1 Tropospheric products of the 2nd European GNSS reprocessing (1996-2014)

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8 Abstract

9 In this paper, we present results of the 2nd reprocessing of all data from 1996 to 2014 from all stations 10 in the European GNSS permanent network as performed at the Geodetic Observatory Pecný (GOP). 11 While the original goal of this research was to ultimately contribute to new realization of the European 12 terrestrial reference system, we also aim to provide a new set of GNSS tropospheric parameter time 13 series with possible applications to climate research. To achieve these goals, we improved a strategy 14 to guarantee the continuity of these tropospheric parameters and we prepared several variants of 15 troposphere modelling. We then assessed all solutions in terms of the repeatability of coordinates as 16 an internal evaluation of applied models and strategies, and in terms of zenith tropospheric delays 17 (ZTD) and horizontal gradients with those of ERA-Interim numerical weather model (NWM) reanalysis. 18 When compared to the GOP Repro1 solution, the results of the GOP Repro2 yielded improvements of 19 approximately 50% and 25% in the repeatability of the horizontal and vertical components, 20 respectively, and of approximately 9% in tropospheric parameters. Vertical repeatability was reduced 21 from 4.14 mm to 3.73 mm when using the VMF1 mapping function, a priori ZHD, and non-tidal 22 atmospheric loading corrections from actual weather data. Raising the elevation cut-off angle from 3° 23 to 7° and then to 10° increased RMS from coordinates' repeatability, which was then confirmed by 24 independently comparing GNSS tropospheric parameters with the NWM reanalysis. The assessment 25 of tropospheric horizontal gradients with respect to the ERA-Interim revealed a strong sensitivity of 26 estimated gradients to the quality of GNSS antenna tracking performance. This impact was 27 demonstrated at the Mallorca station, where gradients systematically grew up to 5 mm during the 28 period between 2003 and 2008, before this behaviour disappeared when the antenna at the station 29 was changed. The impact of processing variants on long-term ZTD trend estimates was assessed at 172 30 EUREF stations with time-series longer than 10 years, resulting in most significant impact, site-specific, 31 due to the non-tidal atmospheric loading followed by the impact of changing elevation cut-off angle 32 from 3° to 10°. The other processing strategy had very small or negligible impact on estimated trends.

Keywords: GPS, reprocessing, zenith tropospheric delay, tropospheric horizontal gradients,
 coordinate time series, reference frame

35 **1 Introduction**

The US Global Positioning System (GPS) became operational in 1995 as the first Global Navigation Satellite System (GNSS). Since that time, this technology has been transformed into a fundamental technique for positioning and navigation in everyday life. Hundreds of GPS permanent stations have been deployed for scientific purposes throughout Europe and the world, and the first stations have collected GPS data for approximately the last two decades. In 1994, a science-driven global network of continuously operating GPS stations was established by the International GNSS Service, IGS 42 (http://www.igs.org) of the International Association of Geodesy (IAG) to support the determination 43 of precise GPS/GNSS orbits and, clocks and earth rotation parameters, which are necessary for 44 obtaining high-accuracy GNSS analyses for scientific applications. A similar network, but regional in its 45 scope, was also organized by the IAG Reference Frame Sub-Commission for Europe (EUREF) in 1996, 46 which was called the EUREF Permanent Network (EPN), http://epncb.oma.be (Bruyninx et al. 2012). 47 Although its primary purpose was to maintain the European Terrestrial Reference System (ETRS), the 48 EPN also attempted to develop a pan-European infrastructure for scientific projects and co-operations 49 (Ihde et al. 2014). Since 1996, the EPN has grown to include approximately 300 operating stations, 50 which are regularly distributed throughout Europe and its surrounding areas. Today, EPN data are 51 routinely analysed by 18 EUREF analysis centres.

Throughout the past two decades, GPS data analyses of both global and regional networks have been affected by various changes in processing strategy and updates of precise models and products, reference frames and software packages. To reduce discontinuities in products, particularly within coordinate time series, homogeneous reprocessing was initiated by the IGS and EUREF on a global and regional scale, respectively. To exploit the improvements in these IGS global products, the 2nd European reprocessing was performed in 2015-2016, with the ultimate goal of providing a newly realized ETRS.

59 Currently, station coordinate parameter time series from reprocessed solutions are mainly used in the 60 solid earth sciences as well as to maintain global and regional terrestrial reference systems. 61 Additionally, from an analytical perspective, the long-term series of estimated parameters and their 62 residuals are useful for assessing the performances of applied models and strategies over a given 63 period. Moreover, tropospheric parameters derived from this GNSS reanalysis could be useful for 64 climate research (Yuan et al., 1993), due to their high temporal resolution and unrivalled relative 65 accuracy for sensing water vapour when compared to other techniques, such as radio sounding, water 66 vapour radiometers, and radio occultation (Ning, 2012). In this context, the GNSS Zenith Tropospheric 67 Delay (ZTD) represents a site-specific parameter characterizing the total signal path delay in the zenith 68 due to both dry (hydrostatic) and wet contributions of the neutral atmosphere, the latter of which is 69 known to be proportional to precipitable water (Bevis et al. 1994).

70 With the 2nd EUREF reprocessing, the secondary goal of the GOP was to support the activity of Working 71 Group 3 of the COST Action ES1206 (http://gnss4swec.knmi.nl), which addresses the evaluation of 72 existing and future GNSS tropospheric products, and assesses their potential uses in climate research. 73 For this purpose, GOP provided several solution variants, with a special focus on optimal tropospheric 74 estimates, including VMF1 vs. GMF mapping functions, the use of different elevation cut-off angles, 75 and estimates of tropospheric horizontal gradients using different time resolutions. Additionally, in 76 order to enhance tropospheric outputs, we improved the processing strategy in a variety of ways 77 compared to the GOP Repro1 solutions (Douša and Václavovic, 2012): 1) by combining tropospheric 78 parameters in midnights and across GPS week breaks, 2) by checking weekly coordinates before their 79 substitutions in order to estimate tropospheric parameters, and 3) by filtering out problematic stations by checking the consistency of daily coordinates. The results of this GOP reprocessing, including all 80 81 available variants, were assessed using internal evaluations of applied models and strategy settings, 82 and external validations with independent tropospheric parameters derived from numerical weather

- 84 The processing strategy used in the 2nd GOP reanalysis of the EUREF permanent network is described
- 85 in Section 2 and, new approach that is developed to guarantee a continuity of estimated tropospheric
- parameters at midnights as well as between different GPS weeks is summarised in Section 3. The
- 87 relationship between mean tropospheric horizontal gradients and the quality of low-elevation GNSS
- tracking is explained in Section 4. The results of internal and external evaluations of GOP solution
- variants and processing models are presented in Section 4 and, the assessment of impacts of specific
 variants on estimated ZTD trends in Section 5. The last section concludes our findings and suggests
- 91 avenues of future research.

92 **2 GOP processing strategy and solution variants**

The EUREF GOP analysis centre was established in 1997, and contributed to operational EUREF 93 94 analyses until 2013 by providing final, rapid, and near real-time solutions. Recently, GOP changed its 95 contributions to that of a long-term homogeneous reprocessing of all data from the EPN historical archive. The GOP solution of the 1st EUREF reanalysis (Repro1) (Völksen, 2011) comprised the 96 97 processing of a sub-network of 70 EPN stations during the period of 1996-2008. In 2011, for the first 98 time, GOP reprocessed the entire EPN network (spanning a period of 1996-2010) in order to validate 99 the European reference frame and to provide the first homogeneous time series of tropospheric 100 parameters for all EPN stations (Douša and Václavovic, 2012).

101 In the 2nd EUREF reprocessing (Repro2), GOP analysed data obtained from the entire EPN network from 102 a period of 1996-2014 using the Bernese GNSS Software V5.2 (Dach et al., 2015). The GOP strategy 103 relies on a network approach utilizing double-difference observations. Only GPS data from the EPN 104 stations were included according to official validity intervals provided by the EPN central Bureau 105 (http://epncb.oma.be). Two products were derived from the reprocessing campaign in order to 106 contribute to a combination at the EUREF level performed by the coordinator of analysis centres and 107 the coordinator of troposphere products: 1) site coordinates and corresponding variance-covariance 108 information in daily and weekly SINEX files and 2) site tropospheric parameters in daily Tro-SINEX files.

109 This GOP processing was clustered into eight subnetworks (Figure 1) and then stacked into daily 110 network solutions with pre-eliminated integer phase ambiguities when ensuring strong ties to IGS08 111 reference frame. This strategy introduced state-of-the-art models (IERS Conventions, 2010) that are 112 recommended as standards for highly accurate GNSS analyses, particularly for the maintenance of the 113 reference frame. Additionally, the use of precise orbits obtained from the 2nd CODE global reprocessing 114 (Dach et al., 2014) guaranteed complete consistency between all models on both the provider and user 115 sides. Characteristics of this GOP data reprocessing strategy and their models are summarized in Table 116 1. Additionally, seven processing variants were performed during the GOP Repro2 analysis for studying 117 selected models or settings: a) applying tropospheric mapping function model GMF (Böhm et al., 118 2006a) vs. VMF1 (Böhm et al., 2006b), the latter based on actual weather information, b) increasing 119 the temporal resolution of tropospheric linear horizontal gradients in the north and east directions, c) 120 using different elevation cut-off angles, d) modelling atmospheric loading effects, and e) modelling 121 higher-order ionospheric effects. Table 2 summarizes the settings and models of solution variants 122 selected for generating coordinate and troposphere products, which are supplemented with variant 123 rationales.

124 Within the processing, we screened station coordinate repeatabilities from weekly combined solutions 125 and we identified any problematic station for which north/east/up residuals exceeded 15/15/30 mm 126 or RMS of north/east/up coordinate component exceeded values 10/10/20 mm. Such station was a 127 priori excluded from the tropospheric product for the corresponding day. There were other standard 128 control procedures within the processing when individual station could have been excluded, e.g. if a) 129 less than 60% of GNSS data available, b) code or phase data revealed poor quality, c) station metadata 130 were found inconsistent with data file header information (receiver, antenna and dome names, 131 antenna eccentricities) and, d) phase residuals were too large for all satellites in the processing period 132 indicating a problem with station. Tropospheric parameters were estimated practically without 133 constrains (a priori sigma greater than 1 m) thus parameter formal errors reflect relative uncertainties 134 of estimates. Usually, large errors indicate the lack of observations contributing to the parameter. 135 During the tropoposheric parameter evaluations, we applied filter for exceeding formal errors of 136 estimated parameters (ZTD sigma greater than 3 mm, normal cases stay below 1 mm).

137 **3 Ensuring ZTD continuity at midnights**

When site tropospheric parameter time series generated from the 2nd EUREF reprocessing are applied 138 139 to climate research, they should be free of artificial offsets in order to avoid misinterpretations (Bock 140 et al., 2014). However, GNSS processing is commonly performed on a daily basis according to adopted 141 standards for data and product dissemination. Thus far, EUREF analysis centres have provided 142 independent daily solutions, although precise IGS products are combined and distributed on a weekly 143 basis. Station coordinates are estimated on a daily basis and are later combined to form more stable 144 weekly solutions. According to the EUREF analysis centre guidelines 145 (http://www.epncb.oma.be/ documentation/guidelines/guidelines analysis centres.pdf), weekly 146 coordinates should be used to estimate tropospheric parameters on a daily basis, but there are no 147 requirements with which to guarantee the continuity of tropospheric parameters at midnights. 148 Additionally, there are also discontinuities on a weekly basis, as neither daily coordinates nor hourly 149 tropospheric parameters are combined across midnights between corresponding adjacent GPS weeks.

150 The impact of a 3-day combination was previously studied when assessing the tropospheric parameters stemming from the 2nd IGS reprocessing campaign 2016 in the GOP-TropDB (Győri and 151 Douša, 2016). We compared two global tropospheric products provided by the analysis centre CODE 152 153 (Centre of Orbit Determination in Europe) differing only in the procedure of combining tropospheric 154 parameters from the daily original solutions. The first product, COF, was based purely on a single-day 155 solution while the second product, COD, on a 3-day combination (Dach et al., 2014). A sub-daily 156 statistics were calculated by comparing 2-hour ZTD estimates from both products during 2013. There 157 were no significant biases observed, but mean standard deviation estimated from differences reached 158 0.8 mm in ZTD over a day, but almost 1.8 mm close to the day boundaries. Similarly, a dispersion 159 characterized by 1-sigma over all stations reached 0.5 mm for the former, but up to 1.2 mm for the 160 latter. Actual differences in ZTDs could even be significantly larger reaching up to several millimetres or more as the middle values of 2-hour ZTD estimates could have been compared only, i.e. at 1:00 UTC 161 and 23:00 UTC. 162

During the 1st GOP reprocessing, there was no way to guarantee tropospheric parameter continuity at
 midnight, as the troposphere was modelled by applying a piecewise constant model. In these cases,
 tropospheric parameters with a temporal resolution of one hour were reported in the middle of the

hour, as was originally estimated. In the 2nd GOP reprocessing, using again hourly estimates, we applied 166 a piecewise linear model for the tropospheric parameters. The parameter continuities at midnights 167 168 were not guaranteed implicitly, but only by an explicit combination of parameters at daily boundaries. 169 For the combination procedure we used three consecutive days while the tropospheric product stems 170 from the middle day. The procedure is done again for three consecutive days shifted by one day. A 171 similar procedure, using the piecewise constant model, was applied for estimating weekly coordinates 172 which aimed to minimize remaining effects in consistency at transition of GPS weeks (at Saturday 173 midnight). The coordinates of the weekly solution corresponding to the middle day of a three-day 174 combination were fixed for the tropospheric parameter estimates. In the last step, we transformed 175 the piecewise linear model to the piecewise constant model expressed in the middle of each hourly 176 interval (HR:30), which was saved in the TRO-SINEX format to support the EUREF combination 177 procedure requiring such sampling. The original piecewise linear parameter model was thus lost and 178 to retain this information in the official product in the TRO-SINEX format, we additionally stored values 179 for full hours (HR:00). Figure 2 summarizes four plots displaying tropospheric solutions with 180 discontinuities in the left panels (a), (c) and enforcing tropospheric continuities in the right panels (b, 181 d). While the upper plots (a), (b) display the piecewise constant model, bottom plots (c), (d) indicates 182 the solution representing the piecewise linear model. The GOP Repro1 implementation is thus 183 represented by Figure 2(a) plot while the GOP Repro2 solution corresponds to Figure 2(d) and, 184 alternatively Figure 2(b).

These theoretical concepts were practically tested using a limited data set in 1996 (Figure 3). The 185 186 panels in Figure 3 follow the organization of the theoretical plots shown in Figure 2; corresponding 187 formal errors are also plotted along with estimated ZTDs. Discontinuities are visible in the left-hand 188 plots and are usually accompanied by increasing formal errors for parameters close to data interval 189 boundaries. As expected, discontinuities disappear in the right-hand plots. Although the values 190 between 23:30 and 00:30 on two adjacent days are not connected by a line in the top-right plot, 191 continuity was enforced for midnight parameters anyway, as seen in the bottom-right plot. Formal 192 errors also became smooth near day boundaries, thus characterizing the contribution of data from 193 both days and demonstrating that the concept behaves as expected in its practical implementation.

4 Quality of the observations and impact on tropospheric gradients

Recently, we have developed a new interactive web interface to conduct tropospheric parameter 195 196 comparisons in the GOP-TropDB (Győri and Douša, 2016), which is being prepared for the IGS 197 Tropospheric Working Group web (<u>http://twg.igs.org/</u>). Using the interface, we observed large 198 systematic tropospheric gradients during specific years at several EPN stations. Generally, from GNSS 199 data, we can only estimate total tropospheric horizontal gradients without being able to distinguish 200 between dry and wet contributions. The former is mostly due to horizontal asymmetry in atmospheric 201 pressure, and the latter is due to asymmetry in the water vapour content. The latter is thus more 202 variable in time and space than the former (Li et al., 2015). Regardless, mean gradients should be close 203 to zero, whereas dry gradients may tend to point slightly more to the equator, corresponding to 204 latitudinal changes in atmosphere thickness (Meindl et al., 2004). Similarly, orography-triggered 205 horizontal gradients can appear due to the presence of high mountain ranges in the vicinity of the 206 station (Morel et al., 2015). Such systematic effects can reach the maximum sub-millimetre level, while 207 a higher long-term gradient (i.e. that above 1 mm), is likely more indicative of issues with site

instrumentation, the environment, or modelling effects. Therefore, in order to clearly identify thesesystematic effects, we also compared our gradients with those calculated from the ERA-Interim.

210 It is beyond the scope of this paper to investigate in detail the correlation between tropospheric 211 horizontal gradients and effects such as, for example, antenna tracking performance. However, we do 212 observe a strong impact in the most extreme case identified when comparing gradients from the GNSS and the ERA-Interim for all EPN stations. Figure 4 shows the monthly means of differences in the north 213 214 and east tropospheric gradients from the MALL station (Mallorca, Spain). These differences increase 215 from 0 mm up to -4 mm and 2 mm for the east and north gradients, respectively, within the period of 216 2003/06 - 2008/10. Such large monthly differences in GNSS and NWM gradients are not realistic, and 217 were attributed to data processing when long-term increasing biases dropped down to zero on 218 November 1, 2008, immediately after the antenna and receiver were changed at the station. During 219 the same period, also yearly mean ZTD differences to ERA-Interim steadily changed from about 3 mm 220 to about -12 mm and immediately dropping down to -2 mm in 2008 after the antenna change.

221 The EPN Central Bureau (http://epncb.oma.be), operating at the Royal Observatory of Belgium (ROB), 222 provides a web service for monitoring GNSS data quality and includes monthly snapshots of the 223 tracking characteristics of all stations. The sequence of plots displayed in Figure 5, representing the 224 interval of interest (2002, 2004, 2006 and 2008), reveals a slow but systematic and horizontally 225 asymmetric degradation of the capability of the antenna to track low-elevation observations at the 226 station. Therefore, we analysed days of the year (DoY) 302 and 306 (corresponding to October 28 and 227 November 1, 2008) with the in-house G-Nut/Anubis software (Václavovic and Douša, 2016) and 228 observed differences in the sky plots of these two days. The left-hand plot in Figure 6 depicts the severe 229 loss of dual-frequency observations up to a 25° elevation cut-off angle in the South-East direction (with 230 an azimuth of 90°-180°), which cause the tropospheric linear gradient of approximately 5 mm to point 231 in the opposite direction. Figure 10 also demonstrates that an increasing loss of second frequency 232 observations appears to occur in the East (represented as black dots). The right-hand plot in this figure 233 demonstrates that both of these effects fully disappeared after the antenna was replaced on October 234 30, 2008 (DoY 304), resulting in the appearance of normal sky plot characteristics and a GLONASS 235 constellation with one satellite providing only single frequency observations (represented as black 236 lines).

This situation demonstrates the high sensitivity of the estimated gradients on data asymmetry, particularly at low-elevation angles. The systematic behaviour of these monthly mean gradients, their variations from independent data and a profound progress over time, seem to be useful indicators of instrumentation-related issues at permanent GNSS stations. It is also considered that gradient parameters can be valuable method as a part of ZTD data screening procedure (Bock et al., 2016).

Although the station MALL represented an extreme case, biases at other stations were observed too, e.g. GOPE (1996-2002), TRAB (1999-2008), CREU (2000-2002), HERS (1999-2001), GAIA (2008-2014) and others. Site-specific, spatially or temporally correlated biases suggest different possible reasons such as site-instrumentation effects including the tracking quality and phase centre variation models, site-environment effects including multipath and seasonal variation (e.g. winter snow/ice coverage), edge-network effects when processing double-difference observations, spatially correlated effects in reference frame realization and possibly others. The problematic stations and periods mentioned above were however still included in comparisons and trend analysis because of the lack of objectivecriteria for their identification, which should be studied in future.

251 **5** Assessment of reprocessing solutions

GOP variants and reprocessing models were assessed by a number of criteria, including those of the internal evaluations of repeatability of station coordinates, residuals at reference stations, and the external validation of ZTDs and tropospheric horizontal gradients with data from numerical weather model (NWM) reanalyses.

256 **5.1 Repeatability of station coordinates**

We used coordinate repeatability to assess the quality of models applied in GNSS analysis. To be as thorough as possible, we not only assessed all GOP Repro2 variants but also assessed two GOP Repro1 solutions in order to discern improvements within the new reanalyses. The two Repro1 solutions differed in their used reference frames and PCV models: IGS05 and IGS08.

261 Table 3 summarizes mean coordinate repeatability in the north, east and up components of all stations 262 from their weekly combinations. All GOP Repro2 solution variants reached approximately 50% and 263 25% of the lower mean RMS of coordinate repeatability when compared to the GOP Repro1/IGS08 264 solution in its horizontal and vertical components, respectively. These values represent even greater 265 improvements when compared to the GOP-Repro1/IGS05 solution. Comparing these two Repro1 266 solutions clearly demonstrates the beneficial impact of the new PCV models and reference frames. The 267 observed differences between Repro2 and Repro1 also indicate an overall improvement of the 268 processing software from V5.0 to V5.2, and the enhanced quality of global precise orbit and earth 269 orientation products.

270 Various GOP Repro2 solutions were also used to assess the selected models. Variants GO0 and GO1 271 differ in their mapping functions (GMF vs VMF1) used to project ZTDs into slant path delays. These 272 comparisons demonstrate that vertical component repeatability improved from 4.14 mm to 3.97 mm, 273 whereas horizontal component repeatability decreased slightly. By increasing the elevation cut-off 274 angle from 3° to 7° (GO2) and 10° (GO3), we observed a slight increase in RMS from repeatability of all 275 coordinates. This can be explained by the positive impact of low-elevation observations on the 276 decorrelation of height and tropospheric parameters, despite the fact that applied models (such as 277 elevation-dependent weighting, PCVs, multipath) are still not optimal for including observations at 278 very low elevation angles. On the other hand, it should be noted that the VMF1 mapping function is 279 particularly tuned to observations at 3° elevation angle which leads to biases at higher elevation angles, 280 Zus et al. (2015).

281 The GO4 solution represents an official GOP contribution to EUREF combined products. It is identical 282 to the variant GO1, but applies a non-tidal atmospheric loading. Steigenberger et al. (2009) discussed 283 the importance of applying non-tidal atmospheric loading corrections together with precise a priori 284 ZHD model. It has been concluded that using mean, or slowly varying, empirical pressure values for 285 estimating a priori ZHD instead of true pressure values results in a partial compensation of atmospheric loading effects which is the case of GO1 solution. A positive 10% improvement in height repeatability 286 287 was observed for the GO4 solution. Our improvement was slightly lower than in a global scope 288 reported by Dach et al. (2011) with an improvement of 10-20% over all stations. As the effect depends on selected stations, a slightly higher impact in a global scale might be attributed to the station
 distribution, particularly differences in term of latitude and altitude.

No impact was observed from the higher-order ionospheric effects (GO4 vs. GO5) in term of coordinate repeatability. As the effect is systematic within the regional network (Fritsche et al., 2005) and it was mostly eliminated by using reference stations in the domains of interest. The combination of tropospheric horizontal gradients from 6-h to 24-h time resolution (GO4 vs. GO6), using the piecewise linear model, had a negligible impact on the repeatability of station coordinates too.

296 **5.2 Reference frame - residuals at fiducial stations**

297 The terrestrial reference frame (Altamimi et al., 2001) is a realization of a geocentric system of 298 coordinates used by space geodetic techniques. To avoid a degradation of GNSS products, differential 299 GNSS analysis methods require a proper referencing of the solution to the system applied in the 300 generation of precise GNSS orbit products. For this purpose, we often use the concept of fiducial 301 stations with precise coordinates well-known in the requested system. Such stations are used to define 302 the geodetic datum while their actual position can be re-adjusted by applying a condition minimizing 303 coordinate residuals. None station is able to guarantee a stable monumentation and unchanged 304 instrumentation during the whole reprocessing period. Thus a set of about 50 stations, with 100 and 305 more time periods for reference coordinates, was carefully prepared for datum definition in the GOP 306 reprocessing. An iterative procedure was applied then for every day by comparing a priori reference 307 coordinates with actually estimated ones and excluding fiducial station exceeding differences by 5, 5 308 and 15 mm in north, east and up components.

309 Figure 7 shows the evolution of the number of actually used fiducial stations (represented as red dots) 310 from all configured fiducial sites (represented as black dots) after applying an iterative procedure of 311 validation on a daily basis. This reprocessing began with the use of 16-20 fiducial stations in 1996, and 312 this number increased to reach a maximum of over 50 during the period from 2003-2011. After 2011, 313 this number decreased, due to a common loss of reference stations available from the last realization 314 of the global terrestrial reference frame without changes in its instrumentation. In most cases, only 2 315 or 3 stations were excluded from the total number, however, this number is lower for some daily solutions, indicating the removal of even more stations. The lowest number of fiducial sites (12) was 316 317 identified on day 209 of the year 1999 while, but low numbers were, generally, observed at the 318 beginning of the reprocessing period, in 1996. We observed consistent mean RMS errors for horizontal, 319 vertical, and total residuals of 6.47, 10.22, and 12.25 mm and 4.83, 7.94, 9.35 mm for daily and weekly 320 solutions, respectively, which demonstrate the stability of the reference system in the reprocessing. 321 The seasonality in height coordinate estimates characterized by the RMS of residuals from the 322 reference frame realization is dominated by errors due to modelling of the troposphere. We believe, 323 the main contribution stems from the insufficiencies in modelling of wet tropospheric delay, as the 324 effect has the most pronounced seasonal signal within the GNSS data analysis. Additionally, the 325 estimated station ZTD parameters and height are difficult to de-correlate. In the next section, the 326 strong seasonal variation in comparing zenith total delays estimated from GNSS and NWM data is 327 clearly visible.

328 5.3 Zenith total delays

We compared all reprocessed tropospheric parameters with respect to independent data from the ERA-Interim global reanalysis (Dee et al. 2011) provided by the European Centre for Medium-Range Weather Forecasts (ECMWF) from 1969 to the present. For the period of 1996-2014, we calculated tropospheric parameters (namely ZTD and tropospheric horizontal linear gradients) from the NWM for all EPN stations using the GFZ (German Research Centre for Geosciences) ray-tracing software (Zus et al., 2014). The comparison of tropospheric parameters was performed by applying the linear interpolation of GNSS parameters to the original NWM 6-hour representation, using the GOP TropoDB (Győri and Douša, 2016). For monthly statistics discussed in this section, we applied an iterative precedure for outlier detection using the 2 sigma criteria calculated from the compared differences

337 procedure for outlier detection using the 3-sigma criteria calculated from the compared differences.

Table 4 summarizes comparisons of GNSS ZTDs, and tropospheric horizontal gradients, from all GOP processing variants with those obtained from the ERA-Interim. Mean biases and standard deviations were first calculated for each stations and each month and then mean and standard deviation of these values were computed, characterizing dispersions of all statistical values over the ensemble of stations.

342 The results in the table indicate a mean ZTD bias -1.8 mm for all comparisons (GNSS – NWM) suggesting 343 ZTDs achieved from the NWM reanalysis are drier than those obtained from GNSS reprocessing. Similar 344 biases have been observed for all other European GNSS re-processing products during the period of 345 1996-2014 (Pacione et al., 2017). On the other hand, when processing the ERA-Interim using two 346 different software and methodologies within the GNSS4SWEC Benchmark campaign (Dousa et al., 347 2016) during May and June of 2013 in Central Europe, and by their comparing to two GNSS reference 348 products based on different processing methods, we observed bias differences within ±0.4 mm in ZTD. 349 As neither GNSS nor NWM is able to sense the troposphere with an absolute accuracy better than the 350 bias that we observed, we cannot make any conclusion, but its independence of the GNSS software. A 351 mixture of common processing aspects such as scope of GNSS network, applied tropospheric model, 352 precise orbit product and others could still cause such a small biases in GNSS analysis at least.

Comparing the results of the official GOP Repro2 solution (GO4) to those of the legacy solution (GO0) demonstrates an overall improvement of 9%, which corresponds to a similar comparison between the EUREF Repro1 and Repro2 products (Pacione et al., 2017). The improvement is assumed to be even larger (indicated by the coordinate repeatability), as the quality of ZTD retrievals are generally lower for NWM compared to GNSS from various intra-/inter-technique comparisons (Douša et al., 2016, Kačmařík et al., 2017, Bock and Nuret, 2009).

359 Comparing the GO1 and GO0 variants demonstrates that the VMF1 mapping function outperforms 360 GMF in term of standard deviation if the elevation cut-off angle of 3° is used. The change of mapping 361 function together with the use of more accurate a priori ZHD, resulted in the ZTD standard deviation improving from 8.8 mm (GO0) to 8.3 mm (GO1). However, bias was slightly increased which could be 362 363 partly attributed to the use of mean pressure model for a priori ZHD calculation and compensating part of the non-tidal atmospheric loading (see Section 5.1). Using non-tidal atmospheric loading 364 365 corrections along with precise modelling of a priori ZHD contributed to a small reduction of the bias 366 from -2.0 mm to -1.8 mm and, mainly, to the improvement by reducing this ZTD accuracy to 8.1 mm 367 (GO4). This corresponds with the previous assessment of the repeatability of station coordinates. Degradation in ZTD precision was also observed when the elevation cut-off angle was raised from 3° 368 369 to 7° (GO2) or 10° (GO3). No impacts on ZTD were, however, visible neither from additional modelling 370 of high-order ionospheric effects (GO5) nor from stacking of 6-hour horizontal gradients into daily 371 piecewise linear estimates (GO6).

372 Figure 8 displays the time series of statistics from comparisons of the GOP official ZTD product (GO4) with respect to the results of the ERA-Interim reanalysis. Mean bias and standard deviation were 373 374 derived from the monthly statistics of the 6-hourly GNSS-ERA differences. A 1-sigma range of the mean 375 values, represented by error bars, are additionally derived from all stations on a monthly basis. 376 Although the time series show homogeneous results over the given time span, a small increase in the 377 mean standard deviation over time likely corresponds with increasing number of EPN sites, rising from 378 approximately 30 to 300. The early years (1996-2001) also display a worse overall agreement in 1-379 sigma range of mean values over all stations, which can be attributed to the varying quality of historical 380 observations and precise orbit products. The mean bias varies from -3 to 1 mm during the period of 1996-2014, with a long-term mean of -1.8 mm (Table 4). The long-term mean is also relatively small 381 382 compared to the ZTD mean 1-sigma range of 3-5 mm.

383 5.4 Tropospheric horizontal linear gradients

Additional GNSS signal delay due to the tropospheric gradients were developed by McMillan (1995). The complete tropospheric model for the line-of-sight delay (ΔD_T) using parameters zenith hydrostatic delay (ZHD), zenith wet delay (ZWD) and first-order horizontal tropospheric gradients G_N and G_E , all expressed in units of length, is described as follows

$$\Delta D_T = mf_h(e)ZHD + mf_w(e)ZWD + mf_g(e)\cot(e)[G_N\cos(A) + G_E\sin(A)]$$
(1)

389 where e and a are observation elevation and azimuth angles and mf_h , mf_w , mf_g are hydrostatic, wet 390 and gradient mapping functions representing the projection from an elevation to the zenith. Horizontal gradients should optimally represent a ZTD change in a distances for north and east directions as it 391 392 could be represented by terms $G_N \cot(e)$ and $G_E \cot(e)$ in the equation. However, the gradients need 393 to be parametrized practically with respect to observation elevation angle instead of the distance 394 theoretically applicable to the tropospheric effect at various elevation angles. The interpretation of 395 the tropospheric horizontal gradients in the Bernese software represents north and east components of angle applied for the tilting the zenith direction in the mapping function with gradients representing 396 397 (in unit of length) the tilting angle multiplied by the delay in zenith (Meindl et al., 2004).

398 Similarly as in case of ZTD and coordinate assessment, Table 4 shows that tropospheric gradients 399 became worse when raising the elevation cut-off angle from 3° to 7° (GO2) or 10° (GO3). Mean 400 standard deviations of the GO2 and GO3 solutions increased by 8% and 12%, respectively, which is 401 valid for the whole period of monthly time series (not showed). No significant differences in temporal 402 variations of mean biases of the north and east tropospheric gradients variants were identified while 403 they shared a higher variability during the years 1996-2001. No impact of modelling of high-order 404 ionospheric effects (GO5) was observed. Statistics of GO4 and GO6 solutions compared to ERA-Interim 405 revealed that standard deviations dropped from 0.38 mm to 0.28 mm and from 0.40 mm to 0.29 mm 406 for the east and north gradients, respectively. Worse performance of the GO4 solution is attributed to 407 the fact that tropospheric horizontal gradients were estimated with a 6-h sampling interval using the 408 piecewise linear model with applying practically no absolute or relative constraints. In such cases, 409 increased correlations of the gradients with other parameters can cause instabilities in processing 410 certain stations at specific times; the gradients absorb some remaining errors in the GNSS analysis 411 model. The mean biases of the tropospheric gradients are considered to be negligible, but it was

demonstrated in Section 4 that some large systematic effects were indeed discovered and attributedto the quality of GNSS signal tracking.

414 Figure 9 displays monthly time series of statistics from comparisons of the GNSS and NWM 415 tropospheric horizontal gradients in north and east directions. Two solutions are highlighted in order 416 to demonstrate the impact of different parameter temporal resolutions; a 6-hour resolution is used 417 for GO4 and a 24-hour resolution is used for GO6. Seasonal variations are mainly pronounced when 418 observing mean standard deviations (top plot), whereas gradual improvement is more pronounced for 419 mean biases (bottom plot). The reduction of the initial mean biases in horizontal gradients, and the 420 corresponding 1-sigma ranges over the values from the ensemble of stations, can be attributed to the 421 improved availability and quality of low elevation observation tracking. Elevation cut-off angles for 422 collecting GNSS observations were initially configured station by station, ranging from 0° to 15°, until 423 2008 when the elevation cut-off angle 0° was recommended for all the stations.

424 Mean standard deviations and their 1-sigma ranges over all stations (Figure 9, top plot) are lower by a 425 factor of 1.3 for the solution with 24-hour resolution (GO6) compared to the 6-hour resolution (GO4); 426 the impact is also pronounced especially in the early years of the dataset. The improvement factor 427 ranges from 1.03 to 1.65 with the mean value of 1.35 overall stations and it is usually higher for years 428 before 2001. Theoretically, with 4 times more observations in GO6 the standard deviation was 429 expected to be divided by a factor of 2. This discrepancy indicates serious correlations in errors which 430 are among others stemming from the errors in precise products and models. Significant improvements, 431 however, indicates possible correlations between tropospheric gradients and other estimated 432 parameters, such as ambiguities, height and zenith total delays, and suggests a careful handling 433 particularly when applying a sub-daily temporal resolution.

434 **5.5 Spatial and temporal ZTD analysis**

We performed spatial and temporal analyses of all processed variants in order to assess the impact of different settings on tropospheric products. Zenith tropospheric delays from all variants were compared in such a way to enable assessing impact of any single processing change: 1) GO1-GO0 for mapping function and more precise a priori ZHD model, 2) GO2-GO1 and GO3-GO1 for different elevation cut-off angle, 3) GO4-GO1 for non-tidal atmospheric corrections, 4) GO5-GO4 for higherorder ionospheric corrections and, 5) GO6-GO4 for temporal resolution tropospheric horizontal gradients. Station-specific behavior is out of this paper and will be studied in future.

442 Geographical maps of spatially distributed biases and standard deviations in ZTDs from all compared 443 variants for the whole network are showed in Figure 10 and Figure 11. Additionally median, minimum 444 and maximum values of station-wise total statistics are provided in Table 5. The comparisons 445 demonstrated that the impact of the higher-order effect is fully negligible. Although overall mean 446 biases in Table 5 are small, the GO1-GO0 comparison indicates a small negative bias over a majority of 447 the stations, see Figure 10. Biases from the comparison of variants with different elevation cut-off 448 angles strongly indicates a station-specific behavior with a positive bias for stations around Poland, 449 which has not been explained yet. According to the table and Figure 11, the highest impact on standard 450 deviations is found in the GO1 vs. GO0 solutions comparison. The effect is latitude dependent and it 451 follows the increasing magnitude of ZTDs towards the equator. Detailed study illustrated in Figure 12, 452 Figure 13 and Figure 14 then illustrates ZTD statistics with respect to the station latitude, ellipsoidal 453 height and time, respectively.

454 Using VMF1 mapping function together with precise a priori ZHD from VMF1 instead of the GMF and 455 GPT models, respectively, see GO1 vs. GO0, we observe biases ranging from -1.52 to 0.70 mm and the 456 median value -0.36 mm and, according to Table 5, with a moderate latitudinal dependence, see Figure 457 12. A similar, but slightly larger negative bias of -0.94±0.28 mm, was reported Kacmarik et al. (2017) 458 studying 400 stations in the central Europe. Standard deviations range from 0.69 mm to 3.82 mm in 459 Table 5 with a profound increase along with the latitude, Figure 12, indicating the GPT performs worse 460 at higher latitudes. This fully corresponds to the results from the paper by Steigenberger et al. (2009) 461 demonstrating a partial compensation of the atmospheric loading effect by using the GPT model. In 462 case the atmospheric loading effect is not corrected for, the errors are mostly assimilated to the zenith 463 total delay parameters if station coordinates are fixed on a weekly basis. Additionally, Figure 14 shows 464 the standard deviation grows with time which can be attributed to the use of blind (GMF) and actual-465 weather (VMF1) mapping functions. The mapping function affects an optimal use of low-elevation 466 observations, which were growing in EUREF permanent network with time as demonstrated for WTZR 467 station in Figure 15.

468 Biases obtained from the comparison of different elevation cut-off angles, i.e. variants 3°/7° (GO2-469 GO1) and 3°/10° (GO3-GO1), range from -0.81 mm to 1.66 mm and -2.22 mm to 2.66 mm, respectively, 470 and standard deviations from 0.15 mm to 1.29 mm and 0.31 to 2.04 mm, see Table 5. Generally, the 471 impact of different elevation cut-off angle doesn't reveal any biases neither with respect to the latitude 472 (Figure 12) nor the station height (Figure 13). As expected, the impact is larger for the GO3-GO1 473 differences and affected particularly some stations. Yearly biases exceeding ±2.5 mm were identified 474 for BELL, DENT, MLVL, MOPS, POLV RAMO and SBG2 stations. Temporal dependences in the GO2-GO1 475 and GO3-GO1 comparisons, Figure 14, show that the scatter of station-specific biases steadily grows 476 in time which is assumed to be related to the higher availability of low-elevation observations. On the 477 other hand, a small impact is observed for the standard deviation compared to the other studied 478 effects. This indicates the elevation cut-off angle affects mainly ZTD biases, which has been also 479 reported by Ning and Elgered (2012).

480 Table 5 shows that biases due to the non-tidal atmospheric loading (GO4-GO1) range from -2.29 mm 481 to 5.55 mm, which is one of the largest impact compared to other comparison variants, and standard 482 deviations range from 0.68 mm to 4.72 mm that represents the second largest impact compared to all 483 other variants. Standard deviation larger than 3 mm was observed at some stations, such as JOZE, 484 MAD2, MADR, MDVO, MOPI, NYAL, SBG2, VENE and WETT. It should be emphasized this comparison 485 reflects differences due to the modelling of atmospheric loading corrections in GO4 and, a partial 486 compensation of the loading effect by zenith tropospheric delay estimates in the GO1 solution variant. 487 The differences are strongly station-dependent, but did not reveal any dependence on latitude, see 488 Figure 12. It shows, however, some degradation in standard deviation during the first years of the 489 reprocessing, see Figure 14. Since a similar degradation has not been observed for other comparison 490 variants, it can be related to the quality of pressure data used to compute atmospheric loading.

The impact of higher-order ionospheric effect (GO5-GO4) is negligible at all stations demonstrating total statistics for all stations within ±0.3 mm when applying the y-range about 10 times smaller than in other panels of Figure 12, Figure 13 and Figure 14. A strong latitudinal dependence is, however, clearly visible in Figure 12 as well as a temporal variability showing yearly statistics up to ±0.4 mm, Figure 14. Both dependences are due to the changing magnitude of ionospheric corrections, generally increasing towards the equator and a daily noon, and along with cycles of the solar magnetic activity,reaching peaks around years 2001 and 2014.

498 The impact of stacking tropospheric gradients from 6-hour to daily estimates (GO6-GO4) is almost 499 negligible in term of biases which stay below ±1 mm, Table 5 and Figure 10. However, standard 500 deviations range from 0.76 mm to 2.46 mm and grow towards the equator, Figure 12. That can be 501 certainly attributed to the more difficult modelling of a local asymmetry in the troposphere, which is 502 generally increasing together with the increasing of the water vapor content. There is no significant 503 temporal variation observed in Figure 14, but a small decrease in standard deviation. It can be 504 attributed to a higher stability of the gradient estimates with time, see Figure 9, when supported with 505 increased number of available low-elevation observations.

506 6 Impact of variants on long-term ZTD trend estimates

We assessed the impact of solution variant on long-term ZTD trend estimates by analysing 172 EUREF
stations providing the time-series of data longer than 10 years. For each station, the trend analysis was
performed without any data homogenization or outlier rejection as our focus was only on assessing
the impact of solution variants on the trend estimates. The ZTD trends were estimated using the least
squares regression method applied on model (Weatherhead et al., 1998)

512
$$Y_t = \mu + \beta X_t + S_t + \varepsilon_t$$
(2)

513 where μ is the constant term of the model, βX_t is the linear trend function with β representing the 514 trend magnitude, S_t represents the term modelled by the sine wave function of time X_t including

annual, 2^{nd} harmonics and daily variations, and finally \mathcal{E}_{t} is the noise in the data.

516 Site-by-site estimated ZTD trends from all the variants are provided in supplementary materials 517 completed by time-span information, number of records and estimated mean formal errors calculated 518 over all variants. In total, trends range from -0.99 to 0.96 mm/year. Although the individual station 519 trend provided in supplements could be compared to other studies, e.g. Baldysz et al. (2016), Klos et 520 al. (2016) or Nilsson and Elgered (2008), however, it should be strongly emphasized here that our 521 trends are estimated without any preceding time-series homogenization and the formal errors of the 522 trend estimates are underestimated by a factor 2-4 (Nilsson and Elgered, 2008).

523 Table 6 summarizes the statistics of estimated trend differences at all 172 stations, always between 524 particular variants as defined in Section 5.5. Interestingly, the most significant impact is observed due 525 to the non-tidal atmospheric loading effects reaching differences up to ±0.55 mm/year in ZTD trends 526 for some extreme cases from the ensemble of 172 stations, and an overall scatter of 0.50 mm/year 527 from the ensemble of stations. Changes in elevation cut-off angle, particularly from 3° to 10°, reveal 528 also a significant impact characterized by differences up to ±0.34 mm/year and the scatter of 0.32 529 mm/year. The impact of mapping function on trend estimates remains small, with a maximum 530 difference of 0.12 mm/year and the scatter below 0.08 mm/year, while other strategy changes, due to time resolution of tropospheric gradients and higher-order ionospheric effects, remains negligible, 531 532 always below ±0.04 mm/year for all 172 stations, with the scatter of the same magnitude. All mean biases over differences stay also below 0.05 mm/year. 533

Finally, we selected 12 stations optimally available over the entire 2nd re-processing period and all 534 estimated trends are displayed in Figure 16. Trends for 12 stations range from -0.05 to 0.38 mm/year 535 536 with formal errors of 0.02-0.04 mm/year. It should be noted the formal errors are underestimated by a factor of 2-4 because the noise is assumed white (Nilsson and Elgered (2008). For the 12 selected 537 538 stations, the most significant impact is observed in the change of elevation cut-off angle when reaching 539 differences up to 0.1 mm/year in estimated ZTD trends. A similar, but more extensive study, was 540 performed by Ning and Elgered (2012) for Integrated Water Vapor content (IWV), roughly equal to 1/6 (ZTD-ZHD) kg.m⁻², using larger differences in the elevation cut-off angle and obtaining highly sensitive 541 542 results in term of estimated IWV trends. Impacts of other strategies are generally below 0.05 mm/year 543 - variants GO4, GO5, and GO6 are very similar, but not consistent again with GO1, meaning the non-544 tidal atmospheric loading has a significant impact on trend estimates for selected stations with the 545 longest data time-series.

546 7 Conclusions

547 In this paper, we present results of the new GOP reanalysis of all stations within the EUREF Permanent 548 network during the period of 1996-2014. This reanalysis was completed during the 2nd EUREF 549 reprocessing to support the realization of a new European terrestrial reference system. In the 2nd 550 reprocessing, we focused on analysing a new product – GNSS tropospheric parameter time-series for 551 applications to climate research. To achieve this goal, we improved our strategy for combining 552 tropospheric parameters at midnights and at transitions in GPS weeks. We also performed seven 553 solution variants to study optimal troposphere modelling; we assessed each of these variants in terms 554 of their coordinate repeatability by using internal evaluations of the applied models and strategies. We 555 also compared tropospheric ZTD and tropospheric horizontal gradients with independent evaluations 556 obtained by numerical weather reanalysis via the ERA-Interim.

Results of the GOP Repro2 yielded improvements of approximately 50% and 25% for their horizontal and vertical component repeatability, respectively, when compared to those of the GOP Repro1 solution. Vertical repeatability was reduced from 4.14 mm to 3.73 mm when using the VMF1 mapping function, a priori ZHD, and non-tidal atmospheric loading corrections from actual weather data. Increasing the elevation cut-off angle from 3° to 7°/10° increased RMS errors of residuals from these coordinates' repeatability. All of these factors were also confirmed by the independent assessment of tropospheric parameters using NWM reanalysis data.

564 We particularly recommend using low-elevation observations along with the VMF1 mapping function, 565 as well as using precise a priori ZHD values together with the consistent model of non-tidal atmospheric 566 loading. While estimating tropospheric horizontal linear gradients improves coordinates' repeatability, 567 6-hour sampling without any absolute or relative constraints revealed a loss of stability due to their correlations with other parameters. On the other hand, 24-h piecewise linear gradients did not indicate 568 569 a worse repeatability of coordinates estimates. For saving the time needed for the processing of 4 570 times less gradient parameters, we could recommend as sufficient using unconstrained 24-h piecewise 571 model for the first-order tropospheric asymmetry.

572 The impact of processing variants on long-term ZTD trend estimates was assessed at 172 EUREF 573 stations with time-series longer than 10 years. The most significant impact was observed due to the 574 non-tidal atmospheric loading effect reaching differences up to ±0.55 mm/year in ZTD trends for some extreme cases from the ensemble of 172 stations. Changes in elevation cut-off angle, particularly from
3° to 10°, revealed also a significant impact reaching differences up to ±0.35 mm/year. The change of
mapping function was observed rather small, with a maximum difference of 0.12 mm/year, while other
strategy changes, due to time resolution of tropospheric gradients and higher-order ionospheric
effects, remained negligible, always below ±0.04 mm/year for all 172 stations.

580 Assessing the tropospheric horizontal gradients with respect to the ERA-Interim reanalysis data 581 revealed some long-term systematic behaviour linked to degradation in antenna tracking quality. We 582 presented an extreme case at the Mallorca station (MALL), in which gradients systematically increased 583 up to 5 mm from 2003-2008 while pointing in the direction of prevailing observations at low elevation 584 angles. However, these biases disappeared when the malfunctioning antenna was replaced. More 585 cases similar to this, although less extreme, have indicated that estimated tropospheric gradients are 586 extremely sensitive to the quality of GNSS antenna tracking, thus suggesting that these gradients can 587 be used to identify problems with GNSS data tracking in historical archives.

One of the main difficulties faced during the 2nd reprocessing was that of the quality of the historical 588 589 data, which contains a large variety of problems. We removed data that caused significant problems 590 in network processing when these could not be pre-eliminated from normal equations during the 591 combination process without still affecting daily solutions. To provide high-accuracy, high-resolution 592 GNSS tropospheric products, the elimination of such problematic data or stations is even more critical 593 considering the targeting static coordinates on a daily or weekly basis for the maintenance of the reference frame or the derivation of a velocity field. Before undertaking the 3rd EUREF reprocessing, 594 which is expected to begin after significant improvements have been made to state-of-the-art models, 595 596 products and software, we need to improve data quality control and clean the EUREF historical archive 597 in order to optimize any future reprocessing efforts and to increase the quality of tropospheric products. These efforts should also include the collection and documentation of all available 598 599 information from each step of the 2nd EUREF reprocessing, including individual contributions, EUREF 600 combinations, time-series analyses and coordinates, and independent evaluations of tropospheric 601 parameters.

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718 Table 1: Characteristics of GOP reprocessing models

Processing options	Description						
Products	CODE precise orbit and earth rotation parameters from the 2^{nd}						
	reprocessing.						
Observations	Dual-frequency code and phase GPS observations from L1 and L2						
	carriers. Elevation cut-off angle 3°, elevation-dependent weighting						
	1/cos ² (zenith), double-difference observations and with 3-minute						
	sampling rate.						
Reference frame	IGb08 realization, core stations set as fiducial after a consistency						
	checking. Coordinates estimated using a minimum constraint.						
Antenna model	GOP: IGS08_1832 model (receiver and satellite phase centre offsets						
	and variations).						
Troposphere	A priori zenith hydrostatic delay/mapping function: GPT/GMFh						
	(GO0) and VMF1/VMF1h (GO1-GO6). Estimated ZWD corrections						
	every hour using VMF1 wet mapping function; 5 m and 1 m for						
	absolute and relative constraints, respectively. Estimated horizontal						
	NS and EW tropospheric gradients every 6 hours (GO0-GO5) or 24						
	hours (GO6) without a priori tropospheric gradients and constraints.						
Ionosphere	Eliminated using ionosphere-free linear combination (GO0-GO6).						
	Applying higher-order effects estimated using CODE global						
	ionosphere product (GO5).						
Loading effects	Atmospheric tidal loading and hydrology loading not applied. Ocean						
	tidal loading FES2004 used. Non-tidal atmospheric loading						
	introduced in advanced variants from the model from TU-Vienna						
	(GO4-GO6).						

720 Table 2: GOP solution variants for the assessment of selected models and settings

Solution ID	Specific settings and differences	Remarks and rationales				
GO0	GMF and 3° cut-off	Legacy solution for Repro1				
GO1	VMF1 and 3° cut-off	New candidate for Repro2				
GO2	=GO1; 7° cut-off	Impact of elevation cut-off angle				
GO3	=GO1; 10° cut-off	Impact of elevation cut-off angle				
GO4	=GO1; atmospheric loading	Non-tidal atmospheric loading applied				
GO5	=GO4; higher-order ionosphere	Higher-order ionosphere effect not applied				
GO6	=GO4; 24-hour gradients	Stacking tropospheric gradients to 24-hour				
		sampling				

Solution	North RMS	East RMS	Up RMS
	[mm]	[mm]	[mm]
GOP-Repro1/IGS05	3.01	2.40	5.08
GOP-Repro1/IGS08	2.64	2.21	4.94
GO0	1.20	1.30	4.14
G01	1.23	1.33	3.97
GO2	1.24	1.33	4.01
GO3	1.26	1.34	4.07
GO4	1.14	1.24	3.73
GO5	1.14	1.24	3.73
GO6	1.14	1.24	3.73

Table 3: Comparison of GOP solution variants for north, east and up coordinate repeatability.

724 725 726 Table 4: Statistics (bias and standard deviations) of ZTD and tropospheric gradients from the seven reprocessing variants

compared to those obtained from the ERA-Interim NWM reanalysis. In addition to the statistics, 1-sigma range over ensemble of stations is provided.

Solution	ZTD bias [mm]	ZTD sdev [mm]	EGRD bias [mm]	EGRD sdev [mm]	NGRD bias [mm]	NGRD sdev [mm]
GO0	-1.5 ± 2.1	8.8 ± 2.0	-0.04 ± 0.08	0.39 ± 0.10	+0.01 ± 0.09	0.43 ± 0.12
GO1	-2.0 ± 2.1	8.3 ± 2.2	-0.04 ± 0.08	0.39 ± 0.10	+0.01 ± 0.09	0.42 ± 0.13
GO2	-1.9 ± 2.2	8.4 ± 2.2	-0.05 ± 0.10	0.41 ± 0.10	+0.00 ± 0.12	0.45 ± 0.12
GO3	-1.8 ± 2.3	8.5 ± 2.1	-0.08 ± 0.13	0.43 ± 0.11	-0.01 ± 0.14	0.49 ± 0.12
GO4	-1.8 ± 2.4	8.1 ± 2.1	-0.04 ± 0.09	0.38 ± 0.10	+0.00 ± 0.09	0.40 ± 0.12
GO5	-1.8 ± 2.4	8.1 ± 2.1	-0.05 ± 0.09	0.38 ± 0.10	+0.01 ± 0.08	0.40 ± 0.12
GO6	-1.8 ± 2.4	8.2 ± 2.1	-0.04 ± 0.08	0.29 ± 0.06	+0.01 ± 0.09	0.28 ± 0.06

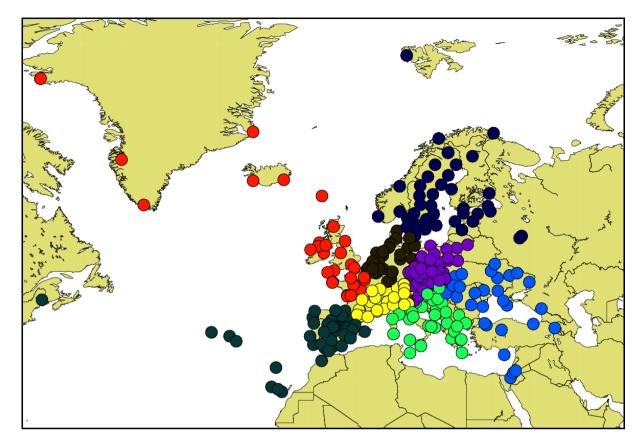
728 729 Table 5: Median, minimum (min) and maximum (max) values of total ZTD biases and standard deviation (sdev) over all

stations. Units are millimetres.

Compared	ZTD bias	ZTD bias	ZTD bias	ZTD sdev	ZTD sdev	ZTD sdev
variants	median	min	max	median	min	max
G01-G00	-0.36	-1.52	+0.70	2.01	0.69	3.82
G02-G01	+0.03	-0.81	+1.66	0.66	0.15	1.29
GO3-GO1	+0.03	-2.22	+2.66	1.10	0.31	2.04
GO4-GO1	+0.05	-3.29	+5.55	1.37	0.68	4.72
G05-G04	-0.02	-0.31	+0.07	0.07	0.04	0.30
G06-G04	-0.02	-0.23	+0.16	1.24	0.76	2.46

731 Table 6: Mean statistics of ZTD trends differences estimated between variants for 172 stations. Units are millimetres/year.

Statistics	GO1-GO0	G02-G01	GO3-GO1	GO4-GO1	GO5-GO4	GO6-GO4
Min	-0.118	-0.141	-0.308	-0.547	-0.017	-0.038
Max	0.045	0.179	0.331	0.452	0.031	0.036
mean	0.036	0.018	0.012	-0.048	0.007	0.001
Sdev	0.081	0.160	0.319	0.499	0.024	0.037





735 Figure 1: EUREF Permanent Network's clusters (designated by different colours) in the 2nd GOP reprocessing.

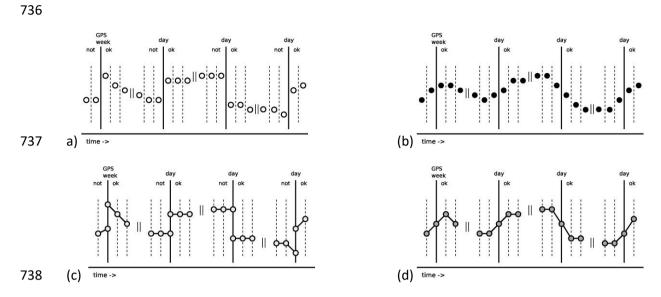


Figure 2: Charts of 4 variations on representations of tropospheric parameters. Right (b), (d) and left (a), (c) panels

739 740 741 display estimates made with and without midnight combinations, respectively. Top (a), (b) and bottom (c), (d) panels display the piecewise constant and the linear model, respectively.

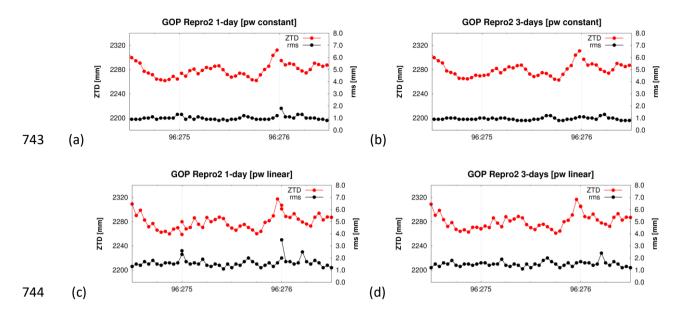


Figure 3: Four variations in representation of tropospheric parameters. Right (b), (d) and left (a), (c)

panels display estimates with and without midnight combinations, respectively. Top (a), (b) and

bottom (c), (d) panels display the piecewise constant and the piecewise linear model, respectively.

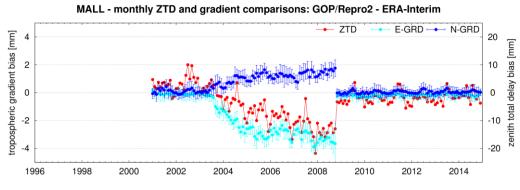


Figure 4: MALL station - monthly mean differences in tropospheric horizontal gradients with respect to the ERA-Interim.

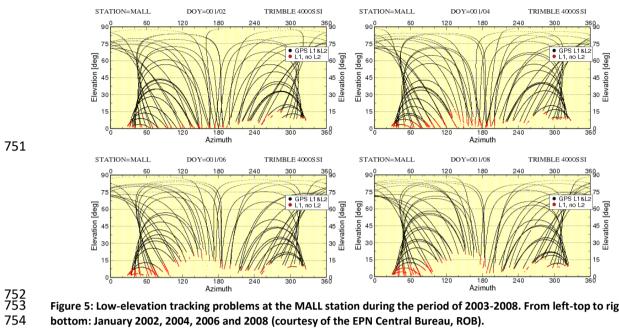
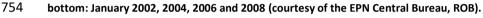
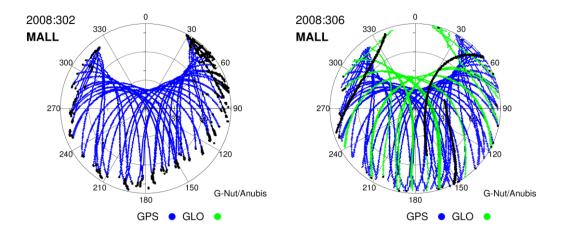
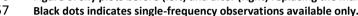


Figure 5: Low-elevation tracking problems at the MALL station during the period of 2003-2008. From left-top to right-





757 Figure 6: Sky plots before (left) and after (right) replacing the malfunctioning antenna at the MALL site (Oct 30, 2008). Black dots indicates single-frequency observations available only.



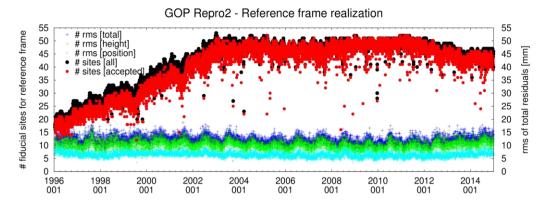


Figure 7: Statistics of the daily reference system realization: a) RMS of residuals at fiducial stations (representing the total, height and position); b) number of stations (all and accepted after an iterative control)

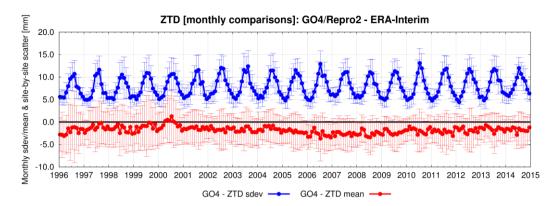
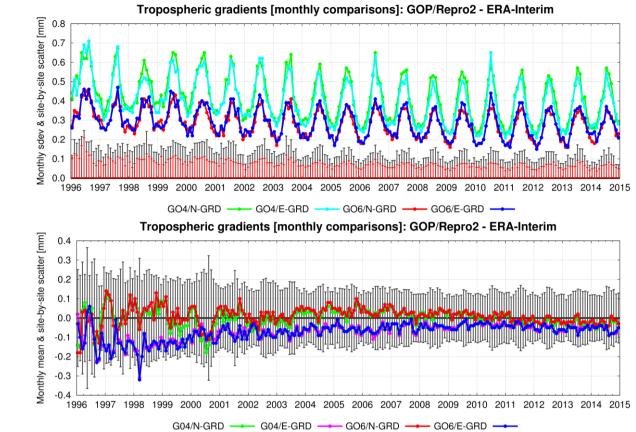


Figure 8: Monthly means of bias and standard deviation of official GOP ZTD product compared to those of the ERA-Interim.
 Error bars indicate standard errors of mean values over all compared stations.



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Figure 9: Monthly means of bias and standard deviation of tropospheric horizontal north (N-GRD) and east (E-GRD) 770 gradients compared to those obtained by ERA-Interim. Note: Similar products are almost superposed. Error bars indicate 771 standard errors of mean values over all compared stations plotted from the zero y-axis to emphasise seasonal variations

772 and trends. Error bars are displayed for north gradients only, however, being representative for the east gradients too.

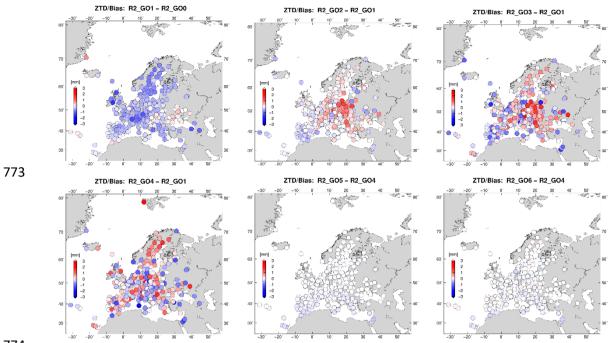




Figure 10: Geographic visualization of biases from inter-comparisons of GOP 2nd reprocessing variants.

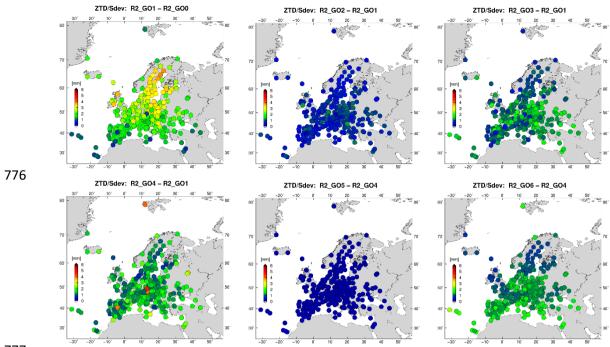




Figure 11: Geographic visualization of standard deviations from inter-comparisons of GOP 2nd reprocessing variants.

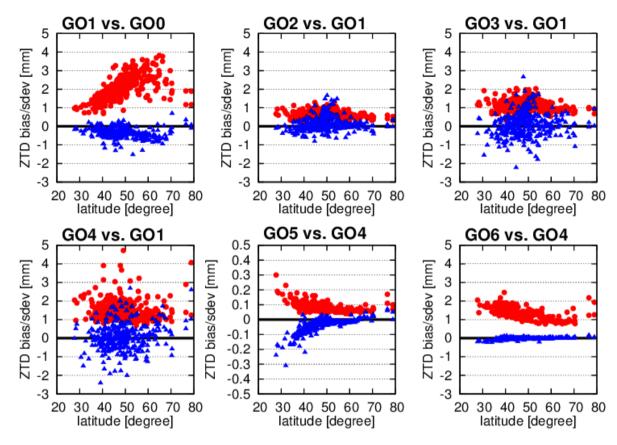


Figure 12: Dependence of ZTD biases (blue) and standard deviations (red) from inter-comparisons of GOP 2nd

781 reprocessing solution variants on station latitude. Note different y-range for the GO5 vs. GO4 comparison.

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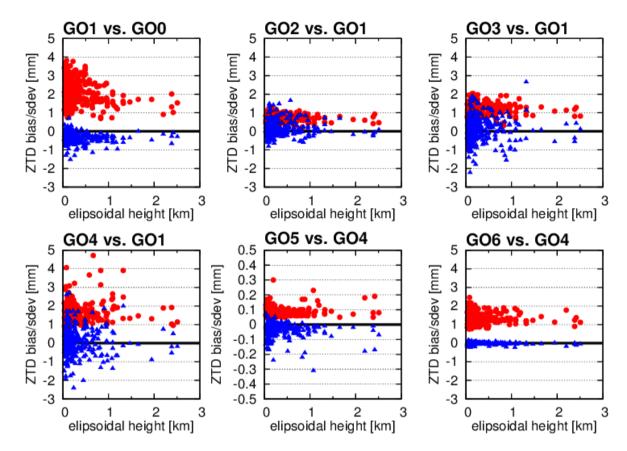
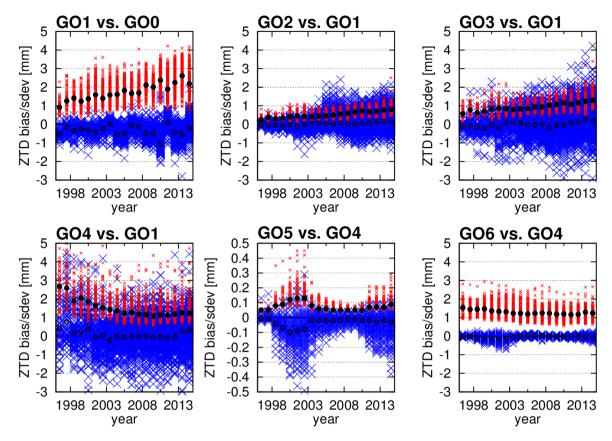


Figure 13: Dependence of ZTD biases (blue) and standard deviations (red) from inter-comparisons of GOP 2nd reprocessing
 solution variants on station ellipsoidal height. Note different y-range for the GO5 vs. GO4 comparison.



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787 Figure 14: Dependence of ZTD biases (blue), mean biases (unfilled black circles), standard deviations (red) and mean

standard deviations (filled black circles) from inter-comparisons of GOP 2nd reprocessing solution variants on year. Note
 different y-range for the GO5 vs. GO4 comparison.

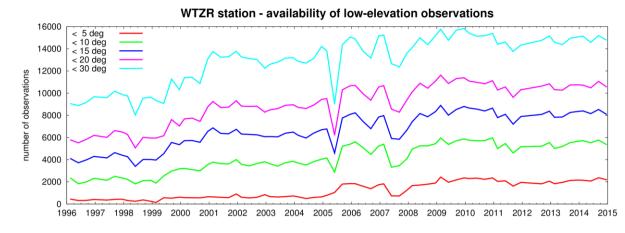


Figure 15: Availability of observations at low-elevation angles (below 5°, 10°, 15°, 20° and 30°) for WTZR station.

