

Interactive comment on “Advanced characterization of aerosol properties from measurements of spectral optical depth using the GRASP algorithm” by B. Torres et al.

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Review for Atmospheric Measurement Techniques

Title: Advanced characterization of aerosol properties from measurements of spectral optical depth using the GRASP algorithm.

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General Comments: This paper describes a new application of the GRASP algorithm that utilizes accurate spectral AOD measurements from sunphotometer instruments

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(both ground-based and airborne) to retrieve aerosol size distributions. The paper is clearly organized and very well written. This new algorithm provides inversions that may be extremely useful in analysis of particle size related aerosol parameters, since these retrievals can be made much more frequently from AOD spectra alone than from the combined input of AOD spectra and sky radiance distributions, such as in almucantar inversions in AERONET (the Dubovik and King (2000) algorithm).

The sensitivity analysis is clear and thorough, however I am surprised that the perturbations in real refractive index applied in this section are so small. Real refractive index variations of ± 0.02 (for fine mode dominated cases; section 3.5.1) are somewhat trivial and it would seem very difficult to provide with such accuracy when day-to-day (and seasonal) variation in real refractive index can be expected at some sites (such as at GSFC). This is due to hygroscopic growth of particles related to variation in relative humidity and low altitude cloud interaction with aerosol, plus relative changes in aerosol species composition variation. Ideally, the analysis should include much larger uncertainty/variation in the initial guess of the real refractive index.

I agree with the editor and other reviewers that you should add the case of $\text{AOD}(440 \text{ nm})=0.1$ to your sensitivity analysis section, since most data in the AERONET network are for AOD levels <0.3 . Especially for the Lanai site the $\text{AOD}=0.3$ is much too high to be representative. This is a background marine site where AOD is predominately <0.15 . The monthly mean AODs at Lanai do not exceed 0.08 at 440 nm except for the dust transport season in spring when they reach a monthly average maximum of 0.12 in April. I also agree with the editor that it would be useful to include very high AOD cases in your sensitivity analysis (1.5 or 2 at 440 nm) for dust and also fine mode smoke, since these cases are important for analysis of major aerosol events.

The estimation of uncertainty in AERONET measured AOD for field instruments of 0.01 in the visible and NIR and increasing to 0.02 in the UV wavelengths should be mentioned on Page 13 or elsewhere in the manuscript. This wavelength dependent uncertainty was taken into account for operational SDA inversions where only the 380

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to 870 nm wavelengths (5 channels) were utilized as input to the algorithm in order to avoid somewhat higher uncertainties in the 340 nm AOD due to both calibration and filter instability (degradation in time), and for 1020 nm due to temperature sensitivity of the detector at this wavelength and water vapor absorption. It seems that you should make some tests of wavelength range used as input to the code (with accounting for realistic uncertainty for each wavelength) to determine the optimal set of AOD spectral data that should be used as input, plus characterize how the uncertainty changes when specific wavelengths are not available. In particular some AERONET Cimels operated by PHOTONS only have the 440, 675, 870 and 1020 nm channels for AOD measurement. How is the algorithm performance affected when only these 4 wavelengths are available as input to the GRASP-AOD algorithm?

For the Real Cases (Section 4) analysis it is important again that you include cases with AOD of 0.1 at 440 nm for all sites, and for the Lanai site a case of $AOD(440)=0.07$ since this is the climatological mean AOD (for non-dust months, i.e. excluding spring season) for that site. Additionally you should include some analysis of sites that are not included in the Dubovik et al. (2002) Table 1. The vast majority of AERONET sites are not characterized in that table so if the algorithm is to be operationally applied a strategy for selection of first guesses of the input parameters is needed. The way the paper is written it sometimes implies using the almucantar retrievals as a source of first guesses for refractive indices and size, therefore it is presented (in some sections) as a companion retrieval to the almucantar sky radiances retrievals and not a fully independent retrieval such as SDA. This should be discussed or clarified in the paper.

Specific Comments:

Title: You say “. . .characterization of aerosol properties. . .” but shouldn't this be more specific, since only aerosol size related properties are retrieved as the refractive indices are assumed a priori. Therefore it seems that perhaps the word 'size' should be inserted between 'aerosol' and 'properties' in the title.

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Abstract lines 10-11: “Differences in the fine mode volume median radii for the GRASP-AOD and AERONET inversions are less than $0.02 \mu\text{m}$ at sites dominated by the fine mode for all cases. . .” Is this the maximum difference or mean difference?

Abstract line 16: I suggest removing the word ‘advance’ in this sentence since it is confusing.

Introduction, Page 2, line 4: I suggest that you remove ‘Unfortunately’ from this sentence, as it is somewhat confusing.

Page 3, lines 4-8: Please provide more details and justification for utilizing bi-modal assumption with r_v and sigmas for each mode to describe the size distributions rather than use 22 size bins as is done in the AERONET standard retrieval. I assume that it is a more stable retrieval than the 22 bins, given only the AOD spectra as input, but it would be informative to provide some more details on why you selected this way to make the retrievals (either here or in Section 3).

Page 3, lines 11-13: “For instance, AERONET includes around forty direct Sun measurements per day in its standardized sequence of measurements (in cloudless conditions); only about eight of these sequences are coincident with sky-radiance measurements that are suitable as input to the AERONET inversion code.” It would be useful to include information on the new Cimels here, where Hybrid scans make sky radiance retrievals possible each hour throughout the day thereby doubling or more the total number of sky-radiance retrievals. Also many new Cimels are run in the turbo mode for AOD measurement where the time interval of AOD sampling is 3 minutes rather than the 15 minutes for the standard scenario in a Cimel sun-sky radiometer. Therefore the ~ 40 measurements per day increases to about 200 AOD spectra per day (depending on latitude and season, therefore day length) with new Cimels.

Page 3, lines 13-14: “. . . a large amount of data containing only direct Sun measurements which are not used apart from the characterization of the aerosol load.” I disagree, since the Angstrom Exponent and its spectral variation are also useful param-

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ters providing basic aerosol size information that are computed by AERONET from the spectral AOD, and are utilized in many studies.

Page 3, lines 14-15: “Moreover, many AERONET sites are plagued by several months of partial cloudiness (especially in wintertime).” I think this is somewhat misleading, since some sites have the seasonal opposite and some tropical sites are cloudy for most seasons. Therefore I suggest dropping the phrase ‘(especially in wintertime)’ from the sentence.

Page 3, lines 14-15: “As a result, the aerosol loads are the only data reported at many sites. . .” Please insert after ‘aerosol loads’ something like: ‘and Angstrom Exponents that parameterize the relative fine versus coarse mode optical influence, depending on the wavelength range used in the computation of AE (Eck et al. 1999; Schuster et al., 2006).’

Page 4, lines 11-12: Please include the reference of van Donkelaar et al. (2010).

Page 6, lines 4-6: For GRASP-AOD please discuss whether there would be any difference in the retrievals if spherical particle shape assumption was utilized/assumed versus spheroidal shape.

Page 7, lines 7-9: In most cases this is true, that lognormal bi-modal size distribution assumptions are sufficient. However, for cloud-processed aerosol a third middle mode sometimes exists, see Eck et al. 2012 for Dubovik almucantar retrievals that are tri-modal, not bi-modal, and these middle modes are supported by independent in situ measurements of fog/cloud processed aerosols. This should be mentioned as a relatively rare case that does occur however.

Page 8, lines 12: Note that Dubovik 2002 presents an AERONET Version 1 database climatology that has been refined significantly with the AERONET Version 2 reprocessing of the entire database that occurred in 2006.

Page 8, lines 16-17: This is weakly absorbing aerosol at GSFC, non-absorbing aerosol

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would have $SSA=1$.

Page 8, lines 17-18: It should be mentioned that this is strongly absorbing biomass burning (BB) aerosol at Mongu, Zambia and that BB aerosol range from very weakly absorbing to strongly absorbing (see Eck et al, 2003, GRL) depending on fuel types and phase of combustion (flaming versus smoldering).

Page 8, lines 27-28: Why would you assume that for a mixed case (Bahrain) that all particles are non-spherical? Solar Village also has many mixed mode aerosol days with predominately spherical fine mode particles from pollution.

Page 8, lines 12-13: Surprising that you do not show the case of $AOD=0.9$ in Figure 2 since this is the one where the curvature for the fine mode cases would become the most obvious, since the fine mode radius is largest. Note that the 1640 nm AOD showing a departure from a 2nd order fit of AOD versus WL in Figure 2 is not observed in the real GSFC site AERONET measurement data. Perhaps you can show the 2nd order fit to your simulated data in Figure 2 and also show the delta AOD departures from the fit.

Page 9, lines 8-9: Note that the Angstrom Exponents you computed for the Lanai site are much too high, you have 1.22 in Table 2, while the measurement data yield averages of ~ 0.6 to 0.8 for Lanai data (see the AERONET climatology tables). This possibly suggests an issue with the Dubovik 2002 climatology parameters for this site. Also note that for the Solar Village site only the 0.6 and 0.9 AOD cases can be considered dust dominated, since the $AOD=0.3$ case is mixed with Angstrom Exponent=0.84 in Table 2.

Page 13, line 29-31: The estimation of uncertainty for AERONET measured AOD for field instruments is 0.01 in the visible and NIR and increasing to 0.02 in the UV wavelengths (Eck et al. 1999). This spectral variation in AOD uncertainty should be mentioned here or elsewhere in the paper.

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Page 14, line 8: This is an awkward way to say that uncertainties increase at lower AOD. For the reader it would be much clearer if you did not use abbreviations such as GSFC1 and GSFC3 in the text but instead referred to the AOD levels and associated uncertainties.

Page 19, line 5: Should insert the Dubovik et al. (2006) reference here since these significant algorithm advances were incorporated in the AERONET operational inversions in 2006.

Page 19, line 6: Please add the reference Smirnov et al. (2000) here for Level 2 data in Version 2.

Page 19, line 22-23: Should say that these sites did not have the 1640 nm channel of the newer Cimels, as many readers will not know what you mean by 'extended' photometer.

Page 19, line 31-33: I suggest that an additional, perhaps more robust way to compare the retrievals of the fine mode fraction (FMF) of optical depth is to provide a scatterplot of all individual (not just daily means) GRASP-AOD retrievals regardless of SZA (not just time matched to almucantars) for many sites (including sites not in the Dubovik 2002 table) compared to the AERONET almucantar retrievals of FMF. This should include all AOD levels measured at each site. Both can be plotted as a function of AE(440-870) on the x-axis. Regression fit and statistics can then also be shown. Other parameters such as volume median radius could also be compared as scatterplots using time-matched individual retrievals from many sites, with the AERONET standard sky-radiance inversion on the x-axis and the GRASP-AOD inversion on the y-axis.

Page 20, lines 10-11: Need to mention the likely reason for this large difference for the GSFC-A case. It is the lowest AOD of all sites with $AOD(440)=0.166$. This is important and needs to be quantified further with cases that have lower AOD values such as 0.10 or lower.

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Page 20, line 13: This is an anomalously large r_v fine for a low AOD case at GSFC. Note that the climatology of GSFC size distributions (as a function of AOD) in Eck et al. (2012; Figure 17 b) and in Dubovik et al. (2002) strongly suggest that this case is an anomaly, therefore a relatively poor choice for a case study. The actual values of the almucantar retrieved individual r_v fine at GSFC for this date (Nov 22, 2009) range from 0.21 to 0.26 micron (still very high for this AOD level) so I don't know where you got the 0.27 micron number from.

Page 21, lines 17-20: Should note here in the text that differences in the definition of modes, radius cutoff in AERONET standard retrievals versus tails of modes included in SDA explain a least a portion of these differences.

Page 22, lines 9-10: Need to note that the moon measurements have a spectral range from 440 to 1640 nm, while earlier in the paper you made the sensitivity analysis for the 340 nm to 1640 nm spectral range (see Table 2).

Page 23, lines 20-21: Are you referring to pollution or smoke from biomass burning or a mixture of the two types of aerosols?

Page 24, lines 8-9: Please give the r_v for both the fine and coarse modes for the AERONET standard almucantar inversion, for comparison purposes to the GRASP-AOD inverted size distributions.

Page 28, lines 2-3: Calibration uncertainty for AERONET master instruments is better than you stated here. The uncertainty in V_0 due to calibration by Langley analysis at Mauna Loa is $\sim 0.25\%$ in the visible and NIR and 0.5% in the UV. The resulting total uncertainty in AOD (additionally including some other sources of uncertainty) is ~ 0.002 to 0.009 for Master Cimels (lower values in visible and NIR and higher in UV ; see Eck et al. 1999). This is for overhead sun ($SZA=0$ and optical airmass=1) and the uncertainty in measured AOD is less as solar zenith angle increases (optical airmass increases).

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Page 28, lines 8-9: Should cite Smirnov et al. (2000) here since the Version 2 Level 2 data are cloud screened and quality controlled as described in this reference.

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