

1 REVIEW REPORT

2

3 Review of amt-2017-401 by RC1

4

5 By Viswanathan Bringi, Merhala Thurai and Darrel Baumgardner

6

7 Manuscript Title – Raindrop Fall Velocities from an Optical Array Probe and 2D-Video Disdrometer

8

9

10 GENERAL COMMENTS

11

12 In the manuscript the Authors analyzed three precipitation events occurred in USA (Colorado and
13 Alabama) that differs for the climatology of the colocations and for the wind conditions. The aim of the
14 study is to evaluate the wind influence on the raindrops terminal fall speed measured by two different
15 type of devices, namely the Meteorological Particle Spectrometer (MPS) and the 2D video disdrometer
16 (2DVD). The manuscript is well organized however I think that the aim of the study and in particular its
17 practical applications should be specified in the Introduction section. As stated also by the Authors more
18 case studies should be added or at least the analysis should be extended to larger drops (see for
19 example comment 6 and 9 below). Furthermore section 2 need to be enlarged with information
20 regarding the data processing (see comment 3 below) and more analytical comparison should be done
21 to confirm the consistency of fall speed measurements from the two devices (see comment 1 below).
22 Finally I have some specific comments, that are shown below.

23 ***We appreciate the reviewer's comments and our response is given in italics while the***
24 ***modifications to the manuscript are highlighted as below.***

25 *General: The first sentence in the Introduction gives three applications with three pertinent*
26 *references. We believe these are also "practical" references. For example, the modeling of*
27 *collisional processes is in part based on assigning a unique terminal fall speed for a given mass*
28 *of raindrop. Same is true for retrieval of DSD from vertical pointing profilers. We have also*
29 *added in a new reference to Yu et al. (2016) in AMT who also suggest that ambient flow and*
30 *turbulence may play a role in modifying drop fall speeds.*

31 *We have added the following sentences to clarify this point in the Introduction:*

32 “In these and other applications it is nearly universally accepted that there is a unique
33 fall speed ascribed to drops of a given mass or diameter and that it equals the terminal
34 speed with adjustment for pressure (e.g., Beard 1976). “

35 “ Thus, there is some evidence that rain drops may not fall at their terminal velocity
36 except under calm conditions and that the concept of a fall speed distribution for a drop
37 of given mass (or, diameter) might need to be considered which is the topic of this
38 paper. The implications are rather profound especially for numerical modeling of
39 collision-coalescence and breakup processes which are important for shaping the drop
40 size distribution.”

41

42 **Our response to SPECIFIC COMMENTS.**

43 SPECIFIC COMMENTS

44 1. Line 24-25: in the manuscript the consistency of fall speed measurements from the two devices
45 is provided only qualitatively (i.e. “excellent visual agreement”) some quantitative results should
46 be provided for all the diameters in the overlapping region.

47

48 *We have added a new Table 1 that compares the mean and standard deviation from MPS and*
49 *2DVD for both sites in the overlap diameter range 0.7 to 2 mm.*

50

51 2. Line 109: please clarify which are the “other factors” that gives the threshold of 0.7 mm for the
52 drop diameter. The 2DVD is able to measure drops with $D < 0.7$ mm. Usually the minimum
53 detectable diameter for 2DVD is considered 0.2 mm or 0.3 mm. In this case the overlapping
54 between the two instruments can be enlarged. Please provide a clarification of this threshold or
55 consider the option of enlarging the overlapping region.

56

57 *In line 107 we have given three references that discuss 2DVD accuracy. In particular,*
58 *Bernauer et al. (2015) carefully document the accuracy of sizing and fall speed for solid*
59 *precipitation. The main problem for $D < 0.7$ mm is that the drop image in Camera A has*
60 *to be “matched” to the corresponding drop image in Camera B which is difficult given the*
61 *resolution of 170 microns. This has been mentioned in second para of Section 3. But we*
62 *have added the “mis-match” problem in original line 109.*

63

64 “Considering the horizontal pixel resolution of 170 μm and other factors (such as
65 “mis-matched” drops), the effective sizing range is $D > 0.7$ mm. To clarify the
66 “mis-matched” drop problem: it is very difficult to match a drop detected in the top
67 light-beam plane of the 2DVD to the corresponding drop in the bottom plane for
68 tiny drops resulting in erroneous fall speeds.”

69

70 *The smallest calibration spheres that are provided by the manufacturer is 0.5 mm and*
71 *these are extremely difficult to drop in the sensor area and to collect them below.*
72 *According to Bernauer et al. accurate sizing with errors <5% is only possible for $D > 1$*
73 *mm based on solid (snow) precipitation. In our experience for rain drops which are*
74 *smooth shaped the threshold is lower at 0.7 mm which is what we have quoted in our*
75 *paper. While the mid-point of the first “bin” of the 2DVD sizing is 0.25 mm, the accuracy*
76 *is very questionable as calibration is not possible. Hence we cannot enlarge the*
77 *overlapping region. The whole point of our paper is to use high resolution MPS and*
78 *lower resolution 2DVD to cover the entire size range with good-to-excellent accuracy.*
79

- 80 3. Line 114-115: As reported in numerous papers in the literature, the 2DVD measures a number of
81 spurious drops that can be usually removed from the data using proper filter criterion, such as
82 the one based on the relation between measured and theoretical fall velocities. Please note that
83 in my experience most of the spurious drops have small diameters ($D < 2$ mm) and therefore are
84 within the range of diameters analyzed in this study. Did the Authors use any kind of criterion to
85 filter out these drops? If yes which is the impact of the filtering on the results. If not, how can
86 the Authors be sure that those drops are real drops and not spurious ones? I think that the
87 Authors should clarify this point in the manuscript because it is crucial for the validity of the
88 results obtained in the study.

89
90 *We clearly state in original lines 114-115 that ...” The only fall velocity threshold used for*
91 *the 2DVD is the lower limit set at 0.5 m s^{-1} in accordance with the manufacturer*
92 *guidelines for rain measurements.” We do not use any velocity filter in our analysis as*
93 *doing so would be counter to the detection of sub- or super terminal fall speeds. Please*
94 *see our response above to point (2) specifically the threshold of 0.7 mm which we use to*
95 *eliminate mis-matched or “spurious” drops. Our confidence of sizing and fall speed*
96 *measurement using 2DVD for $D > 0.7$ mm is reinforced by the agreement with MPS in the*
97 *overlap region (please see new Table 1). Finally, the use of a DFIR wind shield appears*
98 *to have reduced the occurrence of “spurious” drops.*
99

- 100
101 4. Line 119: How do the Authors identify the different rain types?

102 *The identification of rain types was based on CSU-CHILL radar data as described in the*
103 *quoted reference (Thurai et al. 2017).*

- 104 5. Line 131: I suggest to change the word “excellent” with the word “good”. The MPS
105 underestimates the fall velocities for $0.7 \text{ mm} < D < 1 \text{ mm}$ with respect to 2DVD, while the 2DVD
106 overestimates the fall velocities for $1 \text{ mm} < D < 2 \text{ mm}$ with respect to the Gunn and Kinzer fit.
107 Furthermore a more quantitative agreement should be performed.

108
109 *Please see new Table 1 for a quantitative comparison between MPS and 2DVD in the*
110 *overlap region. The agreement between the two instruments is, in our opinion, excellent*

111 for both sites given the quoted accuracies in fall speed measurement for both
112 instruments.

113
114 We have replaced the fit to Gunn-Kinzer data with the 9th order polynomial fit given in
115 Foote and du Toit (1969) as opposed to using the exponential fit of Atlas et al. (1973) in
116 both figs. 1a and 3a. The latter fit is not as accurate for the small drop end (e.g., it does
117 not pass through the origin). We have added in Section 2.1:

118
119 “Also shown is the (Foote and du Toit 1969) (henceforth FT fit) to the terminal fall
120 speed measurements of (Gunn and Kinzer, 1949) at sea level and after applying
121 altitude corrections (Beard, 1976) for the elevation of 1.4 km MSL for Greeley.”

122
123 The pressure adjustment to the Gunn-Kinzer fit to account for the altitude (1.4 km MSL)
124 of the Greeley site is in excellent agreement with the 2DVD measurements. The slight
125 underestimation of the MPS fall speeds relative to the fit of Foote and du Toit to the data
126 of Gunn-Kinzer in the range 0.7-1.5 mm (max deviation of 0.5 m/s or 10%) is puzzling
127 given the excellent agreement for the Huntsville site (sea level). We have added
128 sentence below.

129
130 “However, the altitude-adjusted FT fit is slightly higher than the measured values
131 as shown in Table 1.”

- 132
133 6. Figure 1b and Figure 2b: I suggest to plot the fall velocity histogram also for other drop
134 diameters (let say 0.7 mm and 1.5 mm for example) so the readers can have more cases to
135 evaluate the agreements between 2DVD and MPS.

136
137 Done as suggested...see new panel c in figs. 1 and 3.

- 138
139 7. Line 187: similarly to comment 5, also here the word “excellent” is not appropriate due to the
140 overestimation of MPS with respect to Gunn and Kinzer fit for $D < 0.5$ mm.

141 As noted in our response to (5) above, after replacing the Atlas et al fit by the Foote and
142 du Toit fit in Fig. 3a the agreement between MPS and the latter fit is considered to be
143 excellent.

- 144 8. Figure 3a: can the Authors provide an explanation of the differences in the mean fall velocity
145 between Gunn and Kinzer fit and MPS measurements for $D > 1.5$ mm?

146
147 Please see our response to (7) above. The differences in Fig. 3a are no longer an issue.

- 148
149 9. Figure 5: what about large drops? Which is the effect of wind on large drops? I suggest to use
150 the 2DVD data to made the same analysis for larger D.

151

152 *We have added new panels Fig. 5e,f which show the effect of wind/gusts for 3 mm*
153 *drops. The trend is similar to 2 mm drops. We cannot show larger drops due to very low*
154 *number of samples.*

155
156

157 TECHNICAL CORRECTIONS 1. Line 286: probably “wind range” should be “wide range”.

158 *Corrected. Thank you for pointing this out.*

159
160
161
162
163
164
165
166
167
168
169
170
171
172
173
174
175
176
177
178

179 Interactive comment on “Raindrop Fall Velocities from an Optical Array Probe and 2D-video
180 Disdrometer” by Viswanathan Bringi et al.

181 H. Leijnse (Referee) hidde.leijnse@knmi.nl Received and published: 11 December 2017: RC2

182 This paper describes results from two measurement campaigns with a Meteorological Particle Sensor
183 (MPS) and a 2D-Video Disdrometer (2DVD). The analyses presented in this paper are focussed on the fall
184 speeds of droplets measured by the different instruments, and whether these deviate from results from
185 laboratory experiments (super or sub-terminal fall speeds). Observed sub-terminal fall speeds are then
186 linked to turbulence intensity. I think this is an interesting paper. It contributes to the scientific
187 discussion on the puzzling super-terminal small raindrops by showing results where these were not
188 observed. However, the paper would benefit from a clearer description of its aims, and, if possible,
189 stronger conclusions. As far as I understand it, there are three main messages in the paper: 1) there is
190 no evidence of super-terminal raindrops (contrary to Montero-Martinez et al., 2009 and Larsen et al.,
191 2014); 2) the fall velocities of real drops closely follow the relations found by Gunn and Kinzer (1949)
192 down to very small drops; and 3) there is a clear effect of strong turbulence on the mean and the
193 standard deviation of fallspeeds of drops of a given diameter. If this is indeed the case, then I think this
194 should be more clearly stated in the introduction, and should be discussed more elaborately in the
195 conclusions. So I think that after revisions, this paper is suitable for publication in Atmospheric
196 Measurement Techniques. Specific comments are given below.

197 ***We thank the reviewer for the positive comments. Our response is in italics while the text***
198 ***modifications are highlighted and in quotes.***

199 *We have added in the Introduction the following:*

200 “In these and other applications it is generally accepted that there is a unique fall speed
201 ascribed to drops of a given mass or diameter and that it equals the terminal speed
202 with adjustment for pressure (e.g., Beard 1976).”

203 *We are reluctant to draw firmer or more elaborate conclusions than what is stated in our paper*
204 *since, (a) the database is rather small (3 cases but from two different climatologies), (b) we do*
205 *not have a direct measurement of turbulence and (c) we cannot quantify if the DFIR wind shield*
206 *is affecting the sensor area in some subtle way. These questions will be addressed in the future.*

207 *In Section 3 last para we have clarified as follows:*

208 “One caveat is that the response of the DFIR wind shield to ambient winds in terms of producing
209 subtle vertical air motions near the sensor area is yet to be evaluated as future work. Analysis
210 of further events with direct measurement of turbulent intensity, for example using a 3D-sonic
211 anemometer at the height of the sensor, would be needed to generalize our findings.”

212 Specific comments

213 1. In the introduction it should be more clearly stated what the exact aims of this paper are.

214
215
216

217
218
219
220
221
222

223
224
225
226
227
228
229
230
231
232
233
234
235
236
237
238
239
240
241
242
243
244
245
246
247
248
249
250
251
252
253

We have added the following in the Introduction:

“Thus, there is some evidence that rain drops may not fall at their terminal velocity except under calm conditions and that the concept of a fall speed distribution for a drop of given mass (or, diameter) might need to be considered which is the topic of this paper. The implications are rather profound especially for numerical modeling of collision-coalescence and breakup processes which are important for shaping the drop size distribution.”

2. On lines 66-74, the use of a DFIR is discussed along with its effects on the local windfield. For studying the relation between turbulence intensity and raindrop fall speeds, how does this double fence affect the turbulence just above the instrument? I can imagine that by reducing the average wind speed, the turbulence is also reduced. On the other hand, as stated on line 74, the fence itself also generates up- and downdrafts. I think that the effects of the use of a DFIR on the results presented in this paper should be discussed and, if possible, quantified.

As mentioned in the text we do not have a direct measurement of turbulence at the height of the sensor. Rather we use the wind/gusts from the anemometer at 10 m high tower as a proxy for turbulence. Regarding the DFIR perturbing our results, there is unfortunately not many articles describing the DFIR affect on the ambient flow other than the quoted reference of Theriault et al. (2015). Most articles are related to the “catch efficiency” of snow gages located inside the DFIR relative to gages with standard wind skirts. In the future we will locate one 2DVD inside the DFIR and one outside as well as collocated 3D-sonic anemometers both inside and outside...but this is another field project for which we have to acquire funding. Unfortunately, we cannot quantify the effect of the DFIR in this paper.

3. In Fig.1 (especially panel a) the MPS seems to detect slightly (but systematically) lower fall velocities than the 2DVD in the Greeley data. This is not the case for the Huntsville data (Fig. 3). Please give an explanation for this.

Please see new Table 1 for a quantitative comparison between MPS and 2DVD in the overlap region. The agreement between the two instruments is, in our opinion, excellent for both sites given the quoted accuracies in fall speed measurement for both instruments.

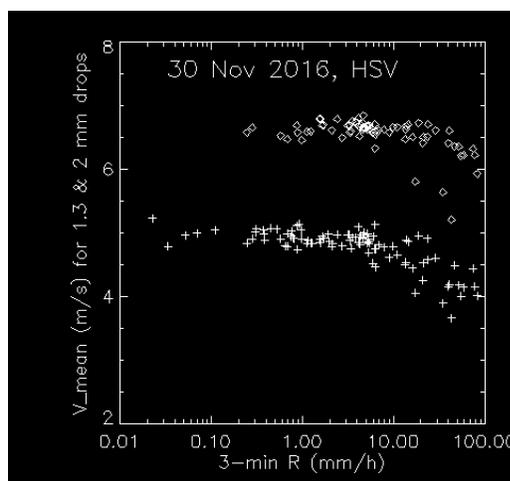
We have replaced the fit to Gunn-Kinzer data with the 9th order polynomial fit given in Foote and du Toit (1969) as opposed to using the exponential fit of Atlas et al. (1973) in both figs. 1a and 3a. The latter fit is not as accurate for the small drop end (e.g., it does not pass through the origin).

254 The pressure adjustment to the Gunn-Kinzer fit to account for the altitude (1.4 km MSL) of the
255 Greeley site is in excellent agreement with the 2DVD measurements. The slight underestimation
256 of the MPS fall speeds relative to the fit of Foote and du Toit to the data of Gunn-Kinzer in the
257 range 0.7-1.5 mm (max deviation of 0.5 m/s or 10%) is puzzling given the excellent agreement
258 for the Huntsville site (sea level). We have re-worded the sentences in Section 2.1 as follows:
259

260 "Panel (a) demonstrates the excellent "visual" agreement between the two
261 instruments in the overlap size range (0.7-2 mm) which is quantified in Table 1.
262 However, the altitude-adjusted FT fit is slightly higher than the measured values
263 as shown in Table 1."
264

265 4. On lines 208-215, the correlation between E on the one hand, and the mean and standard
266 deviation of the raindrop fall speeds on the other is discussed. I agree that this correlation is
267 there. However, judging from Fig. 4, I think there is also some correlation with the rain rate R
268 (especially the peak at 10 UTC). Please elaborate on the role of the rain rate for these
269 correlations.

270
271 The reviewer is correct in that there is a correlation with rain rate (as shown in the figure below
272 for the reviewer's benefit) but this can be misleading in that heavy rain can and does occur
273 during calm conditions and the opposite also occurs. So it is difficult to come to a firm conclusion
274 unless we look at many cases which is reserved for future work.
275



276
277 Figure 1: The mean fall speed versus rain rate for 30 Nov 2016 Huntsville for 1.3 and 2 mm
278 drops.

279 5. On lines 224-225, the observed near-linear decrease of the mean fall speed with turbulence
280 intensity (or at least its proxy E) is mentioned. How significant is this relation?
281

282 *Not sure if what is meant by “significant” relation. We can do a linear fit and compute*
283 *correlation coefficient but this won’t add much to what is already fairly obvious. We do not wish*
284 *to quantify this as E is only a proxy for turbulence.*

285

286 6. On lines 247-254, the results presented by Montero-Martinez et al. (2009) are compared to
287 those presented in this paper. What could be the explanation for this difference? Could it be
288 something similar to what is discussed in the next paragraph (lines 255-262) about the findings
289 of Larsen et al. (2014)? Please elaborate on this.

290

291 *We can only speculate why Montero-Martinez et al. found strongly skewed distribution*
292 *for 0.44 mm drops. The instruments they used, 2D-C and 2D-P, were designed for use*
293 *on aircraft and not as fixed disdrometers. The airspeed clock is about a factor of 10*
294 *higher in their implementation. The calibration method is not discussed in any detail by*
295 *them whereas the MPS can be calibrated as often as needed with a special device.*
296 *They did not use any wind shield as far as we can ascertain.*

297

298 *The Larsen et al. study used a 2DVD which has “mis-matched” drop problem not present*
299 *in the 2D-C,P probes.*

300

301 7. On lines 272-281, the relation to the findings of Stout et al. (1995) are discussed. Is there an
302 empirical relation between E and the rms velocity fluctuations due to turbulence? If so it would
303 be interesting to see whether the 35% reduction in mean velocity is observed at similar rms
304 velocity fluctuations-to-terminal fall speed ratios (0.8).

305

306 *We do not think there is a way to relate our estimate of E based on 3-s wind data to what a 3D-*
307 *sonic anemometer would measure in terms of velocity fluctuations at much higher sampling*
308 *rate. This for a future project where we would locate a sonic anemometer next to the 2DVD/MPS*
309 *inside the DFIR.*

310 Technical comments

311 1. In Figs 2b and 4b, would it be possible to use a second y-axis for R instead of presenting R/10 on
312 the existing y-axis?

313

314 *Done as suggested.*

315

316 2. On line 275, “greater that” should be “greater than”.

317 *Corrected. Thanks for pointing this out.*

318 Interactive comment on Atmos. Meas. Tech. Discuss., doi:10.5194/amt-2017-401, 2017.

319 Review of AMT-2017-401 By V. Bringi, M. Thurai and D. Baumgardner Manuscript Title – Raindrop Fall
320 Velocities from an Optical Array Probe and 2D-Video Disdrometer. RC3

321

322 This manuscript reports on raindrop fall velocity measurements by using two different instruments: a
323 MPS (Meteorological Particle Spectrometer) which measures drops in the 0.1-3 mm range, and the
324 widely used 2DVD (two-Dimensional Video Disdrometer), which measures size and fall velocity of drops
325 between 0 and 10 mm. The MPS and 2DVD were used to measure fall velocity of drops in the 0.1-2 and
326 larger than 0.7 mm diameter range. The overlapping region 0.7-2 mm diameter was used to cross-
327 validate the two measurements. Three different case studies were analyzed in order to relate the
328 properties of the drop fall velocity to different precipitation systems (one stratiform, one squall line and
329 one super-cell case with low and high turbulence associated for the first two and the third case,
330 respectively). The paper is linear and quite easy to read. I have only one major comment that can give a
331 contribution, in my opinion, to the generalization of the results. It is reported below together with minor
332 comments that, once addressed, will allow the publication of the paper on the Atmospheric
333 Measurement Techniques journal.

334 ***The authors appreciate the above general comments and our response is given below in***
335 ***italics whereas the text modifications are highlighted.***

336

337 Major comment.

338 - Section 2.2: in the Section 2.1 the authors investigated a stratiform case, while in the Section 2.2 a
339 squall line and a super-cell case. The squall line case reported generally low rainfall rate and turbulence
340 (comparable to the values registered in the stratiform case). It could be useful, in my opinion add (or
341 substitute) a convective event, a sort a middle point between a convective and tornadic case, in order to
342 have a general overview of the characteristics of drop fall velocity in a broader range of precipitation
343 systems.

344 *In Section 2.2 it is clearly stated that..." About 3 h later several squall-line type storm cells*
345 *passed over the site from 0700-0900 UTC again with strong winds but considerably lower E*
346 *values 2-4 m² s⁻² and maximum R of 80 mm h⁻¹. After 1000 UTC the E values were much*
347 *smaller (< 0.5 m² s⁻²) indicating calm conditions. The peak R is also smaller at 30 mm h⁻¹ at*
348 *1000 UTC." Hence, Fig. 4b already depicts high and moderate-low rain rate conditions.*
349 *Perhaps the reviewer overlooked the rain rate scale (where we plotted R/10) which has now*
350 *been changed, the values now on the right Y-axis without any scaling.*

351

352

353

354

355

356

357 Minor comments.

358

359 - Line 141: what does it mean that the finite bin width causes a spread of 0.5 m/s? Can the authors
360 explain better? The same is reported in other parts of the text.

361 *The bin width for the histograms is ± 0.1 mm about the center value of 0.5 mm , for example in*
362 *Fig. 1b all fall speeds that fall in the range 0.4 to 0.6 mm are included in the histogram. The*
363 *spread of 0.5 m/s is just $V(0.6 \text{ mm}) - V(0.4 \text{ mm})$. In Section 2.1 we have clarified as:*

364 **"The finite bin width used (0.9-1.1 mm) causes a corresponding fall speed "spread" of**
365 **around 0.5 m s^{-1} which is clearly a significant contributor to the measured coefficient of**
366 **variation."**

367 *Please see new Table 1 which quantifies the finite bin width spread in Column 1.*

368 - Lines 211-215: what is the explanation that the authors give to the decrease of fall speed during the
369 most intense wind and rainfall rate? Does it can be related to the presence of ascending flow?

370 *In Section 3 we refer to Stout et al. (1995) who have simulated the effects of turbulence to*
371 *cause a decrease in fall speed relative to still air conditions. The effect is due to increase in the*
372 *non-linear drag due to both vertical and horizontal gusts. We defer to the explanation in Stout et*
373 *al. (1995) but we introduce this reference earlier in Section 2.2.*

374 - Lines 224-225: similar to the previous comment. How do they justify the decrease of fall speed when E
375 (turbulence) increases?

376 *Please refer to our response above.*

377

378 - Panel (b) of Figures 2 and 4: the rain rate should be reported on the right y-axis avoiding the necessity
379 to show its values scaled on a factor ten.

380 *Done as requested.*

381

382 - Figure 2a: the y-axis limit should not exceed 10 m/s to improve the detail of the plot.

383 *Done as requested.*

384

385

386 Interactive comment on “Raindrop Fall Velocities from an Optical Array Probe and 2D-video
387 Disdrometer” by Viswanathan Bringi et al.

388 Anonymous Referee #4 Received and published: 18 December 2017: RC4

389 GERNERAL COMMENTS This is an excellent paper, addressing the issue of in-situ measurements of rain
390 drops over their full size range, with focus on the drops’ fall velocities. Especially the characteristics of
391 very small drops still are not investigated thoroughly up to here. The investigation of the relationship
392 between wind turbulence and rain drops’ fall velocity is convincing and in the future will allow new
393 refinements in the related remote sensing algorithms. The study is presented in a clear and concise way.
394 The reviewer likes to point out, that such work does not only address academic interests, but points to a
395 number of applications, like remote sensing of the atmosphere, as the authors shortly mention in their
396 introduction. Other applications include knowledge of channel characteristics for satellite
397 communications at EHF frequency bands. Thus the study is very relevant for these fields of science and
398 applications. The work is based on the combined analyses of data from 2 instruments at same site, the
399 Meteorological Particle Spectrometer (MPS) and the 2D-video disdrometer (2DVD). Field measurement
400 of rain drops still may be considered as challenge, with quite a few solutions being around, but none of
401 the instruments may claim to be perfect. The presented results do rise a few relevant questions, which
402 are given in below specific comments. Summarizing it is said, that this paper gives an excellent
403 contribution, bringing new aspects for the relevant fields of science and applications.

404 ***We appreciate the very positive comments by the reviewer. Our response is in italics***
405 ***below and modifications to the text are highlighted.***

406 SPECIFIC COMMENTS

407 *** Line 98 / 99: “All possible corrections have been applied, including the removal of artifacts due to
408 splashing,...” Careful preprocessing, verification and validation of data is of utmost importance for the
409 present study. The results presented require correctness of challenging measurement processes. Some
410 of the result given actually leave a few questions open, as addressed in below comment to Fig 1a and
411 Fig3a. To be sufficiently self-contained, thus the paper might shortly describe the mentioned correction
412 algorithms. .

413 *We have replaced the sentence: “All possible corrections have been applied, including the removal of*
414 *artifacts due to splashing, and oversizing that results from out-of-focus droplets (Korolev 2007).” With*

415 *“There are a number of potential artifacts that arise when making measurements with optical array*
416 *probes (Baumgardner et al., 2017): droplet breakup on the probe tips that form satellite droplets,*
417 *multiple droplets imaged simultaneously, and out-of-focus drops whose images are usually larger than*
418 *the actual drop (Korolev, 2007). The measured images have been analyzed to remove satellite droplets*
419 *whose interarrival times are usually too short to be natural drops, multiple drops are detected by shape*

420 *analysis and removed, and out-of-focus drops are detected and size corrected using the technique*
421 *described by Korolev (2007)."*

422 *** Line 157: "The histogram from MPS for the 0.5 mm sizes shows positive skewness" Can the authors
423 give an explanation / discussion / assumption?

424 *The histogram in Fig. 1b for 0.5 mm size is skewed towards higher values...the tail extends*
425 *from 2.6 to 4 m/s . Given that the mean is 1.8 m/s and $\sigma=0.65$ m/s, there is a finite occurrence of*
426 *super-terminal fall speeds that exceed mean+1 σ or 2.45 m/s. Such super-terminal speeds have*
427 *been noted by Montero-Martinez et al using 2DC,P probes for similar sized drops in fact with*
428 *much longer tail or skewness than our observations.*

429 *** Fig. 1a and Fig. 3a: Mean fall velocities from the MPS are read from these figures approximately as:
430 Fig. 1a: D = 0,7 mm, v = 2.55 m/s

431 Fig. 3a: D = 0.7 mm, v = 3.05 m/s

432 That represents an exceedance by more than 16 % (Fig. 3a over Fig. 1a), in spite of the lower pressure in
433 Greeley (Fig 1a) leading to the expectation of faster drops than in Huntsville (Fig 3a). The authors please
434 could discuss this.

435 *We refer to new Table 1 from which we now have:*

436 *Greeley site D=0.7 mm mean MPS=2.6 m/s*

437 *Huntsville D=0.7 mm mean MPS=2.6 m/s*

438 *We have also added new histograms for 0.7 and 1.5 mm from both sites (new Fig. 1c and 3c).*
439 *The 0.7 mm histogram from Greeley is slightly skewed (Fig. 1c) whereas the Huntsville*
440 *histogram is more symmetric. For small drops (< 1 mm) the pressure adjustment is not*
441 *significant (see Fig. 1a) and this is reflected by Table 1 data.*

442

443 *** Fig. 3a: Mean fall velocities from the MPS are read from this figure approximately as:

444 D = 1.5 mm, v = 5.28 m/s

445 D = 1.6 mm, v = 6.31 m/s

446 D = 1.7 mm, v = 6.5 m/s

447 D = 1.8 mm, v = 5.93 m/s

448 These values differ significantly from the fit to Gunn-Kinzer, further from the expected monotonic
449 behaviour. The authors please could discuss this.

450 We have added a new Table 1 which quantifies the measurement comparison between MPS and 2DVD in
 451 the overlap region for the 2 sites. The agreement in the mean fall speed is excellent for both sites. The
 452 mean values quoted above by the reviewer are not accurate but we do appreciate the effort involved to
 453 read values of a graph. In particular at $D=1.7$ mm, from Table 1 the mean is 6 ± 0.3 m/s (and not 5.93 m/s
 454 as quoted above). At $D=1.8$ mm, interpolation gives 6.25 m/s which is monotonic behavior. Table 1 also
 455 gives the expected values using the Foote and du Toit (1969) 9th order polynomial fit (FT) to Gunn-Kinzer
 456 which is more accurate than the exponential fit of Atlas et al especially for the smaller drops. We list
 457 below the FT fit value and mean MPS from Table 1 for Huntsville site.

458	0.7 mm	FT fit=2.9 m/s	MPS=2.6	2DVD=2.5
459	0.9	3.65	3.4	3.3
460	1.1	4.3	4.2	4.1
461	1.3	4.9	4.9	4.9
462	1.5	5.45	5.4	5.4 (see new histogram in Fig. 3c)
463	1.7	5.9	6.0	5.8
464	1.9	6.3	6.5	6.3

465 For 0.7 and 0.9 mm sizes the MPS and 2DVD mean values are systematically lower than the FT fit
 466 by 7-10%. For comparison, Yu et al. (2016; new reference added in revised text) found that their high
 467 speed camera measurement of fall speeds in the interior stair-case of a building also were systematically
 468 lower by 5% (for sizes in the range 0.4 to 1.6 mm) relative to the same FT fit as used here (their fig. 7).

469 TECHNICAL COMMENTS:

470 *** Line 230: "(σf) versus E is shown in panels 6 (b,d)." It probabaly should read as "(σf) versus E is
 471 shown in panels (b,d) of Fig 5."

472 Done...thank you for pointing this out.

473 Interactive comment on Atmos. Meas. Tech. Discuss., doi:10.5194/amt-2017-401, 2017.

474

475

476

477

478

479

480 Raindrop Fall Velocities from an Optical Array Probe and 2D-Video
481 Disdrometer

482 Viswanathan Bringi¹, Merhala Thurai¹ and Darrel Baumgardner²

483 ¹ *Department of Electrical and Computer Engineering, Colorado State University, Fort*
484 *Collins, Colorado, USA*

485 ² *Droplet Measurements Technologies, Longmont, Colorado, USA*

486 Correspondence to: V.N. Bringi

487 Email: bringi@colostate.edu

488

489

490 Abstract

491 We report on fall speed measurements of rain drops in light-to-heavy rain events from
492 two climatically different regimes (Greeley, Colorado, and Huntsville, Alabama) using
493 the high resolution (50 μm) Meteorological Particle Spectrometer (MPS) and a 3rd
494 generation (170 μm resolution) 2D-video disdrometer (2DVD). To mitigate wind-effects,
495 especially for the small drops, both instruments were installed within a 2/3-scale Double
496 Fence Intercomparison Reference (DFIR) enclosure. Two cases involved light-to-
497 moderate wind speeds/gusts while the third case was a tornadic supercell and several
498 squall-lines that passed over the site with high wind speeds/gusts. As a proxy for
499 turbulent intensity, maximum wind speeds from 10-m height at the instrumented site
500 recorded every 3 s were differenced with the 5-min average wind speeds and then
501 squared. The fall speeds versus size from 0.1-2 mm and >0.7 mm were derived from
502 the MPS and the 2DVD, respectively. Consistency of fall speeds from the two
503 instruments in the overlap region (0.7-2 mm) gave confidence in the data quality and
504 processing methodologies. Our results indicate that under low turbulence, the mean fall
505 speeds agree well with fits to the terminal velocity measured in the laboratory by Gunn
506 and Kinzer from 100 μm up to precipitation sizes. The histograms of fall speeds for 0.5,
507 0.7, 1 and 1.5 mm sizes were examined in detail under the same conditions. The
508 histogram shapes for the 1 and 1.5 mm sizes were symmetric and in good agreement
509 between the two instruments with no evidence of skewness or of sub- or super-terminal
510 fall speeds. The histograms of the smaller 0.5 and 0.7 mm drops from MPS while
511 generally symmetric showed that occasional occurrences of sub- and super-terminal fall
512 speeds could not be ruled out. In the supercell case, the very strong gusts and inferred
513 high turbulence intensity caused a significant broadening of the fall speed distributions
514 with negative skewness (for drops of 1.3, 2 and 3 mm). The mean fall speeds were also
515 found to decrease nearly linearly with increasing turbulent intensity attaining values
516 about 25-30% less than the terminal velocity of Gunn-Kinzer, i.e. sub-terminal fall
517 speeds.

518

519

520 1 Introduction

521 Knowledge of the terminal fall speed of raindrops as a function of size is important in
522 modelling collisional break-up and coalescence processes (e.g., *List et al.*, 1987), in the
523 radar-based estimation of rain rate, in retrieval of drop size distribution using Doppler
524 spectra at vertical incidence (e.g., *Sekhon and Srivastava*, 1971) and in soil erosion
525 studies (e.g., *Rosewell* 1986). In these and other applications it is generally accepted
526 that there is a unique fall speed ascribed to drops of a given mass or diameter and that
527 it equals the terminal speed with adjustment for pressure (e.g., *Beard* 1976). The
528 terminal velocity measurements of *Gunn and Kinzer*, 1949) under calm laboratory
529 conditions, and fits to their data (e.g., *Atlas et al.*, 1973; *Footo and du Toit*, 1969; *Beard*
530 *and Pruppacher*, 1969) are still considered the standard against which measurements
531 using more modern optical instruments in natural rain are compared (*Löffler-Mang and*
532 *Joss*, 2000; *Barthazy et al.*, 2004; *Schönhuber et al.*, 2008; *Testik and Rahman*, 2016;
533 *Yu et al.*, 2016). More recently, the broadening and skewness of the fall speed
534 distributions of a given size (3 mm) in one intense rain event were attributed to mixed-
535 mode amplitude oscillations (*Thurai et al.*, 2013). Super- and sub-terminal fall speeds in
536 intense rain shafts have been detected and attributed, respectively, to drop breakup
537 fragments (sizes < 0.5 mm), and high wind/gusts (sizes 1-2 mm) (*Montero-Martinez et*
538 *al.*, 2009; *Larsen et al.*, 2014; *Montero-Martinez and Garcia-Garcia*, 2016). Thus, there
539 is some evidence that rain drops may not fall at their terminal velocity except under
540 calm conditions and that the concept of a fall speed distribution for a drop of given mass
541 (or, diameter) might need to be considered which is the topic of this paper. The
542 implications are rather profound especially for numerical modeling of collision-
543 coalescence and breakup processes which are important for shaping the drop size
544 distribution.

545 The fall speeds and concentration of small drops (< 1 mm) in natural rain are difficult to
546 measure accurately given the poor resolution (>170 μm) of most optical disdrometers
547 and/or sensitivity issues. While cloud imaging probes (with high resolution 25-50 μm)
548 on aircraft have been used for many years they generally cannot measure the fall
549 speeds. A relatively new instrument, the Meteorological Particle Spectrometer (MPS) is
550 a droplet imaging probe that was built by Droplet Measurements Technologies (DMT,
551 Inc.) under contract from the US Weather Service specifically designed for drizzle as
552 small as 50 μm and rain drops up to 3 mm. This instrument in conjunction with a lower
553 resolution 2D-Video Disdrometer (*Schoenhuber et al.*, 2008) is used in this paper to
554 measure fall speed distributions in natural rain.

555 This paper briefly describes the instruments used, presents fall speed measurements
556 from two sites under relatively low wind conditions, and one case from an unusual

557 tornadic supercell with high winds and gusts and ends with a brief discussion and
558 summary of the results.

559

560 2 Instrumentation and Measurements

561 The principal instruments used in this study are the MPS and 3rd generation 2D-video
562 disdrometer (2DVD), both located within a 2/3-scale Double Fence Intercomparison
563 Reference (DFIR; *Rasmussen et al.*, 2012) wind shield. As reported in (*Notaros et al.*,
564 2016), the 2/3-scale DFIR was effective in reducing the ambient wind speeds by nearly
565 a factor of 2-3 based on data from outside and inside the fence. The flow field in and
566 around the DFIR has been simulated by (*Theriault et al.*, 2015) assuming steady
567 ambient winds. They found that depending on the wind direction relative to the
568 octagonal fence, weak vertical motions could be generated above the sensor areas. For
569 5 m/s speeds, the motions could range between -0.4 (down draft) to 0.2 m/s (up draft).

570 The instrument set-up was the same for the two sites (Greeley, Colorado and
571 Huntsville, Alabama). Huntsville has a very different climate from Greeley, and its
572 altitude is 212 m MSL as compared with 1.4 km MSL for Greeley. According to the
573 Köppen–Trewartha climate classification system (*Trewartha and Horn*, 1980), this labels
574 Greeley as a semiarid-type climate, whereas Huntsville is a humid subtropical-type
575 climate (*Belda et al.*, 2014).

576 The MPS is an optical array probe (OAP) that uses the technique introduced by
577 *Knollenberg* (1970, 1976, 1980) and measures drop diameter in the range from 0.05-3.1
578 mm. A 64 element photo-diode array is illuminated with a 660 nm collimated laser
579 beam. Droplets passing through the laser cast a shadow on the array and the decrease
580 in light intensity on the diodes is monitored with the signal processing electronics. A two
581 dimensional image is captured by recording the light level of each diode during the
582 period that the array is shadowed. The fall velocity is derived using two methods. One
583 uses the same approach as described by (*Montero-Martinez et al.*, 2009) where the fall
584 velocity is calculated from the product of the true air speed clock and ratio of the image
585 height -to-width. Note that “width” is the horizontal dimension parallel to the array and
586 “height” is along the vertical. The second method computes the fall velocity from the
587 maximum horizontal dimension (spherical drop shape assumption) divided by the
588 amount of time that the image is on the array, a time measured with a 2 MHz clock. In
589 order to be comparable to the results of (*Montero-Martinez et al.*, 2009), their approach
590 is implemented here for sizes > 250 μm . The fall velocity of smaller, slower moving
591 droplets, is measured using the second technique.

592 The limitations and uncertainties associated with OAP measurements have been well
593 documented (Korolev *et al.*, 1991; 1998; Baumgardner *et al.*, 2017). There are a
594 number of potential artifacts that arise when making measurements with optical array
595 probes (Baumgardner *et al.*, 2017): droplet breakup on the probe tips that form satellite
596 droplets, multiple droplets imaged simultaneously, and out-of-focus drops whose
597 images are usually larger than the actual drop (Korolev, 2007). The measured images
598 have been analyzed to remove satellite droplets whose interarrival times are usually too
599 short to be natural drops, multiple drops are detected by shape analysis and removed,
600 and out-of-focus drops are detected and size corrected using the technique described
601 by (Korolev 2007). The sizing and fall speed errors primarily depend on the digitization
602 error ($\pm 25 \mu\text{m}$). The fall speed accuracy according to the manufacturer (DMT) is $<10\%$
603 for 0.25 mm and $<1\%$ for sizes greater than 1 mm , limited primarily by the accuracy in
604 droplet sizing.

605 The 3rd generation 2DVD is described in detail by (Schoenhuber *et al.*, 2007; 2008) and
606 its accuracy of size and fall speed measurement has been well documented (e.g.,
607 Thurai *et al.*, 2007; 2009; Huang *et al.*, 2008; Bernauer *et al.*, 2015). Considering the
608 horizontal pixel resolution of $170 \mu\text{m}$ and other factors (such as “mis-matched” drops),
609 the effective sizing range is $D > 0.7 \text{ mm}$. To clarify the “mis-matched” drop problem: it is
610 very difficult to match a drop detected in the top light-beam plane of the 2DVD to the
611 corresponding drop in the bottom plane for tiny drops resulting in erroneous fall speeds.
612 The fall velocity accuracy is determined primarily by the accuracy of calibrating the
613 distance between the two orthogonal light “sheets” or planes and is $< 5\%$ for fall velocity
614 $< 10 \text{ m s}^{-1}$. In our application, we utilize the MPS for measurement of small drops
615 with $D < 1.2 \text{ mm}$. The measurements from the MPS are compared with those from the
616 2DVD in the overlap region of $D \approx 0.7\text{--}2.0 \text{ mm}$ to ensure consistency of
617 observations. The only fall velocity threshold used for the 2DVD is the lower limit set at
618 0.5 m s^{-1} in accordance with the manufacturer guidelines for rain measurements.

619 2.1 Fall Speeds from Greeley, Colorado

620 We first consider a long duration (around 20 h) rain episode on 17 April 2015 which
621 consisted of a wide variety of rain types/rates (mostly light stratiform $< 8 \text{ mm h}^{-1}$) as
622 described in Table 2 of (Thurai *et al.*, 2017). Two wind sensors at a height of 1 m were
623 available to measure the winds outside and inside the DFIR. Average wind speeds
624 were, respectively, $< 1.5 \text{ m s}^{-1}$ inside the DFIR and $< 4 \text{ m s}^{-1}$ outside with light gusts.
625 These wind sensors were specific to the winter experiment described in (Notaros *et al.*,
626 2016) and were unavailable for the rain measurement campaign after May 2015.

627 Figure 1(a) shows the fall speeds versus D from the 2DVD (shown as contoured
628 frequency of occurrence), along with mean and $\pm 1\sigma$ standard deviation from the MPS.
629 Also shown is the (Foote and du Toit 1969) (henceforth FT fit) to the terminal fall speed

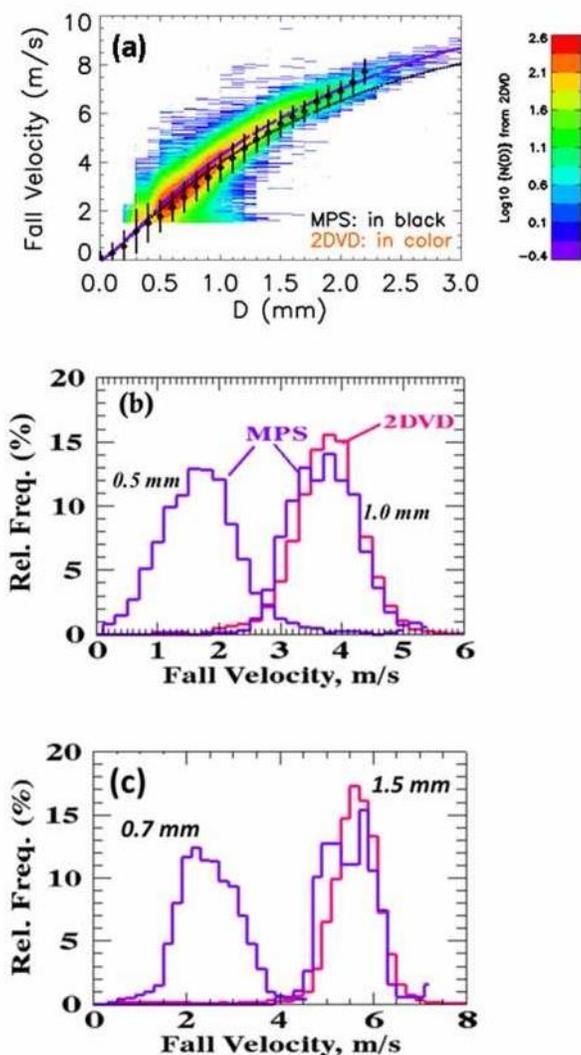
630 measurements of (*Gunn and Kinzer, 1949*) at sea level and after applying altitude
 631 corrections (*Beard, 1976*) for the elevation of 1.4 km MSL for Greeley. Panels (b,c)
 632 shows the histogram of fall speeds for diameter intervals (0.5 ± 0.1) and (1 ± 0.1 mm), and
 633 (0.7 ± 0.1) and (1.5 ± 0.1 mm), respectively. Panel (a) demonstrates the excellent “visual”
 634 agreement between the two instruments in the overlap size range (0.7-2 mm) which is
 635 quantified in Table 1. However, the altitude-adjusted FT fit is slightly higher than the
 636 measured values as shown in Table 1. Notable in Fig. 1a is the remarkable agreement
 637 in mean fall speeds between the FT fit and the MPS for $D < 0.5$ mm down to near the
 638 lower limit of the instrument (0.1 mm). Few measurements have been reported of fall
 639 speeds in this size range.

640 Table 1: Expected fall velocities for various diameter intervals (bin width of 0.2 mm)
 641 from (*Foote and du Toit, 1969*) with altitude adjustment, and the measured mean fall
 642 velocities with $\pm 1\sigma$ (standard deviation)

D range (mm) (Greeley)	Expected ($m\ s^{-1}$) at 1.4 km	MPS ($m\ s^{-1}$) Mean $\pm 1\sigma$	2DVD ($m\ s^{-1}$) Mean $\pm 1\sigma$
0.6 to 0.8	2.6 to 3.5	2.6 ± 0.6	2.5 ± 0.8
0.8 to 1.0	3.5 to 4.3	3.4 ± 0.6	3.3 ± 0.9
1.0 to 1.2	4.3 to 4.9	4.2 ± 0.6	4.1 ± 0.9
1.2 to 1.4	4.9 to 5.5	4.9 ± 0.5	5.0 ± 0.8
1.4 to 1.6	5.5 to 6.1	5.6 ± 0.5	5.7 ± 0.7
1.6 to 1.8	6.1 to 6.6	6.1 ± 0.4	6.2 ± 0.7
1.8 to 2.0	6.6 to 7.0	6.7 ± 0.4	6.6 ± 0.8
D range (mm) (Huntsville)	Expected (m/s) at 0 km	MPS (m/s) Mean \pm Std_dev	2DVD (m/s) Mean \pm Std. dev
0.6 to 0.8	2.5 to 3.3	2.6 ± 0.6	2.5 ± 0.7
0.8 to 1.0	3.3 to 4.0	3.4 ± 0.5	3.3 ± 0.7
1.0 to 1.2	4.0 to 4.6	4.2 ± 0.6	4.1 ± 0.8
1.2 to 1.4	4.6 to 5.2	4.9 ± 0.4	4.9 ± 0.7
1.4 to 1.6	5.2 to 5.7	5.4 ± 0.4	5.4 ± 0.6
1.6 to 1.8	5.7 to 6.1	6.0 ± 0.3	5.8 ± 0.6
1.8 to 2.0	6.1 to 6.5	6.5 ± 0.4	6.3 ± 0.5

643
 644 The histograms in Fig. 1(b,e) show good agreement between 2DVD and MPS for 1 mm
 645 and 1.5 mm drop sizes, respectively, with respect to the mode, symmetry, spectral width
 646 and lack of skewness in the distributions. For the 1 mm size histogram, the mean is 3.8
 647 $m\ s^{-1}$ while the spectral width or standard deviation from MPS data is $0.6\ m\ s^{-1}$. The
 648 corresponding coefficient of variation (ratio of standard deviation to mean) is 15.7%.
 649 The finite bin width used (0.9-1.1 mm) causes a corresponding fall speed “spread” of
 650 around $0.6\ m\ s^{-1}$ which is clearly a significant contributor to the measured coefficient of

651 variation. Similar comments apply to the fall speed histogram for the 1.5 mm size shown
 652 in Fig. 1c. The definition of sub- or super-terminal fall speeds by (Montero-Martinez et
 653 al., 2009) is based on fall speeds that are, respectively, less than 0.7 times the mean
 654 value or greater than 1.3 times the mean value (i.e., exceeding 30% threshold on either
 655 side of the mean terminal fall speed). From examining the 1 mm size fall speed
 656 histogram there is negligible evidence of occurrences with fall speeds $< 2.66 \text{ m s}^{-1}$
 657 (sub) or $> 4.94 \text{ m s}^{-1}$ (super). Similar comment also applies for the 1.5 mm size based
 658 on the corresponding histogram.



659

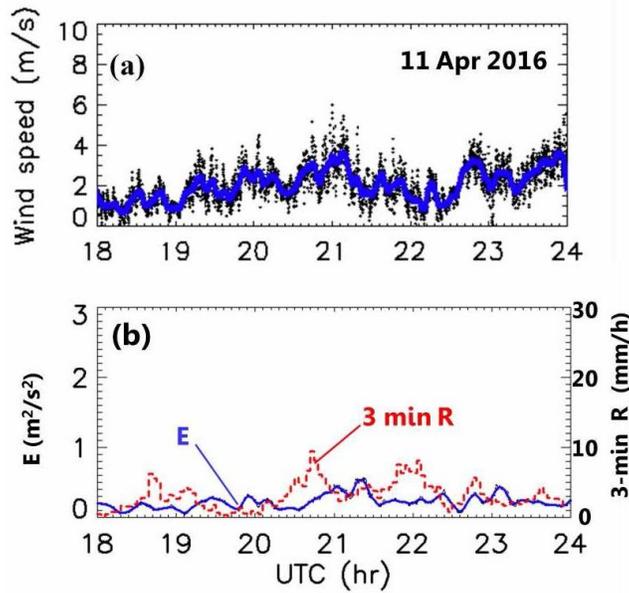
660 *Figure 1.* (a) Fall velocity versus diameter (D). The contoured frequency of occurrence from
 661 2DVD data is shown in color (log scale). The mean fall velocity and $\pm 1\sigma$ standard deviation bars
 662 are from MPS. The dark dashed line is from the fit to the laboratory data of Gunn and Kinzer
 663 (1949) and the purple line is the same except corrected for the altitude of Greeley, CO (1.4 km

664 MSL). (b) Relative frequency histograms of fall velocity for the 0.5 ± 0.1 mm and 1 ± 0.1 mm
665 bins. (c) as in (b) except for the 0.7 ± 0.1 mm and 1.5 ± 0.1 mm bins.

666 The histogram from MPS for the 0.5 mm sizes shows positive skewness with mean of
667 1.8 m s^{-1} , spectral width of 0.65 m s^{-1} and corresponding coefficient of variation nearly
668 doubling to 35% (relative to the 1 mm size histogram). The finite bin width (0.4-0.6 mm)
669 causes a corresponding fall speed “spread” of 0.4 m s^{-1} which contributes to the
670 measured coefficient of variation. Nevertheless, it is not possible to rule out the low
671 frequency of occurrence of sub- or super-terminal fall speeds, respectively, less than
672 1.26 m s^{-1} or exceeding 2.34 m s^{-1} (i.e., exceeding 30% of the mean value) based on
673 our data. Examination of the MPS-based fall speed histogram for the 0.7 mm size
674 indicates negative skewness. As with the 0.5 mm drops it is not possible to rule out the
675 occurrences of fall speeds $< 1.8 \text{ m s}^{-1}$ or $> 3.4 \text{ m s}^{-1}$, i.e., sub- or super-terminal fall
676 speeds.

677 2.2 Fall Speeds from Huntsville, Alabama

678 The first Huntsville event occurred on 11 April 2016 and consisted of precipitation
679 associated with the mesoscale vortex of a developing squall line that moved across
680 northern Alabama between 1800 and 2300 UTC and produced over 25 mm of rainfall in
681 the Huntsville area. Figure 2(a) shows the ambient 10-m height wind speeds (3 s and 5-
682 min averaged) recorded at the site. Maximum speeds were less than 5 m s^{-1} and wind
683 gusts were light. As no direct *in situ* measurement of turbulence was available we use
684 the approach by (Garrett and Yuter, 2014) who estimate the difference between the
685 maximum wind speed, or gust, that was sampled every 3 s, and the average wind
686 speed derived from successive 5 min intervals. The estimated turbulent intensity is
687 proportional to $E = (\text{Gusts} - \text{AverageWind})^2/2$. Figure 2(b) shows the E values which
688 were small (maximum $E < 0.4 \text{ m}^2 \text{ s}^{-2}$) and indicative of low turbulence. Also, shown in
689 Fig. 2(b) is the 2DVD-based time series of rainfall rate (R) averaged over 3 mins; the
690 maximum R is around 10 mm h^{-1} .



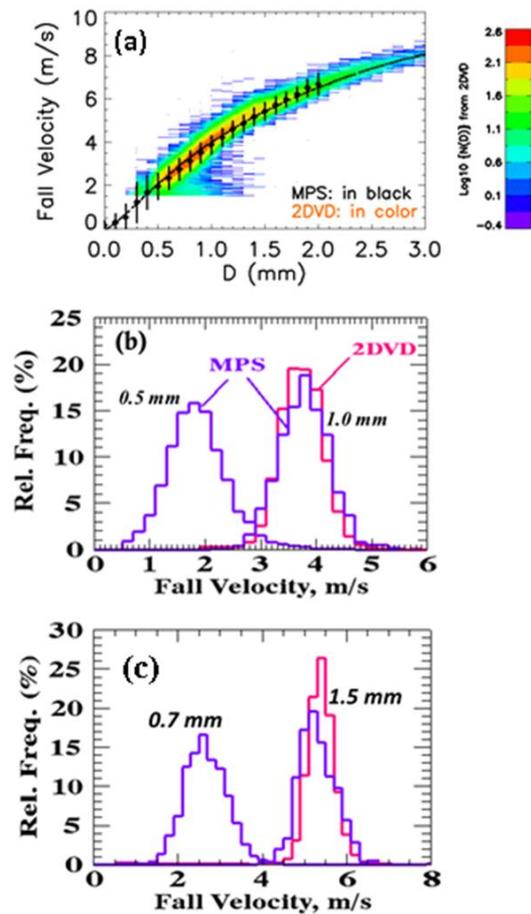
691

692 *Figure 2:* (a) 3-s raw and 5-min averaged wind speeds at 10-m height. (b) turbulent
 693 intensity estimates E , and 3-min averaged R .

694

695 Figure 3(a) shows the fall velocity versus D comparison between the two instruments
 696 while panels (b,c) show the histograms for the 0.5 and 1 mm, and 0.7 and 1.5 mm
 697 sizes, respectively. Similar to the Greeley event, the mean fall speed agreement
 698 between both instruments in the overlap region is excellent (see Table 1) and consistent
 699 with the FT fit to the Gunn-Kinzer laboratory data. As in Fig. 1(a), the MPS data in Fig.
 700 3(a) is in excellent agreement with FT fit for sizes < 0.5 mm.

701 The 0.5 and 1 mm histogram shapes in Fig. 3(b) are quite similar to the Greeley case
 702 shown in Fig. 1(b). The mean and standard deviations from the MPS data for the 0.5
 703 and 1 mm bins are, respectively, $[2 \pm 0.62]$ and $[3.88 \pm 0.44]$ m s^{-1} . The values for the
 704 0.7 and 1.5 mm bins are, respectively, $[2.6 \pm 0.6]$ and $[5.4 \pm 0.4]$ m s^{-1} . There is
 705 negligible evidence of sub- or super-terminal fall speed occurrences based on the 1 and
 706 1.5 mm histograms. The comments made earlier with respect to Fig. 1(b,c) of the
 707 Greeley event for the 0.5 and 0.7 mm histograms are also applicable here, i.e., we
 708 cannot rule out the occasional occurrences of sub- or super-terminal fall speeds based
 709 on our data.



710

711 *Figure 3. (a) as in Fig. 1(a) except for 11 April 2016 event. The dashed line is fit to*
 712 *Gunn-Kinzer at sea level. (b,c) as in Fig. 1(b,c) except for 11 April 2016 event.*

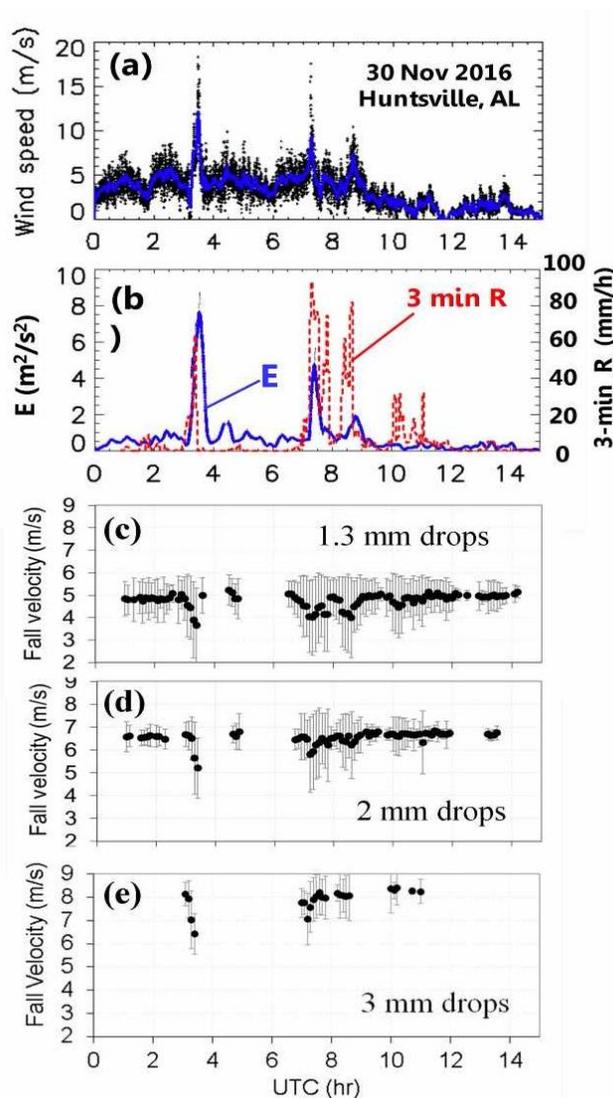
713

714 The second case considered is from 30 November 2016 wherein a supercell passed
 715 over the instrumented site from 0300-0330 UTC producing about 15 mins later a long-
 716 lived EF-2 tornado. Strong winds were recorded at the site with 5-min averaged speeds
 717 reaching 10-12 m s⁻¹ between 0320-0330 and E values in the range to 7-8 m² s⁻²
 718 indicating strong turbulence (Fig. 4a,b). The rain rates peaked at 70 mm h⁻¹ during this
 719 time (Fig. 4b). About 3 h later several squall-line type storm cells passed over the site
 720 from 0700-0900 UTC again with strong winds but considerably lower E values 2-4 m² s⁻²
 721 and maximum R of 80 mm h⁻¹. After 1000 UTC the E values were much smaller (< 0.5
 722 m² s⁻²) indicating calm conditions. The peak R is also smaller at 30 mm h⁻¹ at 1000 UTC.

723 *Figure 4 panels (c), (d) and (e) show the mean and ±1σ of the fall speeds from the*
 724 *2DVD for the 1.3, 2 and 3 mm drop sizes, respectively. The MPS data are not shown*

725 here since during this event it was located outside the DFIR on its turntable and we did
 726 not want to confuse the wind-effects between the two instruments. It is clear from Fig.
 727 4(c) that during the supercell passage (0300-0330 UTC) the mean fall speed for 1.3 mm
 728 drops decreases (from 5 to 3.5 m s⁻¹) and the standard deviation increases (from 0.5 to
 729 1.5 m s⁻¹). The histogram shapes also show increasing negative skewness (not shown).
 730 The same trend can be seen for the subsequent squall-line rain cell passage from 0700-
 731 0900 UTC. Similar trends are noted in panels (d) and and less so in panel (e).

732

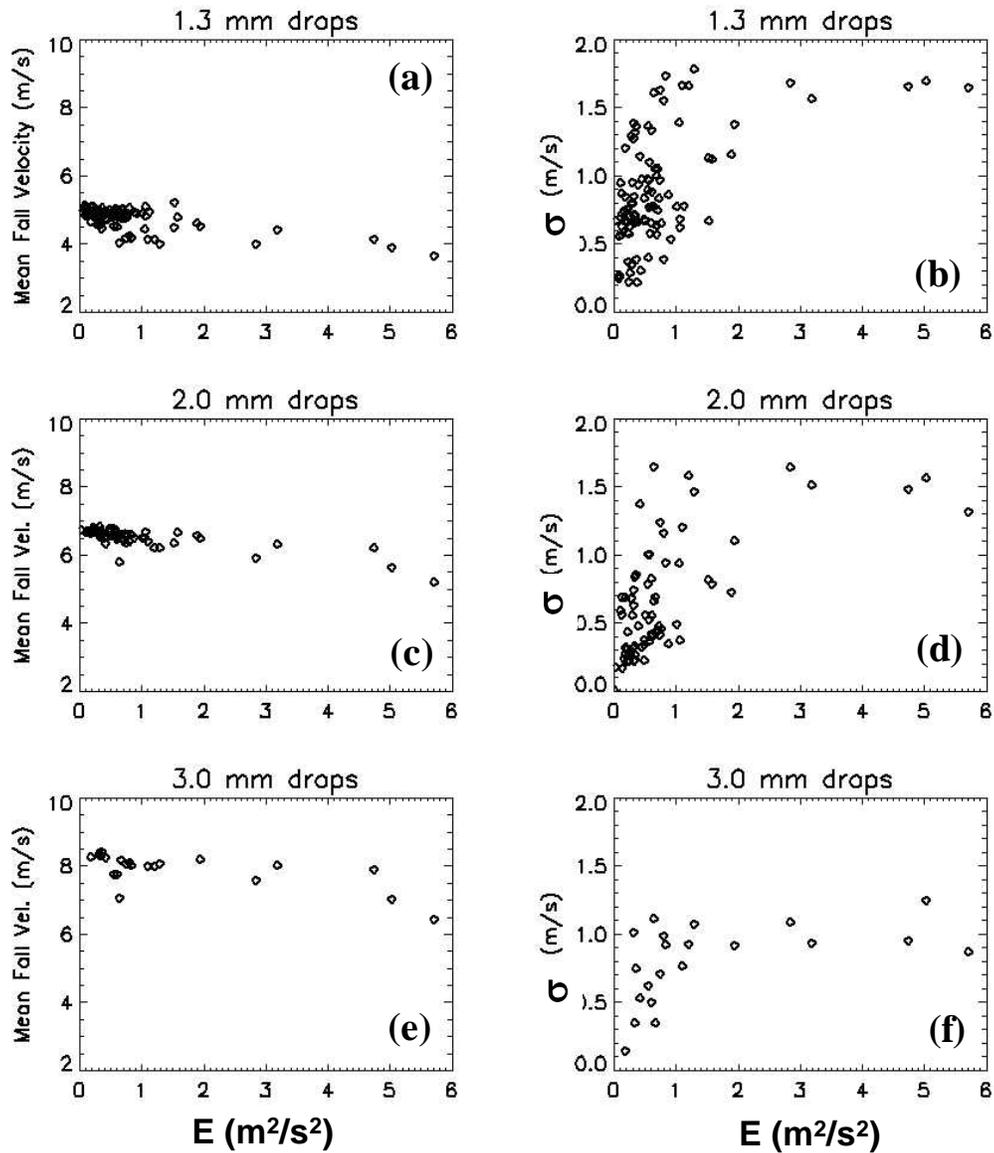


733

734 *Figure 4. (a) as in Fig. 2(a) except for 30 Nov 2016 event. (b) as in Fig. 2(b). (c) mean*
 735 *and $\pm 1\sigma$ standard deviation of fall speeds from 2DVD for 1.3 ± 0.1 mm sizes. (d,e) as in*
 736 *(c) except for 2 ± 0.1 and 3 ± 0.1 mm sizes, respectively.*

737

738 To expand on this observed correlation, Fig. 5 shows scatterplots of the mean fall speed
739 and standard deviation versus E for the 1.3 mm drops (panels a,b), while panels (c,d)
740 and (e,f) show the same but for the 2 and 3 mm drops, respectively. The mean fall
741 speed decreases with increasing E nearly linearly for $E > 1 \text{ m}^2 \text{ s}^{-2}$ but less so for the 3 mm
742 size drops (Stout et al., 1995). This decrease relative to *Gunn-Kinzer* terminal fall
743 speeds is termed as “sub-terminal” and our data is in general agreement with (Montero-
744 Martínez and García-García 2016) who found an increase in the numbers of sub-
745 terminal drops with sizes between 1-2 mm under windy conditions using a 2D-
746 Precipitation probe with resolution of 200 μm (similar to 2DVD) but without a wind fence.
747 The standard deviation of fall speeds (σ_f) versus E is shown in panels 5 (b,d,f). When
748 $E > 1 \text{ m}^2 \text{ s}^{-2}$, the σ_f is nearly constant at 1.5 m s^{-1} for both 1.3 and 2 mm drop sizes and
749 constant at 1 m s^{-1} for the 3 mm size. For $E < 1$, the σ_f is more variable and essentially
750 uncorrelated with E . From the discussion related to Fig. 1(b,c) and 3(b,c), σ_f values
751 exceeding approximately 0.5 m s^{-1} can be attributed to physical, not instrumental or
752 finite bin width effects (see, also, Table 1). Thus, the fall speed distributions are
753 considerably broadened when $E > 1 \text{ m}^2 \text{ s}^{-2}$ due to increasing turbulence levels which is
754 again consistent with the findings of (Montero-Martínez and García-García, 2016) as
755 well as those of (Garrett and Yuter, 2014). The latter observations, however, were of
756 graupel fall speeds in winter precipitation using a multi-angle snowflake camera (Garrett
757 et al., 2012).



758

759

760 *Figure 5.* (a,b) mean fall speed and standard deviation, respectively, versus E for 1.3
 761 mm sizes. (c,d) same but for 2 mm sizes. (e,f) same but for 3 mm.

762

763 3 Discussion and Conclusions

764 We have reported on raindrop fall speed distributions using a high resolution (50 μm)
 765 droplet spectrometer (MPS) collocated with moderate resolution (170 μm) 2DVD (with
 766 both instruments inside a DFIR wind shield) to cover the entire size range (from 0.1 mm
 767 onwards) expected in natural rain. Turbulence intensity (E) was derived from wind/gust

768 data at 10-m height following (Garrett and Yuter, 2014). For low turbulent intensities (E
769 $< 0.4 \text{ m}^2 \text{ s}^{-2}$), in the overlap region of the two instruments (0.7-2 mm), the mean fall
770 speeds were in excellent agreement with each other for both the Greeley, CO and
771 Huntsville, AL sites giving high confidence in the quality of the measurements. For
772 $D < 0.5$ mm and down to 0.1 mm, the mean fall speeds from MPS from both sites were in
773 remarkable agreement with the (Foote and du Toit, 1969) fit to the laboratory data of
774 (Gunn and Kinzer, 1949). In the overlap region, the mean fall speeds from the two
775 instruments were in excellent agreement with the FT fit for the Huntsville site (no
776 altitude adjustment required) and good agreement for the Greeley site (after adjustment
777 for altitude of 1.4 km). For $D > 2$ mm, the mean fall speeds from 2DVD were in excellent
778 agreement with the FT fit at both sites.

779 Our histograms of fall speeds for 1 and 1.5 mm sizes under low turbulence intensity
780 conditions ($E < 0.4 \text{ m}^2 \text{ s}^{-2}$) from both MPS and 2DVD were in good agreement and did
781 not show any evidence of either sub- or super-terminal speeds, rather the histograms
782 were symmetric with mean close to the Gunn-Kinzer terminal velocity with no significant
783 broadening over that ascribed to instrument and/or finite bin width effects. (Note: sub-
784 terminal implies fall speeds < 0.7 times the terminal fall speed whereas super-terminal
785 implies > 1.3 times terminal value; Montero-Martinez et al., 2009). However, for the 0.5
786 and 0.7 mm sizes, from the histogram of fall speeds using the MPS under the same
787 conditions occasional occurrences of both sub- and super-terminal fall speeds, after
788 accounting for instrumental and finite bin width effects, cannot be ruled out.

789 The only comparable earlier study is by (Montero-Martinez et al., 2009) who used
790 collocated 2D-cloud and precipitation probes (2D-C, 2D-P) but restricted their data to
791 calm wind conditions. Their main conclusion was that the distribution of the ratio of the
792 measured fall speed to the terminal fall speed for 0.44 mm size, while having a mode at
793 1 was strongly positively skewed with tails extending to 5 especially at high rain rates. In
794 our data for the 0.5 and 0.7 mm sizes shown in Fig. 1(b,c) and 3(b,c), no such strong
795 positive skewness was observed in the fall speed histograms, and the corresponding
796 ratio of MPS-measured fall speeds to terminal values does not exceed 1.5 to 2.

797 Another study by Larsen et al., (2014) appears to confirm the ubiquitous existence of
798 super-terminal fall speeds for sizes < 1 mm using different instruments one of which
799 was a 2DVD similar to the one used in this study. However, it is well-known that “mis-
800 matched” drops cause erroneous fall speed estimates from 2DVD for drops < 0.5 mm
801 (Schoenhuber et al., 2008; Appendix in Huang et al., 2010; Bernauer et al., 2015). It is
802 not clear if (Larsen et al., 2014) accounted for this problem in their analysis. In addition,
803 their 2DVD was not located within a DFIR-like wind shield.

804 In a later study using only the 2D-P probe, (Montero-Martinez and Garcia-Garcia, 2016)
805 found sub-terminal fall speeds and broadened distributions under windy conditions for

806 1-2 mm sizes in general agreement with our results using the 2DVD. *Stout et al.*, (1995)
807 simulated the motion of drops **subject to non-linear drag** in isotropic turbulence and
808 determined that there would be a significant reduction of the average drop settling
809 velocity (relative to terminal velocity) of greater than 35% for drops around 2 mm size
810 when the ratio of *rms* velocity fluctuations (due to turbulence) relative to drop terminal
811 velocity is around 0.8. Whereas we did not have a direct measure of the *rms* velocity
812 fluctuations, the proxy for turbulence intensity (*E*) related to wind gusts during supercell
813 passage (very large *E* around $7 \text{ m}^2 \text{ s}^{-2}$) **and two squall-line passages (moderate *E***
814 **between $2\text{-}5 \text{ m}^2 \text{ s}^{-2}$) clearly showed a significant reduction in mean fall speeds of 25-**
815 **30% relative to terminal speed for 1.3 and 2 mm sizes (and less so for 3 mm drops),**
816 with significant broadening of the fall speed distributions relative to calm conditions by
817 **nearly a factor of 1.5 to 2.**

818 While our dataset is limited to three events they cover a wide range of rain rates, wind
819 conditions and two different climatologies. **One caveat is that the response of the DFIR**
820 **wind shield to ambient winds in terms of producing subtle vertical air motions near the**
821 **sensor area is yet to be evaluated as future work.** Analysis of further events with direct
822 measurement of turbulent intensity, **for example using a 3D-sonic anemometer at the**
823 **height of the sensor,** would be needed to generalize our findings.

824

825 Data Availability

826 Data used in this paper can be accessed from:

827 ftp://lab.chill.colostate.edu/pub/kennedy/merhala/Bringi_et_al_2017_GRL_datasets/

828

829 Competing interests

830 VNB and MT declare they have no conflict of interest. DB is employed by Droplet
831 Measurements Technologies, Inc. (Longmont, Colorado, USA) who manufacture the
832 Meteorological Particle Spectrometer used in this study.

833 Acknowledgements

834 Two of the authors (VNB and MT) acknowledge support from the U.S. National Science
835 Foundation via grant AGS-1431127. The assistance of Dr. Patrick Gatlin of
836 NASA/MSFC is gratefully acknowledged. Prof. Kevin Knupp and Mr. Carter Hulseley of
837 the University of Alabama in Huntsville processed the wind data.

838

839

840 References

841 Atlas, D., R. C. Srivastava, and R. S. Sekhon (1973). Doppler radar characteristics of
842 precipitation at vertical incidence, *Rev. Geophys.*, *11*, 1–35,
843 doi:<https://doi.org/10.1029/RG011i001p00001>

844 Barthazy, E., S. Göke, R. Schefold, and D. Högl (2004), An optical array instrument for
845 shape and fall velocity measurements of hydrometeors, *J. Atmos. Oceanic*
846 *Technol.*, *21*, 1400–1416, doi:[https://doi.org/10.1175/1520-0426\(2004\)021<1400:AOAIFS>2.0.CO;2](https://doi.org/10.1175/1520-0426(2004)021<1400:AOAIFS>2.0.CO;2).

848 Baumgardner, D., S. Abel, D. Axisa, R. Cotton, J. Crosier, P. Field, C. Gurganus, A.
849 Heymsfield, A. Korolev, M. Krämer, P. Lawson, G. McFarquhar, J. Z Ulanowski, J.
850 Shik Um, 2016: Chapter 9: Cloud Ice Properties - In Situ Measurement
851 Challenges, *AMS Monograph on Ice Formation and Evolution in Clouds and*
852 *Precipitation: Measurement and Modeling Challenges*, Eds. D. Baumgardner, G.
853 McFarquhar, A. Heymsfield, Boston, MA.

854 Beard, K. V. (1976), Terminal velocity and shape of cloud and precipitation drops aloft,
855 *J. Atmos. Sci.*, *33*, 851–864, doi:[https://doi.org/10.1175/1520-0469\(1976\)033<0851:TVASOC>2.0.CO;2](https://doi.org/10.1175/1520-0469(1976)033<0851:TVASOC>2.0.CO;2)

857 Beard, K. V., and H. R. Pruppacher (1969), A determination of the terminal velocity and
858 drag of small water drops by means of a wind tunnel, *J. Atmos. Sci.*, *26*, 1066–
859 1072, doi: [https://doi.org/10.1175/1520-0469\(1969\)026<1066:ADOTTV>2.0.CO;2](https://doi.org/10.1175/1520-0469(1969)026<1066:ADOTTV>2.0.CO;2).

860 Belda, M., E. Holtanová, T. Halenka, and J. Kalvová (2014), Climate classification
861 revisited: From Köppen to Trewartha. *Climate Res.*, *59*, 1–13,
862 doi:<https://doi.org/10.3354/cr01204>.

863 Bernauer, F., K. Hürkamp, W. Rühm, and J. Tschiersch (2015), On the consistency of
864 2-D video disdrometers in measuring microphysical parameters of solid precipitation,
865 *Atmos. Meas. Tech.*, *8*, 3251-3261, doi: <https://doi.org/10.5194/amt-8-3251-2015>

866 Foote, G.B. and P.S. Du Toit (1969), Terminal Velocity of Raindrops Aloft. *J. Appl.*
867 *Meteor.*, *8*, 249–253, [https://doi.org/10.1175/1520-0450\(1969\)008<0249:TVORA>2.0.CO;2](https://doi.org/10.1175/1520-0450(1969)008<0249:TVORA>2.0.CO;2)

869 Garrett, T. J., and S. E. Yuter (2014), Observed influence of riming, temperature, and
870 turbulence on the fallspeed of solid precipitation, *Geophys. Res. Lett.*, *41*, 6515–
871 6522, doi:10.1002/2014GL061016.

872 Garrett, T. J., Fallgatter, C., Shkurko, K., and Howlett, D. (2012), Fall speed
873 measurement and high-resolution multi-angle photography of hydrometeors in free fall,
874 *Atmos. Meas. Tech.*, *5*, 2625-2633, doi:10.5194/amt-5-2625-2012

875 Gunn, R. and G.D. Kinzer (1949), The terminal velocity of fall for water droplets in
876 stagnant air, *J. Meteor.*, *6*, 243–248, [https://doi.org/10.1175/1520-0469\(1949\)006<0243:TTVOFF>2.0.CO;2](https://doi.org/10.1175/1520-0469(1949)006<0243:TTVOFF>2.0.CO;2)

- 878 Huang, G., V.N. Brinji, and M. Thurai (2008), Orientation Angle Distributions of Drops
879 after an 80-m Fall Using a 2D Video Disdrometer, *J. Atmos. Oceanic*
880 *Technol.*, 25,1717–1723, <https://doi.org/10.1175/2008JTECHA1075.1>
- 881 Huang, G., V.N. Brinji, R. Cifelli, D. Hudak, and W.A. Petersen (2010), A Methodology
882 to Derive Radar Reflectivity–Liquid Equivalent Snow Rate Relations Using C-Band
883 Radar and a 2D Video Disdrometer, *J. Atmos. Oceanic Technol.*, 27, 637–
884 651, <https://doi.org/10.1175/2009JTECHA1284.1>
- 885 Joe, P., and R. List (1987), Testing and performance of two-dimensional optical array
886 spectrometer with grey scale, *J. Atmos. Oceanic Technol.*, 4, 139-150,
887 [https://doi.org/10.1175/1520-0426\(1987\)004<0139:TAPOTD>2.0.CO;2](https://doi.org/10.1175/1520-0426(1987)004<0139:TAPOTD>2.0.CO;2)
888
- 889 Knollenberg, R. (1970), The Optical Array: An alternative to scattering or extinction for
890 airborne particle size determination, *J. Appl. Meteorol.*, 9, 86–103,
891 [doi:https://doi.org/10.1175/1520-0450\(1970\)009<0086:TOAAAT>2.0.CO;2](https://doi.org/10.1175/1520-0450(1970)009<0086:TOAAAT>2.0.CO;2).
892
- 893 Knollenberg, R. (1976), Three new instruments for cloud physics measurements: The
894 2–D spectrometer probe, the forward scattering spectrometer probe and the active
895 scattering aerosol spectrometer, in International Conference on Cloud Physics,
896 American Meteorological Society, 554–561.,
897
- 898 Knollenberg, R. (1981), Clouds, Their Formation, Optical Properties and Effects, chap.
899 Techniques for probing cloud microstructure, Academic Press, edited by P.V.
900 Hobbs and A. Deepak. 495 pp
- 901 Korolev, A. V., Kuznetsov, S. V., Makarov, Y. E., and Novikov, V. S. (1991), Evaluation
902 of Measurements of Particle Size and Sample Area from Optical Array Probes, *J.*
903 *Atmos. Oceanic Tech.*, 8, 514–522, [https://doi.org/10.1175/1520-
904 0426\(1991\)008<0514:EOMOPS>2.0.CO;2](https://doi.org/10.1175/1520-0426(1991)008<0514:EOMOPS>2.0.CO;2)
905
- 906 Korolev, A. V., J. W. Strapp, and G. A. Isaac (1998), Evaluation of the Accuracy of
907 PMS Optical Array Probes, *Journal of Atmospheric and Oceanic Technology*, 15,
908 708–720, [https://doi.org/10.1175/1520-0426\(1998\)015<0708:EOTAOP>2.0.CO;2](https://doi.org/10.1175/1520-0426(1998)015<0708:EOTAOP>2.0.CO;2)
909
- 910 Korolev, A. V. (2007), Reconstruction of the Sizes of Spherical Particles from Their
911 Shadow Images. Part I: Theoretical Considerations, *Journal of Atmospheric and*
912 *Oceanic Technology*, 24, 376–389, <https://doi.org/10.1175/JTECH1980.1>
- 913 Larsen, M.L., A.B. Kostinski, and A.R. Jameson (2014), Further evidence for super
914 terminal drops, *Geophysical Research Letters* , 41, 2014GL061397.
- 915 List, R., N.R. Donaldson, and R.E. Stewart (1987), Temporal Evolution of Drop Spectra
916 to Collisional Equilibrium in Steady and Pulsating Rain, *J. Atmos. Sci.*, 44, 362–
917 372, [https://doi.org/10.1175/1520-0469\(1987\)044<0362:TEODST>2.0.CO;2](https://doi.org/10.1175/1520-0469(1987)044<0362:TEODST>2.0.CO;2)

- 918 Löffler-Mang, M. and J. Joss (2000), An Optical Disdrometer for Measuring Size and
 919 Velocity of Hydrometeors, *J. Atmos. Oceanic Technol.*, 17, 130–
 920 139, [https://doi.org/10.1175/1520-0426\(2000\)017<0130:AODFMS>2.0.CO;2](https://doi.org/10.1175/1520-0426(2000)017<0130:AODFMS>2.0.CO;2)
- 921 Montero-Martinez, G. and F. Garcia-Garcia (2016), On the behavior of raindrop fall
 922 speed due to wind, *Quarterly Journal of the Royal Meteorological Society*, 142,
 923 2794, DOI: 10.1002/qj.2794.
- 924 Montero-Martinez, G., A. B. Kostinski, R. A. Shaw, and F. Garcia-Garcia (2009), Do all
 925 raindrops fall at terminal speed?, *Geophysical Research Letters*, 36, 246 L11818,
 926 doi:10.1029/2008GL037111.
- 927 Notaroš, B., V. N. Bringi, C. Kleinkort, P. Kennedy, G-J Huang, M. Thurai, A. J.
 928 Newman, W. Bang and G. Lee (2016), Accurate Characterization of Winter
 929 Precipitation Using Multi-Angle Snowflake Camera, Visual Hull, Advanced
 930 Scattering Methods and Polarimetric Radar, *Atmosphere*, 7(6), 81;
 931 doi:10.3390/atmos7060081
- 932 Rasmussen, R., B. Baker, J. Kochendorfer, T. Meyers, S. Landolt, A.P. Fischer, J.
 933 Black, J.M. Thériault, P. Kucera, D. Gochis, C. Smith, R. Nitu, M. Hall, K. Ikeda,
 934 and E. Gutmann (2012). How Well Are We Measuring Snow: The
 935 NOAA/FAA/NCAR Winter Precipitation Test Bed, *Bull. Amer. Meteor.*
 936 *Soc.*, 93, 811–829, <https://doi.org/10.1175/BAMS-D-11-00052.1>
- 937 Rosewell, C.J. (1986), Rainfall Kinetic Energy in Eastern Australia, *J. Climate Appl.*
 938 *Meteor.*, 25, 1695–1701, [https://doi.org/10.1175/1520-0450\(1986\)025<1695:RKEIEA>2.0.CO;2](https://doi.org/10.1175/1520-0450(1986)025<1695:RKEIEA>2.0.CO;2)
- 940 Schönhuber, M., Günter Lammer, Randeu W.L. (2007), *Advances in Geosciences*, 10,
 941 pp. 85-90, 2007, <https://doi.org/10.5194/adgeo-10-85-2007>.
- 942 Schönhuber, M., G. Lammar, and W. L. Randeu (2008), “The 2D-video-distrometer”,
 943 Chapter 1 in “Precipitation: Advances in Measurement, Estimation and Prediction”,
 944 S. C. Michaelides, Ed., Springer, 3–31.
- 945 Sekhon, R.S. and R.C. Srivastava (1971), Doppler Radar Observations of Drop-Size
 946 Distributions in a Thunderstorm, *J. Atmos. Sci.*, 28, 983–994,
 947 [https://doi.org/10.1175/1520-0469\(1971\)028<0983:DROODS>2.0.CO;2](https://doi.org/10.1175/1520-0469(1971)028<0983:DROODS>2.0.CO;2)
- 948 Stout, J.E., S.P. Arya, and E.L. Genikhovich (1995). The Effect of Nonlinear Drag on the
 949 Motion and Settling Velocity of Heavy Particles, *J. Atmos. Sci.*, 52, 3836–3848,
 950 [https://doi.org/10.1175/1520-0469\(1995\)052<3836:TEONDO>2.0.CO;2](https://doi.org/10.1175/1520-0469(1995)052<3836:TEONDO>2.0.CO;2)
- 951 Testik, F.Y. and M.K. Rahman (2016), High-Speed Optical Disdrometer for Rainfall
 952 Microphysical Observations, *J. Atmos. Oceanic Technol.*, 33, 231–243,
 953 <https://doi.org/10.1175/JTECH-D-15-0098.1>
- 954 Thériault, J.M., R. Rasmussen, E. Petro, J. Trépanier, M. Colli, and L.G. Lanza
 955 (2015), Impact of Wind Direction, Wind Speed, and Particle Characteristics on the

- 956 Collection Efficiency of the Double Fence Intercomparison Reference, *J. Appl.*
957 *Meteor. Climatol.*, 54, 1918–1930, <https://doi.org/10.1175/JAMC-D-15-0034.1>
- 958 Thurai, M. and V. N. Bringi (2005), Drop axis ratios from a 2D video disdrometer, *J.*
959 *Atmos. Oceanic Technol.*, 22, 966–978, <https://doi.org/10.1175/JTECH1767.1>
- 960 M. Thurai, V. N. Bringi, M. Szakáll, S. K. Mitra, K. V. Beard, and S. Borrmann (2009),
961 Drop Shapes and Axis Ratio Distributions: Comparison between 2D Video
962 Disdrometer and Wind-Tunnel Measurements, *Journal of Atmospheric and*
963 *Oceanic Technology*, 26 (7), 1427–1432,
964 <https://doi.org/10.1175/2009JTECHA1244.1>
- 965 Thurai, M., V. N. Bringi, W. A. Petersen, P. N. Gatlin (2013), Drop shapes and fall
966 speeds in rain: two contrasting examples, *J. Appl. Meteor. Climatol.*, 52,
967 2567–2581, <https://doi.org/10.1175/JAMC-D-12-085.1>
- 968 Thurai, M., P. Gatlin, V. N. Bringi, W. Petersen, P. Kennedy, B. Notaroš, and L. Carey
969 (2017), Toward completing the raindrops size spectrum: Case studies involving
970 2D-video disdrometer, droplet spectrometer, and polarimetric radar measurements,
971 *J. Appl. Meteor. Climatol.*, 56, 877–896, <https://doi.org/10.1175/JAMC-D-16-0304.1>
- 972 Trewartha, G. T., and L. H. Horn (1980), *Introduction to Climate*, 5th ed. McGraw Hill,
973 416 pp.
- 974 Yu, C.-K., Hsieh, P.-R., Yuter, S. E., Cheng, L.-W., Tsai, C.-L., Lin, C.-Y., and Chen, Y.:
975 Measuring droplet fall speed with a high-speed camera: indoor accuracy and
976 potential outdoor applications, *Atmos. Meas. Tech.*, 9, 1755-1766,
977 <https://doi.org/10.5194/amt-9-1755-2016>, 2016.
- 978