



1 **A simple insect removal algorithm for 35-GHz cloud radar** 2 **measurements**

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11 **Abstract.** One of the key parameters that must be included in the analysis of atmospheric constituents (gases and
12 particles) and clouds is the vertical structure of the atmosphere. Therefore high-resolution vertical profile
13 observations of the atmospheric targets are required for both theoretical and practical evaluation and as inputs to
14 increase accuracy of atmospheric models. Cloud radar reflectivity profiles can be an important measurement for the
15 investigation of cloud vertical structure in a resourceful way. However, extracting intended meteorological cloud
16 content from the overall measurement often demands an effective technique or algorithm that can reduce error and
17 observational uncertainties in the recorded data. In this work a technique is proposed to identify and separate cloud
18 and non-hydrometeor returns from a cloud radar measurements. Firstly the observed cloud reflectivity profile must
19 be evaluated against the theoretical radar sensitivity curves. This step helps to determine the range of receiver noise
20 floor above which it can be identified as signal or an atmospheric echo. However it should be noted that the signal
21 above the noise floor may be contaminated by the air-borne non-meteorological targets such as insects, birds, or
22 airplanes. The second step in this analysis statistically reviews the continual radar echoes to determine the signal de-
23 correlation period. Cloud echoes are observed to be temporally more coherent, homogenous and have a longer de-
24 correlation period than insects and noise. This step critically helps in separating the clouds from insects and noise
25 which show shorter de-correlation periods. The above two steps ensure the identification and removal of non-
26 hydrometeor contributions from the cloud radar reflectivity profile which can then be used for inferring unbiased
27 vertical cloud structure. However these two steps are insufficient for recovering the weakly echoing cloud
28 boundaries associated with the sharp reduction in cloud droplet size and concentrations. In the final step in order to
29 obtain intact cloud height information, identified cloud echo peak(s) needs to be backtracked along the either sides
30 on the reflectivity profile till its value falls close to the mean noise floor. The proposed algorithm potentially identify
31 cloud height solely through the characterization of high resolution cloud radar reflectivity measurements with the
32 theoretical echo sensitivity curves and observed echo statistics for the cloud tracking (TEST). This technique is
33 found to be more robust in identifying and filtering out the contributions due to insects and noise which may



34 contaminate a cloud reflectivity profile. With this algorithm it is possible to improve monsoon tropical cloud
35 characterization using cloud radar.

36

37 1.0 Introduction

38 Cloud vertical structure (CVS) associated with the monsoon cloud systems is a key parameter needs to be
39 understood specifically in the light of its contrasting seasonal features. This allows in unraveling the role of CVS
40 for better characterization of various types of tropical monsoon cloud systems viz., deep, shallow convective clouds
41 and stratiform cloud systems. The high-resolution vertical profiling measurements of cloud radar are useful resource
42 in understanding of clouds and its evolution. Furthermore this technique will aid in the statistical characterization of
43 CVS associated with the Indian Summer Monsoon (ISM) and its intra-seasonal variability. The cloud radar system
44 uses short millimeter-wavelengths of 3.1 and 8.5 mm which correspond to frequency is ~94 and ~35 GHz
45 respectively. These wavelengths allow the detection of small cloud droplets and ice particles more sensibly at high
46 spatial and temporal resolutions that is used to infer important information on microphysical and dynamical structure
47 of cloud (e.g., Lhermitte, 1987; Frisch et al., 1995; Kollias and Albrecht, 2000; Sassen et al., 1999; Hogan et al.,
48 2005). Cloud radar, especially 35 GHz, is not only sensitive to hydrometeors (cloud particles and rain drops) but
49 also to air-borne biological targets such as birds and insects, and waste plant materials e.g., dry leaves, pollen or dust
50 (also known as “atmospheric plankton” or atmospheric “biota” or simply “insects”; Lhermitte, 1966; Teschke et al.,
51 2006). Although insects are probably the principal contaminants because of their size and dielectric constant,
52 spiders, spider webs, and other organic materials have been detected in the atmosphere through the use of nets and
53 other means (Sekelsky et al., 1998). Furthermore due to reduced scattering efficiency in the Mie region, cloud radar
54 observations at 95 GHz are found to be less sensitive to insects than observations at 35 GHz (Khandwalla et al.,
55 2003). Cloud radar signals frequently encounter this biota, within a couple of kilometers altitude close to the Earth
56 surface, confined mostly to the Atmospheric Boundary Layer (ABL). These echoes from the insect in the ABL have
57 reflectivity values comparable to those from the clouds and precipitation, and thus they contaminate and mask the
58 true cloud returns (Luke et al., 2008). The identification and removal of returns from such non-meteorological
59 targets (biota and receiver noise) is one of the prime tasks that is required to perform before using the meteorological
60 (cloud and precipitation) returns received by the cloud radar data, for the research and analysis purpose. The current
61 work focuses on identifying and filtering non-hydrometeor echoes in order to significantly improve the quality of
62 cloud radar data. This allows for the improved characterization of the tropical CVS.

63

64 Review of previous studies shows that different techniques have been attempted to remove non
65 meteorological echoes, for example, static techniques for the ground clutter (Harrison et al., 2014; 2000), return
66 signal-level correction (Doviak and Zrníć, 1984; Torres and Zrníć, 1999; Nguyen et al., 2008), dynamic filtering
67 (Steiner and Smith, 2002), and operational filtering (Alberoni et al., 2003; Meischner et al., 1997). The



68 aforementioned studies were mostly confined with the use of single polarization radar. However a new possibility
69 has been developed using dual-polarization information to identify the non-meteorological clutter echoes (Zrníc´ and
70 Ryzhkov, 1998; Mueller, 1983; Zhang et al., 2005). With the advent in Doppler spectral processing, it is possible to
71 have improved clutter mask (Bauer-Pfundstein and Görsdorf, 2007; Luke et al., 2008; Warde and Torres, 2009;
72 Unal, 2009). As mentioned one of the non-meteorological echoes is due to the insects and air-borne biota and these
73 unwanted echoes are problematic for studies involving meteorological information such as wind measurements
74 (Muller and Larkin, 1985) and true cloud returns (Martner and Moran, 2001). As a consequence, observations of
75 insects were done using variable polarization and multiple frequency radars operating initially in the centimeter
76 wavelength (Hajovsky et al., 1966; Hardy et al., 1966; Mueller and Larkin, 1985). At millimeter wavelength radar,
77 Bauer-Pfundstein and Görsdorf (2007) showed effective LDR filtering of insects while Khandwalla et al. (2003) and
78 Luke et al. (2008) showed that dual-wavelength ratio filters are more effective than the linear depolarization ratio
79 filters. Dual-polarization also offers a wide variety of methods (e.g., Gourley et al., 2007; Hurtado and Nehorai,
80 2008; Unal, 2009; Chandrasekar et al., 2013). Fuzzy logic classification techniques for the identification and
81 removal of spurious echoes from radar are also in use (e.g., Cho et al., 2006; Dufton and Collier, 2015). From the
82 above summary, it is therefore evident that most of the studies either concentrate on the polarimetric capabilities of
83 radar or off-line spectral processing of radar data to filter out echoes contaminated by non-meteorological targets.
84 The importance of the current work presented here lies in the development of an algorithm that uses solely high
85 spatial and temporal resolution reflectivity measurements. These high spatial and temporal resolution (25 m and 1
86 sec) measurements enable the characterization of irregular echoes associated with the spurious nature of radar
87 returns due to insects. This method is simple and does not require spacious complex spectral data (and associated
88 complicated analysis) or expensive advanced dual-polarimetric or dual-wavelength techniques.

89

90 **2.0 System, Data and Methodology**

91 This investigation employs vertically oriented observations of IITM's Ka-band scanning polarimetric radar
92 (KaSPR) for the study of vertical cloud structure. KaSPR has been providing high resolution (25 m and 1 sec.)
93 resourceful measurements of cloud and precipitation at a tropical site (Mandhardev, 18.0429° N 73.8689° E, 1.35 km
94 AMSL) on a mobile platform since June, 2013. Its main technical features are given in Table 1. KaSPR possess
95 sensitivity of above -60 (-45) dBZ at 1(5) km, it is therefore sensitive to the cloud droplet. According to T-matrix
96 computations, single 0.1 mm size of target at ~35 GHz may have the reflectivity ~ -60 dBZ whereas near one
97 million (63) of 0.01 (0.05) mm size is required to give the same reflectivity. Furthermore in one second if there are
98 500 (pulses per second) hits on the target in the radar scattering volume, the mean of those 500 samples at a range
99 bin (height) will be affected by the mean characteristics of target such as composition, orientation, number density
100 and kinematics associated with it. Therefore it is safer to assume that the atmospheric or meteorological targets (in
101 this case cloud particle) are distributive in nature and passive in the sense that their motion and/or orientation are in
102 resonance with the kinematics of the background atmosphere. By comparison birds and insects are point targets in
103 nature and active in the sense that they can change their motion, direction and orientation within a few seconds. This



104 leads to the irregular nature of intermittent or spurious radar returns characteristic of atmospheric biota due to the
105 much smaller de-correlation time associated with them. This study utilizes the high resolution profile of cloud radar
106 reflectivity factor (Z) to construct the cloud vertical structures by filtering out the returns from the noise and insects.

107 Figure 1a represents the height profiles of Z on 27 Apr 2014 at 2303 UT with various theoretical radar
108 sensitivity (noise-equivalent reflectivity, NER) curves (S0-S5; the range profile correction with the start range
109 sensitivity value of reflectivity, i.e., $r^2 Z_{\text{start range}}$, where r is range or height and Z is reflectivity, for S1, Z is -60
110 dBZ, for example). These different NER or sensitivity curves are utilized to qualify the observed radar returns that
111 are indeed above the NER, the inherent radar receiver noise level. The receiver noise level is the inherent thermal
112 noise associated with electronic components in the receiver chain and it remains approximately constant over the
113 length of the pulse returns. However, range correction is intuitive in the radar equation due to the decrease in echo
114 signal strength with increasing height (for vertical orientation). In order to determine the noise range in every range
115 bin, S0 to S5 are computed and overlaid on Z . This allows for identification and characterization of the signal that
116 overlays the background system noise level. As discussed earlier, the signal at any level may have contributions due
117 to either volumetric meteorological cloud particulates and/or strong non-meteorological point targets (e.g. biota). In
118 Figure 1a the echoes at ~ 3.7 km and below 2 km can be marked as cloud and insects respectively as it exceeds the
119 profile S5. The noise variations around 15 dB are mostly confined in between S0 and S2 with S1 as mean NER.
120 Contrasting echo texture associated with the cloud and atmospheric biota (hereafter insect) is evident from the
121 height-time-intensity (HTI) plot of Z in Figure 1b. This is a weak cloud case having reflectivity ~ -38 dBZ at ~ 3.7
122 km altitude with the presence of intermittent, non homogeneous echo texture from the insects below 2.7 km altitude.
123 Near similar weak cloud case of -38 ± 2 dBZ at 5.4 km altitude is confirmed as cloud with the sharp increase in
124 relative humidity of $\sim 80\%$ at that altitude by collocated GPS-RS measurements but is not shown here. Insect echoes
125 are observed to be confined most densely below 1.7 km and fall in the reflectivity range of -50 to -20 dBZ. The
126 observed standard deviation is always more than 2 and de-correlation period of ~ 4 -5 sec (returns due to insects is
127 found to vanish at an interval of ~ 3 -8 sec). Two sensitivity (S1 and S5) tests have been performed on Z profile to
128 quantify as the meteorological cloud returns. All the tests have been affected due to the presence of non-
129 meteorological echo due to insects even though these are mostly present in the ABL. Reflectivity values associated
130 with the cloud boundaries are very faint and are noticed to be fall within or close to system noise floor by 2-5 dB.
131 The profile S5 seems to be better in screening out the cloud echoes by 10 dBZ higher level than system mean noise
132 floor but this can eliminate significant portion of the weakest reflectivity area at the cloud edge (Figure 1d). Apart
133 from clouds, insects also show higher reflectivity values than S5. Figure 1d is similar to Figure-1b except, it is
134 completely screened out for cloud by applying typical threshold of radar system sensitivity profile, S1 and S5. In
135 addition to this, in case of Figure 1c, contiguous set of four reflectivity profiles have been considered for computing
136 running mean and standard deviation. The method followed to generate Figure 1c is the main objective of this paper
137 and is outlined by the flowchart in Figure 6. This method will be fully explained in the following section. In this
138 case, insect reflectivity values are similar to those of the cloud but their altitude levels are significantly different.
139 The contribution due to insects can therefore be removed by fixing with S5 and leaving the contribution due to



140 clouds untouched (Figure 1d). Thus, for the simultaneous presence of cloud and insect echoes at around same
141 altitude this NER method fails to identify the contributions separately. This NER method also fails whenever there
142 exist sharp reflectivity changes, usually seen with cloud boundaries/edges. This issue therefore demands the
143 development of a robust algorithm that explores the fundamental difference between cloud and insect returns so that
144 it could be identified and separated out these factors automatically.

145 In order to make the algorithm more robust for running it automatically, a close re-inspection of Figure 1b
146 infers that cloud returns are much more regular and near homogeneous when compared to insect's returns, which
147 appears to be spurious or intermittent in occurrence. Therefore, the NER criterion works reasonably well for the case
148 of homogeneous, isolated stable cloud layers but its robustness will be in question whenever there are vigorous and
149 quick changes associated with cloud edge and/or structure (will be explained in the discussion of cloud 1-2 in Figure
150 5). An additional criterion makes the current algorithm robust for complete revival of cloud information from the Z
151 observations by utilizing the de-correlation periods of insects (close to 3-5 sec). During this time interval significant
152 changes are not seen within the cloud. To explore this fact, in the next section the same weak low level cloud case
153 has been chosen further to understand the coherence period associated with cloud and insects.

154 3.0 Results and Discussions

155 Figure 2 takes the same case as in Figure 1 but confined below 4 km and 80-300 s, (left panel). Figure 2
156 reveals three main type of radar echo region namely (1) consistent radar returns characterized by the smooth and
157 gradual change(s) associated with cloud particles (at ~ 3.7 km height), (2) sharp (gradient) and spurious radar returns
158 (at altitude below 2.7 km) due to point target(s) and (3) receiver noise floor. In order to locate the above echo types
159 easily, various sensitivity or NER (i.e., S0-S5) curves have been utilized. The second type of echo is associated with
160 a characteristic point target (which has sharp reflectivity gradient feature due to the target's limited spatial as well as
161 temporal spread associated with the radar scattering volume). The third type, noise floor, is seen to be confined
162 mostly in between S0 and S2. The right panel in Figure 2 corresponds to HTI plot where the echo texture pertinent
163 to the above mentioned three echo types can be clearly visualized. The cloud echoes spreads in the altitude region
164 of approximately 300 m (3.6-3.9 km) with consistent smooth and gradual evolution with its weakest and/or broken
165 structure during 165-190s. In contrast to this the observed irregular point or rounded texture of insects echo spread
166 is seen to be limited temporally around 3-7 seconds and spatially below four range-bins size (i.e., < 100 m) with
167 strongest reflectivity at its center. This indicates that one second temporal resolution might be good enough to see
168 the insects as point or rounded echo texture. When biota density is more in the lower altitude levels, it is difficult to
169 clearly identify the boundary of one point target from another. Such a scenario, though rare, can lead to
170 misidentification as clouds. The coexistence of cloud and transient high density flocks of biota adds complexity
171 which becomes almost impossible to discriminate. However, this issue is observed to be rare and limited to lowest
172 altitudes only.



173 To investigate the similarities and contrasting features associated with various contributions to the cloud
174 reflectivity profile, it is important to explore further the case of Figure 1. Statistical de-correlation times associated
175 with three types of echo have been computed for their identification and separation. Both the cloud at ~3.7 km
176 narrow region and insect returns below ~ 1.5 km in Figure 3 are evident above the maximum noise level. Both cloud
177 and insect parts of the Z profiles are expanded to allow for review of the mean (Figure 3b and 3d) and standard
178 deviation (σ ; Figure 3c and 3e) of Z for every set of consecutive 15 profiles. Figure 3b shows the patterns of the
179 seven mean cloud reflectivity profiles are organized and more consistent or correlated to one another during 105
180 seconds, this is in comparison to less organized reflectivity profiles due to insects that are much less consistent or
181 correlated with one another in figure 3d. Moreover, the corresponding seven σ profiles show differences for cloud
182 that is less than 1.5 σ (figure 3c). By comparison differences in profiles due to insects are more than 4.0 σ most of
183 the time (figure 3e). It is seen that the mean cloud reflectivity peak values gradually extend from 3.7 to 3.8 km
184 where the corresponding standard deviation values are less than 1 σ . In order to further test the minimum de-
185 correlation time associated with cloud and biota, the averaging time is reduced to a set of 5 profiles (5 sec) with the
186 same data (see Figure 4). In this case also, Figure 4c depicts σ for all the seven mean cloud reflectivity profiles are
187 below 1.5 with peak <1 σ . This manifests that volumetric distribution nature of cloud particles is statistically more
188 homogeneous or show less dispersion. However, Z values associated with biota show random behavior with
189 significant dispersion >1.5 σ (Figure 4e). This high dispersion in the Z values infers that the echo due to biota de-
190 correlates quickly within ~5 second time interval (see Figure 4d-4e). It is seen from Figure 3 that for vertical levels
191 from 0.9 km to 1.5, the sharp peaks in reflectivity profiles and strong dispersion of > 3 σ are associated with the
192 return from biota. This is attributed mostly to the observed intermittent point target nature of insect echoes plausibly
193 due to the rambling or meandering motion of insects within the radar sampling volume. Moreover, the inherent radar
194 system noise (random in nature) dispersion is observed to be in between the cloud and biota (1.5-3.0 σ). It is evident
195 from the top panels of Figure 3-4 that cloud reflectivity profiles show relatively consistent trend and correlation
196 among the contiguous mean profiles computed from the set of 15 Z profiles than computed from the 5 profiles. This
197 may be mainly due to the homogeneities or in-homogeneities associated within the chosen data sets those are
198 independent to one and another. Therefore, in order to preserve the real time sequence of observations for the study
199 of cloud evolution as well as to recover underlying smooth trends pertinent to natural clouds, a four-point moving or
200 running average is applied on the time series of Z data instead of deriving a simple average. The four seconds is
201 the optimal moving average time for yielding the best cloud results (Figure 5) by characterizing the cloud to insect
202 echoes coherent to incoherent property during the moving average period. By this four point running average, insect
203 echo become incoherent due to its short de-correlation period (~4 sec) whereas those echoes de-correlating over
204 longer periods indicate the presence of clouds. To understand the degree of dispersion, along with σ the absolute
205 deviations in mean and median values have also been analyzed. Their relation with σ is seen to be as mean absolute
206 deviation slightly smaller than σ as $\sigma/1.253$ where as median absolute deviation smallest as $\sigma/1.483$. This work
207 makes use of the statistical mean and σ but using above relation one can relates the present results with other
208 statistical central tendencies of data distribution. Next, the filtering of noise and insects from the presence of cloud
209 using the cloud radar reflectivity profile will be explored. The segregation has been carried out using theoretical



210 radar echo sensitivity curves and statistically computed echo de-correlation periods and finally tracking the cloud
211 echo peak to its adjacent sides till it is close to the S1 profile for the cloud height. The above set of tasks, Theoretical
212 Echo Sensitivity and observed Echo based Statistics for cloud height Tracking (TEST), is repetitively performed on
213 the cloud radar Z measurements under an algorithm whose flowchart can be seen in Figure 6. The algorithm used in
214 this work is named as TEST and can be summarized below:

- 215 1. Wherever the moving mean Z values in the profile are equal to or above the S5 can be qualified as cloud or
216 insect. This step ensures removal of the system noise floor.
- 217 2. Those altitude regions of the qualified echo are then further scrutinized to identify clouds using the
218 minimum thickness of greater than 100 m (to strictly avoid biota that are found to extend less than 2-4
219 height bins each of 25 m) and mean standard deviation below 1.5σ .
- 220 3. In order to keep the identified cloud's structure, intact, the identified cloud peak(s) are tracked back on
221 either side (towards upper and bottom heights) up to around (preferably 1-2 dBZ) the mean noise profile
222 S1.

223 It is interesting to note that the cloud echo regions are always stronger and above the mean noise fluctuations i.e.,
224 S1. Therefore at the left side of the curve, S0 to S1, always appears as a void region in the 2-dimensional reflectivity
225 plot wherever there is a presence of cloud, no matter weak or strong (just below 4 km in the left panel of Figure 1
226 and 3). This causes sharp boundary gradients between cloud and noise in the vertical profiles of Z and hence with
227 the corresponding σ . This can be used as a visual criterion for detection of cloud.

228 Figure 7 is similar to Figure 1 but it represents a multi layer pre-monsoon cloud system for the period 1200-
229 1205 UT, 29 May 2014. Various labeled altitude regions (biota, noise and cloud) of the vertical reflectivity structure
230 show typical mean features that can be broadly classified the returns into cloud and non-cloud (biota and noise)
231 portion. Furthermore, Figure 7 shows the typical variety of cloud layers existing within the vertical structure of
232 tropical cloud as well as morphological features pertinent to pre-monsoon thunderstorm activity. The cirrus layer at
233 12-14 km shows gradual structural change having peak reflectivity values of ~ 5 dBZ. Here, the high reflectivity
234 values contribute to form single deep convective cloud by merging with the cloud layer that exists at lower heights.

235 Figure 8a and 8b reveal the reflectivity time series associated with the labeled non-cloud and cloud portion
236 of Table 2 respectively. Noise and biota shows max 2 dB fluctuations around the 4-point-running mean reflectivity
237 whereas for biota the max fluctuation is 3-5 dB (bold solid line). It can be understood that noise values increase
238 gradually with altitude with σ values ~ 2.3 whereas sharp boundary gradients associated with biota and ragged
239 shallow cloud regions (cloud 1&2 in Figure 7) also show higher σ values > 3 . Stable or layer cloud regions (cloud 4
240 & 5 in Figure 7) show significantly standard deviation below 2σ . Further, it is interesting to examine the time series
241 plots for the contrasting variations between the insects and noise and cloud regions with Figures 8a and 8b. The
242 range of dBZ variability is 4-10 for insects and 2-4 for noise and for cloud that is less than 1 within an interval of 5-
243 10 seconds. The corresponding variability in standard deviation (S.D) is observed to be 4-10 σ for insects, 1.5-3.5 σ



244 for noise and $\sim 1 \sigma$ for cloud ($< 1 \sigma$ for cloud peak) except for weaker cloud regions. These statistical characteristics
245 of all types of observed cloud echoes have been tabulated in the Table 2.

246 Figure 9 demonstrates the application of the work presented here and illustrates the significant differences
247 between the uncorrected (Figure 9a) and corrected (Figure 9b) reflectivity profiles. The peaks in frequency
248 distribution of uncorrected cloud reflectivity profiles at just below -50 dBZ, in between -50 and -40 and just above -
249 40 dB are the predominant contributions from noise (middle panel of Figure 9a). These noise regions bias severely
250 the corresponding histogram frequency distribution at three different altitude levels that are associated with the
251 Johnson's tri-modal cloud distribution (extreme right panel of Figure 9a). In order to infer the distribution of cloud
252 reflectivity values in the various altitude regions pertinent to tri-modal cloud vertical structure (Johnson et al.,
253 1999), the observed vertical structure is subdivided into warm or low (< 3.6 km), mixed or mid ($3.6 \text{ km} \leq \text{altitude}$
254 ≤ 8.6 km) and ice or high (> 8.6 km) phase and/or level clouds. The plots of uncorrected reflectivity distribution
255 clearly shows skewness towards lowest values of reflectivity (below -50dB, -40 dB and -30 dB for low, mid and
256 high level respectively seen with right panels of Figure 9a). This is mainly due to the predominance of noise
257 contribution except for the low cloud regions where the contribution of insects is also included. After applying the
258 TEST algorithm the corrected reflectivity distribution peaks at -42dB, -35 dB and -22 dB for low, mid and high level
259 respectively (right panel of Figure 9b) reflects the actual scenario of the cloud system. This method is simple and
260 has potential to bring out the statistically significant micro- and macro-physical characteristics from meteorological
261 information (i.e., cloud) and hence for better characterization of the cloud vertical structure over a region.

262 In order to test the merit of the current algorithm on filtering out the non-meteorological contribution with
263 Z profile, the parametric thresholds on Pulse-Pair (PP) processed Z and few polarimetric variables profiles of the
264 cloud radar measurements have also been considered in place of usual Fast Fourier Transformation (FFT) process.
265 The FFT process is capable to provide only polarimetric parameter, i.e., linear depolarization ratio (LDR). Figure 10
266 is similar to the Figure 1 that illustrates FFT (top) and PP (bottom) processed Z profiles on 28 Aug 2014 but are 15
267 minutes apart from one another (0415 and 0400 UT respectively) which causes some dissimilarities in the observed
268 three layer cloud structure between the two plots (upper and lower panel). Minimum range of the noise floor in the Z
269 profiles (2-D plot in the first panel) is seen to be greater for PP than FFT processing. The TEST algorithm performs
270 in a similar way for both the FFT and PP processed Z profiles and is able to isolate the cloud structure as best as
271 possible. Figure 11 explores further the polarimetric capability of the KaSPR in separating out the meteorological
272 contribution with Z by using critical threshold on the PP-polarimetric measurements that correspond to the bottom
273 panels of Figure 10. The top panels of Figure 11 stand for HTI plots of, three polarimetric parameters namely, LDR,
274 Φ_{dp} and K_{dp} . Computation of LDR is inherently limited to the cross polar isolation of the radar system that is -27 dB
275 for KaSPR. Hence, high LDR values above -17 dB are mostly seen with insect and low LDR values below -17 dB
276 are seen with cloud. Low to lower LDR values (i.e., < -17 dB to -25 dB) are strictly confined within the peak values
277 of co-polar reflectivity (> -10 dB) of cloud altitude regions, ~ 8 -10 km. Except the inherent limitations associated
278 with LDR, these results are in agreement with earlier reported results (e.g. Bauer-Pfundstein and G6rsdorf, 2007 and



279 Khandwalla et al., 2003). The LDR, Φ_{dp} and K_{DP} threshold values are set below -17 dB, 56° and $-15^{\circ} \text{ km}^{-1}$
280 respectively, can be used to filter out biota from the corresponding Z profiles that are shown at lower panels of
281 Figure 11. The threshold used for Φ_{dp} and K_{DP} are subjective depending on the observed case for better filtering of
282 insects. These polarimetric threshold methods are although successful in filtering out the non-meteorological
283 contributions but they are bound to sacrifice the weaker portion of the cloud where polarimetric computations are
284 not perfect. Thus, polarimetric method is incapable to preserve the weaker portions of the whole cloud regions
285 where the TEST method is noticed to perform better (bottom right panel of Figure 10). This further proves the
286 efficiency of the proposed TEST method. This has implemented in the post-processing of high resolution reflectivity
287 measurements. The method developed here is far simpler and provides a superior solution to filtering out signal due
288 to noise and biota and preserve cloud data in the form of pure meteorological reflectivity measurements which can
289 be used to infer the true characteristics of clouds.

290 Figure 12a demonstrates further application of the current work on filtered cloud reflectivity profiles
291 (bottom plot) by considering the six hours evolution of variety of tropical cloud systems. On 21 May 2013, a typical
292 convective cloud system present during pre-monsoon season was observed. This event is composed of three
293 systems, first three hours (00:00-03:12 UT) shows stratiform cloud confirmed from bright band occurrence at an
294 altitude of 4 km AGL, convective system around 0500 UT, which is a cumulus congestus initially, and above it
295 cirrus (ice) cloud in the altitude range of 13-14 km. The screened out reflectivity profile can therefore be utilized to
296 fully characterize the tri-modal cloud episode as shown in Figure 12b. The mean reflectivity profile with standard
297 deviation bars reveals the nature of important phase change regions associated with cloud vertical structure. The
298 change in cloud processes in the cloud vertical structure is closely associated with the phase of cloud water that is
299 strongly linked with the predominant change of temperature.

300 4.0 Summary and Conclusions

301 High resolution vertically oriented reflectivity measurements of cloud radar are solely potential to understand the
302 cloud vertical structure after segregate the meteorological and non-meteorological contributions with it. Theoretical
303 noise equivalent reflectivity curves are used to remove the system noise. The simple statistical variance of continual
304 radar echoes show the contrasting different characteristic of signals like high dispersion (more than 2σ) is associated
305 with the highly spurious and intermittent echoes of insects and low dispersion (less than 1σ) is associated with
306 coherent nature of echoes of cloud hydrometeors and for noise it is 1.5-3.0 σ . Furthermore, these characteristic
307 features are mainly used to demarcate the returns of cloud hydrometeor to those from insects and noise. Running
308 mean and standard deviation of reflectivity profiles for ~4-5 seconds that works well to filter out all non-
309 hydrometeor returns. In this way, the de-correlation period associated with biota helps in identifying and filtering
310 out the insect returns. The proposed TEST algorithm evaluates the observed cloud radar reflectivity profiles with
311 combined theoretical radar sensitivity curves and statistical variance of radar echo and then tracks the cloud peak at
312 either side to obtain the complete cloud height profile. In case of azimuth and elevation radar surveillance scans (PPI
313 and RHI, for example), there is a regular change in the radar sampling area that disables to have exclusive set of



314 measurements required to perform the TEST method. But this method is advantageous and easily adaptable for
315 better characterization of any high-resolution vertical profile measurements. The robustness of TEST is also proved
316 through polarimetric methods and found that it works much better, particularly in the weak cloud region, at the
317 cloud radar frequencies. Such scrutinized reflectivity profiles has been further utilized to investigate the important
318 CVS pertinent to the various phases of the Indian Summer Monsoon with the aim of improved prediction. Hence,
319 the proposed TEST algorithm is able to extract the possible unbiased meteorological cloud vertical structure
320 information with the cloud profiling radar. This enables carrying out the pragmatically effective research
321 investigations on the seasonal and epochal tropical cloud characteristics.

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328



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Figure Captions

Figure 1. (a) Vertical looking cloud radar measured sample ten reflectivity height profiles on 27 April 2014 during 2303-2308 UT. S0 to S5 are the theoretical noise equivalent reflectivity curves with their respective threshold values in bracket. HTI plot of (b) the same reflectivity profile for the duration of 306 sec (c) screened out reflectivity profile for the receiver noise floor and the biota (insects) using running average constrained with standard deviation (d) constrained with NER (S5).

Figure 2. (left) Same as 1(a) but for 220 profiles. (right) HTI plot of Z profiles. Smoothly varying homogeneous cloud layer is at altitudes of 3.5-3.8 km and sharp, rounded and spurious kind of echoes below 2.7 km are due to biota.

Figure 3. (a) Same as 1(a) but for 105 profiles. (b) mean and (c) standard deviation of 15 profiles of Z pertinent to cloud height region (3.5-3.9 km) and (d) and (e) same as (b) and (c) but pertinent to insects height region (0.9-1.5 km).

Figure 4. Same as Figure 3 but for total duration 35 sec; the mean and standard deviation profiles are for every 5 second interval.

Figure 5. Same as Figure 3 but for total duration 10 sec; the mean and standard deviation profiles are for 4-point-moving average.

Figure 6. TEST algorithm flow chart that identifies and filter-out the insects and noise echoes for screening-out the cloud contributions with the Z measurements.

Figure 7. (a-c) Same as 1(a-c) but on 29 May 2014 during 1200-1205 UT for the duration of 306 sec. Statistics corresponds to the labels on the Z profile can be seen in Table 2.

Figure 8a. Time series of the mean and standard deviation (S.D) of Z for insects (bottom panels) and four noise floor regions as per Table 2. Bold solid lines are the 5-point-running mean over the actual time series data (lines with symbol).

Figure 8b. Same as Figure 8a but for the cloud regions as per Table 2.

Figure 9a. (Left panel) Uncorrected mean reflectivity profile on 29 May 2014 during 1200-1205 UT superimposed with curves S1 (dashed red line) and S5 (solid green line). Histogram of Z profile (Middle panel). (left three sub panels) for altitude regions of low (<3.6 km), mid (3.6 km ≤ ht < 8.6 km) and high (≥ 8.6 km). The right sub panels each peak of histogram are mapped on to the corresponding three peaks with the whole vertical structure of Z. This infers the noise clearly suppresses the meteorological information.



Figure 9b. Same as 9a but it is corrected by filtering out noise and biota. The correction applied to Z profile allows to pop-up the true meteorological cloud reflectivity distribution.

Figure 10. Same as 7 but for vertical looking KaSPR measurements at 0400 UT on 28 Aug 2014 using (top) FFT processing (bottom) 15 minutes prior one using PP processing. PP case will be used further to evaluate the polarimetric algorithm performance.

Figure 11. HTI plots of (top panel) LDR, Φ_{dp} and K_{DP} parameters pertinent to PP processed data of Figure 10 and (bottom panels) biota filtered reflectivity after applying corresponding polarimetric thresholds of the respective top panels.

Figure 12a. (Top) Same as Figure 7b (uncorrected) and (bottom) same as Figure 7c (corrected) but integrated for duration of 0000-0630 UT taken at an interval of ~ 15 minutes on 21 May 2013

Figure 12b. Same as Figure 9b but excluding middle panel for the corrected Z data of figure 12a.



Table 1: KaSPR specifications

Radar specifications	value
RF output frequency	35.29 GHz
Peak power	2.1 kW
Duty cycle	5 % max.
Pulse widths (selectable)	3.3 ns (50-13000 ns)
Pulse compression ratio	1:10 (1-100)
Transmit polarization	H or V-pol linear; Pulse-to-pulse polarization agility
Receiver polarization	Simultaneous Co- and Cross-polarization linear
Receiver noise figures	2.8 dB min
Sensitivity at 5.0 km	-45 dBZ
Tx & Rx losses	1.15 & 0.3 dB
IF output to digital receiver	90 MHz
Antenna diameter	1.2 m
Antenna Beam width	0.5 ⁰
Antenna gain (includes OMT loss)	49 dB
First side lobe level	-19 dBi min.
Cross-polarization isolation	-27 dB



Table 2: Statistical mean and standard deviation of cloud radar reflectivity corresponds to the selected height regions, which are labeled, on the Figure 7.

Label	Mean Z for 305 sec (4 sec) dBZ	σ for 305 sec (4 sec)
Insects (1.2-1.7 Km)	-54.1(-55.0)	4.08 (3.4)
Noise 1 (2.1-2.4 Km)	-52.9 (-52.4)	2.33 (1.9)
Noise 2 (5.9-6.2 Km)	-44.4 (-44.2)	2.22 (2.3)
Noise 3 (11.1-11.6 Km)	-39.1 (-39.1)	2.30 (2.2)
Noise 4 (14.7-15.2 Km)	-36.7 (-36.9)	2.29 (2.2)
Cloud 1 (3.7-3.9 Km)	-36.2 (-28.3)	5.99 (12.7)
Cloud 2 (4.8-5.1 Km)	-31.8 (-22.7)	5.54 (4.5)
Cloud 3 (6.8-7.2 Km)	-0.4 (0.3)	2.60 (3.5)
Cloud 4 (9.8-10.2 Km)	-10.9 (-9.9)	2.03 (3.1)
Cloud 5 (12.8-13.2 Km)	3.1 (1.4)	0.86 (1.0)

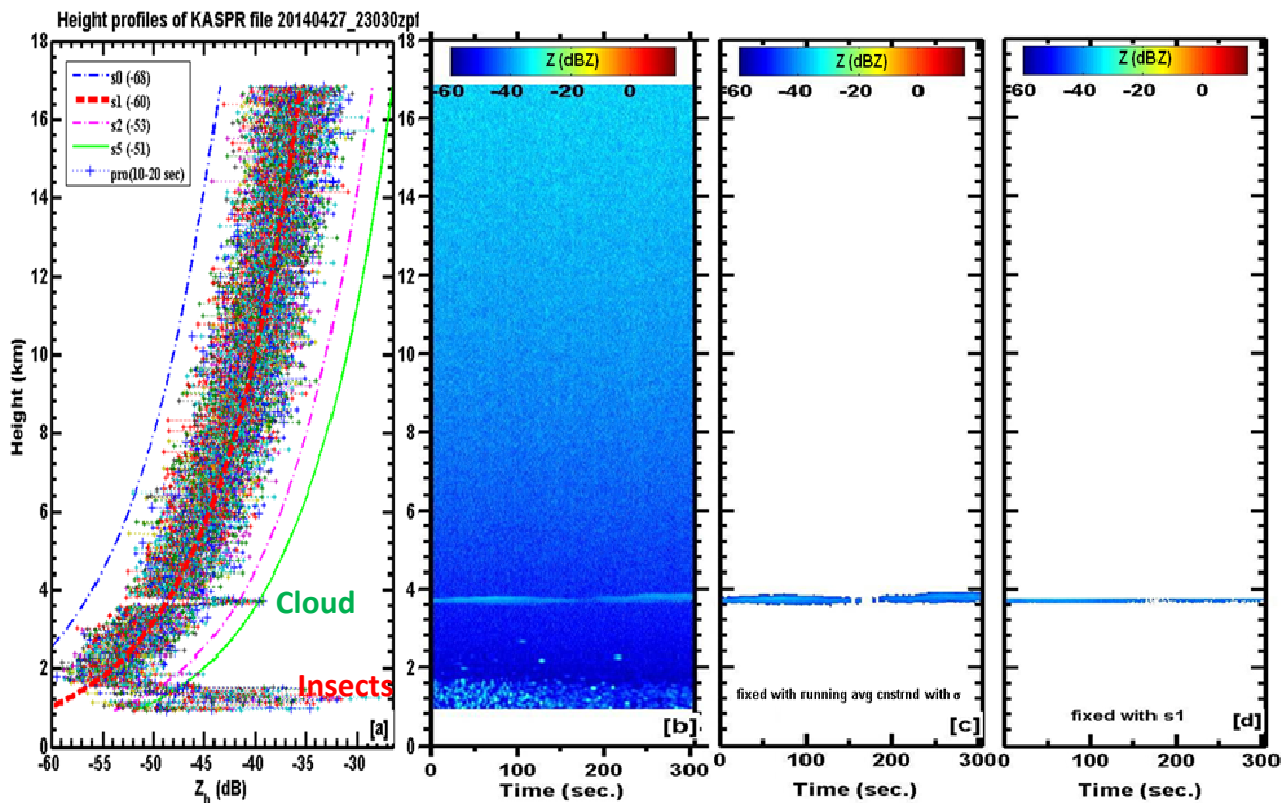


Figure 1: (a) Vertical looking cloud radar measured sample ten reflectivity height profiles on 27 April 2014 during 2303-2308 UT. S_0 to S_5 are the theoretical noise equivalent reflectivity curves with their respective threshold values in bracket. HTI plot of (b) the same reflectivity profile for the duration of 306 sec (c) screened out reflectivity profile for the receiver noise floor and the biota (insects) using running average constrained with standard deviation (d) constrained with NER (S_5).

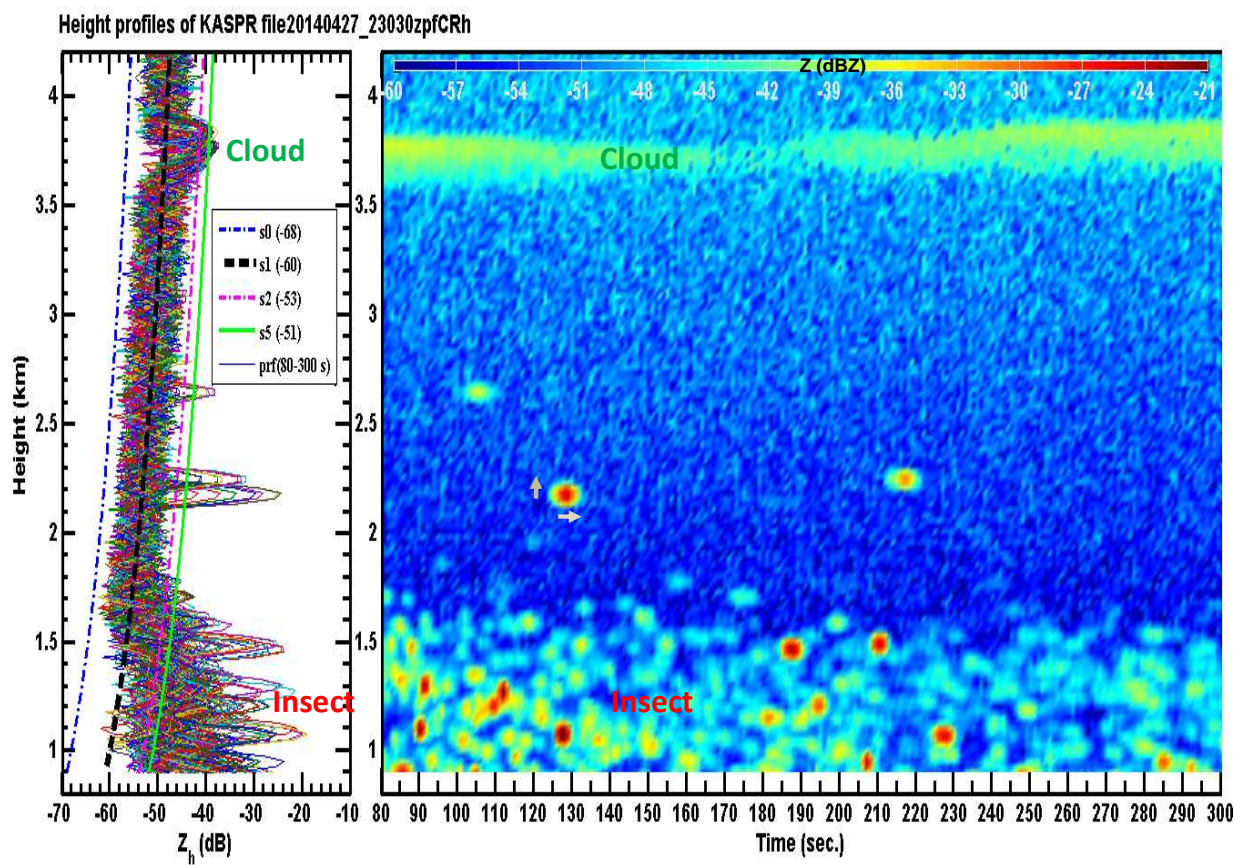


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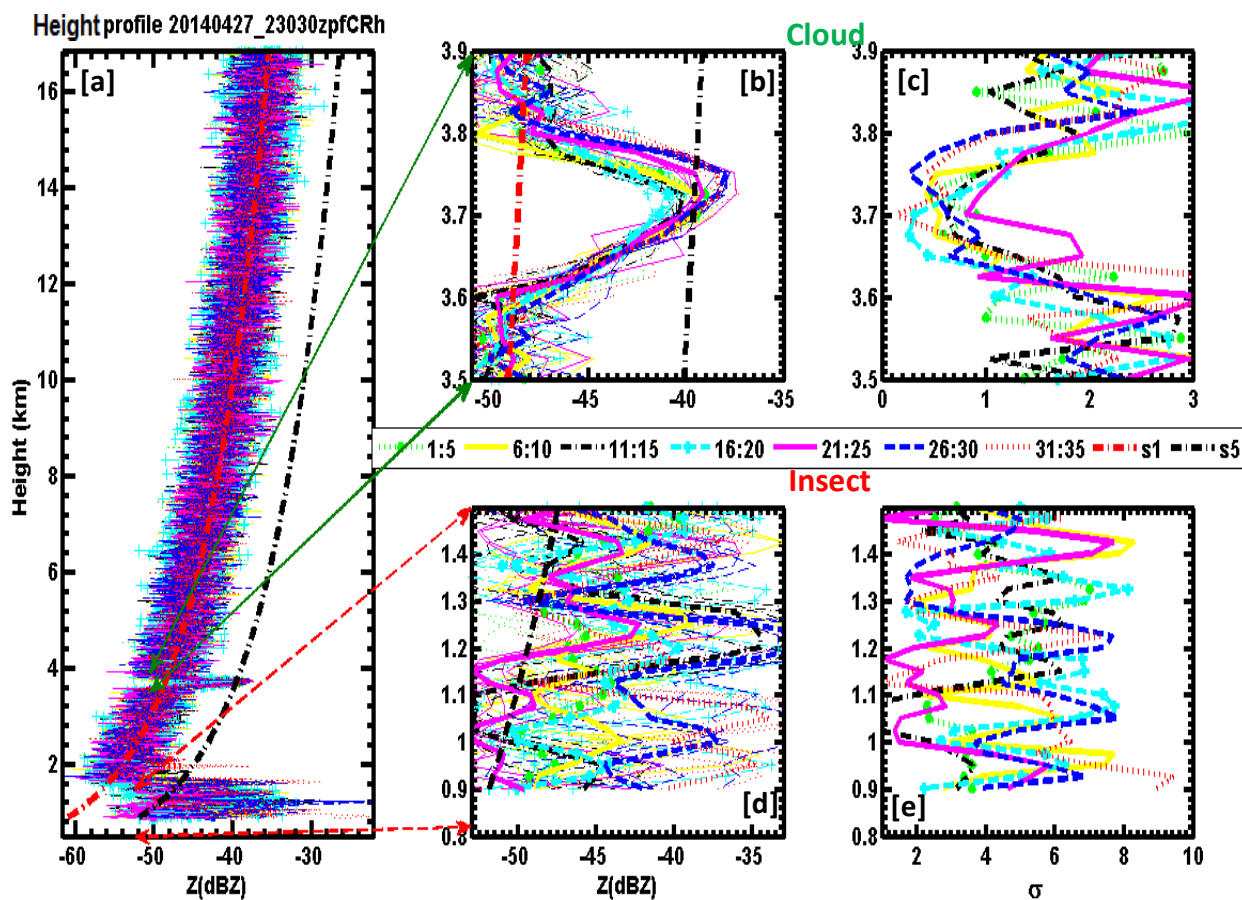


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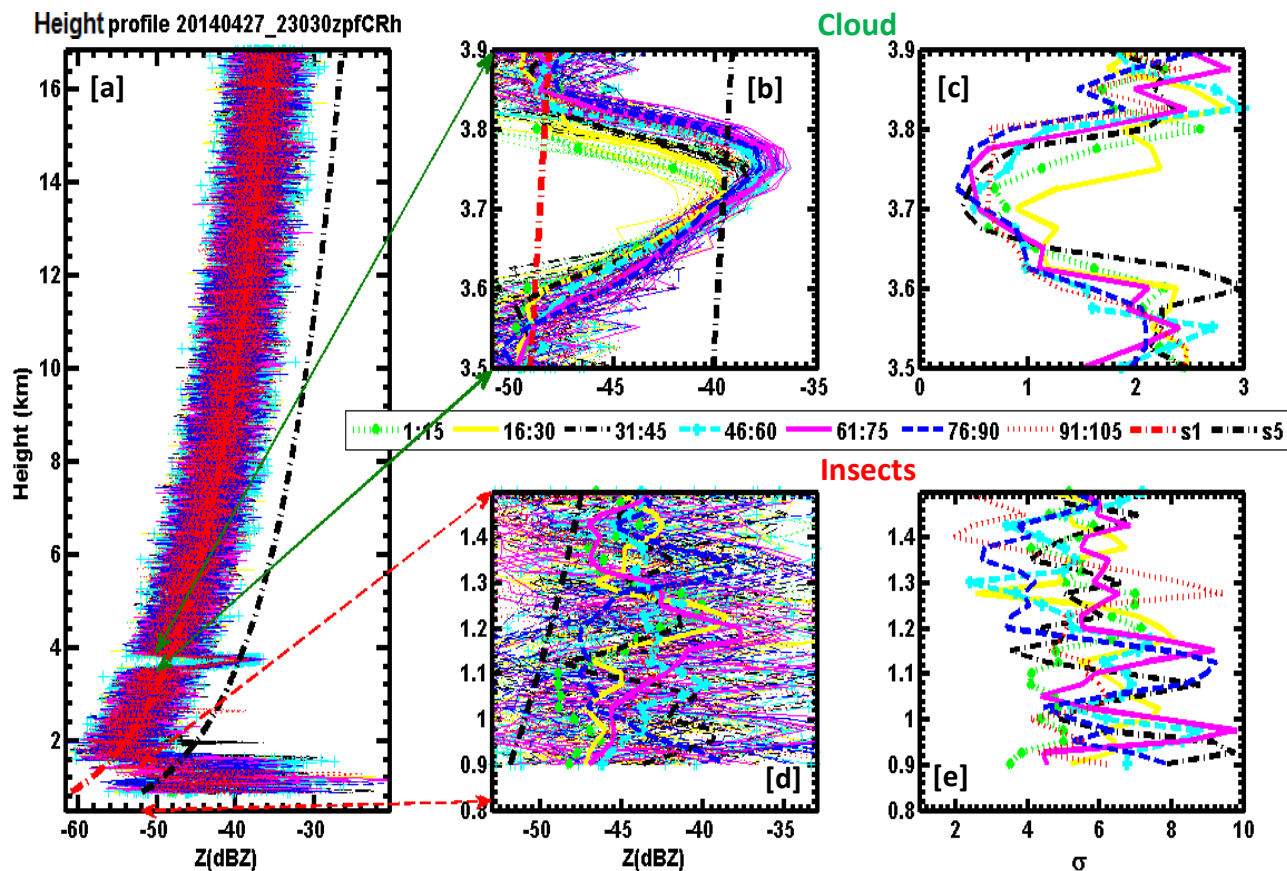


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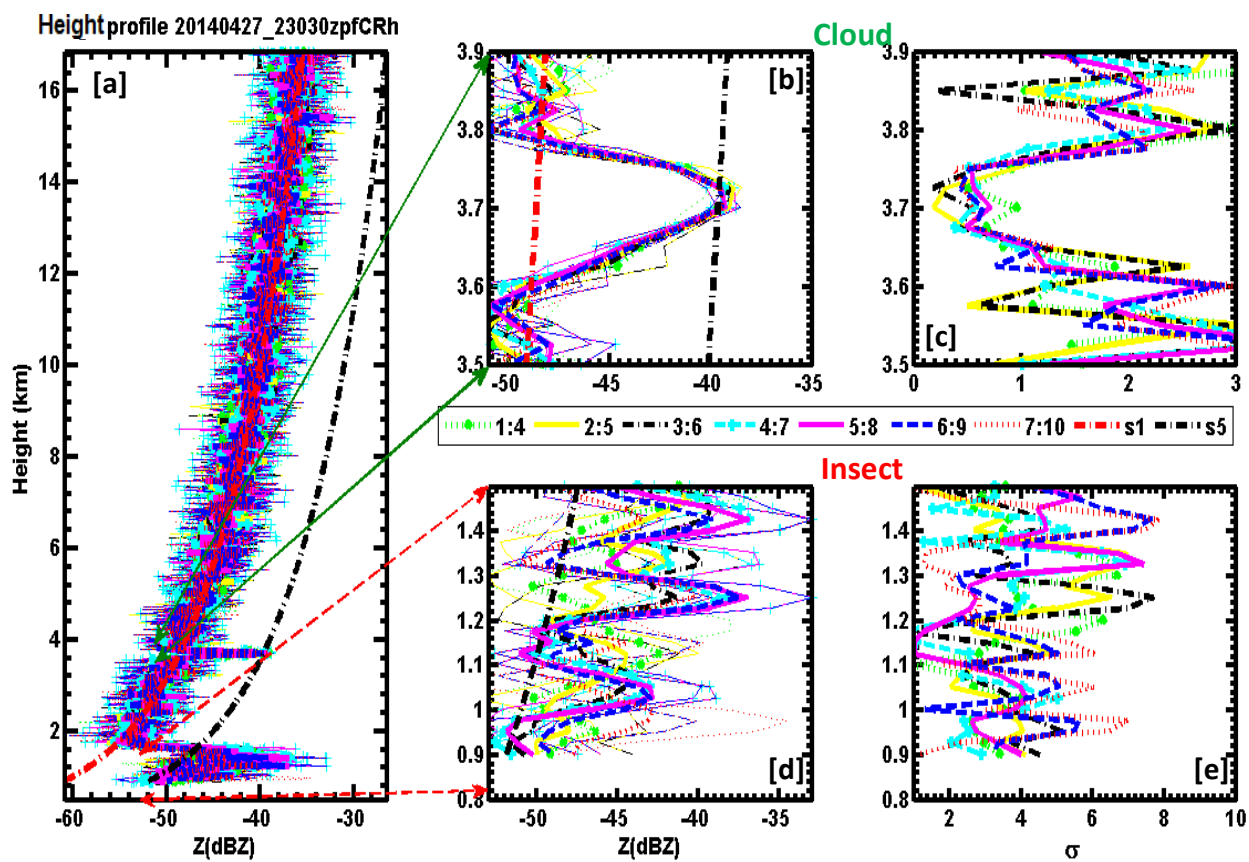


Figure 5: Same as Figure 3 but for total duration 10 sec; the mean and standard deviation profiles are for 4-point-moving average.

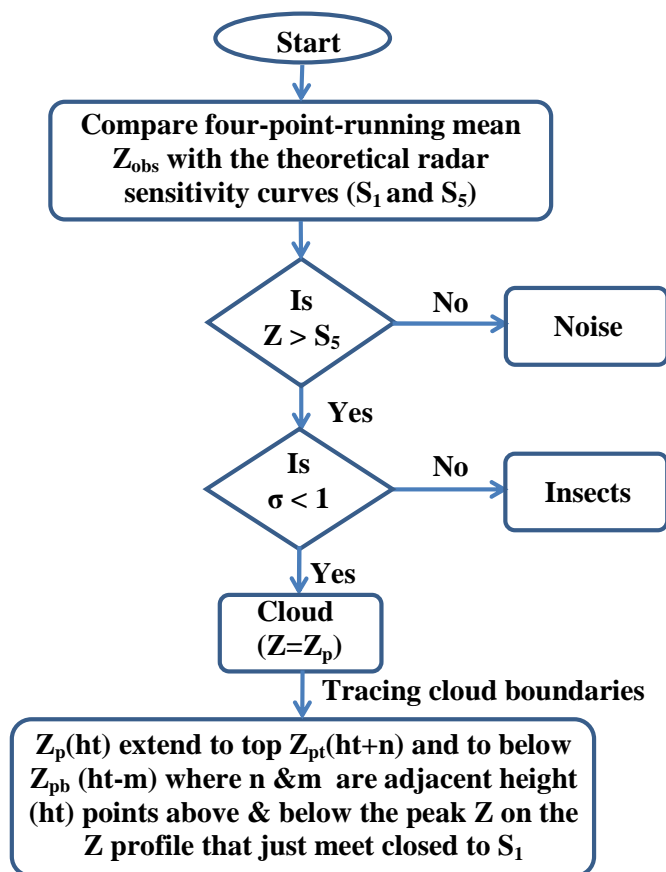


Figure 6: TEST algorithm flow chart that identifies and filter-out the insects and noise echoes for screening-out the cloud contributions with the Z measurements.

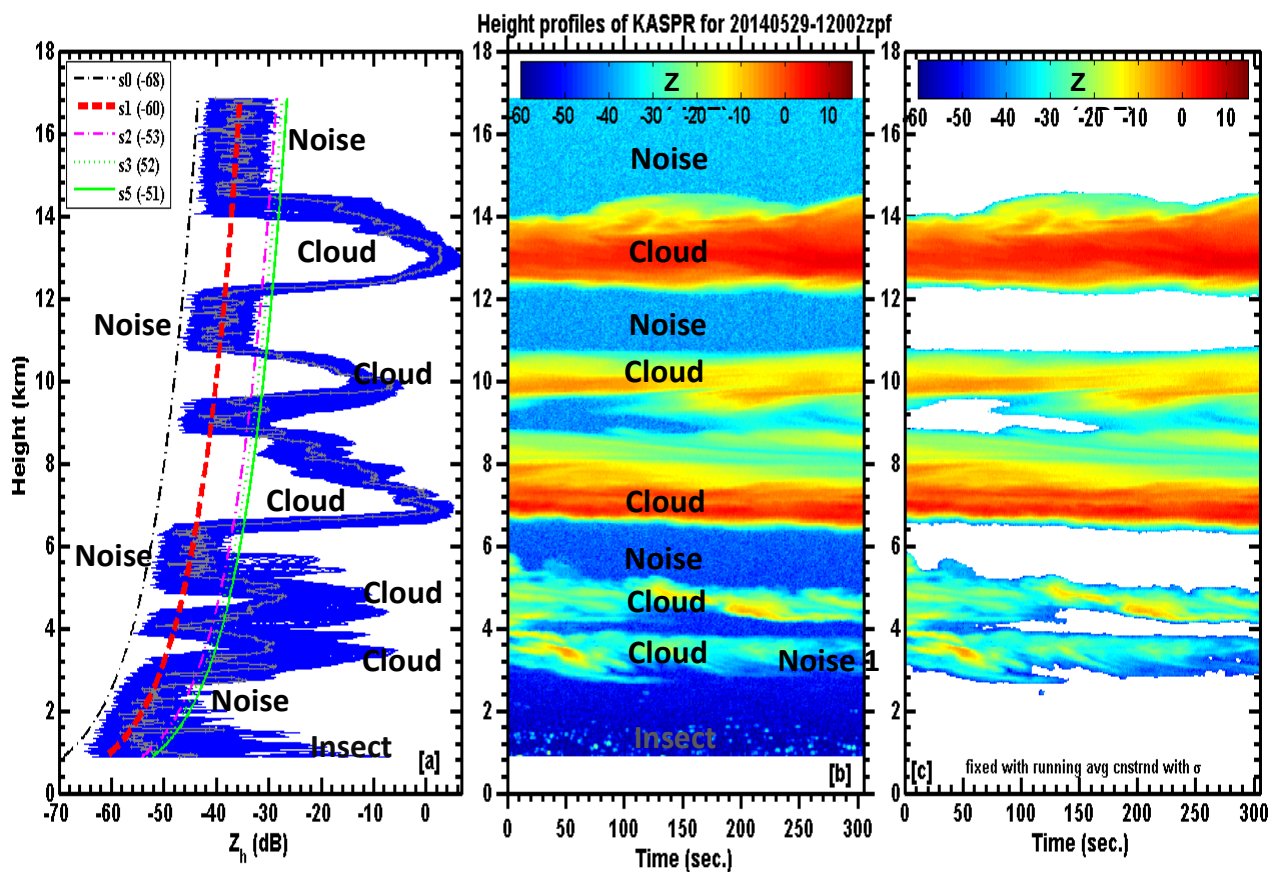


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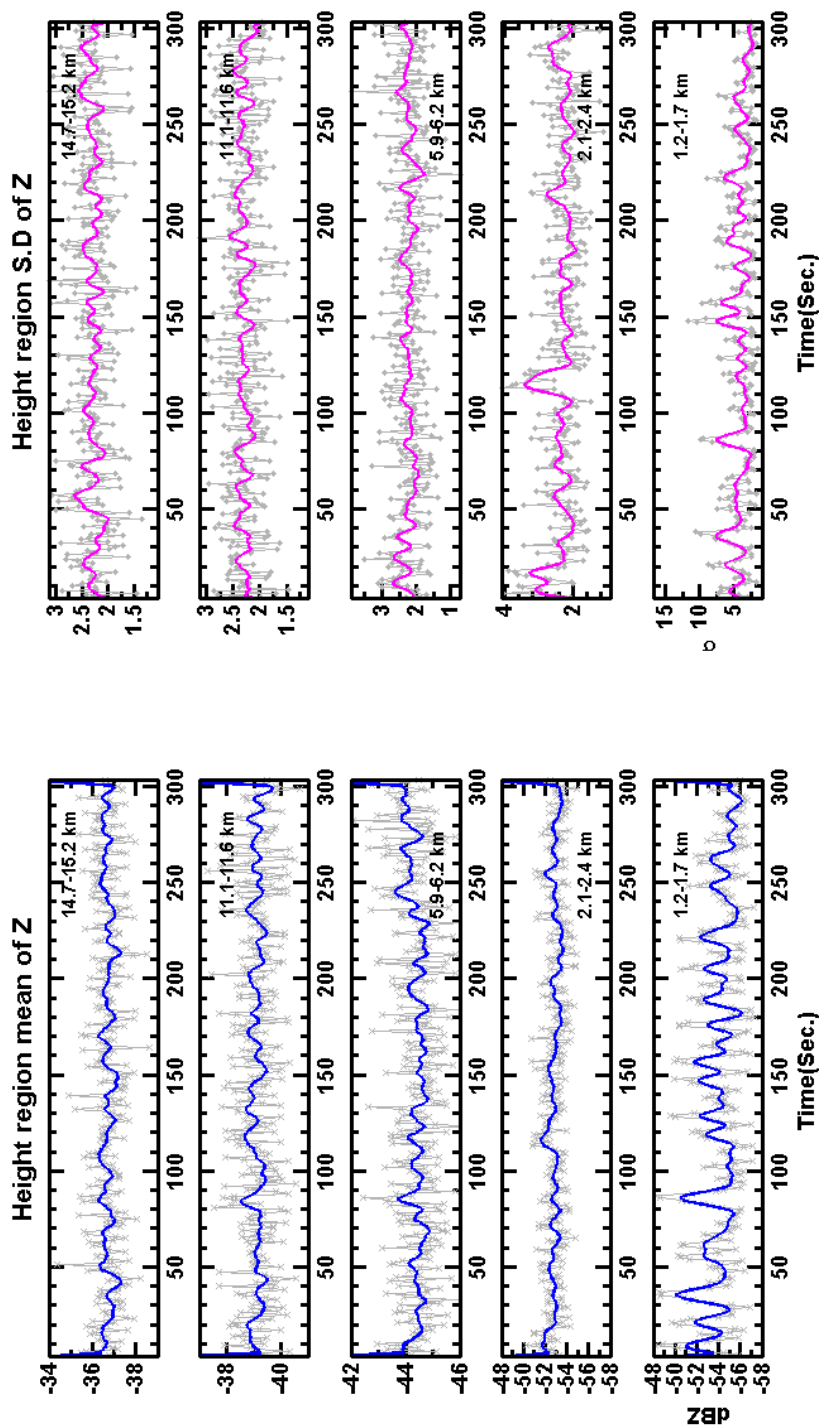


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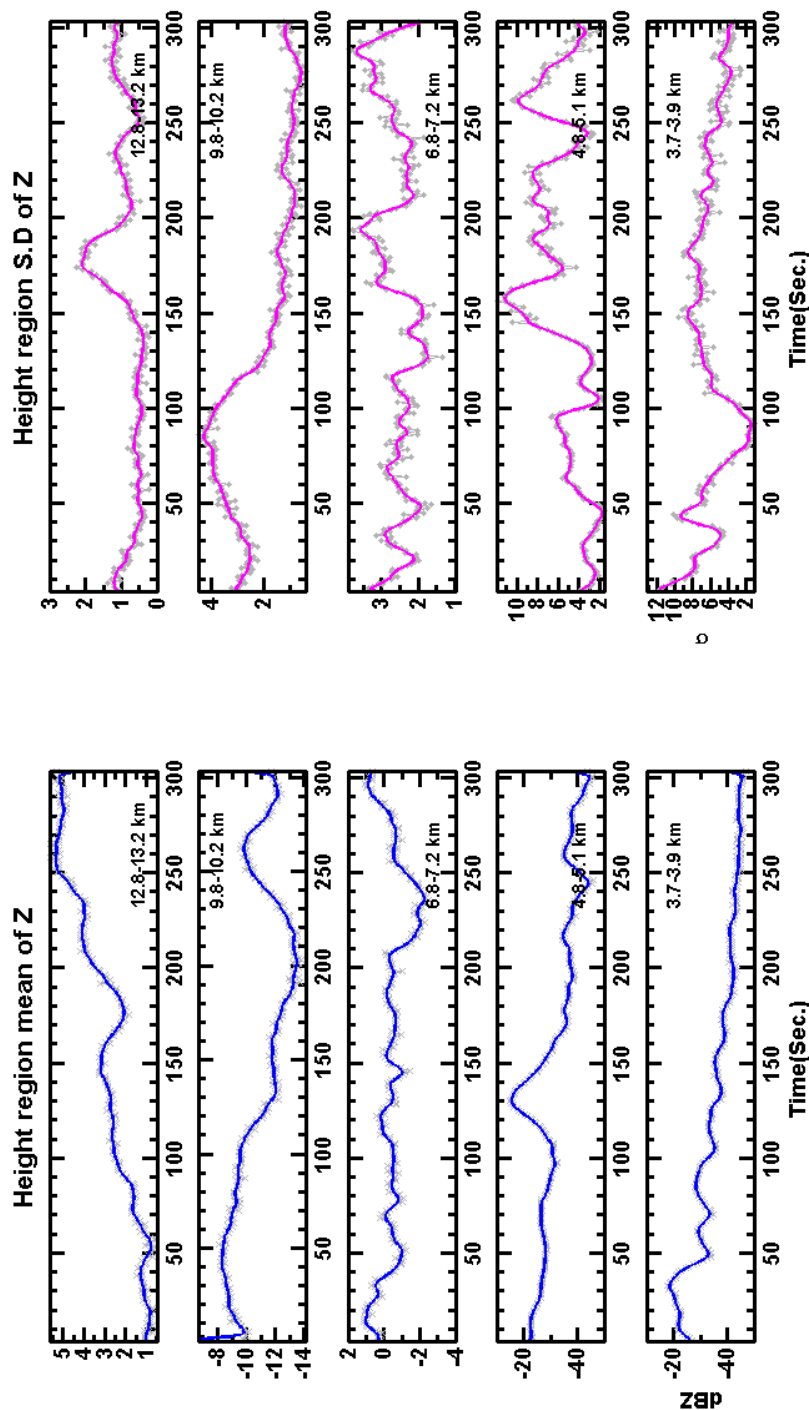


Figure 8b: Same as Figure 8a but for the cloud regions as per Table 2.

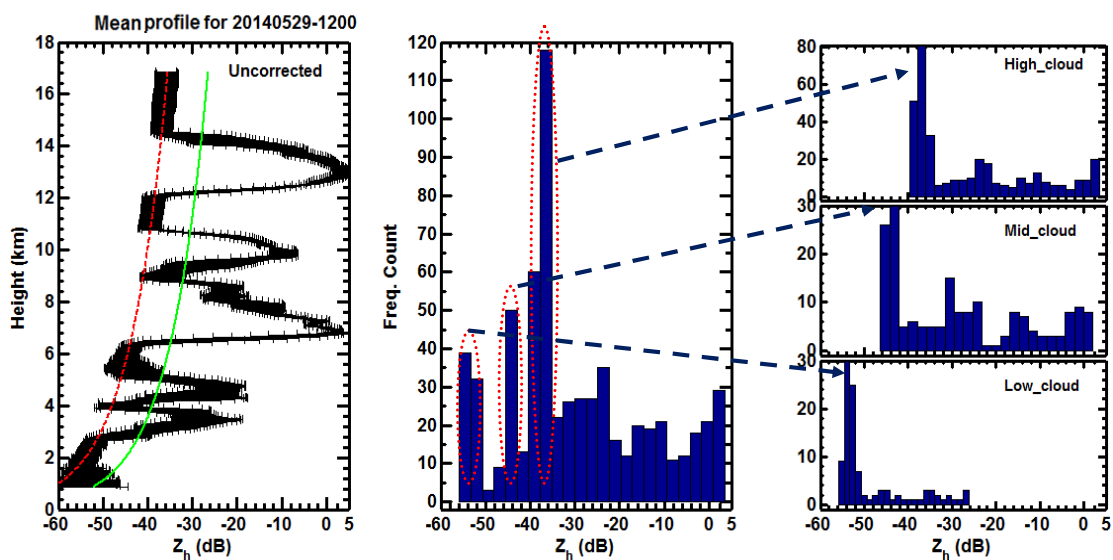


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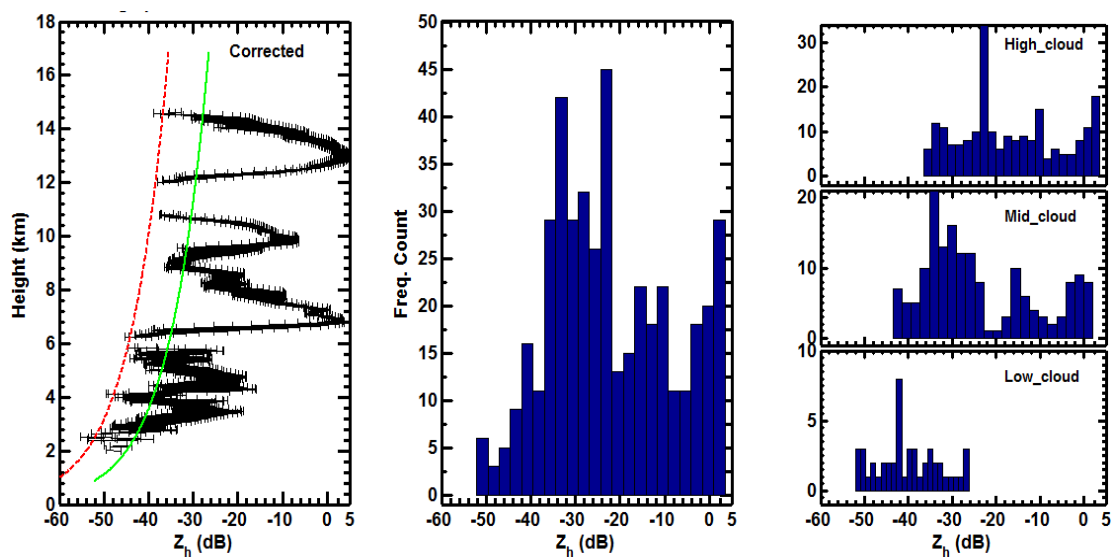


Figure 9b: Same as 9a but it is corrected by filtering out noise and biota. The correction applied to Z profile allows to pop-up the true meteorological cloud reflectivity distribution.

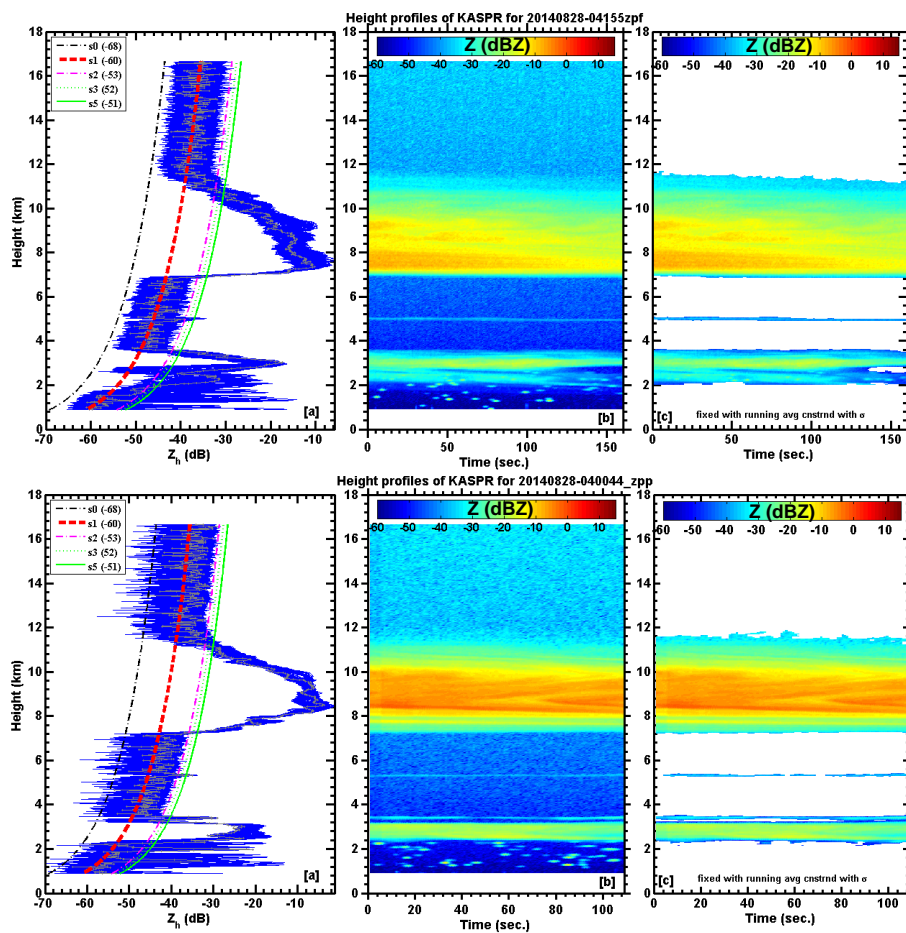


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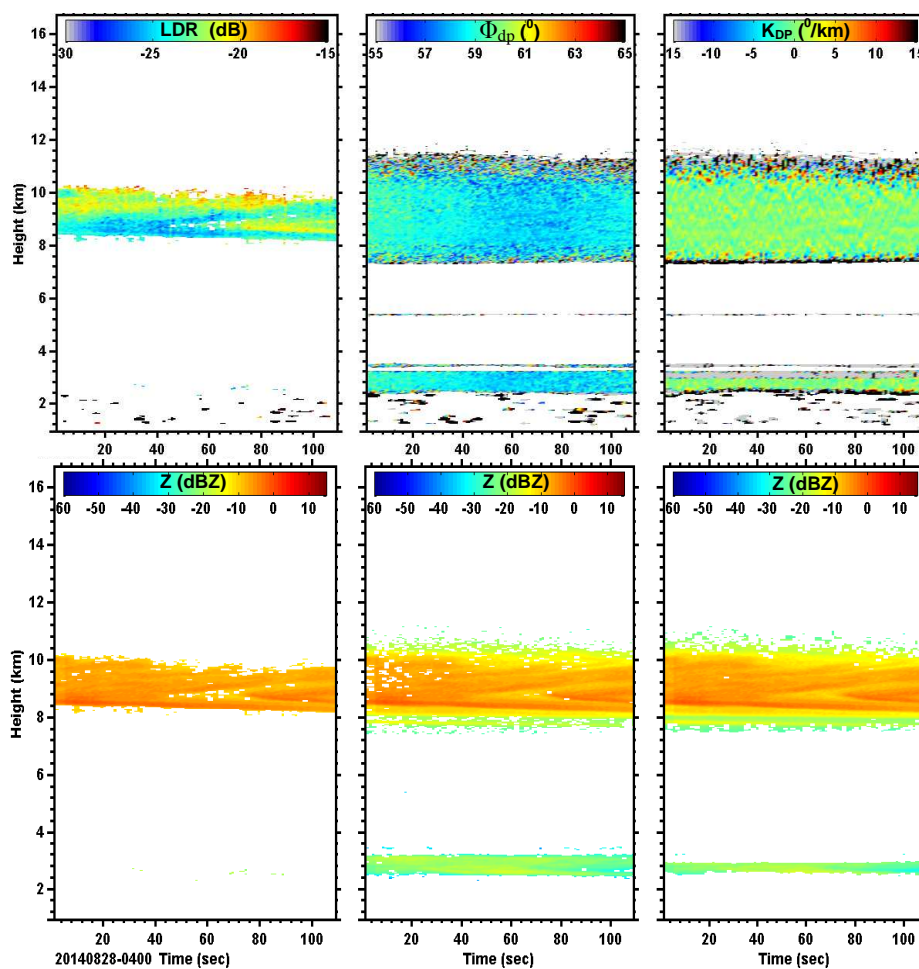


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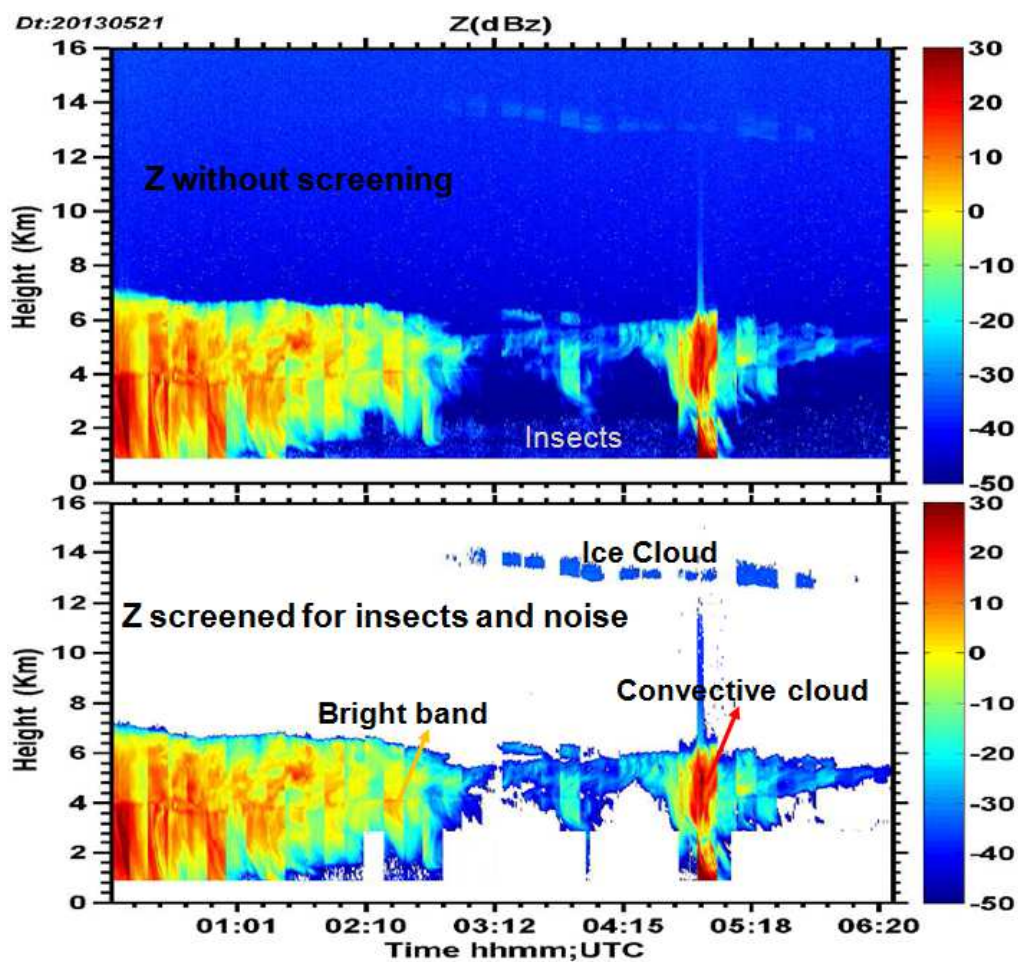


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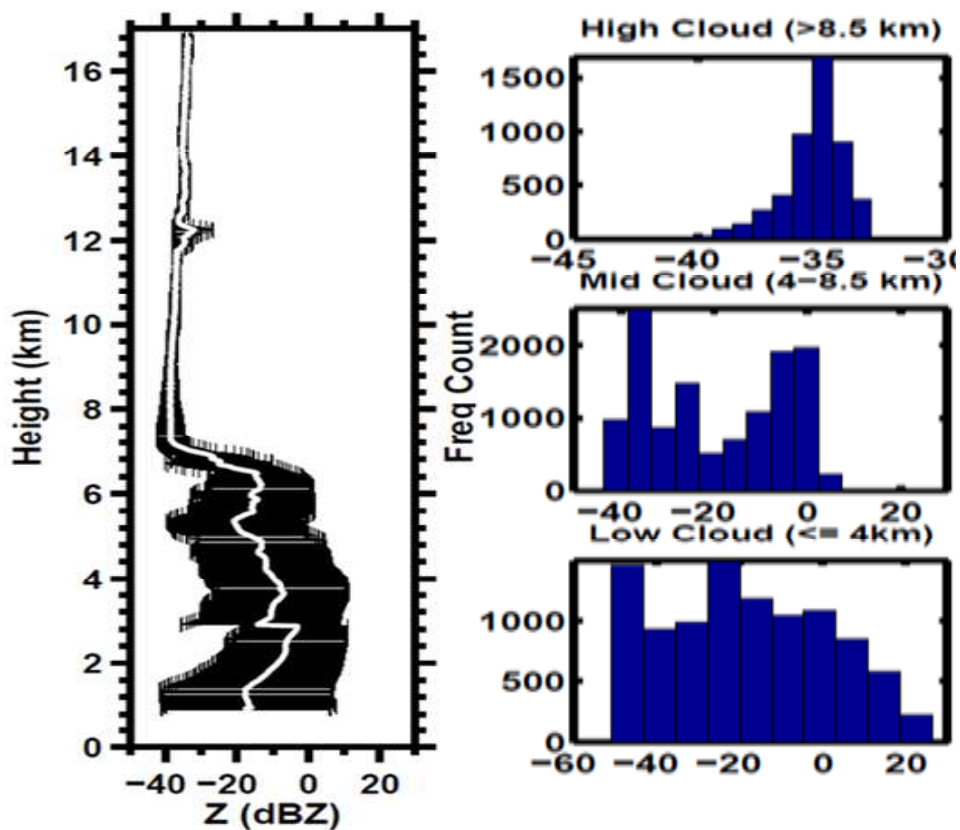


Figure 12b: Screened-out cloud radar reflectivity mean and standard deviation profile with the tri-model cloud reflectivity frequency distribution.

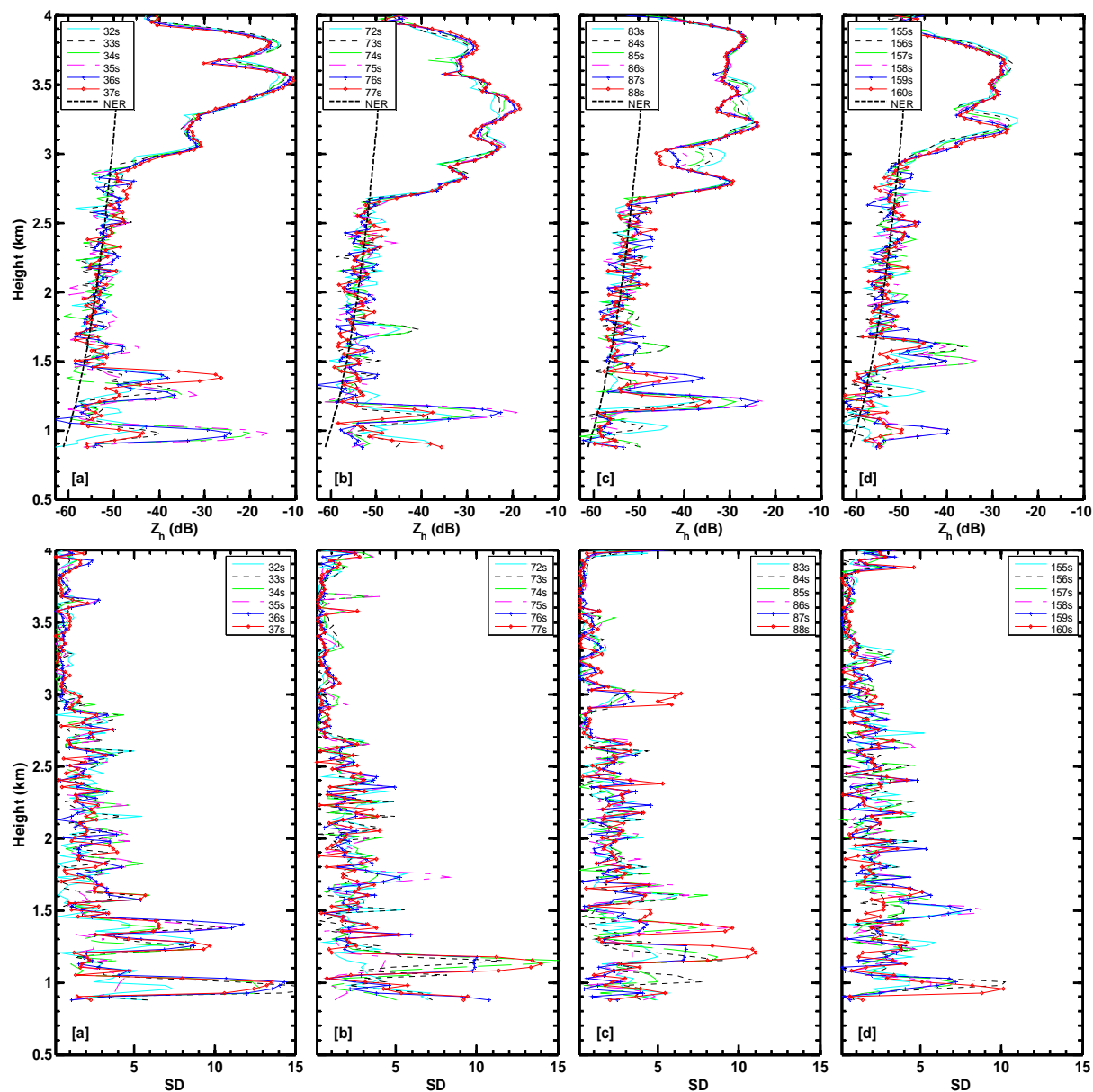


Figure A1: Instantaneous height profiles of Z during 1200-1205 UT on 29 May 2014 with centered numer profile notice to be the strong insect return identified with HTI plot of figure 4b. Bottom panels correspond to standard deviation (SD) from four point running average.

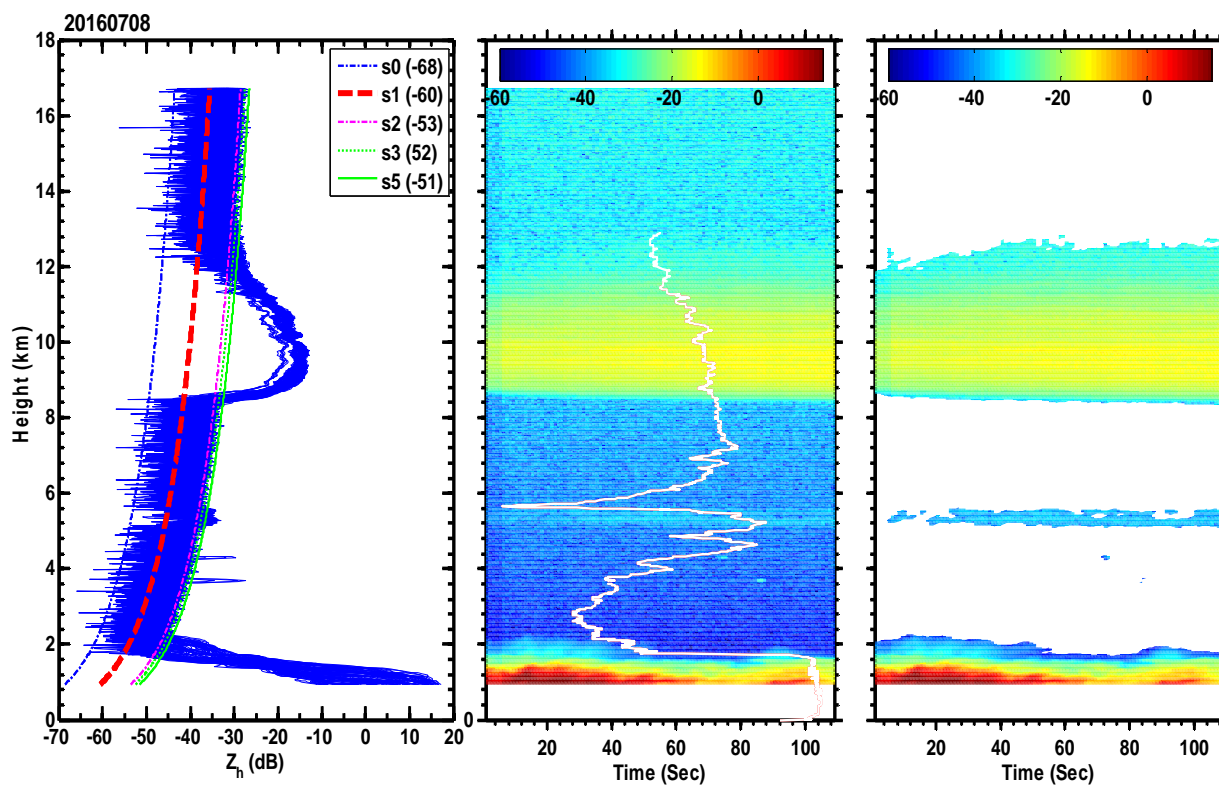


Figure A2: (Right-middle-left) Same as 1(a-c) but on 08 Jul 2016 during 0531 UT for the duration of 108 sec. S0-S5 are NER curves. Collocated GPS-RS relative humidity (%) profile had shown as white solid line in the middle panel.