



An aircraft gas chromatograph-mass spectrometer System for Organic Fast Identification Analysis (SOFIA): design, performance and a case study of Asian monsoon pollution outflow

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10 **Abstract.** Volatile organic compounds (VOC) are important for global air quality and oxidation processes in the troposphere. In addition to ground-based measurements, the chemical evolution of such species during transport can be studied by performing in-situ airborne measurements. Generally, aircraft instrumentation needs to be sensitive, robust and sample at higher frequency than ground based systems while their construction must comply with rigorous mechanical and electrical safety standards. Here, we present a new System for Organic Fast
15 Identification Analysis (SOFIA), which is a custom built fast Gas Chromatography – Mass Spectrometry (GC-MS) system with a time resolution of 2-3 min. The relatively high time resolution is the result of a novel cryogenic pre-concentration unit which rapidly cools ($\sim 6^{\circ}\text{C/s}$) the sample enrichment traps to -140°C , and a new chromatographic oven designed for rapid cooling rates ($\sim 30^{\circ}\text{C/s}$) and subsequent thermal stabilization. SOFIA was installed in the High Altitude and Long Range Research Aircraft (HALO) for the Oxidation Mechanism
20 Observations (OMO) campaign in August 2015, aimed at investigating the Asian monsoon pollution outflow in the tropical upper troposphere. In addition to a comprehensive instrument characterization we present an example monsoon plume crossing flight as a case study to demonstrate the instrument capability. Hydro- and halocarbon data from SOFIA are compared with mixing ratios of carbon monoxide (CO) and methane (CH_4), used to define the pollution plume. By using excess (ExR) and normalized excess mixing ratios (NEMRs) the pollution could
25 be attributed to two air masses of distinctly different origin, identified by back-trajectory analysis. This work endorses the use of SOFIA for aircraft operation and demonstrates the value of relatively high-frequency, multicomponent measurements in atmospheric chemistry research.

1 Introduction

30 Despite their generally low ambient concentrations, organic trace gases can have significant impacts on atmospheric chemistry (Williams, 2004). Halogenated organic compounds are capable of destroying both tropospheric and stratospheric ozone (Molina and Rowland, 1974; Read et al., 2008; Saiz-Lopez et al., 2012; Wang et al., 2015), and thus altering the oxidative capacity of the atmosphere (Parrella et al., 2012; Baker et al., 2016), with prominent examples being the chlorofluorocarbons (CFCs) and the hydrochlorofluorocarbons (HCFCs). Their industrial replacements, i.e. hydrofluorocarbons (HFCs), may have lower stratospheric ozone
35 depletion potentials but they can act as potent greenhouse gases (Velders et al., 2009). In addition, oxygenated volatile organic compounds (OVOCs) and simple hydrocarbons (HCs), whose emissions are associated with both biogenic and anthropogenic sources, play a central role in the production and destruction of key atmospheric oxidants such as the hydroxyl radical (OH) (Kley et. al., 1997; Atkinson, 2000; Monks et. al., 2005; Lelieveld



et al., 2016) and ozone (Pusede and Cohen, 2012). It is therefore essential to monitor their atmospheric abundance and understand their source-sink dynamics particularly in the chemically sensitive but poorly accessible region of the upper troposphere and lower stratosphere.

For over sixty years, organic trace gases in the atmosphere have been measured by acquiring air samples in the field, using pressurized metal/glass containers or on adsorbent filled tubes, and subsequently measuring them “off-line” in the laboratory. A commonly used analytical technique for such samples has been gas chromatography (GC), coupled to detectors such as the mass spectrometer (MS), the flame ionization detector (FID) and the electron capture detector (ECD) (e.g. Haagen-Smit et al., 1953; Colman et al., 2001; Colomb et al., 2006; Williams et al., 2007; Pollmann et al., 2008; Lerner et al., 2017). While this approach allows for straightforward sampling in remote regions, and in the upper troposphere using research aircraft or balloons, the sample frequency is limited by the number of containers available and it is prone to interference from surface related artefacts (especially for the most reactive species (Plass-Dülmer et al., 2006)). To achieve higher time resolution and hence data density in atmospheric measurements, several “on-line” instrumental techniques have been developed based on ionization mass spectrometry (e.g. Arnold and Hauck, 1985; Crutzen et al., 2000; Sprung et al., 2001; Blake et al., 2006; Le Breton et al., 2012). These methods use positively or negatively charged species to ionize the substances of interest prior to their detection in a mass spectrometer. While these techniques enable high resolution sampling (typically 1s – 1min), as they do not require a pre-concentration or separation, they suffer from lower molecular specificity compared to GC-MS, as multiple species may contribute to a given mass signal. Therefore, provided it can be made to measure fast enough, gas chromatography (GC) coupled with a mass spectrometer (MS) can be an extremely powerful tool in atmospheric research.

In recent years, there has been considerable focus on optimizing GC analysis for speed (Mastovska and Lehotay, 2003). This has led to a new generation of fast GC systems that can be applied in airborne research since they combine high time resolution and mass selectivity. Apel et al. (2003) developed the first fast GC-MS system for airborne measurements of VOCs and halocarbons (Trace Organic Gas Analyzer (TOGA); see also Apel 2016, Hornbrook et al., 2011; Apel et al., 2010, 2012). TOGA uses a custom built liquid nitrogen (LN₂) fueled system to cool sample enrichment traps down to -130°C (Apel et al., 2016) enabling a sample every 2-5min. To our knowledge, three other fast GC systems have since been designed for aircraft operation: GhostMS (Sala et al., 2014), HCG (Jäger et al. 2014) and μ Dirac (Gostlow et al., 2010). All three do not require the use of a liquid cryogen and have been successfully operated in aircraft, Zeppelin or balloon campaigns. The use of LN₂ for cryogenic trapping has both advantages and disadvantages that are discussed in section 4.

The objective of this paper is to present a comprehensive description of the new fast GC-MS instrument, called System for Organic Fast Identification Analysis (SOFIA), designed for airborne measurements. In section 2 a detailed description of the major components and sampling characteristics is provided along with the instrument specifications, as defined under laboratory and field conditions. In section 3 we demonstrate the capabilities of SOFIA by presenting an example from the OMO aircraft campaign, flying over the Arabian Peninsula in August 2015 to intercept pollution plumes convected to the upper atmosphere in the Indian monsoon system. We demonstrate the usefulness of fast GC-MS measurements by studying the mixing ratios of selected chloromethanes (chloromethane (CHCl₃), dichloromethane (CH₂Cl₂), chloroform (CHCl₃) and carbon tetrachloride (CCl₄)), hydrocarbons (isoprene (C₅H₈), benzene (C₆H₆) and toluene (C₇H₈)), OVOCs (propanal (C₃H₆O) and acetone (C₃H₆O)) and sulphur containing species (carbon disulphide (CS₂)) in a pollution plume, and discuss the technical and scientific implications.



2 Material and methods

The system described below has been designed and constructed for use on-board the German High Altitude Long Range aircraft (HALO), a Gulfstream G550 aircraft, operated by the German Aerospace Organization (DLR). Operation on-board a high flying aircraft such as HALO requires that the analytical system can perform reliably over the widely varying conditions of external temperature (40°C to -60°C) and pressure (1000mb – 200mb). In addition, the internal cabin conditions also change and all relevant parameters must be measured for subsequent performance assessments. In the case of the SOFIA system, an additional challenge arises through the use of a liquid nitrogen (LN₂) cryogenic system, because the LN₂ evaporation rate in the container is strongly pressure and temperature dependent, thus safe and accurate pressure control within the cryogenic system is essential. The following sections present a detailed description of the major components of the SOFIA GC-MS system. They comprise of a sampling system overview (section 2.1), a detailed description of the cryogenic trapping method (section 2.2) and chromatographic GC oven (section 2.3), a characterization of the detection and quantification unit using standard gas experiments (section 2.4), and a description of the process control and software (section 2.5). The last part (section 2.6) reports the analytical specifications during the example flight presented in the “case study” section, which was part of the first field deployment of the instrument.

2.1 Sampling overview

A schematic overview of the SOFIA system is presented in **Fig. 1**. Air is drawn by the inlet pump (described in section 2.1.1) through the sampling line which extends outside the aircraft fuselage. The sample pre-concentration then follows in three stages (described in section 2.2). First, the air is pumped through the cooled water trap where H₂O is selectively removed from the airstream and then the low temperature enrichment trap where VOC molecules are retained. During sampling, a flow controller (**Fig. 1**; SampleFC) is used to regulate the sampling flow, and the pressure change inside a sampling volume unit (**Fig. 1**; Calibrated volume) is used for accurate sample volume determination. In the second stage, the enrichment trap is heated and the volatiles are transferred to the cryofocus trap which is a narrow (0.25mm) inert chromatography column with low volume to optimise the subsequent injection. In the third stage, the cryofocus trap is heated rapidly and the sample is injected into the chromatography column housed in the oven (denoted GC in **Fig. 1**) where it is separated prior to ionization and detection of the peak sequence by the mass spectrometer (denoted MS in **Fig. 1**).

2.1.1 Inlet system

Air was drawn into the aircraft through a rearward facing Trace Gas Inlet (TGI; Enviroscope GmbH) (Wendisch et. al., 2016). In the cabin, a heated Teflon line (¼" (0.635cm), 2m length, 40°C) brings the air in the inlet MuPo (**Fig. 1**). The inlet system of SOFIA was designed to compress ambient air to standard pressure at a minimum flow rate of 200scm in ~15km altitude. Under such conditions a maximum residence time of 15sec can be maintained at the highest flight altitudes in the flow loop from the aircraft inlet through the zero and calibration system to the sampling point. The compression ratio of single stage all-PFA Teflon diaphragm pumps was not sufficient to fulfil our requirements, and dual stage pumps would push the overall system beyond the weight and size limits. Therefore, three small identical pumps (NMP 850KNDC, KNF Global strategies AG) were run on the low pressure inlet in parallel, backed by a fourth pump connected to the high pressure exit upstream of the sampling system.

Since even small leaks in the low pressure part of the inlet system can cause interferences, the pumps' tightness was tested prior to integration. It was found that the pump heads, as sold, leaked severely across the edge of the



diaphragm. This issue was solved by fitting chronometer O-rings around the diaphragm of the individual pumps. The pressure in the sampling system was maintained by a mechanical diaphragm pressure controller downstream of the sampling point (see 2.2.4).

2.1.2 Zero / Calibration unit

5 Directly downstream of the inlet system, ambient air was either sent directly to the sampling loop or through an oxidizing catalyst tube, depending on the 3-way zero valve (**Fig. 1**; ZV) position. The platinum bead packed catalyzer was maintained at 350°C and the outflow used as VOC free air, since under these conditions all VOC measured were converted entirely to CO₂. The catalyzer does not significantly affect ambient water concentrations and so keeps sample humidity in calibration and zero modes the same as ambient conditions. In
10 calibration mode, the calibration valve (**Fig. 1**; CV) was activated to allow a multicomponent calibration gas mixture (Apel-Riemer Environmental, Inc., USA) to enter via a 5ml flow controller (Bronkhorst, Germany) into the line exiting the unit.

2.1.3 Multi-position valves (MuPos)

During operation, two multiple position valves (**Fig 1**; Source MuPo, Traps MuPo; VICI, Germany) were
15 operated in concert. The Source MuPo is a 4-port valve that switches gaseous input either to a zero/calibration unit, a helium line or to a plug. The plugged position is used for diagnostic purposes, in particular for leak tests during flight preparation. When running continuously, the Source MuPo valve was set to helium source except during sampling mode. The Traps MuPo is a 6-port valve that connects the water, enrichment, and focus traps. The Traps MuPo valve switches between SAMPLE and INJECT modes. In the SAMPLE position, preparation
20 of the next sampling cycle and sampling is done in parallel, with the last sample driven through the isolated cryofocus trap to the GC column and the detector. In INJECT position, the enrichment part of the sampling loop is connected to the injection loop to transfer the sample to the cryofocus trap, while the water trap is purged in parallel to vent.

2.1.4 The sampling volume unit

25 The sample volume measurement unit is connected to the Traps MuPo downstream of the traps. It consists of a 100ml flow controller, a calibration volume tank (432ml), 2 solenoid valves (ET-2, 24V, Clippard) to direct the sample flow into the sampling volume or bypass it, a sample diaphragm pump (Pfeiffer MVP 006-4) and a NTC temperature sensor. The process controller isolates the pump from the sampling volume during sampling, monitors the pressure difference accumulated and calculates the air volume sampled, additionally taking the
30 temperature reading into account. When not in sampling mode, the sample volume is connected to the sample pump to prepare for the next sampling cycle.

2.2 Cryogenic trapping

2.2.1 Design and implementation

The main objective of the cryogenic concentration systems is to achieve the minimum cycle time between the
35 required temperature set-points (i.e. -160°C to 120°C) with minimum LN₂ consumption. The total operation time between LN₂ refilling is 17h and covers the maximum flight duration (10 hours) and associated pre-flight ground tests (2 hours). The cryogenic concentration system (**Fig. 2**) is positioned on top of the LN₂ container. It consists of the pressure tight container-top plate with three cooler assemblies mounted on top of each other. In accordance



with the target temperatures of each trap, the cryofocus trap is housed at the bottom, closest to the LN₂, the middle one is used for the enrichment trap and the upper one supports the water trap. Good thermal insulation of the whole arrangement is achieved by thin walled, stainless steel tubing and housings surrounded by an Aerogel powder filling up a powder-tight, 3D-printed elastomer enclosure.

- 5 Enrichment and cryofocus trap housings have a tube reaching down into a LN₂ container directly below the top plate of the container, from where LN₂ is drawn if the outlet pressure of the housing is ~50mbar lower than the pressure in the container. The water trap housing takes cold N₂ gas off of the headspace of the container, because its cooling power demands are much smaller. The cooling tubes immersed in the LN₂, trap operations do not interfere with each other, because the change of liquid level is minimal on changing individual cooler power.
- 10 The system can therefore transport LN₂ to rapidly cool down the enrichment and cryofocus traps (**Fig. 2**). This approach makes use of the latent heat energy (as opposed to cooled gas solutions), optimizing the cycle time. The three traps are made of straight thin walled stainless steel tubing. The inner diameter of the water trap is 1/8''(0.318 cm) and the enrichment trap 1/16''(0.159 cm). The cryofocus trap (1/16'') acts as the housing of a 0.25mm inert chromatographic column (Agilent Technologies). Trap temperature measurement is achieved by use of an 100 µm (Omega Engineering, Germany) wire thermocouple attached via a thin walled PFA tube.
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2.2.2 Trap temperature control

Heating of the traps is achieved via direct current from a custom built 4V/30A, 0-100% DC/DC converters. 1/16'' tubing heating rates are ~100°C/s and hard to determine accurately because of the thermocouple reading delay. 1/8'' (water trap) heating rates are ~30°C/s.

- 20 Cooling of the traps is accomplished by controlling the flow rate of N₂ downstream of the cooler housings by proportional valves in series with 1/16'' capillaries of appropriate length. The individual trap systems are self-stabilizing, because higher LN₂ levels in each branch immediately lead to higher evaporation rates, thus higher pressures forcing the LN₂ column down again. With the current setup, cooling from +100 to a stable -180°C is typically possible within ~50s.
- 25 Depending on sample state, the process controller commands either the cooler valves or the trap tubes DC heater power to activate in order to maintain pre-set high or low temperatures. Typical temperature set points for water trap operation were between -40° and +120°C, the enrichment trap was run between -140° and +120°C and the cryofocus trap between -160° and +120°C.

2.2.3 Liquid nitrogen container

- 30 The LN₂ container (LN₂-Badkryostat, CryoVac, Germany) has a usable volume of 10.5l which is sufficient for ~17h of continuous operation at a cycle time of 3.1 min. It is a vacuum isolated stainless steel device equipped with a mechanical overpressure valve and a rupture diaphragm (required for aircraft certification). Additionally, there is a pressure tight cylindrical double-wall tube insert built into the container, which widens into a cone at the top, establishing the LN₂ reservoir from which the trap coolers are supplied (see **Fig. 2**).
- 35 In continuous operation the inside pressure is maintained at ~50mbar higher than the sampling system exhaust pressure (~1000 mbar) by the pressure regulating system (see 2.2.4). The small overpressure drives the LN₂ column up into the cone, where it evaporates with a rate proportional to the cone area immersed. This way the LN₂ is evaporated directly below the traps, taking up the latent heat where it is actually needed, improving the efficiency of the system. With the filling valve open, the LN₂ column remains at the LN₂ level, which determines the idle LN₂ consumption. By closing the filling valve, the mechanical pressure regulating system stabilizes the
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system into a standby state. The level of LN₂ in the container is determined with capacitive LN₂ level probe, which is connected through a custom built voltage-frequency converter to the process controller.

2.2.4 Pressure regulation

5 The main design objectives of the pressure regulating system were a) the establishment of constant sampling pressure independent of sample air flow or outside pressure and b) simple and robust control of the pressure inside the LN₂ system even without electrical power supplied to the instrument. Therefore, the system was built based on pneumatic components. Since no commercial components of reasonable weight and size were available, the diaphragm pressure controllers were custom built. The diaphragms were taken from 50mbar commercial camping gas equipment. By adapting the springs, the set points and dynamic ranges could easily be adapted to
10 our needs.

The so-called constant pressure output regulator (CPOR) has a 100cm³ reference volume attached, which automatically resets to mean ambient pressure with a time constant of 1-2 days. It is used to keep the sampling system base pressure (which is also the LN₂ container control reference pressure) approximately constant at a slightly higher level than ambient (~50mbar). This pressure is monitored throughout continuous measurements
15 to correct for possible systematic effects on overall system performance.

The main uncertainty here is pressure variation caused by temperature changes, which could lead to reference pressure changes in the field on the order of 10%. For use in SOFIA on board a research aircraft, this volume had to be isolated from changes in aircraft internal pressure (ca. 120mbar), whereas on ground or on ships it could be opened to remove drifts caused by temperature fluctuations.

20 Upstream of the CPOR a differential pressure regulator of the same type was used to control both the LN₂ level and the standby flow with the filling valves of the LN₂ container closed. The filling level self-regulation operated analogously to the trap coolers. If the level in the LN₂ container rises, the warmer part of the wall leads to enhanced evaporation which immediately increases the pressure across the exhaust capillary, depressing the level again. The regulation of the LN₂ level was very important in removing the feedbacks of the regulator action on
25 the rest of the system.

2.3 Chromatographic oven

The key factors in achieving the shortest possible cycle time of a fast GC system are a) the reproducible heating and b) extremely fast cooling to maximize the chromatogram run time while minimizing cycle time. To achieve these goals, the thermal masses of the oven, the column mandrel, the fan and the heater (**Fig. 3**) have to be
30 minimized or eliminated. Furthermore, the surface to volume ratio of the mandrel has to be maximized and the insulation surface, mass and performance must be optimized. With the limited cooling power available for the GC oven, and to avoid condensation issues, the chromatography was run above ambient temperatures.

The oven was designed as a horizontally mounted toroidal structure (**Fig. 3**) to minimize internal surface area and thermal convection influence. The oven insulation is made of 3 double-walled hollow rings with wall
35 thickness ~20µm, making up the inner, outer and top insulation. They were manufactured by galvanizing nickel onto a 3D printed plastic core, which was dissolved afterwards by a solvent. Under this configuration the mass, heat capacity and heat conductivity were insignificant compared to convection and heat distribution.

The column mandrel (130mm OD) consists of thin wires holding two concentric sheath metal heaters (0.1mm * 12mm cross section) enveloping the column in shape. The column (DB-624, 10m, 0.25mm, 1.4µm; Agilent
40 Technologies) was nested between the heaters. A thermocouple was placed on the column to get a fast response



measurement of the column temperature. The sheath heaters are directly heated with a 10V/10A max, 0-100% controllable custom built DC-DC converter.

The fan (EBM Papst, 56 Watts, ~ 200mm OD) was taken out of the heated zone of the oven to remove its thermal mass. To sustain stirring of the air inside the oven and to minimize possible temperature gradients by strong thermal convection, the fan was mounted below the column, with ~40% of the oven bottom area being open facing the fan blades.

For rapid cooling of the column, the oven top cover has to be opened in order to enable a high air flow rate through the ring gap around both sides of the mandrel (see **Fig. 3b**). During test experiments, it was found that the ~50g top cover could be lifted by the airstream of the fan, resettling down to the oven side wall rings when the fan was turned down. Since it is inherently difficult to control low heat capacity / high power systems, regulator parameters should be different at various temperatures to guarantee stability and minimize stabilization time. Our custom built system is capable of controlling the system in real time. Low temperature settling time and stability were much improved by running the fan at low power along with the heater. Unfortunately, the fans implemented cannot run down to very low revolution rates to make a smooth fadeout of the fan possible, eliminating the minor disturbances on the temperature ramp (see **Fig. 3c**).

2.4 Detector

The GC column is connected to a quadrupole mass spectrometer MS (Agilent Technologies 5973) via a heated transfer line (143mm deactivated column of 0.25mm; Agilent Technologies). The pump of the MS was replaced by a high-power turbomolecular pump (EXT7DX, Edwards Vacuum) that can operate under higher gravitational forces to avoid problems during turbulence and on landing. The pre-pump was replaced with an oil free membrane pump (MVP006-4, Pfeifer), to avoid potential contamination. In order to achieve higher peak resolution and therefore increase precision, the electronics board of the MS was exchanged with a fast, commercial version (Sideboard PCA G3169-65015, Agilent Technologies) and the MS is operated in Selected Ion Mode (SIM) that substantially improves the detection limits while clearly separating the eluting peaks (**Table 1**). The dwell time for the individual ions selected was 10ms and a complete chromatogram run for 2.4min (**Fig. 4**). The linearity of the 5973MSD signal has been shown in multiple studies during the past decade (e.g. Yassaa et al., 2012). We observed a linear relationship for all species investigated ($R^2 > 0.9$) for mixing ratios ranging from few ppt up to 3 ppb. Average precision of the measurements ranges between 3.9 and 12.4% with the respective total uncertainties is between 8.7 and 14.3% under laboratory conditions using calibration standard (Apel-Riemer Environmental Inc.).

2.5 Process controller hardware and software

Due to limitations in the number of instrument operators on board a research aircraft, the system was designed to be fully automated. All electronic units and sequential sampling processes are controlled with electronics software that has been developed in-house (V25, MPIC). The V25 was additionally coupled with an external computer to trigger the MS data acquisition software Chemstation™.

Faster GC-MS cycling and therefore a higher number of chromatograms acquired creates the need of robust peak integration software. While most of the peaks could be analysed with IAU-Chrom software (Sala et al., 2014), great attention has been given to the separation of peaks that elute in very close retention times and are not clearly separated. Therefore, additional effort was put into the development of software that could clearly separate co-eluting peaks. MPIC-Chrom is a new peak integration software, written in IGOR, and was used for the separation of acetone and propanal peaks.



2.6 Specifications during OMO campaign

The system was installed on board the HALO aircraft after its final configuration and certification. At first, a limited number of compounds (11) was monitored in order to ensure reliable quantification (**Table 1**). At the start of the OMO campaign, high sampling flows (100sccm) resulted in inefficient water removal and hence poor and non-reproducible chromatographic peaks. The solution was to operate the system with a lower sampling flow (40sccm) and only at high altitudes where low dew point temperatures do not affect the sampling procedure. The sampling time was 1min so a total volume of 40 ± 6 ml was collected into the traps. During sample collection, the water trap temperature was set to -30 ± 0.3 °C and the enrichment trap to -140 ± 4 °C. During the sample transfer, the cryofocus trap was set to -160 ± 1 °C. All traps were then heated to 120°C to ensure that all volatiles were desorbed efficiently from each trap.

The GC oven temperature was programmed to start from 50°C, hold this temperature for 20sec and thereafter increase at a rate of 2°C/sec until it stabilizes at 80°C for 10sec and subsequently increase to 150°C at a rate of 1°C/sec (**Fig. 3**). Under the initial configuration of the GC oven, imperfect reproducibility of the temperature ramp resulted in small retention time shifts that nonetheless remained within the selected SIM time windows. Post-campaign improvements on the oven ventilation control system resulted in more reproducible temperature profiles largely eliminating the retention time shifts.

In the first flight campaign a sample was acquired every 3.1 minutes with an MS run of 2.4min. The software was programmed to perform three calibration steps or three zero air measurements after 20 consecutive ambient samples. The calibration points were used to derive an in-flight pressure calibration curve that was used for the quantification of ambient air samples. During flight, higher uncertainty values were determined as a result of lower precision (**Table 2**).

3. Case study of Asian monsoon outflow

3.1 Flight 20170813

The scientific aim of the OMO campaign was to investigate the oxidation processes in the convectively lifted air masses that originate in polluted areas of South Asia, and are then transported within the Indian monsoon anticyclonic flow system in the upper troposphere. The base of operation in the eastern Mediterranean was Paphos airport, in Cyprus, and flights generally headed east over the Arabian Peninsula to intercept monsoon flows heading west. Methane, measured by IR absorption spectroscopy (Schiller et al., 2008, Tadic et al., 2017) was used to identify monsoon outflow influenced air masses. A threshold methane value was derived to identify such plumes based on the average of profiles (4-10km height; 17 flights) over Cyprus, Italy and Germany, which represented the European background without monsoon influence. The threshold was calculated as the sum of the averaged observations plus two times the standard deviation (threshold = average + 2σ = 1879.8ppb). Methane mixing ratios above this threshold were assumed to be influenced by the Indian summer monsoon system.

On August 13th, 2015, a flight was performed to traverse the anticyclonic system that was forecasted to extend over the eastern part of Arabic Peninsula at high altitudes (see **Fig. 5**). The results from this flight will be used as a case study in order to demonstrate the instrument performance.



3.2 Pollution plume characteristics

According to the aforementioned CH₄ threshold, monsoon influenced air masses were mainly encountered on the eastern part of the flight track, with a few additional areas located above continental Saudi Arabia. On average, CH₄ increased by 66.2±23.2 ppb and CO by 29.5±12.3 ppb in the pollution plumes (see red dots in **Fig. 6a**). On the return flight, very clean air masses were encountered and these mixing ratios were considered as background (12:30-13:30 UTC). We define the excess mixing ratios (ExR) as the difference between the observations obtained under background conditions (bg) and the respective mixing ratios measured within the pollution plume (Yokelson et al., 2013).

Increased mixing ratios of all hydro- and chlorocarbons were observed in the pollution plume except for carbon tetrachloride (CCl₄) which remained constant during the entire flight ([CCl₄]_{BG} = 124±13 ppt, [CCl₄]_{plume} = 123±14 ppt). The largest increases were observed for benzene (ExR_{C₆H₆} = 269%) and acetone (ExR_{C₃H₆O} = 225%) while strong increases were also observed for chloromethane (ExR_{CH₃Cl} = 140%) and chloroform (ExR_{CHCl₃} = 55%). Dichloromethane mixing ratios were higher by 18% on average in the pollution plumes, however, this falls within the uncertainty range of the measurement on this flight. Besides the hydro- and chlorocarbons, carbon disulphide was more than double as high in the plume (ExR_{CS₂} = 109%) with a maximum measured mixing ratio of 23ppt.

The highest correlation coefficient between the commonly used pollution marker CO and the monitored hydrocarbons was observed for benzene (CC_{CO,C₆H₆} = 0.85; **Fig. 7**) and acetone (CC_{CO,C₆H₆} = 0.75). High correlations were also observed between benzene and chloroform (CC_{C₆H₆,CHCl₃} = 0.78) as well as benzene and carbon disulphide (CC_{C₆H₆,CS₂} = 0.61), and even higher for the air masses marked as pollution plumes (CC_{C₆H₆,CHCl₃} = 0.81, CC_{C₆H₆,CS₂} = 0.92) as shown in **Fig. 8**.

3.3 Air mass separation

The highest mixing ratios of CH₄ and CO were observed between 09:53-10:40 UTC over Oman (**Fig. 6** and **Fig. 11**) and with relatively stable plume delineator abundance ([CH₄] = 1925±14 ppb ; [CO] = 109±7 ppb) which is indicative of a large scale pollution plume. Interestingly, the main species measured by SOFIA reveal markedly different mixing ratios for the first and second part of the plume. This suggests that, what appears to be one plume in the relatively unspecific marker compounds (CO and CH₄), has two distinctly different composition regions when volatile organic compounds are considered. To investigate the chemical differences between these air masses, the plume was subdivided into two separate plumes termed as P1 and P2.

In **Fig. 9** we illustrate the ExRs of the main VOC species measured over P1 and P2 against the respective ExR of CH₄ and CO. Benzene, acetone and chloroform increased in both plumes, moderately correlated with the respective increase in the CO mixing ratios (R²_{C₆H₆,CO} = 0.6 , R²_{C₃H₆O,CO} = 0.61, R²_{CHCl₃,CO} = 0.52).

Benzene is the most abundant aromatic hydrocarbon in the atmosphere (Martín-Reviejo and Wirtz, 2005) and its gas phase chemistry is dominated by the reaction with OH radical (Bloss et al., 2005). Differences between biomass burning and pollution outflow can be traced by the benzene/CO ratio. Scheeren et al. (2003) reported a ratio of 0.23 ppt ppb⁻¹ inside the monsoon outflow while Andreae and Merlet (2001) derived a ratio of 1.3 ppt ppb⁻¹ for biomass burning plumes. In our measurements, the benzene/CO ratio was significantly increased in P2 (Benz/CO_{P1} = 0.4 ppt ppb⁻¹, Benz/CO_{P2} = 0.7 ppt ppb⁻¹). In another study, Hornbrook et al. (2011) used the excess ratios (ΔBenz/ΔCO) for all biomass burning influenced air masses and reported values that range between 0.5 and 2.5 ppt ppb⁻¹. The respective values were increased by ~80% in P2 (compared with P1), with an average



ratio of 1.44 ± 0.19 ppt ppb^{-1} indicating the influence of biomass burning emissions in the second part of the plume.

Acetone is the most abundant OVOC in the upper troposphere with mixing ratios that can reach 2ppb (Pöschl et al., 2001) with strong seasonal variations at the mid-latitude tropopause (Sprung and Zahn, 2010). It has both anthropogenic and biogenic sources (Jacob et al., 2002; Khan et al., 2015) but it can be also formed by the oxidation of precursor compounds such as propane (Jacob et al., 2002; Fischer et al., 2012). Biomass burning is another direct source of acetone (Holzinger et al., 2005). We have observed values up to 1.5 ppb in the Asian monsoon plume with highly elevated mixing ratios in P2 ($[\text{C}_3\text{H}_6\text{O}]_{\text{P1}} = 995 \pm 246$ ppt, $[\text{C}_3\text{H}_6\text{O}]_{\text{P2}} = 1397 \pm 18$ ppt), identified as biomass burning influence. At the same time, propanal was moderately increased from 19 ± 4 ppt in P1 to 29 ± 5 ppt in P2. In a recent study, Fischbeck et al. (2017) showed that the increase of acetone in such pollution plumes is more likely to depend on the initial mixing ratios at the source rather than on secondary production.

Chloroform has both anthropogenic and biogenic sources, while biomass burning is considered to be an important contributor (Laternus et al., 2002). We observed an increase of chloroform in P1 ($[\text{CHCl}_3]_{\text{P1}} = 15 \pm 3$ ppt) and a stronger increase in P2 ($[\text{CHCl}_3]_{\text{P2}} = 22.4 \pm 2$ ppt) compared to the background measurements ($[\text{CHCl}_3]_{\text{BG}} = 10 \pm 1$ ppt). In general, chloroform increased in a linear relationship with benzene as shown in **Fig. 8**.

For most species we observed enhanced mixing ratios in the second part of the plume (P2). The only compound that displayed a different tendency is chloromethane, which was higher in P1 compared to P2 ($[\text{CH}_3\text{Cl}]_{\text{P1}} = 1485 \pm 274$ ppt, $[\text{CH}_3\text{Cl}]_{\text{P2}} = 988 \pm 136$ ppt). While biomass burning is also a significant source of chloromethane (Rudolph et al., 1995; Andreae and Merlet, 2001; Keppler et al., 2005; Umezawa et al., 2014), we identified stronger emissions from the non-burning monsoon outflow, which indicates that different strong sources are present. Scheeren et al. (2003) have shown that a ratio of 10ppt ppb^{-1} is indicative of Asian pollution over the eastern Mediterranean. Hence the higher $\text{CH}_3\text{Cl}/\text{CO}$ ratios observed in P1 ($[\text{CH}_3\text{Cl}/\text{CO}]_{\text{P1}} = 14.1 \pm 2.6$ ppt ppb^{-1}) compared with P2 ($[\text{CH}_3\text{Cl}/\text{CO}]_{\text{P2}} = 8.7 \pm 1.5$ ppt ppb^{-1}) are indicative for pollution outflow, which may have been influenced by other strong sources such as tropical vegetation (Gebhardt et al., 2008), biofuel use (Lobert et al., 1999) or the burning of agricultural residues, waste and dung (Scheeren et al., 2002).

To further investigate the air mass characteristics, normalized excess mixing ratios (NERMs) relative to CO, were calculated and presented in **Fig. 10**. The observed NERMs between benzene and chloromethane reveal a distinct separation of the relationships between the two parts of the plume. The linear relationship for chloromethane inside P2 ($R^2 = 0.74$), in combination with the range of $\text{CH}_3\text{Cl}/\text{CO}$ ratios (6.7 to 10.9 ppt ppb^{-1}) support the assumption that the air was influenced by biomass burning (Scheeren et al., 2003). The NERMs of chloroform and carbon disulphide increased in a similar linear manner. Especially chloroform enhancement ratios, relative to CO, were uniquely correlated with the respective enhancement of benzene ($R^2 = 0.98$; **Fig. 10**). We further examine the distinct plume characteristics by calculating 10 day back-trajectories using the FLEXible PARTicle dispersion model (FLEXPART; Stohl et al., 1998). As shown in **Fig. 11**, when HALO descended by 1 km at 10:21 UTC, a different air mass was measured. P1 originated over north India and at higher altitudes (~ 8 km) and was transported within the prevailing anticyclonic system. In contrast, the air mass P2 emerged from lower altitudes over central India with some influence from Bangladesh, Bhutan and west Myanmar, were biomass burning (Streets et al., 2003; van der Werf et al., 2006) and fuel consumption (van der Werf et al., 2010) are more prominent.



4. Discussion

The need for high time resolution monitoring in the upper troposphere has given rise to a new generation of custom made, fast GC instruments that were designed to operate on board a research aircraft (**Table 2**). The most critical parameter for fast GC instrumentation on board a research aircraft is time resolution. As demonstrated in the case study section, the aircraft can rapidly cross air masses with different characteristics and high resolution monitoring is essential to interpret the underlying atmospheric phenomena.

The main constraints in achieving high time resolution sampling for GC-MS systems are the cooling and heating rates of the traps and GC oven, together with the inherent limitations of chromatographic separation time. During the first campaign (OMO), our system was operated with a time resolution of 3.1min, while a reduction to 2min is feasible due to the high cooling rates of both the GC oven and traps' housing. Stable and reproducible temperature are essential in to ensure reproducible retention times that ease the analysis.

Quadrupole MS has been the preferred method for quantitative detection despite the inherent restrictions on the number of species monitored due to limits in the number of measurement points needed to accurately define the fast eluting peaks and the numbers of ions that can be sequentially monitored. In the selected ion monitoring mode (SIM) several quadrupole systems have proven to be sufficiently sensitive, robust, with good reproducibility and a high degree of linearity over a wide range of mixing ratios (e.g. Apel et al., 2003). Importantly, when targeting fast measurement times the MS allows the separation of co-eluting substances that can be distinguished by their different mass and fragmentation patterns. While considerable improvements on the mass resolving power and speed could be achieved with a time-of-flight (TOF-MS) mass spectrometer that can simultaneously measure all mass-to-charge signals at high frequency, detector non-linearities of the instrument sensitivity have been observed (Hoker et al., 2015; Obersteiner et al., 2016a).

Generally for all systems, prior to detection, the substances of interest are pre-concentrated in an enrichment trap, where low temperatures are used to retain them. The range of trapping temperatures required is highly dependent on the target substances. SOFIA utilizes LN₂ in order to achieve large cooling capacity allowing high vapor pressure VOCs such as methyl chloride to be trapped without the need of an absorbent material that may induce artifacts and memory effects (Apel et al., 2003). The main disadvantage of using LN₂ as cryogen is the significant volumes required in combination with safety restrictions and supply in remote locations. The SOFIA system can be used on routine flights that do not exceed 17h. However, multiple flights that include an extended layover or remote site landing for an overnight stay will restrict the operation of SOFIA to the outward flight, unless a re-supply of LN₂ is available in the host airport.

Alternative cryogen free approaches have also been realized. Compression coolers containing refrigerant (e.g. 1,1,1,2-tetrafluoroethane) in combination with a tube evaporator that cools a liquid such as 50% ethanol have been applied (Jäger, 2015). The disadvantage of this method is the reduced cooling capacity with associated limitations in retaining high vapor pressure substances. A further solution has been the use of stirling coolers as demonstrated by Obersteiner et al. (2016b). These only require electrical power and hence are eminently suitable for long term use at remote locations. The cooling rates and minimum attainable temperatures are comparable (albeit slightly higher) than with the LN₂ systems, even if they are not as powerful. In any case, low adsorption temperatures will result in trapping CO₂ which can induce chromatographic and detection problems depending on the selected ions monitored. Since the species monitored do not have interferences with the parent CO₂ ion, our measurements were not influenced by ambient CO₂.

Last but not least, any instrument that is mounted on a research aircraft is subject to weight limitations. Our system is comparatively compact and light (120 kg), despite the fact that is equipped with a liquid nitrogen



5 container. Summarizing, the different cryogenic methods and instrument configurations chosen for each fast GC are tailored to the specific research objectives of each instrument and come with both advantages and disadvantages concerning their application for measurements in the upper troposphere. The SOFIA instrument has been compactly built to meet the following specifications: fit one aircraft rack (65 x 55 x 163 cm) at maximum 120 kg, and operate for 17 hours on liquid nitrogen, which enables even species such as methyl chloride to be effectively trapped in-situ at an altitude range of 0-15km.

5. Conclusions

10 We have developed a new fast GC-MS instrument for airborne measurements of volatile organic compounds including hydro- and halocarbons, OVOC and sulphur species. The system incorporates a novel cryogen-conservative VOC enrichment system that is based on the differential pressure between a LN₂ dewar and the trap housing to transport the cryogen as a liquid, and rapidly cool the traps to the desired temperatures. In addition, we have developed a new chromatographic oven with exceptionally high cooling rates (30°C/min) and rapid stabilization which helps achieve relatively high measurement time resolution. SOFIA was operated on-board the HALO research aircraft during the OMO campaign to study the convectively transported pollutants within the Indian monsoon anticyclone system. The high time resolution allowed the investigation of a seemingly uniform pollution plume and distinguish two notably different air masses. We have confirmed the distinct origins of these air masses with the use of a back-trajectory transport model, and conclude that the use of on-line, high resolution monitoring is essential for the adequate characterization of air masses and atmospheric processes that take place in the upper troposphere.

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7. Tables and figures

Compound	Formula	SIM ion	Detection limit (ppt)	Average precision of measurements (%)	Total uncertainty (%)	RT	RTstd
Methyl chloride	CH ₃ Cl	50,52	3 (6)	5,2 (3,7)	8,8 (8)	0,43	0,005
Methyl bromide	CH ₃ Br	94	1	5,2	8,8	0,52	0,008
Trichlorofluoromethane	CCl ₃ F	101	1	5,4	8,9	0,61	0,02
Isoprene	C ₅ H ₈	67	3 (6)	6,6 (10,5)	9,7 (12,6)	0,68	0,02
Propanal	C ₃ H ₆ O	58	11 (19)	8,3 (5,1)	10,9 (8,7)	0,72	0,02
Ethane,1,1,2-trichloro-1,2,2-trifluoro	C ₂ Cl ₃ F ₃	101	1	3,9	8,1	0,73	0,02
Acetone	C ₃ H ₆ O	58	12 (21)	7,5 (6,1)	10,3 (9,3)	0,74	0,02
Methyl iodine	CH ₃ I	142	1	5,1	8,7	0,76	0,02
Carbon disulfide	CS ₂	76	1 (4)	5,3 (11,6)	8,8 (13,6)	0,78	0,02
Dichloromethane	CH ₂ Cl ₂	49	3 (9)	5,4 (11,3)	8,9 (13,3)	0,84	0,02
Butanone	CH ₃ CCH ₂ CH ₃	43	10	7,4	10,2	1,02	0,02
Chloroform	CHCl ₃	83	1 (4)	6,1 (10,4)	9,3 (12,6)	1,06	0,03
1,1,1-Trichloroethane	C ₂ H ₃ Cl ₃	97	1	6,5	9,6	1,46	0,05
Cyclohexane	C ₆ H ₁₂	56	4	5,2	8,8	1,48	0,05
Carbon tetrachloride	CCl ₄	117	1 (5)	5,9 (6,5)	9,2 (6,5)	1,52	0,05
Benzene	C ₆ H ₆	78	1 (4)	7,7 (6,4)	10,5 (9,5)	1,60	0,03
1,2-Dichloroethane	C ₂ H ₄ Cl ₂	62	5	10,6	12,7	1,62	0,02
Trichloroethylene	C ₂ HCl ₃	130	2	9,6	11,9	1,82	0,03
Toluene	C ₇ H ₈	91	3 (8)	6,9 (5,2)	11,7 (8,8)	2,20	0,03
Tetrachloroethylene	C ₂ Cl ₄	166	3 (7)	12,4 (10,2)	14,3 (12,4)	2,35	0,03

5 **Table 1.** Selective Ion Mode (SIM) was used to measure the listed compounds. The total uncertainty takes into account the measurement precision, the 5% uncertainty of the standard gas and the 5% uncertainty of the sampling volume. Retention times (RT) and their standard deviations from 2.4min chromatograms are given in the last two columns. The values in brackets indicate the performance during flight 20150813.

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Instrument	Detector	Cryogenic method	Time resolution (min)	Adsorption temp. (°C)	Lowest DL (ppt)	Highest precision (%)	Weight (kg)	Reference
SOFIA	MSD 5973	LN ₂	3 (2)	-140	1	4	120	This study
TOGA	MSD 5973	LN ₂	2	-130	1	3	<200	Apel (2016)
GhostMS	MSD 5975	stirling cooler	4,3	-100	0,001	2	n.r.	Sala et al. (2014)
HGC	MSD 5975	compressed refrigerant	9	30	1	<1	128	Jäger (2015)
µDirac	ECD	n.r.	8-15	15-25	0,5	1	11	Gostlow et al. (2010)

Table 2. Custom built fast GC instruments for on-line monitoring of organic trace gases. Non reported information is denoted with n.r.

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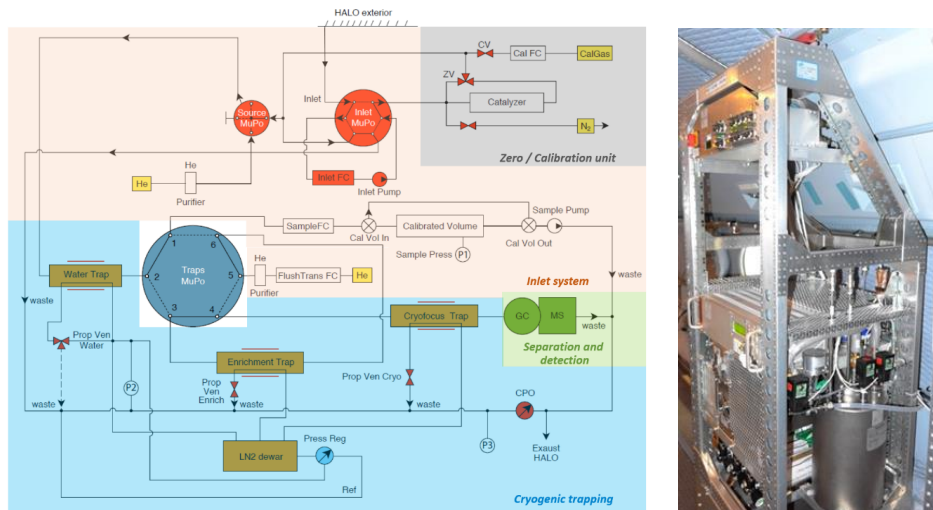


Figure 1: Schematic overview of SOFIA (left) and a photo of the instrument on-board HALO aircraft (right).

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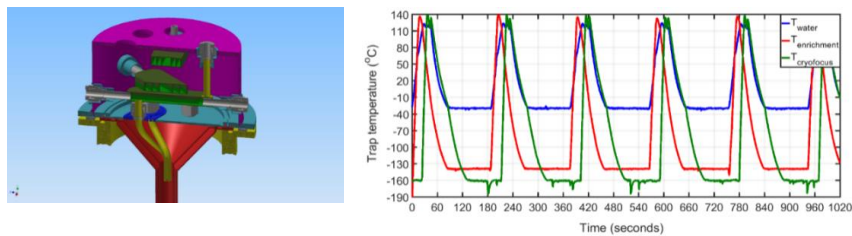


Figure 2. Cryocooler and trap temperatures. Schematic of the trap housing (left) and illustration of temperatures over six cycles (right). The water trap (T_{water}) is shown in blue, the enrichment trap ($T_{\text{enrichment}}$) in red and the cryofocus trap (T_{cryo}) in green.

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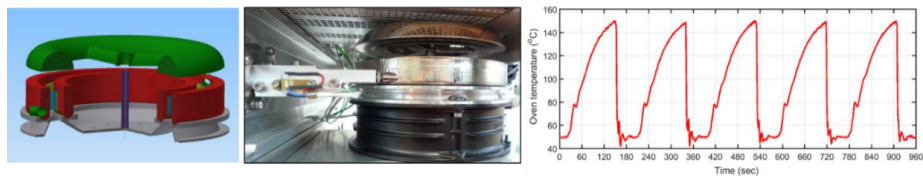


Figure 3. The chromatographic oven. Schematic design of the oven (left), photo of the oven during cooling (middle) and illustration of the oven temperature from six consecutive cycles during the case study flight (right).

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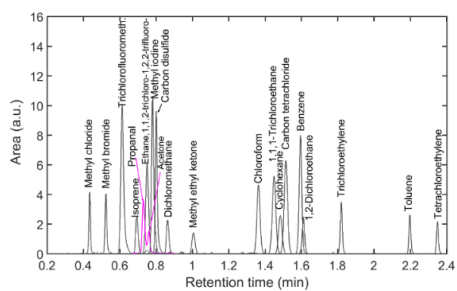


Figure 4. Chromatogram using standard gas. Elution times during a 3min sample. The signals of propanal and acetone (purple) are illustrated as 10x the raw signals. The monitored compounds are listed in Table 1.

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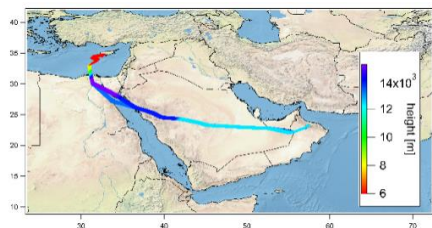


Figure 5. Route of research flight 20150813

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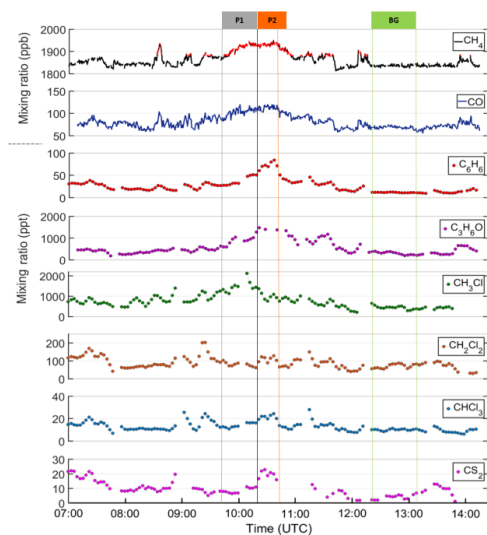


Figure 6. Time series of eight compounds from flight 20150813.

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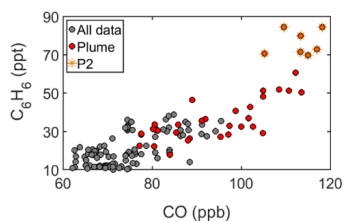


Figure 7. Correlation between CO and benzene mixing ratios during flight 20150813. The data points that were impacted by the Asian monsoon outflow are shown by red bullets, and those that were not in grey. The increased mixing ratios observed at P2 (see Fig. 6) are shown by the orange symbols.

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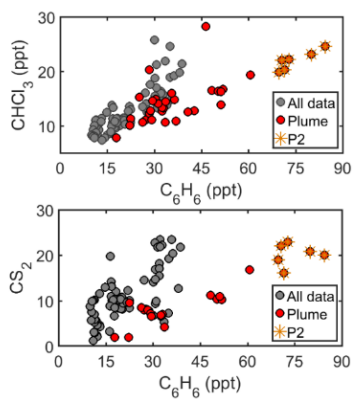


Figure 8. Correlation between benzene and chloromethane (top), and benzene and carbon disulphide (Bottom). Symbols as in Fig. 7.

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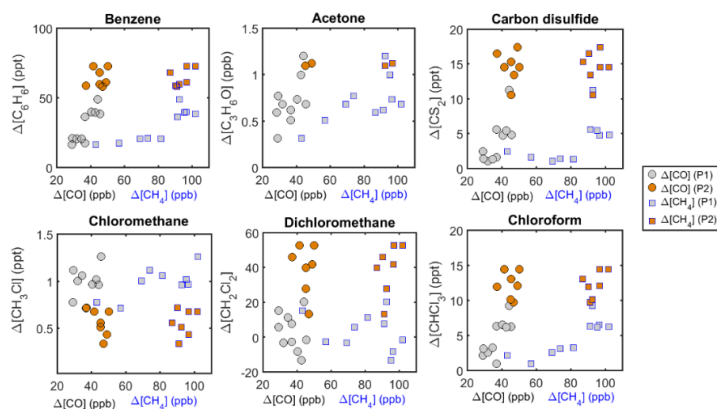


Figure 9. Enhancement mixing ratio correlations in the two plumes. Correlations of EnR between CO (cycles) and CH₄ (squares) with benzene, acetone, carbon disulfide, chloromethane, dichloromethane and chloroform.

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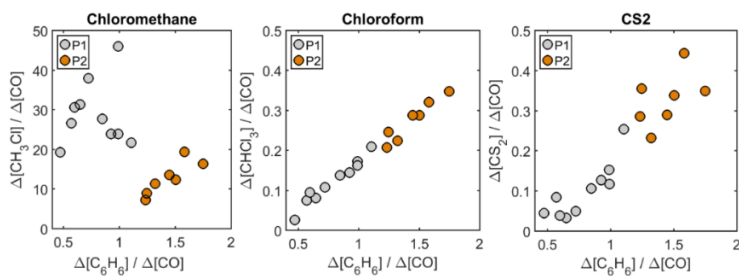


Figure 10. Comparison plots of NEMRs to CO for chloromethane, chloroform and carbon disulphide. The first part of the pollution plume P1 is indicated by grey bullets and the second part P2 with orange bullets.

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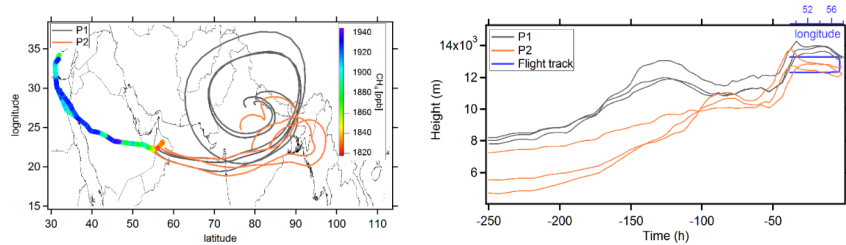


Figure 11. Back-trajectories for the plumes P1 and P2 according to geographic location (left) and in time (right).

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References

- Andreae, M. O., and Merlet, P.: Emission of trace gases and aerosols from biomass burning, *Global Biogeochemical Cycles*, 15, 955-966, [10.1029/2000GB001382](https://doi.org/10.1029/2000GB001382), 2001.
- 5 Apel, E. C., Hills, A. J., Lueb, R., Zindel, S., Eisele, S., and Riemer, D. D.: A fast-GC/MS system to measure C2 to C4 carbonyls and methanol aboard aircraft, *Journal of Geophysical Research: Atmospheres*, 108, [10.1029/2002JD003199](https://doi.org/10.1029/2002JD003199), 2003.
- 10 Apel, E. C., Emmons, L. K., Karl, T., Flocke, F., Hills, A. J., Madronich, S., Lee-Taylor, J., Fried, A., Weibring, P., Walega, J., Richter, D., Tie, X., Mauldin, L., Campos, T., Weinheimer, A., Knapp, D., Sive, B., Kleinman, L., Springston, S., Zaveri, R., Ortega, J., Voss, P., Blake, D., Baker, A., Warneke, C., Welsh-Bon, D., de Gouw, J., Zheng, J., Zhang, R., Rudolph, J., Junkermann, W., and Riemer, D. D.: Chemical evolution of volatile organic compounds in the outflow of the Mexico City Metropolitan area, *Atmos. Chem. Phys.*, 10, 2353-2375, [10.5194/acp-10-2353-2010](https://doi.org/10.5194/acp-10-2353-2010), 2010.
- 15 Apel, E. C., Olson, J. R., Crawford, J. H., Hornbrook, R. S., Hills, A. J., Cantrell, C. A., Emmons, L. K., Knapp, D. J., Hall, S., Mauldin Iii, R. L., Weinheimer, A. J., Fried, A., Blake, D. R., Crouse, J. D., Clair, J. M. S., Wennberg, P. O., Diskin, G. S., Fuelberg, H. E., Wisthaler, A., Mikoviny, T., Brune, W., and Riemer, D. D.: Impact of the deep convection of isoprene and other reactive trace species on radicals and ozone in the upper troposphere, *Atmos. Chem. Phys.*, 12, 1135-1150, [10.5194/acp-12-1135-2012](https://doi.org/10.5194/acp-12-1135-2012), 2012.
- 20 Apel, E., & UCAR/NCAR - Earth Observing Laboratory. Trace Organic Gas Analyzer (TOGA) for HIAPER. UCAR/NCAR - Earth Observing Laboratory. <https://doi.org/10.5065/D6DF6P9Q> Retrieved December 20, 2016.
- 25 Arnold, F., and Hauck, G.: Lower stratosphere trace gas detection using aircraft-borne active chemical ionization mass spectrometry, *Nature*, 315, 307-309, 1985.
- Atkinson, R.: Atmospheric chemistry of VOCs and NOx, *Atmospheric Environment*, 34, 2063-2101, [http://dx.doi.org/10.1016/S1352-2310\(99\)00460-4](http://dx.doi.org/10.1016/S1352-2310(99)00460-4), 2000.
- 30 Baker, A., Sauvage, C., Thorentz U. R., van Velthoven, P., Oram, D. E., Zahn, A., Brenninkmeier, C. A. M., and Williams, J.: Evidence for strong, widespread chlorine radical chemistry associated with pollution outflow from continental Asia, *Nat. Sci. Rep.*, 6, 36821, [doi:10.1038/srep36821](https://doi.org/10.1038/srep36821), 2016.
- 35 Blake, R. S., Wyche, K. P., Ellis, A. M., and Monks, P. S.: Chemical ionization reaction time-of-flight mass spectrometry: Multi-reagent analysis for determination of trace gas composition, *International Journal of Mass Spectrometry*, 254, 85-93, <http://dx.doi.org/10.1016/j.ijms.2006.05.021>, 2006.
- Bloss, C., Wagner, V., Jenkin, M. E., Volkamer, R., Bloss, W. J., Lee, J. D., Heard, D. E., Wirtz, K., Martin-Reviejo, M., Rea, G., Wenger, J. C., and Pilling, M. J.: Development of a detailed chemical mechanism (MCMv3.1) for the atmospheric oxidation of aromatic hydrocarbons, *Atmos. Chem. Phys.*, 5, 641-664, [10.5194/acp-5-641-2005](https://doi.org/10.5194/acp-5-641-2005), 2005.
- 40 Colman, J. J., Swanson, A. L., Meinardi, S., Sive, B. C., Blake, D. R., and Rowland, F. S.: Description of the Analysis of a Wide Range of Volatile Organic Compounds in Whole Air Samples Collected during PEM-Tropics A and B, *Analytical Chemistry*, 73, 3723-3731, [10.1021/ac010027g](https://doi.org/10.1021/ac010027g), 2001.
- Colomb, A., Williams, J., Crowley, J., Gros, V., Hofmann, R., Salisbury, G., Klüpfel, T., Kormann, R., Stickler, A., Forster, C., and Lelieveld, J.: Airborne Measurements of Trace Organic Species in the Upper Troposphere Over Europe: the Impact of Deep Convection, *Environmental Chemistry*, 3, 244-259, <https://doi.org/10.1071/EN06020>, 2006.
- 50 Crutzen, P. J., Williams, J., Pöschl, U., Hoor, P., Fischer, H., Warneke, C., Holzinger, R., Hansel, A., Lindinger, W., Scheeren, B., and Lelieveld, J.: High spatial and temporal resolution measurements of primary organics and their oxidation products over the tropical forests of Surinam, *Atmos. Environ.*, 34, 1161-1165, 2000.
- Fischbeck, G., Bönisch, H., Neumaier, M., Brenninkmeijer, C. A. M., Orphal, J., Brito, J., Becker, J., Sprung, D., van Velthoven, P. F. J., and Zahn, A.: Acetone-CO enhancement ratios in the upper troposphere based on 7 years of CARIBIC data: new insights and estimates of regional acetone fluxes, *Atmos. Chem. Phys.*, 17, 1985-2008, [10.5194/acp-17-1985-2017](https://doi.org/10.5194/acp-17-1985-2017), 2017.
- 60 Gebhardt, S., Colomb, A., Hofmann, R., Williams, J., and Lelieveld, J.: Halogenated organic species over the tropical South American rainforest, *Atmos. Chem. Phys.*, 8, 3185-3197, [doi:10.5194/acp-8-3185-2008](https://doi.org/10.5194/acp-8-3185-2008), 2008.



- Fischer, E. V., Jacob, D. J., Millet, D. B., Yantosca, R. M., and Mao, J.: The role of the ocean in the global atmospheric budget of acetone, *Geophysical Research Letters*, 39, 10.1029/2011GL050086, 2012.
- 5 Gostlow, B., Robinson, A. D., Harris, N. R. P., O'Brien, L. M., Oram, D. E., Mills, G. P., Newton, H. M., Yong, S. E., and A Pyle, J.: μ -Dirac: an autonomous instrument for halocarbon measurements, *Atmos. Meas. Tech.*, 3, 507-521, 10.5194/amt-3-507-2010, 2010.
- Haagen-Smit A. J., Bradley C. E., Fox M. M.: Ozone formation in the photooxidation of organic substances. *Ind. Eng. Chem.*, 9, 48, 1484-1487, 10.1021/ie50525a044, 1956.
- 10 Hoker, J., Obersteiner, F., Bönisch, H., and Engel, A.: Comparison of GC/time-of-flight MS with GC/quadrupole MS for halocarbon trace gas analysis, *Atmos. Meas. Tech.*, 8, 2195-2206, 10.5194/amt-8-2195-2015, 2015.
- 15 Holzinger, R., Williams, J., Salisbury, G., Klüpfel, T., de Reus, M., Traub, M., Crutzen, P. J., and Lelieveld, J.: Oxygenated compounds in aged biomass burning plumes over the Eastern Mediterranean: evidence for strong secondary production of methanol and acetone, *Atmos. Chem. Phys.*, 5, 39-46, 10.5194/acp-5-39-2005, 2005.
- Hornbrook, R. S., Blake, D. R., Diskin, G. S., Fried, A., Fuelberg, H. E., Meinardi, S., Mikoviny, T., Richter, D., Sachse, G. W., Vay, S. A., Walega, J., Weibring, P., Weinheimer, A. J., Wiedinmyer, C., Wisthaler, A., Hills, A., Riemer, D. D., and Apel, E. C.: Observations of nonmethane organic compounds during ARCTAS − Part 1: Biomass burning emissions and plume enhancements, *Atmos. Chem. Phys.*, 11, 11103-11130, 10.5194/acp-11-11103-2011, 2011.
- 20 25 Jacob, D. J., Field, B. D., Jin, E. M., Bey, I., Li, Q., Logan, J. A., Yantosca, R. M., and Singh, H. B.: Atmospheric budget of acetone, *Journal of Geophysical Research: Atmospheres*, 107, ACH 5-1-ACH 5-17, 10.1029/2001JD000694, 2002.
- Jäger, J.: Airborne VOC measurements on board the Zeppelin NT during PEGASOS campaigns in 2012 deploying the improved Fast-GC-MSD System, PhD, Forschungszentrum Jülich GmbH, Jülich, Germany, 2014.
- 30 35 Keppler, F., Harper, D. B., Röckmann, T., Moore, R. M., and Hamilton, J. T. G.: New insight into the atmospheric chloromethane budget gained using stable carbon isotope ratios, *Atmos. Chem. Phys.*, 5, 2403-2411, 10.5194/acp-5-2403-2005, 2005.
- Khan, M. A. H., Cooke, M. C., Utembe, S. R., Archibald, A. T., Maxwell, P., Morris, W. C., Xiao, P., Derwent, R. G., Jenkin, M. E., Percival, C. J., Walsh, R. C., Young, T. D. S., Simmonds, P. G., Nickless, G., O'Doherty, S., and Shallcross, D. E.: A study of global atmospheric budget and distribution of acetone using global atmospheric model STOCHEM-CRI, *Atmospheric Environment*, 112, 269-277, <http://dx.doi.org/10.1016/j.atmosenv.2015.04.056>, 2015.
- 40 45 Kley, D.: Tropospheric Chemistry and Transport, *Science*, 276, 1043-1044, 10.1126/science.276.5315.1043, 1997.
- Laternus, F., Haselmann, K. F., Borch, T., and Grøn, C.: Terrestrial natural sources of trichloromethane (chloroform, CHCl₃) – An overview, *Biogeochemistry*, 60, 121-139, 10.1023/a:1019887505651, 2002.
- 50 Le Breton, M., McGillen, M. R., Muller, J. B. A., Bacak, A., Shallcross, D. E., Xiao, P., Huey, L. G., Tanner, D., Coe, H., and Percival, C. J.: Airborne observations of formic acid using a chemical ionization mass spectrometer, *Atmos. Meas. Tech.*, 5, 3029-3039, 10.5194/amt-5-3029-2012, 2012.
- Lelieveld, J., Gromov, S., Pozzer, A., and Taraborrelli, D.: Global tropospheric hydroxyl distribution, budget and reactivity, *Atmos. Chem. Phys.*, 16, 12477-12493, 2016.
- 55 60 Lerner, B. M., Gilman, J. B., Aikin, K. C., Atlas, E. L., Goldan, P. D., Graus, M., Hendershot, R., Isaacman-VanWertz, G. A., Koss, A., Kuster, W. C., Lueb, R. A., McLaughlin, R. J., Peischl, J., Sueper, D., Ryerson, T. B., Tokarek, T. W., Warneke, C., Yuan, B., and de Gouw, J. A.: An improved, automated whole air sampler and gas chromatography mass spectrometry analysis system for volatile organic compounds in the atmosphere, *Atmos. Meas. Tech.*, 10, 291-313, 10.5194/amt-10-291-2017, 2017.
- Lobert, J. M., Keene, W. C., Logan, J. A., and Yevich, R.: Global chlorine emissions from biomass burning: Reactive chlorine emissions inventory, *J. Geophys. Res.*, 104(D7), 8373-8389, 1999.



- Martín-Reviejo, M., and Wirtz, K.: Is Benzene a Precursor for Secondary Organic Aerosol?, *Environmental Science & Technology*, 39, 1045-1054, 10.1021/es049802a, 2005.
- 5 Mastovska, K., and Lehotay, S. J.: Practical approaches to fast gas chromatography-mass spectrometry, *Journal of chromatography. A*, 1000, 153-180, 2003.
- Molina, M. J., and Rowland, F. S.: Stratospheric sink for chlorofluoromethanes: chlorine atomc-atylsed destruction of ozone, *Nature*, 249, 810-812, 1974.
- 10 Monks, P. S.: Gas-phase radical chemistry in the troposphere, *Chemical Society Reviews*, 34, 376-395, 10.1039/B307982C, 2005.
- Obersteiner, F., Bönisch, H., and Engel, A.: An automated gas chromatography time-of-flight mass spectrometry instrument for the quantitative analysis of halocarbons in air, *Atmos. Meas. Tech.*, 9, 179-194, 10.5194/amt-9-179-2016, 2016a.
- 15 Obersteiner, F., Bönisch, H., Keber, T., O'Doherty, S., and Engel, A.: A versatile, refrigerant- and cryogen-free cryofocusing-thermodesorption unit for preconcentration of traces gases in air, *Atmos. Meas. Tech.*, 9, 5265-5279, 10.5194/amt-9-5265-2016, 2016b.
- 20 Parrella, J. P., Jacob, D. J., Liang, Q., Zhang, Y., Mickley, L. J., Miller, B., Evans, M. J., Yang, X., Pyle, J. A., Theys, N., and Van Roozendael, M.: Tropospheric bromine chemistry: implications for present and pre-industrial ozone and mercury, *Atmos. Chem. Phys.*, 12, 6723-6740, 10.5194/acp-12-6723-2012, 2012.
- 25 Plass-Dülmer, C., N. Schmidbauer, J. Slemr, F. Slemr, and H. D'Souza: European hydrocarbon intercomparison experiment AMOHA part 4: Canister sampling of ambient air, *J. Geophys. Res.*, 111, D04306, doi:10.1029/2005JD006351, 2006.
- Pollmann, J., Helmig, D., Hueber, J., Plass-Dülmer, C., and Tans, P.: Sampling, storage, and analysis of C2-C7 non-methane hydrocarbons from the US National Oceanic and Atmospheric Administration Cooperative Air Sampling Network glass flasks, *Journal of Chromatography A*, 1188, 75-87, <http://dx.doi.org/10.1016/j.chroma.2008.02.059>, 2008.
- 30 Pöschl, U., Williams, J., Hoor, P., Fischer, H., Crutzen, P. J., Warneke, C., Holzinger, R., Hansel, A., Jordan, A., Lindinger, W., Scheeren, H. A., Peters, W., and Lelieveld, J.: High Acetone Concentrations throughout the 0-12 km Altitude Range over the Tropical Rainforest in Surinam, *Journal of Atmospheric Chemistry*, 38, 115-132, 10.1023/a:1006370600615, 2001.
- 35 Pusede, S. E., and Cohen, R. C.: On the observed response of ozone to NO_x and VOC reactivity reductions in San Joaquin Valley California 1995-present, *Atmos. Chem. Phys.*, 12, 8323-8339, 10.5194/acp-12-8323-2012, 2012.
- 40 Read, K. A., Mahajan, A. S., Carpenter, L. J., Evans, M. J., Faria, B. V. E., Heard, D. E., Hopkins, J. R., Lee, J. D., Moller, S. J., Lewis, A. C., Mendes, L., McQuaid, J. B., Oetjen, H., Saiz-Lopez, A., Pilling, M. J., and Plane, J. M. C.: Extensive halogen-mediated ozone destruction over the tropical Atlantic Ocean, *Nature*, 453, 1232-1235, http://www.nature.com/nature/journal/v453/n7199/supinfo/nature07035_S1.html, 2008.
- 45 Rudolph, J., Khedim, A., Koppmann, R., and Bonsang, B.: Field study of the emissions of methyl chloride and other halocarbons from biomass burning in Western Africa, *Journal of Atmospheric Chemistry*, 22, 67-80, 10.1007/bf00708182, 1995.
- 50 Saiz-Lopez, A., Lamarque, J. F., Kinnison, D. E., Tilmes, S., Ordóñez, C., Orlando, J. J., Conley, A. J., Plane, J. M. C., Mahajan, A. S., Sousa Santos, G., Atlas, E. L., Blake, D. R., Sander, S. P., Schauffler, S., Thompson, A. M., and Brasseur, G.: Estimating the climate significance of halogen-driven ozone loss in the tropical marine troposphere, *Atmos. Chem. Phys.*, 12, 3939-3949, 10.5194/acp-12-3939-2012, 2012.
- 55 Sala, S., Bönisch, H., Keber, T., Oram, D. E., Mills, G., and Engel, A.: Deriving an atmospheric budget of total organic bromine using airborne in situ measurements from the western Pacific area during SHIVA, *Atmos. Chem. Phys.*, 14, 6903-6923, 10.5194/acp-14-6903-2014, 2014.
- 60 Scheeren, H. A., Lelieveld, J., de Gouw, J. A., van der Veen, C., and Fischer, H.: Methyl chloride and other chlorocarbons in polluted air during INDOEX, *Journal of Geophysical Research: Atmospheres*, 107, INX2 14-11-INX12 14-18, 10.1029/2001JD001121, 2002.



- 5 Scheeren, H. A., Lelieveld, J., Roelofs, G. J., Williams, J., Fischer, H., de Reus, M., de Gouw, J. A., Warneke, C., Holzinger, R., Schlager, H., Klüpfel, T., Bolder, M., van der Veen, C., and Lawrence, M.: The impact of monsoon outflow from India and Southeast Asia in the upper troposphere over the eastern Mediterranean, *Atmos. Chem. Phys.*, 3, 1589-1608, [10.5194/acp-3-1589-2003](https://doi.org/10.5194/acp-3-1589-2003), 2003.
- 10 Schiller C.I., Bozem H., Gurk C., Parchatka U., Königstedt R., Harris G.W., Lelieveld J., and Fischer H.: Applications of quantum cascade lasers for sensitive trace gas measurements of CO, CH₄, N₂O and HCHO, *Appl. Phys. B*, 92, 419-430, 2008.
- 15 Sprung, D., Jost, C., Reiner, T., Hansel, A., and Wisthaler, A.: Acetone and acetonitrile in the tropical Indian Ocean boundary layer and free troposphere: Aircraft-based intercomparison of AP-CIMS and PTR-MS measurements, *Journal of Geophysical Research: Atmospheres*, 106, 28511-28527, [10.1029/2000JD900599](https://doi.org/10.1029/2000JD900599), 2001.
- 20 Sprung, D., and Zahn, A.: Acetone in the upper troposphere/lowermost stratosphere measured by the CARIBIC passenger aircraft: Distribution, seasonal cycle, and variability, *Journal of Geophysical Research: Atmospheres*, 115, [10.1029/2009JD012099](https://doi.org/10.1029/2009JD012099), 2010.
- Stohl, A., Hittenberger, M., and Wotawa, G.: Validation of the lagrangian particle dispersion model FLEXPART against large-scale tracer experiment data, *Atmospheric Environment*, 32, 4245-4264, [http://dx.doi.org/10.1016/S1352-2310\(98\)00184-8](http://dx.doi.org/10.1016/S1352-2310(98)00184-8), 1998.
- 25 Streets, D. G., Yarber, K. F., Woo, J.-H., and Carmichael, G. R.: Biomass burning in Asia: Annual and seasonal estimates and atmospheric emissions, *Global Biogeochem. Cy.*, 17, 1099, [doi:10.1029/2003GB002040](https://doi.org/10.1029/2003GB002040), 2003.
- 30 Tadic, I., Parchatka, U., Königstedt, R., and Fischer, H.: In-flight stability of quantum cascade laser based on infrared absorption spectroscopy measurements of atmospheric carbon monoxide, submitted to *Appl. Phys. B*, 2017.
- 35 Umezawa, T., Baker, A. K., Oram, D., Sauvage, C., O'Sullivan, D., Rauthe-Schöch, A., Montzka, S. A., Zahn, A., and Brenninkmeijer, C. A. M.: Methyl chloride in the upper troposphere observed by the CARIBIC passenger aircraft observatory: Large-scale distributions and Asian summer monsoon outflow, *Journal of Geophysical Research: Atmospheres*, 119, 5542-5558, [10.1002/2013JD021396](https://doi.org/10.1002/2013JD021396), 2014.
- 40 van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Kasibhatla, P. S., and Arellano Jr., A. F.: Interannual variability in global biomass burning emissions from 1997 to 2004, *Atmos. Chem. Phys.*, 6, 3423-3441, [doi:10.5194/acp-6-3423-2006](https://doi.org/10.5194/acp-6-3423-2006), 2006.
- 45 van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P. S., Morton, D. C., DeFries, R. S., Jin, Y., and van Leeuwen, T. T.: Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009), *Atmos. Chem. Phys.*, 10, 11707-11735, [doi:10.5194/acp-10-11707-2010](https://doi.org/10.5194/acp-10-11707-2010), 2010.
- 50 Velders, G. J. M., Fahey, D. W., Daniel, J. S., McFarland, M., and Andersen, S. O.: The large contribution of projected HFC emissions to future climate forcing, *Proceedings of the National Academy of Sciences*, 106, 10949-10954, [10.1073/pnas.0902817106](https://doi.org/10.1073/pnas.0902817106), 2009.
- 55 Wang, S., Schmidt, J. A., Baidar, S., Coburn, S., Dix, B., Koenig, T. K., Apel, E., Bowdalo, D., Campos, T. L., Eloranta, E., Evans, M. J., DiGangi, J. P., Zondlo, M. A., Gao, R.-S., Haggerty, J. A., Hall, S. R., Hornbrook, R. S., Jacob, D., Morley, B., Pierce, B., Reeves, M., Romashkin, P., ter Schure, A., and Volkamer, R.: Active and widespread halogen chemistry in the tropical and subtropical free troposphere, *Proceedings of the National Academy of Sciences*, 112, 9281-9286, [10.1073/pnas.1505142112](https://doi.org/10.1073/pnas.1505142112), 2015.
- 60 Wendisch, M., Pöschl, U., Andreae, M. O., Machado, L. A. T., Albrecht, R., Schlager, H., et al.: The ACRIDICON-CHUVA campaign: Studying tropical deep convective clouds and precipitation over Amazonia using the new German research aircraft HALO, *B. Am. Meteorol. Soc.*, 97, 1885–1908, [doi:10.1175/BAMS-D-14-00255.1](https://doi.org/10.1175/BAMS-D-14-00255.1), 2016.
- Williams, J.: Organic Trace Gases in the Atmosphere: An Overview, *Environmental Chemistry*, 1, 125-136, <https://doi.org/10.1071/EN04057>, 2004.
- Williams, J., Yassaa, N., Bartenbach, S., and Lelieveld, J.: Mirror image hydrocarbons from Tropical and Boreal forests, *Atmos. Chem. Phys.*, 7, 973-980, [10.5194/acp-7-973-2007](https://doi.org/10.5194/acp-7-973-2007), 2007.



Yassaa, N., Song, W., Lelieveld, J., Vanhatalo, A., Bäck, J., and Williams, J.: Diel cycles of isoprenoids in the emissions of Norway spruce, four Scots pine chemotypes, and in Boreal forest ambient air during HUMPPA-COPEC-2010, *Atmos. Chem. Phys.*, 12, 7215-7229, doi:10.5194/acp-12-7215-2012, 2012.

- 5 Yokelson, R. J., Andreae, M. O., and Akagi, S. K.: Pitfalls with the use of enhancement ratios or normalized excess mixing ratios measured in plumes to characterize pollution sources and aging, *Atmos. Meas. Tech.*, 6, 2155-2158, 10.5194/amt-6-2155-2013, 2013.