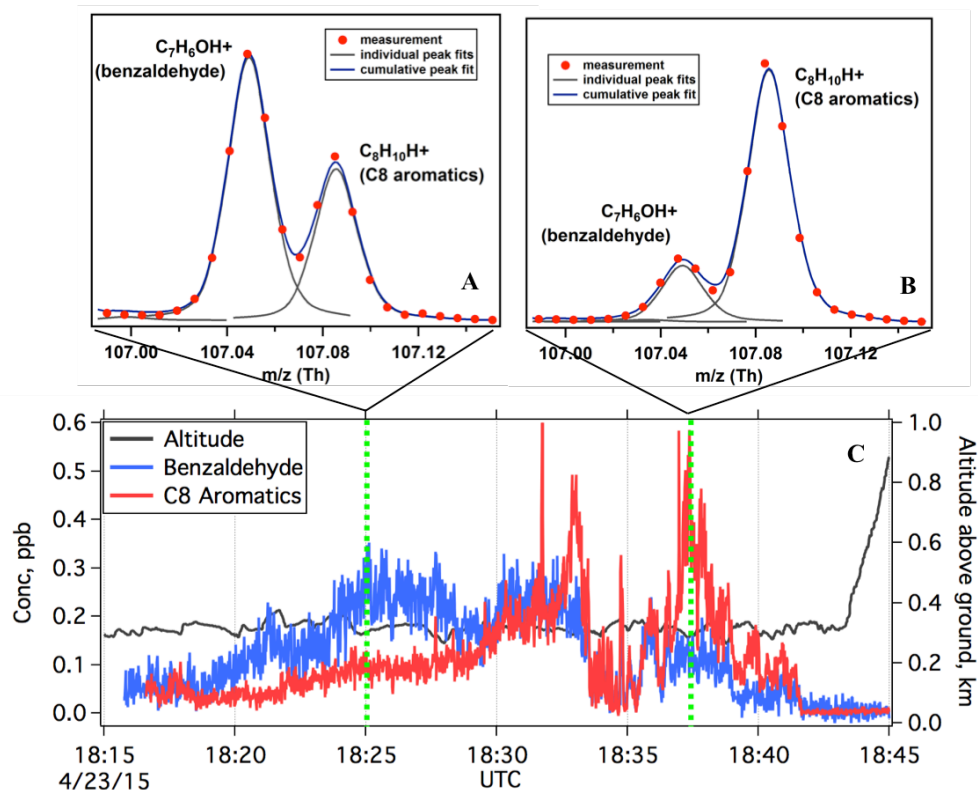
969 Figure 7. (A) The ratios of hydrated ions to protonated ions for several OVOC species as
 970 a function of $R_{37/19}$ ratios. (B and C) Normalized signals relative to dry condition for
 971 protonated and hydrated ions of formic acid (HCOOH, B) and isocyanic acid (HNCO, C)
 972 as a function of $R_{37/19}$ ratios. The ratios of hydrated ions to protonated ions for formic
 973 acid and isocyanic acid versus $R_{37/19}$ ratios are also shown in (B) and (C), respectively.
 974 Vertical dashed lines in each panel indicate the level of $R_{37/19}$ under dry condition.

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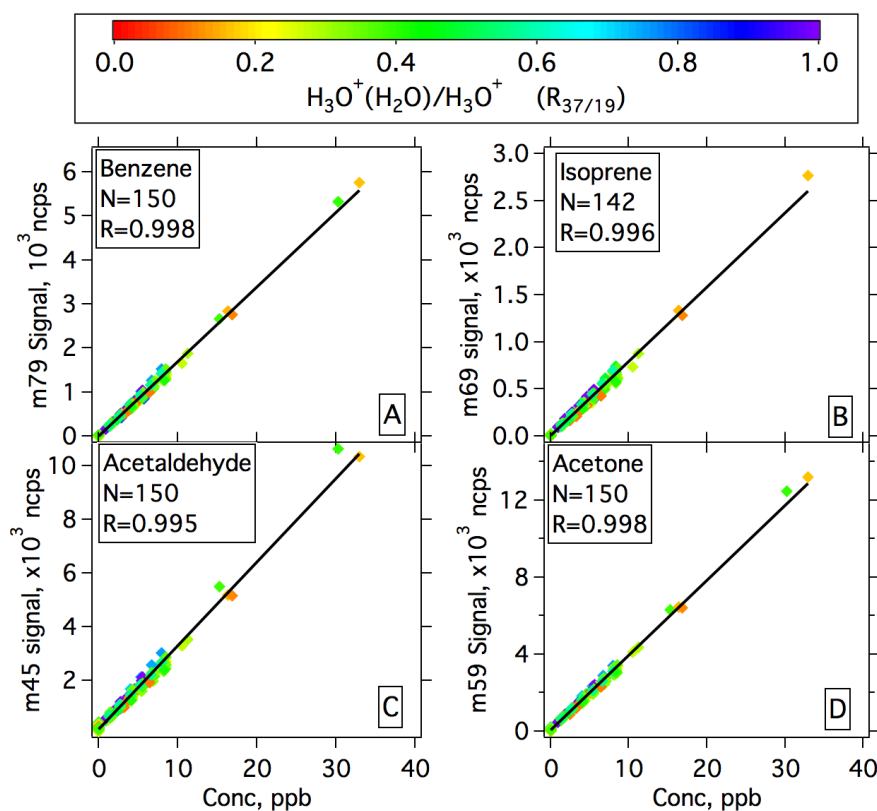
979

980 Figure 8. Separation of benzaldehyde and C8 aromatics from the nominal mass at m/z
981 107 in a period (18:15-18:45 UTC) during the flight over the Permian Basin on April 23,
982 2015. (A and B) HR peak fittings of mass spectra at m/z 107 at 18:25 and 18:37,
983 respectively. (C) Time series of benzaldehyde and C8 aromatics during the period. Flight
984 altitude of the NOAA WP-3D aircraft is included in C.

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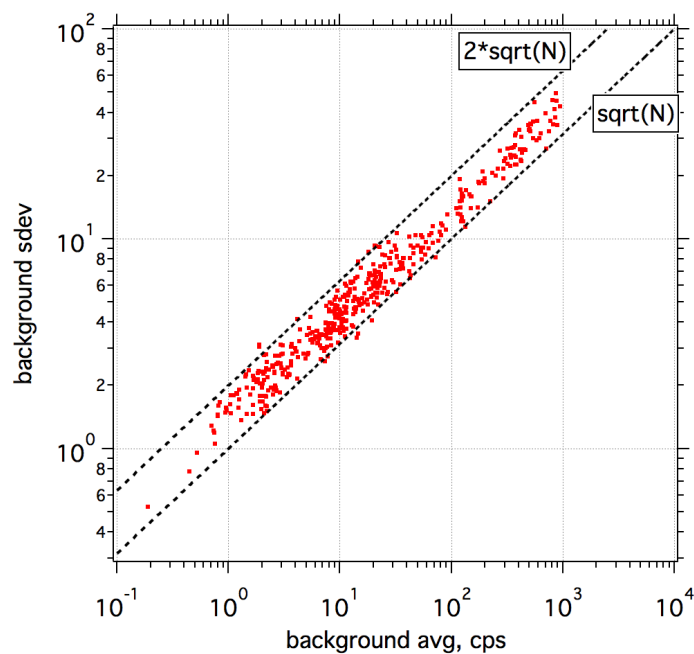


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989 Figure 9. Inflight calibration results of benzene (A), isoprene (B), acetaldehyde (C) and
990 acetone (D) during the SONGNEX campaign. Signals for the species have been humidity
991 corrected. Data points are color-coded using $R_{37/19}$ ratios. The total times (N) when gas
992 standard was introduced into the instrument during SONGNEX and correlation
993 coefficients (R) are shown in the textbox.

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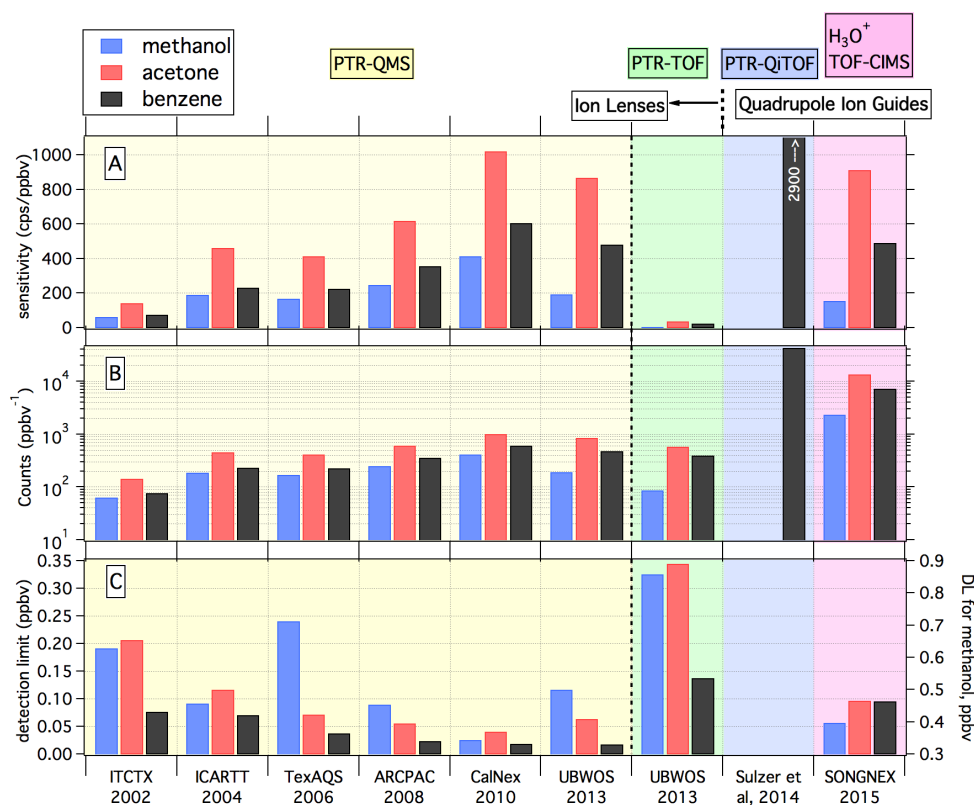
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997 Figure 10. Scatterplot of the standard deviations of background signals versus the
998 measured background signals from a flight on April 27, 2015 during SONGNEX. In this
999 graph, the signals are not corrected for the ToF duty cycle to better reflect the counting
1000 statistics of the ToF detector. The two dashed lines are \sqrt{N} and $2\times\sqrt{N}$, respectively.

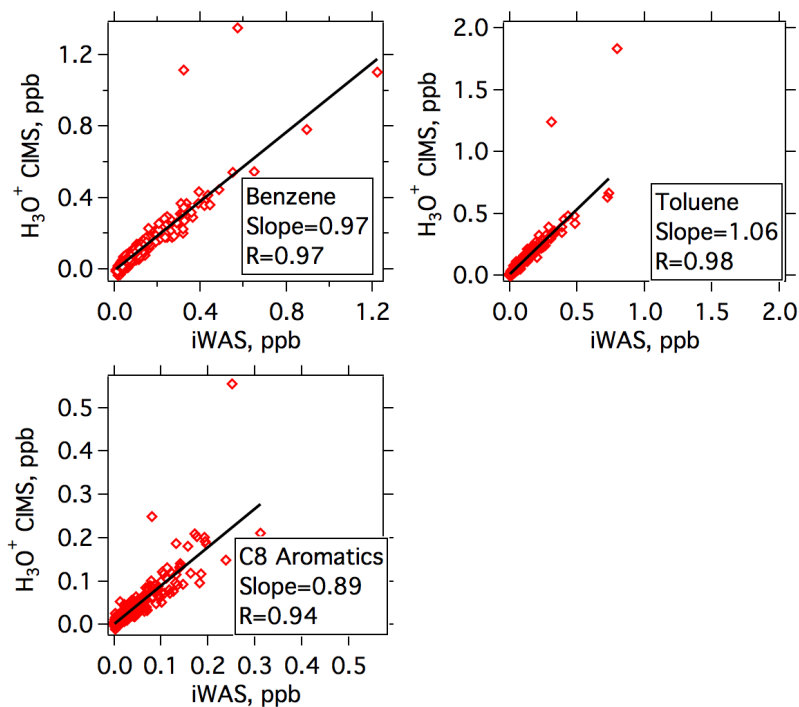
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1003 Figure 11. Comparison of the sensitivities (A), total ion counts from measurements in a
 1004 period of 15 s (B) and 1-s detection limits (C) of methanol, acetone and benzene for the
 1005 NOAA PTR-QMS (Warneke et al., 2011), a PTR-TOF during UBWOS 2013 (Warneke
 1006 et al., 2015), the PTR-QiTOF in Sulzer et al. (2014) and the H₃O⁺ ToF-CIMS presented
 1007 here during SONGNEX. In (B), we assume that the PTR-QMS was operated at the
 1008 typical cycle length of 15 s with 1-s dwell time for the three compounds as have been the
 1009 norm during our aircraft measurements. Note that the detection limits (DL) of methanol
 1010 in panel C are drawn on the right axis.

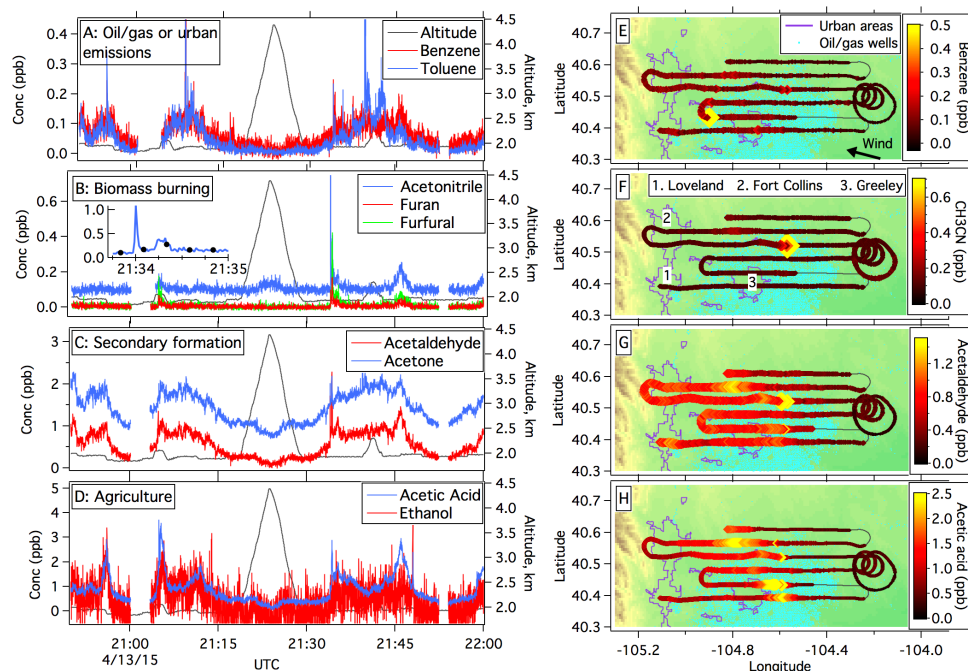
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1013 Figure 12. Comparison of measured mixing ratios of benzene, toluene and C8 aromatics
1014 between the H_3O^+ ToF-CIMS and iWAS (Lerner, in prep) during the SONGNEX
1015 campaign.

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1018 Figure 13. VOC measurements by the H_3O^+ ToF-CIMS for the characterization of air
1019 masses and atmospheric photochemistry during a part of the flight (20:50–22:00 UTC) on
1020 April 13, 2015 over the Denver-Julesburg basin. Time series of various VOC species
1021 from oil/gas or urban emissions (A), biomass burning (B), secondary formation (C) and
1022 agricultural emissions (D) are shown. The inserted plot in B highlights acetonitrile
1023 around 21:34 UTC, where the black dots indicate re-sampled acetonitrile data every 15 s,
1024 reflecting measurement results if a PTR-QMS had been deployed onboard the NOAA
1025 WP-3D. Flight tracks color-coded using mixing ratios of benzene (E), acetonitrile (F),
1026 acetaldehyde (G) and acetic acid (H) are shown in the right panels. Urban regions and
1027 locations of oil and gas wells are indicated in the right panels. The arrow in E indicates
1028 wind direction in the boundary layer during this period.

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