

# Comparison of three aerosol chemical characterization techniques utilizing PTR-ToF-MS: A study on freshly formed and aged biogenic SOA

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## Abstract

An inter-comparison of different aerosol chemical characterization techniques has been performed as part of a chamber study of biogenic SOA formation and aging at the atmosphere simulation chamber SAPHIR. Three different aerosol sampling techniques, the aerosol collection module (ACM), the chemical analysis of aerosol on-line (CHARON) and the collection thermal desorption unit (TD) were connected to Proton Transfer Reaction Time of Flight Mass Spectrometers (PTR-ToF-MS) to provide chemical characterization of the SOA. The techniques were compared among each other and to results from an Aerosol Mass Spectrometer (AMS) and a Scanning Mobility Particle Sizer (SMPS). The experiments investigated SOA formation from the ozonolysis of  $\beta$ -pinene, limonene, a  $\beta$ -pinene/limonene mix and real plant emissions from *Pinus sylvestris* L. (Scots pine). The SOA was subsequently aged by photooxidation except for limonene SOA which was aged by NO<sub>3</sub> oxidation.

Despite significant differences in the aerosol collection and desorption methods of the PTR based techniques, the determined chemical composition, i.e. the same major contributing signals were found by all instruments for the different chemical systems studied. These signals could be attributed to known products expected from the oxidation of the examined monoterpenes. The sampling and desorption method of ACM and TD, provided additional information on the volatility of individual compounds and showed relatively good agreement.

Averaged over all experiments, the total aerosol mass recovery compared to an SMPS varied from  $80 \pm 10\%$ ,  $51 \pm 5\%$  and  $27 \pm 3\%$  for CHARON, ACM and TD, respectively. Comparison to the oxygen to carbon ratios (O:C) obtained by AMS showed that all PTR based techniques observed lower O:C ratios indicating a loss of molecular oxygen either during aerosol sampling or detection. The differences in total mass recovery and O:C between the three instruments resulted predominantly from differences in the field strength (E/N) in the drift-tube reaction ionization chambers of the PTR-ToF-MS instruments and from dissimilarities in the collection/desorption of

38 aerosols. Laboratory case studies showed that PTR-ToF-MS E/N conditions influenced fragmentation which  
40 resulted in water and further neutral fragment losses of the detected molecules. Since ACM and TD were operated in  
42 higher E/N compared to CHARON this resulted to higher fragmentation, thus affecting primarily the detected  
oxygen and carbon content and therefore also the mass recovery. Overall, these techniques have been shown to  
provide valuable insight on the chemical characteristics of BSOA, and can address unknown thermodynamic  
properties such as partitioning coefficient values and volatility patterns down to a compound specific level.

44

## 1 Introduction

46 Atmospheric organic aerosols (OA) represent a major contribution to the submicrometer particulate matter (PM<sub>1</sub>)  
thus playing a key role in climate change and air quality (Kanakidou et al., 2005). OA are either directly emitted  
48 through e.g. combustion processes (primary OA, POA) or formed through the oxidation of volatile organic  
compounds (VOCs), called secondary OA (SOA) (Seinfeld and Pandis, 2006). SOA constitute a major fraction of  
50 OA (Jimenez et al., 2009) with biogenic VOC oxidation products affecting their global contribution (Guenther et al.,  
2012). Due to thousands of individual compounds involved in SOA, the chemical characterization of OA still  
52 presents a huge analytical challenge (Goldstein and Galbally, 2007). The ability of these compounds to condense to  
the particulate-phase or partition between the gas and particle phase as well as their volatility are thermodynamic  
54 parameters of interest that determine their atmospheric fate.

Various techniques have been established in order to better quantify and chemically characterize SOA (Hallquist et  
56 al., 2009). These techniques optimize and compromise for time, size or chemical resolution combined with the  
percentage of OA mass they can detect. Off-line techniques, based on filter measurement, provide detailed  
58 information on functional groups or individual chemical species while having low time resolution (hours to days)  
and size information. On-line techniques, like e.g. the Aerodyne aerosol mass spectrometer (AMS) (Canagaratna et  
60 al., 2007), provide high time resolution and size resolved data while less specific chemical composition information  
or molecular identification of the OA compounds is acquired.

62 In recent years attempts to develop new techniques that combine both chemical identification but also improved  
time resolution have been established. These techniques use different pre-concentration methods in order to detect  
64 the particulate-phase compounds. Filter based techniques like the Filter Inlet for Gases and AEOROsols  
(FIGAERO) (Lopez-Hilfiker et al., 2014) provide highly effective collection of particles on filters, under high flow  
66 rates (30 standard Liters per minute, sLpm), thus low collection times. Thermal desorption of the sampled particles  
on the filter is performed with the disadvantage of sampling artefacts from gas-phase compounds that may condense  
68 on the large surface area of the filter and contribute to the overall signal. Other techniques, like the thermal  
desorption aerosol gas chromatograph (TAG) (Kreisberg et al., 2009; Williams et al., 2006) or the collection thermal  
70 desorption unit (TD) (Holzinger et al., 2010b), utilize the concept of particle collection on an impaction surface by  
means of humidification and inertial impaction, followed by desorption. TAG and TD provide hourly time  
72 resolution measurements, and when combined with a gas-phase denuder avoid sampling of additional gas-phase  
constituents on their collection thermal desorption (CTD) cell. Due to the particle humidification step these  
74 techniques may bias collection efficiency towards water soluble compounds. The aerosol collection module (ACM)

(Hohaus et al., 2010) collects aerosols by passing them through an aerodynamic lens for particle collimation (Liu et al., 1995a; b), further through a vacuum system (comparable in design to the AMS), and finally impacting the particle phase on a cooled sampling surface. Although the ACM has a low time resolution (3-4 h), its design makes it applicable for the investigation of compound specific thermodynamic properties e.g. partitioning coefficient and volatility (Hohaus et al., 2015). The chemical analysis of aerosol online (CHARON) (Eichler et al., 2015) is a technique that provides on-line real time measurements by passing the particles through a denuder to strip off the gas-phase. Particles are sampled through an aerodynamic lens combined with an inertial sampler for the particle-enriched flow, and a thermodesorption unit for particle volatilization prior to chemical analysis. The enrichment factor of this system is known by performing calibrations, thus reducing the quantification uncertainty. All the above pre-concentration systems detect the compounds originating from the particulate-phase that underwent evaporation to the gas-phase by desorption, thus introducing possible thermal break down of analytes.

A variety of detection instruments have been coupled to these inlet techniques, providing different functionality and chemical composition information. The proton transfer reaction time of flight mass spectrometer (PTR-ToF-MS) (Jordan et al., 2009) is a soft ionization technique with low detection limits and high time resolution (ms), that can cover a wide volatility range, from VOCs to low-volatility VOCs (LVOCs), depending on the inlet used (Eichler et al., 2017). Techniques utilizing a PTR-ToF-MS are capable of measuring a large fraction of the OA mass, ranging from 20 to 100% (Eichler et al., 2015; Mensah et al., 2012), and provide additional information on the elemental composition of the organic compounds; however, the compound's molecular identity attribution is challenging. On the contrary, gas chromatography mass spectrometry is considered ideal for detailed compound specific structural analysis. Techniques like the TAG have been applied utilizing a gas chromatograph, to provide non-polar and low-polarity tracers identification while the modified semi-volatile TAG (SV-TAG) has broadened this range to highly polar oxygenates, mostly seen in the atmosphere, by using online derivatization (Isaacman et al., 2014; Zhao et al., 2013). The volatility and polarity separator (VAPS) is a similar technique that provides volatility- and polarity-resolved OA information by using a modified 2-dimensional gas chromatography (2D-GC) approach combined with high resolution time-of-flight mass spectrometry (Martinez et al., 2016). Although these techniques provide chemical speciation and lower time resolution, they can only do so for a small fraction of the OA mass (10 - 40%).

The specificity of the above newly developed techniques is still to be explored in detail. In this work, an inter-comparison campaign was performed in the atmosphere simulation chamber SAPHIR (Rohrer et al., 2005) to investigate biogenic SOA (BSOA) formation and aging. The focus of this work is on the comparison of three different aerosol characterization techniques, the ACM - PTR-ToF-MS, the TD - PTR-ToF-MS and the CHARON - PTR-ToF-MS. The OA mass fraction these techniques were able to detect combined with the OA chemical characteristics and volatility trends were investigated and compared.

## 2 Methods and instrumentation

### 2.1 Facilities

Experiments were conducted in the atmospheric simulation chamber SAPHIR (Simulation of Atmospheric PHotochemistry In a large Reaction chamber) located in Jülich, Germany. The chamber consisted of twin FEP

112 Teflon foils with a volume of 270 m<sup>3</sup>, resulting in a surface to volume ratio of approximately 1 m<sup>-1</sup>. High purity  
nitrogen (99.9999% purity) was flushed at all times to the space between the twin walls and a pressure gradient was  
114 maintained in order to prevent contamination from outside. A high flow (150 to 200 m<sup>3</sup> h<sup>-1</sup>) of air was introduced in  
order to clean the chamber and reach aerosol and trace gases concentrations below detection limits before each  
116 experiment was initiated. A low flow (8 m<sup>3</sup> h<sup>-1</sup>) was used to replenish SAPHIR during experiments from losses due  
to leaks and sampling of the instruments. The chamber is equipped with a louvre system thus experiments could be  
performed under dark conditions focusing on O<sub>3</sub> and NO<sub>3</sub> oxidation (roof closed) or as photooxidation experiments  
118 utilizing sun light (roof open). More details on SAPHIR can be found in Rohrer et al. (2005).

A PLant chamber Unit for Simulation (PLUS) was recently coupled to SAPHIR to investigate the impact of real  
120 plant emissions on atmospheric chemistry (Hohaus et al., 2016). PLUS is an environmentally controlled, flow  
through plant chamber where continuous measurements and adjustments of important environmental parameters  
122 (e.g., soil relative humidity, temperature, photosynthetic active radiation) are performed. To simulate solar  
radiation and control the tree emissions in PLUS, 15 light-emitting diode (LED) panels were used with an average  
124 photosynthetically active radiation value (PAR) of 750 μmol photons m<sup>-2</sup> s<sup>-1</sup> and an average temperature of 25 °C . In this study,  
BVOC emissions were generated from 6 *Pinus sylvestris* L. (Scots pine) trees.

126 A set of standard instrumentation was coupled to the simulation chamber SAPHIR. Air temperature was measured  
by an ultrasonic anemometer (Metek USA-1, accuracy 0.3 K) and humidity was determined with a frost point  
128 hygrometer (General Eastern model Hygro M4). NO and NO<sub>2</sub> measurements were performed with a  
chemiluminescence analyser (ECO PHYSICS TR480) equipped with a photolytic converter (ECO PHYSICS  
130 PLC760). Ozone was measured by an UV absorption spectrometer (ANSYCO model O341M). Particle size  
distribution was measured using a Scanning Mobility Particle Analyser (SMPS TSI, TSI Classifier model 3080, TSI  
132 DMA 3081, TSI Water CPC 3786), measuring in the 10 - 450 nm range with a time resolution of 8.5 min and an  
accuracy of 12% (Wiedensohler et al., 2012). A High-Resolution Time-of-Flight Aerosol Mass Spectrometer (HR-  
134 ToF-AMS) (Canagaratna et al., 2007; DeCarlo et al., 2006) was used to determine the total organic mass and  
composition of the SOA formed with an accuracy of 31% (Aiken et al., 2008). High resolution mass spectra were  
136 analyzed using the software packages SQUIRREL (v1.57) and PIKA (v1.15Z). Oxygen to carbon ratios were  
calculated based on the newly developed “Improved-Ambient” method by Canagaratna et al. (2015).

## 138 2.2 Experimental procedure

SOA was formed through the ozonolysis of different monoterpenes using the simulation chamber SAPHIR.  
140 Experimental starting conditions varied from the injection of β-pinene and limonene, as single compounds or as a  
mixture, to the injection of real plant emissions from 6 *Pinus sylvestris* L. (Scots pine), provided from SAPHIR-  
142 PLUS (Section 2.1). For the tree emissions experiment the BVOCs consisted of 42% δ<sup>3</sup>-carene, 38% α-pinene, 5%  
β-pinene, 4% myrcene, 3% terpinolene and 8% other monoterpenes, as determined by GC-MS measurements. The  
144 details of the experiments are given in Table 1. The chamber was initially humidified (55% RH, 295 – 310 K) and  
background measurements for all instruments were performed. Monoterpenes were injected either with a Hamilton  
146 syringe injection and subsequent evaporation into the replenishment flow of SAPHIR, or by SAPHIR-PLUS (real

tree emissions). After one hour, ozone was introduced in the system to initiate chemistry. The ozonolysis of  
148 monoterpene and the tree emissions was performed under low NO<sub>x</sub> conditions (10 – 60 pptV) in the absence of an  
OH scavenger. For the limonene experiment, 8 hours after the ozone injection, an addition of 30 ppbV of NO was  
150 introduced into the dark chamber. The reaction of NO<sub>2</sub> with remaining ozone in the dark chamber resulted in the  
generation of NO<sub>3</sub>, thus initiating the NO<sub>3</sub> oxidation chemistry. In all other experiments the chamber was  
152 illuminated 20 hours after the ozone injection, exposing the SOA to real sunlight, thus initiating photo-oxidation by  
OH radicals. Finally, for the real tree emissions, after 11 hours of ozone exposure, additional biogenic VOCs  
154 (BVOCs) were re-introduced into the SAPHIR chamber to generate fresh SOA which was subsequently aged by  
photooxidation for additional 6 hours. The duration of the experiments varied from 17 to 36 hours, providing ample  
156 time to experimentally investigate the aging of the biogenic SOA.

### 158 2.3 PTR-ToF-MS aerosol chemical characterization techniques

Three independent aerosol chemical characterization techniques utilizing PTR-ToF-MS were compared, the aerosol  
160 collection module (ACM – PTR-ToF-MS, referred to as “ACM” hereafter), the chemical analysis of aerosol online  
(CHARON – PTR-ToF-MS, referred to as “CHARON” hereafter) and the collection thermal desorption unit (TD –  
162 PTR-ToF-MS, referred to as “TD” hereafter). Their characteristics and differences are provided in Table 2 and  
discussed in detail in this section. The time resolution of the techniques varied from CHARON providing online  
164 measurements to the TD and ACM having increased collection times of 30 and 240 min, respectively. CHARON  
was operated at a constant temperature and lower pressure (< 1 atm) while ACM and TD, operated at 1 atm,  
166 introduced temperature ramps during desorption thus providing more detailed volatility information. The pre-  
concentration factor for ACM and TD was calculated from the ratio of the volume sampled during collection to the  
168 volume evaporated during desorption, assuming a 3 min desorption time for an individual compound. The limit of  
detection (LOD), dependent on the different pre-concentration factors for each technique, resulted in TD having the  
170 lowest LOD (0.02 ng m<sup>-3</sup>), followed by the CHARON (1.4 ng m<sup>-3</sup>), while ACM showed the highest  
values (35 ng m<sup>-3</sup>). It should be noted that for the individual PTRMS the LOD for gas-phase measurements,  
172 bypassing any pre-concentration step, agreed within a factor of two. Different electric field strength (V cm<sup>-1</sup>) to  
buffer gas density (molecules cm<sup>-3</sup>) ratio (E/N) conditions were applied to the PTR-ToF-MS of each aerosol  
174 chemical characterization technique. Lower E/N set values resulted in longer ion residence times in the drift tube of  
the PTR-ToF-MS thus higher sensitivity due to enhanced proton transfer reaction times. Ions were introduced to a  
176 lower kinetic energy system, thus resulting in reduced fragmentation during ionization while the cluster ion  
distribution was changed when lowering the E/N, supporting more H<sub>3</sub>O<sup>+</sup>(H<sub>2</sub>O)<sub>n</sub> (n=1,2,3..) cluster ion generation (de  
178 Gouw and Warneke, 2007). Since the proton affinity of H<sub>3</sub>O<sup>+</sup>(H<sub>2</sub>O)<sub>n</sub> is higher than that of H<sub>3</sub>O<sup>+</sup>, a certain range of  
organic compounds could not be ionized in such operating conditions. An overview of the primary ion distribution is  
180 provided in Figure S 1. Normalization of the signal was performed based on the sum of  
500 \* H<sub>3</sub>O<sup>+</sup> + 250 \* H<sub>3</sub>O<sup>+</sup>(H<sub>2</sub>O) for all PTRMS. ACM and TD showed more than 98 % of the primary ions  
182 originating from H<sub>3</sub>O<sup>+</sup> while for CHARON, when operated at 100 Td (1 Td = 10<sup>-17</sup> V cm<sup>-2</sup> molecule<sup>-1</sup>), around 65%

184 originated from  $\text{H}_3\text{O}^+$  and 35% from  $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ , and for CHARON at 65 Td, around 20% from  $\text{H}_3\text{O}^+$  and 75%  
186 from  $\text{H}_3\text{O}^+(\text{H}_2\text{O})$ . Based on the uncertainty in the reaction rate coefficient of the organic compounds with  $\text{H}_3\text{O}^+$  the  
PTR-ToF-MS was assumed to introduce a  $\pm 40\%$  uncertainty on the volume mixing ratios of uncalibrated  
188 compounds for CHARON and TD. The ACM used an average sensitivity of 15 ncps/ppbV with an uncertainty of  
 $\pm 50\%$  ( $\pm 1\sigma$ ) where ncps accounted for the normalized to the primary ions signal.  
190 All PTR-ToF-MS used in this campaign were of the model PTR-TOF 8000, manufactured from Ionicon Analytik  
GmbH, Innsbruck, Austria. Although being the same models, minor differences in the design e.g. the TOF interface  
192 existed, related mostly to ACM when compared to CHARON and TD. These differences introduced additional  
fragmentation and affected the resolution of the PTRMS as reflected from Table 2. Nevertheless, the sensitivity of  
194 all PTRMS when using acetone as a calibration compound was in a similar range as observed in Figure S1. When  
calculating the sensitivity using the cps instead of the ncps, observed differences suggested lower primary ion signal  
and reaction times for ACM and TD when compared to CHARON. In the following subsections the principle of  
operation and operating conditions of the different inlets and PTRMS systems used in this study is reported.

### 196 2.3.1 ACM – PTR-ToF-MS

The ACM is an aerosol collection inlet with subsequent sample evaporation coupled to a gas-phase detector  
198 designed for in situ, compound specific chemical analysis. The ACM can be adapted to work with different gas-  
phase analysers and has previously been used coupled to a GC-MS (Hohaus et al., 2010). In this work, the ACM  
200 was coupled to a PTR-ToF-MS (model PTR-TOF 8000; Ionicon Analytik GmbH, Innsbruck, Austria).

In brief, ambient air was sampled through an aerodynamic lens (Liu et al., 1995a; b) with a flow rate of  $80 \text{ ml min}^{-1}$ .  
202 Within the aerodynamic lens the gas and particle phase of an aerosol were separated and the particles were  
collimated into a narrow beam. The particle beam was directed through a high vacuum environment ( $10^{-5}$  torr) to a  
204 cooled ( $-5^\circ\text{C}$ ) sampling surface made of Siltek®/Sulfurnert®-treated stainless steel. After collection was completed  
(a collection time of 4 h was used in this study) the particles were thermally desorbed by heating up the collector.  
206 The evaporated compounds were transferred to the PTR-ToF-MS through a coated stainless steel line of 0.8 mm  
inner diameter and 30 cm length, constantly kept at  $300^\circ\text{C}$ . Nitrogen was used as carrier gas with a flow of  
208  $300 \text{ ml min}^{-1}$ , resulting in a residence time of 60 ms. For this study, the collector temperature was ramped by  
 $100^\circ\text{C min}^{-1}$  to a maximum of  $250^\circ\text{C}$ , with 3-minute isothermal sections at  $100^\circ\text{C}$ ,  $150^\circ\text{C}$  and  $250^\circ\text{C}$ ,  
210 respectively. During the final temperature step of  $250^\circ\text{C}$ , desorption time was extended for additional 7 minutes to  
ensure complete evaporation of the sample. These temperature steps provided enough time for compounds to  
212 undergo evaporation within a defined volatility range. The signal dropped close to zero before each temperature step  
was completed, making the ACM-PTR-ToF-MS ideal for compound specific volatility trend analysis. Parallel to the  
214 ACM particulate-phase collection, a bypass line was used, coupled to the same PTR-ToF-MS, measuring the gas-  
phase during particle phase sampling time. An example of the gas and particulate-phase measurements is given in  
216 Figure S2. During the campaign, the aerosol-phase sampling line was a stainless steel line  
(total length: 4 m, OD: 1/4") with a flow of  $0.7 \text{ L min}^{-1}$ .

218 Assuming a collection efficiency of 100% (Hohaus et al., 2010) for all particles in the aerosol sample, measured  
 PTR-ToF-MS signals could be converted to particulate mass concentrations by applying PTR calibrations as  
 220 described in the following. Normalization of the PTR-ToF-MS counts per second was performed based on the  $\text{H}_3\text{O}^+$   
 signal, resulting in ncps. The ACM was corrected for mass discrimination. The mass discrimination function was  
 222 determined based on the ratio of the measured over the theoretical sensitivity of acetaldehyde, acetone, butanone,  
 benzene, toluene, xylene and mesitylene. The instrument was calibrated for a total of 15 compounds including  
 224 aromatics (benzene, toluene, xylene, chlorobenzene), oxygenates (acetaldehyde, acetone, 2-butanone, 3-pentanone,  
 MVK, nopinone, methanol, 1-butanol), pure hydrocarbons (isoprene,  $\alpha$ -pinene) and acetonitrile. Calibration was  
 226 performed by coupling the PTR-ToF-MS to a calibration unit (LCU, Ionicon Analytik GmbH, Innsbruck, Austria)  
 and measuring known concentration of the compounds in the gas-phase. For signals observed at uncalibrated masses  
 228 the average sensitivity of acetaldehyde, acetone, MVK, Butanon, pentanone and nopinone was applied resulting in  
 15 ncps/ppb. The mass concentration of an aerosol compound  $z_i$  in the air sample was calculated based on the  
 230 mixing ratios the PTR-MS measures:

$$\mathbf{mz}_{i,(\mu\text{g}/\text{m}^3)} = \frac{\mathbf{mz}_{i,(\text{ppb})} \times \mathbf{MW}_i}{\mathbf{T} \times \mathbf{R}} \times \frac{\mathbf{F}_{\text{N}_2} \times \mathbf{t}_{\text{meas}}}{\mathbf{F}_{\text{col}} \times \mathbf{t}_{\text{col}}}, \quad (1)$$

232 where  $\mathbf{mz}_{i,(\mu\text{g}/\text{m}^3)}$  is the aerosol concentration of compound  $i$  in  $\mu\text{g m}^{-3}$ ,  $\mathbf{mz}_{i,(\text{ppb})}$  the background corrected  
 arithmetic mean of the mixing ratio during the aerosol analysis in the nitrogen flow in ppb,  $\mathbf{MW}_i$  is the molecular  
 234 weight of compound  $i$  in  $\text{gmol}^{-1}$ ,  $\mathbf{R}$  is the universal gas law constant,  $\mathbf{T}$  the ambient temperature of the SAPHIR  
 chamber in Kelvin,  $\mathbf{F}_{\text{N}_2}$  the flow of the carrier gas in standard liter per minute,  $\mathbf{t}_{\text{meas}}$  the aerosol desorption duration,  
 236  $\mathbf{F}_{\text{col}}$  the collection flow rate of the aerosol to the ACM in standard liter per minute and  $\mathbf{t}_{\text{col}}$  the aerosol collection  
 duration. The volume ratio correction  $\left(\frac{\mathbf{F}_{\text{N}_2} \times \mathbf{t}_{\text{meas}}}{\mathbf{F}_{\text{col}} \times \mathbf{t}_{\text{col}}}\right)$  was applied in order to account for the ACM collection  
 238 preconcentration step. The mass concentration was calculated by taking into account only the signal above the  
 instrument noise ( $> 2\sigma$ ) for each compound at each desorption.

240 Background measurements were performed before and after every experiment ( $\sim 2$  times per day) by heating up the  
 collector, without depositing particles on the surface beforehand. The signal derived from the background  
 242 measurements at each temperature step was then interpolated and subtracted from all desorptions for all compounds.  
 Two major factors could affect the background signal, gas-phase interference and aerosol residual remaining at the  
 244 collector after each desorption cycle. Due to the aerodynamic lens set-up the ACM design prevents gas-phase  
 contamination (removal  $> 99.9999\%$ ). Background measurements throughout this study show no residual  
 246 compounds on the collector in the desorption temperature range studied.

PTR-ToF-MS operation conditions were kept constant throughout the campaign. It was operated at  $E/N = 120$  Td.  
 248 The drift tube was kept at a temperature of  $100^\circ\text{C}$  and a pressure of  $2.30$  mbar. The mass resolving power of this  
 PTR-ToF-MS was  $m/\Delta m \sim 2500$  ( $\Delta m$  is full width at half maximum). Mass spectra were collected up to  $m/z$  400 at  
 250  $10$  s signal integration time. Analysis of the raw data was performed using the PTR-TOF Data Analyzer (version  
 4.40) software (Müller et al., 2013). In brief, an integration time of  $90$  s was chosen for the software and  $m/z$   
 252 calibration peaks were assigned based on the peaks of  $21.02$ ,  $59.05$  and  $180.94$  accounting for  $\text{H}_3[18\text{O}]^+$ , protonated

254 acetone and trichlorobenzene respectively. Trichlorobenzene was used as an internal standard throughout the  
campaign. The chemical composition assignment was derived from the measured exact mass assuming a molecular  
formula of  $C_xH_yO_zN_a$  and attributing the isotopic pattern when possible.

### 256 2.3.2 CHARON – PTR-ToF-MS

258 The analyzer deployed by the University of Innsbruck consisted of a Chemical Analysis of Aerosol Online  
(CHARON) inlet interfaced to a PTR-ToF-MS.

260 The CHARON inlet (Eichler et al., 2015) consists of a gas-phase denuder for stripping off gas-phase analytes, an  
aerodynamic lens for particle collimation combined with an inertial sampler for the particle-enriched flow, and a  
thermodesorption unit for particle volatilization prior to chemical analysis. The monolithic charcoal denuder (Mast  
262 Carbon International Ltd., Guilford, UK) used in this study was 25 cm long, had an outer diameter of 3 cm and a  
channel density of 585 channels per inch (cpi). The thermodesorption unit consisted of a heated Siltek®/Sulfinert®-  
264 treated stainless steel tube kept at a temperature of 140 °C and a pressure on the order of a few mbar. A HEPA filter  
(ETA filter model HC01-5N-B, Aerocolloid LLC, Minneapolis, MN, USA) was periodically placed upstream of the  
266 gas-phase denuder for determining the instrumental background. More details on the performance of the CHARON  
inlet are given in Eichler et al. (2015).

268 The CHARON inlet was interfaced to a commercial PTR-ToF-MS instrument (model PTR-TOF 8000; Ionicon  
Analytik GmbH, Innsbruck, Austria). PTR-ToF-MS mass spectra were collected up to  $m/z$  500 at 10 s signal  
270 integration time. The PTR-TOF Data Analyzer (version 4.40) software was used for data analysis (Müller et al.,  
2013). During the tree emissions experiment the electric field applied to the drift tube was periodically switched in  
272 300 s intervals, i.e. measurements were performed at alternating E/N-values of 65 Td (referred to as “CHARON<sub>65</sub>”  
hereafter) and 100 Td (referred to as “CHARON<sub>100</sub>” hereafter), respectively ( $1 \text{ Td} = 10^{-17} \text{ V cm}^{-2} \text{ molecule}^{-1}$ ). For all  
274 other experiments the E/N-value analysed was at 100 Td. The drift tube was kept at a temperature of 120 °C and a  
pressure of 2.40 mbar. Continuous permeation of 1,2-diiodobenzene was performed into the drift tube for generating  
276 mass axis calibration signals at  $m/z$  203.943 and  $m/z$  330.847. The PTR-ToF-MS was characterized using a 16-  
compound gas mixture that included aromatics (benzene, toluene, o-xylene, mesitylene, chlorobenzene), oxygenate  
278 compounds (acetaldehyde, acetone, 2-butanone, 3-pentanone, MVK, nopinone, methanol, 1-butanol), pure  
hydrocarbons (isoprene,  $\alpha$ -pinene) and acetonitrile. The mass resolving power of this PTR-ToF-MS was  $m/\Delta m$   
280 4500-5000.

The entire CHARON setup was calibrated using size-selected ammonium nitrate particles as described in Eichler et  
282 al. (2015). A sensitivity model based on Su and Chesnavich’s parameterized reaction rate theory and a chemical  
composition based parameterization of polarizabilities at a constant dipole moment of  $\mu_D = 2.75 \text{ D}$  (between 1 – 4.5  
284 D for most oxygenated organic compounds) was applied to calculate sensitivities of unknown compounds. This  
resulted in an  $m/z$  independent sensitivity accuracy of about  $\pm 25\%$ . For compounds without assigned elemental  
286 composition the polarizability of acetone was applied with an accuracy of  $\pm 40\%$ . Derived volume mixing ratios  
were transformed to mass concentrations using the molecular  $m/z$  information at Normal Temperature and Pressure  
288 (NTP) conditions (293.15 K, 101.325 kPa). Quantification was hampered by two events (power failure, partial



obstruction of the aerodynamic lens) which resulted in a higher than usual variability of the particle enrichment in the aerodynamic lens. Results from two experiments (limonene ozonolysis/NO<sub>3</sub> oxidation and limonene/β-pinene mixture ozonolysis) were particularly affected as will be shown and discussed in section 3.

The CHARON-PTR-ToF-MS setup was interfaced to the SAPHIR chamber using Siltek®/Sulfinert®-treated stainless steel tubing (total length: 600 cm, 50 cm extending into the chamber, ID: 5.33 mm). During the β-pinene ozonolysis and limonene ozonolysis/NO<sub>3</sub> oxidation experiments, the inlet flow was kept at 0.6 l min<sup>-1</sup> resulting in a sample residence time of 13.4 s. During the β-pinene/limonene mixture ozonolysis and the real tree emissions ozonolysis experiments, the inlet flow was increased to 1.6 l min<sup>-1</sup> resulting in a sample residence time of 5.0 s.

### 2.3.3 TD – PTR-ToF-MS

The Thermal-Desorption unit was coupled to a commercial PTR-TOF8000 instrument (Ionicon Analytik GmbH, Austria). The TD is a dual aerosol inlet system consisting of impact collection thermal desorption cells. The setup was already used in several campaigns as described by Holzinger et al. (2013); (2010a).

In short, the centrepiece of both aerosol inlets is a Collection Thermal Desorption cell (CTD, Aerosol Dynamics, Berkeley, CA, USA), on which humidified ambient particles in the size range of 70 nm to 2 μm at an air sample flow rate of ~6 L min<sup>-1</sup> are collected by impaction onto a stainless steel collection surface using a sonic jet impactor. The humidification of the aerosol sample flow to approximately 70% is achieved by a Nafion based humidifier and reduces particle rebound. All tubing in contact with volatilized aerosol compounds (i.e. the CTD cell, and all transfer tubing and valves) is coated to increase the chemical inertness of the surface. The CTD cell coating is AMCX (AMCX, L.L.C., Lemont PA, USA); all other parts received the Siltek®/Sulfinert®- treatment. The transfer lines are operated at elevated temperatures of 200 °C to avoid re-condensation of desorbed aerosol compounds.

In this study, aerosols were sampled from the chamber through a ~5 m long copper line (ID=6.5 mm). The operation of the system was fully automated. One cycle was completed in 2.5 h and included the analysis of (i) the first aerosol inlet (namely inlet A), (ii) the second aerosol inlet (namely inlet B), (iii) inlet A and (iv) inlet B that sampled particle-filtered chamber air, and (v) the analysis of gas-phase in conventional PTR-MS mode. The duration of each section was 30 min. Due to lab air contamination the conventional PTR-MS gas-phase measurements of the chamber air were not available from the TD-PTR. In addition, inlet A data quality was affected by a systematic change of the PTR-MS conditions (E/N fluctuation during background measurements caused by a malfunctioning valve). Consequently, inlet A data were excluded from this campaign.

The aerosols were pre-concentrated onto the CTD cell for 30 min with a flow of 6 L min<sup>-1</sup> before thermal desorption into the PTR-MS. After collection, a small flow of ~ 10 mL min<sup>-1</sup> of nitrogen carrier gas transported all compounds desorbing from the CTD cell directly into the PTR-MS. Aerosol compounds were thermally released from the CTD-cell by ramping the temperature up to 350 °C from room temperature (normally, 25 °C). Temperature ramped continuously at a rate of ~15 °C min<sup>-1</sup> for ~21 minutes until 350 °C followed by a dwell time of 3 minutes (at 350 °C). After a cool down period of 6 min a new collection was initiated. For the last experiment (tree emissions), a denuder was installed on inlet B to constrain a possible artefact from gas-phase compounds adsorbing on the CTD cell.

326 The aerosol background was measured every other run by passing the airstream through a Teflon membrane filter  
(Zefluor 2.0  $\mu\text{m}$ , Pall Corp.) that removed the particles from the air stream (sections: iii and iv mentioned above).  
328 The effective removal of particles was confirmed by test measurements with a condensation particle counter (TSI,  
WCPC Model 3785). While particles are removed by the Teflon filter, gas-phase compounds should be less affected.  
Filter samples to determine the aerosol background have been taken by turns: in each cycle, inlet A and inlet B  
330 sampled successively for 30 min of each, then the samples collected through the two inlets were analysed  
successively as well.

332 The PTRMS measures mixing ratios of compounds desorbed from aerosols in a nitrogen carrier gas. The mass  
concentration of an aerosol compound in the air sample is calculated according to

$$334 \quad n_{\text{aer},x} = C_{\text{mean},x} \times \frac{F_{\text{N}_2} \times t_{\text{meas}}}{22.4 \times F_{\text{col}} \times t_{\text{col}}}, \quad (2)$$

where  $n_{\text{aer},x}$  is the aerosol concentration of compound X in  $\mu\text{g m}^{-3}$ ,  $C_{\text{mean},x}$  its (arithmetic) mean mixing ratio during  
336 the aerosol analysis in the nitrogen carrier gas in  $\text{nmol mol}^{-1}$ ,  $MW_x$  the molecular weight of compound X in  $\text{g mol}^{-1}$ ,  
 $F_{\text{N}_2}$  the flow of the carrier gas in standard liters per minute,  $t_{\text{meas}}$  the duration of the aerosol measurement in minutes,  
338  $F_{\text{col}}$  the flow rate at which the aerosols are collected in standard liters per minute,  $t_{\text{col}}$  the duration of aerosol  
collection in minutes and 22.4 the volume one mole of an ideal gas will occupy in liters. Mixing ratios of most  
340 compounds were calculated according to the method described in Holzinger et al. (2010b), which involves the use of  
default reaction rate constants ( $3 \times 10^{-9} \text{ cm}^3 \text{ s}^{-1} \text{ molecule}^{-1}$ ),

342 Specific conditions during the campaign were as follows:  $E/N = 1.6 \times 10^{-19} \text{ V m}^2 \text{ molec}^{-1}$  (i.e. 160 Townsend units) to  
ensure ionization only by  $\text{H}_3\text{O}^+$ , temperature of the drift tube  $T_d = 120 \text{ }^\circ\text{C}$ , and a mass resolution of  $m/\Delta m \approx 4000$ .

344 Mass spectra were obtained on a 5s time resolution. The data were processed using the PTRwid software (Holzinger,  
2015). The software has several unique features including autonomous and accurate calibration of mass scale and  
346 the export of a uniform peak list which avoids the same ion being attributed to a slight different mass within the  
limits of precision. In total, 543 organic ions represented in the “unified mass list” have been obtained.

348

### 3 Results and discussions

350 In order to compare the different measurement techniques a time synchronization of the three data sets was  
performed. All data presented in this work have been synchronized to the ACM time with a time resolution of 4  
352 hours. The presented time is the center of the sampling interval for all experiments.

#### 3.1 Comparison of PTR-based aerosol measurement techniques to SMPS and AMS

354 Comparison of the different aerosol chemical characterization techniques to the AMS and SMPS was performed by  
means of linear regression (Figure 1). Since no collection efficiency (CE) was applied to the PTR-based aerosol  
356 measurement techniques, AMS data were treated the same way throughout this work, thus no AMS CE was  
enforced. SMPS organic mass concentration was calculated assuming a density of  $1.4 \text{ g cm}^{-3}$ , a valid assumption for  
358 SOA (Cross et al., 2007), that represented more than 98 % of the mass as observed from AMS. Each aerosol  
technique was collecting/detecting particles in different size ranges (Table 2). The volume distribution derived from

360 SMPS measurements (Figure S3) covered a particle diameter range of 100 to 400 nm which is within the size  
detection limits of all applied aerosol techniques.

362 A least orthogonal distance regression linear fit function, included in the IGOR extension ODRPack95, was used for  
each instrument related to SMPS data. Results suggested that the measured fraction compared to the SMPS mass  
364 was constant for each technique throughout the campaign. Due to experimental flaws CHARON<sub>100</sub> introduced a  
higher than usual variability of the particle enrichment in the aerodynamic lens during two experiments, the  $\beta$ -pinene  
366 ozonolysis and limonene ozonolysis/NO<sub>3</sub> oxidation (Section 2.3.2). These experiments were excluded when applying  
the linear fit. CHARON<sub>100</sub> was able to measure 80% ( $1\sigma = \pm 10\%$ ) of the SMPS mass. ACM and AMS measured  
368 51% ( $\pm 5\%$ ) and 67% ( $\pm 10\%$ ) while TD measured 27% ( $\pm 3\%$ ) of the SMPS, respectively. TD and ACM showed  
the lowest slope variability ( $\leq 5\%$ ), thus the highest stability in terms of recovery or overall detection efficiency.  
370 CHARON<sub>100</sub> and AMS followed with accuracy of  $\sim 10\%$ , but at higher recovery rates. All instruments showed  
linear fit offset values close to zero when taking into account the error of the fit ( $\pm 3\sigma$ ).

372 For the PTR based techniques and AMS a mass recovery underestimation could be expected due to a variety of  
processes from (i) the unideal CE during particle enrichment, (ii) thermal dissociation during desorption, (iii)  
374 incomplete evaporation or transmission, (iv) ionic dissociation in the ionization region and (v) the inability to ionize  
the reactant/fragment. The extent to which these processes affect the different techniques was investigated in detail  
376 and presented in the following by tracking the path of the particles from collection to detection.

It is well known that AMS derived mass concentrations have to be corrected for CE due to particle bounce signal  
378 loss on the vaporizer (Canagaratna et al., 2007). Fresh biogenic SOA though have a high CE (Kiendler-Scharr et al.,  
2009) and reduced bouncing effect, also observed from the relatively high AMS CE in this work ( $\sim 0.7$ ). ACM and  
380 TD utilize a collection surface as well and therefore introduce a CE uncertainty with the TD setup reducing even  
further the bouncing effects by humidifying the particles prior to collection. CHARON is an on-line technique  
382 avoiding the latter loss processes thus increasing the ability of the instrument to measure the mass concentration of  
the compounds generated during these experiments.

384 During desorption, thermal dissociation of molecules could introduce two or more fragmentation products.  
Canagaratna et al. (2015) reported that in the AMS organics gave rise to H<sub>2</sub>O<sup>+</sup>, CO<sup>+</sup> and CO<sub>2</sub><sup>+</sup> signal due to surface  
386 evaporation and thermal break down of organic molecules at vaporizer operating temperatures down to 200 °C  
(under vacuum conditions). Although neutral dissociation products like H<sub>2</sub>O, CO and CO<sub>2</sub> could be ionized by the  
388 AMS, their proton affinities are lower than that of H<sub>2</sub>O, thus PTR techniques would no longer ionize and detect  
them. On the contrary, remaining smaller organic fragmentation products with proton affinities higher than H<sub>2</sub>O  
390 would still be visible to the PTR-MS. A lack of detection of certain neutral fragments formed during thermal  
desorption could introduce an underestimation of the total mass, oxygen and carbon concentration for the PTR based  
392 techniques. It should be noted that decarboxylation and dehydration reactions are strongly dependent on the  
temperature, pressure and the heat exposure time of the molecules. CHARON was operated at the lowest  
394 temperature of 140 °C, under a few mbars of pressure and with the lowest heat exposure time thus avoiding the  
latter reactions. On the contrary, ACM and TD were operated at 1 bar and up to 250 °C and 350 °C respectively with  
396 longer heat exposure times. To further assess whether surface evaporation for ACM and TD had an additional effect

398 on the measurements, focus was given on the experimental case studies performed by Salvador et al. (2016) using  
the TD-PTR-ToF-MS. Five authentic standard substances (phthalic acid, levoglucosan, arabitol, *cis*-pinonic and  
400 glutaric acid) were utilized to examine the response of the sampling device. If the compounds would only fragment  
in the PTR-ToF-MS due to ionic dissociation, then the detected fragments should have the same volatility trend as  
the parent compounds since both originate from the latter. During desorption of the collected samples, fragment ions  
402 were found to represent different volatility trends compared to their parent ions (Arabitol, *cis*-Pinonic Acid). These  
thermogram differences, originating from the same substance, promoted a certain amount of neutral  
404 fragmentation/pyrolysis in the hot TD cell.

The thermal desorption process varied for the different PTR-based inlet techniques with different desorption  
406 residence times, desorption temperatures and pressure conditions (see section 2.3). Although CHARON was  
operated at lower temperatures compared to ACM and TD, its reduced pressure compensated for the temperature  
408 difference thus increasing the volatility range down to LVOC (Eichler et al., 2017). It could still be though that a  
fraction of the SOA mass in the extremely low volatility OC (ELVOC) range will not evaporate during desorption  
410 from any of the systems studied. If this effect would be significant it would be more pronounced in the presence of  
high percentages of ELVOCs in the aerosol, i.e. during periods with increased O:C ratios (indicated in Figure 2). A  
412 non-linear relationship between SMPS and the PTR based techniques would be the result, which has not been  
observed (Figure 1). We therefore concluded that incomplete evaporation of ELVOC constitutes a minor  
414 contribution to the mass recovery underestimation. Transmission losses of OA vapours on the pathway from  
evaporation to detection could occur on cold spots in between the evaporation zone and the drift tube. All  
416 components were heated to higher temperatures than the evaporation zone in order to avoid these losses. Within the  
drift tube of the PTR the temperature is lower than in the evaporation zone but the lower pressure will reduce but not  
418 exclude the possibility of re-condensation of organic vapours.

Ionic dissociation in the ionization region of the PTR-MS is strongly affected by the PTR operating conditions and  
420 in particular the E/N applied (Section 2.3). The lower mass concentration detected by the TD unit compared to the  
other techniques could be partly explained by the different E/N used, with TD operated at the highest E/N = 160 Td.  
422 This high potential of fragmentation losses during quantification would be given as:



424 where  $(R^+)^*$  is the unstable protonated reactant,  $F^+$  is the protonated fragment and N is the neutral product.  
Commonly occurring neutral fragments are H<sub>2</sub>O from organic hydroxyl functional groups or HNO<sub>3</sub> from organic  
426 nitrate functional groups. While the former is often observed, during our studies organic nitrate fragmentation has  
not been observed as their formation is hindered during our experiments due to low NO<sub>x</sub>-conditions. This has been  
428 supported by AMS derived organic nitrate measurements being below 10% (Figure S4). By increasing the  
fragmentation potential the neutral products would increase thus lowering the total mass concentration detected.  
430 This could also lead to an underestimation of the ACM mass concentration compared to CHARON<sub>100</sub> (ACM  
operated at 120 Td and CHARON<sub>100</sub> at 100 Td) and is discussed in detail in Section 3.2. It should be noted that the  
432 mass underestimation of the ACM due to ionic and thermal dissociation could be higher than 16% (the mass  
difference between the ACM and AMS). This would imply that ACM CE was higher compared to the AMS CE

434 during this campaign, a possible result in view of the differences of vaporizer/collector geometry (Hohaus et al.,  
2010).

436 Additional comparison between the AMS and the PTR-ToF-MS based techniques was examined by determining the  
bulk oxygen to carbon ratio (O:C) for all instruments (Figure 2). AMS O:C values were calculated based on the  
438 method by Canagaratna et al. (2015). All instruments followed similar trends. O:C ratios increased with  
photochemistry initiation (chamber illumination) or NO<sub>3</sub> oxidation (limonene experiment/NO injection). On the  
440 contrary, O:C values decreased when fresh BVOC was introduced into SAPHIR and additional SOA was formed  
during the tree BVOCs re-emission stage (11 – 22 h after ozone injection). When compared to AMS, all PTR-ToF-  
442 MS based techniques showed lower O:C values. Good agreement was found between the ACM and TD O:C values  
( $< 3\%$  difference). CHARON<sub>100</sub> measured higher O:C compared to ACM and TD (ACM lower by  $\sim 20\text{-}35\%$ ), an  
444 indication that during this campaign CHARON<sub>100</sub> was capable of detecting more oxygenated compounds. When  
comparing the  $\beta$ -pinene and limonene experiments, CHARON<sub>100</sub> had increased O:C values for experiments that  
446 incorporated  $\beta$ -pinene while ACM had the opposite behavior, with higher O:C during the limonene experiment. For  
the tree emissions experiment the BVOC system resulted in SOA that showed increased O:C values for all  
448 instruments introducing compounds with higher oxygen content in the particulate-phase. During this experiment  
CHARON was operated at different E/N operating conditions thus providing further insights of the influence of E/N  
450 on O:C values (Figure S5). Results showed that O:C increased by approximately 10% when changing the CHARON  
E/N from 100 Td to 65 Td, thus providing softer ionization conditions.

452 Although nearly all C<sub>x</sub>H<sub>y</sub>O<sub>z</sub> ions can be identified and quantified within the AMS mass spectra, AMS O:C  
calculation based on Canagaratna et al. (2015) has several sources of uncertainties due to correction factors applied.  
454 As stated by Canagaratna et al. (2015), the overall errors observed in elemental ratios calculations would introduce  
an upper uncertainty of 28%. In contrast to AMS data O:C ratios for the PTR based techniques were calculated with  
456 no additional correction factors thus explaining their lower values when compared to AMS.

PTR-ToF-MS is considered a soft ionization technique which suffers less from fragmentation and therefore should  
458 provide O:C ratios closer to the true values compared to uncorrected AMS data. Nevertheless, water clustering and  
fragmentation could occur, either increasing or decreasing O:C ratios. When proton transfer reactions induce  
460 fragmentation a neutral fragment is lost. For oxygenated organics it has been shown that the loss of water as neutral  
fragment is a common fragmentation pathway (de Gouw and Warneke, 2007). This could explain the lower O:C  
462 values seen from CHARON, ACM and TD compared to the AMS. Intercomparison of the PTR based techniques  
further showed that CHARON<sub>100</sub> was more sensitive to oxygenated compounds compared to ACM and TD. Higher  
464 O:C ratios were observed when comparing CHARON<sub>65</sub> to CHARON<sub>100</sub> indicating that low E/N values can decrease  
the loss of neutral fragments such as water or carbon containing compounds with O:C ratios  $>1$  (e.g. CO<sub>2</sub>, HCOOH).  
466 This factor does affect the ACM and TD O:C ratios even more, since they are operated at even higher E/N (120 Td  
and TD at 160 Td, respectively) than CHARON. It should be noted that lower E/N values could also increase the  
468 tendency to detect water clusters, i.e. AH<sup>+</sup>(H<sub>2</sub>O)<sub>n</sub>, where A is the ionized organic compound, bearing the risk to bias  
the O:C ratio high which is explored further in the next section.

470 As previously discussed, AMS  $\text{H}_2\text{O}^+$ ,  $\text{CO}^+$  and  $\text{CO}_2^+$  signals are generated due to surface evaporation at  
temperatures exceeding 200 °C (under vacuum conditions). These fragment signals cannot be detected from ACM  
472 and TD (that also undergo surface evaporation compared to CHARON), thus an additional underestimation of their  
O:C values could not be excluded. To assess the extent of surface fragmentation, further re-calculation of the AMS  
474 O:C, excluding the  $\text{H}_2\text{O}^+$ ,  $\text{CO}^+$  and  $\text{CO}_2^+$  peaks (Figure S6) was performed and compared to the PTR-based  
techniques. Results showed that AMS O:C ratios were lower than O:C ratios of ACM and TD. When only excluding  
476 the  $\text{H}_2\text{O}^+$  signal, AMS O:C ratios were higher than those of ACM and TD. These results suggest that CO and  $\text{CO}_2$   
loss by thermal dissociation in the ACM and TD play a less significant role compared to AMS due to their lower  
478 operating evaporation temperatures and higher pressure.

When comparing experiments incorporating  $\beta$ -pinene or limonene, the different behavior of the O:C ratios found for  
480 the CHARON<sub>100</sub> ( $\text{O:C}_{\text{CHARON, limonene}} < \text{O:C}_{\text{CHARON, } \beta\text{-pinene}}$ ) and ACM ( $\text{O:C}_{\text{ACM, limonene}} > \text{O:C}_{\text{ACM, } \beta\text{-pinene}}$ ) could be  
due to different fragmentation patterns of the particulate-phase functional groups or due to their volatility  
482 differences. Since limonene SOA are less volatile than  $\beta$ -pinene SOA (Lee et al., 2011) a fraction of the OA  
oxygenated mass that would evaporate at higher temperatures could be lost, thus leading to lower O:C values  
484 compared to the  $\beta$ -pinene experiments. However, ACM showed only minor volatility differences when comparing  
the  $\beta$ -pinene to the limonene experiments, as seen in Figure S7. Although CHARON was operated at lower  
486 temperatures compared to ACM, its reduced pressure compensated for the temperature difference thus increasing the  
volatility range down to LVOC (Eichler et al., 2017). These results conclude that differences in the O:C trends of  
488 ACM and CHARON could not be explained by changes of the SOA volatility. The ionic and thermal dissociation  
patterns of the different particulate-phase functional groups could play a role in these findings and has to be  
490 examined in future studies.

### 3.2 Classification of SOA composition

492 Further comparison of the aerosol chemical characterization techniques was performed with a focus on the different  
chemical characteristics (oxygen content, carbon content, molecular weight) of the SOA composition. A desorption  
494 period from the tree emissions experiment, 25 hours after the ozone injection (Figure 2 (d)), was chosen in order to  
highlight the instrument performance differences, shown in Figure 3. The mass concentration of all compounds  
496 containing the same carbon number was calculated. These carbon fractions were then further separated depending  
on the number of oxygen atoms the compounds contained. The molecular weights (MW) of the SOA constituents  
498 was separated in five different  $m/z$  range groups, from  $m/z$  30 - 50,  $m/z$  50 - 100,  $m/z$  100 - 150,  $m/z$  150 - 250,  $m/z$   
>250. All instruments showed similar carbon content distributions, with the highest concentration introduced from  
500 C8 compounds. CHARON was able to measure compounds in the C10 - C20 range while ACM and TD only  
detected up to C13 compounds. The overall OA mass concentration decreased when moving from lower  
502 (CHARON<sub>65</sub> and CHARON<sub>100</sub>) to higher E/N values (ACM at 120 Td and TD at 160 Td). The same trend was seen  
for the oxygen content of compounds; with a characteristic example being the compounds containing 5 oxygen  
504 atoms that decreased by a factor of 2 with the same instrument but different operational parameters for the PTR-  
ToF-MS (CHARON<sub>65</sub> vs. CHARON<sub>100</sub>). In ACM and TD compounds containing 5 oxygens were negligible. A

506 similar trend was observed for  $m/z$  range distributions, with a higher fraction of low  $m/z$  compounds observed at  
increasing E/N values. ACM and TD results indicated that the main fraction of compounds was detected for MW <  
508 100 amu (70 and 75% of the overall mass concentration, respectively).

These results clearly show the high dependency of the overall mass concentration detection as well as the carbon,  
510 oxygen and MW content determination being strongly affected by the PTR-ToF-MS E/N operating conditions. As  
the E/N values increased, fragmentation increased leading to undetected neutral fragments. This loss of information  
512 directly affects the overall mass concentration and MW detection range. Comparing the ACM to the TD MW pie  
charts showed that, although ACM was operated at lower E/N conditions (120 Td) than the TD (160 Td) the  
514 contribution in the lower MW range was higher for the ACM. The reason for this dissimilarity could be due to the  
lower resolution and the higher limit of detection of the PTR-ToF-MS used for the ACM (see Table 2) leading to  
516 lower detection of the higher molecular weight compounds. Since water loss is the major fragmentation occurring in  
the PTR-ToF-MS, the oxygen content is affected the strongest. This could explain why compounds with 5 oxygens  
518 were nearly undetectable for ACM and TD compared to CHARON.

To further assess the differences in chemical classification by each instrument the relative OA mass concentration of  
520 molecular carbon, oxygen and weight (box-and-whiskers including all data points throughout the campaign) were  
used, as seen in Figure 4. ACM and TD showed similar distributions for all contributions throughout the campaign  
522 with only minor differences (< 3%). On the contrary, their comparison to CHARON<sub>100</sub> showed a clear difference.  
Compounds in the lower MW range (<  $m/z$  150), containing lower molecular carbon (< 9 carbon atoms) and oxygen  
524 (< 2 oxygen atoms) showed higher contributions for the ACM and TD compared to CHARON<sub>100</sub>. A detailed  
comparison of CHARON's different E/N conditions during the tree emissions experiment (Figure S8) was also  
526 performed. Results indicated that for lower E/N, an absolute difference of 2%, 5% and 10% for the molecular  
carbon, weight and oxygen contributions were observed, respectively, suggesting that in this E/N range (from 65 to  
528 100 Td) fragmentation is dominated by oxygen containing functional groups loss.

The above results strongly suggest that the E/N settings play a key role to the fragmentation patterns. By increasing  
530 the drift tube voltage, the velocity of the ions increased, leading to higher kinetic energy in ion molecule and  
therefore stronger buffer gas collision. This energy increase was translated to an increase in fragmentation. On the  
532 contrary, the lower the E/N was set, the higher the sensitivity due to enhanced reaction times but also the stronger  
the cluster ion distribution change, supporting more  $H_3O^+(H_2O)_n$  ( $n=1,2,3$ ) cluster ion generation (de Gouw and  
534 Warneke, 2007). In order to quantify whether the PTR-ToF-MS E/N conditions were a major factor for the  
differences seen during this campaign, a case study of pinonic acid was performed in the lab. Monodisperse pinonic  
536 acid particles were generated (900 – 1100 particles/cm<sup>3</sup>) and directed to a CHARON-PTR-ToF-MS, changing E/N  
values from 60 to 170 Td (Figure S9). Results showed that the relative intensity of the parent ion decreased rapidly  
538 when increasing the E/N values. At the same time, the relative intensity of the lightweight fragments was increasing.  
The effect of the parent ion clustering with water was negligible suggesting no overestimation of the CHARON  
540 oxygen content at low E/N (65 Td). By assuming a uniform sensitivity and calculating the total signal (parent ion  
and fragments, assuming all  $m/z$  represent parent molecules) the mass fraction of pinonic acid particles was  
542 calculated (Figure S10). The higher the E/N values were set, the less the PTR-ToF-MS measured compared to the

SMPS. These results confirmed our previous findings that fragmentation losses lead to an underestimation of the overall mass concentration. Therefore the different E/N conditions of the detection systems (PTR-ToF-MS) could explain in a large fraction the differences between the CHARON, ACM and TD oxygen and carbon content (results seen in Figure 2 and Figure 4) as well as their differences in the overall detectable mass (results seen in Figure 1 and Figure 3). A clear influence of the aerosol sampling technique on the differences of these parameters cannot be determined nor excluded (Salvador et al., 2016).

### 3.3 Volatility comparison

During the campaign, CHARON was operated at a constant temperature (140 °C) while ACM and TD ramped through different temperatures during desorption of the collected aerosol samples (see Section 2). The ramping of ACM and TD provided the possibility of a detailed comparison of the compound dependent volatility trends. In Figure 5 the timeseries of ACM and TD for the  $\beta$ -pinene, the  $\beta$ -pinene/limonene mixture and the tree emissions experiments were investigated. The limonene ozonolysis and NO<sub>3</sub> oxidation was excluded from this comparison, due to TD operational problems. For both instruments high contributions of the aerosol mass concentration evaporated at lower temperatures when fresh SOA were generated (initial hours of the experiments and tree emissions A<sub>0</sub> stage), hence higher SOA volatility values were observed. As oxidation continued the relative contributions of aerosol mass evaporating at low temperatures and therefore the overall volatility decreased. When illuminating the chamber, SOA volatility decreased suggesting that photochemical aging of the SOA took place leading to a change of the chemical composition and volatility distribution. For experiments having  $\beta$ -pinene as a precursor for the subsequent SOA formation, TD showed a decreasing volatility as the experiment evolved while ACM reached a plateau after 5 to 10 hours of aging.

The volatility changes for both instruments, during the initial hours of the experiments and during the re-introduction of BVOCs for the trees experiment, could be attributed to the high concentration semi-volatile organic compounds (SVOCs) in the gas-phase that had the maximum available surface to condense on (SMPS at its maximum surface area and mass concentration). Under these conditions, these compounds would partition more to the particulate-phase thus increasing their contribution during the highest concentration periods. These easier to evaporate SVOCs could change the volatility patterns by a change of the thermograms during the maximum concentration periods, as observed from both techniques. Discrepancies between the ACM and TD, with the latter having a steadily changing desorption temperature with time, could be affected by several operating differences. During evaporation ACM was ramped by 100 °C min<sup>-1</sup> to a maximum of 250 °C, with 3-minute isothermal sections at 100 °C, 150 °C and 250 °C, respectively, while TD was ramped continuously at a rate of ~15 °C min<sup>-1</sup> for ~21 minutes until 350 °C. The higher volatility resolution of TD compared to ACM could introduce an increased sensitivity to volatility changes thus increase the TD variability compared to ACM. Differences could be partly attributed to the different design of the instruments. ACM ensured complete separation of the particulate from the gas-phase (> 99.9999 gas-phase removal) while TD was corrected for gas-phase contamination by performing background measurements (Section 2). As the collection of the particulate-phase compounds was performed for the TD, the collector was exposed to high concentration of SVOCs from the gas-phase, thus increasing the absorption of



these compounds to the particulate-phase. As the gas-phase concentrations decreased the TD volatility decreased. This could thus indicate a possible background correction artifact mostly affecting compounds in the higher volatility range, evaporating in the first temperature steps (100 °C). It should be noted that after the  $\beta$ -pinene initial hours of consumption, secondary reactions in the absence of light and the presence of ozone should be negligible due to the lack of unsaturated reactants. The expected temporal volatility behavior would thus be shifted towards a more stable instead of changing volatility system.

To further assess the volatility differences of ACM and TD, focus was given on the molecular oxygen number based on the assumption that oxygen number correlates to volatility (Jimenez et al., 2009). Box-and-whiskers, including all campaign desorption periods, were generated for each molecular oxygen number at each temperature, as seen in Figure 6. The data were normalized to the sum of the measured mass concentration from each molecular oxygen number in all temperatures (top equation in Figure 6). Results showed that TD had a broader range in fractional contribution for all oxygen bins when compared to the ACM. A characteristic temperature showing this difference was at 150 °C, where TD showed results in the range of 0.2 to 0.55 while ACM was in the range from 0.15 to 0.25. Despite the differences in relative contribution, both instruments showed similar trends. As the collector temperature increased oxygenated compounds (2, 3 and 4 oxygens) contributed more than lower oxygenates. On the contrary, at lower temperatures compounds containing 0 and 1 oxygen were the dominant factor. Overall, for ACM around 20% of the SOA evaporated at 100 °C, 20% at 150 °C and 60% at 250 °C. TD showed similar volatility trends with 15 to 20% of the SOA evaporating at 100 °C, 35% at 150 °C and 50 to 55% at 250 °C.

According to observations and theory (Jimenez et al., 2009) oxygenated compounds are expected to have lower volatility thus evaporating at higher temperatures. TD and ACM described the expected volatility trends during the performed experiments based on compound specific information in accordance to theory. The variability of TD compared to ACM reflected the differences in the design and operation of the individual systems described previously. The higher volatility resolution but also the higher E/N conditions of TD could explain most of the observed discrepancies. Fragmentation due to ionic dissociation after the evaporation could influence the volatility molecular oxygen content distribution by loss of neutral oxygen containing fragments. This could further affect the volatility distribution when the oxidation product concentrations change with time, reflected by the increase of the O:C ratios (see Figure 2). Furthermore, the ability of ACM to achieve complete gas to particle separation resulted in a lower thermogram uncertainty in the higher volatility range thus smaller variations. These results show the applicability of both techniques to study BSOA volatility trends in a compound specific level.

### 3.4 Compound detection comparison and tracers attribution

The molecular formula ( $C_xH_yO_zN_a$ ) was attributed to each detected signal derived from the exact molecular mass (see Section 2) determined by the TOF-MS for all 3 techniques throughout the campaign. In order to assess whether major contributing molecules with the same chemical characteristics were determined by all instruments, a comparison of the dominant signals was performed i.e. the molecular formulas that (i) were measured by all techniques during each experiment and (ii) were within the 80 highest signal concentrations. Figure 7 shows the respective results from the BSOA detected in the C7 to C10 range with varying oxygen content (from 0 to 4

oxygen). Although these techniques could provide the molecular formula of the compounds, the molecular structures are unknown. In order to derive further information, comparison to previous publications was performed for the major oxidation products from (a) the  $\beta$ -pinene ozonolysis (Chen and Griffin, 2005; Hohaus et al., 2015; Jenkin, 2004; Yu et al., 1999), (b) limonene ozonolysis and  $\text{NO}_3$  oxidation (Chen and Griffin, 2005; Jaoui et al., 2006; Kundu et al., 2012; Leungsakul et al., 2005a; Leungsakul et al., 2005b) and (c) tree emissions ozonolysis with  $\alpha$ -pinene and  $\Delta^3$ -carene being the major reactants (Chen and Griffin, 2005; Praplan et al., 2014; Yu et al., 1999). Results showed that all techniques were able to detect most of the expected molecules. Details on the molecular formula and suggested structure are provided in more detail in Table S1. Due to fragmentation most of the compounds were not detected at the parent ion molecular weight but underwent water loss in accordance to the findings that O:C ratios are observed to be reduced by ACM, TD and CHARON compared to the AMS (see Section 3.1). These compounds corresponded to a large fraction of the BSOA mass measured from each technique (bars in Figure 7). On average, 70%, 60%, and 40% of the measured mass was contributed from these compounds, for ACM, CHARON and TD respectively. When comparing the above compounds concentration to the SMPS total mass, around 30%, 50% and 10% of the SMPS mass for ACM, CHARON and TD respectively was explained. The overlapping of detected compounds to previous publications (theoretical and experimental work) and their high contribution (up to 50%) to the overall BSOA mass concentration strongly promotes the use of PTR-ToF-MS aerosol measurement techniques to gain valuable insight on the chemical characteristics of BSOA.

#### 632 4 Conclusions

A comparison of three different aerosol chemical characterization techniques has been performed as part of a chamber study on fresh and photochemically aged BSOA, formed from the ozonolysis of monoterpenes. The aerosol collection module (ACM), the chemical analysis of aerosol on-line (CHARON) and the collection thermal desorption unit (TD) are different aerosol sampling inlets utilizing a PTR-ToF-MS. These techniques were deployed in a set of chamber experiments at the atmosphere simulation chamber SAPHIR to investigate SOA formation and aging from different monoterpenes ( $\beta$ -pinene, limonene) and from real plant emissions (*Pinus sylvestris L.*).

The total aerosol concentration recovery of the PTR based techniques, compared to an SMPS, was  $80 \pm 10\%$ ,  $51 \pm 5\%$  and  $27 \pm 3\%$  for CHARON, ACM and TD respectively. In contrast, an AMS concurrently operated and with no collection efficiency correction applied, showed a recovery of 67%. The three PTR based techniques were capable of measuring the same major contributing signals for the different monoterpene oxidation products studied. These attributed compounds corresponded to a high fraction of the overall SOA mass concentration with 30%, 50% and 10% of the overall mass being explained for ACM, CHARON and TD respectively. Additional comparison to previous publications showed that these compounds corresponded to known products of the monoterpenes studied. Both the ACM and TD collection and thermal desorption design provided additional information on their volatility and showed similar trends. Compounds containing higher molecular oxygen number ( $\geq 2$ ) contributed more to the aerosol fraction desorbed at high temperatures than lower oxygenated compounds (molecular oxygen number  $< 2$ ) which were more efficiently desorbed at low temperatures.

650 Oxygen to carbon ratios (O:C) increased while SOA production and aging proceeded. All instruments had  
comparable O:C trends during the course of an experiment. Good agreement was found for the ACM and TD O:C  
652 values (< 3% difference) while CHARON showed 20 to 35% higher O:C ratios.

Despite significant difference in the aerosol collection and desorption techniques, the major reason for the  
654 discrepancies was the different operating conditions of the PTR-ToF-MS. Laboratory case studies supported that  
E/N conditions played a crucial role in fragmentation leading to lower O:C ratios at high E/N. Since ACM and TD  
656 were operated at higher E/N compared to CHARON this resulted to higher fragmentation, thus affecting their  
oxygen and carbon content and mass recovery. Compared to AMS, PTRMS is a soft ionization technique even at  
658 high E/N and therefore less prone to fragmentation. AMS requires correction factors (Canagaratna et al., 2015), to  
determine O:C ratios whereas for PTRMS corrections were omitted. Determination of O:C ratios for the PTR based  
660 techniques was thus underestimated, explaining their difference to the HR-ToF-AMS (30 to 50% higher).  
Differences in the sampling and evaporation technique might introduce also deviations between the chemical  
662 characterizations i.e. due to thermal decomposition. This has to be studied in detail in future comparisons by  
operating the PTR-ToF-MS instruments under the same E/N conditions.

664 The ability of all PTR based techniques to measure compounds, supported from previous publications, strongly  
promotes their use. These techniques can provide valuable insight on the chemical characteristics of freshly formed  
666 and aged BSOA, and on thermodynamic properties such as partitioning coefficient values and volatility patterns on a  
compound specific level.

668

#### **Author contribution**

670 RT, RH, AW and AKS designed the experiments. TH and RT operated the chambers. SHS, PS, PE, MM, KM, GIG,  
RW, MK, AW, RH and RT conducted the data collection and evaluation for AMS, TD, CHARON, ACM, PTR and  
672 GC-MS. MM designed and carried out the laboratory characterization experiments. GIG, RT, TH and AKS did the  
data analysis. GIG did the data interpretation and prepared the manuscript with contributions from all co-authors.

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#### **References**

682 Aiken, A. C., P. F. DeCarlo, J. H. Kroll, D. R. Worsnop, J. A. Huffman, K. S. Docherty, I. M. Ulbrich, C. Mohr, J.  
R. Kimmel, D. Sueper, et al.: O/C and OM/OC Ratios of Primary, Secondary, and Ambient Organic Aerosols with  
684 High-Resolution Time-of-Flight Aerosol Mass Spectrometry, *Environ Sci Technol*, 42(12), 4478-4485,  
doi:10.1021/es703009q, 2008.

686

688 Canagaratna, M. R., J. T. Jayne, J. L. Jimenez, J. D. Allan, M. R. Alfarra, Q. Zhang, T. B. Onasch, F. Drewnick, H.  
Coe, A. Middlebrook, et al.: Chemical and microphysical characterization of ambient aerosols with the aerodyne  
690 aerosol mass spectrometer, *Mass Spectrom Rev*, 26(2), 185-222, doi:10.1002/mas.20115, 2007.

692 Canagaratna, M. R., J. L. Jimenez, J. H. Kroll, Q. Chen, S. H. Kessler, P. Massoli, L. Hildebrandt Ruiz, E. Fortner,  
L. R. Williams, K. R. Wilson, et al.: Elemental ratio measurements of organic compounds using aerosol mass  
694 spectrometry: characterization, improved calibration, and implications, *Atmospheric Chemistry and Physics*, 15(1),  
253-272, doi:10.5194/acp-15-253-2015, 2015.

696 Chen, J., and R. Griffin: Modeling secondary organic aerosol formation from oxidation of  $\alpha$ -pinene,  $\beta$ -pinene, and  
limonene, *Atmospheric Environment*, 39(40), 7731-7744, doi:10.1016/j.atmosenv.2005.05.049, 2005.

698 Cross, E. S., J. G. Slowik, P. Davidovits, J. D. Allan, D. R. Worsnop, J. T. Jayne, D. K. Lewis, M. Canagaratna, and  
700 T. B. Onasch: Laboratory and Ambient Particle Density Determinations using Light Scattering in Conjunction with  
Aerosol Mass Spectrometry, *Aerosol Science and Technology*, 41(4), 343-359, doi:10.1080/02786820701199736,  
702 2007.

704 de Gouw, J., and C. Warneke: Measurements of volatile organic compounds in the earth's atmosphere using proton-  
transfer-reaction mass spectrometry, *Mass Spectrom Rev*, 26(2), 223-257, doi:10.1002/mas.20119, 2007.

706 DeCarlo, P. F., J. R. Kimmel, A. Trimborn, M. J. Northway, J. T. Jayne, A. C. Aiken, M. Gonin, K. Fuhrer, T.  
708 Horvath, K. S. Docherty, et al.: Field-Deployable, High-Resolution, Time-of-Flight Aerosol Mass Spectrometer,  
*Analytical Chemistry*, 78(24), 8281-8289, doi:10.1021/ac061249n, 2006.

710 Eichler, P., M. Müller, B. D'Anna, and A. Wisthaler: A novel inlet system for online chemical analysis of semi-  
712 volatile submicron particulate matter, *Atmospheric Measurement Techniques*, 8(3), 1353-1360, doi:10.5194/amt-8-  
1353-2015, 2015.

714 Eichler, P., M. Müller, C. Rohmann, B. Stengel, J. Orasche, R. Zimmermann, and A. Wisthaler: Lubricating Oil as a  
716 Major Constituent of Ship Exhaust Particles, *Environmental Science & Technology Letters*, 4(2), 54-58,  
doi:10.1021/acs.estlett.6b00488, 2017.

718 Goldstein, A. H., and I. E. Galbally: Known and Unexplored Organic Constituents in the Earth's Atmosphere,  
720 *Environ Sci Technol*, 41(5), 1514-1521, doi:10.1021/es072476p, 2007.

722 Guenther, A. B., X. Jiang, C. L. Heald, T. Sakulyanontvittaya, T. Duhl, L. K. Emmons, and X. Wang: The Model of  
Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an extended and updated framework for  
724 modeling biogenic emissions, *Geoscientific Model Development*, 5(6), 1471-1492, doi:10.5194/gmd-5-1471-2012,  
2012.

726 Hallquist, M., J. C. Wenger, U. Baltensperger, Y. Rudich, D. Simpson, M. Claeys, J. Dommen, N. M. Donahue, C.  
728 George, A. H. Goldstein, et al.: The formation, properties and impact of secondary organic aerosol: current and  
emerging issues, *Atmos. Chem. Phys.*, 9(14), 5155-5236, doi:10.5194/acp-9-5155-2009, 2009.

730 Hohaus, T., I. Gensch, J. R. Kimmel, D. R. Worsnop, and A. Kiendler-Scharr: Experimental determination of the  
732 partitioning coefficient of  $\beta$ -pinene oxidation products in SOAs, *Physical Chemistry Chemical Physics*, 17(22),  
14796-14804, doi:10.1039/C5CP01608H, 2015.

734 Hohaus, T., U. Kuhn, S. Andres, M. Kaminski, F. Rohrer, R. Tillmann, A. Wahner, R. Wegener, Z. Yu, and A.  
736 Kiendler-Scharr: A new plant chamber facility, PLUS, coupled to the atmosphere simulation chamber SAPHIR,  
*Atmospheric Measurement Techniques*, 9(3), 1247-1259, doi:10.5194/amt-9-1247-2016, 2016.

738 Hohaus, T., D. Trimborn, A. Kiendler-Scharr, I. Gensch, W. Laumer, B. Kammer, S. Andres, H. Boudries, K. A.  
740 Smith, D. R. Worsnop, et al.: A new aerosol collector for quasi on-line analysis of particulate organic matter: the  
Aerosol Collection Module (ACM) and first applications with a GC/MS-FID, *Atmospheric Measurement  
742 Techniques*, 3(5), 1423-1436, doi:DOI 10.5194/amt-3-1423-2010, 2010.

744 Holzinger, R.: PTRwid: A new widget tool for processing PTR-TOF-MS data, *Atmospheric Measurement*  
746 *Techniques*, 8(9), 3903-3922, doi:10.5194/amt-8-3903-2015, 2015.

748 Holzinger, R., A. H. Goldstein, P. L. Hayes, J. L. Jimenez, and J. Timkovsky: Chemical evolution of organic aerosol  
750 in Los Angeles during the CalNex 2010 study, *Atmospheric Chemistry and Physics*, 13(19), 10125-10141,  
752 doi:10.5194/acp-13-10125-2013, 2013.

754 Holzinger, R., A. Kasper-Giebl, M. Staudinger, G. Schauer, and T. Röckmann: Analysis of the chemical  
756 composition of organic aerosol at the Mt. Sonnblick observatory using a novel high mass resolution thermal-  
758 desorption proton-transfer-reaction mass-spectrometer (hr-TD-PTR-MS), *Atmospheric Chemistry and Physics*,  
760 10(20), 10111-10128, doi:10.5194/acp-10-10111-2010, 2010a.

762 Holzinger, R., J. Williams, F. Herrmann, J. Lelieveld, N. M. Donahue, and T. Röckmann: Aerosol analysis using a  
764 Thermal-Desorption Proton-Transfer-Reaction Mass Spectrometer (TD-PTR-MS): a new approach to study  
766 processing of organic aerosols, *Atmos. Chem. Phys.*, 10(5), 2257-2267, doi:10.5194/acp-10-2257-2010, 2010b.

768 Isaacman, G., N. M. Kreisberg, L. D. Yee, D. R. Worton, A. W. H. Chan, J. A. Moss, S. V. Hering, and A. H.  
770 Goldstein: Online derivatization for hourly measurements of gas- and particle-phase semi-volatile oxygenated  
772 organic compounds by thermal desorption aerosol gas chromatography (SV-TAG), *Atmospheric Measurement*  
774 *Techniques*, 7(12), 4417-4429, doi:10.5194/amt-7-4417-2014, 2014.

776 Jaoui, M., E. Corse, T. E. Kleindienst, J. H. Offenberg, M. Lewandowski, and E. O. Edney: Analysis of Secondary  
778 Organic Aerosol Compounds from the Photooxidation of d-Limonene in the Presence of NOX and their Detection in  
780 Ambient PM2.5, *Environ Sci Technol*, 40(12), 3819-3828, doi:10.1021/es052566z, 2006.

782 Jenkin, M. E.: Modelling the formation and composition of secondary organic aerosol from  $\alpha$ - and  $\beta$ -pinene  
784 ozonolysis using MCM v3 *Atmos. Chem. Phys.*, 4, 1741-1757, 2004.

786 Jimenez, J. L., M. R. Canagaratna, N. M. Donahue, A. S. H. Prevot, Q. Zhang, J. H. Kroll, P. F. DeCarlo, J. D.  
788 Allan, H. Coe, N. L. Ng, et al.: Evolution of Organic Aerosols in the Atmosphere, *Science*, 326(1525),  
790 doi:10.1126/science.1180353, 2009.

792 Jordan, A., S. Haidacher, G. Hanel, E. Hartungen, L. Märk, H. Seehauser, R. Schottkowsky, P. Sulzer, and T. D.  
794 Märk: A high resolution and high sensitivity proton-transfer-reaction time-of-flight mass spectrometer (PTR-TOF-  
796 MS), *International Journal of Mass Spectrometry*, 286(2-3), 122-128,  
doi:<http://dx.doi.org/10.1016/j.ijms.2009.07.005>, 2009.

798 Kanakidou, M., J. Seinfeld, S. Pandis, I. Barnes, F. Dentener, M. Facchini, R. V. Dingenen, B. Ervens, A. Nenes,  
800 and C. Nielsen: Organic aerosol and global climate modelling: a review, *Atmospheric Chemistry and Physics*, 5(4),  
802 1053-1123, 2005.

804 Kiendler-Scharr, A., Q. Zhang, T. Hohaus, E. Kleist, A. Mensah, T. F. Mentel, C. Spindler, R. Uerlings, R.  
806 Tillmann, and J. Wildt: Aerosol Mass Spectrometric Features of Biogenic SOA: Observations from a Plant Chamber  
808 and in Rural Atmospheric Environments, *Environ Sci Technol*, 43(21), 8166-8172, doi:10.1021/es901420b, 2009.

810 Kreisberg, N. M., S. V. Hering, B. J. Williams, D. R. Worton, and A. H. Goldstein: Quantification of Hourly  
812 Speciated Organic Compounds in Atmospheric Aerosols, Measured by an In-Situ Thermal Desorption Aerosol Gas  
814 Chromatograph (TAG), *Aerosol Science and Technology*, 43(1), 38-52, doi:10.1080/02786820802459583, 2009.

816 Kundu, S., R. Fisseha, A. L. Putman, T. A. Rahn, and L. R. Mazzoleni: High molecular weight SOA formation  
818 during limonene ozonolysis: insights from ultrahigh-resolution FT-ICR mass spectrometry characterization,  
820 *Atmospheric Chemistry and Physics*, 12(12), 5523-5536, doi:10.5194/acp-12-5523-2012, 2012.

798 Lee, B.-H., J. R. Pierce, G. J. Engelhart, and S. N. Pandis: Volatility of secondary organic aerosol from the  
800 ozonolysis of monoterpenes, *Atmospheric Environment*, 45(14), 2443-2452, doi:10.1016/j.atmosenv.2011.02.004,  
2011.

802 Leungsakul, S., M. Jaoui, and R. M. Kamens: Kinetic Mechanism for Predicting Secondary Organic Aerosol  
804 Formation from the Reaction of d-Limonene with Ozone, *Environ Sci Technol*, 39(24), 9583-9594,  
doi:10.1021/es0492687, 2005a.

806 Leungsakul, S., H. E. Jeffries, and R. M. Kamens: A kinetic mechanism for predicting secondary aerosol formation  
808 from the reactions of d-limonene in the presence of oxides of nitrogen and natural sunlight, *Atmospheric  
Environment*, 39(37), 7063-7082, doi:10.1016/j.atmosenv.2005.08.024, 2005b.

810 Liu, P., P. J. Ziemann, D. B. Kittelson, and P. H. McMurry: Generating Particle Beams of Controlled Dimensions  
812 and Divergence: I. Theory of Particle Motion in Aerodynamic Lenses and Nozzle Expansions, *Aerosol Science and  
Technology*, 22(3), 293-313, doi:10.1080/02786829408959748, 1995a.

814 Liu, P., P. J. Ziemann, D. B. Kittelson, and P. H. McMurry: Generating Particle Beams of Controlled Dimensions  
816 and Divergence: II. Experimental Evaluation of Particle Motion in Aerodynamic Lenses and Nozzle Expansions,  
*Aerosol Science and Technology*, 22(3), 314-324, doi:10.1080/02786829408959749, 1995b.

818 Lopez-Hilfiker, F. D., C. Mohr, M. Ehn, F. Rubach, E. Kleist, J. Wildt, T. F. Mentel, A. Lutz, M. Hallquist, D.  
820 Worsnop, et al.: A novel method for online analysis of gas and particle composition: description and evaluation of a  
Filter Inlet for Gases and AEROSols (FIGAERO), *Atmospheric Measurement Techniques*, 7(4), 983-1001, doi:DOI  
10.5194/amt-7-983-2014, 2014.

822 Martinez, R. E., B. J. Williams, Y. Zhang, D. Hagan, M. Walker, N. M. Kreisberg, S. V. Hering, T. Hohaus, J. T.  
824 Jayne, and D. R. Worsnop: Development of a volatility and polarity separator (VAPS) for volatility- and polarity-  
resolved organic aerosol measurement, *Aerosol Science and Technology*, 50(3), 255-271,  
doi:10.1080/02786826.2016.1147645, 2016.

826 Mensah, A. A., R. Holzinger, R. Otjes, A. Trimborn, T. F. Mentel, H. ten Brink, B. Henzing, and A. Kiendler-  
828 Scharr: Aerosol chemical composition at Cabauw, The Netherlands as observed in two intensive periods in May  
2008 and March 2009, *Atmospheric Chemistry and Physics*, 12(10), 4723-4742, doi:10.5194/acp-12-4723-2012,  
830 2012.

832 Müller, M., T. Mikoviny, W. Jud, B. D'Anna, and A. Wisthaler: A new software tool for the analysis of high  
834 resolution PTR-TOF mass spectra, *Chemometrics and Intelligent Laboratory Systems*, 127, 158-165,  
doi:10.1016/j.chemolab.2013.06.011, 2013.

836 Praplan, A. P., S. Schobesberger, F. Bianchi, M. P. Rissanen, M. Ehn, T. Jokinen, H. Junninen, A. Adamov, A.  
838 Amorim, J. Dommen, et al.: Elemental composition and clustering of  $\alpha$ -pinene oxidation products for different  
oxidation conditions, *Atmospheric Chemistry and Physics Discussions*, 14(22), 30799-30833, doi:10.5194/acpd-14-  
30799-2014, 2014.

840 Rohrer, F., B. Bohn, T. Brauers, D. Brüning, F. J. Johnen, A. Wahner, and J. Kleffmann: Characterisation of the  
842 photolytic HONO-source in the atmosphere simulation chamber SAPHIR, *Atmos. Chem. Phys.*, 5(8), 2189-2201,  
doi:10.5194/acp-5-2189-2005, 2005.

844 Salvador, C. M., T. T. Ho, C. C. K. Chou, M. J. Chen, W. R. Huang, and S. H. Huang: Characterization of the  
846 organic matter in submicron urban aerosols using a Thermo-Desorption Proton-Transfer-Reaction Time-of-Flight  
Mass Spectrometer (TD-PTR-TOF-MS), *Atmospheric Environment*, 140, 565-575,  
848 doi:10.1016/j.atmosenv.2016.06.029, 2016.

850 Seinfeld, J., and S. Pandis, ATMOSPHERIC CHEMISTRY AND PHYSICS: From Air Pollution to Climate  
852 Change, second edition, in *WILEY-INTERSCIENCE PUBLICATION*, edited, 2006.

854 Wiedensohler, A., W. Birmili, A. Nowak, A. Sonntag, K. Weinhold, M. Merkel, B. Wehner, T. Tuch, S. Pfeifer, M.  
856 Fiebig, et al.: Mobility particle size spectrometers: harmonization of technical standards and data structure to  
facilitate high quality long-term observations of atmospheric particle number size distributions, *Atmos. Meas. Tech.*,  
5(3), 657-685, doi:10.5194/amt-5-657-2012, 2012.

858 Williams, B. J., A. H. Goldstein, N. M. Kreisberg, and S. V. Hering: An In-Situ Instrument for Speciated Organic  
860 Composition of Atmospheric Aerosols: Thermal Desorption Aerosol GC/MS-FID (TAG), *Aerosol Science and  
Technology*, 40(8), 627-638, doi:10.1080/02786820600754631, 2006.

862 Yu, J., D. R. Cocker, R. J. Griffin, R. C. Flagan, and J. H. Seinfeld: Gas-Phase Ozone Oxidation of Monoterpenes:  
864 Gaseous and Particulate Products, *Journal of Atmospheric Chemistry*, 34(2), 207-258,  
doi:10.1023/a:1006254930583, 1999.

866 Zhao, Y., N. M. Kreisberg, D. R. Worton, A. P. Teng, S. V. Hering, and A. H. Goldstein: Development of an In  
868 Situ Thermal Desorption Gas Chromatography Instrument for Quantifying Atmospheric Semi-Volatile Organic  
Compounds, *Aerosol Science and Technology*, 47(3), 258-266, doi:10.1080/02786826.2012.747673, 2013.

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892 **Table 1: Experimental conditions for each experiment. For the tree emissions experiment there were two VOC injection periods.**

<b>Experiment</b>	<b>Monoterpenes (ppb)</b>	<b>Ozone (ppb)</b>	<b>Duration (h)</b>	<b>Maximum SOA formed (<math>\mu\text{g}/\text{m}^3</math>)</b>	<b>SOA formation conditions</b>	<b>SOA aging Conditions</b>
<b><math>\beta</math>-Pinene</b>	120	700	34	130	Ozonolysis	Photochemical oxidation for 10 h
<b>Limonene</b>	25	150	17	50	Ozonolysis	Continuous $\text{NO}_3$ oxidation for 8 h
<b><math>\beta</math>-Pinene/Limonene mixture</b>	60/12	300	26	60	Ozonolysis	Photochemical oxidation for 4 h
<b>Tree emissions</b> 1 <sup>st</sup> inj. / 2 <sup>nd</sup> inj.	65/10	300	30	80	Ozonolysis	Photochemical oxidation for 6 h

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916 **Table 2: Instruments operating conditions.**

<b>INSTRUMENT CHARACTERISTICS</b>	<b>ACM (in situ)</b>	<b>CHARON (online)</b>	<b>TD (in situ)</b>
<b>Time resolution (min)</b>	240	1	120
<b>Gas/particle separation</b>	High vacuum	Denuder	Denuder and/or blank correction (filtered air)
<b>Pre-concentration factor</b>	21 <sup>a</sup>	44	6000 <sup>b</sup>
<b>LOD<sup>c</sup> (ng/m<sup>3</sup>)</b>	35 <sup>d</sup>	1.4 <sup>e</sup>	0.02 <sup>b</sup>
<b>Temperature range (°C)</b>	25 – 250	140	25 – 350
<b>Heating rate (°C / min)</b>	100	0	15
<b>Temperature steps (°C)</b>	100, 150, 250 (3 min)	none	None
<b>Desorption pressure (atm)</b>	1	< 1	1
<b>Particle range (nm)</b>	70 – 1000	70 – 1000	70 - 2000
<b>PTR-ToF-MS model</b>	8000	8000	8000
<b>Drift tube Temperature (°C) / Pressure (mbar) / Voltage (V)</b>	90 / 2.3 / 550	120 / 2.4 / 400 and 240	120 / 2.25 / 600
<b>PTR-ToF-MS E/N (Td)</b>	120	65 / 100	160
<b>PTR-ToF-MS mass resolution (m/Δm)</b>	2500	4500-5000	4000

<sup>a</sup> based on 240 min sampling at 80 mL/min and 3 min desorption at 300 mL/min

918 <sup>b</sup> based on 30 min sampling at 6 L/min and 3 min desorption at 10 mL/min a typical value for most ions based on the method in (Holzinger et al., 2010a)

920 <sup>c</sup> Limit of detection

<sup>d</sup> For signal on *m/z* 139 and 10 sec integration time

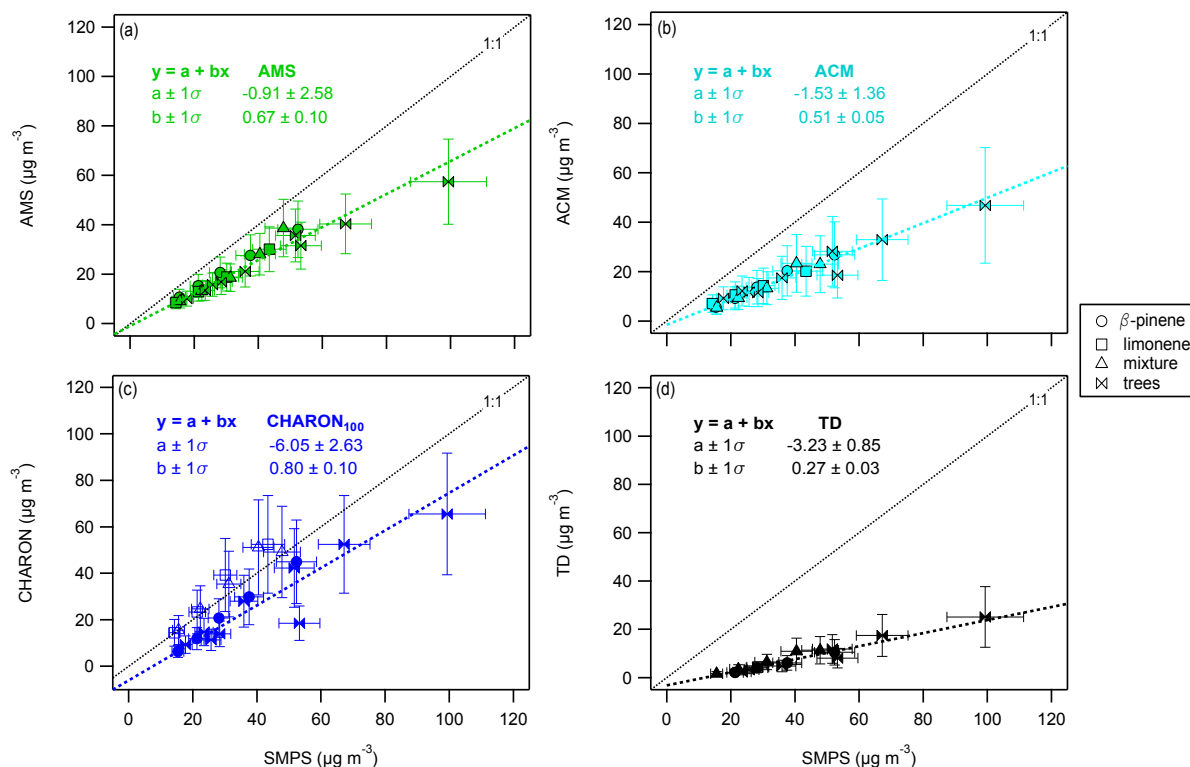
922 <sup>e</sup> For signals around *m/z* 200 and 1 min integration time

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932 **Figure 1: Comparison of the organic mass concentration of (a) AMS (green), (b) ACM (ciel), (c) CHARON<sub>100</sub> (blue) and**  
 934 **(d) TD (black), to the SMPS (x-axis). Markers correspond to the different experiments with the mixture experiment**  
 936 **accounting for the mixture of  $\beta$ -pinene and limonene. AMS data presented are not corrected for collection efficiency.**  
 938 **CHARON<sub>100</sub> corresponds to data taken only at 100 Td E/N operating condition. Error bars provide the uncertainty of**  
 each instrument (details in Section 2.3). A least orthogonal distance regression linear fit is applied for every instrument,  
 taking into account all campaign measurement points. Exception is the CHARON limonene and mixture data (unfilled  
 markers) that were excluded due to experimental flaws. Details of the coefficient values and their standard deviation are  
 given on the upper left of each graph.

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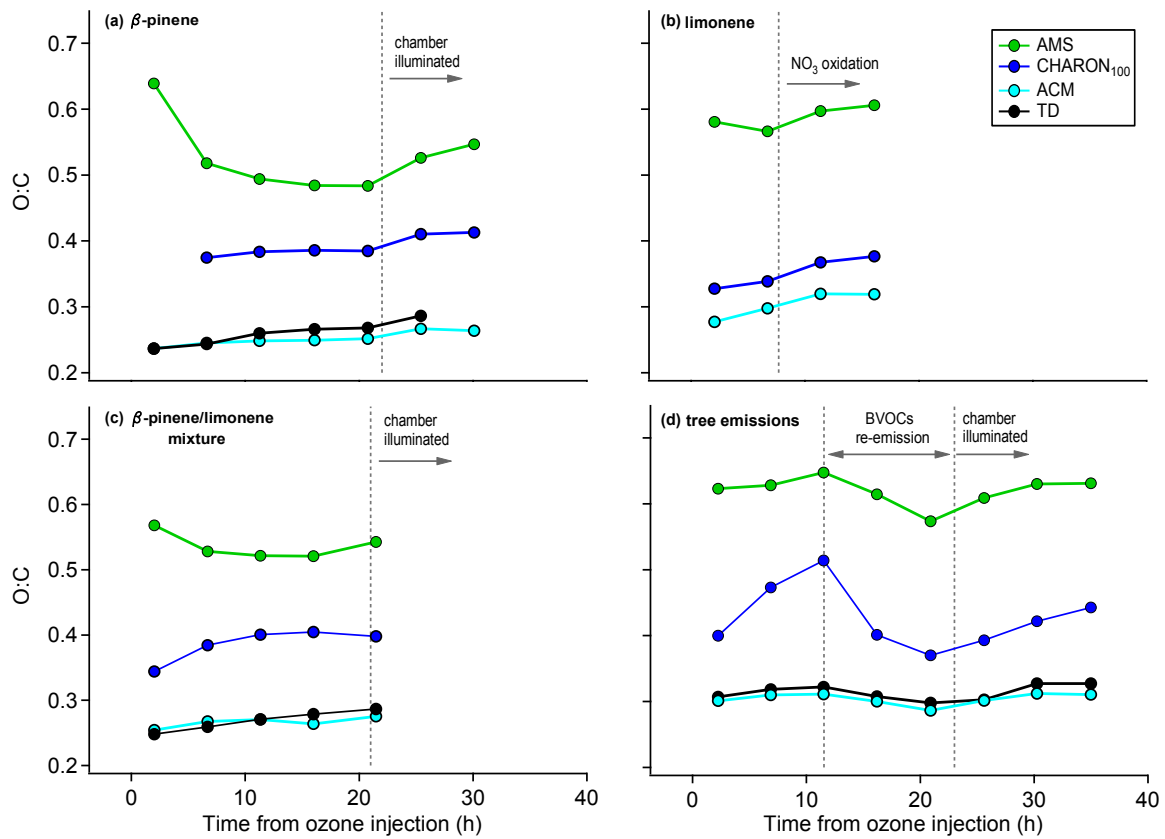


Figure 2: Bulk oxygen to carbon ratio comparison for the different instruments (CHARON<sub>100</sub>: blue, AMS: green, ACM: cyan, TD: black) versus the time from ozone injection. Experimental description details are provided in Table 1.

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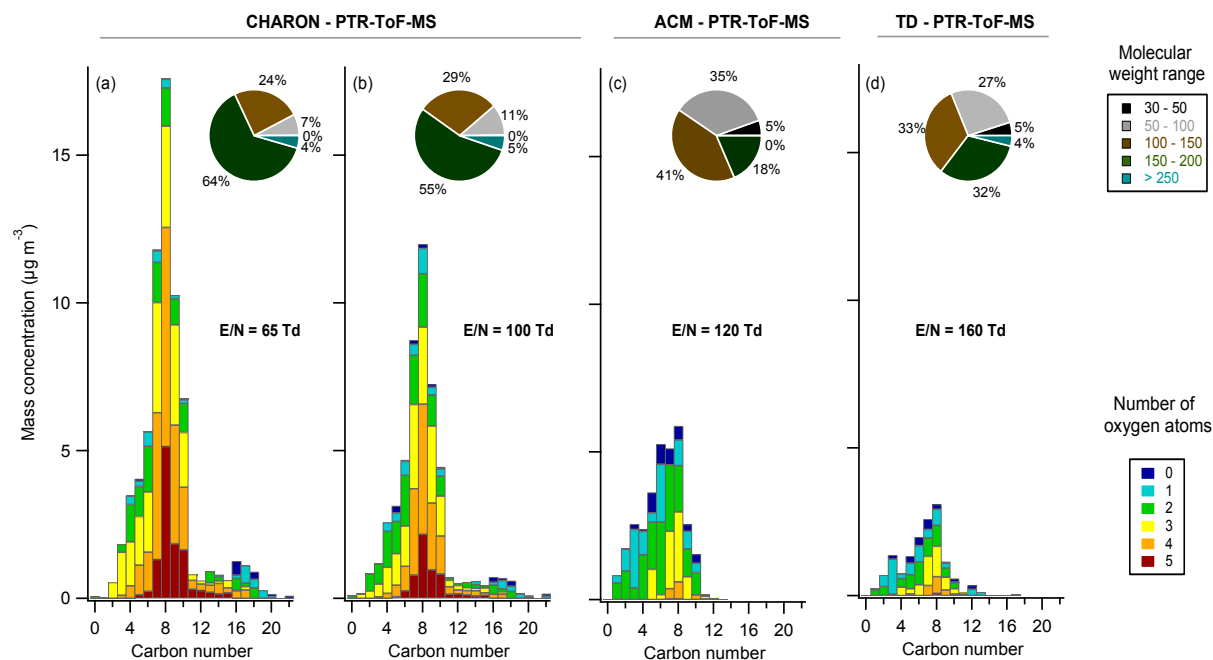
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968 **Figure 3: OA mass concentration (y-axis) distributed based on the number of carbon atoms (x-axis). Bar colours**  
 970 **correspond to the contribution of oxygen atoms starting from 0 (blue) to 5 (red) for each carbon group when (a)**  
 972 **CHARON was operated at E/N = 65 Td, (b) CHARON operated at 100 Td, (c) ACM operated at 120 Td and (d) TD**  
 974 **operated at 160 Td. Pie charts correspond to the molecular weight contribution to the overall mass starting from  $m/z$  30 –**  
 976 **50 (black) up to  $m/z$  > 250 (ciel). Results shown in this graph are from the tree emissions experiment at a high OA mass**  
 978 **concentration, 25 h after the ozone injection (Figure 2 (d)).**

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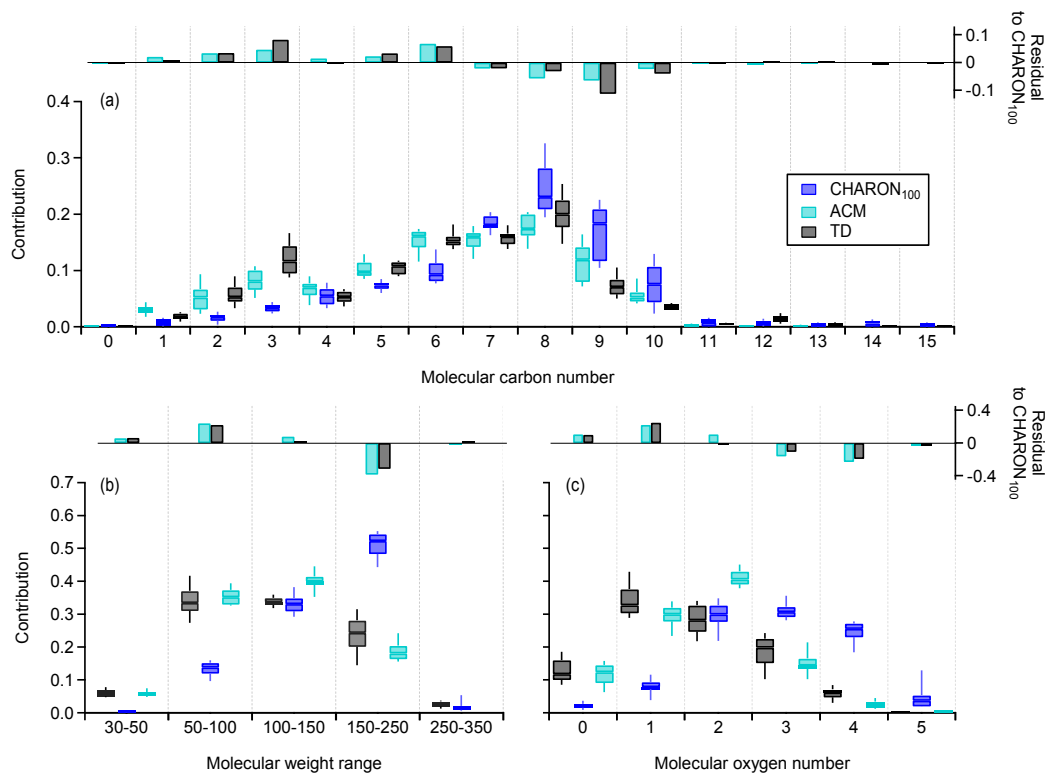
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994 **Figure 4: Box-and-whisker plots showing the relative OA mass concentration distribution dependent on (a) molecular**  
 996 **carbon number, (b) molecular weight and (c) molecular oxygen number for the different instruments, indicated with**  
 998 **different colours (CHARON<sub>100</sub> blue, ACM ciel and TD black). Each box-and-whisker corresponds to the median, 25<sup>th</sup> and**  
 1000 **75<sup>th</sup> percentile levels of all data throughout the campaign. Upper graphs indicate the difference between the ACM and TD**  
 1002 **to the CHARON<sub>100</sub> median values defined as residual to CHARON<sub>100</sub>.**

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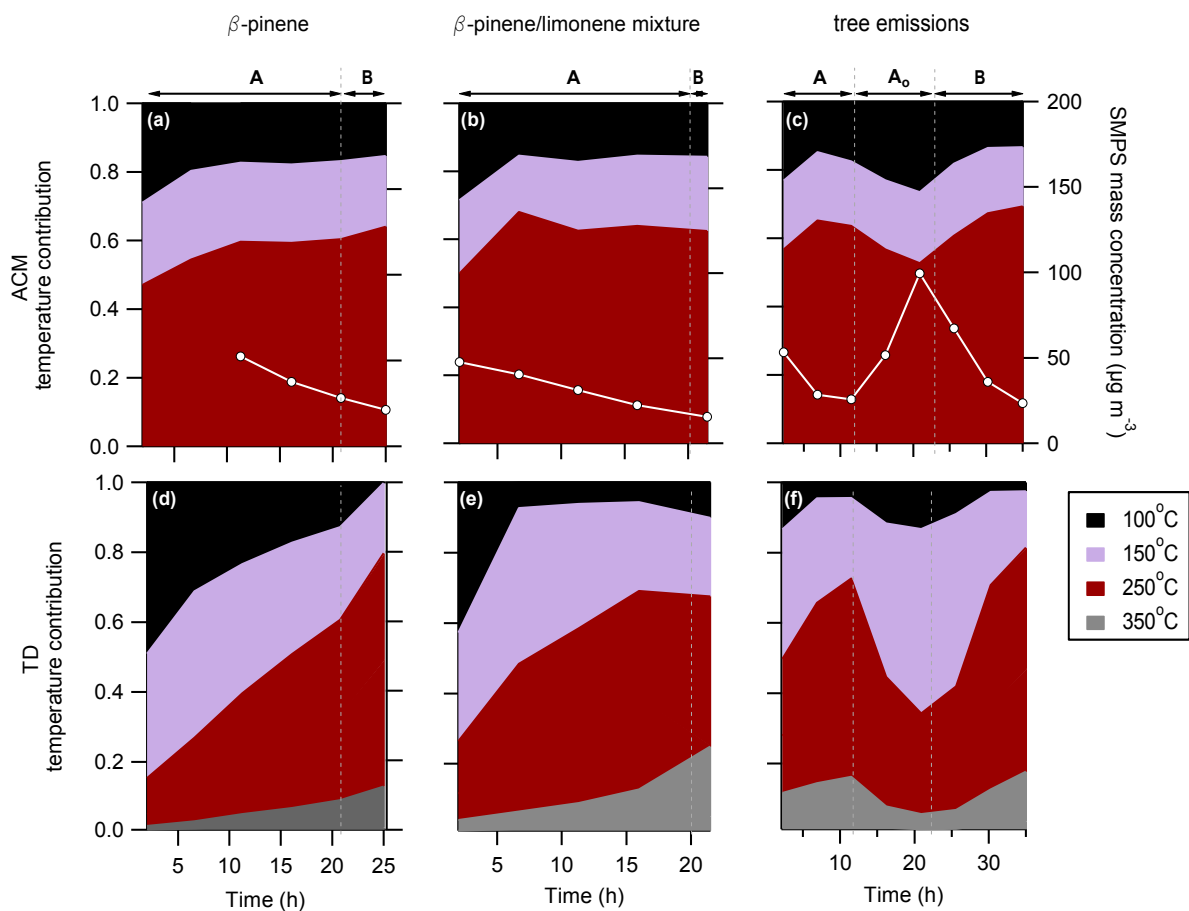
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1018 **Figure 5: Temperature dependent mass concentration contribution (left y-axis) of ACM (upper plots: a, b, c) and TD**  
 1019 **(lower plots: d, e, f) for  $\beta$ -pinene (a, d),  $\beta$ -pinene and limonene mixture (b, e) and real tree emissions (c, f) versus the time**  
 1020 **since ozone injection (x-axis). White lines and circle markers (right y-axis) represent the SMPS mass concentration**  
 1021 **during each experiment. Dash vertical lines indicate the different experimental periods with A: the ozonolysis and SOA**  
 1022 **formation period, B: the chamber illumination and photo-oxidation period and  $A_0$ : the tree emissions BVOCs re-injection**  
 1023 **to the SAPHIR chamber.**

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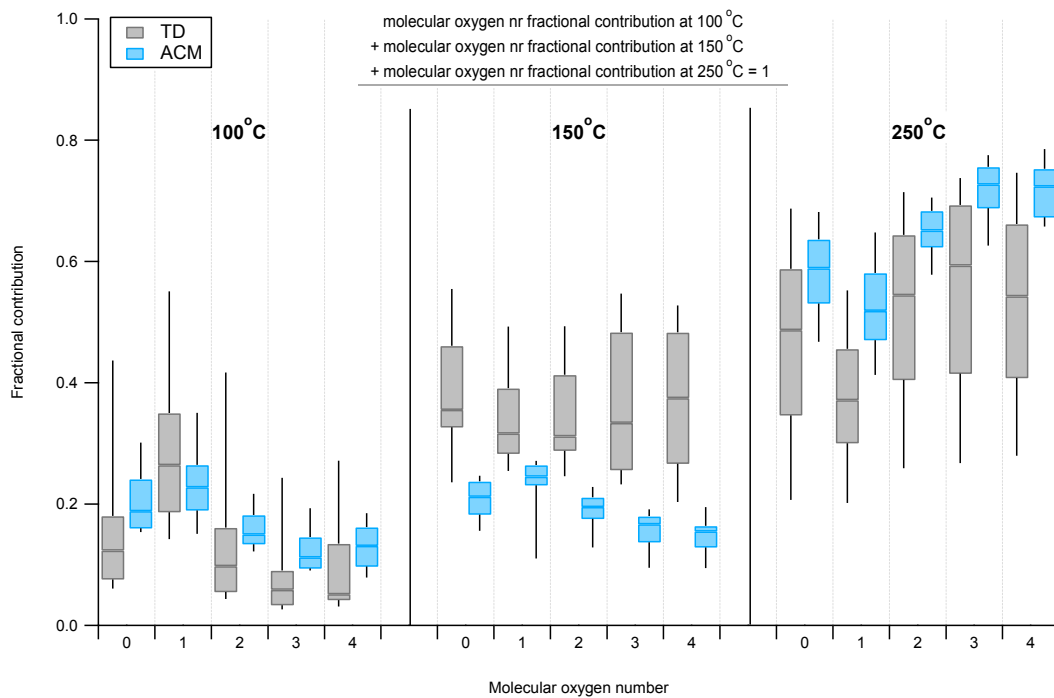
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1038 **Figure 6: Box-and-whisker plots showing the distribution of the molecular oxygen number (x-axis), for the different**  
 1040 **temperature steps (100 °C, 150 °C, 250 °C) of ACM (ciel) and TD (black). Each box-and-whisker corresponds to the**  
 1042 **median, 25<sup>th</sup> and 75<sup>th</sup> percentile levels of all desorption points throughout the campaign. Upper equation indicates how the**  
 1044 **contribution of each molecular oxygen number, at each temperature, corresponds to unity.**

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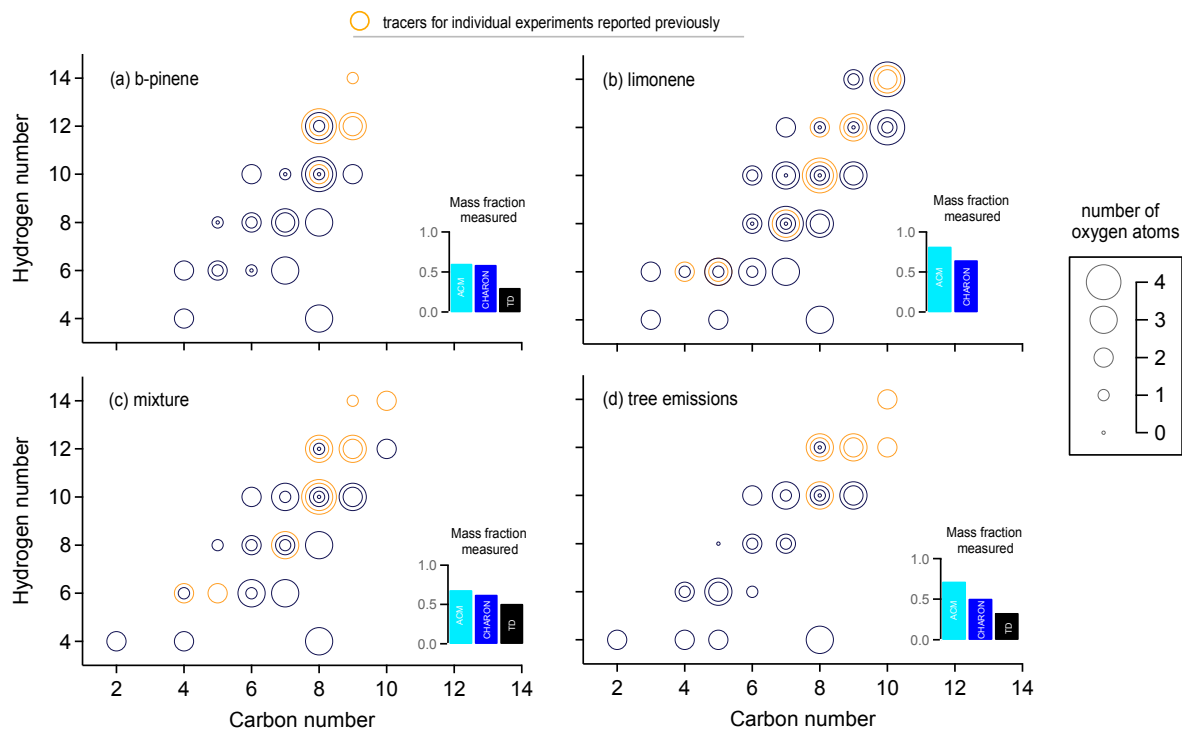
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1060 **Figure 7: Chemical formula attribution based on the molecular carbon number (x-axis), hydrogen number (y-axis) and**  
 1062 **oxygen number (markers size) for (a) the ozonolysis of  $\beta$ -pinene, (b) ozonolysis and  $\text{NO}_3$  oxidation of limonene, (c)**  
 1064 **ozonolysis of the  $\beta$ -pinene/limonene mixture and (d) ozonolysis of real tree emissions (Scotts pine). Markers correspond to**  
 1066 **compounds measured from all techniques (ACM, CHARON and TD) at high concentrations (within the 80 compounds**  
**observing highest concentration). Each circle corresponds to one compound. Orange markers indicate tracer compounds**  
**supported from previous publications (for details refer to Table S1). Bars indicate the fraction of mass explained when**  
**accounting only the presented compounds, for each instrument (ACM cie1, CHARON<sub>100</sub> blue and TD black) based on**  
**their total aerosol mass measured.**