#### **Response to Anonymous Reviewer #1:**

Comment: The CALIPSO aerosol optical depth (AOD) for a particular profile is the sum of the extinction of various features identified as aerosol layers within it. Optically thin or diffuse layers may be missed due to detection limitations, leading to an underestimate of the total aerosol loading. The data product contains retrieval fill values (RFVs) when no aerosol layer is detected. Whether these are counted as zeroes or not when creating aerosol climatologies affects the results. This study builds on previous work to quantify the occurrence of RFVs and estimate, via comparison to MODIS and AERONET, what the missing optical depth in (daytime over-water) CALIPSO data as a result of this is.

This work is within scope for AMT and since CALIPSO is one of only two spaceborne lidar providing aerosol data (the other being CATS, which likely has similar issues for the same reasons), so understanding and correcting for biases, which some users may be unaware of, is unimportant. The authors do a thorough job and have a fairly rigorous approach. I recommend publication following minor revisions. I would be happy to review the next version, although a re-review may not be necessary.

## Response: We thank the reviewer for his/her comments and encouragement.

Comment: A general comment is that the authors have done this analysis with data versions which are all becoming out of date around now (and this is something they acknowledge). For example CALIOP version 4 products are partially or fully released already (and are examined briefly in the paper in section 3.7); MODIS Collection 6.1 and AERONET version 3 are both available within the coming weeks or months. I agree with the authors' statements that this paper is a quantification of the problem and not the final word on the issue, but as some aspects of the data relevant to the analysis change in the latest data versions (e.g. CALIOP version 4 has improved a few detection and calibration issues, AERONET version 3 does a better job of screening out cirrus and not screening out smoke), rerunning the analysis with the latest data versions before final publication then that would be good to keep everything up to date. This is particularly relevant because AERONET, CALIOP, and MODIS all have fairly infrequent update schedules so these new versions are likely to be the latest for several years. Otherwise in the coming few years it may not be clear to readers how quantitatively transferable the results of this analysis are to the data products available at that time.

Response: We thank the reviewer for this suggestion. The reason for using Collection 6 MODIS, Version 3 (V3) CALIOP and Level 2.0 AERONET data is because these were the data sources available when we conducted the study. When the manuscript was in preparation for journal submission to JGR (where it was under review for over 4 months), the Version 4 (V4) CALIPSO products were in the process of being released. In anticipation of these new products, we added a section describing a two-month case study to check the frequency of occurrence of all-RFV profiles in V4 CALIPSO L2 aerosol data. We found only minor differences between V3 and V4, and therefore did not rerun the analysis with the new data. At the time of this writing, the Collection 6.1 MODIS data have just begun release. Still, checking the change log for over water Dark TargetMODISproducts(https://modis-atmosphere.gsfc.nasa.gov/sites/default/files/ModAtmo/C061\_Aerosol\_Dark\_Target\_v2.pdf), the only major change is the modification of the sedimentation mask, which isunlikely to make a significant change to the conclusions of the study. In addition, theMODIS Collection 6.1 data have been only partially released a month ago, and thus arenot used in this study (https://modis-atmos.gsfc.nasa.gov/).Lastly, as we recentlychecked, the V3 Level 2.0 AERONET products are still not available for downloading (asof November 28<sup>th</sup>, 2017).

All of that aside, the point of the paper is not necessarily a quantitative evaluation of the current products on offer. Instead, our primary goal is a conceptualization of the problem, both for future missions and for science inquiries at high latitudes that rely on three-dimensional aerosol information (i.e., radiative forcing inquiries). Within that context, we consider our approach, the consideration of Version 3 and approximation of the effect within Version 4, wholly reasonable.

Comment: A second general comment is that the map projection used in mapped figures (e.g. Figure 3, but all of the maps) is a strange one. It distorts to give a disproportionately high weight to high-latitude areas, which is not only a comparatively small fraction of the Earth but the portion with fewest interesting aerosol features. For example the data gap resulting from Antarctica covers about the same amount of page space as the whole African/Asian dust outflow region. I suggest a better map projection is used. Even the regular equal-angle projection would be better, if not an equal-area projection. Otherwise the eye is naturally drawn away from these areas of most interest.

*Response: We agree with the reviewer that the current map projection distorts the highlatitude areas of the globe. We have created new Figs. 3, 5, and 7 with a new map projection (i.e., Robinson), and added them to the manuscript.* 

Comment: Line 186: For MODIS, I think it makes sense to use the data set Effective\_Optical\_Depth\_Average\_Ocean rather than Effective\_Optical\_Depth\_Best\_Ocean, as the 'average' solution is the one which is used to generate level 3 aerosol products (which are perhaps more heavily used than the level 2 products). There has not to my knowledge been much evaluation of the best vs. average MODIS ocean AODs, but Table 5 in Sayer et al (AMT 2012, doi:10.5194/acp-12-8889-2012), in comparison with limited ship-based data, suggests that the 'average' solution may have smaller bias and RMSE than the 'best' solution. That is relevant given the present study's attempt to use MODIS AODs to quantify the missing aerosol from CALIOP RFVs.

Response: Thank you for the suggestion. We have conducted a one-month (January 2008) case study comparing the 'best' and 'average' solutions for C6 MODIS AOTs, and found little difference between the two. For example, the mean MODIS AOTs for all collocated MODIS/CALIOP points are 0.118 and 0.122 for the 'best' and 'average' solutions, respectively. For only those MODIS points collocated with CALIOP all-RFV profiles, the mean MODIS AOTs are 0.088 and 0.092 for the 'best' and 'average'

solutions, respectively. Thus, due to only minor differences between the datasets, we would like to leave the analysis as currently presented in the paper.

Comment: Line 196 (and also 425): It is true that AERONET data can suffer from cirrus contamination, but this is also true for the satellite products; perhaps an explicit mention of that is warranted. Related to my general comment, the AERONET team's presentations suggest this screening is better in AERONET version 3 than the version 2 the authors are using.

Response: We agree with the reviewer that mentioning this possibility should be stated in the paper. While CALIOP has the capability of detecting optically thin clouds, thin cirrus cloud contamination may exist in MODIS products. We have made the following changes to the text:

Modified a sentence in Section 2.2: "Also, thin cirrus contamination may exist in the MODIS aerosol products (e.g., Toth et al., 2013)."

Comment: Line 201: I believe the AERONET team like people to cite Smirnov et al (RSE 2000, doi:10.1016/S0034-4257(00)00109-7) when discussing the AERONET cloud screening and quality assurance procedures.

# Response: We have added this citation to the paper.

Comment: Section 3.3: I wonder if this information about how collocation is achieved could be moved earlier in the manuscript. It is cumbersome to have results in Section 3.1 refer forward in the paper to a method in Section 3.3. Method description should come before results so the reader can understand what is done without having to flip forwards and then backwards again.

#### Response: We agree, and have moved the collocation description to Section 3.1.

Comment: Line 379: I am not sure why Ichoku et al (2003), which is about MODIS aerosol retrievals in Africa, is being cited in the context of limited over-ocean sampling at high ocean latitudes due to sea ice? This reference should be updated to a more appropriate reference, or removed (since almost all readers will know MODIS does not provide AOD retrievals over sea ice, and that sea ice is common near the poles).

## Response: We have removed the reference.

Comment: Line 474: From the earlier discussion of MODIS, shouldn't this be 0.02-0.04, not 0.02-0.03? Given the MODIS ocean uncertainty estimate for near-zero AOD ranges from -0.02 to +0.04.

Response: Yes. We have made the correction to the text.

Comment: Lines 475-478: This sentence (known biases in V3 CALIOP calibration which have been addressed in V4) is another example why it would be better to update the study to use V4 CALIOP products instead. As the authors note V4 was released last year, and the authors include some CALIOP team members, so I don't understand why the study was performed and submitted using an outdated CALIOP data version.

Response: Thank you for the comment. When the study was originally designed and conducted, V3 was the current version of the CALIOP data. The V4.1 data were released in November 2016, and the initial version of this manuscript was submitted to JGR in February 2017 (where it was in review for over 4 months). We have checked a few months of V4 data, and found no major differences in our results between V3 and V4. Thus, we have left the analysis in the paper using V3 data, as the fundamental conclusions of our study remain unchanged.

Comment: Figure 6: It would be clearer to present this as one panel with three different colored lines (one per region). That would aid the reader in making the comparison between the different latitude ranges.

*Response: We have made the suggested changes and inserted the new figure into the manuscript. The text and figure captions have been edited accordingly.* 

Comment: Table 3: The left column is quite awkward, especially as the descriptions require two additional subscripts in the caption to differentiate certain rows. Perhaps this can be redrawn as a set of check-boxes, i.e. check a box if non-all-RFVs are corrected, check another box if all-RFVs are set to zero, another box if all-RFVs are ignored, etc. That would more clearly and directly show the permutations. Also, the right column is somewhat redundant given it is just column 2 subtracted from column 3. Perhaps some additional statistics could be presented here.

Response: Thank you for this suggestion. We have made the recommended changes to the table (i.e., check boxes and included the standard deviations of each dataset).

#### Paper cited:

Toth, T. D. and coauthors: Investigating enhanced Aqua MODIS aerosol optical depth retrievals over the mid-to-high latitude Southern Oceans through intercomparison with co-located CALIOP, MAN, and AERONET data sets, Journal of Geophysical Research: Atmospheres, 118(10), 4700-4714, 2013.

#### **Response to Reviewer #2 (J. Yorks):**

Comment: This paper identifies the frequencies in which the CALIPSO L2 algorithms fail to detect tenuous aerosol layers (AOT < 0.05) and reports retrieval fill values (RFV) for extinction for the entire column. It also compares these profiles to collocated MODIS and AERONET data to determine AOT is being undetermined/underestimated by CALIPSO. Finally, a method to remedy these RFV profiles is presented. As noted in the conclusion, the main impact of the results shown in the paper, from a data product and lidar algorithm standpoint, is that the CALIPSO L2 aerosol products (AOT, extinction) are underestimated. The method presented for correcting these RFVs is a novel concept and valid method. The paper is well written, clear, and gives proper credit to related work. It deserves to be published with a few minor revisions.

#### Response: Thank you for your positive comments and encouraging review of the paper.

Comment: I have 2 main comments that I believe would strengthen the paper:

1) The "scientific" impact of the work presented in the paper is not well stated. The impacts on lidar data products and processing algorithms are well stated and important, but not everyone that reads the paper will be a "lidar expert". High aerosol loading critically impacts the Earth's radiation budget and air quality, but what is the influence of aerosols at AOTs less than 0.05? To put it bluntly, why should a non-lidar expert care about AOTs of less than 0.05? I think the answer is that, from a climate perspective, they are so frequent that they become important if we ever want to decrease the uncertainties in aerosol radiative effects. I suggest adding a figure that shows the MODIS detection frequencies of AOTs < 0.05 in cloud-free retrievals relative to all cloud-free retrievals (for a few months or even a year of data if possible). Then add a few sentences discussing the figure and point to the potential cumulative impact of these low aerosol loading profiles global aerosol models (Koffi on et al 2012: http://onlinelibrary.wiley.com/doi/10.1029/2011JD016858/full) and the global/regional radiation budget (Use something like Figure 4 from Yang et al. 2009 to determine radiative impact; http://onlinelibrary.wiley.com/doi/10.1029/2009GL039801/full).

Response: We thank the reviewer for these suggestions. The recommended figure (similar to another reviewer's request) has already been published in a previous paper. Figure 14 of Levy et al. (2013) provide histograms of C6 MODIS AOT over oceans. While the 0.05 MODIS AOT bin exhibits the largest frequency, AOTs less than 0.05 comprise roughly 20-25% of the total sample (i.e., estimating from the figure). Another way of looking at this is through Maritime Aerosol Network (MAN) sun photometer derived AOT over oceans, the histograms of which are found in Fig. 4 of Smirnov et al. (2011). Most areas of the global ocean show occurrence frequencies of AOTs less than 0.05 between 10 and 20%. These are much larger for the Southern Ocean (>80%; cleaner aerosol conditions) and smaller (~2%) for the Baltic, Black, and Mediterranean Seas (subject to more air pollution). Thus, we do not show similar histograms in this paper, but have added the following text to the manuscript (Conclusions section):

"Note that this conclusion hints that CALIOP may not detect very thin aerosol layers (i.e., AOTs < 0.05), which account for ~10-20% of the AOT spectrum and are of climatological importance (e.g., Smirnov et al., 2011; Levy et al., 2013). Also, these CALIOP-undetected thin aerosol layers are important for various applications, ranging from data assimilation to aerosol indirect effects."

Comment: 2) The section discussing the anticipation of CALIPSO V4 data products is lacking some important details. The study uses V3 CALIPSO data and V6 MODIS data, but new releases have been made (CALIPSO) or will be made shortly (MODIS). Section 3.7 shows that the frequency of RFV profiles doesn't change dramatically with CALIPSO V4 data products, and points out the important improvements to the L1 calibration and impacts. However, do any of the improvements to L2 retrievals impact your study? Surely changes in cloud-aerosol discrimination or surface detection can also impact aerosol detection and likely play a role in some of the differences in all-RFV frequencies observed. Please add a few sentences in section 3.7 on this impact. Also, there is no discussion about how MODIS V6.1 may change the statistics of MODIS AOT for CALIPSO RFV profiles. Since that data hasn't been released yet, you can't do a reanalysis yet, but please add a few sentences on this topic. I'm not too familiar with what changes will be made for MODIS, so it is possible that none of the changes will impact your results. If that is the case, please let the reader know because that strengthens your paper.

Response: Thank you for the comment. As mentioned to Reviewer 1, checking the change log for over water Dark Target MODIS products (https://modisatmosphere.gsfc.nasa.gov/sites/default/files/ModAtmo/C061\_Aerosol\_Dark\_Target\_v2.p df), the only major change is the modification of the sedimentation mask, which is unlikely to make a significant change to the conclusions of the study. We did not add new comments to the paper as the MODIS Collection 6.1 data have been only partially released a month ago, after submission of this paper (https://modisatmos.gsfc.nasa.gov/).

For CALIOP data, we have revised the sentence in Section 3.7 to:

"Specifically, V4 data feature improved calibrations of Level 1 (L1) backscatter, as well as improved cloud-aerosol discrimination and surface detection, that may increase the detection sensitivity of diffuse aerosol layers that are reflected in L2 aerosol extinction retrievals."

Further, and as we describe to Reviewer 1, the purpose of the paper was never necessarily a quantitative evaluation of the current products on offer. Instead, we are really stressing a conceptualization of the problem, both for future missions and for science inquiries at high latitudes that rely on three-dimensional aerosol information (i.e., radiative forcing inquiries). We recognize the relative inconsistences. But, in that primary context, we still think that the consideration of Version 3 CALIPSO and approximation of the effect within Version 4 is reasonable. Thanks again.

Comment: Minor comments/suggestions:

Line 112: The phrase "believed likeliest" is awkward to read. I suggest rewording it.

Response: We have edited this phrase to "it is likely".

Comment: Line 445: The fixed lidar ratio of 29 sr is appropriate, but I would include the standard deviation computed in Kim et al. 2017 along with a few words about the fact that the value was derived from constrained lidar ratios over ocean and represents background aerosols within the entire tropospheric column. Otherwise, the reader has to look up the paper to find out that information. (Note: for the future, it would be interesting to see the values of 532 nm lidar ratios that are measured by the LaRC HSRL during NAAMES).

Response: Thank you for the comment. We have revised the sentence to:

"The aerosol extinction profiles for all-RFVs are derived in two steps. First, using an assumed lidar ratio of 29 sr (standard deviation of 10 sr; derived from constrained lidar ratios over ocean and represents background aerosols for the entire atmospheric column; Kim et al., 2017), an unconstrained extinction solution is generated from 20 km to the top of the surface-attached layer (3.5 km)."

Comment: Table 2: I suggest adding columns for the standard deviation of the MODIS and AERONET distributions.

Response: The standard deviations of the MODIS and AERONET distributions were added to the table as suggested.

Comment: Tables 3: I suggest highlighting rows 2 and 3 because it is a key result of your work. I also suggest adding columns for the standard deviation of the MODIS and CALIPSO AOT distributions.

*Response: Table 3 has been edited to account for both of these suggestions. We have also added the following sentence to the table caption.* 

"Key results are highlighted in yellow."

Papers cited:

Levy, R. C., Mattoo, S., Munchak, L. A., Remer, L. A., Sayer, A. M., Patadia, F., and Hsu, N. C.: The Collection 6 MODIS aerosol products over land and ocean, Atmos. Meas. Tech., 6, 29893034, doi:10.5194/amt-6-2989-2013, 2013.

Smirnov, A. and coauthors: Maritime aerosol network as a component of AERONET – first results and comparison with global aerosol models and satellite retrievals, Atmos. Meas. Tech., 4, 583-597, doi:10.5194/amt-4-583-2011, 2011.

#### **Response to Anonymous Reviewer #3:**

Comment: The paper presents and discusses the CALIOP detectability problem of tenuous aerosol layers with backscatter below the algorithm noise floor. This technical issue is critical since it propagates into CALIOP climatological AOT studies and based on the selected approach introduces artificial underestimations or overestimations to detected AOT features. The paper is not only limited to addressing the issue. The paper quantifies the related AOT of retrieval fill values (RFV) over ocean (daytime), through comparison with MODIS Aqua DT AOD and AERONET coastal sites, performs a proof-of-concept exercise to correct the artificial effect of RFV values, introduces the nighttime problem and refers to the CALIPSO improved V4.

The study falls within the scope of AMT. The authors have done a thorough job and have a rigorous approach. The manuscript is well-written/structured, the presentation clear, the language fluent and the quality of the figures high. Furthermore, the authors give credit to related work and the results support the conclusions. I recommend publication following minor revisions.

## Response: We thank the reviewer for his/her comments and warm encouragement.

Comment: 1) CALIOP methodology: The description of the methodology is not sufficient. In the Datasets section the authors state that "prior to analysis, advanced QA procedures are performed on the L2\_05kmAProf product. This QA scheme is similar to that employed in Campbell et al. (2012) and Winker et al. (2013), detailed descriptions of which are also outlined in our most recent CALIOP-based study (Toth et al., 2016)". This section is of high importance since the scientific methods, assumptions, the validity of the conclusions are based on the preprocessing of the CALIOP data. Although proper reference is given, a short summary of the methodology would help the reader to follow.

#### Response: This is a nice suggestion. We have revised the sentence to:

"This QA scheme is similar to that employed in Campbell et al. (2012) and Winker et al. (2013), and involves several parameters included in the L2\_05kmAProf product: Extinction\_Coefficient\_532 ( $\geq 0$  and  $\leq 1.25$  km<sup>-1</sup>), Extinction\_QC\_532 (= 0, 1, 2, 16, or 18), CAD\_Score ( $\geq -100$  and  $\leq -20$ ), and Extinction\_Coefficient\_Uncertainty\_532 ( $\leq 10$  km<sup>-1</sup>). The Integrated\_Attenuated\_Backscatter\_532 ( $\leq 0.01$  sr<sup>-1</sup>) parameter from the L2 5 km Aerosol Layer (L2\_05kmALay) product is also used as a QA metric."

Comment: 2) In page 11, lines 256-258 and for Figure 2c the authors state that "... L2 CALIOP profiles collocated with MODIS AOT between 0.03 and 0.07." The reason of the selected boundaries 0.03/0.07 is not clear.

Response: These boundaries were arbitrarily selected to represent low aerosol loading scenarios in order to limit the influence of retrieval failures during high AOT cases. The following sentence appears in the text: "This restriction is meant to limit the influence of layer misclassifications and occasional QA failures, and in particular relatively high

AOT cases where unusually high TAB could influence the mean profile." To be clearer, the following was added to the first occurrence of the 0.03-0.07 MODIS AOT restriction: "(i.e., arbitrarily selected for low aerosol loading scenarios)".

Comment: 3) I would suggest the authors to provide similar histograms of all over ocean Aqua MODIS AOT (#) for the same domains (used in figure 4) and of AERONETnumber of AOT (presented in figure 8), in order for a reader to be able to visualize the differences between the different sensors, apart from just MODIS and AERONET statistical values (mean/median-table 2). This would strengthen the scientific question through the simple visual comparison of the different histograms. The figures could be added either in comparison with the already existing figures or as supplementary files in the end of the paper.

Response: Thank you for this suggestion. The histograms of MODIS and AERONET AOT the reviewer is referring to have been published in other papers. For example, C6 MODIS AOT histograms are shown in Fig. 14 in Levy et al. (2013). Also, histograms of Maritime Aerosol Network (MAN; a component of AERONET) sun photometer derived over-ocean AOT are shown in Fig. 4 of Smirnov et al. (2011). Thus, we do not include similar histograms in this paper. However, we have added the following to make the reader aware of these plots:

In Section 3.3: "We also note, for the reference of the reader, that histograms of C6 MODIS AOT (not collocated with CALIOP) are provided in Levy et al. (2013)."

In Section 3.4: "We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al. (2011)."

Comment: 4) Figure 4. The exhibited distributions between the three domains are not similar in terms of the first MODIS AOT bin, between 0 and 0.01. Figures 4a and 4c are characterized by large number of CALIOP profiles, both all-RFV and all, larger than the following bin between 0.01 and 0.02. This characteristic reverses for the 30S to 30N domain. This feature is interesting and may deserve some justification.

Response: Thank you for bringing this to our attention. We have checked on the MODIS AOT distribution in the first bin (0 to 0.01). The vast majority (>99%) of these points are MODIS AOTs = 0. Thus, this feature in the histograms indicates that MODIS AOTs = 0 are more frequent in the 30N to 60N and 60S to 30S domains compared to the 30S to 30N region, likely because the 30S to 30N region is more aerosol polluted (as indicated by Fig. 5b).

Comment: 5) Figure 6. The exhibited distributions between the three domains are not similar in terms of the last AOT Aqua MODIS bins in the 60S-30S domain. Figures 6a and 6b are characterized by a decreasing percentage with increasing Aqua MODIS AOT values (0.2-0.3). This characteristic reverses in Figure 6c. This feature is interesting and may deserve more attention.

Response: Thank you for the comment. We suspect this is due to the relatively "clean" aerosol environment in the Southern Hemisphere (SH) compared to that of the Northern Hemisphere (NH). Thus, MODIS AOTs greater than 0.2 are less frequent in the SH than NH (see the histograms in Fig. 4), resulting in an increasing percentage of CALIOP all-RFV profiles due to the limited number of data points at this higher AOT range.

Comment: 6) Although the paper's purpose is the description of the RFV problem, the quantification of the RFV problem in CALIPSO V4 may be more interesting than for the outdated CALIPSO V3.

Response: Thank you for the comment. As mentioned in the responses to the other reviewers, when the analysis was originally performed, the V4 products were in initial release. We have completed a case study using the V4 products and included the results in Section 3.7. Minor differences in the results were found between V3 and V4, and the overall conclusions of the study do not change.

Papers cited:

Levy, R. C., Mattoo, S., Munchak, L. A., Remer, L. A., Sayer, A. M., Patadia, F., and Hsu, N. C.: The Collection 6 MODIS aerosol products over land and ocean, Atmos. Meas. Tech., 6, 29893034, doi:10.5194/amt-6-2989-2013, 2013.

Smirnov, A. and coauthors: Maritime aerosol network as a component of AERONET – first results and comparison with global aerosol models and satellite retrievals, Atmos. Meas. Tech., 4, 583-597, doi:10.5194/amt-4-583-2011, 2011.

# Minimum Aerosol Layer Detection Sensitivities and their Subsequent Impacts on Aerosol Optical Thickness Retrievals in CALIPSO Level 2 Data Products

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16 Abstract.

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18 Due to instrument sensitivities and algorithm detection limits, Level 2 (L2) Cloud-19 Aerosol Lidar with Orthogonal Polarization (CALIOP) 532 nm aerosol extinction profile 20 retrievals are often populated with retrieval fill values (RFVs), which indicate the absence of detectable levels of aerosol within the profile. In this study, using four years 21 22 (2007-2008 and 2010-2011) of CALIOP Version 3 L2 aerosol data, the occurrence 23 frequency of daytime CALIOP profiles containing all RFVs (all-RFV profiles) is studied. In the CALIOP data products, the aerosol optical thickness (AOT) of any all-RFV profile 24 is reported as being zero, which may introduce a bias in CALIOP-based AOT 25 climatologies. For this study, we derive revised estimates of AOT for all-RFV profiles 26 27 using collocated Moderate Resolution Imaging Spectroradiometer (MODIS) Dark Target 28 (DT) and, where available, Aerosol Robotic Network (AERONET) data. Globally, all-29 RFV profiles comprise roughly 71% of all daytime CALIOP L2 aerosol profiles (i.e., 30 including completely attenuated profiles), accounting for nearly half (45%) of all daytime 31 cloud-free L2 aerosol profiles. The mean collocated MODIS DT (AERONET) 550 nm 32 AOT is found to be near 0.06 (0.08) for CALIOP all-RFV profiles. We further estimate a 33 global mean aerosol extinction profile, a so-called "noise floor", for CALIOP all-RFV 34 profiles. The global mean CALIOP AOT is then recomputed by replacing RFV values with the derived noise floor values for both all-RFV and non-all-RFV profiles. This 35 36 process yields an improvement in the agreement of CALIOP and MODIS over-ocean 37 AOT.

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# 39 1 Introduction and Motivation

40 Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) measurements provide critical information on aerosol vertical distribution for studies involving aerosol 41 42 modeling (e.g., Campbell et al., 2010; Sekiyama et al., 2010; Yu et al., 2010; Zhang et al., 43 2011; 2014), air quality (e.g., Martin, 2008; Prados et al., 2010; Toth et al., 2014), aerosol 44 climatic effects (e.g., Huang et al., 2007; Chand et al., 2009; Tesche et al., 2014; Thorsen 45 and Fu, 2015; Alfaro-Contreras et al., 2016;), and aerosol climatologies (Pappalardo et al., 2010; Wandinger et al., 2011; Amiridis et al., 2015; Toth et al., 2016). In addition, the 46 47 column-integrated aerosol optical thickness (AOT) derived from Level 2 (L2) CALIOP 48 532 nm observations is also widely used, in comparing and combining with passive-49 based L2 aerosol retrievals, for a comprehensive understanding of regional and global 50 aerosol optical properties (e.g., Redemann et al., 2012). Two such passive-based systems

51 are Aqua Moderate Resolution Imaging Spectroradiometer (MODIS), due to its 52 proximity to CALIOP in the "A-Train" satellite constellation (Levy et al., 2013), and 53 Aerosol Robotic Network (AERONET) sun photometers, which is the primary means for 54 validation of satellite AOT retrievals (Holben et al., 1998).

55 It is well-documented that a discrepancy exists between CALIOP-derived AOTs and those from MODIS data (i.e., CALIOP retrievals lower than MODIS counterparts), 56 57 albeit invoking varying quality-assurance (QA)/quality control (QC) procedures across 58 different timeframes and spatial domains (e.g., Kacenelenbogen et al., 2011; Kittaka et al., 59 2011; Redemann et al., 2012; Kim et al., 2013; Ma et al., 2013). These studies tend to attribute the AOT differences to either uncertainties/cloud contamination in the MODIS 60 61 retrieval, or incorrect selection of the lidar ratio (extinction-to-backscatter ratio; 62 Campbell et al., 2013) when deriving CALIOP aerosol extinction, and subsequent AOT. 63 In a similar fashion, CALIOP AOTs have been evaluated against AERONET-derived AOTs, with the disparities (CALIOP lower) attributed to incorrect CALIOP lidar ratio 64 65 assumptions, cloud contamination, and differences in instrument viewing angles 66 (Schuster et al., 2012; Omar et al., 2013).

While some studies cite the failure to detect tenuous aerosol layers as a possible factor in the aforementioned AOT discrepancy (Kacenelenbogen et al., 2011; Rogers et al., 2014), the extent to which these layer detection failures contribute to the AOT differences between multiple sensors has not been fully quantified. For L2 CALIOP profiles, an extinction coefficient retrieval is performed only for those range bins where aerosol backscatter is detected above the algorithm noise floor. Otherwise, the bins are assigned fill values (retrieval fill values, or RFVs) within the corresponding profile (i.e., -

74 9999.00s; Vaughan et al., 2009; Winker et al., 2013). In fact, all L2 CALIOP extinction 75 profiles contain a non-zero percentage of RFVs. It is thus critical to recognize that since 76 lidar-derived AOTs reflect the integration of range-resolved extinction retrievals, in the absence of multi-spectral instruments (i.e., Raman and high spectral resolution lidars 77 78 [HSRLs]), there will always be range bins where aerosol is present below the detection 79 thresholds of the instrument. Indeed, even in relatively "clean conditions", low 80 extinction but geometrically deep aerosol loadings can integrate to significant AOT 81 contributions (Reid et al., 2017).

82 For a fairly large subset of CALIOP daytime measurements, no aerosol is 83 detected anywhere within a column and hence no aerosol extinction retrieved. This 84 results in an aerosol extinction profile consisting entirely of RFVs (defined as CALIOP all-RFV profiles in this study). Assigning aerosol extinction coefficients to 0.0 km<sup>-1</sup> to 85 86 replace fill values during integration of the extinction coefficient profile results in a 87 corresponding column AOT equal to zero. Note that this scenario further includes those 88 profiles reduced to fill values in the process of applying QA procedures on a per-bin basis 89 (e.g., Campbell et al., 2012; Winker et al., 2013). Thus, it is plausible that a column 90 exhibiting significant AOT may be underestimated in those cases where the aerosol 91 backscatter is both highly diffuse and unusually deep, and thus consistently falls below 92 the algorithm detection threshold.

93 The RFV issue is essentially a layer detectability problem, which has been
94 previously investigated in regional validation studies. For example, Rogers et al. (2014)
95 evaluated CALIOP layer and total-column AOT with the use of collocated HSRL data.
96 Minimum detection thresholds for aerosol extinction were estimated as 0.012 km<sup>-1</sup> at

97 night and 0.067 km<sup>-1</sup> during daytime (in a layer median context). From a column98 integrated perspective, CALIOP algorithms were found to underestimate AOT by about
99 0.02 during nighttime (attributed to tenuous aerosol layers in the free troposphere).
100 During daytime, due to the influence of the solar background signal, CALIOP algorithms
101 were unable to detect about half of weak (AOT < 0.1) aerosol profiles.</li>

102 At first glance, the RFV issue may seem superfluous, and one easily resolved in a 103 subsequent study. In fact, the issue has already caused some confusion within the 104 literature. For example, some studies (e.g., Redemann et al., 2012; Kim et al., 2013; and 105 Winker et al., 2013) include all-RFV profiles (i.e., AOT = 0) for analysis when evaluating 106 climatological AOT characteristics. Campbell et al. (2012; 2013) and Toth et al. (2013; 107 2016), on the other hand, do not include all-RFV profiles while generating climatological 108 averages. Clearly, the first approach introduces an artificial underestimation of mean 109 AOT by including profiles where AOT was not retrieved. The latter, however, 110 presumably leads to an overestimation, since it is likely that all-RFV profiles reflect 111 relatively low AOT cases (i.e., lower than any apparent mean sample value) where 112 CALIOP layer detection exhibits a lack of sensitivity to diffuse aerosol presence that 113 caused nothing to be reported within the column. As a result, Kim et al. (2013) and 114 Winker et al. (2013) report global mean CALIOP AOTs lower than those from Campbell et al. (2012) that does not include the profiles. Other factors (e.g., different temporal 115 116 domains and QA metrics invoked) also contribute to the observed disparity in these global mean AOT computations. This state of affairs indicates a clear need to carefully 117 118 quantify the occurrence frequency of all-RFV profiles on a global scale, and, if possible, 119 derive representative column-integrated AOT values for RFV profiles.

120 Further, and as introduced above, for non-all-RFV profiles there remain range 121 bins with RFVs where low aerosol extinction is likely present (the sum of which, 122 however, can result in a relatively significant AOT). Though some QA can filter obvious 123 cases of attenuation-limited profiles (e.g., require aerosol presence within 250 m of the 124 surface as in Campbell et al., 2012; 2013), the only current remedy otherwise is to accept 125 RFV bins as equal to zero extinction, then integrating to obtain a column AOT estimate. 126 It is compelling to investigate, in a manner similar to Rogers et al. (2014), what this 127 quantitative effect is for climatological analysis.

128 In this paper, using four years (2007-2008 and 2010-2011) of daytime 129 observations from CALIOP, Aqua MODIS, and AERONET, we investigate the RFV 130 issue with an emphasis on the following questions:

(1) What is the frequency of occurrence of all-RFV profiles in the daytime cloud-freeCALIOP data set?

- (2) By collocating MODIS and AERONET AOTs with CALIOP cloud-free all-RFV
  profiles, what is the modal AOT associated with this phenomenon and how
  randomly are the data distributed as a function of passive-derived AOT?
- 136 (3) What is the quantitative underestimation in CALIOP AOT due to RFVs in profiles

137 where extinction is retrieved?

- 138 (4) How much of the discrepancy between MODIS and CALIOP L2 over-ocean AOT
- 139 retrievals can be explained by RFVs and all-RFV profiles?
- 140 We note that the primary CALIOP laser failed in March 2009, forcing the Cloud-Aerosol
- 141 Lidar and Infrared Pathfinder Satellite Observations (CALIPSO) mission team to switch
- 142 to a secondary laser. Therefore, two years of CALIOP aerosol data are analyzed prior to

143 (2007-2008) and after (2010-2011) the switch to investigate any discernible difference in

144 RFV statistics between the two lidar profiles.

145

146 2 Datasets

147 2.1 CALIOP

148 Orbiting aboard the CALIPSO satellite within the "A-Train" constellation 149 (Stephens et al., 2002), CALIOP is a two-wavelength (532 and 1064 nm) polarization-150 sensitive (at 532 nm) elastic backscatter lidar, observing the vertical distribution of 151 aerosols and clouds in Earth's atmosphere since June 2006 (Winker et al., 2010). The 152 532 nm backscatter profiles measured by CALIOP are used to detect aerosol and cloud 153 features and then retrieve corresponding particle extinction and subsequent AOTs (i.e., 154 column-integrated extinction; Young and Vaughan, 2009) within layer boundaries 155 determined by a multi-resolution layer detection scheme (Vaughan et al., 2009) and the assumption of a lidar ratio based upon aerosol or cloud type (Omar et al., 2005; 2009). 156 157 For this study, 532 nm aerosol extinction coefficient data from the Version 3 (V3) 158 CALIPSO L2 5 km Aerosol Profile (L2 05kmAProf) product are utilized (Winker et al., 159 2009; hereafter, all references to CALIOP data imply the 532 nm channel/product). 160 These aerosol profiles are reported in 5 km segments and feature a vertical resolution of 60 m below an altitude of 20.2 km above mean sea level (AMSL). Only CALIOP data 161 collected during daytime conditions are considered for this study, such that comparison 162 with aerosol observations from MODIS and AERONET can be accomplished. 163

Prior to analysis, advanced QA procedures are performed on the L2\_05kmAProf product. This QA scheme is similar to that employed in Campbell et al. (2012) and

Winker et al. (2013), and involves several parameters included in the L2\_05kmAProf 166 product: Extinction Coefficient 532 ( $\geq 0$  and  $\leq 1.25$  km<sup>-1</sup>), Extinction QC 532 (= 0, 1, 2, 167 168 16, or 18), CAD Score ( $\geq$  -100 and  $\leq$  -20), and Extinction Coefficient Uncertainty 532 169  $(\leq 10 \text{ km}^{-1})$ . The Integrated Attenuated Backscatter 532 ( $\leq 0.01 \text{ sr}^{-1}$ ) parameter from 170 the L2 5 km Aerosol Layer (L2 05kmALay) product is also used as a QA metric. A 171 detailed description of these QA checks is also outlined in our most recent CALIOP-172 based study (Toth et al., 2016). Extinction retrievals reported in the CALIOP data 173 products that do not pass the full suite of QA tests are converted to RFVs. To limit the 174 influence of clouds on our analysis (i.e., in order to ensure that the RFV issue is occurring 175 due to layer detection sensitivity and not because of attenuation effects caused by cloud 176 presence), each aerosol profile is cloud-screened using the Atmospheric Volume 177 Description (AVD) parameter. We implement the strictest cloud-screening possible, as 178 profiles are flagged "cloudy" if any of the bins within the CALIOP column are classified 179 as cloud.

180

# 181 2.2 Aqua MODIS

As an integral part of the payloads for NASA's Terra and Aqua satellites, MODIS is a 36 channel spectroradiometer with wavelengths ranging from 0.41 microns to 15 microns. Seven of these channels (0.47-2.13 microns) are used to retrieve aerosol optical properties, such as AOT (e.g., Levy et al., 2013). MODIS L2 aerosol products are reported at a spatial resolution of 10 x 10 km<sup>2</sup> at nadir, with a reported over-ocean expected error of (-0.02, - 10%), (+0.04, + 10%) (Levy et al., 2013). However, uncertainties for individual retrievals may be larger (Shi et al., 2011). Also, thin cirrus

Deleted: \*AOT Deleted: \*AOT 191 contamination may exist in the MODIS aerosol products (e.g., Toth et al., 2013). In this 192 study, the Effective\_Optical\_Depth\_Best\_Ocean (550 nm) parameter in the L2 Collection 193 6 (C6) Aqua MODIS aerosol product (MYD04\_L2; Levy et al., 2013) is utilized. Only 194 those retrievals flagged as "Good" and "Very Good" are considered for analysis, as 195 determined by the Quality\_Assurance\_Ocean parameter within the MYD04\_L2 files.

196

# 197 **2.3 AERONET**

198 Developed for the purpose of furthering aerosol research and validating satellite 199 retrievals, NASA's AERONET program is a federated worldwide system of ground-200 based sun photometers that collect measurements of aerosol optical and radiative 201 properties (Holben et al., 1998). With a reported uncertainty of  $\pm 0.01 - 0.02$  (although 202 this estimate is low in the presence of unscreened cirrus clouds; e.g., Chew et al., 2011), 203 AOTs are derived at several wavelengths ranging from 340 nm to 1640 nm. Due to the 204 lack of retrievals at the CALIOP wavelength, AOTs at 532 nm are computed from interpolation of those derived at the 500 and 675 nm channels using an Angstrom 205 relationship (e.g., Shi et al., 2011; Toth et al., 2013). The highest quality V2.0 206 207 AERONET data (Level 2.0) are used in this study, as these are both cloud-screened and 208 quality-assured (Smirnov et al., 2000). Also, only observations from coastal/island 209 AERONET sites are considered for comparison with over-ocean CALIOP profiles, 210 despite the potential overestimation of CALIOP AOT in coastal regions due to the 211 CALIPSO aerosol typing algorithms (e.g., Kanitz et al., 2014).

212

213 3 Results and Discussion

# 215 3.1 Demonstrating how CALIOP backscatter distribution can render profiles of all 216 RFVs 217

218 To demonstrate the nature of the RFV problem, Fig. 1 shows an example of 219 cloud free all-RFV CALIOP profiles embedded within curtain plots of total attenuated 220 backscatter (TAB; Fig. 1a) and matching vertical feature mask (VFM; Fig. 1b). Both 221 plots were obtained from the CALIPSO Browse Images website [https://www-222 calipso.larc.nasa.gov/products/lidar/browse\_images/production/], and the data were collected from CALIOP during daytime on July 2<sup>nd</sup>, 2010 over the Arctic. The VFM 223 224 shows that the range bins within the white box are classified as either surface or clear air 225 features, and thus the corresponding L2 aerosol extinction coefficient profiles (not 226 shown) are all-RFVs (i.e., the AOT=0 scenario).

227 However, even under pristine conditions, aerosol particles are still present in the atmosphere. For example, the baseline maritime AOT is estimated to be  $0.06 \pm 0.01$ 228 229 (Kaufman et al., 2005; Smirnov et al., 2011). Thus, aerosol particles are likely present 230 and yet undetected for the all-RFV cases shown in Fig. 1. Similar issues can also exist 231 for profiles for which some aerosol is detected. This scenario is represented by the white 232 arrow in the TAB and VFM plots, and the associated L2 aerosol extinction coefficient 233 profile is depicted in Fig. 1c. An aerosol layer is evident from about 1.5 to 2.5 km AMSL, 234 leaving the remainder of the column as RFVs.

To further demonstrate the RFV phenomenon in the CALIOP dataset, we next examine differences in TAB found in profiles where all-RFV were reported and those where some extinction was retrieved. The CALIPSO Lidar Level 1.5 data product (L1.5) is specifically leveraged for this task, as TAB for the all-RFVs class of data is not included in L2 datasets. The L1.5 product is a merging of the L1 and L2 products, cloud240 cleared, screened for non-aerosol features (e.g., surface, subsurface, totally attenuated, invalid, etc.), and available at 20 km (horizontal) and 60 m (vertical) resolutions 241 242 (Vaughan et al., 2011). One month (February 2008) of daytime L1.5 TAB profiles over all global oceans were collocated with CALIOP AOTs derived from the L2 05kmAProf 243 244 product. The data were limited to only those L1.5 averages that contain either four 245 contiguous 5 km L2 all-RFV profiles, or, conversely, four contiguous profiles where 246 extinction was retrieved in each. The selected TAB profiles were then averaged to a 20 247 km resolution for each altitude range (i.e., to obtain over global ocean mean TAB 248 profiles).

249 The results of this analysis are shown in Fig. 2. Profiles of mean TAB over global 250 oceans for February 2008 are shown in Fig. 2a; blue lines show all-RFV profiles and red 251 lines show those where some extinction was retrieved (i.e., non-all-RFVs). For most of 252 the troposphere, little difference is observed between the two profiles (i.e., "clear sky" in 253 the aggregate). However, the profiles begin to deviate below 3 km AMSL, as larger TAB are found for the extinction-retrieved sample (peak TAB is ~0.0031 km<sup>-1</sup> sr<sup>-1</sup>) compared 254 255 to those profiles consisting of all-RFVs (peak TAB value is  $\sim 0.0017$  km<sup>-1</sup> sr<sup>-1</sup>). An 256 additional analysis was conducted (not shown) using data over the Pacific Ocean to check 257 for influences of geographic sampling (i.e., aerosol distribution) on the mean TAB profiles. Both the all-RFV and non-all-RFV mean TAB profiles increase at similar 258 259 magnitudes after implementing this restriction, thus resulting in only a minor difference between the profiles. 260

261 Figure 2c shows a second pair of mean TAB profiles, but now restricted to only

those L2 CALIOP profiles collocated with MODIS AOTs between 0.03 and 0.07 (i.e.,

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265	arbitrarily selected for low aerosol loading scenarios). The collocation method applied	Formatted: Font:Not Italic
266	here is the same as the one used by Toth et al. (2013), where the midpoint of a $10 \times 10$	Moved (insertion) [1]
267	km <sup>2</sup> (at nadir) over-ocean MODIS AOT pixel is required to be within 8 km of the	
268	temporal midpoint of a 5 km L2 CALIOP aerosol profile. Observations outside this	
269	range are not considered. Whereas below, the modal MODIS AOT for passive retrievals	
270	collocated with all-RFV CALIOP profiles is about 0.05, this restriction (i.e., 0.03-0.07	Deleted: (see Sect. 3.3 for rationale)
271	MODIS AOTs) is meant to investigate a more nuanced question. The presence of all-	
272	RFV profiles is the result of several processes that can work either independently or in	
273	tandem. The dominant cause is, as described above, detection failure. RFVs also occur	
274	when the cloud-aerosol discrimination algorithm mistakenly classifies an aerosol layer as	
275	a cloud, and again when the extinction coefficients retrieved for a detected aerosol layer	
276	fail any of the QA metrics (e.g., an out-of-range extinction QC flag). This restriction is	
277	meant to limit the influence of layer misclassifications and occasional QA failures, and in	
278	particular relatively high AOT cases where unusually high TAB could influence the mean	
279	profile. Including such samples would degrade the accuracy of the TAB noise floor	
280	estimate that we will use in subsequent analyses described in Sec. 3.5. Relatively	
281	speaking, though, the profiles in Fig. 2c are fairly similar to those of Fig. 2a. However,	
282	the relative deviation between the two samples now occurs below 2 km AMSL, and the	
283	peak value of TAB for non-all-RFVs lowers to around 0.0025 $\rm km^{-1}~sr^{-1}$ (illustrating the	
284	effect of the MODIS AOT restriction). Also, for context, we include corresponding	
285	profiles of attenuated scattering ratio (TAB/molecular attenuated backscatter) for both	
286	analyses in Figs. 2b and 2d.	

288 The initial point of this comparison is that the mean TAB for all-RFV profiles is, 289 as expected, lower than in those profiles where extinction is retrieved above and within 290 the planetary boundary layer. Thus, the figures represent a simple conceptual model of 291 how profiles consisting of all-RFV cases arise with respect to diffuse aerosol backscatter 292 structure and inherently lower signal-to-noise ratios (SNRs). While there are several 293 possible strategies for mitigating this issue for future global satellite lidar missions 294 (discussed in the concluding remarks), the goal for this initial part of the study is to 295 simply depict how the situation is manifested in the base backscatter product measured 296 by the sensor.

297

## 298 3.2 Frequency of occurrence for L2 CALIOP all-RFV aerosol profiles

299 The next step of the analysis is to determine the frequency of occurrence of all-300 RFV profiles in the daytime CALIOP L2 05kmAProf archive. As these data will be 301 collocated with both MODIS and AERONET data for subsequent analysis, no nighttime 302 data are considered here. Table 1 summarizes the statistics of this analysis. For the 303 2010-2011 period, all-RFV profiles make up about 71% (66%) of all daytime CALIOP 304 L2 05kmAProf profiles globally (global oceans-only). However, these statistics include 305 those profiles for which the CALIOP signal was totally attenuated (e.g., by an opaque cloud layer), thus inhibiting aerosol detection near the surface. For context, the 2010-306 307 2011 occurrence frequencies of CALIOP not detecting the surface are 39.9% (46.1%) globally (global oceans-only). Roughly 30% of the full archive corresponds with cloud-308 309 free conditions (where again, as described in Sec. 2.1, "cloud-free" refers to the 310 implementation of the strictest CALIOP cloud-screening possible where no clouds are classified in the entire profile). Approximately 45% of all cloud-free profiles, and 25%
of cloud-free over ocean profiles, are also all-RFV profiles (~15% and 8%, respectively,
in absolute terms). The over-ocean sample is next considered below, given the relatively
higher fidelity expected in the collocated MODIS AOT data (e.g., Levy et al., 2013).

315 We note that due to the primary CALIOP laser failing in 2009, Table 1 also 316 includes results from a two-year period (2007-2008) before the laser switch to examine 317 any differences in the statistics of the RFV issue between the two lasers. The global frequency of occurrence of all-RFV profiles is consistent for both time periods (i.e., 318 319 70.4% for 2007-2008 and 71.1% for 2010-2011), and thus the remainder of this paper 320 focuses on the 2010-2011 analysis alone. We find no evidence to suggest that laser 321 performance exhibits any significant influence on the occurrence of per-range bin RFVs 322 and all-RFV profiles within the L2 archive.

323 The spatial distribution of daytime over-ocean cloud-free all-RFV profiles is shown in Fig. 3. The percentage of cloud-free CALIOP all-RFV aerosol profiles relative 324 to all cloud-free CALIOP aerosol profiles is computed and presented on a 2° x 5° 325 326 latitude/longitude grid (Fig. 3a). Here we again restrict the analysis to cloud-free scenes 327 to avoid ambiguities in RFV occurrence that are introduced by the presence of clouds. 328 Regions with the largest occurrence frequencies of all-RFV profiles (>75%) include the 329 high latitudes of both the Northern and Southern Hemispheres (NH and SH, respectively). 330 In fact, over snow surfaces, over 80% of CALIOP aerosol profiles are all-RFVs. Over 331 permanent ice (e.g., Greenland), ~99% are all-RFVs. In contrast, the Tropics exhibit the lowest RFV profile occurrence frequencies (<25%). The CALIOP archive contains a 332 333 significant fraction of all-RFV profiles in polar regions, which is an important result with

many ramifications for NASA Earth Observing System science. It is likely that all-RFVs
correlate with both low aerosol loading scenarios and high albedo surfaces (e.g., snow
and sea ice).

Figure 3 also includes the spatial distribution of mean cloud-free CALIOP-337 338 derived AOT (2° x 5° latitude/longitude resolution) without (Fig. 3b) and with (Fig. 3c) 339 all-RFV profiles, demonstrating the quantitative impact of adding all-RFV AOT=0 340 profiles to the relative analysis. As mentioned above, both approaches have been 341 implemented in past studies. Comparison of the plots reveals that including the all-RFV 342 profiles in the average naturally lowers the mean AOT. To determine the areas for which 343 mean AOTs are most impacted by all-RFVs, the ratio of mean AOT without and with all-344 RFV profiles (i.e., the ratio of Fig. 3b to 3c) is shown in Fig. 3d. Little change in mean 345 AOT is found for most of the oceans, with the exception of the high latitudes of each 346 hemisphere. Overall, global ocean cloud-free mean AOT values of  $\sim 0.09$  and  $\sim 0.07$  are 347 found, without and with all-RFV profiles, respectively. Such decrease of mean AOT is 348 expected, as 27% of CALIOP L2 over-ocean cloud-free aerosol profiles are all-RFVs. 349 Also, regions with the largest all-RFV occurrence frequencies (i.e., high latitudes of both 350 the NH and SH) correspond with a greater lowering of mean AOT, compared with those 351 regions (i.e., the Tropics) where small all-RFV occurrence frequencies dominate.

352

# 353 3.3 Collocation of MODIS AOT for over-ocean CALIOP all-RFV cases

By collocating MODIS over-ocean AOT retrievals with CALIOP all-RFV profiles, we can estimate the distribution of AOT when algorithm detection/retrieval performance has been compromised. After collocation was performed (as described in

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**Moved up [1]:** The collocation method applied here is the same as the one used by Toth et al. (2013), where the midpoint of a  $10 \times 10 \text{ km}^2$  (at nadir) over-ocean MODIS AOT pixel is required to be within 8 km of the temporal midpoint of a 5 km L2 CALIOP aerosol profile. Observations outside this range are not considered.

364 Sec. 3.1), the number of all cloud-free CALIOP all-RFV profiles were binned by MODIS 365 AOT in 0.01 increments (as depicted in Fig. 4), and separated into three latitude bands: the NH mid-latitudes (30° to 60° N; Fig. 4a), the Tropics (-30° to 30° N; Fig. 4b), and the 366 SH mid-latitudes (-60° to -30° N; Fig. 4c) where coincident data densities are reasonably 367 368 sufficient. For example, see Fig. 5a for numbers of valid MODIS over-ocean AOT data 369 points available for collocation at 2° x 5° latitude/longitude, based on "Good" or "Very 370 Good" over-ocean L2 MODIS AOT retrievals, relative to all corresponding retrievals. 371 For context, Fig. 5b shows the associated spatial distribution of mean L2 MODIS AOT. 372 We note that this includes only those MODIS points collocated with CALIOP, and thus 373 the AOT distributions shown in Fig. 5b are likely different from distributions derived 374 using the full MODIS data record (e.g., Levy et al., 2013). We also note, for the 375 reference of the reader, that histograms of C6 MODIS AOT (not collocated with 376 CALIOP) are provided in Levy et al. (2013).

377 Modal values of MODIS AOT for all-RFV profiles are found between 0.03 and 378 0.04, with the exception of the 30° to 60° N band for which the greatest number of all-379 RFV profiles coincide with MODIS AOTs between 0.04 and 0.05. Thus, the primary 380 mode of CALIOP RFV profiles is 0.03-0.05 from the perspective of MODIS. 381 Corresponding mean and median MODIS AOTs for collocated CALIOP all-RFV profiles 382 are presented in Table 2, with a mean value of 0.07 for the Tropics and NH mid-latitudes, 383 and 0.05 for the SH mid-latitudes band (global mean of 0.06). Median AOTs are similar, 384 though slightly lower, with a global median of 0.05, reflecting the impact of the tail 385 toward higher AOT in the sample distributions. We expect several modes of algorithm response contributing to these distributions, which are borne out in the CALIOP data: 386

layer detection failures due to sensitivity limits, random noise in the attenuatedbackscatter measurement, and extinction retrieval failures.

389 While a similar distribution is exhibited for each region, the number of total observations for the Tropics is much greater than that of the other two regions. Thus, the 390 391 results of Fig. 4b are more robust, which is primarily due to MODIS AOT data 392 availability and collocation (Fig. 5a). Total MODIS occurrence frequencies are greatest 393 in the Tropics (generally >50%), decreasing poleward. The mid-latitude regions exhibit 394 occurrence frequencies less than 25%, with near-zero frequencies observed in the high 395 latitudes of the NH and SH. We note the low number of valid MODIS AOT retrievals in 396 the high Northern and Southern latitudes, due at least partly to sea ice extent in these 397 regions, presents a limitation for our study. That is, the areas for which all-RFV profiles 398 occur most frequently (Fig. 3a) are the same areas with the least numbers of valid 399 MODIS AOT retrievals. Note that in these regions, even for valid MODIS AOT 400 retrievals, biases due to sub-pixel sea ice contamination may still exist. 401 All-RFV profile occurrence frequencies are computed as a function of MODIS 402 AOT, in order to quantify the amount of CALIOP-derived AOT underestimation at a

409 mid-latitude regions (the red and black curves, respectively, of Fig. 6), respectively) for

given MODIS-based AOT. Achieved by division of corresponding data counts in Fig. 4,

this underestimation (expressed as a percentage) is shown in line plots in Fig. 6. The

same regional sorting and MODIS AOT binning procedures from Fig. 4 are applied. A

similar distribution is found for all three latitude bands, with the 0.01-0.02 MODIS AOT

bin exhibiting the largest underestimation percentage that gradually lowers toward higher

MODIS AOT. CALIOP all-RFV underestimation near 50% is found for the NH and SH

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412	MODIS AOTs between 0.01 to 0.02, and this value increases to about 70% for the	
413	Tropics (the blue curve of Fig. 6). This implies that 70% of all CALIOP aerosol profiles	Deleted: Fig. 6b
414	in this MODIS AOT range are underestimated (i.e., CALIOP reports all-RFV profiles	
415	70% of the time for MODIS AOTs between 0.01 and 0.02).	
416	While the distribution for the Tropics is considered most robust, due to MODIS	Deleted: of Fig. 6b
417	AOT availability in this region, it is important to note that increasingly lower AOTs (i.e.,	
418	below ~0.03) are within the uncertainty range of MODIS AOT retrievals, and thus these	
419	results should be interpreted within the context of this caveat. Also, the relatively low	
420	underestimation percentages corresponding with MODIS AOTs less than 0.02 are	
421	believed to be an error, likely resulting from an artifact in the MODIS AOT	
400	retrievals/products	
422	Terro (ullo) producto.	
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422 423 424	3.4 Collocation of CALIOP all-RFV Profiles with AERONET	Deleted: s
422 423 424 425	<ul> <li>3.4 Collocation of CALIOP all-RFV Profiles with AERONET</li> <li>AERONET data are considered the benchmark for satellite AOT retrievals</li> </ul>	Deleted: s
422 423 424 425 426	<ul> <li>3.4 Collocation of CALIOP all-RFV Profiles with AERONET</li> <li>AERONET data are considered the benchmark for satellite AOT retrievals</li> <li>(Holben et al., 1998). Thus, similar to the over-ocean MODIS analysis above, CALIOP</li> </ul>	Deleted: s
<ul> <li>422</li> <li>423</li> <li>424</li> <li>425</li> <li>426</li> <li>427</li> </ul>	<ul> <li>3.4 Collocation of CALIOP all-RFV Profiles with AERONET</li> <li>AERONET data are considered the benchmark for satellite AOT retrievals</li> <li>(Holben et al., 1998). Thus, similar to the over-ocean MODIS analysis above, CALIOP</li> <li>AOT and all-RFV profiles are examined using collocated AOTs derived from</li> </ul>	Deleted: s
<ul> <li>422</li> <li>423</li> <li>424</li> <li>425</li> <li>426</li> <li>427</li> <li>428</li> </ul>	<ul> <li>3.4 Collocation of CALIOP all-RFV Profiles with AERONET</li> <li>AERONET data are considered the benchmark for satellite AOT retrievals</li> <li>(Holben et al., 1998). Thus, similar to the over-ocean MODIS analysis above, CALIOP</li> <li>AOT and all-RFV profiles are examined using collocated AOTs derived from</li> <li>measurements collected at coastal and island AERONET sites. Ninety-three sites are</li> </ul>	Deleted: s
<ul> <li>422</li> <li>423</li> <li>424</li> <li>425</li> <li>426</li> <li>427</li> <li>428</li> <li>429</li> </ul>	3.4 Collocation of CALIOP all-RFV Profiles with AERONET AERONET data are considered the benchmark for satellite AOT retrievals (Holben et al., 1998). Thus, similar to the over-ocean MODIS analysis above, CALIOP AOT and all-RFV profiles are examined using collocated AOTs derived from measurements collected at coastal and island AERONET sites. Ninety-three sites are used, the locations of which are depicted globally in Fig. 7. Similar to Sec. 3.2, CALIOP	Deleted: s
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<ul> <li>422</li> <li>423</li> <li>424</li> <li>425</li> <li>426</li> <li>427</li> <li>428</li> <li>429</li> <li>430</li> <li>431</li> </ul>	<b>3.4 Collocation of CALIOP all-RFV Profiles with AERONET</b> AERONET data are considered the benchmark for satellite AOT retrievals (Holben et al., 1998). Thus, similar to the over-ocean MODIS analysis above, CALIOP AOT and all-RFV profiles are examined using collocated AOTs derived from measurements collected at coastal and island AERONET sites. Ninety-three sites are used, the locations of which are depicted globally in Fig. 7. Similar to Sec. 3.2, CALIOP L2_05kmAProf data are spatially (within 0.4° latitude/longitude) and temporally (within 30 minutes) collocated with Level 2.0 AERONET data. Note that we include all four	Deleted: s
<ul> <li>422</li> <li>423</li> <li>424</li> <li>425</li> <li>426</li> <li>427</li> <li>428</li> <li>429</li> <li>430</li> <li>431</li> <li>432</li> </ul>	<b>3.4 Collocation of CALIOP all-RFV Profiles with AERONET</b> AERONET data are considered the benchmark for satellite AOT retrievals (Holben et al., 1998). Thus, similar to the over-ocean MODIS analysis above, CALIOP AOT and all-RFV profiles are examined using collocated AOTs derived from measurements collected at coastal and island AERONET sites. Ninety-three sites are used, the locations of which are depicted globally in Fig. 7. Similar to Sec. 3.2, CALIOP L2_05kmAProf data are spatially (within 0.4° latitude/longitude) and temporally (within 30 minutes) collocated with Level 2.0 AERONET data. Note that we include all four years (2007-2008 and 2010-2011) for this analysis, as there are far fewer AERONET data	Deleted: s

433 points available in contrast to MODIS (e.g., Omar et al., 2013).

437	Figure 8 summarizes the results of the CALIOP/AERONET collocation. In a	
438	similar manner as Fig. 4, Fig. 8a is a histogram of the number of cloud-free CALIOP	
439	aerosol profiles (all-RFV profiles and all available) for each 0.01 AERONET AOT bin.	
440	The overall distribution observed here is comparable to that from MODIS (Fig. 4), but	
441	noticeably noisier due to the limited AERONET data sample size. However, peak counts	
442	of all-RFV profiles occur for AERONET AOTs between 0.04 and 0.05, which is roughly	
443	consistent with the MODIS comparisons. The corresponding mean AERONET AOTs of	
444	collocated CALIOP all-RFV profiles are generally higher than those found from MODIS,	
445	with values of 0.1 and 0.09 for the Tropics and NH mid-latitudes, respectively (Table 2),	
446	and a global mean (median) value of $0.08$ (0.07). We note that this analysis may be	
447	influenced by residual cloud contamination of subvisible cirrus in the AERONET dataset	
448	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer	Deleted: for a non-CALIOP-collocated perspective,
448 449	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean	Deleted: for a non-CALIOP-collocated perspective,
448 449 450	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al.	Deleted: for a non-CALIOP-collocated perspective,
448 449 450 451	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al. (2011).	Deleted: for a non-CALIOP-collocated perspective,
448 449 450 451 452	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al. (2011). Fig. 8b shows all-RFV profile occurrence frequencies as a function of AERONET	Deleted: for a non-CALIOP-collocated perspective,
448 449 450 451 452 453	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al. (2011). Fig. 8b shows all-RFV profile occurrence frequencies as a function of AERONET AOT, computed by dividing the respective counts in Fig. 8a. Again, a noisier overall	Deleted: for a non-CALIOP-collocated perspective,
448 449 450 451 452 453 454	<ul> <li>(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer</li> <li>derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean</li> <li>component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al.</li> <li>(2011).</li> <li>Fig. 8b shows all-RFV profile occurrence frequencies as a function of AERONET</li> <li>AOT, computed by dividing the respective counts in Fig. 8a. Again, a noisier overall</li> <li>distribution is found compared with the line plots of Fig. 6. As expected, the 0.01-0.02</li> </ul>	Deleted: for a non-CALIOP-collocated perspective,
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<ul> <li>448</li> <li>449</li> <li>450</li> <li>451</li> <li>452</li> <li>453</li> <li>454</li> <li>455</li> <li>456</li> </ul>	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al. (2011). Fig. 8b shows all-RFV profile occurrence frequencies as a function of AERONET AOT, computed by dividing the respective counts in Fig. 8a. Again, a noisier overall distribution is found compared with the line plots of Fig. 6. As expected, the 0.01-0.02 bin exhibits the largest underestimation percentage. However, while this value is 70% for the MODIS analysis (the blue curve of Fig. 6), it increases to 100% for AERONET,	Deleted: for a non-CALIOP-collocated perspective,
<ul> <li>448</li> <li>449</li> <li>450</li> <li>451</li> <li>452</li> <li>453</li> <li>454</li> <li>455</li> <li>456</li> <li>457</li> </ul>	<ul> <li>(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer</li> <li>derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean</li> <li>component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al.</li> <li>(2011).</li> <li>Fig. 8b shows all-RFV profile occurrence frequencies as a function of AERONET</li> <li>AOT, computed by dividing the respective counts in Fig. 8a. Again, a noisier overall</li> <li>distribution is found compared with the line plots of Fig. 6. As expected, the 0.01-0.02</li> <li>bin exhibits the largest underestimation percentage. However, while this value is 70%</li> <li>for the MODIS analysis (the blue curve of Fig. 6), it increases to 100% for AERONET,</li> <li>and we again conclude that an artifact is likely present in the MODIS retrievals for very</li> </ul>	Deleted: for a non-CALIOP-collocated perspective,
<ul> <li>448</li> <li>449</li> <li>450</li> <li>451</li> <li>452</li> <li>453</li> <li>454</li> <li>455</li> <li>456</li> <li>457</li> <li>458</li> </ul>	(e.g., Chew et al., 2011; Huang et al., 2012). We note that histograms of sun photometer derived AOT from Maritime Aerosol Network (MAN) observations (i.e., over-ocean component of AERONET; not collocated with CALIOP data) are shown in Smirnov et al. (2011). Fig. 8b shows all-RFV profile occurrence frequencies as a function of AERONET AOT, computed by dividing the respective counts in Fig. 8a. Again, a noisier overall distribution is found compared with the line plots of Fig. 6. As expected, the 0.01-0.02 bin exhibits the largest underestimation percentage. However, while this value is 70% for the MODIS analysis (the blue curve of Fig. 6), it increases to 100% for AERONET, and we again conclude that an artifact is likely present in the MODIS retrievals for very low aerosol loading cases. While the sample size is small, in the 4-year data set	Deleted: for a non-CALIOP-collocated perspective,

- 462 collocated CALIOP aerosol profiles contained only RFVs.
- 463

#### 464 **3.5 Reconciling CALIOP AOT Underestimation**

465 In this part of the study, we describe a proof-of-concept analysis that uses one-466 month of data with the same spatio-temporal domain and conditions introduced in Sec. 467 3.1 to estimate the nominal underestimation of CALIOP AOT due to RFVs in otherwise 468 high-fidelity L2 retrievals (i.e., those where extinction is derived and the profile passes all 469 QA/QC tests). This is achieved by retrieving extinction profiles from the mean global TAB profiles previously constructed from all-RFV profiles (i.e., as presented in Fig. 2). 470 471 Characterizing these profiles, including those derived for all corresponding/collocated 472 MODIS AOT (Fig. 2a, with an average MODIS AOT of 0.067) and MODIS AOT 473 between 0.03 and 0.07 (Fig. 2c, with an average MODIS AOT of 0.045) to suppress the 474 influence of random algorithm failure events at relatively high AOT, as TAB "noise 475 floors", we then replace RFV bins with corresponding extinction and calculate column-476 integrated AOT. The premise here assumes that the distribution of aerosol depicted in 477 the TAB noise floors is constant globally. This is highly uncertain, and we strongly 478 caution that the purpose is to provide an initial demonstration of a practical way to 479 correct RFVs in the CALIOP archive.

The aerosol extinction profiles for all-RFVs are derived in two steps. First, using
an assumed lidar ratio of 29 sr (standard deviation of 10 sr; derived from constrained
lidar ratios over ocean and represents background aerosols for the entire atmospheric
column; Kim et al., 2017), an unconstrained extinction solution is generated from 20 km

484 to the top of the surface-attached layer (3.5 km). In this step, the molecular and aerosol 485 attenuation in the measured backscatter is accounted for at each range bin (from a top-486 down approach) by taking into account the overlying molecular and aerosol loading. The aerosol backscatter is then calculated by subtracting the unattenuated molecular 487 488 backscatter from the newly derived aerosol-and-molecular-attenuation-corrected 489 backscatter, from which the aerosol extinction is derived by multiplication of the lidar 490 ratio. The top of surface-attached layer is determined by inspection of the ratio between 491 the measured backscatter and the modeled molecular attenuated backscatter, as provided in the CALIPSO L1.5 product. Integrating this extinction profile provides an estimate of 492 the AOT overlying the surface-attached layer (AOT<sub>upper</sub>). The derived AOT<sub>upper</sub> values 493 494 are ~0.015 and ~0.01 for the total all-RFV sample and AOT-limited sample, respectively. 495 These values are not surprising, as they are in agreement with AERONET measurements 496 obtained at the Mauna Loa site (elevation of ~3. 5 km AMSL; Alfaro-Contreras et al., 497 2016).

498 Next, a constrained extinction solution and optimized estimate of the lidar ratio 499 are generated from 3.5 km to the surface using the AOT of this layer (i.e., column AOT -500 AOT<sub>upper</sub>). This step is similar to the above-mentioned approach, except now an iterative 501 process is implemented to derive a lidar ratio for the layer. Resulting surface-attached 502 layer lidar ratios are 43 sr and 30 sr, for the first and second case respectively, with the latter value comparing reasonably well with the coastal marine lidar ratio of ~28 derived 503 from AERONET analyses (Sayer et al., 2012). However, the lidar ratio solved for the 504 505 all-RFV sample case is higher than that typical of marine aerosols (i.e., ~26; Dawson et 506 al., 2015), which may be a result of uncertainties in both MODIS and CALIOP datasets.

508 of -0.02 - 0.04 (Levy et al., 2013). These lidar ratios are also likely biased high due to biases in the daytime CALIOP V3 calibration scheme: the V3 daytime calibration 509 510 coefficients are typically 10% to as much as 30% higher than their V4 counterparts, 511 depending on location and season (Getzewich et al., 2016). Additionally, some all-RFV 512 profiles may include non-marine aerosols, which would further contribute to the high 513 biases in the retrieved lidar ratios.

For example, the uncertainty of the lower end of MODIS AOT retrievals is on the order

507

514 Despite these caveats, the resultant all-RFV extinction profiles are shown in Fig. 9, 515 with values peaking near the surface and decreasing exponentially with height. These are 516 thus considered the corresponding/approximated CALIOP extinction-based noise floors. 517 Next, for those cloud-free, over-ocean, L2 05kmAProf CALIOP profiles from the same 518 month (February 2008), RFV bins for profiles where some measure of extinction has 519 been observed and passed QA/QC were replaced with the corresponding extinction noise-520 floor values solved for the two TAB samples. Profiles were then reintegrated to yield 521 RFV-corrected AOTs.

522 The results of this exercise are summarized in Table 3. The first result, 523 representing the inclusion of all-RFV profiles as is within bulk global samples (i.e., 524 adding cases of AOT=0 to a given sample) shows a difference of 0.033 between collocated CALIOP and MODIS AOT. The noise floor correction applied to both all-525 RFV profiles and those where some extinction was solved yields AOT differences (i.e., 526 MODIS-CALIOP) of -0.009 and 0.006 depending on the correction sample, which is an 527 528 improvement (~20% in absolute value) in the agreement of CALIOP and MODIS AOTs. 529 If profiles with nominal extinction are not corrected and all-RFV profiles are ignored, a

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mean AOT difference of 0.025 is found with MODIS. Applying the noise-floor 532 533 corrections for this scenario results in AOT differences of -0.013 and 0.001, or a  $\sim$ 10-534 20% improvement (in absolute value) in the disparity in mean AOT between the two sensors. Lastly, we emphasize to the reader that this section describes only an initial 535 536 attempt to resolve the RFV issue, and can likely be improved in future studies. For 537 example, the noise floor extinction profile is derived using data from global oceans, while 538 a regional dependency is possible. Also, longer spatial and temporal averages of 539 CALIOP data would likely increase the SNRs and reduce the frequency of occurrence of 540 the RFV issue.

541

# 542 3.6 Case study: Nighttime CALIOP all-RFV profile occurrence frequencies

543 The analyses in this paper use daytime CALIOP data to allow for comparison 544 with passively-sensed aerosol observations from MODIS and AERONET. However, for 545 context, in this section we conduct a case study for a two-month (January and February 546 2008) period to investigate the occurrence frequencies of CALIOP all-RFV profiles 547 during nighttime conditions. The same CALIOP products and QA procedures as 548 described earlier are used here, and Table 4 summarizes the results of this analysis. 549 During nighttime, about half of all global CALIOP aerosol profiles for this period are all-RFVs, but this statistic decreases to about 22% when restricted to cloud-free conditions. 550 551 This percentage lowers even further for over-ocean profiles. Depending on the analysis, absolute decreases between daytime and nighttime all-RFV occurrence frequencies range 552 553 from  $\sim 8\%$  to  $\sim 25\%$ . These findings are expected, as the lack of solar background signal 554 during nighttime allows for an increased SNR and improves the ability of the CALIOP Deleted:

556 algorithms to detect aerosol layers.

557

#### 558 3.7 Anticipating Version 4 CALIOP Aerosol Products

559 Version 4 (V4) CALIOP L2 aerosol products were publicly released in 560 November 2016. A case study was thus performed to assess changes in RFV impacts 561 using these new products, again considering cloud-free over-global-ocean observations 562 during daytime conditions. Whereas the broader point of the paper is a conceptualization of the lower-threshold sensitivity of CALIOP to aerosol presence, and the global 563 564 distribution and impact on overall archive availability, this analysis is included for general consistency. Specifically, V4 data feature improved calibrations of Level 1 (L1) 565 566 backscatter, as well as improved cloud-aerosol discrimination and surface detection, that 567 may increase the detection sensitivity of diffuse aerosol layers that are reflected in L2 568 aerosol extinction retrievals. This may then result in a possible decrease in the 569 occurrence of all-RFV profiles overall.

570 A two-month V4 (January and February of 2008) analysis using QA aerosol 571 profile data (L2 05kmAPro-Standard-V4-10) reveals a 4% relative decrease (1% 572 absolute decrease) in global all-RFV profile occurrence frequencies between V3 and V4. 573 Without QA screening (Sec. 2.1), a 15% relative decrease (2% absolute decrease) is 574 found in the occurrence frequency of all-RFV profiles between versions. A supplemental analysis was also conducted, through the use of the CALIOP aerosol layer product 575 (L2 05kmALay-Standard-V4-10) with alternative cloud screening (i.e., cloud optical 576 577 depth = 0 instead of the AVD parameter), the results of which are consistent with those 578 from the L2 05kmAPro-Standard-V4-10 test. Though this is an initial look at this Deleted: , and Deleted: thus important new dataset, it appears that improvements in instrument calibration are likely having some positive influence on retrieval sensitivity, though the broader impact of all-RFV profiles as a limiting factor on the breadth of the CALIOP archive, particularly at the poles, mostly remains.

585

# 586 4 Conclusions

587 Since June 2006, the NASA Cloud-Aerosol Lidar with Orthogonal Polarization 588 (CALIOP) instrument has provided a unique global space-borne view of aerosol vertical 589 distribution in Earth's atmosphere. As indicated by this study, a significant portion of 590 Level 2 (L2) CALIOP 532 nm aerosol profiles consist of retrieval fill values (RFVs) 591 throughout the entire range-resolved column (i.e., all-RFVs), overwhelmingly the result 592 of instrument sensitivity and algorithm layer detection limits. The relevant impact of the 593 all-RFV profile is a subsequent column-integrated aerosol optical thickness (AOT) equal 594 to zero.

595 Using four years (2007-2008 and 2010-2011) of daytime CALIOP Version 3 L2 aerosol products, the frequency of occurrence of all-RFV profiles within the CALIOP 596 597 archive is quantified. L2 retrieval underestimation and lower detectability limits of 598 CALIOP-derived AOT are assessed using collocated L2 aerosol retrievals from over-599 ocean Aqua Moderate Resolution Imaging Spectroradiometer (MODIS) and 600 coastal/island Aerosol Robotic Network (AERONET) measurements. The results are partitioned into three latitude bands: Northern Hemisphere mid-latitudes (30° to 60° N), 601 Tropics (-30° to 30° N), and Southern Hemisphere mid-latitudes (-60° to -30° N). The 602 603 primary findings of this study are:

604	1.	Analysis of CALIOP Level 1.5 attenuated backscatter data reveals that all-RFV
605		profiles are primarily the result of diffuse aerosol layers with inherently lower
606		signal-to-noise ratios (SNRs) that are below CALIOP layer detection limits.
607	2.	All-RFV profiles make up 71% (66%) of all daytime CALIOP L2 aerosol profiles
608		globally (global oceans-only), although this includes completely attenuated
609		columns. For cloud-free CALIOP L2 aerosol profiles, 45% (27%) globally
610		(global oceans-only) are all-RFV profiles. The largest relative all-RFV profile
611		occurrence frequencies (>75%) are found in the high latitudes of both
612		hemispheres, and are smallest (<25%) in the Tropics. The results of this study
613		indicate that there is a significant daytime observational gap in CALIOP aerosol
614		products near the poles, which is a critically important finding for community
615		awareness.
616	3.	The primary mode of CALIOP all-RFV profiles corresponds with MODIS AOTs
617		of 0.03-0.05, which is largely consistent with an AERONET-based analysis.
618		Also we found that a small fraction of AERONET data have AOTs lower than

nd that a small fraction of AERONET have AOTs lower 618 Also, data 619 0.02, of which all collocated CALIOP L2 profiles are all-RFVs. This finding is 620 consistent with the lowest detectable CALIOP aerosol optical depth range of 0.02-621 0.04, as hypothesized by Kacenelenbogen et al. (2011). Note that this conclusion 622 hints that CALIOP may not detect very thin aerosol layers (i.e., AOTs < 0.05), 623 which account for ~10-20% of the AOT spectrum and are of climatological 624 importance (e.g., Smirnov et al., 2011; Levy et al., 2013). Also, these CALIOP-625 undetected thin aerosol layers are important for various applications, ranging from 626 data assimilation to aerosol indirect effects.

4. As a preliminary study, aerosol extinction coefficient values for two distinct
CALIOP all-RFV profile samples are derived using an inversion algorithm
applied to corresponding attenuated backscatter data, and a collection of RFVcorrected mean CALIOP AOTs are estimated for a one-month case study. The
mean over-ocean CALIOP AOTs increase 10-20% (in absolute value) after
correction, with a closer match to collocated Aqua MODIS mean over-ocean
AOT.

634 5. A small decrease in all-RFV profile occurrence is found from Version 4 CALIOP
635 data, which are undergoing widespread release at the time of this writing. Still,
636 the larger-scale impact of all-RFV profiles remains.

637 This research demonstrates that all-RFV profiles exert a significant influence on 638 the L2 CALIOP AOT archive, as these data compose nearly half of global cloud-free 639 CALIOP aerosol points. Disagreements exist in the literature on the manner for which to 640 handle all-RFV profiles when generating Level 3 AOT statistics. Some studies have set 641 the integrated AOTs of all-RFV profiles to zero, for instance, and included them. 642 However, analyses with passive-based sensors presented in this study reveal these AOTs 643 are most certainly non-zero (global mean values of 0.06 for MODIS and 0.08 for 644 AERONET). These findings are not surprising, as this is the baseline AOT range expected under clean maritime conditions (Kaufman et al., 2001; 2005). 645

646 This research also shows that CALIOP RFVs caused by lower backscatter 647 threshold sensitivities to highly diffuse aerosols, contribute significantly to the 648 discrepancy between CALIOP AOT and those derived from passive sensors like MODIS. 649 Previous studies have mostly attributed this offset to selection of the CALIOP lidar ratio 650 (extinction-to-backscatter ratio) or errors in passive aerosol retrievals. Multi-spectral
651 lidar measurements can begin to close the gap, but will experience SNR issues of their
652 own.

By characterizing lower detection limits of CALIOP-derived extinction and AOT, the potential exists for innovations in instrumentation design and algorithm development of future lidar missions, such as those affiliated with the NASA Aerosol-Clouds-Ecosystems (ACE) mission or the signal processing effort of Mariais et al. (2016). Specifically, increasing the intensity of the lidar signal or implementing larger spatial averaging schemes may help to lower the occurrence frequency of all-RFV profiles and relative RFV occurrence per range bin in L2 products. Questions, however, arise in terms of developing datasets with sufficient spatial and temporal resolution versus needs for optimal data densities, and which is more significant for a given project. Regardless of the potential solution, science teams of current and future lidar systems should carefully consider the existence of RFVs in project datasets. 

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# 883 Figure and Table Captions

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885 Figure 1: For data collected during daytime on July 2<sup>nd</sup>, 2010 over the Arctic, browse 886 image curtain plots of CALIPSO (a) 532 nm total attenuated backscatter (km<sup>-1</sup> sr<sup>-1</sup>) and 887 (b) corresponding vertical feature mask (VFM). The white box represents an example 888 segment of the granule for which range bins in the associated Level 2 (L2) aerosol 889 extinction coefficient profile are all retrieval fill values (RFVs), as the VFM classified 890 these bins as either surface (green) or clear air (blue) features. The white arrow indicates 891 a column in which some aerosol has been detected (orange), and the resultant L2 aerosol 892 extinction profile for this column is shown in (c).

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Figure 2: For February 2008, mean profiles of (a, c) Level 1.5 total attenuated backscatter (TAB) and (b, d) attenuated scattering ratio (TAB/molecular attenuated backscatter) over global oceans, corresponding to Level 2 all-RFV (in blue) and non-all-RFV (AOT > 0; in red) profiles. The left column is from an analysis of all cloud-free CALIOP points over global oceans and the right column represents only those collocated with MODIS AOTs between 0.03 and 0.07.

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901 Figure 3: For 2010-2011, (a) the frequency of occurrence (%) of cloud-free CALIOP
902 profiles at 2° x 5° latitude/longitude grid spacing. Also shown are the corresponding
903 cloud-free mean CALIOP column AOTs (b) without and (c) with all-RFV profiles, and
904 (d) the ratio of (b) to (c).

906	Figure 4: For 2010-2011, histograms of all over-ocean cloud-free CALIOP profiles (in
907	green) and all-RFV profiles (in purple) as a function of collocated Aqua MODIS AOT
908	(0.01 bins), for (a) $30^{\circ}$ to $60^{\circ}$ N, (b) $-30^{\circ}$ to $30^{\circ}$ N, and (c) $-60^{\circ}$ to $-30^{\circ}$ N.
909	
910	Figure 5: For 2010-2011, (a) frequency of occurrence (%) of valid ("Good" or "Very
911	Good") over-ocean Level 2 (L2) MODIS AOT retrievals, relative to all over-ocean L2
912	MODIS AOT retrievals, for every 2° x 5° latitude/longitude grid box. Also shown is (b)

913 the corresponding spatial distribution of mean L2 MODIS AOT for the same time period.

914 This analysis includes only those MODIS points collocated with CALIOP.

916 Figure 6: 2010-2011 frequency of occurrence (%) of over-ocean cloud-free CALIOP all-

917 RFV profiles, relative to all cloud-free CALIOP profiles, as a function of collocated

918 Aqua MODIS AOT (0.01 bins), for 30° to 60° N (in red), -30° to 30° N (in blue), and -

919 <u>60° to -30° N (in black).</u>

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Figure 7: Map of the ninety-three coastal/island AERONET sites with Level 2.0 data, for the 2007-2008 and 2010-2011 periods, used for collocation with over-ocean CALIOP aerosol observations. **Deleted:** Figure 6: 2010-2011 frequency of occurrence (%) of over-ocean cloud-free CALIOP all-RFV profiles, relative to all cloud-free CALIOP profiles, as a function of collocated Aqua MODIS AOT (0.01 bins), for (a) 30° to 60° N, (b) -30° to 30° N, and (c) -60° to -30° N.  $\cdot$ 

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Figure 8: For the 2007-2008 and 2010-2011 periods, (a) histograms of all cloud-free CALIOP profiles (in green) and all-RFV profiles (in purple), and (b) corresponding frequency of occurrence (%) of cloud-free CALIOP all-RFV profiles, relative to all cloud-free CALIOP profiles, both as a function of collocated coastal/island AERONET AOT (0.01 bins).

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936	Figure 9: For February 2008 over cloud-free global oceans, the all-RFV aerosol		
937	extinction coefficient profiles derived from the inversion algorithm. The black curve		
938	represents all cloud-free CALIOP profiles over global oceans, while the green curve is		
939	from an analysis restricted to only those CALIOP points collocated with MODIS AOTs		
940	between 0.03 and 0.07.		
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942	Table 1: Statistical summary of the results for this study, for the 2007-2008 and 2010-		
943	2011 periods, both globally and for global oceans only. The values in bold and		
944	parentheses represent the percentages of each category relative to the entire CALIOP		
945	aerosol profile archive for each respective period.		
946			
947	Table 2: Mean, median, and standard deviation of AOTs derived from Aqua MODIS	$\leq$	Deleted: and
948	(2010-2011) and AERONET (2007-2008; 2010-2011), both independently collocated		Deleted: values
949	with CALIOP all-RFV profiles.		
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952	deviation of collocated CALIOP and MODIS AOTs for various scenarios related to the		
953	treatment of non-all-RFV and all-RFV CALIOP aerosol profiles. For those scenarios that		
954	involve correction, [1] refers to analyses including all cloud-free CALIOP profiles over		
955	global oceans, while [2] refers to analyses restricted to CALIOP points collocated with		
956	MODIS AOTs between 0.03 and 0.07. The corresponding aerosol extinction profiles		
957	used for RFV correction are shown in Fig. 9. Key results are highlighted in yellow.		

Table 4: All-RFV CALIOP occurrence frequencies for two months (January and February 2008) from various analyses using daytime and nighttime data, as well as their corresponding absolute differences.



Figure 1: For data collected during daytime on July  $2^{nd}$ , 2010 over the Arctic, browse image curtain plots of CALIPSO (a) 532 nm total attenuated backscatter (km<sup>-1</sup> sr<sup>-1</sup>) and (b) corresponding vertical feature mask (VFM). The white box represents an example segment of the granule for which range bins in the associated Level 2 (L2) aerosol extinction coefficient profile are all retrieval fill values (RFVs), as the VFM classified these bins as either surface (green) or clear air (blue) features. The white arrow indicates a column in which some aerosol has been detected (orange), and the resultant L2 aerosol extinction profile for this column is shown in (c).

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Figure 2: For February 2008, mean profiles of (a, c) Level 1.5 total attenuated backscatter (TAB) and (b, d) attenuated scattering ratio (TAB/molecular attenuated backscatter) over global oceans, corresponding to Level 2 all-RFV (in blue) and non-all-RFV (AOT > 0; in red) profiles. The left column is from an analysis of all cloud-free CALIOP points over global oceans and the right column represents only those collocated with MODIS AOTs between 0.03 and 0.07.



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Figure 5: For 2010-2011, (a) frequency of occurrence (%) of valid ("Good" or "Very Good") over-ocean Level 2 (L2) MODIS AOT retrievals, relative to all over-ocean L2 MODIS AOT retrievals, for every 2° x 5° latitude/longitude grid box. Also shown is (b) the corresponding spatial distribution of mean L2 MODIS AOT for the same time period. This analysis includes only those MODIS points collocated with CALIOP.



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Figure 8: For the 2007-2008 and 2010-2011 periods, (a) histograms of all cloud-free CALIOP profiles (in green) and all-RFV profiles (in purple), and (b) corresponding frequency of occurrence (%) of cloud-free CALIOP all-RFV profiles, relative to all cloud-free CALIOP profiles, both as a function of collocated coastal/island AERONET AOT (0.01 bins).



Figure 9: For February 2008 over cloud-free global oceans, the all-RFV aerosol extinction coefficient profiles derived from the inversion algorithm. The black curve represents all cloud-free CALIOP profiles over global oceans, while the green curve is from an analysis restricted to only those CALIOP points collocated with MODIS AOTs between 0.03 and 0.07.

Globe **Global Oceans** Number of 5 km CALIOP Profiles 2007-2008 2010-2011 2007-2008 2010-2011 41,929,328 41,188,208 27,198,000 Total 27,742,947 All-RFV 29,503,781 (70.4%) 29,297,919 (**71.1%**) 18,190,188 (65.6%) 18,026,930 (66.3%) 7,812,682 (**28.7%**) Cloud-free 8,006,719 (28.9%) 13,317,918 (**31.8%**) 13,190,530 (**32.0%**) Cloud-free & all-RFV 5,764,098 (13.7%) 5,899,221 (**14.3%**) 2,089,865 (7.5%) 2,101,155 (7.7%) Cloud-free, all-RFV, & MODIS 791,570 (1.9%) 814,514 (**2.0%**) 781,983 (**2.8%**) 803,546 (**3.0%**) AOT≥0

Table 1: Statistical summary of the results for this study, for the 2007-2008 and 2010-2011 periods, both globally and for global oceans only. The values in bold and parentheses represent the percentages of each category relative to the entire CALIOP aerosol profile archive for each respective period.

Decien	MODIS			AERONET			
Region	Mean	Median	Standard Deviation	Mean	Median	Standard Deviation	
90°S to 60°S	0.05	0.04	0.10	-	-	-	
60°S to 30°S	0.05	0.04	0.11	0.04	0.04	0.01	
30°S to 30°N	0.07	0.06	0.11	0.10	0.10	0.19	
30°N to 60°N	0.07	0.06	0.13	0.09	0.08	0.07	
60°N to 90°N	0.07	0.06	0.17	0.05	0.04	0.04	
Globe	0.06	0.05	0.12	0.08	0.07	0.11	

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	Region	Me
	90°S to 60°S	0.
	60°S to 30°S	0.
	30°S to 30°N	0.
	30°N to 60°N	0.
	60°N to 90°N	0.
	Globe	0.
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Table 2: Mean, median, and standard deviation of AOTs derived from Aqua MODIS (2010-2011) and AERONET (2007-2008; 2010-2011), both independently collocated with CALIOP all-RFV profiles.

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Scenario						CALIOP AOT		DIS AOT		
Corrected non-All-RFVs?	All-RFVs set to zero?	All-RFVs ignored?	All-RFVs corrected?	Correction Subset	Mean	Standard Deviation	Mean	Standard Deviation	ΔΑΟΤ (MODIS-CALIOP)	
	1				0.084	0.113	0.117	0.133	0.033	
1			1	[1]	0.126	0.107	0.117	0.133	-0.009	
1			1	[2]	0.111	0.109	0.117	0.133	0.006	
		1			0.098	0.116	0.123	0.123	0.025	
1		1		[1]	0.136	0.112	0.123	0.123	-0.013	
1		1		[2]	0.122	0.114	0.123	0.123	0.001	

Scenario
Uncorrected non-all-RFVs; all-R
Corrected non-all-RFVs; correct
Corrected non-all-RFVs; correct
Uncorrected non-all-RFVs; all-R
Corrected non-all-RFVs; all-RFV
Corrected non-all-RFVs; all-RFV

Table 3: For February 2008 over cloud-free global oceans, <u>the mean and standard deviation of collocated CALIOP and</u> MODIS AOTs for various scenarios related to the treatment of non-all-RFV and all-RFV CALIOP aerosol profiles. For those scenarios that involve correction, [1] refers to analyses including all cloud-free CALIOP profiles over global oceans, while [2] refers to analyses restricted to CALIOP points collocated with MODIS AOTs between 0.03 and 0.07. The corresponding aerosol extinction profiles used for RFV correction are shown in Fig. 9. Key results are highlighted in yellow. Deleted: values

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Analysis	All Points	Cloud-free		
Dautima	Globe	70.7%	46.7%	
Daytime	Global Oceans	63.4%	21.8%	
Nightting	Globe	53.5%	22.0%	
Nightume	Global Oceans	52.2%	14.0%	
Nighttime Deutime	Globe	-17.2%	-24.7%	
Nighttime - Daytime	Global Oceans	-11.2%	-7.8%	

Table 4: All-RFV CALIOP occurrence frequencies for two months (January and February 2008) from various analyses using daytime and nighttime data, as well as their corresponding absolute differences.

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Figure 7: Map of the ninety-three coastal/island AERONET sites with Level 2.0 data, for the 2007-2008 and 2010-2011 periods, used for collocation with over-ocean CALIOP aerosol observations.

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