



SEALDH-II – a calibration-free transfer standard for airborne water vapor measurements: Pressure dependent absolute validation from 5 – 1200 ppmv at a metrological humidity generator

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Abstract

Highly accurate water vapor measurements are indispensable for understanding a variety of scientific questions as well as industrial processes. While in metrology water vapor concentrations can be defined, generated and measured with relative uncertainties in the single percentage range, field deployable airborne instruments deviate even under quasi-static laboratory conditions up to 10-20%. The novel SEALDH-II hygrometer, a calibration-free, tuneable diode laser spectrometer, bridges this gap by implementing an entirely new concept to achieve higher accuracy levels in the field. Here we present the absolute validation of SEALDH-II at a traceable humidity generator during 23 days of permanent operation at 15 different H₂O concentration levels between 5 and 1200 ppmv. At each concentration level, we studied the pressure dependence at 6 different gas pressures between 65 and 950 hPa. Further, we describe the setup for this metrological validation, the challenges to overcome when assessing water vapor measurements on a high accuracy level, as well as the comparison results. With this validation, SEALDH-II is the first metrologically validated humidity transfer standard which links several scientific airborne and laboratory measurement campaigns to the international metrological water vapor scale.

1. Introduction

Water vapor affects, like no other substance, nearly all atmospheric processes (Ludlam, 1980; Möller et al., 2011; Ravishankara, 2012). Water vapor represents not only a large direct feedback to global warming when forming clouds, but also plays a major role in atmospheric chemistry (Held and Soden, 2000; Houghton, 2009; Kiehl and Trenberth, 1997). Changes in the water distribution, as vapor or in condensed phases (e.g. in clouds), have a large impact on the radiation balance of the atmosphere. This justifies that water vapor is often mentioned as the most important greenhouse gas and one of the most important parameters in climate research (Ludlam, 1980; Maycock et al., 2011). Water vapor is often needed for other in-situ atmospheric analyzers to correct for their water vapor cross-interference. The high (spatial and temporal) variability of atmospheric water vapor, its large dynamic range (3 – 40 000 ppmv), and its broad



39 spectroscopic fingerprint typically require complex multi-dimensional calibrations, in particular for
40 spectroscopic sensors. These calibrations often embrace the water vapor content of the gas flow to be
41 analyzed as one of the key calibration parameters even if the instrument (e.g. for CO₂), is not intended to
42 measure water vapor at all.

43 In particular for field weather stations, water vapor analyzers often are seen as part of the standard
44 instrumentation in atmospheric research. This seems reasonable due to several reasons: slow H₂O
45 concentration change over hours, the typical mid-range humidity levels (approx. above 5000 ppmv), no
46 significant gas pressure or temperature change, target accuracy on the order of 15%, and the absence of
47 “non-typical atmospheric components” such as soot or hydrophobic substances. Water vapor
48 measurements under these conditions can be performed by a variety of different devices (Wiederhold,
49 1997): Capacitive polymer sensors such as (Salasmaa and Kostamo, 1986) are frequently deployed in low
50 cost (field) applications. Standardized spectral absorption devices such as (Petersen et al., 2010) are often
51 used in research campaigns. Dew-point mirror hygrometers (DPM) are known for their high accuracy.
52 However, this is only true if they are regularly calibrated at a high accuracy (transfer-) standards in
53 specialized hygrometry laboratories such as in metrology institutes (Heinonen et al., 2012).

54 As soon as hygrometers have to be deployed in harsh environments (e.g. on airborne platforms), this
55 situation changes entirely: The ambient gas pressure (70 – 1000 hPa) and gas temperature (-80 – 40°C)
56 ranges are large and both values change rapidly, the required H₂O measurement range is set by the ambient
57 atmosphere (3 – 40000 ppmv), mechanical stress and vibrations occur, and the sampled air contains
58 additional substances from condensed water (ice, droplets), particles, or even aircraft fuel vapor (e.g. on
59 ground). These and other impacts complicate reliable, accurate, long-term stable H₂O measurements and
60 briefly outline why water vapor measurements remain a quite difficult in-situ measurement in the field,
61 even if they are nearly always needed in atmospheric science. Up to now, the lack of sufficient accuracy
62 may have limited important scientific interpretations (Krämer et al., 2009; Peter et al., 2006; Scherer et al.,
63 2008; Sherwood et al., 2014).

64 Over the last decades, numerous hygrometers were developed and deployed on aircraft (Buck, 1985; Busen
65 and Buck, 1995; Cerni, 1994; Desjardins et al., 1989; Diskin et al., 2002; Durry et al., 2008; Ebert et al., 2000;
66 Gurlit et al., 2005; Hansford et al., 2006; Helten et al., 1998; Hunsmann et al., 2008; Karpechko et al., 2014;
67 Kley and Stone, 1978; May, 1998; Meyer et al., 2015; Ohtaki and Matsui, 1982; Roths and Busen, 1996;
68 Salasmaa and Kostamo, 1986; Schiff et al., 1994; Silver and Hovde, 1994a, 1994b; Thornberry et al., 2014;
69 Webster et al., 2004; Zöger et al., 1999a, 1999b) (non-exhaustive list), but those often show results which are
70 not sufficient for validation or falsification of atmospheric models in terms of the required absolute
71 accuracy, precision, temporal resolution, long-term stability, comparability, etc. These problems can be
72 grouped into two major categories: accuracy linked problems and time response linked problems. The latter
73 one is in particular important for investigations in strongly, spatially structured regions in the lower
74 troposphere as well as for investigations in clouds. In these regions, even two on average agreeing
75 instruments with different response times yield local, large, relative deviations on the order of up to 30%
76 (Smit et al., 2014). In contrast to time response studies, accuracy linked problems in flight are difficult to



77 isolate since they are always covered by the spatial variability (which leads to temporal variability for
78 moving aircraft) of atmospheric H₂O distribution. Comparing hygrometer in flight, such as, for example in
79 (Rollins et al., 2014), does not facilitate a clear accuracy assessment.

80 Therefore in 2007, an international intercomparison exercise named “AquaVIT” (Fahey et al., 2014) was
81 carried out to compare airborne hygrometers under quasi-static, laboratory-like conditions for upper
82 tropospheric and stratospheric humidity levels. AquaVIT (Fahey et al., 2014) encompassed 22 instruments
83 from 17 international research groups. The instruments were categorized in well-validated, often deployed
84 “core” instruments (APiC, FISH, FLASH, HWV, JLH, CFH) and “younger” non-core instruments.

85 AquaVIT revealed in the important 1 to 150 ppmv H₂O range, that -even under quasi-static conditions- the
86 deviation between the core instrument’s readings and their averaged group mean was on the order of ±10
87 %. This result fits to the typical interpretation problems of flight data where instruments often deviate from
88 each other by up to 10%, which is not covered by the respective uncertainties of the individual instruments.

89 AquaVIT was a unique first step to document and improve the accuracy of airborne measurements in order
90 to make them more comparable. However, no instrument could claim after AquaVIT that its accuracy is
91 higher than any other AquaVIT instrument, since no “gold standard” was part of the campaign, i.e., a
92 metrological transfer standard (JCGM 2008, 2008; Joint Committee for Guides in Metrology (JCGM), 2009)
93 traced back to the SI units. There is no physical argument for the average being better than the measured
94 value of a single instrument. Instead, many arguments speak for systematic deviations of airborne
95 hygrometers: Most hygrometers have to be calibrated. Even for a perfect instrument, the accuracy issue is
96 entirely transferred to the calibration source and its gas handling system, which in this case leads to two
97 major concerns: First, one has to guarantee that the calibration source is accurate and stable under field
98 conditions, i.e., when using it before or after a flight on the ground. This can be challenging especially for
99 the transportation of the source with all its sensitive electronics/mechanics and for the deviating ambient
100 operation temperature from the ambient validation temperature (hangar vs. laboratory). Even more prone
101 to deviations are calibration sources installed inside of an aircraft due to changing ambient conditions such
102 as cabin temperature, cabin pressure, orientation angle of instrument (important, if liquids are used for
103 heating or cooling). Secondly, the gas stream with a highly defined amount of water vapor has to be
104 conveyed into the instrument. Especially for water vapor, which is a strongly polar molecule, this gas
105 transport can become a critical step. Changing from high to low concentrations or even just changing the
106 gas pressure or pipe temperature can lead to signal creep due to slow adsorption and desorption processes,
107 which can take long to equilibrate. In metrology, this issue is solved by a long validation/calibration time
108 (hours up to weeks, depending on the H₂O concentration level), a generator without any connectors/fittings
109 (everything is welded) and piping made out of electro-polished, stainless steel to ensure that the
110 equilibrium is established before the actual calibration process is started. However, this calibration
111 approach is difficult to deploy and maintain for aircraft/field operations due to the strong atmospheric
112 variations in gas pressure and H₂O concentrations, which usually leads to a multi-dimensional calibration
113 pattern (H₂O concentration, gas pressure, sometimes also gas temperature) in a short amount of calibration
114 time (hours). Highly sensitive, frequently flown hygrometers like (Zöger et al., 1999a) are by their physical



115 principle, not as long-term stable as it would be necessary to take advantage of a long calibration session.
116 Besides the time issue to reach a H₂O equilibrium between source and instrument, most calibration
117 principles for water vapor are influenced by further issues. A prominent example is the saturation of air in
118 dilution/saturation based water vapor generators: gas temperature and pressure defines the saturation level
119 (described e.g. by Sonntag's Equation (Rollins et al., 2014)), however, it is well-known that e.g. 100.0%
120 saturation is not easily achievable. This might be one of the impact factors for a systematic offset during
121 calibrations in the field. The metrology community solves this for high humidity levels with large, multi-
122 step saturation chambers which decrease the temperature step-wise to force the water vapor to condense in
123 every following step. These few examples of typical field-related problems show, that there is a reasonable
124 doubt that deviations in field situations are norm-distributed. Hence, the mean during AquaVIT might be
125 biased, i.e. not the correct H₂O value.
126 The instruments by themselves might actually be more accurate than AquaVIT showed, but deficiencies of
127 the different calibration procedures (with their different calibration sources etc.) might mask this. To
128 summarize, AquaVIT documented a span of up to 20% relative deviation between the world's best airborne
129 hygrometers – but AquaVIT could not assess absolute deviations nor explain them, since a link to a
130 metrological H₂O primary standard (i.e., the definition of the international water vapor scale) was missing.
131 Therefore, we present in this paper the first comparison of an airborne hygrometer (SEALDH-II) with a
132 metrological standard for the atmospheric relevant gas pressure (65 – 950 hPa) and H₂O concentration
133 range (5 – 1200 ppmv). We will discuss the validation setup, procedure, and results. Based on this
134 validation, SEALDH-II is by definition the first airborne transfer standard for water vapor.
135

136 2. SEALDH-II

137 2.1. System description

138 This paper focuses on the metrological accuracy validation of the Selective Extractive Airborne Laser Diode
139 Hygrometer (SEALDH-II). SEALDH-II is the airborne successor of the proof-of-concept spectrometer study
140 published in (Buchholz et al., 2014), which showed the possibility and the achievable accuracy level for
141 calibration-free dTDLAS hygrometry. The publication (Buchholz et al., 2014) demonstrates this for the
142 600 ppmv to 20000 ppmv range at standard ambient pressure). SEALDH-II integrates numerous different
143 principles, concepts, modules, and novel parts, which contribute to or enable the results shown in this
144 paper. SEALDH-II's high internal complexity does not allow a full, detailed discussion of the entire
145 instrument in this paper; for more details the reader is referred to (Buchholz et al., 2016). The following
146 brief description covers the most important technical aspects of the instrument from a user's point of view:
147
148 SEALDH-II is a compact (19" rack 4 U (=17.8 cm)) closed-path, absolute, directly Tunable Diode Laser
149 Absorption Spectroscopy (dTDLAS) hygrometer operating at 1.37 μ m. With its compact dimensions and the



moderate weight (24 kg), it is well suited for space- and weight-limited airborne applications. The internal optical measurement cell is a miniaturized White-type cell with an optical path length of 1.5 m. It is connected to the airplane's gas inlet via an internal gas handling system comprising a temperature exchanger, multiple temperature sensors, a flow regulator, and two gas pressure sensors. Approximately 80 different instrument parameters are controlled, measured, or corrected by SEALDH-II at any time to provide a holistic view on the spectrometer status. This extensive set of monitoring data ensures reliable and well-characterized measurement data at any time. The knowledge about the instruments status strongly facilitates metrological uncertainties calculations. SEALDH-II's calculated linear measurement uncertainty is 4.3%, with an additional offset uncertainty of ± 3 ppmv (further details in (Buchholz et al., 2016)). The precision of SEALDH-II was determined via the Allan-variance approach and yielded 0.19 ppmv ($0.17 \text{ ppmv} \cdot \text{m} \cdot \text{Hz}^{-1/2}$) at 7 Hz repetition rate and an ideal precision of 0.056 ppmv ($0.125 \text{ ppmv} \cdot \text{m} \cdot \text{Hz}^{-1/2}$) at 0.4 Hz. In general, SEALDH-II's time response is limited by the gas flow through the White-type multi-pass measurement cell with a volume of 300 ccm. With the assumption of a bulk flow of 7 SLM at 200 hPa through the cell, the gas exchange time is 0.5 seconds. SEALDH-II's measurement range covers 3 – 40000 ppmv. The calculated mixture fraction offset uncertainty of ± 3 ppmv defines the lower detection limit. This offset uncertainty by itself is entirely driven by the capability of detecting and minimizing parasitic water vapor absorption. The concept, working principle, and its limits are described in (Buchholz and Ebert, 2014). The upper limit of 40000 ppmv is defined by the lowest internal instrument temperature, which has to always be higher than the dew point temperature to avoid any internal condensation. From a spectroscopic perspective, SEALDH-II could handle concentrations up to approx. 100000 ppmv before spectroscopic problems like saturation limit the accuracy and increase the relative uncertainty beyond 4.3%.

2.1. Calibration-free evaluation approach

SEALDH-II's data treatment works differently from nearly all other published TDLAS spectrometers. Typically, instruments are setup in a way that they measure the absorbance or a derivative measurand of absorbance, and link it to the H_2O concentration. This correlation together with a few assumptions about long-term stability, cross interference, gas temperature dependence, gas pressure dependence is enough to calibrate a system (Muecke et al., 1994). Contrarily, a calibration-free approach requires a fully featured physical model describing the absorption process entirely. The following description is a brief overview; for more details see e.g. (Buchholz et al., 2014, 2016; Ebert and Wolfrum, 1994; Schulz et al., 2007). In a very simplified way, our physical absorption model uses the *extended* Lambert-Beer equation (Equation 1) which describes the relationship between the initial light intensity $I_0(\lambda)$ before the absorption path (typically being in the few mW-range) and the transmitted light intensity $I(\lambda)$.

$$\text{Equation 1: } I(\lambda) = E(t) + I_0(\lambda) \cdot \text{Tr}(t) \cdot \exp[-S(T) \cdot g(\lambda - \lambda_0) \cdot N \cdot L]$$

The parameter $S(T)$ describes the line strength of the selected molecular transition. In SEALDH-II's case, the spectroscopic multi-line fit takes into account 19 transition lines in the vicinity of the target line at 1370 nm



(energy levels: 110 – 211, rotation-vibrational combination band). The other parameters are the line shape function $g(\lambda - \lambda_0)$, the absorber number density N , the optical path length L and corrections for light-type background radiation $E(t)$ and broadband transmission losses $Tr(t)$.

Equation 1 can be enhanced with the ideal gas law to calculate the H_2O volume mixing ratio c :

Equation 2:
$$c = - \frac{k_B \cdot T}{S(T) \cdot L \cdot p} \int \ln \left(\frac{I(v) - E(t)}{I_0(v) \cdot Tr(t)} \right) \frac{dv}{dt} dt$$

The additional variables in Equation 2 are: constant entities like the Boltzmann constant k_B ; the optical path length L ; molecular constants like the line strength $S(T)$ of the selected molecular transition; the dynamic laser tuning coefficient $\frac{dv}{dt}$, which is a constant laser property; continuously measured entities such as gas pressure (p), gas temperature (T) and photo detector signal of the transmitted light intensity $I(v)$ as well as the initial light intensity $I_0(v)$, which is retrieved during the evaluation process from the transmitted light intensity $I(v)$.

Equation 2 facilitates an evaluation of the measured spectra without any instrument calibration at any kind of water vapor reference (Buchholz et al., 2014; Ebert and Wolfrum, 1994; Schulz et al., 2007) purely based on first principles. Our concept of a fully calibration-free data evaluation approach (this excludes also any referencing of the instrument to a water standard in order to correct for instrument drift, offsets, temperature dependence, pressure dependence, etc.) is crucial for the assessment of the results described in this publication. It should be noted that the term “calibration-free” is frequently used in different communities with dissimilar meanings. We understand this term according to the following quote (JCGM 2008, 2008): “calibration (...) in a first step, establishes a relation between the measured values of a quantity with measurement uncertainties provided by a measurement standard (...), in a second step, [calibration] uses this information to establish a relation for obtaining a measurement result from an indication (of the device to be calibrated)”. Calibration-free in this sense means, that SEALDH-II does not use any information from “calibration-, comparison-, test-, adjustment-” runs with respect to a higher accuracy “water vapor standard” to correct or improve any response function of the instrument. SEALDH-II uses as described in (Buchholz et al., 2016) only spectroscopic parameters and the 80 supplementary parameters as measurement input to calculate the final H_2O concentration. The fundamental difference between a calibration approach and this stringent concept is that only effects which are part of our physical model are taken into account for the final H_2O concentration calculation. All other effects like gas pressure or temperature dependencies, which cannot be corrected with a well-defined physical explanation, remain in our final results even if this has the consequence of slightly uncorrected results deviations. This strict philosophy leads to measurements which are very reliable with respect to accuracy, precision and the instrument’s over-all performance. The down-side is a relatively computer-intensive, sophisticated evaluation. As SEALDH-II stores all the raw spectra, one could – if needed for whatever reason – also calibrate the instrument by referencing it to a high accuracy water vapor standard and transfer the better accuracy e.g. of a metrological standard onto the instrument. Every calibration-free instrument can be calibrated since pre-requirements for a calibration are just a subset of the requirements for a calibration-free



instrument. However, a calibration can only improve the accuracy for the relatively short time between two calibration-cycles by adding all uncertainty contributions linked to the calibration itself to the system. This is unpleasant or even intolerable for certain applications and backs our decision to develop a calibration-free instrument to enable a first principles, long-term stable, maintenance-free and autonomous hygrometer for field use e.g. at remote sites or aircraft deployments.

3. SEALDH-II validation facility

3.1. Setup

Figure 1 right shows the validation setup. As a well-defined and highly stable H₂O vapor source, we use a commercial Thunder scientific model (TSM) 3900, similar to (Thunder-Scientific, 2016). This source saturates pre-dried air at an elevated gas pressure in an internally ice covered chamber. The gas pressure in the chamber and the chamber's wall temperature are precisely controlled and highly stable and thus define the absolute water vapor concentration via the Sonntag equation (Sonntag, 1990). After passing through the saturator, the gas expands to a pressure suitable for the subsequent hygrometer. The pressure difference between the saturation chamber pressure and the subsequent step give this principle its name "two pressure generator". The stable H₂O concentration range of the TSM is 1 – 1300 ppmv for these specific deployment conditions. This generator provides a stable flow of approximately 4 – 5 SLM. Roughly 0.5 SLM are distributed to a frost/dew point hygrometer, D/FPH, (MBW 373) (MBW Calibration Ltd., 2010). SEALDH-II is fed with approx. 3.5 SLM, while 0.5 SLM are fed to an outlet. This setup ensures that the dew point mirror hygrometer (DPH)¹ operates close to the ambient pressure, where its metrological primary calibration is valid, and that the gas flow is sufficiently high in any part of the system to avoid recirculation of air. The vacuum pump is used to vary the gas pressure in SEALDH-II's cell with a minimized feedback on the flow through the D/FPH and the TSM. This significantly reduces the time for achieving a stable equilibrium after any gas pressure change in SEALDH-II's chamber. SEALDH-II's internal electronic flow regulator limits the mass flow at higher gas pressures and gradually opens towards lower pressures (vacuum pumps usually convey a constant volume flow i.e., the mass flow is pressure dependent). We termed this entire setup "traceable humidity generator", THG, and will name it as such throughout the text.

3.1. Accuracy of THG

The humidity of the gas flow is set by the TSM generator but the absolute H₂O values are traceably determined with the dew point mirror hygrometer (D/FPH). The D/FPH, with its primary calibration, thus guarantees the absolute accuracy in this setup. The D/FPH is not affected by the pressure changes in SEALDH-II's measurement cell and operates at standard ambient gas pressure and gas temperature where

¹ The used dew point mirror hygrometer can measure far below 0°C; therefore, it is a dew point mirror above > 0°C and a frost point mirror as soon as there is ice on the mirror surface. We will use both DPH and D/FPH abbreviations interchangeably.



its calibration is most accurate. The D/FPH was calibrated (Figure 2) at the German national standard for mid-range humidity (green, 600 – 8000 ppmv) as well as at the German national standard for low-range humidity (blue, for lower values 0.1 – 500 ppmv). The two national standards work on different principles: The two pressure principle (Buchholz et al., 2014) currently supplies the lower uncertainties (green, “±”-values in Figure 2). Uncertainties are somewhat higher for the coulometric generator (Mackrodt, 2012) in the lower humidity range (blue). The “Δ”-values in Figure 2 show the deviations between the readings of the D/FPH and the “true” values of the national primary standards.

4. SEALDH-II validation procedure

4.1. Mid-term multi-week permanent operation of SEALDH-II

One part of the validation was a permanent operation of SEALDH-II over a time scale much longer than the usual air or ground based scientific campaigns. In this paper, we present data from a permanent 23 day long (550 operation hours) operation in automatic mode. Despite a very rigorous and extensive monitoring of SEALDH-II’s internal status, no malfunctions of SEALDH-II could be detected. One reason for this are the extensive internal control and error handling mechanisms introduced in SEALDH-II, which are mentioned above and described elsewhere (Buchholz et al., 2016). Figure 3 shows an overview of the entire validation. The multi-week validation exercise comprises 15 different H₂O concentration levels between 2 and 1200 ppmv. At each concentration level, the gas pressure was varied in six steps (from 65 to 950 hPa) over a range which is particularly interesting for instruments on airborne platforms operating from troposphere to lower stratosphere. Figure 3 (top) shows the comparison between SEALDH-II (black line) and the THG setup (red). Figure 3 (bottom) shows the gas pressure (blue) and the gas temperature (green) in SEALDH-II measurement cell. The gas temperature increase in the second week was caused by a failure of the laboratory air conditioner that led to a higher room temperature and thus higher instrument temperature. Figure 4 shows the 200 hPa section of the validation in Figure 3. To avoid any dynamic effects from time lags, hysteresis of the gas setup, or the instruments themselves, every measurement at a given concentration/pressure combination lasted at least 60 min. The data from the THG (red) show that there is nearly no feedback of a gas pressure change in SEALDH-II’s measurement cell towards the D/FPH, respectively the entire THG. The bottom subplot in Figure 4 shows the relative deviation between the THG and SEALDH-II. This deviation is correlated to the absolute gas pressure level and can be explained by deficiencies of the Voigt lines shape used to fit SEALDH-II’s spectra (Buchholz et al., 2014)(Buchholz et al., 2016). The Voigt profile, a convolution of Gaussian (for temperature broadening) and Lorentzian (pressure broadening) profiles used for SEALDH-II’s evaluation, does not include effects such as Dicke Narrowing, which become significant at lower gas pressures. Neglecting these effects cause systematic, but long-term stable and fully predictable deviations from the reference value in the range from sub percent at atmospheric gas pressures to less than 5 % at the lowest gas pressures described here. We have chosen not to implement any higher order line shape (HOLS) models as the spectral reference data needed are not



288 available at sufficient accuracy. Further, HOLS would force us to increase the number of free fitting
289 parameters, which would destabilize our fitting procedure, and lead to reduced accuracy/reliability (i.e.,
290 higher uncertainty) as well as significantly increased computational efforts. This is especially important for
291 flight operation where temporal H₂O fluctuations (spatial fluctuations result in temporal fluctuations for a
292 moving device) occur with gradients up to 1000 ppmv/s.
293 These well understood, systematic pressure dependent deviations will be visible in each further result plot
294 of this paper. The impact and methods of compensation are already discussed in (Buchholz et al., 2014). The
295 interested reader is referred to this publication for a more detailed analysis and description.
296 SEALDH-II's primary target areas of operations are harsh field environments. Stability and predictability is
297 to be balanced with potential, extra levels of accuracy which might not be required or reliably achievable
298 for the intended application. Higher order line shape models are therefore deliberately traded for a stable,
299 reliable, and unified fitting process under all atmospheric conditions. This approach leads to systematic,
300 predictable deviations in the typical airborne accessible atmospheric gas pressure range (125 – 900 hPa) of
301 less than 3%. One has to compare these results for assessment to the non-systematic deviations of 20%
302 revealed during the mentioned AquaVIT comparison campaign (Fahey et al., 2014). Hence, for
303 field/airborne purposes, the 3% seems to be fully acceptable – especially in highly H₂O structured
304 environments.
305

306 4.1. Assessment of SEALDH-II's mid-term accuracy: Dynamic effects

307 Besides the pressure dependence discussed above, SEALDH-II's accuracy assessment is exacerbated by the
308 differences in the temporal behavior between the THG's dew/frost point mirror hygrometer (D/FPH) and
309 SEALDH-II: Figure 5 (left) shows an enlarged 45 min. long section of measured comparison data. SEALDH-
310 II (black) shows a fairly large water vapor variation compared to the THG (red). The precision of SEALDH-
311 II (see chapter 2) is 0.056 ppmv at 0.4 Hz (which was validated at a H₂O concentration of 600 ppmv
312 (Buchholz et al., 2016)) yielding a signal to noise ratio of 10700. Therefore, SEALDH-II can very precisely
313 detect variations in the H₂O concentration. Contrarily, the working principle of a D/FPH requires an
314 equilibrated ice/dew layer on the mirror. As an indirect, inertia, thermal adjustment process, the response
315 time of a dew/frost point mirror hygrometer has certain limitations due to this principle (the dew/frost
316 point temperature measurement is eventually used to calculate the final H₂O concentration), whereas the
317 optical measurement principle of SEALDH-II is only limited by the gas transport, i.e., the flow (exchange
318 rate) through the measurement cell. The effect of those different response times is clearly visible from 06:00
319 to 06:08 o'clock in Figure 5. The gas pressure of SEALDH-II's measurement cell (blue), which is correlated
320 to the gas pressure in the THG's ice chamber, shows an increase of 7 hPa – caused by the regulation cycle of
321 the THG's generator (internal saturation chamber gas pressure change). The response in the THG frost
322 point measurement (green, red) shows a significant time delay compared to SEALDH-II, which detects
323 changes approx. 20 seconds faster. This signal delay is also clearly visible between 06:32 to 06:40 o'clock,
324 where the water vapor variations detected by SEALDH-II are also visible in the smoothed signals of the



THG. Figure 5 right shows such a variation in detail (5 min). The delay between the THG and SEALDH-II is here also approximately 20 seconds. If we assume that SEALDH-II measures (due to its high precision) the true water vapor fluctuations, the relative deviation can be interpreted as overshooting and undershooting of the D/FPH's controlling cycle, which is a commonly known response behavior of slow regulation feedback loops to fast input signal changes. The different time responses lead to "artificial" noise in the concentration differences between SEALDH-II and THG. Theoretically, one could characterize this behavior and then try to correct/shift the data to minimize this artificial noise. However, a D/FPH is fundamentally insufficient for a dynamic characterization of a fast response hygrometer such as SEALDH-II. Thus, the better strategy is to keep the entire system as stable as possible and calculate mean values by using the inherent assumption that under- and overshoots of the DPM affect the mean statically and equally. With this assumption, the artificial noise can be seen in the first order as Gaussian distributed noise within each pressure step (Figure 4) of at least 60 min. The error induced by this should be far smaller than the above discussed uncertainties of the THG (and SEALDH-II).

5. Results

The results of this validation exercise are categorized in three sections according to the following conditions in atmospheric regions: mid-tropospheric range: 1200 – 600 ppmv (Figure 6), upper tropospheric range: 600 – 20 ppmv (Figure 7), and lower stratospheric range: 20 – 5 ppmv (Figure 8). This categorization is also justified by the relative influence of SEALDH-II's calculated offset uncertainty of ± 3 ppmv (Buchholz and Ebert, 2014): At 1200 ppmv, its relative contribution of 0.25% is negligible compared to the 4.3% linear uncertainty of SEALDH-II. At 5 ppmv, the relative contribution of the offset uncertainty is 60% and thus dominates the linear uncertainty. Before assessing the following data, it should be emphasized again that SEALDH-II's spectroscopic first-principles evaluation was designed to rely on accurate spectral data instead of a calibration. SEALDH-II was never calibrated or referenced to any kind of reference humidity generator or sensor.

5.1. The 1200 – 600 ppmv range

Figure 6 shows the summary of the pressure dependent validations in the 1200 – 600 ppmv range. Each of the 48 data points represents the mean over one pressure measurement section of at least 60 min (see Figure 4). A cubic polynomial curve fitted to the 600 ppmv results (blue) serves as an internal quasi-reference to connect with the following graphs. The 600 ppmv data (grey) are generated via a supplementary comparison at a different generator: The German national mid-humidity primary generator (PHG). This primary generator data at 600 ppmv indicate a deviation between PHG and THG of about 0.35 %, which is compatible with the uncertainties of the THG (see chapter 3.1) and the PHG (0.4%) (Buchholz et al., 2014). The PHG comparison data also allow a consistency check between the absolute values of (see Figure 2) the PHG (primary standard = calibration-free), the THG (DPM calibrated) and SEALDH-II (calibration-free).



5.2. The 600 – 20 ppmv range

In this range, the linear uncertainty (4.3%) and the offset uncertainty (± 3 ppmv) have both a significant contribution. Figure 7 shows a clear trend: The lower the concentration, the higher the deviation. We believe this is being caused by SEALDH-II's offset variation and will be discussed in the 20 – 5 ppmv range.

5.3. The 20 – 5 ppmv range

The results in this range (Figure 8) are dominated by the offset uncertainty. It is important to mention at this point, that the ± 3 ppmv uncertainties are calculated based on assumptions, design innovations, and several independent, synchronous measurements which are automatically done while the instrument is in operation mode (see publication (Buchholz et al., 2016; Buchholz and Ebert, 2014)). Hence, the calculated uncertainties resemble an upper uncertainty threshold; the real deviation could be lower than 3 ppmv. A clear assessment is fairly difficult since at low concentrations (i.e., low optical densities) several other effects occur together such as, e.g., optical interference effects like fringes caused by the very long coherence length of the used laser. However, Figure 9 (left) allows a rough assessment of the offset instability. This plot shows all the data below 200 ppmv, grouped by the gas pressure in the measurement cell. If one ignores the 65 hPa and 125 hPa measurements, which are clearly affected by higher order line shape effects (see above), the other measurements fit fairly well in a ± 1 ppmv envelope function (grey). In other words, SEALDH-II's combined offset "fluctuations" are below 1 ppmv H₂O. All validation measurements done with SEALDH-II during the last years consistently demonstrated a small offset variability so that the observed offset error is around 0.6 ppmv – i.e., only 20% of the calculated ± 3 ppmv.

5.4. General evaluation

Figure 9 presents a summary of all 90 analyzed concentration/pressure-pairs during the 23 days of validation. The calculated uncertainties (linear 4.3% and offset ± 3 ppmv) of SEALDH-II are plotted in purple. This uncertainty calculation doesn't include line shape deficiencies and is therefore only valid for a pressure range where the Voigt profile can be used to represent all major broadening effects of absorption lines (Dicke, 1953; Maddaloni et al., 2010). This is the case above 250 hPa. The results at 950, 750, 500, 250 hPa show that the maximum deviations, derived from these measurements, can be described with one single performance statement: linear +2.5%, offset -0.6 ppmv.

To prevent further interpretations, it should be noted that this result doesn't change the statement about SEALDH-II's uncertainties, since these are calculated and not based on any validation/calibration process. This is a significantly different approach: The holistic control/overview is one of the most important and essential differences between calibration-free instruments such as SEALDH-II and other classical spectroscopic instruments which rely on sensor calibration. SEALDH-II can guarantee correctness of measurement values within its uncertainties because any effect which causes deviations has to be included in the evaluation model – otherwise it is not possible to correct for it.

As mentioned before, any calibration-free instrument can be calibrated too (see e.g. (Buchholz et al., 2013)).



395 However by doing so, one must accept to a certain extent loss of control over the system, especially in
396 environments which are different from the calibration environment. For example, if a calibration was used
397 to remove an instrumental offset, one has to ensure that this offset is long-term stable, which is usually
398 quite difficult, as - shown by the example of parasitic water offsets in fiber coupled diode laser hygrometers
399 (Buchholz and Ebert, 2014). Another option is to choose the recalibration frequency high enough; i.e.,
400 minimizing the drift amplitude by minimizing the time between two calibrations. This, however, reduces
401 the usable measurement time and leads to considerable investment of time and money into the calibration
402 process. For the case of SEALDH-II, a calibration of the pressure dependence – of course tempting and easy
403 to do – would directly “improve” SEALDH-II’s laboratory overall performance level from $\pm 4.3\% \pm 3$ ppmv
404 to $\pm 0.35\% \pm 0.3$ ppmv. At first glance, this “accuracy” would then be an improvement by a factor of 55
405 compared to the mentioned results of AquaVIT (Fahey et al., 2014). However, it is extremely difficult – if
406 not impossible – to guarantee this performance and the validity of the calibration under harsh field
407 conditions; instead SEALDH-II would “suffer” from the same typical calibration associated problems in
408 stability and in predictability. Eventually, the calibration-free evaluation would define the trusted values
409 and the “improvement”, achieved by the calibration, would have to be used very carefully and might
410 disappear eventually.

411 6. Conclusion and Outlook

412 The SEALDH-II instrument; a novel, compact, airborne, calibration-free hygrometer which implements a
413 holistic, first-principles directly tuneable diode laser absorption spectroscopy (dTDLAS) approach was
414 stringently validated at a traceable water vapor generator at the German national metrology institute (PTB).
415 The pressure dependent validation covered a H₂O range from 5 to 1200 ppmv and a pressure range from
416 65 hPa to 950 hPa. In total, 90 different H₂O concentration/pressure levels were studied within 23 days of
417 permanent validation experiments. Compared to other comparisons of airborne hygrometers - such as those
418 studied in the non-metrological AquaVIT campaign (Fahey et al., 2014), where a selection of the best “core”
419 instruments still showed an accuracy scatter of at least $\pm 10\%$ without an absolute reference value - our
420 validation exercise used a traceable reference value derived from instruments directly linked to the
421 international water scale. This allowed a direct assessment of SEALDH-II’s absolute performance with a
422 relative accuracy level in the sub percent range. Under these conditions, SEALDH-II showed an excellent
423 absolute agreement within its uncertainties which are 4.3% of the measured value plus an offset of ± 3 ppmv
424 (valid at 1013 hPa). SEALDH-II showed at lower gas pressures - as expected - a stable, systematic, pressure
425 dependent offset to the traceable reference, which is caused by the line shape deficiencies of the Voigt line
426 shape: e.g. at 950 hPa, the systematic deviation of the calibration-free evaluated results could be described
427 by (linear +0.9%, offset -0.5 ppmv), while at 250 hPa the systematic deviations could be described by (linear
428 +2.5%, offset -0.6 ppmv). If we suppress this systematic pressure dependence, the purely statistical
429 deviation is described by linear scatter of $\pm 0.35\%$ and an offset uncertainty of ± 0.3 ppmv.



430 Due to its extensive internal monitoring and correction infrastructure, SEALDH-II is very resilient against a
431 broad range of external disturbances and has an output signal temperature coefficient of only 0.026%/K,
432 which has already been validated earlier (Buchholz et al., 2016). Therefore, these results can be directly
433 transferred into harsh field environments. With this metrological, mid and upper atmosphere focused
434 validation presented here, we believe SEALDH-II to be the first directly deployable, metrologically
435 validated, airborne transfer standard for atmospheric water vapor. Having already been deployed in
436 several airborne and laboratory measurement campaigns, SEALDH-II thus directly links for the first time,
437 scientific campaign results to the international metrological water vapor scale.

438 **Data availability**

439 *The underlying data for the results shown in this paper are raw spectra (time vs. photo current), which are compressed*
440 *to be compatible with the instruments data storage. In the compressed state the total amount is approximately 6GB of*
441 *binary data. Uncompressed data size is approx. 60 GB. We are happy to share these data on request.*

442

443 **Author Contributions**

444 *Bernhard Buchholz and Volker Ebert conceived and designed the experiments. Bernhard Buchholz performed the*
445 *experiments; Bernhard Buchholz and Volker Ebert analyzed the data and wrote the paper.*

446

447 **Conflicts of Interest**

448 *The authors declare no conflict of interest*

449

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456 *measurements.*

457



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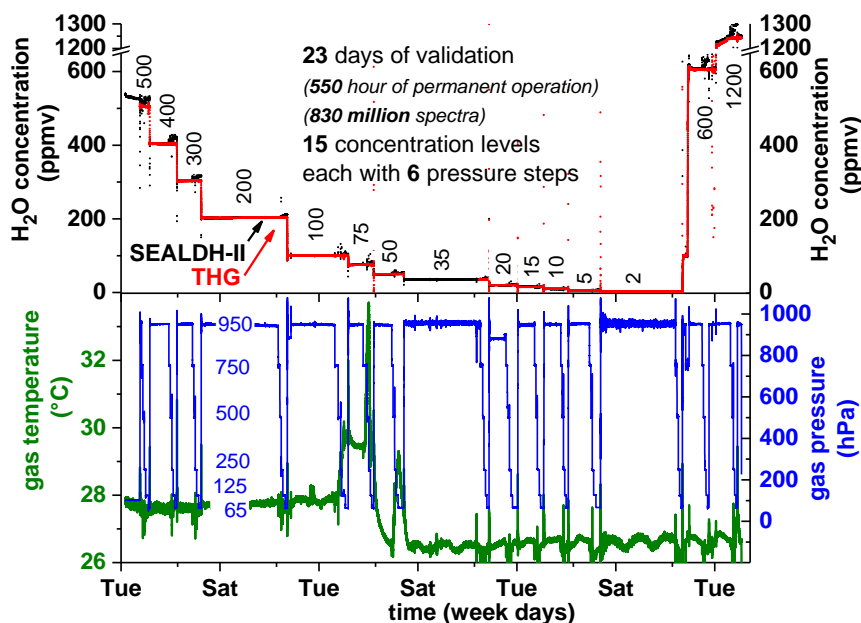


Figure 3: Overview showing all data recorded over 23 days of validation experiments. Measurements of the traceable humidity generator (THG) are shown in red, SEALDH-II data in black, gas pressure and gas temperature in SEALDH-II's measurement cell are shown in blue and green. Note: SEALDH-II operated the entire time without any malfunctions; the THG didn't save data in the 35 ppmv section; the temperature increase during the 75 ppmv section was caused by a defect of the air conditioning in the laboratory.

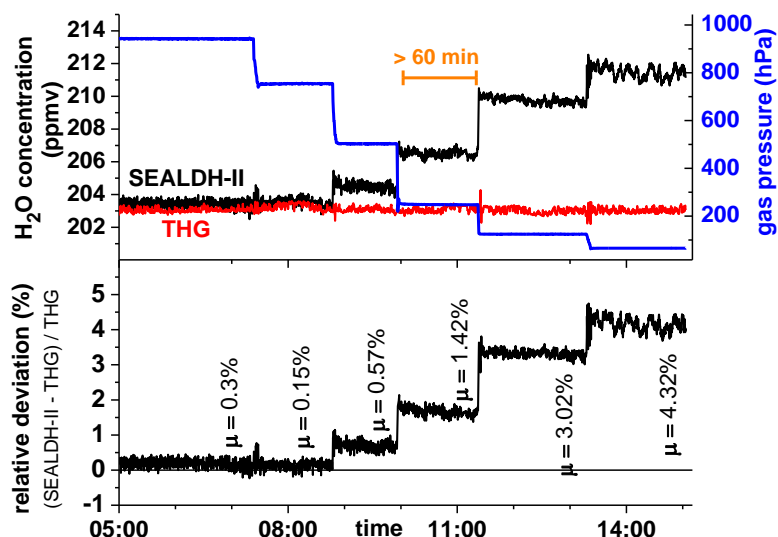


Figure 4: Detailed plot of the validation at 200 ppmv with six gas pressure steps from 50 to 950 hPa. Each individual pressure level was maintained for at least 60 minutes in order to avoid any dynamic or hysteresis effects and to facilitate clear accuracy assessments.

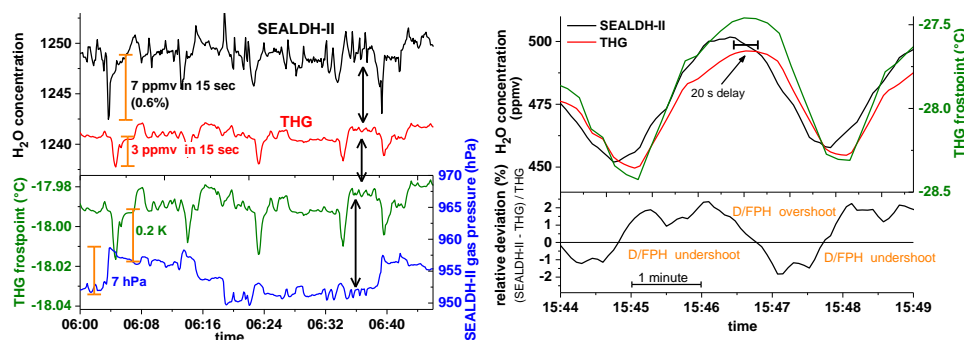


Figure 5: Short term H_2O fluctuations in the generated water vapor flow measured by SEALDH-II and the dew/frost point mirror hygrometer (D/FPH) of the traceable humidity generator (THG). The different dynamic characteristics of SEALDH-II (fast response time) and THG (quite slow response) lead in a direct comparison to artificial noise.

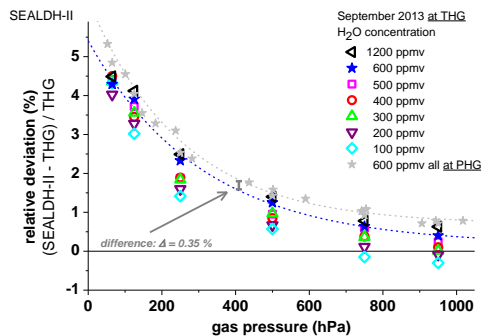


Figure 6: Gas pressure dependent comparison between SEALDH-II and THG over a H₂O concentration range from 600 to 1200 ppmv and a pressure range from 50 to 950 hPa. The 600 ppmv values (in grey) are measured directly at the national primary humidity generator (PHG) of Germany; all other H₂O concentration values are measured at and compared to the traceable humidity generator (THG). All SEALDH-II spectra were evaluated with a calibration-free first principles evaluation based on absolute spectral parameters. No initial or repetitive calibration of SEALDH-II with respect to any “water reference” source was used.

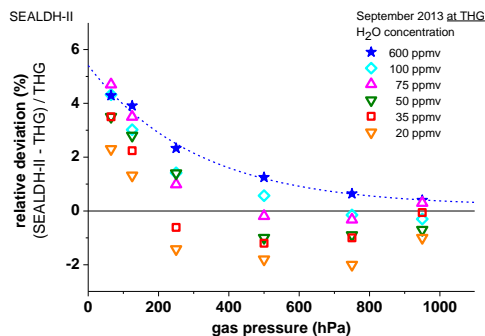


Figure 7: Comparison results as in Figure 6 but for the 200 – 600 ppmv range.

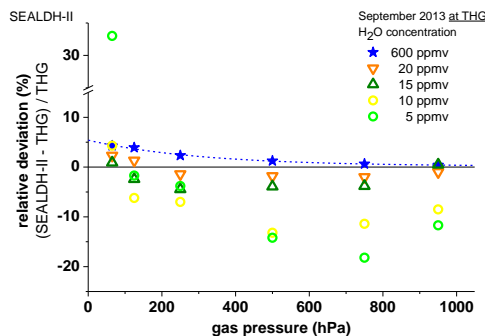




Figure 8: Comparison results as in Figure 6 and Figure 7 but for the 5 – 20 ppmv range. All spectra are determined with a calibration-free first principles evaluation concept. The major contribution to the higher fluctuations at lower concentrations is the accuracy of the offset determination (details see text).

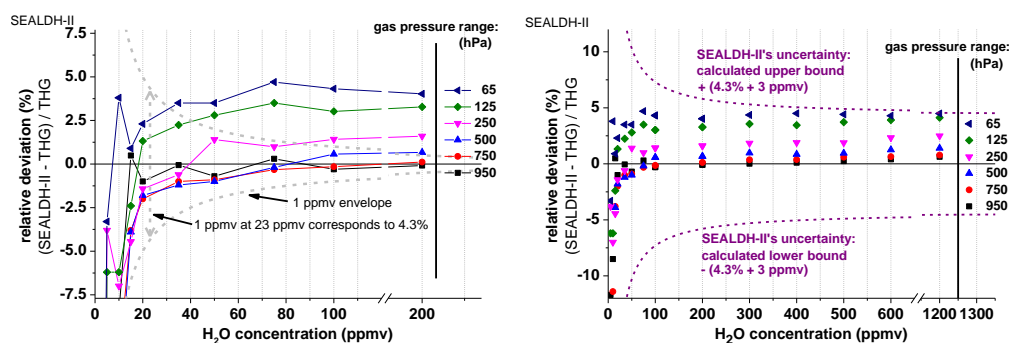


Figure 9: Direct comparison of SEALDH-II versus THG for H_2O concentrations between 5 and 200 ppmv and gas pressures from 65 to 950 hPa. Both figures show the relative deviations between SEALDH-II and THG grouped and color-coded by gas-pressure. Left plot: relative deviations of SEALDH-II versus THG below 200 ppmv; the grey line indicates the computed relative effect in SEALDH-II's performance caused by ± 1 ppmv offset fluctuation. This line facilitates a visual comparison between an offset impact and the 4.3% linear uncertainty of SEALDH-II. Right plot: relative deviations for all measured data in the same concentration range. Also shown is SEALDH-II's total uncertainty of $4.3\% \pm 3$ ppmv (calculated for 1013 hPa) as a dashed line.