



1    **Uncertainties of ground-based microwave radiometer retrievals in**  
2    **zenith and off-zenith methods under snow conditions**

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13   **Abstract.** This paper is to investigate the uncertainties of microwave radiometer  
14   (MWR) retrievals in snow conditions and also explore the discrepancies of MWR  
15   retrievals in zenith and off-zenith methods. The MWR retrievals were averaged in the  
16    $\pm 15$  min period centered at sounding times of 00:00 and 12:00 UTC and compared  
17   with the radiosonde observations (RAOBs). In general, the MWR retrievals have a  
18   better correlation with RAOB profiles in off-zenith method than in zenith method,  
19   and the biases (MWR observations minus RAOBs) and root mean square errors  
20   (RMSEs) between MWR and RAOB are also clearly reduced in off-zenith method.  
21   The biases of temperature, relative humidity, and vapor density decrease from 4.6 K,  
22   9 %, and  $1.43 \text{ g m}^{-3}$  in zenith method to -0.6 K, -2 %, and  $0.10 \text{ g m}^{-3}$  in off-zenith



1 method, respectively. The discrepancies between the MWR retrievals and the RAOB  
2 profiles along with the altitude present the same situation. Case studies show that the  
3 impact of snow on accuracies of the MWR retrievals is more serious in heavy  
4 snowfall than that in light snowfall, but the off-zenith method can mitigate the impact  
5 of snowfall. The MWR measurements become less accurate in snowfall is mainly due  
6 to the retrieving method which does not consider the effect of snow, and the  
7 accumulated snow on the top of radome increases the signal noise of MWR  
8 measurement. As the snowfall drops away by gravity in the sides of the radome and  
9 the off-zenith observations are more representative of the atmospheric conditions for  
10 RAOBs.

11

12 **Key words:** Microwave radiometer, Retrieval uncertainties, Off-zenith method,  
13 Snowfall

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## 15 **1. Introduction**

16 Atmospheric profiles of temperature, relative humidity, and vapor density can be  
17 retrieved from ground-based microwave radiometer (MWR) measurements (Sánchez  
18 et al. 2013; Ware et al. 2013). These profiles are available nearly continuously and are  
19 extensively utilized in the forecasting and analysis of intense convective weather, also  
20 they have been assimilated into numerical weather prediction models (Marzano et al.  
21 2005; Knupp et al. 2009; Löhnert et al. 2012; Madhulatha et al. 2013). The instability  
22 indices calculated from the MWR-retrieved thermodynamic atmospheric profiles are



1 also employed in operational meteorology (Chan et al. 2010; Cimini et al. 2015;  
2 Leena et al. 2015). However, since the radiative transfer model used in the MWR  
3 does not consider the impact of precipitation on the MWR brightness temperature  
4 measurements, the MWR retrievals become less accurate under precipitation  
5 conditions (Ware et al. 2004; Xu et al. 2014). To improve the accuracy of MWR  
6 retrievals in rainy conditions, some methods are performed to minimize the influence  
7 of liquid water on MWR measurements. The MWR is equipped with a hydrophobic  
8 radome and a special blower system, which can sweep water beads and snow away  
9 from the radome (Chan 2009). A method based on linear regression is also employed  
10 to reduce the discrepancy between the MWR retrievals and the radiosonde  
11 observation (RAOB) profiles (Sánchez et al. 2013). Recently, the off-zenith method is  
12 applied in MWR observations and off-zenith retrievals provide higher accuracy  
13 during precipitation by minimizing the effect of liquid water on the radiometer  
14 radome (Cimini et al. 2011, 2015; Ware et al. 2013; Xu et al. 2014).

15 Snow, a special type of precipitation, has distinct scattering characteristics in the  
16 microwave. Some methods are explored to investigate these characteristics and  
17 discuss their utilization on the snow measurements (Matrosov et al. 2008; Löhnert et  
18 al. 2011; Xie et al. 2012). The scattering signal of snow is highly dependent on the  
19 assumption of snow shape and snow size distribution (SSD), especially for  
20 large-sized parameters (Kneifel et al. 2010). Some studies have demonstrated that  
21 snowfall can significantly reduce the measurement accuracy of MWR (Knupp et al.  
22 2009; Cimini et al. 2011; Ware et al. 2013). However, few studies are reported on the



1 improvements of MWR measurement accuracies in snow conditions. Moreover, in  
2 contrast with rain, snow usually freezes on the top of the radome, and it is not easily  
3 blown away from the radome by the blower system attached on the MWR. Since  
4 MWR retrieval accuracies generally are better in off-zenith method than in zenith  
5 method under precipitation conditions (Xu et al., 2014) and snow does not easily  
6 accumulate on the sides of a radome, we attempt to employ off-zenith method to  
7 improve the MWR retrieval accuracies during snowfall.

8 This paper is organized as follows: Section 2 will briefly describe the data and  
9 methodology employed in this study; Section 3 compares the MWR-retrieved  
10 atmospheric profiles of temperature, relative humidity and vapor density with RAOB  
11 profiles obtained at Wuhan station, then discusses the accuracies of MWR retrievals  
12 under snow conditions and the effect of off-zenith method on it; and Section 4 gives  
13 some conclusions.

14

## 15 **2. DATA AND METHODOLOGY**

16 The data used in this study are collected in the Wuhan operational station ( $30.6^\circ$   
17 N,  $114.1^\circ$  E, and 23 m above sea level), including RAOB data, meteorological  
18 observation data and MWR data. The distances between them are all less than 30 m.  
19 RAOB data is the operational data, which is obtained at 00:00 and 12:00 UTC every  
20 day. The profiles of temperature and relative humidity are obtained by the Chinese  
21 GTS1-2 digital radiosonde at a high vertical resolution of 10 m, and the profiles of  
22 vapor density can be calculated from them. The meteorological observation data are



1 used to confirm the snowfall cases. The MWR data used in this paper is provided by  
2 a MP-3000A unit manufactured by Radiometrics, observing at 2 elevation angles  
3 (zenith and  $15^\circ$  elevation) up to 10 km. The MWR data has a higher temporal  
4 resolution of  $\sim 3$  min, and the vertical intervals are 50 m from the surface to 500 m,  
5 100 m to 2 km, and 250 m to 10 km (Ware et al. 2013; Xu et al. 2014).

6 The MP-3000A unit observes brightness temperature at up to 35 channels,  
7 including 21 K-band (22–30 GHz) and 14 V-band (51–59 GHz). Moreover, an  
8 infrared radiation thermometer (IRT) is equipped on the MWR, which measures sky  
9 infrared temperature at one zenith infrared (9.6–10.5  $\mu\text{m}$ ) channel and gives  
10 information on cloud-base temperature (Ware et al. 2013; Cimini et al. 2015; Xu et al.  
11 2015). Meteorological sensors attached to the MWR can obtain ambient temperature,  
12 pressure, and relative humidity at the instrument level. The retrieved algorithm  
13 developed by the factory can automatically convert the microwave, infrared, and  
14 surface meteorological measurements into temperature, humidity, and liquid profiles  
15 using radiative transfer equations with the aid of neural networks (Xu et al. 2015).  
16 The neural network retrieval method uses historical radiosondes to characterize states  
17 of the atmosphere that commonly occur at a particular location (Ware et al. 2013). A  
18 five-year data set of historical radiosondes in Wuhan was used for neural network  
19 training (Xu et al. 2014).

20 Three snow cases (shown in Table 1) are selected to present the comparison of the  
21 profiles between MWR and RAOB under snow conditions, and the effect of  
22 off-zenith method on improving the MWR measurement accuracy during snowfall is



1 explored. All cases in this study include at least one RAOB profile during snowfall.  
2 Since it takes 30 minutes for the balloon from the surface to 10 km altitude in  
3 sounding, the MWR retrievals were averaged in the  $\pm 15$  minute period centered at  
4 sounding times of 00:00 and 12:00 UTC and compared with the RAOB profiles.  
5 Considering the vertical resolution of the RAOB profiles is not consistent with that of  
6 MWR retrievals, the RAOB profiles are interpolated to the height levels of the MWR  
7 retrievals. Based on the above process, there are eight temporal pairs of MWR and  
8 RAOB profiles for comparison in this study. Methods used in this study are simply  
9 employed to calculate the correlation coefficients, bias (MWR observation minus  
10 RAOB), and root mean square error (RMSE) between the MWR and the RAOB for  
11 each parameter in zenith and off-zenith methods. The discrepancies between MWR  
12 retrievals and RAOB profiles at different heights are also calculated to explore how  
13 the MWR retrievals accuracies vary with height.

14

### 15 **3. RESULTS ANALYSIS**

#### 16 **3.1 Uncertainties of MWR retrievals in zenith and off-zenith methods under** 17 **snow conditions**

18 To explore the effect of off-zenith method on MWR measurement accuracy, the  
19 simultaneous MWR zenith and off-zenith retrievals around the time of 00:00 and  
20 12:00 UTC are compared with the RAOB profiles. Table 2 presents the comparison  
21 of MWR retrievals against RAOB profiles in zenith and off-zenith methods under  
22 snow conditions without considering the level division in altitude. All the MWR



1 retrievals have a better correlation in off-zenith method than that in zenith method  
2 especially for relative humidity, and the biases and RMSEs are also clearly reduced in  
3 off-zenith method. For temperature, the MWR zenith observations have a warm bias  
4 of 4.6 K against RAOBs while in off-zenith method the bias decrease to -0.6 K, with  
5 RMSE also decreasing from 5.7 K to 2.0 K. The MWR-retrieved relative humidity  
6 has poor agreement with RAOB relative humidity in zenith method but reasonable in  
7 off-zenith method, and the bias and RMSE also decrease from 10 % and 33% in  
8 zenith method to -2 % and 20 % in off-zenith method, respectively. For vapor density,  
9 the correlation coefficient between MWR observations and RAOBs increases from  
10 0.7130 in zenith method to 0.9389 in off-zenith method. In zenith method, the bias is  
11  $1.43 \text{ g m}^{-3}$  with a RMSE of  $2.14 \text{ g m}^{-3}$ , while in off-zenith method both of them  
12 decrease to  $0.10 \text{ g m}^{-3}$  and  $0.66 \text{ g m}^{-3}$ , respectively. Obviously, the MWR retrievals  
13 have better accuracies against RAOBs in off-zenith method than in zenith method.

14 To further compare the uncertainties of MWR retrievals against RAOBs in  
15 zenith and off-zenith methods, the discrepancies between the MWR retrievals and the  
16 RAOB profiles along with the altitude under snow conditions are also investigated.  
17 As shown in Fig. 1, the temperature correlation coefficients in zenith method are  
18 smaller than those in off-zenith method below 6 km especially around 3.75 km where  
19 the correlation coefficient rapidly increases from 0.01 to 0.92, but the situation is  
20 opposite above 6 km. The MWR temperature shows a warm bias against RAOB in  
21 zenith method and the bias is larger than 3 K at most heights, while in off-zenith the  
22 bias becomes cold and within -1 K at most heights. Both the MWR temperature



1 RMSEs in zenith and off-zenith methods approximately increase with height, but the  
2 RMSE is clear smaller in off-zenith method. The MWR temperature RMSE is greater  
3 than 4 K above 0.5 km in zenith method while in off-zenith method it is within 2 K at  
4 most heights.

5 Fig. 2 presents the results for the relative humidity profiles. The correlation  
6 coefficients between MWR observations and RAOBs are negative at most heights  
7 below 2.5 km. Compared with zenith observations, off-zenith observations have well  
8 agreement with RAOBs above 4.5 km. The correlation coefficient cannot be  
9 calculated in some altitudes because the compared RAOB relative humidity remains  
10 constant at these altitudes, so some breakpoints are shown in the Fig. 2a. The biases  
11 of zenith and off-zenith observations are negative below 5 km and there are no  
12 distinct differences between them. Above 6 km, both the biases in zenith and  
13 off-zenith methods increase with height, but the bias is clear smaller in off-zenith  
14 method. It is the same situation for the RMSE, the RMSE differences between zenith  
15 and off-zenith observations are not evident below 5 km, while above 5 km the RMSE  
16 is clearly smaller in off-zenith observations.

17 The comparison results for the vapor density profiles are shown in Fig. 3. It can  
18 be seen that the correlation coefficient in zenith observation is positive below 3.5 km  
19 but mostly negative above 3.5 km, while in off-zenith observation it is positive except  
20 around 3 km. In general, the correlation coefficient is more reasonable in off-zenith  
21 method than in zenith method. The bias of vapor density in zenith observation  
22 increases from  $0 \text{ g m}^{-3}$  at surface to  $5.51 \text{ g m}^{-3}$  at 2 km and then decreases to near  $0 \text{ g}$





1  $\text{m}^{-3}$  at 10 km again, but in off-zenith observation the bias is clear smaller with a value  
2 within  $\pm 1.0 \text{ g m}^{-3}$ . Both the RMSEs in zenith and off-zenith observations vary  
3 similarly with height, in which the RMSE in zenith (off-zenith) observation firstly  
4 increases to 3 km (2.3 km) and then decreases to near  $0 \text{ g m}^{-3}$  at 10 km. Although the  
5 RMSE has a close value in zenith and off-zenith observations, it is also clear smaller  
6 in off-zenith observation. The RMSE in zenith observation is mostly greater than 1.0  
7  $\text{g m}^{-3}$  with a peak of  $2.60 \text{ g m}^{-3}$ , yet it is generally smaller than  $1.0 \text{ g m}^{-3}$  with a peak  
8 of  $1.47 \text{ g m}^{-3}$ .

9       Based on the above analysis, it is clearly that snowfall has a significant impact  
10 on MWR measurement accuracy, and off-zenith method can improve the accuracies  
11 of MWR retrievals under snow conditions, especially for the temperature and vapor  
12 density retrievals. Snowfall, one of precipitation, does not be considered in the MWR  
13 retrieving method, so the MWR-retrieved atmospheric profiles in snow conditions are  
14 not reasonable as those in non-precipitation conditions (Xu et al, 2014). Although a  
15 special blower system is used to sweep water beads and snow away from the radome,  
16 snowfall, particularly heavy snowfall will always freeze on the radome in the low  
17 temperature situation. Snow produces a strong scattering signal in the microwave  
18 region and the snow ice will increase signal noise of MWR measurement, so the  
19 frozen snow on the radome will have great influence on the MWR measurement of  
20 brightness temperature. Compared to zenith method, off-zenith method has better  
21 measurement accuracies under snow conditions. This is mainly because that the  
22 MWR observes at  $15^\circ$  elevation through vertical sections of the inverted “U” shaped



1 radome that are more readily cleared of snow/water droplets by gravity than the  
2 horizontal sections observed at zenith. Moreover, the MWR accuracies are related  
3 with the balloon drifting in sounding due to the wind in atmosphere (Xu et al., 2015),  
4 and the off-zenith observations are more representative of the conditions in which  
5 radiosonde observations are also taken (Xu et al., 2014), thus the MWR measurement  
6 accuracies are generally better in off-zenith method than in zenith method.

### 7 **3.2 Case study**

8 To better understand the effect of off-zenith method on the improvement of  
9 MWR retrieval accuracy, the comparison between the time series of the MWR  
10 retrievals in a heavy snowfall and a light snowfall are performed. The heavy snowfall  
11 happens from 00:07 UTC 5 February to 04:15 UTC 7 February in 2014 with  
12 cumulative snowfall of 28.0 mm and the light snowfall happens from 07:16 UTC 8  
13 February to 04:22 UTC 9 February in 2014 with cumulative snowfall of 2.3 mm.

14 As shown in Fig. 4, the MWR-retrieved temperature in zenith method presents a  
15 clear increase at ~2.5 km in the heavy snowfall, but the increase is not clear in  
16 off-zenith method. The MWR-retrieved temperature in zenith method is about 10 K  
17 warmer than that in off-zenith method when the snowfall happens, and the greater  
18 temperature is well accordant with the snowfall time. The clearly warmer temperature  
19 disappeared in 1 h after the end of heavy snowfall. Fig. 5 illustrates the situation in  
20 the light snowfall. The MWR temperature discrepancies between zenith and  
21 off-zenith methods are not significant as those in the heavy snowfall, and the MWR  
22 temperatures in zenith method are about 3 K warmer than those in off-zenith method



1 at ~2.5 km when the snowfall happens. The greater temperature is not obvious when  
2 light snowfall maybe due to the light snow on the radome is blown away immediately  
3 by the special blower system.

4 The MWR-retrieved temperatures have well agreement between zenith and  
5 off-zenith method in light snow condition, while they have poor agreement in heavy  
6 snowfall. Although a special blower system is used to sweep water beads and snow  
7 away from the radome, the heavy snowfall is hardly blown away and will easily froze  
8 on the radome. Frozen snow will have great influence on the MWR measurement of  
9 brightness temperature, so heavy snowfall has more effect on the MWR observations  
10 comparing with light snowfall. The greater temperature in zenith method is probably  
11 caused by the discrepancies of MWR-measured brightness temperature, and this will  
12 helpful to explain why the greater temperature is significant in heavy snow condition.  
13 Off-zenith method significantly minimizes contamination from ice and snow, so the  
14 MWR-retrieved temperature in zenith method is more reasonable especially when  
15 heavy snowfall.

16 The MWR relative humidity discrepancies in zenith and off-zenith methods are  
17 also significant in the heavy snowfall (Fig. 6). Although the MWR relative humidity  
18 presents good agreement in zenith and off-zenith methods below 2.5 km, the MWR  
19 relative humidity retrievals in zenith method are clear larger than those in off-zenith  
20 method above 5 km, about 40 % at 7 km. Greater MWR relative humidity appears  
21 above 7 km in zenith method and also well consistent with the timing of the heavy  
22 snowfall, while this situation disappears in the off-zenith method. However, the



1 discrepancies between zenith and off-zenith methods are not clear in the light  
2 snowfall, and the variation of the relative humidity is also more stable (Fig. 7). The  
3 bottom of atmosphere is almost saturated when snowfall happens and we attribute  
4 that snow will easily sublimation in the blowing of the special thermodynamic blower  
5 system.

6 The situation for the vapor density is the same as the temperature. As shown in  
7 Fig. 8, the MWR vapor density retrievals in zenith method are significantly larger  
8 than those in off-zenith method at  $\sim 2.5$  km in the heavy snowfall, and the time of  
9 vapor density increasing is also consistent with the heavy snowfall time. The heavy  
10 snowfall will also reduce the retrieval accuracies of vapor density by influencing the  
11 brightness temperature measurements of MWR, thus the trend of vapor density  
12 variation in zenith method is similar to that of temperature with heavy snowfall.  
13 While in off-zenith method, the MWR vapor density retrievals are more reasonable  
14 without the significantly larger area. In the light snowfall (Fig. 9), the MWR vapor  
15 density retrievals present a similar trend in zenith and off-zenith methods, but the  
16 former is clearly larger than the latter below 3 km.

17 Obviously, the MWR retrieval discrepancies between zenith and off-zenith  
18 methods are greater in heavy snowfall than that in light snowfall, As mentioned  
19 before, this is mainly because that the snowfall is more easy to freeze on the radome  
20 top in heavy snowfalls and the signal noise caused by snowfall increases, while in the  
21 sides of the radome the snowfall drops to the ground for gravity, so the MWR  
22 retrieval discrepancies are greater in heavy snowfalls. However, in light snowfalls,



1 the blower system can sweep some snowfall away, so the impact of snowfall is not  
2 greater as that in heavy snowfalls.

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#### 5 **4. CONCLUSIONS**

6 In this paper, the MWR retrieval accuracies in snow conditions are discussed by  
7 comparing with the RAOBs and improvements of off-zenith method are also  
8 investigated when snowfall happens. We also present two snowfall cases to explore  
9 the MWR retrieval accuracy in heavy and light snow conditions. Based on the above  
10 analysis, we draw the following conclusions:

11 1. Without considering the division of altitude, all the MWR retrievals have a  
12 better correlation with RAOB profiles in off-zenith method than that in zenith method  
13 especially for relative humidity when snowfall happens, and the biases and RMSEs  
14 are also clearly reduced in off-zenith method. The temperature bias and RMSE  
15 decrease from 4.6 K and 5.7 K in zenith method to -0.6 K and 2.0 K in off-zenith  
16 method, respectively. The relative humidity bias and RMSE also decrease from 10%  
17 and 33% in zenith method to -2 % and 20 % in off-zenith method, respectively, while  
18 the correlation coefficient increases from 0.2531 to 0.7997. For vapor density, the  
19 bias is  $1.43 \text{ g m}^{-3}$  with a RMSE of  $2.14 \text{ g m}^{-3}$  in zenith method, while in off-zenith  
20 method the bias decreases to  $0.10 \text{ g m}^{-3}$  with a smaller RMSE of  $0.66 \text{ g m}^{-3}$ .

21 2. The discrepancies between the MWR retrievals and the RAOB profiles along  
22 with the altitude under snow conditions are also investigated. The MWR temperature



1 shows a warm bias against RAOB in zenith method and the bias is larger than 3 K at  
2 most heights, while in off-zenith the bias becomes cold and is within -1 K at most  
3 heights. The temperature RMSE is greater than 4 K above 0.5 km in zenith method  
4 while in off-zenith method it is within 2 K at most heights. The vapor density  
5 retrievals show the same situation, the bias and RMSE are clear smaller in off-zenith  
6 method than in zenith method at most height. The off-zenith relative humidity  
7 retrievals show a better agreement with RAOBs above 4.5 km but the correlation  
8 coefficients are negative in zenith method. Although the differences between zenith  
9 and off-zenith methods in relative humidity bias and RMSE are insignificant below 5  
10 km, the bias and RMSE are clearly smaller in off-zenith method above 6 km.

11 3. Case studies show that the heavy snowfall has an obvious impact on the  
12 accuracies of MWR retrievals by influencing the MWR brightness temperature  
13 measurements, and the off-zenith method greatly mitigates the impact of snowfall.  
14 The zenith retrievals have an increase trend during heavy snowfall process, but the  
15 MWR retrievals in off-zenith method are smooth without the higher retrievals  
16 appearing in zenith method.

17 4. The MWR measurements become less accurate in snowfall is mainly due to  
18 the retrieving method which does not consider the effect of snow, moreover, the  
19 snowfall accumulating on the radome especially in heavy snowfalls also increases the  
20 signal noise of MWR measurement. As the snowfall drops away by gravity in the  
21 sides of the radome and the off-zenith observations are more representative of the  
22 atmospheric conditions for RAOB, the off-zenith method makes a positive effect on



1 mitigating the impact of snowfall,

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3

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 12

13 **Tables:**

14

15 Table 1. Details of three snow cases used in this study

Start time of snowfall	End time of snowfall	Cumulated snowfall (mm)
00:07 UTC 5 Feb 2014	04:15 UTC 7 Feb 2014	28.0
07:16 UTC 8 Feb 2014	04:22 UTC 9 Feb 2014	2.3
12:00 UTC 17 Feb 2014	01:38 UTC 18 Feb 2014	11.1

16

17 Table 2. Comparison of MWR retrievals against RABOs in zenith and off-zenith  
 18 methods under snow conditions when not considering the level division in altitude.



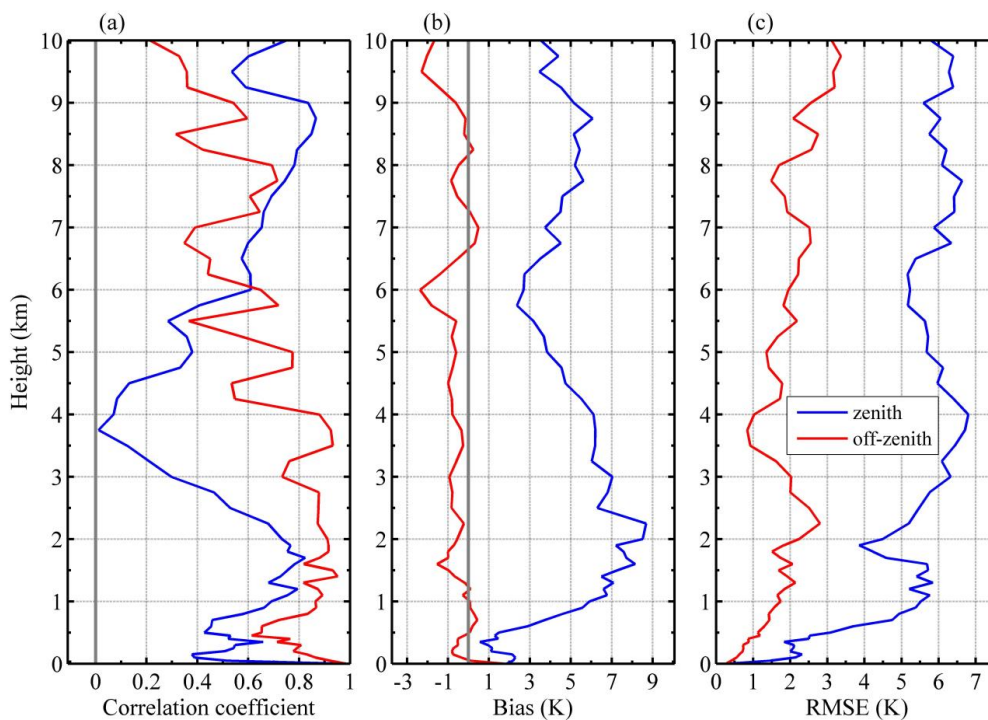
Parameters	Observation mode	Number of samples	Correlation coefficient	Bias	RMSE
Temperature	Zenith	464	0.9239	4.6 K	5.7 K
	Off-zenith	464	0.9890	-0.6 K	2.0 K
Relative Humidity	Zenith	464	0.2531	8.9 %	33.1 %
	Off-zenith	464	0.7997	-2.2 %	20.2 %
Vapor density	Zenith	464	0.7130	1.43 g m <sup>-3</sup>	2.14 g m <sup>-3</sup>
	Off-zenith	464	0.9389	0.10 g m <sup>-3</sup>	0.66 g m <sup>-3</sup>

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2

3 **Figures:**

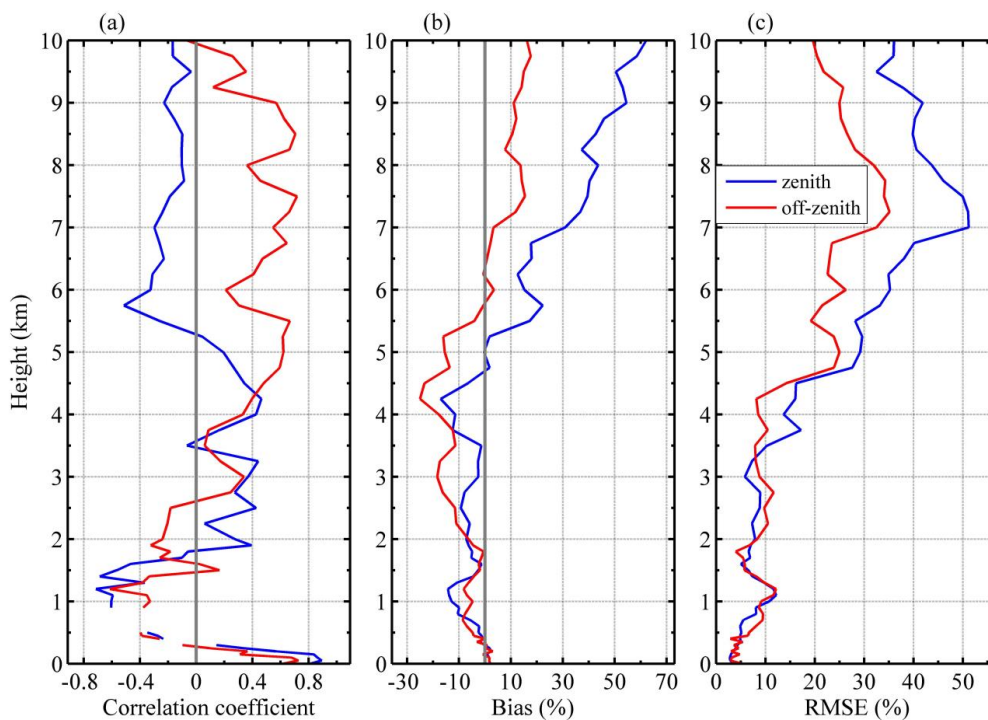
4



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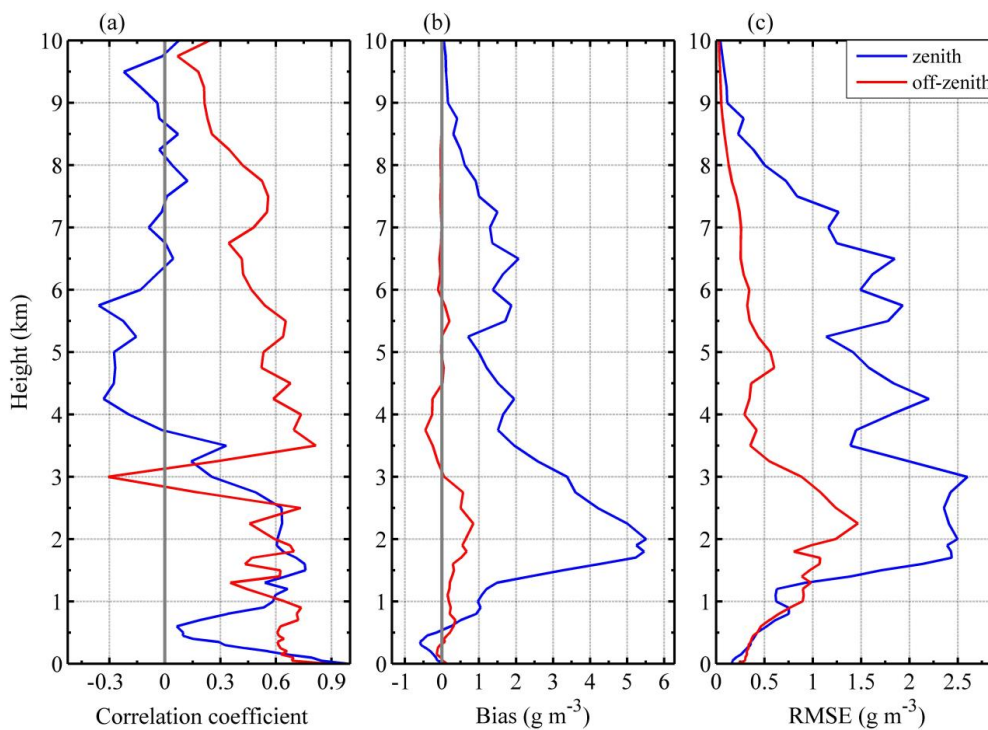
2 Figure 1. The correlation coefficient (a), bias (b) and RMSE (c) between the MWR

3 and RAOB temperature in zenith (blue) and off-zenith (red) observations.



1

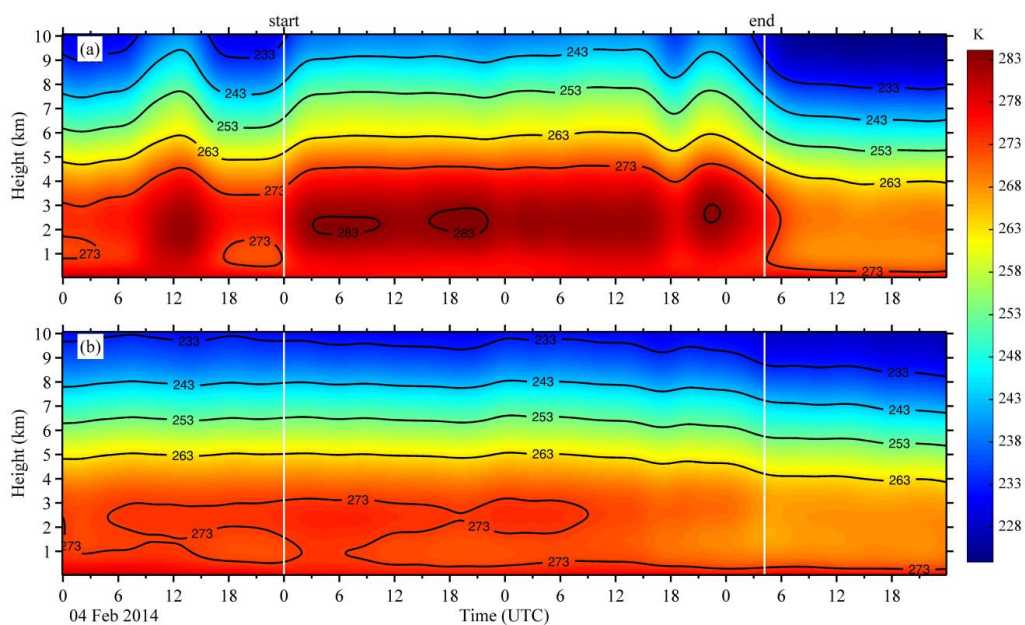
2 Figure 2. Same as Fig. 1 but for relative humidity profiles. Some breakpoints are  
3 shown in Fig. 2a because the compared RAOB relative humidity remains constant at  
4 these altitudes.



1

2 Figure 3. Same as Fig. 1 but for vapor density profiles.

3



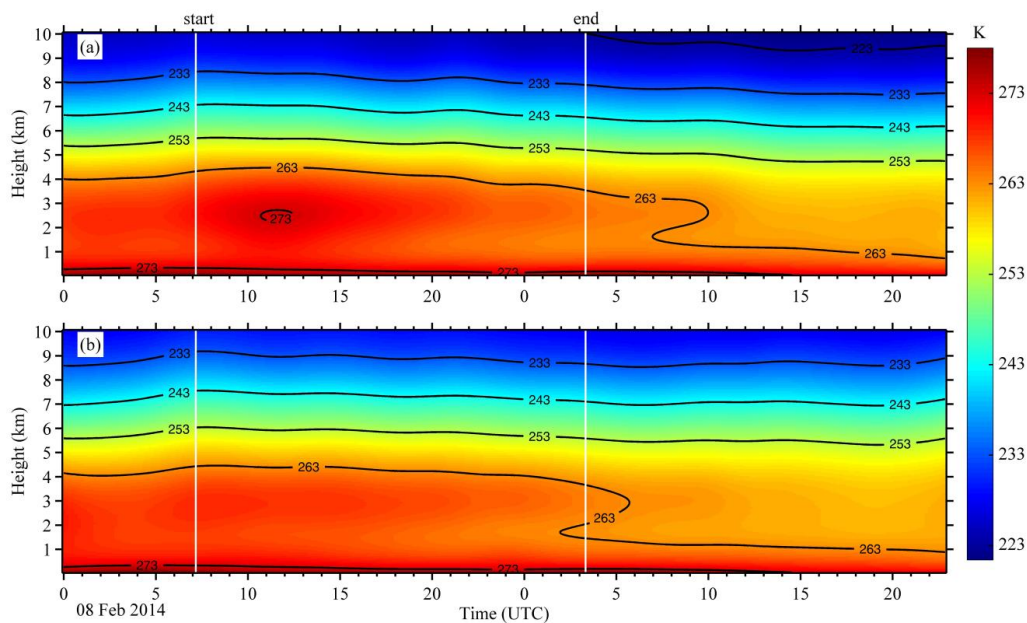
1

2 Figure 4 Comparison of temperature retrievals between zenith (a) and off-zenith (b)

3 observation in heavy snow condition. The start and end times of snowfall are

4 indicated by the vertical lines. The time series starts at 00:00 UTC 04 Feb 2014.

5

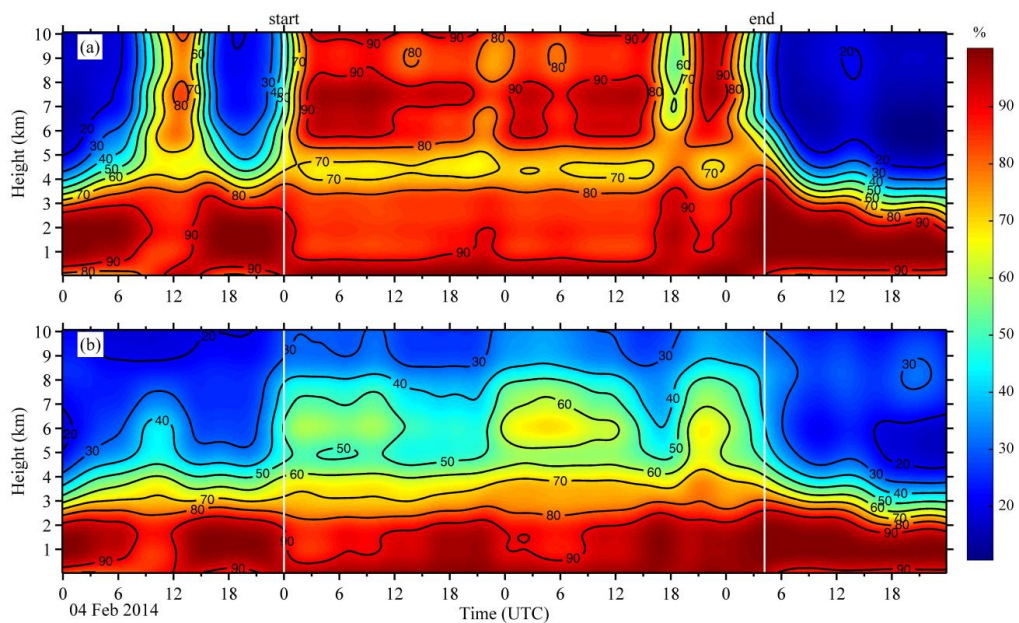


1

2 Figure 5 Comparison of temperature retrievals between zenith (a) and off-zenith (b)  
3 observation in light snow condition. The start and end times of snowfall are indicated  
4 by the vertical lines. The time series starts at 00:00 UTC 08 Feb 2014.

5

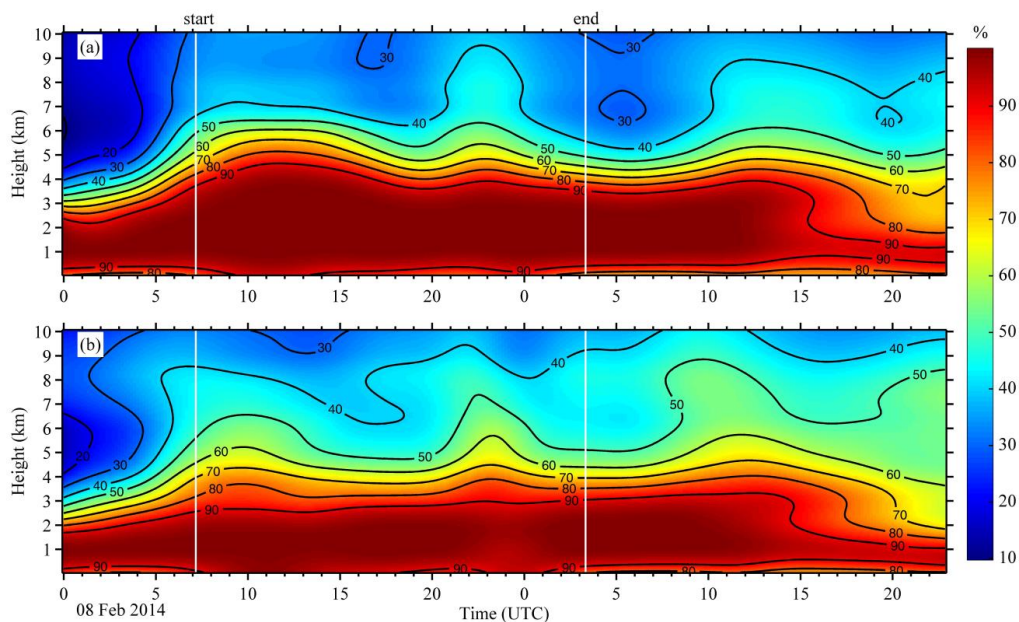




1

2 Figure 6 Same as Fig. 4 but for relative humidity retrievals.

3

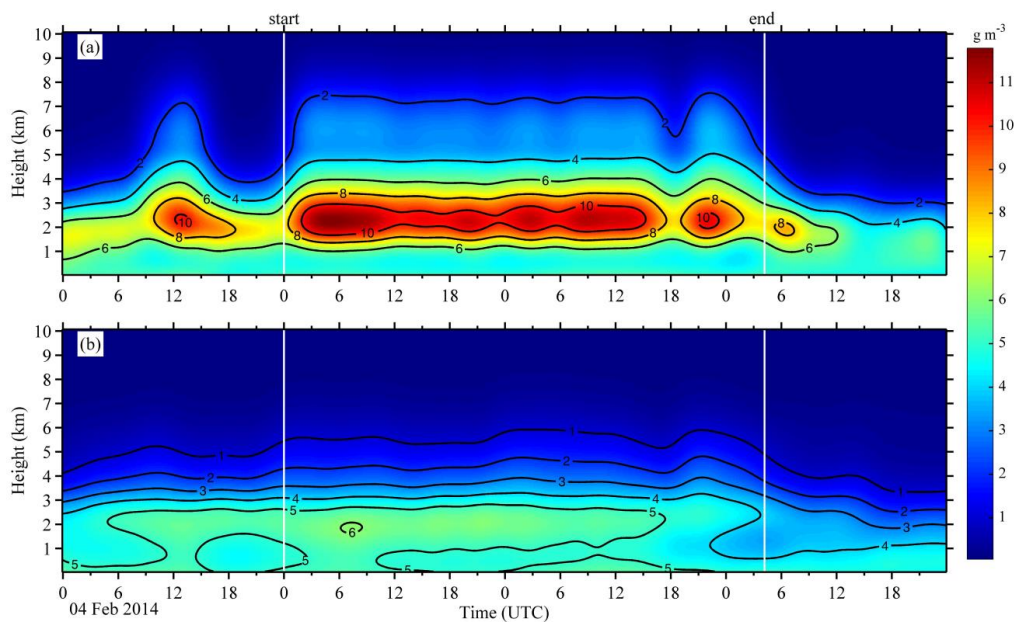


4



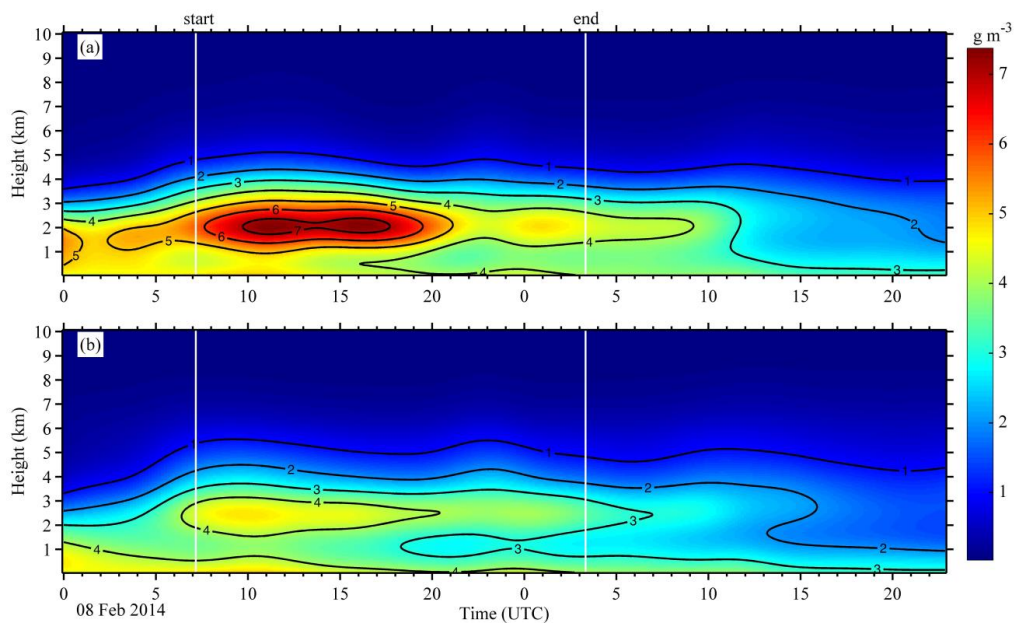
1 Figure 7 Same as Fig. 5 but for relative humidity retrievals.

2



3

4 Figure 8 Same as Fig. 4 but for vapor density retrievals.



1

2 Figure 9 Same as Fig. 5 but for vapor density retrievals.