

Response to Reviewer #2 Comments

We thank the reviewer for their helpful comments. We have incorporated as many of the reviewers' suggestions as possible into the revised manuscript. All reviewer comments are in italics and the author's responses are in standard font.

In this paper, the authors compare ozone profile data from three TOLNet stations across the USA with A) an ozone climatology and ozone data from various transport models, and B) a simulated retrieval result where the climatology and models are used as a priori values.

The authors use a formula from the book by Rodgers (2000) to linearise the calculations of the effect of the a priori on a potential retrieval. While Rodgers uses the formula in Chap 3 and Chap 10 of his book, using this formula to make a selection on a preferred a priori brushes over the potential issues you often get with real satellite data.

The first question that comes to mind is: how representative are these simulated retrievals for real world situations. Or is Eq 1 limited to be used for an error / sensitivity study? My impression is that the error component is not used in the paper. Please give your reasons for using this method.

The reviewer is correct in fact that the application of Eq. (1) in our study is representative of a sensitivity study to determine the potential impact of a priori ozone (O_3) profiles on TEMPO retrievals in the troposphere. The actual "real-world" TEMPO O_3 retrieval algorithm will be non-linear and iterative (Liu et al., 2010). However, the linear approach used in our study has been shown in numerous studies as a good first-order approximation of satellite retrievals of O_3 profiles (e.g., Bowman et al., 2002; Worden et al., 2007; Kulawik et al., 2006, 2008; Natraj et al., 2011; Zoogman et al., 2014). The reviewer is also correct that we do not include random retrieval errors (ϵ) in Eq. (1), however, measurement random-noise error covariance and a priori covariance matrix are included in the calculation of the averaging kernels (AKs) used during this study. Additional text has been added to Sect. 2.2 of the updated manuscript to state these points.

In retrievals of ozone profiles, an a priori consists of a profile shape and an associated error profile. Because the retrieval of ozone profiles is under-determined (more than one profile shape can be retrieved from the same spectrum), an a priori is used in an Optimal Estimation (OE) based retrieval to constrain the outcome to reasonable values. The a priori profile shape is a reference, and the profile error gives the retrieval the freedom to differ from that reference shape based on the input spectrum to minimize the cost function.

In a real retrieval, when either the error on the a priori is set to zero, or the error on the measured/simulated spectrum is too large then the OE retrieval result will reproduce the a priori almost exactly. In this case no information is gained from the spectrum during the retrieval. In other words: the spectrum contains no useful information and the degrees of freedom from the signal (DFS) will be low.

We agree with the reviewer that the a priori and measurement error are important aspects when calculating retrieval sensitivity. The calculation of the AKs used in this study are described in Zoogman et al. (2017) and the a priori profile and associated error are derived from the TB-Clim

product. Overall, a priori and measurement error terms are taken into account during the calculation of TEMPO AKs applied in Eq. (1).

The authors seem to come to the conclusion that TEMPO ozone profile retrievals in the troposphere and LMT require an a priori that already matches the general shape of the observations in order for the required accuracy to be obtained in the retrieval. If the a priori already needs to be so close to the shape and magnitude of the outcome of the retrieval, then one could conclude that the TEMPO spectra do not contain sufficient information for the retrieval, or the retrieval is over-constrained.

How do the authors see these issues, in light of the need of their conclusion that the a priori needs to be close to the true profile? Please clarify.

Another way of looking at it is by looking at Eq 1. If the a priori X_a closely matches the true X_t , then what is 'retrieved' is mainly the a priori, as the second term in the equation falls to zero. It is therefore not surprising that an a priori that more closely matches the true profile will also do well in the simulated retrievals. Those a priori profiles already have the advantage in Eq 1. How does this advantage play out with real retrievals? Is it really necessary to have an a priori so close to the true state of the atmosphere to get a good retrieval? If so, what is the added value of a retrieval in this case?

The results of this study showing that more accurate a priori trace gas profile assumptions lead to more accurate satellite retrievals in the troposphere are not surprising/novel. The sensitivity of satellite trace gas retrievals to a priori profiles has been clearly stated and demonstrated in numerous studies (e.g., Martin et al., 2002; Luo et al., 2007; Kulawik et al., 2008; Zhang et al., 2010, Bak et al., 2013; many others). These studies, in addition to many others, show that in vertical extents of the atmosphere where satellite sensitivity is low (i.e., middle to lower troposphere for satellites retrieving O₃) the retrieved state will be highly dependent on the vertical shape of the a priori. However, our study suggests that the magnitude of the tropospheric-average column O₃ abundance will be accurately retrieved by TEMPO regardless of the a priori. This suggests that the magnitude of tropospheric O₃ will be largely controlled by the retrieval. The shape of the a priori itself will have a large impact on the shape of the retrieved tropospheric O₃ profile and therefore lowermost tropospheric (LMT) O₃ magnitudes where satellite sensitivity is low.

The importance of our study is focusing on TEMPO tropospheric O₃ retrievals, which due to the system design (geostationary orbit and UV+VIS wavelength retrievals will provide observations with high spatio-temporal variability with increased sensitivity to lower tropospheric O₃) will for the first time provide air quality relevant space-borne information. Since TEMPO tropospheric profile O₃ data is expected to be assimilated into chemical transport (CTM) and air quality models and LMT data will be used for air quality and event-specific monitoring/research, it is critical to understand methods to improve the quality/accuracy of this retrieved information. Our study demonstrates that to produce TEMPO retrievals of O₃ in the LMT with increased accuracy it is necessary to have accurate a priori profile shape assumption. The results from our study also indicate that of all the potential sources of a priori O₃ profile data which can be used in satellite retrievals evaluated during this work (climatology data products (e.g., TB-Clim), near-real-time

data assimilation models (e.g., GEOS-5 FP), reanalysis models (e.g., MERRA2), or CTM predictions (e.g., GEOS-Chem), CTM simulated data result in the most accurate retrievals.

Textual/other remarks:

Line 104: You mention an error margin of the TOLNet measurements of 10% in the lower troposphere and 20% in the upper troposphere. The words 'lower' and 'upper' are not defined in this context, while you use the terms LMT (0-2km) and tropospheric (0-10km). Please be more specific about the applicable altitude ranges of the errors of the TOLNet DIAL lasers.

We thank the reviewer for this comment as it has led to conversations resulting in an updated and improved statement of TOLNet data uncertainty. The uncertainty of TOLNet O₃ retrievals is dependent on numerous factors such as individual instrument specifications, vertical and temporal integration/averaging methods, sampling environment characteristics, etc. Since the TOLNet measurement data used in this study are hourly-averaged and all below 10 km above ground level the updated manuscript has been revised to read: “Uncertainty in TOLNet O₃ measurements due to systematic error will be approximately 4-5% for all instruments at almost all altitudes. Precision will vary from 0% to > 20% and is dependent on individual instrument characteristics, time of day, and temporal and vertical averaging (precision typically degrades with height for altitudes above 8-10 km) (Kuang et al., 2013; Sullivan et al., 2015b; Leblanc et al., 2016). Since TOLNet observations used during this study are hourly-averaged and below 10 km above ground-level, overall uncertainty can be assumed to be $\leq 10\%$.”

In sect 2.2/2.2.1 it would be helpful to have a little more information on the input data. Please elaborate on the setup you use to generate the artificial/simulated TEMPO data (the AK's, the Gain matrices, etc). What other relevant sources of information did you use, like temperature, albedo's, cross sections, solar and viewing angles, reference spectra, etc.

As stated in the manuscript: “The UV+VIS AKs applied during this study are based on TEMPO retrieval sensitivity studies that play a key role in determining the instrument requirements and verification of the retrieval performance (Zoogman et al., 2017).” In Sect. 7.3 of Zoogman et al. (2017) information is provided about the GEOS-5 meteorological data and GEOS-Chem modeled trace gases and aerosols used to calculate AK values. Viewing geometry, radiance spectra, and weighting functions with respect to aerosols and trace gases are all calculated based on TEMPO specifications as described in Zoogman et al. (2017). Surface albedo values are from the Global Ozone Monitoring Experiment (GOME) albedo database. As mentioned earlier, TB-Clim climatological a prior mean and error covariance matrixes are used in the calculation of TEMPO AKs. To better emphasize the information regarding the AKs that are used during this work that is provided in Zoogman et al. (2017), the following text has been added to Sect. 2.2 of the updated manuscript: “For detailed information about the TEMPO retrieval sensitivity studies, and the input variables, used to derive AKs used during this study see Zoogman et al. (2017).”

In section 2.2 the authors mention the use adaptation of the SAO retrieval algorithm for TEMPO to do retrievals. But it is not clear to me whether the SAO algorithm played a role in this paper at all. In the second part of 2.2 a simple vector/matrix based formula is used to calculate the simulated retrieved profile. Did the authors use the SAO model for any of the ozone profile

retrievals or was it used in the set-up of the kernels? If it was not used, is it then relevant to for this paper?

The manuscript states “TEMPO will adapt the current SAO OMI UV-only O₃ profile algorithm (Liu et al., 2010) to derive O₃ profiles from joint UV+VIS measurements based on the optimal estimation technique.” to provide an explanation of the TEMPO retrieval algorithm. The SAO algorithm is not used to calculate simulated O₃ profile retrievals in this study and are instead approximated using Eq. (1). Please see the above comments which better describe how the AKs used during this study are derived.

Line 141: In Eq 1, there is a component for the effect of noise. Please explain how you treat the last term in the equation. How does this component affect the retrievals and what are the expectations on its effect on the ranking of the a priori sources used?

Please see our earlier response that we do not include random retrieval errors (ϵ) in Eq. (1). This component will add noise to the linear retrievals. Neglecting this will not affect the rankings of the a priori sources. Additional text has been added to Sect. 2.2 of the updated manuscript to clarify this.

Line 168 and Fig 3: Yellow is a color that is hard to see on a white background. Please use a color with more contrast.

The yellow line in Fig. 3 has been changed to green in the updated manuscript.

Line 193: 'due to data constraints'. What kind of data constraints? Is it an issue of lack of sensitivity at the lower troposphere of most existing satellite instruments? Please clarify.

Both GEOS-5 FP and MERRA2 O₃ vertical profiles are primarily driven by the assimilation of Ozone Monitoring Instrument (OMI) and Microwave Limb Sounder (MLS) satellite data. The reviewer is correct in the fact that these satellite products have limited sensitivity in the lower troposphere, and therefore the O₃ values from GEOS-5 FP and MERRA2 are most trusted in the upper troposphere and stratosphere. This section of the updated manuscript now reads: “Both GEOS-5 FP and MERRA2 O₃ vertical profiles are driven by the assimilation of OMI and Microwave Limb Sounder (MLS) satellite data. Predictions of O₃ from these products are most trusted in the upper troposphere and stratosphere due to OMI and MLS having limited sensitivity in the lower troposphere”.

Line 244: In this section you evaluate the straight model output with the TOLNet profiles, outside the context of use as an a priori. The remark that GEOS-Chem is the 'the only potential source of a priori profiles ...' is out of place here. You address the use of the various models as an a priori in sections 3.2.x.

This has been corrected in the updated manuscript.

In lines 248 and 249 the authors give a few aspects that may be the reasons why GEOS-Chem compares better to TOLNet than the other models. It would be insightful to the reader to learn which of these aspects contributes the most to the better comparison.

It would be difficult, and outside the scope of this study, to determine the single reason, out of many, why CTM predictions from GEOS-Chem compare better to O₃ observations compared to other data sources evaluated during this study. However, we present the main reasons why one would expect a CTM to predict O₃ more accurately compared to GEOS-5 FP, MERRA2, and a climatology product and they are “data-assimilated meteorological fields, comprehensive atmospheric chemistry mechanisms, and state-of-the-art trace gas and aerosol emissions data”. We describe in the manuscript that GEOS-5 FP and MERRA2 O₃ predictions do not take into account complex atmospheric chemistry routines or emission inventories. Since O₃ is a highly reactive trace gas in the troposphere, which has numerous emission sources and production/loss processes, these chemistry routines and emission inventories are necessary to accurately replicate O₃ measured in nature.

Section 3.1.2: In this section the authors make an evaluation of how well the climatology and the models can reproduce the daily variability of the lidar measurements. Please elaborate on the time step/time resolution of the models. Is there a reasonable expectation that the models can actually follow the daily cycle, or are the climatology and model fields spaced to far apart in time?

In Sect. 2.2.2 the TB-Clim product is described to provide monthly-mean O₃ profiles and in Sect. 2.3 the temporal resolution of GEOS-5 FP and MERRA2 are stated to be 3 hours and 10-minute in GEOS-Chem. In Sect. 2.4 it is stated that all measured, modeled, and climatology products are averaged or interpolated to match an hourly temporal resolution for evaluation. The monthly-mean nature of TB-Clim is one of the main reasons why it is unable to replicate the daily and diurnal variability of observed tropospheric O₃. The modeled products all have temporal variability of ≤ 3 hours and therefore have the capability to capture the diurnal variability of O₃. However, tropospheric O₃ mixing ratios are highly dependent on the diurnal variability of emissions, deposition, and atmospheric chemistry and therefore would be expected to be best replicated from a CTM.

Please consider enlarging your time series plots.

This has been done to the best of our ability.

References

- Bak, J., Liu, X., Wei, J. C., Pan, L. L., Chance, K., and Kim, J. H.: Improvement of OMI ozone profile retrievals in the upper troposphere and lower stratosphere by the use of a tropopause-based ozone profile climatology, *Atmos. Meas. Tech.*, 6, 2239-2254, <https://doi.org/10.5194/amt-6-2239-2013>, 2013.
- Bowman, K. W., Worden, J., Steck, T., Worden, H. M., Clough, S. and Rodgers, C.: Capturing time and vertical variability of tropospheric ozone: A study using TES nadir retrievals, *J. Geophys. Res.-Atmos.*, 107(D23), 4723, doi:10.1029/2002JD002150, 2002.

- Granados-Muñoz, M. J. and Leblanc, T.: Tropospheric ozone seasonal and long-term variability as seen by lidar and surface measurements at the JPL-Table Mountain Facility, California, *Atmos. Chem. Phys.*, 16, 9299-9319, doi:10.5194/acp-16-9299-2016, 2016.
- Kuang, S., Newchurch, M. J., Burris, J., and Liu, X.: Ground-based lidar for atmospheric boundary layer ozone measurements, *Appl. Opt.*, 52, 3557-3566, <https://doi.org/10.1364/AO.52.003557>, 2013.
- Kulawik, S. S., Worden, H., Osterman, G., Luo, M., Beer, R., Kinnison, D. E., Bowman, K. W., Worden, J., Eldering, A., Lampel, M., Steck, T., and Rodgers, C. D.: TES atmospheric profile retrieval characterization: An orbit of simulated observations, *IEEE T. Geosci. Remote*, 44, 1324-1333, 2006.
- Kulawik, S. S., Bowman, K.W., Luo, M., Rodgers, C. D., and Jourdain, L.: Impact of nonlinearity on changing the a priori of trace gas profile estimates from the Tropospheric Emission Spectrometer (TES), *Atmos. Chem. Phys.*, 8, 3081–3092, 2008, <http://www.atmos-chem-phys.net/8/3081/2008/>.
- Liu, X., Bhartia, P. K., Chance, K., Spurr, R. J. D., and Kurosu, T. P.: Ozone profile retrievals from the Ozone Monitoring Instrument, *Atmos. Chem. Phys.*, 10, 2521-2537, doi:10.5194/acp-10-2521-2010, 2010.
- Martin, R. V., Chance, K., Jacob, D. J., et al.: An improved retrieval of tropospheric nitrogen dioxide from GOME, *J. Geophys. Res.*, 107, 4437, doi:10.1029/2001JD001027, 2002.
- Natraj, V., Liu, X., Kulawik, S., Chance, K., Chatfield, R., Edwards, D. P., Eldering, A., Francis, G., Kurosu, T., Pickering, K., Spurr, R., and Worden, H.: Multi-spectral sensitivity studies for the retrieval of tropospheric and lowermost tropospheric ozone from simulated clear-sky GEO-CAPE measurements, *Atmos. Environ.*, 45, 7151-7165, 2011.
- Sullivan, J. T., McGee, T. J., Leblanc, T., Sumnicht, G. K., and Twigg, L. W.: Optimization of the GSFC TROPOZ DIAL retrieval using synthetic lidar returns and ozonesondes – Part 1: Algorithm validation, *Atmos. Meas. Tech.*, 8, 4133-4143, doi:10.5194/amt-8-4133-2015, 2015b.
- Worden, J., Liu, X., Bowman, K., Chance, K., Beer, R., Eldering, A., Gunson, M., and Worden, H. M.: Improved tropospheric ozone profile retrievals using OMI and TES radiances, *Geophys. Res. Lett.*, 34, L01809, doi:10.1029/2006GL027806, 2007.
- Zhang, L., Jacob, D. J., Liu, X., Logan, J. A., Chance, K., Eldering, A., and Bojkov, B. R.: Intercomparison methods for satellite measurements of atmospheric composition: application to tropospheric ozone from TES and OMI, *Atmos. Chem. Phys.*, 10, 4725-4739, <https://doi.org/10.5194/acp-10-4725-2010>, 2010.
- Zoogman, P., Jacob, D. J., Chance, K., Liu, X., Lin, M., Fiore, A., and Travis, K.: Monitoring high-ozone events in the US Intermountain West using TEMPO geostationary satellite

observations, *Atmos. Chem. Phys.*, 14, 6261-6271, <https://doi.org/10.5194/acp-14-6261-2014>, 2014.

Zoogman, P., Liu, X., Suleiman, R., Pennington, W., Flittner, D., Al-Saadi, J., Hilton, B., Nicks, D., Newchurch, M., Carr, J., Janz, S., Andraschko, M., Arola, A., Baker, B., Canova, B., Miller, C. C., Cohen, R., Davis, J., Dussault, M., Edwards, D., Fishman, J., Ghulam, A., Abad, G. G., Grutter, M., Herman, J., Houck, J., Jacob, D., Joiner, J., Kerridge, B., Kim, J., Krotkov, N., Lamsal, L., Li, C., Lindfors, A., Martin, R., McElroy, C., McLinden, C., Natraj, V., Neil, D., Nowlan, C., O'Sullivan, E., Palmer, P., Pierce, R., Pippin, M., Saiz-Lopez, A., Spurr, R., Szykman, J., Torres, O., Veefkind, J., Veihelmann, B., Wang, H., Wang, J., and Chance, K.: Tropospheric emissions: Monitoring of pollution (TEMPO), *J. Quant. Spectrosc. Ra.*, 186, 17-39, <https://doi.org/10.1016/j.jqsrt.2016.05.008>, 2017.