1 Tropospheric products of the 2nd <u>GOP</u> European GNSS reprocessing (1996-2014)

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8 Abstract

9 In this paper, we present results of the 2nd reprocessing of all data from 1996 to 2014 from all stations 10 in the European GNSS permanent network as performed at the Geodetic Observatory Pecný (GOP). 11 While the original goal of this research was to ultimately contribute to new realization of the European 12 terrestrial reference system, we also aim to provide a new set of GNSS tropospheric parameter time 13 series with possible applications to climate research. To achieve these goals, we improved a strategy 14 to guarantee the continuity of these tropospheric parameters and we prepared several variants of 15 troposphere modelling. We then assessed all solutions in terms of the repeatability of coordinates as 16 an internal evaluation of applied models and strategies, and in terms of zenith tropospheric delays 17 (ZTD) and horizontal gradients with those of ERA-Interim numerical weather model (NWM) reanalysis. 18 When compared to the GOP Repro1 solution, the results of the GOP Repro2 yielded improvements of 19 approximately 50% and 25% in the repeatability of the horizontal and vertical components, 20 respectively, and of approximately 9% in tropospheric parameters. Vertical repeatability was reduced 21 from 4.14 mm to 3.73 mm when using the VMF1 mapping function, a priori ZHD, and non-tidal 22 atmospheric loading corrections from actual weather data. Raising the elevation cut-off angle from 3° 23 to 7° and then to 10° increased RMS from coordinates' repeatability, which was then confirmed by 24 independently comparing GNSS tropospheric parameters with the NWM reanalysis. The assessment 25 of tropospheric horizontal gradients with respect to the ERA-Interim revealed a strong sensitivity of 26 estimated gradients to the quality of GNSS antenna tracking performance. This impact was 27 demonstrated at the Mallorca station, where gradients systematically grew up to 5 mm during the 28 period between 2003 and 2008, before this behaviour disappeared when the antenna at the station 29 was changed. The impact of processing variants on long-term ZTD trend estimates was assessed at 172 30 EUREF stations with time-series longer than 10 years, resulting in most significant impact, site-specific, 31 due to the non-tidal atmospheric loading followed by the impact of changing elevation cut-off angle 32 from 3° to 10°. The other processing strategy had very small or negligible impact on estimated trends.

Keywords: GPS, reprocessing, zenith tropospheric delay, tropospheric horizontal gradients,
 coordinate time series, reference frame

35 **1** Introduction

The US Global Positioning System (GPS) became operational in 1995 as the first Global Navigation Satellite System (GNSS). Since that time, this technology has been transformed into a fundamental technique for positioning and navigation in everyday life. Hundreds of GPS permanent stations have been deployed for scientific purposes throughout Europe and the world, and the first stations have collected GPS data for approximately the last two decades. In 1994, a science-driven global network of continuously operating GPS stations was established by the International GNSS Service, IGS 42 (http://www.igs.org) of the International Association of Geodesy (IAG) to support the determination 43 of precise GPS/GNSS orbits and, clocks and earth rotation parameters, which are necessary for 44 obtaining high-accuracy GNSS analyses for scientific applications. A similar network, but regional in its 45 scope, was also organized by the IAG Reference Frame Sub-Commission for Europe (EUREF) in 1996, 46 which was called the EUREF Permanent Network (EPN), http://epncb.oma.be (Bruyninx et al. 2012). 47 Although its primary purpose was to maintain the European Terrestrial Reference System (ETRS), the 48 EPN also attempted to develop a pan-European infrastructure for scientific projects and co-operations 49 (Ihde et al. 2014). Since 1996, the EPN has grown to include approximately 300 operating stations, 50 which are regularly distributed throughout Europe and its surrounding areas. Today, EPN data are 51 routinely analysed by 18 EUREF analysis centres.

Throughout the past two decades, GPS data analyses of both global and regional networks have been affected by various changes in processing strategy and updates of precise models and products, reference frames and software packages. To reduce discontinuities in products, particularly within coordinate time series, homogeneous reprocessing was initiated by the IGS and EUREF on a global and regional scale, respectively. To exploit the improvements in these IGS global products, the 2nd European reprocessing was performed in 2015-2016, with the ultimate goal of providing a newly realized ETRS.

59 Currently, station coordinate parameter time series from reprocessed solutions are mainly used in the 60 solid earth sciences as well as to maintain global and regional terrestrial reference systems. 61 Additionally, from an analytical perspective, the long-term series of estimated parameters and their 62 residuals are useful for assessing the performances of applied models and strategies over a given 63 period. Moreover, tropospheric parameters derived from this GNSS reanalysis could be useful for 64 climate research (Yuan et al., 1993), due to their high temporal resolution and unrivalled relative 65 accuracy for sensing water vapour when compared to other techniques, such as radio sounding, water 66 vapour radiometers, and radio occultation (Ning, 2012). In this context, the GNSS Zenith Tropospheric 67 Delay (ZTD) represents a site-specific parameter characterizing the total signal path delay in the zenith 68 due to both dry (hydrostatic) and wet contributions of the neutral atmosphere, the latter of which is 69 known to be proportional to precipitable water (Bevis et al. 1994).

70 With the 2nd EUREF reprocessing, the secondary goal of the GOP was to support the activity of Working 71 Group 3 of the COST Action ES1206 (http://gnss4swec.knmi.nl), which addresses the evaluation of 72 existing and future GNSS tropospheric products, and assesses their potential uses in climate research. 73 For this purpose, GOP provided several solution variants, with a special focus on optimal tropospheric 74 estimates, including VMF1 vs. GMF mapping functions, the use of different elevation cut-off angles, 75 and estimates of tropospheric horizontal gradients using different time resolutions. Additionally, in 76 order to enhance tropospheric outputs, we improved the processing strategy in a variety of ways 77 compared to the GOP Repro1 solutions (Douša and Václavovic, 2012): 1) by combining tropospheric 78 parameters in midnights and across GPS week breaks, 2) by checking weekly coordinates before their 79 substitutions in order to estimate tropospheric parameters, and 3) by filtering out problematic stations by checking the consistency of daily coordinates. The results of this GOP reprocessing, including all 80 81 available variants, were assessed using internal evaluations of applied models and strategy settings, 82 and external validations with independent tropospheric parameters derived from numerical weather 83 reanalyses.

- 84 The processing strategy used in the 2nd GOP reanalysis of the EUREF permanent network is described
- in Section 2 and, new approach that is developed to guarantee a continuity of estimated tropospheric
- 86 parameters at midnights as well as between different GPS weeks is summarised in Section 3. The
- relationship between mean tropospheric horizontal gradients and the quality of low-elevation GNSS
- tracking is explained in Section 4. The results of internal and external evaluations of GOP solution
- variants and processing models are presented in Section 4 and, the assessment of impacts of specific
 variants on estimated ZTD trends in Section 5. The last section concludes our findings and suggests
- 91 avenues of future research.

92 **2 GOP processing strategy and solution variants**

The EUREF GOP analysis centre was established in 1997, and contributed to operational EUREF 93 94 analyses until 2013 by providing final, rapid, and near real-time solutions. Recently, GOP changed its 95 contributions to that of a long-term homogeneous reprocessing of all data from the EPN historical archive. The GOP solution of the 1st EUREF reanalysis (Repro1) (Völksen, 2011) comprised the 96 97 processing of a sub-network of 70 EPN stations during the period of 1996-2008. In 2011, for the first 98 time, GOP reprocessed the entire EPN network (spanning a period of 1996-2010) in order to validate 99 the European reference frame and to provide the first homogeneous time series of tropospheric 100 parameters for all EPN stations (Douša and Václavovic, 2012).

101 In the 2nd EUREF reprocessing (Repro2), GOP analysed data obtained from the entire EPN network from 102 a period of 1996-2014 using the Bernese GNSS Software V5.2 (Dach et al., 2015). The GOP strategy 103 relies on a network approach utilizing double-difference observations. Only GPS data from the EPN 104 stations were included according to official validity intervals provided by the EPN central Bureau 105 (http://epncb.oma.be). Two products were derived from the reprocessing campaign in order to 106 contribute to a combination at the EUREF level performed by the coordinator of analysis centres and 107 the coordinator of troposphere products: 1) site coordinates and corresponding variance-covariance 108 information in daily and weekly SINEX files and 2) site tropospheric parameters in daily Tro-SINEX files.

109 This GOP processing was clustered into eight subnetworks (Figure 1Figure 1) and then stacked into 110 daily network solutions with pre-eliminated integer phase ambiguities when ensuring strong ties to 111 IGS08 reference frame. This strategy introduced state-of-the-art models (IERS Conventions, 2010) that 112 are recommended as standards for highly accurate GNSS analyses, particularly for the maintenance of 113 the reference frame. Additionally, the use of precise orbits obtained from the 2nd CODE global 114 reprocessing (Dach et al., 2014) guaranteed complete consistency between all models on both the 115 provider and user sides. Characteristics of this GOP data reprocessing strategy and their models are 116 summarized in Table 1 Table 1. Additionally, seven processing variants were performed during the GOP 117 Repro2 analysis for studying selected models or settings: a) applying tropospheric mapping function 118 model GMF (Böhm et al., 2006a) vs. VMF1 (Böhm et al., 2006b), the latter based on actual weather 119 information, b) increasing the temporal resolution of tropospheric linear horizontal gradients in the 120 north and east directions, c) using different elevation cut-off angles, d) modelling atmospheric loading 121 effects, and e) modelling higher-order ionospheric effects. Table 2 summarizes the settings and 122 models of solution variants selected for generating coordinate and troposphere products, which are 123 supplemented with variant rationales.

124 Within the processing, we screened station coordinate repeatabilities from weekly combined solutions 125 and we identified any problematic station for which north/east/up residuals exceeded 15/15/30 mm 126 or RMS of north/east/up coordinate component exceeded values 10/10/20 mm. Such station was a 127 priori excluded from the tropospheric product for the corresponding day. There were other standard 128 control procedures within the processing when individual station could have been excluded, e.g. if a) 129 less than 60% of GNSS data available, b) code or phase data revealed poor quality, c) station metadata 130 were found inconsistent with data file header information (receiver, antenna and dome names, 131 antenna eccentricities) and, d) phase residuals were too large for all satellites in the processing period 132 indicating a problem with station. Tropospheric parameters were estimated practically without 133 constrains (a priori sigma greater than 1 m) thus parameter formal errors reflect relative uncertainties 134 of estimates. Usually, large errors indicate the lack of observations contributing to the parameter. 135 During the tropoposheric parameter evaluations, we applied filter for exceeding formal errors of 136 estimated parameters (ZTD sigma greater than 3 mm, normal cases stay below 1 mm).

137 **3 Ensuring ZTD continuity at midnights**

When site tropospheric parameter time series generated from the 2nd EUREF reprocessing are applied 138 139 to climate research, they should be free of artificial offsets in order to avoid misinterpretations (Bock 140 et al., 2014). However, GNSS processing is commonly performed on a daily basis according to adopted 141 standards for data and product dissemination. Thus far, EUREF analysis centres have provided 142 independent daily solutions, although precise IGS products are combined and distributed on a weekly 143 basis. Station coordinates are estimated on a daily basis and are later combined to form more stable 144 weekly solutions. According to the EUREF analysis centre guidelines 145 (http://www.epncb.oma.be/ documentation/guidelines/guidelines analysis centres.pdf), weekly 146 coordinates should be used to estimate tropospheric parameters on a daily basis, but there are no 147 requirements with which to guarantee the continuity of tropospheric parameters at midnights. 148 Additionally, there are also discontinuities on a weekly basis, as neither daily coordinates nor hourly 149 tropospheric parameters are combined across midnights between corresponding adjacent GPS weeks.

150 The impact of a 3-day combination was previously studied when assessing the tropospheric parameters stemming from the 2nd IGS reprocessing campaign 2016 in the GOP-TropDB (Győri and 151 Douša, 2016). We compared two global tropospheric products provided by the analysis centre CODE 152 153 (Centre of Orbit Determination in Europe) differing only in the procedure of combining tropospheric 154 parameters from the daily original solutions. The first product, COF, was based purely on a single-day 155 solution while the second product, COD, on a 3-day combination (Dach et al., 2014). A sub-daily 156 statistics were calculated by comparing 2-hour ZTD estimates from both products during 2013. There 157 were no significant biases observed, but mean standard deviation estimated from differences reached 158 0.8 mm in ZTD over a day, but almost 1.8 mm close to the day boundaries. Similarly, a dispersion 159 characterized by 1-sigma over all stations reached 0.5 mm for the former, but up to 1.2 mm for the 160 latter. Actual differences in ZTDs could even be significantly largerr reaching up to several millimetres 161 or more as the middle values of 2-hour ZTD estimates could have been compared only, i.e. at 1:00 UTC and 23:00 UTC, because this case used approximations leading to smooth low-resolution values close 162 163 to the day boundaries.

During the 1st GOP reprocessing, there was no way to guarantee tropospheric parameter continuity at midnight, as the troposphere was modelled by applying a piecewise constant model. In these cases,

166 tropospheric parameters with a temporal resolution of one hour were reported in the middle of the hour, as was originally estimated. In the 2nd GOP reprocessing, using again hourly estimates, we applied 167 168 a piecewise linear model for the tropospheric parameters. The parameter continuities at midnights 169 were not guaranteed implicitly, but only by an explicit combination of parameters at daily boundaries. 170 For the combination procedure we used three consecutive days while the tropospheric product stems 171 from the middle day. The procedure is done again for three consecutive days shifted by one day. A 172 similar procedure, using the piecewise constant model, was applied for estimating weekly coordinates 173 which aimed to minimize remaining effects in consistency at transition of GPS weeks (at Saturday 174 midnight). The coordinates of the weekly solution corresponding to the middle day of a three-day 175 combination were fixed for the tropospheric parameter estimates. In the last step, we transformed 176 the piecewise linear model to the piecewise constant model expressed in the middle of each hourly 177 interval (HR:30), which was saved in the TRO-SINEX format to support the EUREF combination 178 procedure requiring such sampling. The original piecewise linear parameter model was thus lost and 179 to retain this information in the official product in the TRO-SINEX format, we additionally stored values 180 for full hours (HR:00). Figure 2 Figure 2 summarizes four plots displaying tropospheric solutions with 181 discontinuities in the left panels (a), (c) and enforcing tropospheric continuities in the right panels (b, d). While the upper plots (a), (b) display the piecewise constant model, bottom plots (c), (d) indicates 182 183 the solution representing the piecewise linear model. The GOP Repro1 implementation is thus 184 represented by Figure 2Figure 2(a) plot while the GOP Repro2 solution corresponds to Figure 2Figure 185 2(d) and, alternatively Figure 2 Figure 2(b).

These theoretical concepts were practically tested using a limited data set in 1996 (Figure 3). The 186 187 panels in Figure 3 follow the organization of the theoretical plots shown in Figure 2Figure 2; corresponding formal errors are also plotted along with estimated ZTDs. Discontinuities are visible in 188 189 the left-hand plots and are usually accompanied by increasing formal errors for parameters close to 190 data interval boundaries. As expected, discontinuities disappear in the right-hand plots. Although the 191 values between 23:30 and 00:30 on two adjacent days are not connected by a line in the top-right plot, 192 continuity was enforced for midnight parameters anyway, as seen in the bottom-right plot. Formal 193 errors also became smooth near day boundaries, thus characterizing the contribution of data from 194 both days and demonstrating that the concept behaves as expected in its practical implementation.

4 Quality of the observations and impact on tropospheric gradients

196 Recently, we have developed a new interactive web interface to conduct tropospheric parameter 197 comparisons in the GOP-TropDB (Győri and Douša, 2016), which is being prepared for the IGS 198 Tropospheric Working Group web (<u>http://twg.igs.org/</u>). Using the interface, we observed large 199 systematic tropospheric gradients during specific years at several EPN stations. Generally, from GNSS 200 data, we can only estimate total tropospheric horizontal gradients without being able to distinguish 201 between dry and wet contributions. The former is mostly due to horizontal asymmetry in atmospheric 202 pressure, and the latter is due to asymmetry in the water vapour content. The latter is thus more 203 variable in time and space than the former (Li et al., 2015). Regardless, mean gradients should be close 204 to zero, whereas dry gradients may tend to point slightly more to the equator, corresponding to latitudinal changes in atmosphere thickness (Meindl et al., 2004). Similarly, orography-triggered 205 206 horizontal gradients can appear due to the presence of high mountain ranges in the vicinity of the 207 station (Morel et al., 2015). Such systematic effects can reach the maximum sub-millimetre level, while 208 a higher long-term gradient (i.e. that above 1 mm), is likely more indicative of issues with site instrumentation, the environment, or modelling effects. Therefore, in order to clearly identify thesesystematic effects, we also compared our gradients with those calculated from the ERA-Interim.

211 It is beyond the scope of this paper to investigate in detail the correlation between tropospheric 212 horizontal gradients and effects such as, for example, antenna tracking performance. However, we do 213 observe a strong impact in the most extreme case identified when comparing gradients from the GNSS 214 and the ERA-Interim for all EPN stations. Figure 4Figure 4 shows the monthly means of differences in 215 the north and east tropospheric gradients from the MALL station (Mallorca, Spain). These differences 216 increase from 0 mm up to -4 mm and 2 mm for the east and north gradients, respectively, within the 217 period of 2003/06 - 2008/10. Such large monthly differences in GNSS and NWM gradients are not 218 realistic, and were attributed to data processing when long-term increasing biases dropped down to 219 zero on November 1, 2008, immediately after the antenna and receiver were changed at the station. 220 During the same period, also yearly mean ZTD differences to ERA-Interim steadily changed from about 221 3 mm to about -12 mm and immediately dropping down to -2 mm in 2008 after the antenna change.

222 The EPN Central Bureau (http://epncb.oma.be), operating at the Royal Observatory of Belgium (ROB), 223 provides a web service for monitoring GNSS data quality and includes monthly snapshots of the 224 tracking characteristics of all stations. The sequence of plots displayed in Figure 5, representing 225 the interval of interest (2002, 2004, 2006 and 2008), reveals a slow but systematic and horizontally 226 asymmetric degradation of the capability of the antenna to track low-elevation observations at the 227 station. Therefore, we analysed days of the year (DoY) 302 and 306 (corresponding to October 28 and 228 November 1, 2008) with the in-house G-Nut/Anubis software (Václavovic and Douša, 2016) and 229 observed differences in the sky plots of these two days. The left-hand plot in Figure 6-Figure 6 depicts 230 the severe loss of dual-frequency observations up to a 25° elevation cut-off angle in the South-East 231 direction (with an azimuth of 90°-180°), which cause the tropospheric linear gradient of approximately 232 5 mm to point in the opposite direction. Figure 10 also demonstrates that an increasing loss of second 233 frequency observations appears to occur in the East (represented as black dots). The right-hand plot 234 in this figure demonstrates that both of these effects fully disappeared after the antenna was replaced 235 on October 30, 2008 (DoY 304), resulting in the appearance of normal sky plot characteristics and a 236 GLONASS constellation with one satellite providing only single frequency observations (represented as 237 black lines).

This situation demonstrates the high sensitivity of the estimated gradients on data asymmetry, particularly at low-elevation angles. The systematic behaviour of these monthly mean gradients, their variations from independent data and a profound progress over time, seem to be useful indicators of instrumentation-related issues at permanent GNSS stations. It is also considered that gradient parameters can be valuable method as a part of ZTD data screening procedure (Bock et al., 2016).

Although the station MALL represented an extreme case, biases at other stations were observed too, e.g. GOPE (1996-2002), TRAB (1999-2008), CREU (2000-2002), HERS (1999-2001), GAIA (2008-2014) and others. Site-specific, spatially or temporally correlated biases suggest different possible reasons such as site-instrumentation effects including the tracking quality and phase centre variation models, site-environment effects including multipath and seasonal variation (e.g. winter snow/ice coverage), edge-network effects when processing double-difference observations, spatially correlated effects in reference frame realization and possibly others. The problematic stations and periods mentioned

- above were however still included in comparisons and trend analysis because of the lack of objective
- 251 criteria for their identification, which should be studied in future.

252 **5** Assessment of reprocessing solutions

GOP variants and reprocessing models were assessed by a number of criteria, including those of the internal evaluations of repeatability of station coordinates, residuals at reference stations, and the external validation of ZTDs and tropospheric horizontal gradients with data from numerical weather model (NWM) reanalyses.

257 **5.1 Repeatability of station coordinates**

We used coordinate repeatability to assess the quality of models applied in GNSS analysis. To be as thorough as possible, we not only assessed all GOP Repro2 variants but also assessed two GOP Repro1 solutions in order to discern improvements within the new reanalyses. The two Repro1 solutions differed in their used reference frames and PCV models: IGS05 and IGS08.

262 Table 3Table 3 summarizes mean coordinate repeatability in the north, east and up components of all 263 stations from their weekly combinations. All GOP Repro2 solution variants reached approximately 50% 264 and 25% of the lower mean RMS of coordinate repeatability when compared to the GOP Repro1/IGS08 265 solution in its horizontal and vertical components, respectively. These values represent even greater 266 improvements when compared to the GOP-Repro1/IGS05 solution. Comparing these two Repro1 267 solutions clearly demonstrates the beneficial impact of the new PCV models and reference frames. The 268 observed differences between Repro2 and Repro1 also indicate an overall improvement of the 269 processing software from V5.0 to V5.2, and the enhanced quality of global precise orbit and earth 270 orientation products.

271 Various GOP Repro2 solutions were also used to assess the selected models. Variants GO0 and GO1 272 differ in their mapping functions (GMF vs VMF1) used to project ZTDs into slant path delays. These 273 comparisons demonstrate that vertical component repeatability improved from 4.14 mm to 3.97 mm, 274 whereas horizontal component repeatability decreased slightly. By increasing the elevation cut-off 275 angle from 3° to 7° (GO2) and 10° (GO3), we observed a slight increase in RMS from repeatability of all 276 coordinates. This can be explained by the positive impact of low-elevation observations on the 277 decorrelation of height and tropospheric parameters, despite the fact that applied models (such as 278 elevation-dependent weighting, PCVs, multipath) are still not optimal for including observations at 279 very low elevation angles. On the other hand, it should be noted that the VMF1 mapping function is 280 particularly tuned to observations at 3° elevation angle which leads to biases at higher elevation angles, 281 Zus et al. (2015).

282 The GO4 solution represents an official GOP contribution to EUREF combined products. It is identical 283 to the variant GO1, but applies a non-tidal atmospheric loading. Steigenberger et al. (2009) discussed 284 the importance of applying non-tidal atmospheric loading corrections together with precise a priori 285 ZHD model. It has been concluded that using mean, or slowly varying, empirical pressure values for 286 estimating a priori ZHD instead of true pressure values results in a partial compensation of atmospheric 287 loading effects which is the case of GO1 solution. A positive 10% improvement in height repeatability 288 was observed for the GO4 solution. Our improvement was slightly lower than in a global scope 289 reported by Dach et al. (2011) with an improvement of 10-20% over all stations. As the effect depends on selected stations, a slightly higher impact in a global scale might be attributed to the stationdistribution, particularly differences in term of latitude and altitude.

No impact was observed from the higher-order ionospheric effects (GO4 vs. GO5) in term of coordinate repeatability. As the effect is systematic within the regional network (Fritsche et al., 2005) and it was mostly eliminated by using reference stations in the domains of interest. The combination of tropospheric horizontal gradients from 6-h to 24-h time resolution (GO4 vs. GO6), using the piecewise linear model, had a negligible impact on the repeatability of station coordinates too.

297 **5.2 Reference frame - residuals at fiducial stations**

298 The terrestrial reference frame (Altamimi et al., 2001) is a realization of a geocentric system of 299 coordinates used by space geodetic techniques. To avoid a degradation of GNSS products, differential 300 GNSS analysis methods require a proper referencing of the solution to the system applied in the 301 generation of precise GNSS orbit products. For this purpose, we often use the concept of fiducial 302 stations with precise coordinates well-known in the requested system. Such stations are used to define 303 the geodetic datum while their actual position can be re-adjusted by applying a condition minimizing 304 coordinate residuals. None station is able to guarantee a stable monumentation and unchanged 305 instrumentation during the whole reprocessing period. Thus a set of about 50 stations, with 100 and 306 more time periods for reference coordinates, was carefully prepared for datum definition in the GOP 307 reprocessing. An iterative procedure was applied then for every day by comparing a priori reference 308 coordinates with actually estimated ones and excluding fiducial station exceeding differences by 5, 5 309 and 15 mm in north, east and up components.

310 Figure 7 Figure 7 shows the evolution of the number of actually used fiducial stations (represented as 311 red dots) from all configured fiducial sites (represented as black dots) after applying an iterative 312 procedure of validation on a daily basis. This reprocessing began with the use of 16-20 fiducial stations 313 in 1996, and this number increased to reach a maximum of over 50 during the period from 2003-2011. 314 After 2011, this number decreased, due to a common loss of reference stations available from the last 315 realization of the global terrestrial reference frame without changes in its instrumentation. In most 316 cases, only 2 or 3 stations were excluded from the total number, however, this number is lower for 317 some daily solutions, indicating the removal of even more stations. The lowest number of fiducial sites 318 (12) was identified on day 209 of the year 1999 while, but low numbers were, generally, observed at 319 the beginning of the reprocessing period, in 1996. We observed consistent mean RMS errors for 320 horizontal, vertical, and total residuals of 6.47, 10.22, and 12.25 mm and 4.83, 7.94, 9.35 mm for daily 321 and weekly solutions, respectively, which demonstrate the stability of the reference system in the 322 reprocessing. The seasonality in height coordinate estimates characterized by the RMS of residuals 323 from the reference frame realization is dominated by errors due to modelling of the troposphere. We 324 believe, the main contribution stems from the insufficiencies in modelling of wet tropospheric delay, 325 as the effect has the most pronounced seasonal signal within the GNSS data analysis. Additionally, the 326 estimated station ZTD parameters and height are difficult to de-correlate. In the next section, the 327 strong seasonal variation in comparing zenith total delays estimated from GNSS and NWM data is 328 clearly visible.

329 5.3 Zenith total delays

330 We compared all reprocessed tropospheric parameters with respect to independent data from the 331 ERA-Interim global reanalysis (Dee et al. 2011) provided by the European Centre for Medium-Range 332 Weather Forecasts (ECMWF) from 1969 to the present. For the period of 1996-2014, we calculated 333 tropospheric parameters (namely ZTD and tropospheric horizontal linear gradients) from the NWM for 334 all EPN stations using the GFZ (German Research Centre for Geosciences) ray-tracing software (Zus et 335 al., 2014). The comparison of tropospheric parameters was performed by applying the linear 336 interpolation of GNSS parameters to the original NWM 6-hour representation, using the GOP TropoDB 337 (Győri and Douša, 2016). For monthly statistics discussed in this section, we applied an iterative 338 procedure for outlier detection using the 3-sigma criteria calculated from the compared ZTD or 339 gradient differences.

Table 4Table 4 summarizes comparisons of GNSS ZTDs, and tropospheric horizontal gradients, from all GOP processing variants with those obtained from the ERA-Interim. Mean biases and standard deviations were first calculated for each stations and each month and then mean and standard deviation of these values were computed, characterizing dispersions of all statistical values over the ensemble of stations.

345 The results in the table indicate a mean ZTD bias -1.8 mm for all comparisons (GNSS – NWM) suggesting 346 ZTDs achieved from the NWM reanalysis are drier than those obtained from GNSS reprocessing. Similar 347 biases have been observed for all other European GNSS re-processing products during the period of 348 1996-2014 (Pacione et al., 2017). On the other hand, when processing the ERA-Interim using two 349 different software and methodologies within the GNSS4SWEC Benchmark campaign (Dousa et al., 350 2016) during May and June of 2013 in Central Europe, and by their comparing to two GNSS reference 351 products based on different processing methods, we observed bias differences within ± 0.4 mm in ZTD. 352 As neither GNSS nor NWM is able to sense the troposphere with an absolute accuracy better than the 353 bias that we observed, we cannot make any conclusion, but its independence of the GNSS software. A 354 mixture of common processing aspects such as scope oscale of GNSS network, applied tropospheric 355 model, precise orbit product and others could still cause such a small biases in GNSS analysis at least.

Comparing the results of the official GOP Repro2 solution (GO4) to those of the legacy solution (GO0) demonstrates an overall improvement of 9% in term of accuracy, which corresponds to a similar comparison between the EUREF Repro1 and Repro2 products (Pacione et al., 2017). The improvement is assumed to be even larger (indicated by the coordinate repeatability) since the comparison of tropospheric parameters is , as the quality of ZTD retrievals limited by a lower quality of reference products derived from NWMare generally lower for NWM compared to GNSS from various intra /intertechnique comparisons data (Douša et al., 2016, Kačmařík et al., 2017, Bock and Nuret, 2009).

363 Comparing the GO1 and GO0 variants demonstrates that the VMF1 mapping function outperforms 364 GMF in term of standard deviation if the elevation cut-off angle of 3° is used. The change of mapping function together with the use of more accurate a priori ZHD, resulted in the ZTD standard deviation 365 366 improving from 8.8 mm (GO0) to 8.3 mm (GO1). However, bias was slightly increased which could be 367 partly attributed to the use of mean pressure model for a priori ZHD calculation and compensating 368 part of the non-tidal atmospheric loading (see Section 5.1). Using non-tidal atmospheric loading corrections along with precise modelling of a priori ZHD contributed to a small reduction of the bias 369 370 from -2.0 mm to -1.8 mm and, mainly, to the improvement by reducing this ZTD accuracy to 8.1 mm 371 (GO4). This corresponds with the previous assessment of the repeatability of station coordinates. 372 Degradation in ZTD precision was also observed when the elevation cut-off angle was raised from 3° 373 to 7° (GO2) or 10° (GO3). No impacts on ZTD were, however, visible neither from additional modelling of high-order ionospheric effects (GO5) nor from stacking of 6-hour horizontal gradients into dailypiecewise linear estimates (GO6).

376 Figure 8 Figure 8 displays the time series of statistics from comparisons of the GOP official ZTD product 377 (GO4) with respect to the results of the ERA-Interim reanalysis. Mean bias and standard deviation were 378 derived from the monthly statistics of the 6-hourly GNSS-ERA differences. A 1-sigma range of the mean 379 values, represented by error bars, are additionally derived from all stations on a monthly basis. 380 Although the time series show homogeneous results over the given time span, a small increase in the 381 mean standard deviation over time likely corresponds with increasing number of EPN sites, rising from 382 approximately 30 to 300. The early years (1996-2001) also display a worse overall agreement in 1-383 sigma range of mean values over all stations, which can be attributed to the varying quality of historical 384 observations and precise orbit products. The mean bias varies from -3 to 1 mm during the period of 385 1996-2014, with a long-term mean of -1.8 mm (Table 4Table 4). The long-term mean is also relatively 386 small compared to the ZTD mean 1-sigma range of 3-5 mm.

387 5.4 Tropospheric horizontal linear gradients

Additional GNSS signal delay due to the tropospheric gradients were developed by McMillan (1995). The complete tropospheric model for the line-of-sight delay (ΔD_T) using parameters zenith hydrostatic delay (ZHD), zenith wet delay (ZWD) and first-order horizontal tropospheric gradients G_N and G_E , all expressed in units of length, is described as follows

$$\Delta D_T = mf_h(e)ZHD + mf_w(e)ZWD + mf_e(e)\cot(e)[G_N\cos(A) + G_E\sin(A)]$$
(1)

where e and a are observation elevation and azimuth angles and mf_h , mf_w , mf_q are hydrostatic, wet 393 394 and gradient mapping functions representing the projection from an elevation to the zenith. Horizontal 395 gradients should optimally represent a ZTD change in a distances for north and east directions as it could be represented by terms $G_N \cot(e)$ and $G_E \cot(e)$ in the equation. However, the gradients need 396 397 to be parametrized practically with respect to observation elevation angle instead of the distance 398 theoretically applicable to the tropospheric effect at various elevation angles. The interpretation of 399 the tropospheric horizontal gradients in the Bernese software represents north and east components 400 of angle applied for the tilting the zenith direction in the mapping function with gradients representing 401 (in unit of length) the tilting angle multiplied by the delay in zenith (Meindl et al., 2004).

402 Similarly as in case of ZTD and coordinate assessment, Table 4 shows that tropospheric gradients 403 became worse when raising the elevation cut-off angle from 3° to 7° (GO2) or 10° (GO3). Mean 404 standard deviations of the GO2 and GO3 solutions increased by 8% and 12%, respectively, which is valid for the whole period of monthly time series (not showedshown). No significant differences in 405 temporal variations of mean biases of the north and east tropospheric gradients variants were 406 407 identified while they shared a higher variability during the years 1996-2001. No impact of modelling of 408 high-order ionospheric effects (GO5) was observed. Statistics of GO4 and GO6 solutions compared to 409 ERA-Interim revealed that standard deviations dropped from 0.38 mm to 0.28 mm and from 0.40 mm 410 to 0.29 mm for the east and north gradients, respectively. Worse performance of the GO4 solution is 411 attributed to the fact that tropospheric horizontal gradients were estimated with a 6-h sampling 412 interval using the piecewise linear model with applying practically no absolute or relative constraints. 413 In such cases, increased correlations of the gradients with other parameters can cause instabilities in processing certain stations at specific times; the gradients absorb some remaining errors in the GNSS 414

analysis model. The mean biases of the tropospheric gradients are considered to be negligible, but it
was demonstrated in Section 4 that some large systematic effects were indeed discovered and
attributed to the quality of GNSS signal tracking.

418 Figure 9-Figure 9 displays monthly time series of statistics from comparisons of the GNSS and NWM 419 tropospheric horizontal gradients in north and east directions. Two solutions are highlighted in order 420 to demonstrate the impact of different parameter temporal resolutions; a 6-hour resolution is used 421 for GO4 and a 24-hour resolution is used for GO6. Seasonal variations are mainly pronounced when 422 observing mean standard deviations (top plot), whereas gradual improvement is more pronounced for 423 mean biases (bottom plot). The reduction of the initial mean biases in horizontal gradients, and the 424 corresponding 1-sigma ranges over the values from the ensemble of stations, can be attributed to the 425 improved availability and quality of low elevation observation tracking. Elevation cut-off angles for collecting GNSS observations were initially configured station by station, ranging from 0° to 15°, until 426 427 2008 when the elevation cut-off angle 0° was recommended for all the stations.

428 Mean standard deviations and their 1-sigma ranges over all stations (Figure 9-Figure 9, top plot) are 429 lower by a factor of 1.3 for the solution with 24-hour resolution (GO6) compared to the 6-hour 430 resolution (GO4); the impact is also pronounced especially in the early years of the dataset. The 431 improvement factor ranges from 1.03 to 1.65 with the mean value of 1.35 overall stations and it is 432 usually higher for years before 2001. Theoretically, with 4 times more observations in GO6 the 433 standard deviation was expected to be divided by a factor of 2. This discrepancy indicates serialous 434 correlations in errors which are among others stemming from the errors in precise products and 435 models. Significant improvements, however, indicates possible correlations between tropospheric 436 gradients and other estimated parameters, such as ambiguities, height and zenith total delays, and 437 suggests a careful handling particularly when applying a sub-daily temporal resolution.

438 **5.5 Spatial and temporal ZTD analysis**

We performed spatial and temporal analyses of all processed variants in order to assess the impact of different settings on tropospheric products. Zenith tropospheric delays from all variants were compared in such a way to enable assessing impact of any single processing change: 1) GO1-GO0 for mapping function and more precise a priori ZHD model, 2) GO2-GO1 and GO3-GO1 for different elevation cut-off angle, 3) GO4-GO1 for non-tidal atmospheric corrections, 4) GO5-GO4 for higherorder ionospheric corrections and, 5) GO6-GO4 for temporal resolution tropospheric horizontal gradients. Station-specific behavior is out of this paper and-will be studied in future.

446 Geographical maps of spatially distributed biases and standard deviations in ZTDs from all compared 447 variants for the whole network are showed shown in Figure 10 Figure 10 and Figure 11 Figure 11. 448 Additionally median, minimum and maximum values of station-wise total statistics are provided in 449 Table 5Table 5. The comparisons demonstrated that the impact of the higher order effect is fully 450 negligible. Although overall mean biases in Table 5 are small, the GO1-GO0 comparison indicates a small negative bias over a majority of the stations, see Figure 10. Biases from the comparison of 451 452 variants with different elevation cut-off angles strongly indicates a station specific behavior with a 453 positive bias for stations around Poland, which has not been explained yet. According to the table and 454 Figure 11, the highest impact on standard deviations is found in the GO1 vs. GO0 solutions comparison. 455 The effect is latitude dependent and it follows the increasing magnitude of ZTDs towards the equator. 456 Detailed study illustrated in Figure 12 Figure 12, Figure 13 Figure 13 and Figure 14 Figure 14 then 457 illustrates ZTD statistics with respect to the station latitude, ellipsoidal height and time, respectively.

458 Using VMF1 mapping function together with precise a priori ZHD from VMF1 instead of the GMF and 459 GPT models, respectively, see GO1 vs. GO0, we observe biases ranging from -1.52 to 0.70 mm and the 460 median value -0.36 mm and, according to Table 5Table 5, with a moderate latitudinal dependence, see 461 Figure 12Figure 12. A similar, but slightly larger negative bias of -0.94±0.28 mm, was reported Kacmarik 462 et al. (2017) studying 400 stations in the central Europe. Standard deviations range from 0.69 mm to 463 3.82 mm in Table 5Table 5 with a profound-marked increase along with the latitude, Figure 12Figure 464 12, indicating the GPT performs worse at higher latitudes. This fully corresponds to the results from is 465 consistent with the paper by Steigenberger et al. (2009) demonstrating a partial compensation of the atmospheric loading effect by using the GPT model. In case the atmospheric loading effect is not 466 467 corrected for, the errors are mostly assimilated to the zenith total delay parameters if station 468 coordinates are fixed on a weekly basis. Additionally, Figure 14Figure 14 shows the standard deviation 469 grows with time which might be explained by increased number of can be attributed to the use of blind 470 (GMF) and actual-weather (VMF1) mapping functions. The mapping function affects an optimal use of 471 low-elevation observations with time, which were growing in EUREF permanent network with time-as 472 demonstrated for WTZR station in Figure 15Figure 15.

473 Biases obtained from the comparison of different elevation cut-off angles, i.e. variants 3°/7° (GO2-474 GO1) and 3°/10° (GO3-GO1), range from -0.81 mm to 1.66 mm and -2.22 mm to 2.66 mm, respectively, 475 and standard deviations from 0.15 mm to 1.29 mm and 0.31 to 2.04 mm, see Table 5 Table 5. Generally, 476 the impact of different elevation cut-off angle doesn't reveal any biases neither with respect to the 477 latitude (Figure 12Figure 12) nor the station height (Figure 13Figure 13). As expected, the impact is 478 larger for the GO3-GO1 differences and affected particularly some stations. Yearly biases exceeding 479 ±2.5 mm were identified for BELL, DENT, MLVL, MOPS, POLV RAMO and SBG2 stations. Temporal 480 dependences in the GO2-GO1 and GO3-GO1 comparisons, Figure 14Figure 14, show that the scatter 481 of station-specific biases steadily grows in time which is assumed to be related to the higher availability 482 of low-elevation observations. On the other hand, a small impact is observed for the standard deviation 483 compared to the other studied effects. This indicates the elevation cut-off angle affects mainly ZTD

484 biases, which has been also reported by Ning and Elgered (2012).

485 Table 5Table 5 shows that biases due to the non-tidal atmospheric loading (GO4-GO1) range from -486 2.29 mm to 5.55 mm, which is one of the largest impact compared to other comparison variants, and 487 standard deviations range from 0.68 mm to 4.72 mm that represents the second largest impact 488 compared to all other variants. Standard deviation larger than 3 mm was observed at some stations, 489 such as JOZE, MAD2, MADR, MDVO, MOPI, NYAL, SBG2, VENE and WETT. It should be emphasized this 490 comparison reflects differences due to the modelling of atmospheric loading corrections in GO4 and, 491 a partial compensation of the loading effect by zenith tropospheric delay estimates in the GO1 solution 492 variant. The differences are strongly station-dependent, but did not reveal any dependence on 493 latitude, see Figure 12Figure 12. It shows, however, some degradation in standard deviation during 494 the first years of the reprocessing, see Figure 14Figure 14. Since a similar degradation has not been 495 observed for other comparison variants, it can be related to the quality of pressure data used to 496 compute atmospheric loading.

497 The impact of higher-order ionospheric effect (GO5-GO4) is negligible at all stations demonstrating total statistics for all stations within ±0.3 mm when applying the y-range about 10 times smaller than 498 499 in other panels of Figure 12 Figure 12, Figure 13 Figure 13 and Figure 14 Figure 14. A strong latitudinal 500 dependence is, however, clearly visible in Figure 12 Figure 12 as well as a temporal variability showing 501 vearly statisticpeaks up to ±0.4 mm, Figure 14Figure 14. Both dependences are due to the changing 502 magnitude of ionospheric corrections, generally increasing towards the equator and a daily noon, and 503 along with quasi periodic cycles of the solar magnetic activity, reaching peaks around years 2001 and 504 2014.

505 The impact of stacking tropospheric gradients from 6-hour to daily estimates (GO6-GO4) is almost 506 negligible in term of biases which stay below ±1 mm, Table 5Table 5 and Figure 10Figure 10. However, 507 standard deviations range from 0.76 mm to 2.46 mm and grow towards the equator, Figure 12Figure 508 12. That can be certainly attributed to the more difficult modelling of a local asymmetry in the 509 troposphere, which is generally increasing together with the increasing of the water vapor content. 510 There is no significant temporal variation observed in bias, in Figure 14, but a small decrease in 511 standard deviation (Figure 14). It can be attributed to a higher stability of the gradient estimates with 512 time, see Figure 9-Figure 9, when supported with increased number of available low-elevation 513 observations.

6 Impact of variants on long-term ZTD trend estimates

515 We assessed the impact of solution variant on long-term ZTD trend estimates by analysing 172 EUREF 516 stations providing the time-series of data longer than 10 years. For each station, the trend analysis was 517 performed without any data homogenization or outlier rejection as our focus was only on assessing 518 the impact of solution variants on the trend estimates. The ZTD trends were estimated using the least 519 squares regression method applied on model (Weatherhead et al., 1998)

$$Y_t = \mu + \beta X_t + S_t + \mathcal{E}_t \tag{2}$$

521 where μ is the constant term of the model, βX_t is the linear trend function with β representing the 522 trend magnitude, S_t represents the term modelled by the sine wave function of time X_t including 523 annual, 2nd harmonics and daily variations, and finally ε_t is the noise in the data.

524 Site-by-site estimated ZTD trends from all the variants are provided in supplementary materials 525 completed by time-span information, number of records and estimated mean formal errors calculated 526 over all variants. In total, trends range from -0.99 to 0.96 mm/year. Although the individual station 527 trend provided in supplements could be compared to other studies, e.g. Baldysz et al. (2016), Klos et 528 al. (2016) or Nilsson and Elgered (2008), however, it should be strongly emphasized here that our 529 trends are estimated without any preceding time-series homogenization and the formal errors of the 530 trend estimates are underestimated by a factor 2-4 (Nilsson and Elgered, 2008).

<u>Table 6</u> summarizes the statistics of estimated trend differences at all 172 stations, always
 between particular variants as defined in Section 5.5. Interestingly, the most significant impact is
 observed due to the non-tidal atmospheric loading effects reaching differences up tobelow ±0.55
 mm/year in ZTD trends for some extreme cases from the ensemble of 172 stations, and an overall <u>1-</u>

535 sigma scatter of 0.50 mm/year from the ensemble of stations. Changes in elevation cut-off angle, 536 particularly from 3° to 10°, reveal also a significant impact characterized by differences up tobelow 537 ± 0.34 mm/year and the scatter of 0.32 mm/year. The impact of mapping function on trend estimates 538 remains small, with a maximum difference of 0.12 mm/year and the 1-sigma scatter below 0.08 539 mm/year, while other strategy changes, due to time resolution of tropospheric gradients and higher-540 order ionospheric effects, remains negligible, always below ±0.04 mm/year for all 172 stations, with 541 the scatter of the same magnitude. All mean biases over differences stay also below 0.05 mm/year. 542 These results are consistent with a study performed by Ning and Elgered (2012) spaning a broader 543 span of cutoff angles. They demonstrated a significant impact of this parameter on Integrated Water 544 Vapor trend estimates.

Finally, we selected 12 stations-optimally available over the entire 2nd re-processing period-and a. All 545 estimated trends are displayed in Figure 16-Figure 16, Trends for 12 stations range- ranging from -0.05 546 to 0.38 mm/year. with formal errors of 0.02-0.04 mm/year. It should be noted the formal errors are 547 548 underestimated by a factor of 2-4 because the noise is assumed white (Nilsson and Elgered (2008). 549 Consistent with the overall results reported in Table 5For the 12 selected stations, the most significant 550 impact for the selected 12 stations is observed in the change of elevation cut-off angle (GO2, GO3 vs. 551 GO1) and atmospheric loading (GO4 vs. GO1) when reaching differences up to 0.1 mm/year in 552 estimated ZTD trends. A similar, but more extensive study, was performed by Ning and Elgered (2012) 553 for Integrated Water Vapor content (IWV), roughly equal to 1/6 (ZTD-ZHD) kg.m⁻², using larger 554 differences in the elevation cut off angle and obtaining highly sensitive results in term of estimated 555 IWV trends. Impacts of other strategies are generally below 0.05 mm/year – variants GO4, GO5, and 556 GO6 are very similar, but not consistent again with GO1, meaning the non-tidal atmospheric loading 557 has a significant impact on trend estimates for selected stations with the longest data time-series.

558 **7** Conclusions

559 In this paper, we present results of the new GOP reanalysis of all stations within the EUREF Permanent network during the period of 1996-2014. This reanalysis was completed during the 2nd EUREF 560 reprocessing to support the realization of a new European terrestrial reference system. In the 2nd 561 562 reprocessing, we focused on analysing a new product – GNSS tropospheric parameter time-series for 563 applications to climate research. To achieve this goal, we improved our strategy for combining 564 tropospheric parameters at midnights and at transitions in GPS weeks. We also performed seven 565 solution variants to study optimal troposphere modelling; we assessed each of these variants in terms 566 of their coordinate repeatability by using internal evaluations of the applied models and strategies. We 567 also compared tropospheric ZTD and tropospheric horizontal gradients with independent evaluations 568 obtained by numerical weather reanalysis via the ERA-Interim.

Results of the GOP Repro2 yielded improvements of approximately 50% and 25% for their horizontal and vertical component repeatability, respectively, when compared to those of the GOP Repro1 solution. Vertical repeatability was reduced from 4.14 mm to 3.73 mm when using the VMF1 mapping function, a priori ZHD, and non-tidal atmospheric loading corrections from actual weather data. Increasing the elevation cut-off angle from 3° to 7°/10° increased RMS errors of residuals from these coordinates' repeatability. All of these factors were also confirmed by the independent assessment of tropospheric parameters using NWM reanalysis data. 576 We particularly recommend using low-elevation observations along with the VMF1 mapping function,

- as well as using precise a priori ZHD values together with the consistent model of non-tidal atmospheric
- 578 loading. While estimating tropospheric horizontal linear gradients improves coordinates' repeatability,
- 579 6-hour sampling without any absolute or relative constraints revealed a loss of stability due to their 580 correlations with other parameters. On the other hand, 24-h piecewise linear gradients did not indicate
- 581 a worse repeatability of coordinates estimates. For saving the time needed for the processing of 4
- times less gradient parameters, we could recommend as sufficient using unconstrained 24-h piecewise
- 583 model for the first-order tropospheric asymmetry.

584 The impact of processing variants on long-term ZTD trend estimates was assessed at 172 EUREF 585 stations with time-series longer than 10 years. The most significant impact was observed due to the 586 non-tidal atmospheric loading effect reaching differences up tobelow ±0.55 mm/year in ZTD trends for 587 some extreme cases from the ensemble of 172 stations. Changes in elevation cut-off angle, particularly 588 from 3° to 10°, revealed also a significant impact reaching differences up tobelow ±0.35 mm/year. The 589 change of mapping function was observed rather small, with a maximum difference of 0.12 mm/year, 590 while other strategy changes, due to time resolution of tropospheric gradients and higher-order 591 ionospheric effects, remained negligible, always below ±0.04 mm/year for all 172 stations.

592 Assessing the tropospheric horizontal gradients with respect to the ERA-Interim reanalysis data 593 revealed some long-term systematic behaviour linked to degradation in antenna tracking quality. We 594 presented an extreme case at the Mallorca station (MALL), in which gradients systematically increased 595 up to 5 mm from 2003-2008 while pointing in the direction of prevailing observations at low elevation 596 angles. However, these biases disappeared when the malfunctioning antenna was replaced. More 597 cases similar to this, although less extreme, have indicated that estimated tropospheric gradients are 598 extremely sensitive to the quality of GNSS antenna tracking, thus suggesting that these gradients can 599 be used to identify problems with GNSS data tracking in historical archives.

One of the main difficulties faced during the 2nd reprocessing was that of the quality of the historical 600 601 data, which contains a large variety of problems. We removed data that caused significant problems 602 in network processing when these could not be pre-eliminated from normal equations during the 603 combination process without still affecting daily solutions. To provide high-accuracy, high-resolution 604 GNSS tropospheric products, the elimination of such problematic data or stations is even more critical 605 considering the targeting static coordinates on a daily or weekly basis for the maintenance of the reference frame or the derivation of a velocity field. Before undertaking the 3rd EUREF reprocessing, 606 607 which is expected to begin after significant improvements have been made to state-of-the-art models, 608 products and software, we need to improve data quality control and clean the EUREF historical archive 609 in order to optimize any future reprocessing efforts and to increase the quality of tropospheric 610 products. These efforts should also include the collection and documentation of all available information from each step of the 2nd EUREF reprocessing, including individual contributions, EUREF 611 612 combinations, time-series analyses and coordinates, and independent evaluations of tropospheric 613 parameters.

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730 Table 1: Characteristics of GOP reprocessing models

Processing options	Description						
Products	CODE precise orbit and earth rotation parameters from the 2^{nd}						
	reprocessing.						
Observations	Dual-frequency code and phase GPS observations from L1 and L2						
	carriers. Elevation cut-off angle 3°, elevation-dependent weighting						
	1/cos ² (zenith), double-difference observations and with 3-minute						
	sampling rate.						
Reference frame	IGb08 realization, core stations set as fiducial after a consistency						
	checking. Coordinates estimated using a minimum constraint.						
Antenna model	GOP: IGS08_1832 model (receiver and satellite phase centre offsets						
	and variations).						
Troposphere	A priori zenith hydrostatic delay/mapping function: GPT/GMFh						
	(GO0) and VMF1/VMF1h (GO1-GO6). Estimated ZWD corrections						
	every hour using VMF1 wet mapping function; 5 m and 1 m for						
	absolute and relative constraints, respectively. Estimated horizontal						
	NS and EW tropospheric gradients every 6 hours (GO0-GO5) or 24						
	hours (GO6) without a priori tropospheric gradients and constraints.						
Ionosphere	Eliminated using ionosphere-free linear combination (GO0-GO6).						
	Applying higher-order effects estimated using CODE global						
	ionosphere product (GO5).						
Loading effects	Atmospheric tidal loading and hydrology loading not applied. Ocean						
tidal loading FES2004 used. Non-tidal atmospheric							
	introduced in advanced variants from the model from TU-Vienna						
	(GO4-GO6).						

732 Table 2: GOP solution variants for the assessment of selected models and settings

Solution ID	Specific settings and differences	Remarks and rationales			
GO0	GMF and 3° cut-off	Legacy solution for Repro1			
GO1	VMF1 and 3° cut-off	New candidate for Repro2			
GO2	=GO1; 7° cut-off	Impact of elevation cut-off angle			
GO3	=GO1; 10° cut-off	Impact of elevation cut-off angle			
GO4	=GO1; atmospheric loading	Non-tidal atmospheric loading applied			
GO5	=GO4; higher-order ionosphere	Higher-order ionosphere effect not applied			
GO6	=GO4; 24-hour gradients	Stacking tropospheric gradients to 24-hour			
		sampling			

Solution	North RMS	East RMS	Up RMS
	[mm]	[mm]	[mm]
GOP-Repro1/IGS05	3.01	2.40	5.08
GOP-Repro1/IGS08	2.64	2.21	4.94
GO0	1.20	1.30	4.14
G01	1.23	1.33	3.97
GO2	1.24	1.33	4.01
GO3	1.26	1.34	4.07
GO4	1.14	1.24	3.73
GO5	1.14	1.24	3.73
GO6	1.14	1.24	3.73

Table 3: Comparison of GOP solution variants for north, east and up coordinate repeatability.

736 737 738 Table 4: Statistics (bias and standard deviations) of ZTD and tropospheric gradients from the seven reprocessing variants

compared to those obtained from the ERA-Interim NWM reanalysis. In addition to the statistics, 1-sigma range over ensemble of stations is provided.

Solution	ZTD bias [mm]	ZTD sdev [mm]	EGRD bias [mm]	EGRD sdev [mm]	NGRD bias [mm]	NGRD sdev [mm]
GO0	-1.5 ± 2.1	8.8 ± 2.0	-0.04 ± 0.08	0.39 ± 0.10	+0.01 ± 0.09	0.43 ± 0.12
GO1	-2.0 ± 2.1	8.3 ± 2.2	-0.04 ± 0.08	0.39 ± 0.10	+0.01 ± 0.09	0.42 ± 0.13
GO2	-1.9 ± 2.2	8.4 ± 2.2	-0.05 ± 0.10	0.41 ± 0.10	+0.00 ± 0.12	0.45 ± 0.12
GO3	-1.8 ± 2.3	8.5 ± 2.1	-0.08 ± 0.13	0.43 ± 0.11	-0.01 ± 0.14	0.49 ± 0.12
GO4	-1.8 ± 2.4	8.1 ± 2.1	-0.04 ± 0.09	0.38 ± 0.10	+0.00 ± 0.09	0.40 ± 0.12
GO5	-1.8 ± 2.4	8.1 ± 2.1	-0.05 ± 0.09	0.38 ± 0.10	+0.01 ± 0.08	0.40 ± 0.12
GO6	-1.8 ± 2.4	8.2 ± 2.1	-0.04 ± 0.08	0.29 ± 0.06	+0.01 ± 0.09	0.28 ± 0.06

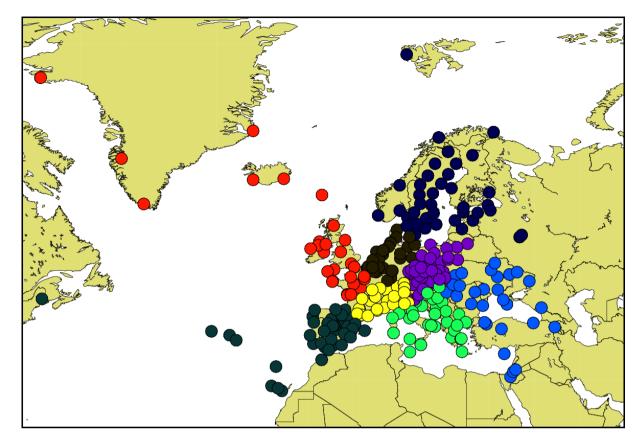
740 741 Table 5: Median, minimum (min) and maximum (max) values of total ZTD biases and standard deviation (sdev) over all

stations. Units are millimetres.

Compared	ZTD bias	ZTD bias	ZTD bias	ZTD sdev	ZTD sdev	ZTD sdev
variants	median	min	max	median	min	max
G01-G00	-0.36	-1.52	+0.70	2.01	0.69	3.82
G02-G01	+0.03	-0.81	+1.66	0.66	0.15	1.29
G03-G01	+0.03	-2.22	+2.66	1.10	0.31	2.04
G04-G01	+0.05	-3.29	+5.55	1.37	0.68	4.72
G05-G04	-0.02	-0.31	+0.07	0.07	0.04	0.30
G06-G04	-0.02	-0.23	+0.16	1.24	0.76	2.46

743 Table 6: Mean statistics of ZTD trends differences estimated between variants for 172 stations. Units are millimetres/year.

Statistics	GO1-GO0	GO2-GO1	GO3-GO1	GO4-GO1	GO5-GO4	GO6-GO4
Min	-0.118	-0.141	-0.308	-0.547	-0.017	-0.038
Max	0.045	0.179	0.331	0.452	0.031	0.036
mean	0.036	0.018	0.012	-0.048	0.007	0.001
Sdev	0.081	0.160	0.319	0.499	0.024	0.037





747 Figure 1: EUREF Permanent Network's clusters (designated by different colours) in the 2nd GOP reprocessing.

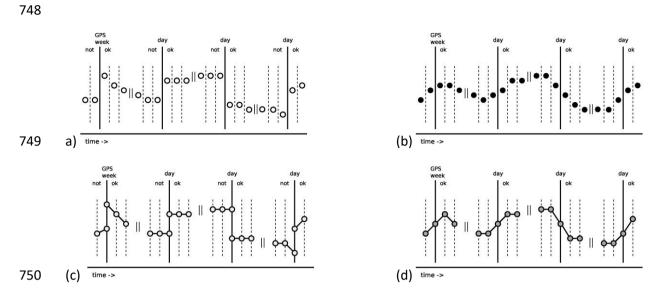


Figure 2: Charts of 4 variations on representations of tropospheric parameters. Right (b), (d) and left (a), (c) panels

751 752 753 display estimates made with and without midnight combinations, respectively. Top (a), (b) and bottom (c), (d) panels display the piecewise constant and the linear model, respectively.

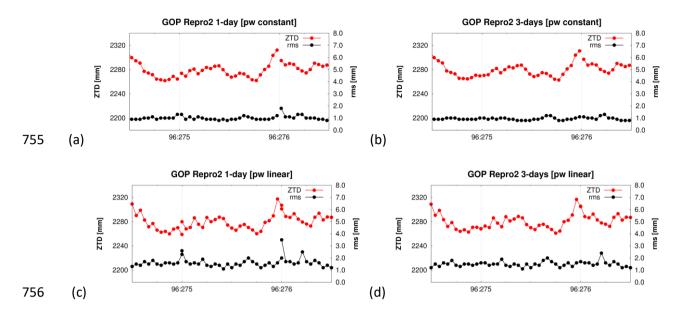


Figure 3: Four variations in representation of tropospheric parameters. Right (b), (d) and left (a), (c) panels display estimates with and without midnight combinations, respectively. Top (a), (b) and

bottom (c), (d) panels display the piecewise constant and the piecewise linear model, respectively.

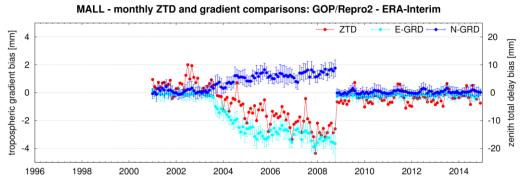


Figure 4: MALL station - monthly mean differences in tropospheric horizontal gradients with respect to the ERA-Interim.

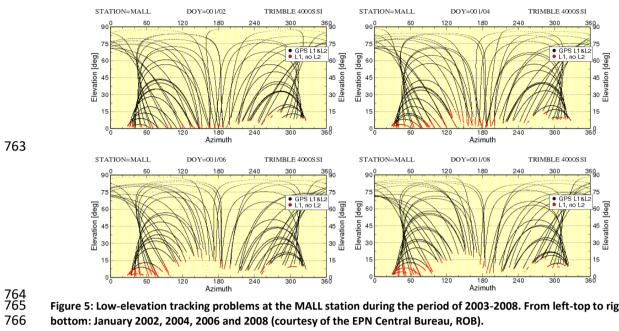
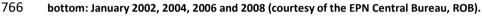


Figure 5: Low-elevation tracking problems at the MALL station during the period of 2003-2008. From left-top to right-



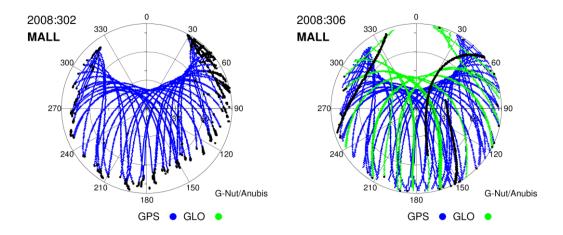


Figure 6: Sky plots before (left) and after (right) replacing the malfunctioning antenna at the MALL site (Oct 30, 2008).
 Black dots indicates single-frequency observations available only.

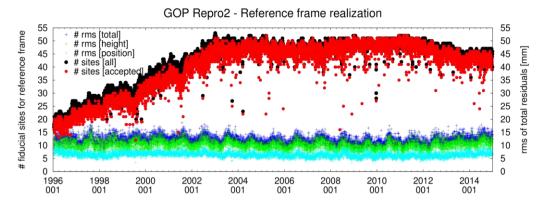


Figure 7: Statistics of the daily reference system realization: a) RMS of residuals at fiducial stations (representing the
 total, height and position); b) number of stations (all and accepted after an iterative control)

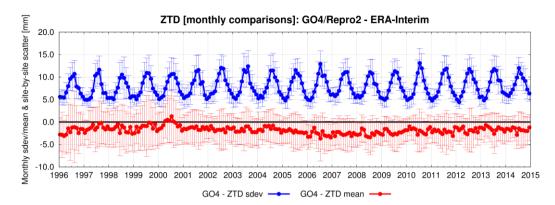
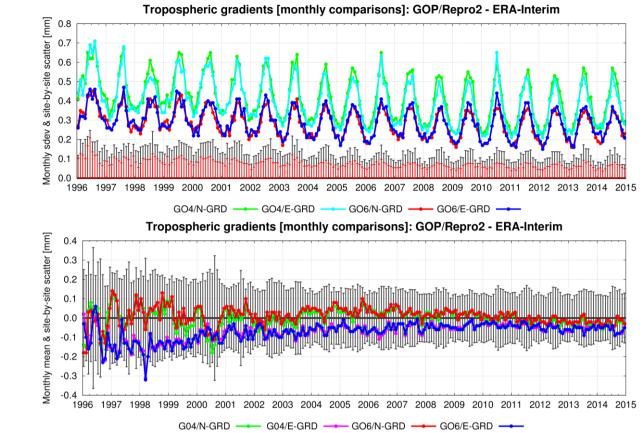


Figure 8: Monthly means of bias and standard deviation of official GOP ZTD product compared to those of the ERA-Interim.
 Error bars indicate standard errors of mean values over all compared stations.



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Figure 9: Monthly means of bias and standard deviation of tropospheric horizontal north (N-GRD) and east (E-GRD) 782 gradients compared to those obtained by ERA-Interim. Note: Similar products are almost superposed. Error bars indicate 783 standard errors of mean values over all compared stations plotted from the zero y-axis to emphasise seasonal variations

784 and trends. Error bars are displayed for north gradients only, however, being representative for the east gradients too.

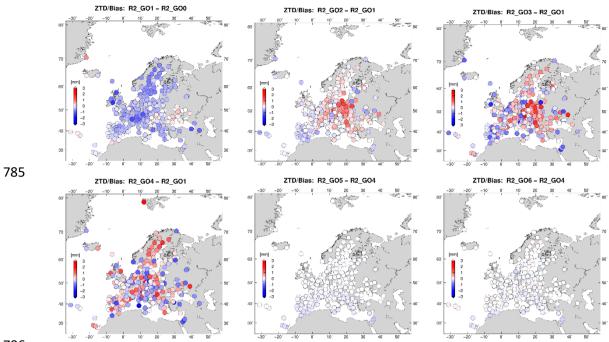




Figure 10: Geographic visualization of biases from inter-comparisons of GOP 2nd reprocessing variants.

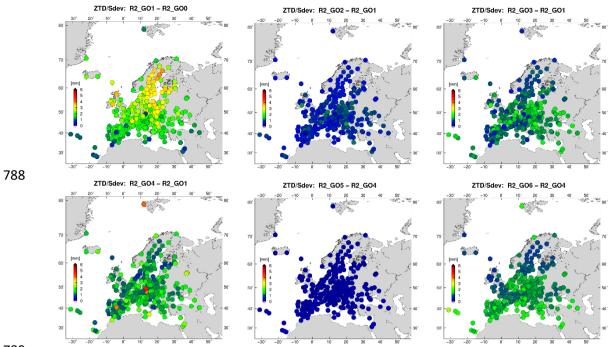


Figure 11: Geographic visualization of standard deviations from inter-comparisons of GOP 2nd reprocessing variants.

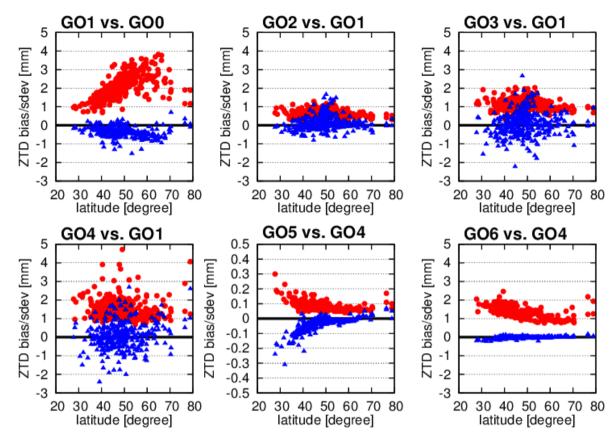


Figure 12: Dependence of ZTD biases (blue) and standard deviations (red) from inter-comparisons of GOP 2nd

reprocessing solution variants on station latitude. Note different y-range for the GO5 vs. GO4 comparison.

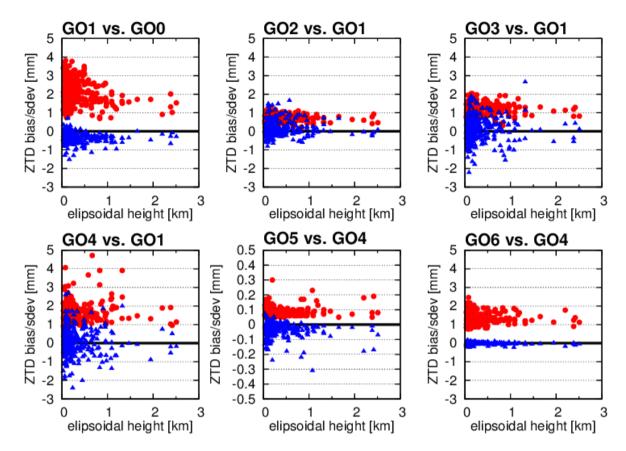
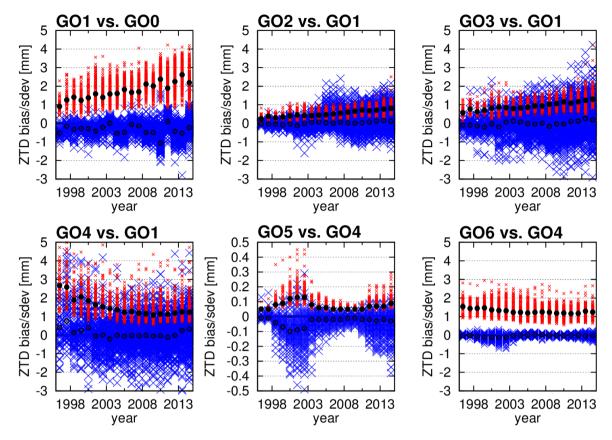


Figure 13: Dependence of ZTD biases (blue) and standard deviations (red) from inter-comparisons of GOP 2nd reprocessing solution variants on station ellipsoidal height. Note different y-range for the GO5 vs. GO4 comparison.



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Figure 14: Dependence of ZTD biases (blue), mean biases (unfilled black circles), standard deviations (red) and mean

800 standard deviations (filled black circles) from inter-comparisons of GOP 2nd reprocessing solution variants on year. Note 801 different y-range for the GO5 vs. GO4 comparison.

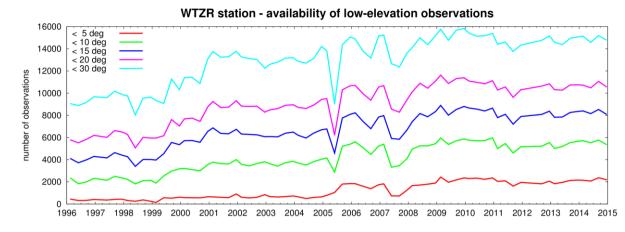
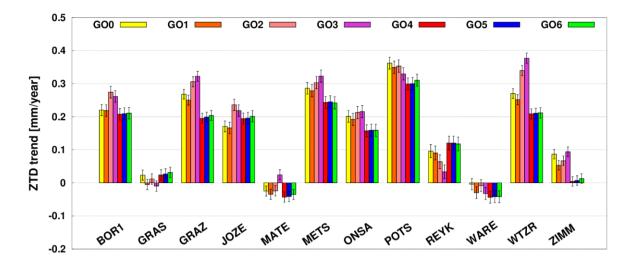


Figure 15: Availability of observations at low-elevation angles (below 5°, 10°, 15°, 20° and 30°) for WTZR station.





807 Figure 16: Long-term ZTD trend estimates and their formal errors (error bars) for all processing variants