



1 Potential sources of a priori ozone profiles for TEMPO 2 tropospheric ozone retrievals

3 Matthew S. Johnson¹, Xiong Liu², Peter Zoogman^{2,*}, John Sullivan³, Michael J. Newchurch⁴, Shi
4 Kuang⁵, Thierry Leblanc⁶, Thomas McGee³

5 ¹Earth Science Division, NASA Ames Research Center, Moffett Field, CA, USA.

6 ²Harvard-Smithsonian Center for Astrophysics, Cambridge, MA, USA.

7 ³Atmospheric Chemistry and Dynamics Laboratory, NASA Goddard Space Flight Center, Greenbelt, Maryland, USA.

8 ⁴Atmospheric Science Department, University of Alabama in Huntsville, Huntsville, AL, USA.

9 ⁵Earth System Science Center, University of Alabama in Huntsville, Huntsville, AL, USA.

10 ⁶Table Mountain Facility, California Institute of Technology, Wrightwood, CA, USA.

11 *also at Minerva Schools at KGI, San Francisco, CA, USA.

12 *Correspondence to:* Matthew S. Johnson (matthew.s.johnson@nasa.gov)

13 **Abstract.** Potential sources of a priori ozone (O₃) profiles for use in Tropospheric Emissions: Monitoring of Pollution
14 (TEMPO) satellite tropospheric O₃ retrievals are evaluated with observations from multiple Tropospheric Ozone Lidar
15 Network (TOLNet) systems in North America. An O₃ profile climatology (tropopause-based O₃ climatology (TB-
16 Clim), currently proposed for use in TEMPO O₃ retrieval algorithms) based on ozonesonde observations and O₃
17 profiles from three separate models (operational Goddard Earth Observing System (GEOS-5) Forward Processing
18 (FP) product, reanalysis product from Modern-Era Retrospective analysis for Research and Applications version 2
19 (MERRA2), and the GEOS-Chem chemical transport model (CTM)) were: 1) evaluated with TOLNet measurements
20 on various temporal scales (seasonally, daily, hourly) and 2) implemented as a priori information in theoretical
21 TEMPO tropospheric O₃ retrievals in order to determine how each a priori impacts the accuracy of retrieved
22 tropospheric (0-10 km) and lowermost tropospheric (LMT, 0-2 km) O₃ columns. We found that all potential sources
23 of a priori profiles evaluated in this study generally reproduced the vertical structure of summer-averaged observations
24 of O₃ profiles. However, larger differences between the a priori profiles and lidar observations were observed when
25 evaluating inter-daily and diurnal variability of tropospheric O₃. The TB-Clim O₃ profile climatology was unable to
26 replicate observed inter-daily and diurnal variability of O₃ while model products, in particular GEOS-Chem
27 simulations, displayed more skill in reproducing these features. Due to the ability of models, primarily the CTM used
28 in this study, on average to capture the inter-daily and diurnal variability of tropospheric and LMT O₃ columns, using
29 a priori profiles from these model simulations resulted in TEMPO retrievals with the best statistical comparison with
30 lidar observations. Furthermore, important from an air quality perspective, when high LMT O₃ values are observed,
31 using GEOS-Chem a priori profiles resulted in TEMPO LMT O₃ retrievals with the least bias.

32 1 Introduction

33 Ozone (O₃) is an important atmospheric constituent for air quality as concentrations above natural levels can have
34 detrimental health impacts (US EPA, 2006) and the United States (US) Environmental Protection Agency (EPA)
35 enforces surface-level mixing ratios under the National Ambient Air Quality Standards (NAAQS). In 2015, the



36 NAAQS for O₃ was reduced from prior levels of 75 parts per billion (ppb) to 70 ppb, requiring that 3-year averages
37 of the annual fourth-highest daily maximum 8-hour mean mixing ratio must be ≤ 70 ppb (US EPA, 2015).
38 Tropospheric and surface-level O₃ mixing ratios are controlled by a complex system of photo-chemical reactions
39 involving numerous trace gas species (e.g., carbon monoxide (CO), methane, volatile organic compounds, and
40 nitrogen oxides (NO_x = nitric oxide and nitrogen dioxide (NO + NO₂)) emitted from anthropogenic and natural sources
41 (Atkinson, 1990; Lelieveld and Dentener, 2000). Furthermore, a portion of tropospheric O₃ is also contributed from
42 the downward transport from the stratosphere, commonly referred to as stratosphere-to-troposphere exchange (STE)
43 (e.g., Stohl et al., 2003; Lin et al., 2015). Due to the complex chemistry and transport processes controlling O₃ mixing
44 ratios, and the continued reduction of NAAQS levels, it is increasingly important to improve the ability to
45 monitor/study tropospheric and surface-level O₃.

46 The monitoring of air quality in North America is typically conducted by using ground-based in situ
47 measurement networks. However, in recent years, observations of tropospheric O₃ and precursor gases (e.g., CO, NO₂,
48 formaldehyde (HCHO)) have been made from space-borne platforms which have led to the better understanding of
49 the tropospheric O₃ budget (Sauvage et al., 2007; Martin, 2008; Duncan et al., 2014). Total column (stratosphere +
50 troposphere) O₃ has been routinely measured by numerous space-based sensors since the launch of the Total Ozone
51 Mapping Spectrometer (TOMS) in 1978. Tropospheric column O₃ has been derived from total column retrievals using
52 strategies such as residual-based approaches which subtract the stratospheric column O₃ from total O₃ (Fishman et al.,
53 2008 and references therein). Tropospheric O₃ profiles have also been directly retrieved from hyperspectral Ultraviolet
54 (UV) (e.g., Liu et al., 2005, 2010) and Thermal Infrared (TIR) (e.g., Bowman et al., 2006) measurements. Currently,
55 sensors measuring tropospheric O₃, such as those using UV measurements from the Ozone Monitoring Instrument
56 (OMI) and TIR measurements from the Tropospheric Emission Spectrometer (TES) (Beer, 2006), are from low earth
57 orbit (LEO). While LEO provides global coverage, the observation of tropospheric O₃ is limited by coarse spatial
58 resolution, limited temporal frequency (once or twice per day), and inadequate sensitivity to lower tropospheric and
59 planetary boundary layer (PBL) O₃ (Fishman et al., 2008; Natraj et al., 2011). These limitations restrict the ability to
60 apply these space-borne observations in air quality policy and monitoring.

61 The Tropospheric Emissions: Monitoring of Pollution (TEMPO) satellite, which will be launched between
62 2019-2021 to geostationary orbit (GEO), is designed to address some of the limitations of current O₃ remote-sensing
63 instruments (Chance et al., 2013; Zoogman et al., 2017). TEMPO will provide critical measurements such as vertical
64 profiles of O₃, total column O₃, NO₂, sulfur dioxide, HCHO, glyoxal, and aerosol/cloud parameters over North
65 America. These data products will be provided hourly at a native spatial resolution of $\sim 2.1 \times 4.4$ km² (at the center of
66 the field of regard) except at the required spatial resolution of 8.4×4.4 km² for the O₃ profile product (four pixels
67 combined to increase signal to noise ratios and reduce computational resources). TEMPO's domain will encompass
68 the region of North America from Mexico City to the Canadian oil sands and from the Atlantic to the Pacific Ocean.
69 TEMPO will have increased sensitivity to lower tropospheric O₃ compared to past/current satellite data by combining
70 measurements from both UV (290-345 nm) and visible (VIS, 540-650 nm) wavelengths (Natraj et al., 2011; Zoogman
71 et al., 2017). The operational TEMPO O₃ product will provide vertical profiles and partial O₃ columns at ~ 24 -30 layers
72 from the surface to ~ 60 km above ground level. This product will also include total, stratospheric, tropospheric, and a



73 0–2 km above ground level O₃ columns. TEMPO’s high spatial and temporal resolution measurements, including the
 74 0–2 km O₃ column, will provide a wealth of information to be used in air quality monitoring and research.

75 Vertical O₃ profile retrievals from TEMPO will be based on the Smithsonian Astrophysical Observatory
 76 (SAO) O₃ profile algorithm which was developed for use in the Global Ozone Monitoring Experiment (GOME) (Liu
 77 et al., 2005), OMI (Liu et al., 2010), GOME-2 (Cai et al., 2012), and the Ozone Mapping and Profiler Suite (Bak et
 78 al., 2017). The SAO O₃ algorithm for TEMPO will apply the tropopause-based O₃ climatology (TB-Clim) developed
 79 in Bak et al. (2013) as the a priori profiles, which was demonstrated to improve OMI O₃ retrievals near the tropopause
 80 compared to calculations using the Labow-Logan-McPeters (LLM) O₃ climatology (a priori used for OMI) (McPeters
 81 et al., 2007). During this work, we evaluate the representativeness of the vertical O₃ profiles from TB-Clim.
 82 Additionally, we evaluate simulated O₃ profiles from a near-real-time (NRT) data assimilation model product
 83 (National Aeronautics and Space Administration (NASA) Global Modeling and Assimilation Office (GMAO)
 84 Goddard Earth Observing System (GEOS-5) Forward Processing (FP)), a reanalysis data product (NASA GMAO
 85 Modern-Era Retrospective analysis for Research and Applications version 2 (MERRA2)), and a chemical transport
 86 model (CTM) (GEOS-Chem). The climatology and model O₃ profiles were evaluated with ground-based lidar data
 87 from the Tropospheric Ozone Lidar Network (TOLNet) at various locations of the US during the summer of 2014.
 88 This evaluation was focused on the performance of each product compared to summer-, daily-, and hourly-averaged
 89 lowermost tropospheric (LMT, 0–2 km) and tropospheric (0–10 km) O₃ columns to demonstrate the effectiveness of
 90 using the TB-Clim and additional products as a priori in the TEMPO O₃ profile algorithm.

91 This paper is organized as follows. Section 2 describes the tropospheric lidar O₃ measurements, TB-Clim and
 92 model products, theoretical TEMPO retrievals, and data evaluation techniques applied during this study. Section 3
 93 provides the results of the comparison of the TB-Clim and modeled a priori profile products with TOLNet observations
 94 and the impact of each product, when applied as a priori, on TEMPO tropospheric O₃ profile retrievals. Finally, Sect.
 95 4 concludes this study.

96 **2 Data and methods**

97 **2.1 TOLNet**

98 TOLNet provides Differential Absorption Lidar (DIAL)-derived vertically-resolved O₃ mixing ratios at 6 different
 99 locations of North America (<http://www-air.larc.nasa.gov/missions/TOLNet/>). TOLNet data have been used
 100 extensively in atmospheric chemistry research on topics such as STE, air pollution transport, nocturnal O₃
 101 enhancements, PBL pollution entrainment, source attribution of O₃ lamina, and the impact of wildfire and lightning
 102 NO_x on tropospheric O₃ (e.g., Kuang et al., 2011; Sullivan et al., 2015a, Johnson et al., 2016; Granados-Muñoz et al.,
 103 2017; Langford et al., 2017). Past analysis has demonstrated the high accuracy of TOLNet O₃ retrievals with errors
 104 typically estimated to be around ±10% in the lower troposphere and ±20% in the upper troposphere (Kuang et al.,
 105 2013; Sullivan et al., 2015b; Granados-Muñoz and Leblanc, 2016). TOLNet data will be applied in this study to
 106 evaluate the TB-Clim and model-predicted profiles which could potentially be used as TEMPO a priori information.
 107 Furthermore, theoretical TEMPO O₃ retrievals in the troposphere and LMT were calculated using the



climatology/model profiles as a priori with TOLNet data representing the “true” atmospheric O₃ profiles (see Sect. 2.2).

During this study, vertical O₃ profiles from 3 separate TOLNet sites during the summer (July–August) of 2014 were applied. Figure 1 shows the location of the Goddard Space Flight Center (GSFC) TROPospheric OZone (TROPOZ), Jet Propulsion Laboratory (JPL) Table Mountain Facility (TMF), and the University of Alabama in Huntsville (UAH) Rocket-city O₃ Quality Evaluation in the Troposphere (RO3QET) TOLNet systems which provided the observations used during this work. These 3 sites were selected due to data availability (<http://www-air.larc.nasa.gov/missions/TOLNet/data.html>) and to represent differing parts of North America, which will be observed by TEMPO, with varying topography, meteorology, and atmospheric chemistry conditions (overview information for each station is presented in Table 1). The RO3QET system is located in the southeast US where the air quality is impacted by both anthropogenic and natural emission sources, complex chemistry, and multiple transport pathways (e.g., Hidy et al., 2014; Johnson et al., 2016; Kuang et al., 2017). During the summer of 2014 this lidar system measured O₃ profiles from the surface to ~5 km above ground level during the daytime hours. The TROPOZ system, which is typically operated at NASA GSFC, was remotely stationed in Colorado to support the Deriving Information on Surface Conditions from COrumn and VERTically Resolved Observations Relevant to Air Quality (DISCOVER-AQ) Colorado and Front Range Air Pollution and Photochemistry Experiment (FRAPPÉ) field campaigns between July–August 2014. The TROPOZ system was arranged to take daytime observations of O₃ profiles in the intermountain west region of the US alongside the frontal range of the Rocky Mountains. The air quality of this location is impacted by large anthropogenic emission sources, complex local transport, and common STE events (e.g., Sullivan et al., 2015a; Vu et al., 2016). Finally, the TOLNet system at the JPL TMF is representative of the western US and remote high-elevation locations. This location has O₃ profiles largely controlled by long-range transport and STEs typical of remote high-elevation locations in the US (e.g., Granados-Muñoz and Leblanc, 2016; Granados-Muñoz et al., 2017). During the summer of 2014, the JPL TMF lidar only conducted measurements during the nighttime hours and therefore will only be used for daily-averaged comparisons to TB-Clim and model predictions.

2.2 TEMPO O₃ profile retrieval

TEMPO will adapt the current SAO OMI UV-only O₃ profile algorithm (Liu et al., 2010) to derive O₃ profiles from joint UV+VIS measurements based on the optimal estimation technique (Rodgers, 2000). Partial O₃ columns at different altitudes, along with other retrieved variables, are iteratively derived by simultaneously minimizing the differences between measured and simulated radiances and between the retrieved and a priori state vectors. For this study, we use the linear estimate approach to perform theoretical TEMPO retrievals and evaluate the impact of a priori profiles on these retrievals. This linear estimation approach is a good approximation of the non-linear retrieval and has been used in past research (e.g., Natraj et al., 2011). In this approach, shown in Eq. (1), the retrieved O₃ profile (X_r) is derived as:

$$X_r = X_a + A(X_t - X_a) + G\varepsilon, \quad (1)$$

where X_a is the a priori O₃ profile, A is the averaging kernel (AK) matrix, X_t is the true O₃ profile, G is the gain matrix, and ε is the measurement noise.



144 2.2.1 TEMPO averaging kernels

145 The UV+VIS AKs applied during this study are based on TEMPO retrieval sensitivity studies that play a key role in
 146 determining the instrument requirements and verification of the retrieval performance (Zoogman et al., 2017). The
 147 production of these AKs involved: 1) radiative transfer model simulations of TEMPO radiance spectra and weighting
 148 functions, 2) retrieval AKs and errors constrained by the TB-Clim a priori error covariance matrix, and 3) measurement
 149 errors estimated using the TEMPO signal to noise ratio model. To represent TEMPO hourly measurements throughout
 150 the year, the retrieval sensitivity calculation was performed hourly for 12 days (15th day of each month) over the
 151 TEMPO domain at a spatial resolution of $2.0^{\circ} \times 2.5^{\circ}$ (latitude \times longitude) using hourly GEOS-Chem model fields.
 152 During this study, we use the UV+VIS O₃ retrieval AKs corresponding to the month and location of TOLNet systems
 153 representative of near clear-sky conditions. Figure 2 shows an example of the UV+VIS AK matrix at the UAH
 154 RO3QET site for 20 UTC in August. The enhanced sensitivity of TEMPO retrievals in the lower troposphere, in
 155 particular the lowest ~2 km, is demonstrated by the large values of A (normalized to 1 km, degrees of freedom (DFS)
 156 per km) in Fig. 2 (> 0.20). When including VIS with UV wavelengths, O₃ retrievals can be greater than a factor of 2
 157 more sensitive in the first 2 km of the troposphere in comparison to just using UV wavelengths. This is particularly
 158 important as accurate O₃ observations between 0-2 km above the surface is a key requirement of TEMPO to be a
 159 sufficient data source for air quality research/monitoring (Zoogman et al., 2017).

160 2.2.2 TB-Clim

161 During this study, TB-Clim is evaluated with observations to determine the ability of these profiles to represent the
 162 spatio-temporal variability of tropospheric O₃ in North America. A detailed description of the data and procedures
 163 used to derive TB-Clim can be found in Bak et al. (2013). The climatology provides monthly-averaged O₃ profiles
 164 with 1 km vertical resolution relative to the tropopause in 18 10° -latitude bins (Bak et al., 2013). During this study,
 165 hourly TB-Clim O₃ profiles were derived by applying hourly-averaged GEOS-5 FP tropopause heights. Figure 3
 166 illustrates the monthly-averaged vertical structure of TB-Clim that will be evaluated at the RO3QET, TROPOZ, and
 167 JPL TMF system locations representative of various regions of the US in July-August 2014. At the location of the
 168 RO3QET system (Fig. 3, yellow line), O₃ values are ~55 ppb near the surface during July and August and steadily
 169 increase to ~95 ppb at 10 km. For the location of the TROPOZ system (Fig. 3, black line), O₃ values are ~40-45 ppb
 170 near the surface and increase to ~80 ppb at 10 km. Finally, at the location of the JPL TMF lidar system (Fig. 3, red
 171 line), O₃ values are ~50-55 ppb near the surface and increase to 80-95 ppb at 10 km.

172 2.3 Simulated O₃ profile data

173 Satellite O₃ retrieval algorithms typically apply climatologies derived from observational data (i.e., ozonesondes) as
 174 a priori information (Liu et al., 2005, 2010; Cai et al., 2012). However, some satellites, such as TES operational
 175 retrievals, apply climatological O₃ profiles from global CTMs as a priori information (Worden et al., 2007). During
 176 this work, we evaluate O₃ profile information from a NRT operational data assimilation model (GEOS-5 FP),
 177 reanalysis model (MERRA2), and a CTM (GEOS-Chem) using TOLNet data and investigate how model products
 178 impact theoretical TEMPO O₃ retrievals when applied as a priori information. These simulated products were selected



179 to represent model predictions of O₃ with highly varying complexity in atmospheric chemistry calculations, emissions
 180 information, data assimilation techniques, and spatial resolution.

181 **2.3.1 GEOS-5 FP and MERRA2**

182 The GEOS-5 atmospheric general circulation model (AGCM) and data assimilation system (DAS) is a product of the
 183 GMAO and is described in Rienecker et al. (2008) with most recent updates presented in Molod et al. (2012). Aerosol
 184 and trace gases are transported in the GEOS-5 AGCM using a finite-volume dynamics scheme implemented with
 185 various physics packages (Putman and Lin, 2007; Bacmeister et al., 2006) and turbulently mixed using the Lock et al.
 186 (2000) PBL scheme. The GEOS-5 AGCM ADS assimilates roughly 2×10^6 observations for each analysis using the
 187 Gridpoint Statistical Interpolation (GSI) three dimensional variational (3DVar) analysis technique (Wu et al., 2002).
 188 A product from the GEOS-5 AGCM is the operationally provided GEOS-5 FP data which offers NRT DAS predictions
 189 (typically within 24 hours) of O₃ vertical profiles at a $0.25^\circ \times 0.3125^\circ$ spatial resolution and 72 vertical levels.
 190 Additionally, we apply MERRA2 reanalysis O₃ profiles which are also produced using the GEOS-5 AGCM (Molod
 191 et al., 2012) and provided at a $0.50^\circ \times 0.667^\circ$ spatial resolution and 72 vertical levels. Both GEOS-5 FP and MERRA2
 192 O₃ vertical profiles are driven by the assimilation of OMI and Microwave Limb Sounder (MLS) satellite data.
 193 Predictions of O₃ from these products are most trusted in the upper troposphere and stratosphere due to data constraints
 194 predominantly occurring in these altitude ranges (e.g., Wargan et al., 2015; Ott et al., 2016). The work by Wargan et
 195 al. (2015) shows that due to highly simplified atmospheric chemistry and lack of surface emissions in the GEOS-5
 196 AGCM, O₃ predictions in the middle to lower troposphere tend to be biased. However, during this work these 3 hour-
 197 averaged products are applied to understand how NRT DAS and reanalysis models could be used as a priori
 198 information in TEMPO O₃ retrievals.

199 **2.3.2 GEOS-Chem**

200 GEOS-Chem (v9-02) was applied in this work as a proxy to determine how a full CTM or air quality model could
 201 potentially be used as a priori information in TEMPO O₃ retrieval algorithms. The purpose of this work is not to
 202 evaluate the performance of the GEOS-Chem model, or to suggest GEOS-Chem as the only model to provide a priori
 203 information for TEMPO, but to simply evaluate how CTM predictions impact the accuracy of theoretical TEMPO O₃
 204 retrievals. The CTM is driven by GEOS-5 FP meteorological data in a nested regional mode for July and August 2014,
 205 after a 2-month spin-up period, at a $0.25^\circ \times 0.3125^\circ$ spatial resolution and 47 hybrid terrain following vertical levels
 206 for the North American domain (130° - 60° W, 9.75° - 60° N). GEOS-Chem includes detailed O₃-NO_x-hydrocarbon-
 207 aerosol chemistry coupled to H₂SO₄-HNO₃-NH₃ aerosol thermodynamics (Bey et al., 2001). Furthermore, aerosol and
 208 trace gas transport are calculated using the TPCORE parameterization (Lin and Rood, 1996) and dry and wet
 209 deposition (Wang et al., 1998; Amos et al., 2012) is simulated on a 10-minute time-step. A detailed description of the
 210 version of GEOS-Chem, and emission inventories, applied during this study can be found in Johnson et al. (2016).

211

212



213 2.4 Data evaluation

214 The evaluation of TB-Clim and model O₃ profiles was done for summer-, daytime- (6am - 6pm local time), and hourly-
 215 averages at the RO3QET and TROPOZ system locations during July and August 2014. Due to the hours of operation,
 216 the evaluation at the JPL TMF lidar location was conducted for summer- and daily-averages. To determine the ability
 217 of a NRT DAS, reanalysis, and CTM model to replicate TOLNet-observed O₃, GEOS-5 FP, MERRA2, and GEOS-
 218 Chem data will be evaluated simultaneously with TB-Clim. For all evaluation and inter-comparisons, TB-Clim, model
 219 data, TOLNet observations, and TEMPO calculations are hourly-averaged and averaged/interpolated to the vertical
 220 grid of the TEMPO AKs during all times/locations when/where TOLNet measurements were obtained. TB-Clim and
 221 model data used as a priori and resulting X_r calculations will be evaluated using statistical parameters (correlation (R),
 222 bias, bias standard deviation (1σ), mean normalized bias (MNB), root mean squared error (RMSE)) and time-series
 223 analysis for tropospheric (0-10 km, 0-5 km for RO3QET) and LMT (0-2 km) columns. Tropospheric column values
 224 are considered to extend from the surface to 10 km in this study based on the fact that TOLNet systems typically only
 225 measured to ~10 km.

226 3 Results

227 3.1 Evaluation of TB-Clim and model-predicted tropospheric O₃ profiles

228 In terms of summertime-averaged tropospheric O₃ profiles, TB-Clim and the GEOS-5 FP, MERRA2, and GEOS-
 229 Chem models could generally replicate the vertical structure of tropospheric O₃ measured by TOLNet lidars. However,
 230 the evaluation of these products as a priori in TEMPO O₃ retrievals at a seasonal/monthly average is insufficient as
 231 TEMPO will provide hourly, high spatial resolution, tropospheric and LMT O₃ values. Therefore, in the following
 232 sections we evaluate these products for daily- and hourly-averages to focus on inter-daily and diurnal variability.

233 3.1.1 Daily-averaged tropospheric O₃ profiles

234 This section focuses on evaluating the ability of TB-Clim and the GEOS-5 FP, MERRA2, and GEOS-Chem models
 235 to reproduce observed daily variability of O₃ in the troposphere and near the surface. Figure 4 shows the daily-averaged
 236 tropospheric and LMT O₃ columns from TB-Clim and models compared to that observed by TOLNet at all 3 sites
 237 with comparison statistics displayed in Table 2. Some slight inter-daily variability can be seen in TB-Clim tropospheric
 238 O₃ due to varying time-dependent tropopause heights, however, the variability in LMT values is mostly due to only
 239 sampling values in the vertical layers and times when TOLNet observations were obtained (vertical layers of TOLNet
 240 observations varied between hours and days). Due to the zonal and monthly mean nature of TB-Clim, this dataset is
 241 unable to replicate inter-daily O₃ observations consistently displaying negative correlation values with daily TOLNet
 242 observations in the troposphere (R range between -0.09 and -0.35) and near the surface (R range between -0.15 and -
 243 0.68). The models demonstrate a better ability to replicate the daily variability of observed tropospheric O₃ at the
 244 TOLNet system locations. Overall, CTM predictions from GEOS-Chem was the only potential source of a priori
 245 profiles which consistently displayed moderate to high positive correlation (all R values > 0.47) compared to all
 246 TOLNet observations in the troposphere and near the surface. This result is not overly surprising as a full CTM



includes aspects necessary to reproduce the spatio-temporal tropospheric O₃ variability occurring in nature such as data-assimilated meteorological fields, comprehensive atmospheric chemistry mechanisms, and state-of-the-art trace gas and aerosol emissions data.

Figure 4a, b shows larger variability of daily-averaged LMT O₃ (44 to 68 ppb) from the RO3QET system than that in the tropospheric column (48 to 64 ppb). From Table 2 it can be seen that TB-Clim was generally high compared to lidar-measured tropospheric O₃ mixing ratios (average bias = 3.7 ppb) with large bias standard deviations and RMSE values (> 6 ppb). MERRA2 displayed good agreement in tropospheric O₃ (negative bias ~0.7 ppb) while GEOS-5 FP and GEOS-Chem resulted in moderate high biases (average bias 2.8 and 1.7 ppb, respectively). GEOS-Chem had moderate high biases but with smaller bias standard deviation and RMSE values (< 4.5 ppb) in comparison to the other products due to the ability to better capture inter-daily tropospheric O₃ variability ($R = 0.61$). LMT O₃ observations by the RO3QET lidar were best replicated by the CTM product resulting in the smallest average bias (-1.3 ppb) and bias standard deviation and RMSE values (4.4 ppb) compared to the other products. MERRA2 was consistently low compared to LMT O₃ observations (bias = -4.9 ppb) while TB-Clim and GEOS-5 FP resulted in moderate biases (2.9 and -2.9 ppb, respectively) with all of these products having large bias standard deviations and RMSE (≥ 8.0 ppb).

At the TROPOZ system location, large variability in tropospheric (47 to 83 ppb) and LMT O₃ values (41 to 73 ppb) was observed. From Fig. 4c, d and Table 2 it can be seen that TB-Clim is unable to replicate the inter-daily tropospheric O₃ variability and is generally higher in comparison to observations with large bias standard deviations (bias \pm standard deviation = 2.2 ± 9.7 ppb). GEOS-Chem best replicates the daily variability of tropospheric O₃ with the largest correlation ($R = 0.82$) and small average bias and standard deviations (2.4 ± 6.0 ppb). GEOS-5 FP and MERRA2 data displayed low positive correlations ($R < 0.40$) and larger average biases and standard deviations 3.3 ± 10.0 and -4.6 ± 9.1 ppb, respectively. In comparison to TROPOZ LMT O₃ observations, TB-Clim and all model products displayed large negative biases. The TB-Clim product resulted in the largest negative biases and bias standard deviations compared to LMT O₃ observations (-11.1 ± 7.5 ppb) and model products displayed smaller biases and standard deviations. GEOS-5 FP data displayed the lowest average bias (-4.4 ppb) compared to TROPOZ observations, however, was unable to replicate the inter-daily variability of LMT O₃ ($R = -0.09$) resulting in large bias standard deviations (7.3 ppb). Overall, GEOS-Chem was the only product which was able to capture the inter-daily variability of LMT O₃ ($R = 0.47$) resulting in moderate low biases and the lowest bias standard deviation (-6.7 ± 6.2 ppb).

Figure 4e, f illustrates that large inter-daily variability of tropospheric (46 to 129 ppb) and LMT (35 to 76 ppb) column O₃ was observed at the JPL TMF site during the summer of 2014. This figure and Table 2 shows that TB-Clim is able to represent the average magnitude of tropospheric O₃ (bias = 0.3 ppb) but with large bias standard deviation and RMSE values (>18 ppb) due to the inability to replicate observed inter-daily variability ($R = -0.35$). The GEOS-Chem model also captures the average magnitude of tropospheric O₃ (bias = -0.5 ppb) but with smaller bias standard deviations (14.6 ppb) compared to TB-Clim due to the ability to replicate the inter-daily availability ($R = 0.72$). GEOS-5 FP and MERRA2 demonstrated negative biases compared to JPL TMF lidar observed tropospheric O₃ (-5.0 and -10.6 ppb, respectively) with relatively low bias standard deviations (~13-14 ppb) compared to the other



products. The large RMSE values for all products is due to the very large variability in daily-averaged O_3 observations which was not well captured by all products. Near the surface, the GEOS-Chem model clearly best captures the variability of daily-averaged LMT O_3 indicated by the smallest bias and standard deviations (0.9 ± 10.4 ppb) and RMSE (~ 10.25 ppb) values.

3.1.2 Diurnal cycle of tropospheric O_3 profiles

TEMPO retrievals will produce hourly tropospheric and LMT O_3 values each day for the entire North America domain. Therefore, this section focuses on evaluating the ability of TB-Clim and the GEOS-5 FP, MERRA2, and GEOS-Chem models to reproduce the observed diurnal variability of O_3 measured at the RO3QET and TROPOZ system locations in the troposphere and near the surface. Figure 5 shows the average diurnal time-series of hourly-averaged tropospheric and LMT O_3 (from all days of observation) from the O_3 climatology and models compared to that observed during the summer of 2014 (statistics displayed in Table 3).

Figure 5a, b shows that larger diurnal variability of O_3 was observed for LMT values (48 to 59 ppb) compared to tropospheric values (55 to 60 ppb) at the RO3QET lidar location. All the potential sources of a priori profiles, excluding the CTM predictions, demonstrate very little diurnal variation in tropospheric and LMT O_3 at the RO3QET lidar location during the summer of 2014. The GEOS-Chem model was the only product able to replicate the diurnal variability of observed tropospheric O_3 ($R = 0.68$). MERRA2 resulted in the lowest bias (-1.2 ppb), GEOS-5 FP and GEOS-Chem displayed modest biases (~ 2.0 – 2.5 ppb), and TB-Clim had the largest bias (3.5 ppb) compared to RO3QET tropospheric O_3 data. Diurnal RO3QET LMT O_3 data was best replicated by CTM predictions resulting in the highest correlation ($R = 0.76$), lowest bias and standard deviations (0.3 ± 2.6 ppb), and RMSE values (2.45 ppb). The TB-Clim product resulted in modest biases compared to LMT O_3 data (1.9 ppb) while GEOS-5 FP and MERRA2 were consistently low (negative bias > 3.0 ppb).

Figure 5c, d shows the diurnal variability of O_3 that was observed for tropospheric and LMT column values at the TROPOZ lidar location during the summer of 2014. In the troposphere, O_3 values varied between ~ 58 to 69 ppb with largest values occurring in the afternoon. Larger diurnal variability was observed near the surface with LMT O_3 values ranging from ~ 56 to 75 ppb with largest values occurring between 21 and 05 UTC. GEOS-Chem data is the only product which could replicate the diurnal variability of TROPOZ lidar tropospheric O_3 observations ($R = 0.78$). The TB-Clim, GEOS-5 FP, and GEOS-Chem products demonstrate moderate high biases (2.2–3.3 ppb) compared to the observations while MERRA2 is consistently low (bias = -5.1 ppb). For comparison of near-surface O_3 values (see Fig. 5d), none of the products sufficiently captured the magnitude and degree of diurnal variability of LMT O_3 at the TROPOZ lidar location. The TB-Clim product displayed a small positive correlation ($R = 0.26$) and large negative biases (-12.6 ppb), bias standard deviation (6.9 ppb), and RMSE values (14.25 ppb). The GEOS-5 FP and GEOS-Chem models display the lowest bias (negative bias between 7.5 ppb and 7.7 ppb), however, the CTM is more highly correlated ($R = 0.92$) and resulted in lower bias standard deviations (4.8 ppb) and RMSE values (9.01 ppb). This indicates that while no product reproduced the magnitude or degree of diurnal variability of near-surface O_3 observed by the TROPOZ lidar, the GEOS-Chem CTM does the best job on average.



320 3.2 Prior O₃ vertical profile impact on TEMPO retrievals

321 This section focuses on how the TB-Clim, GEOS-5 FP, MERRA2, and GEOS-Chem O₃ profiles impact theoretical
 322 TEMPO tropospheric O₃ profile retrievals when applied as the a priori information in Eq. (1). The evaluation is focused
 323 on how different sources of a priori impacted the overall accuracy of TEMPO tropospheric O₃ retrievals and the ability
 324 to meet the required precision of tropospheric and LMT O₃ observations of 10 ppb (Zoogman et al., 2017). The
 325 requirement for TEMPO tropospheric O₃ is that retrieval errors (root square sum of retrieval precision and smoothing
 326 errors) or overall biases should be < 10 ppb, and, therefore, we quantify the number of occurrences when total error
 327 or bias standard deviation/RMSE exceeds this 10 ppb limit. TEMPO will provide tropospheric and LMT O₃ at high
 328 temporal resolution and therefore, X_r values from Eq. (1), using the individual a priori sources, will be evaluated on
 329 a daily-averaged and diurnal cycle time scale.

330 3.2.1 Tropospheric O₃ TEMPO retrievals

331 Figure 6 shows the time-series of daily-averaged tropospheric and LMT X_r column values and bias calculations when
 332 using TB-Clim and model data as a priori information when compared to observed O₃ at all 3 TOLNet sites (statistics
 333 in Table 4). When focusing on the accuracy of the theoretical TEMPO retrievals for tropospheric X_r columns (left
 334 column in Fig. 6), it can be seen that: 1) X_r values using all a priori profiles are similar and 2) X_r values compare well
 335 to observations with tropospheric X_r values typically falling within the 10 ppb bias requirement at all 3 TOLNet
 336 locations. From Table 4 it can be seen that daily-averaged tropospheric column biases exceeded the 10 ppb level on 1
 337 and 2 days when using TB-Clim/GEOS-5 FP and MERRA2 data, respectively, as a priori when compared to TROPOZ
 338 observations, and for 1 day at the JPL TMF location when using all O₃ products as a priori.

339 Table 4 illustrates that applying TB-Clim as the a priori resulted in the largest tropospheric column X_r biases
 340 and modest bias standard deviations (1.4 ± 2.3 ppb) and the MERRA2 data led to the lowest overall bias and modest
 341 bias standard deviation (-0.2 ± 2.5 ppb) at the RO3QET lidar location. Using GEOS-Chem a priori profiles resulted
 342 in modest biases and the lowest bias standard deviations (1.0 ± 2.0 ppb) and RMSE values (2.17 ppb). At the TROPOZ
 343 system site, the lowest tropospheric column X_r biases and standard deviation were calculated when applying GEOS-
 344 Chem as the a priori (-0.5 ± 2.7 ppb). GEOS-5 FP data also resulted in low mean X_r biases but the largest bias standard
 345 deviations (-0.6 ± 4.8 ppb) and MERRA2 data led to larger mean X_r biases but lower bias standard deviations ($-2.2 \pm$
 346 4.4 ppb). The use of TB-Clim resulted in modest mean bias and standard deviations (-0.9 ± 4.2 ppb). Finally, at the
 347 JPL TMF location all a priori profile sources resulted in average tropospheric column X_r biases of < 1.0 ppb, excluding
 348 MERRA2 (bias = -1.7 ppb), with similar bias standard deviations and RMSE values (ranging between 3.0 to 4.0 ppb).
 349 Much larger daily variability of tropospheric O₃ was observed at the JPL TMF site compared to the other TOLNet
 350 system locations and tropospheric column X_r values from theoretical TEMPO retrievals successfully captured this
 351 variability using all the sources of a priori information. These results suggest that TEMPO, using UV+VIS
 352 wavelengths, will likely be able to accurately retrieve highly variable tropospheric column O₃ values using a variety
 353 of sources of a priori profiles.

354



3.2.2 LMT O₃ TEMPO retrievals

The third column of Fig. 6 shows that much larger differences in daily-averaged LMT column X_r values were calculated, compared to tropospheric X_r values, when using different sources of a priori in Eq. (1). It can be seen from this figure that at the RO3QET site, daily variability of near-surface O₃ are clearly best captured by LMT X_r values using GEOS-Chem CTM a priori profiles. While the TB-Clim product resulted in LMT X_r values with the smallest mean bias (0.2 ppb), it also led to large RMSE values (5.88 ppb) and the largest bias standard deviations (6.1 ppb) (see Table 4). Table 4 illustrates that LMT column X_r values calculated using CTM a priori profiles had modest mean bias (-2.2 ppb) and the lowest bias standard deviations (2.5 ppb) and RMSE (3.26 ppb). Applying the GEOS-5 FP and MERRA2 model products as a priori profiles resulted in the largest mean biases in LMT X_r values (negative biases \geq 3.4 ppb) along with largest RMSE values (\geq 6.0 ppb). From an air quality perspective, it is important to note that LMT column X_r values using a priori data other than GEOS-Chem are unable to replicate the larger surface O₃ values occurring in the southeast US (see Fig. 6). A few LMT O₃ accuracy/precision requirement exceedances were calculated at the RO3QET lidar location using all a priori products except for GEOS-Chem predictions. The ability of GEOS-Chem to best reproduce the magnitude of the daily LMT O₃ variability resulted in LMT X_r values with the smallest RMSE and bias standard deviations, no accuracy/precision requirement exceedances, and the best ability to capture the range in daily observed O₃.

At the location of the TROPOZ lidar, it can be seen from Fig. 6 that LMT X_r values, with the use of TB-Clim a priori, are consistently underestimated in comparison to lidar observations. These LMT X_r values have an average negative bias of > 10.0 ppb and largest RMSE values (~ 13.0 ppb) resulting in 10 days with accuracy/precision requirement exceedances (see Table 4). These large errors are because the a priori profiles provided by TB-Clim are not able to replicate the highly variable vertical O₃ profiles observed at the TROPOZ lidar location. The GEOS-5 FP, MERRA2, and GEOS-Chem models were better able to replicate these highly variable vertical O₃ profiles providing a priori information more accurately representing O₃ in the intermountain west region of the US. This better representation from model data resulted in LMT X_r values with lower negative mean biases (< 6.5 ppb) and smaller RMSE values (< 9.0 ppb) and bias standard deviations (< 6.5 ppb), and also fewer accuracy/precision requirement exceedances. Overall, CTM-predicted a priori information resulted in LMT X_r values with the least bias and bias standard deviation (-4.8 ± 4.8 ppb), RMSE (6.71 ppb), and accuracy/precision exceedances.

At the location of the JPL TMF lidar, much larger daily variability in LMT O₃ mixing ratios were observed during the summer of 2014 compared to the other TOLNet systems. LMT X_r values, using all sources of data as a priori information, had difficulty in replicating this large variability (see Fig. 6). From Table 4, it can be seen that despite relatively low biases for all sources of a priori (< 5.0 ppb), the inability of LMT X_r values to capture the dynamic daily variability resulted in large bias standard deviations and RMSE values (> 12.5 ppb). Furthermore, 6-10 accuracy/precision requirement exceedances out of 26 total days were calculated. Despite 6 error exceedances (the least of all profile products), applying GEOS-Chem predictions as a priori information resulted in the lowest mean biases (1.0 ppb) and RMSE values (12.54 ppb). Typically, large underestimations of LMT X_r values occurred when the lidar observed large O₃ enhancements near the surface and significant overestimations of LMT X_r values were calculated when the lidar observed very large O₃ lamina (> 150 ppb) aloft. This indicates that the shape of the a priori



O₃ vertical profile used in TEMPO tropospheric O₃ retrievals are very important in order to capture X_r values for both the tropospheric and LMT column and this will be discussed in Sect. 3.2.3.

Figure 6 and Table 4 demonstrate that in general X_r values in the troposphere and near the surface are more accurately retrieved when applying model predictions, and in particular CTM values from GEOS-Chem, at all 3 TOLNet system locations. Also, from this figure it can be seen that in general when large daily-averaged LMT O₃ mixing ratios are observed (here defined as days with daily-averaged LMT O₃ > 65 ppb), which are important for air quality purposes, LMT X_r values display less bias when applying GEOS-Chem a priori profile information compared to all other products. For the 11 days in which daily-averaged LMT O₃ mixing ratios exceeded 65 ppb, 64%, 9%, and 27% of the LMT X_r values had the smallest bias using GEOS-Chem, GEOS-5 FP, and MERRA2 a priori profiles, respectively. This suggests that applying CTM predictions as a priori profile information will allow TEMPO to observe air quality relevant pollution concentrations of LMT O₃ more accurately compared to TB-Clim and models with limited chemistry and emission schemes evaluated during this work.

3.2.3 Importance of a priori vertical profile shape

Figure 7 displays examples of why climatological a priori information in theoretical TEMPO retrievals resulted in large daily-averaged LMT column X_r biases. The first example in Fig. 7a shows the daily-averaged vertical profiles of X_a and X_r with the use of TB-Clim and GEOS-Chem a priori on 08 July 2014 at the JPL TMF site when the lidar observed large LMT O₃ values above EPA NAAQS levels. This case study illustrates how CTMs are more likely to be able to replicate surface O₃ enhancements compared to climatological products. The GEOS-Chem a priori information resulted in more accurate TEMPO X_r values for the tropospheric and LMT O₃ column values. When using GEOS-Chem model predictions as a priori information, TEMPO LMT column X_r retrievals (65.1 ppb) were closer in magnitude to observations (70.2 ppb) compared to when using TB-Clim a priori (54.7 ppb). Furthermore, when using GEOS-Chem a priori information, TEMPO retrievals for the troposphere (65.8 ppb) were also more similar in magnitude to lidar observations (64.2 ppb) compared to using a priori data from TB-Clim (68.2 ppb).

Another example is illustrated in Fig. 7b which shows X_a and X_r when using TB-Clim and GEOS-5 FP predictions as a priori profiles in TEMPO retrievals on 21 August 2014 at the JPL TMF lidar location. On this day, a STE event was likely occurring as tropospheric O₃ mixing ratios were measured to be > 200 ppb between 6–9 km. This case study illustrates how a NRT DAS model, GEOS-5 FP, displayed some ability to replicate the large O₃ lamina in the middle/upper troposphere due to being constrained with upper atmospheric observations. The GEOS-5 FP a priori information resulted in more accurate TEMPO X_r values for the tropospheric and LMT O₃ column values. When using GEOS-5 FP data as a priori information, TEMPO X_r values for tropospheric O₃ of 130.4 ppb compared closely to the JPL TMF lidar observations (135.6 ppb) while TB-Clim data resulted in much lower values (112.4 ppb). However, the large adjustment needed to correct the a priori profiles to match tropospheric column O₃ observations led to noticeable overestimations of TEMPO LMT X_r values. Since the GEOS-5 FP a priori data was able to better replicate the STE event compared to TB-Clim, the LMT X_r overestimation of observed LMT O₃ values (48.8 ppb) is much less when applying GEOS-5 FP (77.6 ppb) than when applying TB-Clim (99.1 ppb). Overall, these results demonstrate that because TEMPO will only have ~1.5 DFS in the troposphere (only ~0.2 DFS in the 0–2 km level), it



is important for a priori profiles to match the general shape of observations, throughout the entire troposphere and LMT, in order to accurately retrieve both total tropospheric and LMT O₃ values.

3.2.4 Diurnal cycle of tropospheric TEMPO retrievals

This section focuses on evaluating the ability of TEMPO to retrieve hourly-averaged tropospheric O₃ applying TB-Clim and the GEOS-5 FP, MERRA2, and GEOS-Chem models as a priori profile information. This evaluation was conducted for one day each at the RO3QET and TROPOZ sites where constant lidar measurements were obtained in the troposphere/LMT and near-surface O₃ enhancements with potential air quality relevant impacts were observed. Figure 8 shows the time-series of hourly-averaged tropospheric and LMT column X_r retrievals when using TB-Clim and models as a priori compared to that observed by RO3QET on 07 August 2014 and by TROPOZ on 22 July 2014. This figure also displays the a priori vertical O₃ profiles used in TEMPO retrievals for the hour of largest LMT O₃ observations from the TOLNet systems (20 UTC at the RO3QET location and 22 UTC at the TROPOZ site location).

In comparison to lidar measurements by RO3QET, TEMPO retrievals, with all sources of a priori profiles, are able to reproduce the diurnal pattern of tropospheric and LMT column O₃ values (all R values > 0.98) (see Fig. 8). Table 5 shows that all a priori products allowed TEMPO to retrieve average tropospheric column O₃ with minimal biases, however, GEOS-Chem was the only product which resulted in LMT X_r values comparable to observations. This is because GEOS-Chem a priori profiles allow for more dynamic O₃ retrievals for the entire troposphere and LMT. This is demonstrated by the fact that the daily-mean and standard deviation (1σ) of hourly LMT O₃ from TEMPO using GEOS-Chem a priori information (62.1 ± 5.4 ppb) compared the closest to RO3QET observations (65.2 ± 9.3 ppb). The daily-mean and standard deviations for LMT X_r retrievals, using the other a priori profiles, underpredicted the magnitude and diurnal variability to a higher degree compared to predictions using GEOS-Chem a priori.

Similar results are displayed in Fig. 8 and Table 5 when evaluating the case study at the TROPOZ site location. Once again, TEMPO retrievals with all sources of a priori profiles are generally able to reproduce the diurnal pattern of tropospheric and LMT column O₃ values (all R values > 0.51) but all show large negative biases compared to LMT observations. However, Table 5 shows that GEOS-Chem model a priori data allows TEMPO to retrieve hourly tropospheric and LMT O₃ with the least bias. LMT X_r values using the TB-Clim, GEOS-5 FP, and MERRA2 a priori information displayed too little diurnal variability (nearly a factor of 2 lower standard deviation compared to TEMPO retrievals using GEOS-Chem a priori data) and a consistent underestimate of observations. During both case studies, a priori profile shape was critical for TEMPO retrievals to accurately retrieve both tropospheric and LMT O₃. Figure 8 shows a priori profiles from all products for the hour of each day where largest LMT O₃ observations occurred. This figure further emphasizes that GEOS-Chem CTM simulations are able to better capture the dynamic vertical O₃ profiles observed by the lidars compared to the other a priori profile sources. While the GEOS-Chem X_a profiles underestimate the large LMT O₃ enhancements, the ability to replicate the general shape greatly improves tropospheric and LMT column TEMPO X_r values.



463 4 Conclusions

464 This study evaluated the a priori vertical O₃ profiles currently suggested to be used in TEMPO tropospheric profile
 465 retrievals (TB-Clim) and simulated profiles from operational (GEOS-5 FP), reanalysis (MERRA2), and CTM
 466 predictions (GEOS-Chem). The spatio-temporal representativeness of the vertical profiles from each product was
 467 evaluated using TOLNet lidar observations of tropospheric O₃ during the summer (July-August) of 2014. The TOLNet
 468 sites used in this study are situated in areas which represent the southeastern US (RO3QET), intermountain west
 469 (TROPOZ), and remote high-elevation locations in the western US (JPL TMF). Because TEMPO will provide high
 470 spatial resolution tropospheric (0-10 km) and LMT (0-2 km) O₃ values on an hourly time scale, potential sources of a
 471 priori profiles must be able to replicate inter-daily variability and the diurnal cycle of observed vertical tropospheric
 472 O₃ profiles.

473 When evaluating summertime-averaged tropospheric O₃ profiles, it was found that TB-Clim, GEOS-5 FP,
 474 MERRA2, and GEOS-Chem data could generally replicate the vertical structure of tropospheric O₃ measured by
 475 TOLNet lidars. However, the seasonal/monthly evaluation is insufficient as TEMPO will provide hourly, high spatial
 476 resolution, tropospheric and LMT O₃ values. The evaluation of daily-averaged tropospheric and LMT column O₃
 477 values from these products using lidar observations resulted in varying statistical comparisons. Overall, at all 3
 478 TOLNet system locations, GEOS-Chem provided the only data product which consistently captured the inter-daily
 479 variability of tropospheric and LMT column O₃ observations. Furthermore, due to the monthly- and zonal-mean nature
 480 of TB-Clim, this product was unable to reproduce the inter-daily variability of tropospheric O₃. The ability of the
 481 models, in particular GEOS-Chem, to better replicate the temporal variability of O₃ observations led to better statistical
 482 comparison to daily-averaged TOLNet data. An important fact demonstrated in this study is that models, primarily
 483 GEOS-Chem CTM predictions, displayed better skill in reproducing the largest peaks in daily-averaged near surface
 484 O₃ observations which have important implications for air quality. This is partially because GEOS-Chem data best
 485 replicated the diurnal cycle of observations of tropospheric and LMT column O₃ from observations. Overall, the
 486 GEOS-Chem CTM predictions had the best statistical comparison to daily- and hourly-averaged tropospheric and
 487 LMT column O₃ observations.

488 The importance of different a priori profile products for TEMPO tropospheric O₃ retrievals was evaluated
 489 during this study. The results demonstrate that since TEMPO only has ~1.5 DFS in the troposphere (and ~0.2 in the
 490 0-2 km column), the ability of the a priori profile to replicate the actual shape of the “true” O₃ vertical structure
 491 (throughout the entire troposphere and LMT) is important in order for the satellite to accurately retrieve both
 492 tropospheric column and near surface O₃ values. Although TEMPO X_r values using all a priori data were able to
 493 accurately retrieve highly variable column tropospheric O₃ values, there were large differences in LMT X_r values. In
 494 general, LMT column X_r values were more accurately retrieved with model a priori profiles, especially with GEOS-
 495 Chem predictions. The better performance of TEMPO LMT X_r values, with GEOS-Chem a priori profiles, is because
 496 it better reproduces the dynamic vertical structures and inter-daily/diurnal variability of tropospheric O₃. Most
 497 importantly from an air quality perspective is that when large daily-averaged LMT O₃ mixing ratios were observed,
 498 X_r values near the surface with GEOS-Chem a priori displayed the least bias. Overall, this study suggests that applying



499 a CTM as a priori will likely allow TEMPO retrievals to observe air quality relevant O₃ concentrations more accurately
 500 than TB-Clim and other models with limited atmospheric chemistry and emission schemes.

501 This study is a first step in determining what source of a priori vertical O₃ profiles should be applied to best
 502 enhance the ability of TEMPO to retrieve tropospheric and LMT column O₃ in North America. It demonstrates that
 503 model simulations, in particular those from a CTM, improve TEMPO tropospheric O₃ retrievals over TB-Clim data.
 504 However, there are instances where CTM predictions do not improve TEMPO retrieved values compared to the TB-
 505 Clim data. Furthermore, out of the 59 total days of TOLNet observations analyzed during this study, LMT column X_r
 506 values using GEOS-Chem a priori profiles show biases greater than the TEMPO 10 ppb accuracy requirement for
 507 ~15% of the days. It should be noted that this number of LMT column X_r error exceedances is the least compared to
 508 when using all the sources of a priori and greater than a factor of 2 smaller than when applying TB-Clim a priori. The
 509 main reason for the majority of error exceedances is because the a priori profiles cannot capture the dynamic vertical
 510 O₃ profile observed by the TOLNet lidars. Therefore, further work is needed to identify the source of a priori O₃
 511 profiles for use in TEMPO O₃ retrievals which can best capture the shape of tropospheric O₃ profiles in North America.

512 *Acknowledgements.* This work is supported by the TOLNet program within NASA's Science Mission Directorate. X.
 513 Liu and P. Zoogman were supported by the NASA Earth Venture Instrument TEMPO project (NNL13AA09C). The
 514 authors would also like to thank the Harvard University Atmospheric Chemistry Modeling Group for providing the
 515 GEOS-Chem model and the NASA GMAO for providing the GEOS-5 FP and MERRA2 products used during our
 516 research. Resources supporting this work were provided by the NASA High-End Computing (HEC) Program through
 517 the NASA Advanced Supercomputing (NAS) Division at NASA Ames Research Center. All the authors express
 518 gratitude to the support from NASA's Earth Science Division at Ames Research Center. Finally, the views, opinions,
 519 and findings contained in this report are those of the authors and should not be construed as an official NASA or
 520 United States Government position, policy, or decision.



521 References

- 522 Amos, H. M., Jacob, D. J., Holmes, C. D., Fisher, J. A., Wang, Q., Yantosca, R. M., Corbitt, E. S., Galarneau, E.,
523 Rutter, A. P., Gustin, M. S., Steffen, A., Schauer, J. J., Graydon, J. A., St. Louis, V. L., Talbot, R. W., Edgerton,
524 E. S., Zhang, Y., and Sunderland, E. M.: Gas-particle partitioning of atmospheric Hg(II) and its effect on global
525 mercury deposition, *Atmos. Chem. Phys.*, 12, 591-603, <https://doi.org/10.5194/acp-12-591-2012>, 2012.
- 526 Atkinson, R.: Gas-phase Tropospheric Chemistry of Organic Compounds: A Review, *Atmos. Environ.*, 26, 1, 1-41,
527 doi:10.1016/0960-1686(90)90438-S, 1990.
- 528 Bacmeister, J. T., Suarez, M. J., and Robertson, F. R.: Rain Re-evaporation, Boundary Layer Convection Interactions,
529 and Pacific Rainfall Patterns in an AGCM, *J. Atmos. Sci.*, 63, 3383-3403, <https://doi.org/10.1175/JAS3791.1>,
530 2006.
- 531 Bak, J., Liu, X., Wei, J. C., Pan, L. L., Chance, K., and Kim, J. H.: Improvement of OMI ozone profile retrievals in
532 the upper troposphere and lower stratosphere by the use of a tropopause-based ozone profile climatology, *Atmos.*
533 *Meas. Tech.*, 6, 2239-2254, <https://doi.org/10.5194/amt-6-2239-2013>, 2013.
- 534 Bak, J., Liu, X., Kim, J.-H., Haffner, D. P., Chance, K., Yang, K., and Sun, K.: Characterization and correction of
535 OMPS nadir mapper measurements for ozone profile retrievals, *Atmos. Meas. Tech.*, 10, 4373-4388,
536 <https://doi.org/10.5194/amt-10-4373-2017>, 2017.
- 537 Beer, R.: TES on the aura mission: scientific objectives, measurements, and analysis overview, *IEEE Trans. Geosci.*
538 *Remote Sens.*, 44, 1102-1105, doi:10.1109/TGRS.2005.863716, 2006.
- 539 Bey, I., Jacob, D. J., Yantosca, R. M., Logan, J. A., Field, B., Fiore, A. M., Li, Q., Liu, H., Mickley, L. J., and Schultz,
540 M.: Global modeling of tropospheric chemistry with assimilated meteorology: Model description and evaluation,
541 *J. Geophys. Res.*, 106, 23073-23095, doi:10.1029/2001JD000807, 2001.
- 542 Bowman, K. W., Rodgers, C. D., Kulawik, S. S., Worden, J., Sarkissian, E., Osterman, G., Steck, T., Lou, M.,
543 Eldering, A., Shephard, M., Worden, H., Lampel, M., Clough, S., Brown, P., Rinsland, C., Gunson, M., and Beer,
544 R.: Tropospheric Emission Spectrometer: retrieval method and error analysis, *IEEE Trans. Geosci. Remote Sens.*,
545 44, 1297-1307, doi:10.1109/TGRS.2006.871234, 2006.
- 546 Cai, Z., Liu, Y., Liu, X., Chance, K., Nowlan, C. R., Lang, R., Munro, R., and Suleiman, R.: Characterization and
547 correction of Global Ozone Monitoring Experiment 2 ultraviolet measurements and application to ozone profile
548 retrievals?, *J. Geophys. Res.*, 117, D07305, doi:10.1029/2011JD017096, 2012.
- 549 Chance, K., Liu, X., Suleiman, R. M., Flittner, D. E., Al-Saadi, J., and Janz, S. J.: Tropospheric emissions: Monitoring
550 of pollution (TEMPO), *Earth Observing Systems XVIII*, Paper 88660D, doi:10.1117/12.2024479, 2013.
- 551 Duncan, B. N., Prados, A. I., Lamsal, L. N., Liu, Y., Streets, D. G., Gupta, P., Hilsenrath, E., Kahn, R. A., Nielsen, J.
552 E., Beyersdorf, A. J., Burton, S. P., Fiore, A. M., Fishman, J., Henze, D. K., Hostetler, C. A., Krotkov, N. A., Lee,
553 P., Lin, M. Y., Pawson, S., Pfister, G., Pickering, K. E., Pierce, R. B., Yoshida, Y., and Ziemba, L. D.: Satellite
554 data of atmospheric pollution for US air quality applications: Examples of applications, summary of data end-user
555 resources, answers to FAQs, and common mistakes to avoid, *Atmos. Environ.*, 94, 647-662,
556 doi:10.1016/j.atmosenv.2014.05.061, 2014.



- 557 Fishman, J., Bowman, K. W., Burrows, J. P., Richter, A., Chance, K. V., Edwards, D. P., Martin, R. V., Morris, G.
558 A., Pierce, R. B., Ziemke, J. R., Al-Saadi, J. A., Creilson, J. K., Schaack, T. K., and Thompson, A.M.: Remote
559 sensing of tropospheric pollution from space, *Bulletin of the American Meteorological Society*, 89, 805-821,
560 <https://doi.org/10.1175/2008BAMS2526.1>, 2008.
- 561 Granados-Muñoz, M. J. and Leblanc, T.: Tropospheric ozone seasonal and long-term variability as seen by lidar and
562 surface measurements at the JPL-Table Mountain Facility, California, *Atmos. Chem. Phys.*, 16, 9299-9319,
563 doi:10.5194/acp-16-9299-2016, 2016.
- 564 Granados-Muñoz, M. J., Johnson, M. S., and Leblanc, T.: Influence of the North American monsoon on Southern
565 California tropospheric ozone levels during summer in 2013 and 2014, *Geophys. Res. Lett.*, 44, 6431-6439,
566 doi:10.1002/2017GL073375, 2017GL073375, 2017.
- 567 Hidy, G. M., Blanchard, C. L., Baumann, K., Edgerton, E., Tanenbaum, S., Shaw, S., Knipping, E., Tombach, I.,
568 Jansen, J., and Walters, J.: Chemical climatology of the southeastern United States, 1999-2013, *Atmos. Chem.*
569 *Phys.*, 14, 11893-11914, doi:10.5194/acp-14-11893-2014, 2014.
- 570 Johnson, M. S., Kuang, S., Wang, L., and Newchurch, M. J.: Evaluating Summer-Time Ozone Enhancement Events
571 in the Southeast United States, *Atmosphere*, 7, 108, doi:10.3390/atmos7080108, 2016.
- 572 Kuang, S., Newchurch, M. J., Burris, J., Wang, L., Buckley, P., Johnson, S., Knupp, K., Huang, G., Phillips, D., and
573 Cantrell, W.: Nocturnal ozone enhancement in the lower troposphere observed by lidar, *Atmos. Environ.*, 45, 6078-
574 6084, <https://doi.org/10.1016/j.atmosenv.2011.07.038>, 2011.
- 575 Kuang, S., Newchurch, M. J., Burris, J., and Liu, X.: Ground-based lidar for atmospheric boundary layer ozone
576 measurements, *Appl. Opt.*, 52, 3557-3566, <https://doi.org/10.1364/AO.52.003557>, 2013.
- 577 Kuang, S., Newchurch, M. J., Johnson, M. S., Wang, L., Burris, J., Pierce, R. B., Eloranta, E. W., Pollack, I. B., Graus,
578 M., de Gouw, J., Warneke, C., Ryerson, T. B., Markovic, M. Z., Holloway, J. S., Pour-Biazar, A., Huang, G., Liu,
579 X., and Feng, N.: Summertime tropospheric ozone enhancement associated with a cold front passage due to
580 stratosphere- to-troposphere transport and biomass burning: simultaneous ground-based lidar and airborne
581 measurements, *J. Geophys. Res.*, 122, 1293-1311, doi:10.1002/2016JD026078, 2017.
- 582 Langford, A. O., Alvarez, R. J., Brioude, J., Fine, R., Gustin, M. S., Lin, M. Y., Marchbanks, R. D., Pierce, R. B.,
583 Sandberg, S. P., Senff, C. J., Weickmann, A. M., and Williams, E. J.: Entrainment of stratospheric air and Asian
584 pollution by the convective boundary layer in the southwestern U.S., *J. Geophys. Res.*, 122, 1312-1337,
585 doi:10.1002/2016JD025987, 2017.
- 586 Lelieveld, J. and Dentener, F. J.: What controls tropospheric ozone?, *J. Geophys. Res.*, 105, 3531-3551,
587 doi:10.1029/1999JD901011, 2000.
- 588 Lin, M., Fiore, A. M., Horowitz, L. W., Langford, A. O., Oltmans, S. J., Tarasick, D., and Reider, H. E.: Climate
589 variability modulates western US ozone air quality in spring via deep stratospheric intrusions, *Nature*
590 *Communications*, 6, 7105, doi:10.1038/ncomms8105, 2015.
- 591 Lin, S. J. and Rood, R. B.: Multidimensional flux form semi-Lagrangian transport schemes, *Mon. Weather Rev.*, 124,
592 2046-2070, [https://doi.org/10.1175/1520-0493\(1996\)124<2046:MFFSLT>2.0.CO;2](https://doi.org/10.1175/1520-0493(1996)124<2046:MFFSLT>2.0.CO;2), 1996.



- 593 Liu, X., Chance, K., Sioris, C. E., Spurr, R. J. D., Kurosu, T. P., Martin, R. V., and Newchurch, M. J.: Ozone profile
594 and tropospheric ozone retrievals from the Global Ozone Monitoring Experiment: Algorithm description and
595 validation, *J. Geophys. Res.-Atmos.*, 110, D20307, doi:10.1029/2005JD006240, 2005.
- 596 Liu, X., Bhartia, P. K., Chance, K., Spurr, R. J. D., and Kurosu, T. P.: Ozone profile retrievals from the Ozone
597 Monitoring Instrument, *Atmos. Chem. Phys.*, 10, 2521-2537, doi:10.5194/acp-10-2521-2010, 2010.
- 598 Lock, A. P., Brown, A. R., Bush, M. R., Martin, G. M., and Smith, R. N. B.: A new boundary layer mixing scheme.
599 Part I: Scheme description and single-column model tests, *Mon. Weather Rev.*, 128, 3187-3199,
600 [https://doi.org/10.1175/1520-0493\(2000\)128<3187:ANBLMS>2.0.CO;2](https://doi.org/10.1175/1520-0493(2000)128<3187:ANBLMS>2.0.CO;2), 2000.
- 601 Martin, R.: Satellite remote sensing of surface air quality, *Atmos. Environ.*, 42, 7823-7843
602 <http://dx.doi.org/10.1016/j.atmosenv.2008.07.018>, 2008.
- 603 McPeters, R. D., Labow, G. J., and Logan, J. A.: Ozone climatological profiles for satellite retrieval algorithms, *J.*
604 *Geophys. Res.*, 112, D05308, doi:10.1029/2005JD006823, 2007.
- 605 Molod, A., Takacs, L. L., Suarez, M. J., Bacmeister, J. T., Song, I.-S., and Eichmann, A.: The GEOS-5 Atmospheric
606 General Circulation Model: Mean Climate and Development from MERRA to Fortuna, NASA Tech. Memo.
607 104606, 28, Tech. Rep. Series on Global Modeling and Data Assimilation, edited by: Suarez, M. J., 117 pp., 2012.
- 608 Natraj, V., Liu, X., Kulawik, S., Chance, K., Chatfield, R., Edwards, D. P., Eldering, A., Francis, G., Kurosu, T.,
609 Pickering, K., Spurr, R., and Worden, H.: Multi-spectral sensitivity studies for the retrieval of tropospheric and
610 lowermost tropospheric ozone from simulated clear-sky GEO-CAPE measurements, *Atmos. Environ.*, 45, 7151-
611 7165, 2011.
- 612 Ott, L. E., Duncan, B. N., Thompson, A. M., Diskin, G., Fasnacht, Z., Langford, A. O., Lin, M., Molod, A. M., Nielsen,
613 J. E., Pusede, S. E.; et al.: Frequency and impact of summertime stratospheric intrusions over Maryland during
614 DISCOVER-AQ (2011): New evidence from NASA's GEOS-5 simulations, *J. Geophys. Res. Atmos.*, 121,
615 doi:10.1002/2015JD024052, 2016.
- 616 Putman, W. M. and Lin, S.-J.: Finite-volume transport on various cubed-sphere grids, *J. Comput. Phys.*, 227, 55-78,
617 doi:10.1016/j.jcp.2007.07.022, 2007.
- 618 Rienecker, M. M., Suarez, M. J., Todling, R., Bacmeister, J., Takacs, L., Liu, H.-C., Gu, W., Sienkiewicz, M., Koster,
619 R. D., Gelaro, R., Stajner, I., and Nielsen, J. E.: The GEOS-5 Data Assimilation System – Documentation of
620 Versions 5.0.1, 5.1.0, and 5.2.0, NASA/TM-2008-104606, 27, 101 pp, 2008.
- 621 Rodgers, C. D.: *Inverse Methods for Atmospheric Sounding*, World Scientific, River Edge, New Jersey, 2000.
- 622 Sauvage, B., Martin, R. V., van Donkelaar, A., Liu, X., Chance, K., Jaeglé, L., Palmer, P. I., Wu, S., and Fu, T.-M.:
623 Remote sensed and in situ constraints on processes affecting tropical tropospheric ozone, *Atmos. Chem. Phys.*, 7,
624 815-838, <https://doi.org/10.5194/acp-7-815-2007>, 2007.
- 625 Stohl, A., Bonasoni, P., Cristofanelli, P., Collins, W., Feichter, J., Frank, A., Forster, C., Gerasopoulos, E., Gaggeler,
626 H., James, P., Kentarchos, T., Kromp-Kolb, H., Kruger, B., Land, C., Meloan, J., Papayannis, A., Priller, A.,
627 Seibert, P., Sprenger, M., Roelofs, G. J., Scheel, H. E., Schnabel, C., Siegmund, P., Tobler, L., Trickl, T., Wernli,
628 H., Wirth, V., Zanis, P., and Zerefos, C.: Stratosphere-troposphere exchange: A review, and what we have learned
629 from STACCATO, *J. Geophys. Res.*, 108(D12), 8516, doi:10.1029/2002JD002490, 2003.



- Sullivan, J. T., McGee, T. J., Thompson, A. M., Pierce, R. B., Sumnicht, G. K., Twigg, L. W., Eloranta, E., and Hoff, R. M.: Characterizing the lifetime and occurrence of stratospheric-tropospheric exchange events in the rocky mountain region using high-resolution ozone measurements, *J. Geophys. Res. Atmos.*, 120, 12410-12424, doi:10.1002/2015JD023877, 2015a.
- Sullivan, J. T., McGee, T. J., Leblanc, T., Sumnicht, G. K., and Twigg, L. W.: Optimization of the GSFC TROPOZ DIAL retrieval using synthetic lidar returns and ozonesondes – Part 1: Algorithm validation, *Atmos. Meas. Tech.*, 8, 4133-4143, doi:10.5194/amt-8-4133-2015, 2015b.
- U.S. Environmental Protection Agency: Air Quality Criteria for Ozone and Related Photochemical Oxidants (2006 Final), U.S. Environmental Protection Agency, Washington, DC, EPA/600/R-05/004aF-cF, 2006.
- U.S. Environmental Protection Agency. National Ambient Air Quality Standards for Ozone - Final Rule, Federal Register 80, 65292-65468, available at <https://www.gpo.gov/fdsys/pkg/FR-2015-10-26/pdf/2015-26594.pdf>, 2015.
- Vu, K. T., Dingle, J. H., Bahreini, R., Reddy, P. J., Apel, E. C., Campos, T. L., DiGangi, J. P., Diskin, G. S., Fried, A., Herndon, S. C., Hills, A. J., Hornbrook, R. S., Huey, G., Kaser, L., Montzka, D. D., Nowak, J. B., Pusede, S. E., Richter, D., Roscioli, J. R., Sachse, G. W., Shertz, S., Stell, M., Tanner, D., Tyndall, G. S., Walega, J., Weibring, P., Weinheimer, A. J., Pfister, G., and Flocke, F.: Impacts of the Denver Cyclone on regional air quality and aerosol formation in the Colorado Front Range during FRAPPÉ 2014, *Atmos. Chem. Phys.*, 16, 12039-12058, doi:10.5194/acp-16-12039-2016, 2016.
- Wang, Y. H., Jacob, D. J., and Logan, J. A.: Global simulation of tropospheric O₃-NO_x-hydrocarbon chemistry: 1. Model formulation, *J. Geophys. Res.*, 103, 10713-10725, doi:10.1029/98JD00158, 1998.
- Wargan, K., Pawson, S., Olsen, M. A., Witte, J. C., Douglass, A. R., Ziemke, J. R., Strahan, S. E., and Nielsen, J. E.: The global structure of upper troposphere-lower stratosphere ozone in GEOS-5: A multiyear assimilation of EOS Aura data, *J. Geophys. Res. Atmos.*, 120, 2013-2036, doi:10.1002/2014JD022493, 2015.
- Worden, H. M., Logan, J. A., Worden, J. R., Beer, R., Bowman, K., Clough, S. A., Eldering, A., Fisher, B. M., Gunson, M. R., Herman, R. L., Kulawik, S. S., Lampel, M. C., Luo, M., Megretskaia, I. A., Osterman, G. B., and Shephard, M. W.: Comparisons of Tropospheric Emission Spectrometer (TES) ozone profiles to ozonesondes: Methods and initial results, *J. Geophys. Res.*, 112, D03309, doi:10.1029/2006JD007258, 2007.
- Wu, W.-S., Purser, R. J., and Parrish, D. F.: Three-dimensional variational analysis with spatially inhomogeneous covariances, *Mon. Wea. Rev.*, 130, 2905-2916, [https://doi.org/10.1175/1520-0493\(2002\)130<2905:TDVAWS>2.0.CO;2](https://doi.org/10.1175/1520-0493(2002)130<2905:TDVAWS>2.0.CO;2), 2002.
- Zoogman, P., Jacob, D. J., Chance, K., Liu, X., Lin, M., Fiore, A., and Travis, K.: Monitoring high-ozone events in the US Intermountain West using TEMPO geostationary satellite observations, *Atmos. Chem. Phys.*, 14, 6261-6271, <https://doi.org/10.5194/acp-14-6261-2014>, 2014.
- Zoogman, P., Liu, X., Suleiman, R., Pennington, W., Flittner, D., Al-Saadi, J., Hilton, B., Nicks, D., Newchurch, M., Carr, J., Janz, S., Andraschko, M., Arola, A., Baker, B., Canova, B., Miller, C. C., Cohen, R., Davis, J., Dussault, M., Edwards, D., Fishman, J., Ghulam, A., Abad, G. G., Grutter, M., Herman, J., Houck, J., Jacob, D., Joiner, J., Kerridge, B., Kim, J., Krotkov, N., Lamsal, L., Li, C., Lindfors, A., Martin, R., McElroy, C., McLinden, C., Natraj,



667 V., Neil, D., Nowlan, C., O'Sullivan, E., Palmer, P., Pierce, R., Pippin, M., Saiz-Lopez, A., Spurr, R., Szykman,
668 J., Torres, O., Veefkind, J., Veihelmann, B., Wang, H., Wang, J., and Chance, K.: Tropospheric emissions:
669 Monitoring of pollution (TEMPO), J. Quant. Spectrosc. Ra., 186, 17-39,
670 <https://doi.org/10.1016/j.jqsrt.2016.05.008>, 2017.



671 Tables

672 **Table 1. Information about the TOLNet systems applied during this study.**

System Name	Latitude (°N)	Longitude (°W)	Elevation (m) ^a	# of observations ^b
TROPOZ	40.6	105.1	1569.0	21
JPL TMF	34.4	117.7	2285.0	26 ^c
RO3QET	34.7	86.6	206.0	12 ^d

673 ^aElevation of the topography above sea level.

674 ^bNumber of days of lidar observations between July - August 2014.

675 ^cJPL TMF lidar observations only taken during nighttime hours between July-August 2014.

676 ^dRO3QET lidar observations only taken from the surface to ~5 km between July-August 2014.

677



Table 2. Time-series evaluation of TB-Clim, GEOS-5 FP, MERRA2, and GEOS-Chem daily-averaged tropospheric and LMT column O_3 with the RO3QET, TROPOZ and JPL TMF lidars. The statistics include correlation (R), mean bias, bias standard deviation (1σ), and root mean squared error (RMSE).

RO3QET	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-5 km)</i>				
Correlation (R)	-0.09	0.23	-0.10	0.61
Bias $\pm 1\sigma$ (ppb)	3.7 ± 6.0	2.8 ± 5.6	-0.7 ± 5.8	1.7 ± 4.2
RMSE (ppb)	6.81	6.14	5.61	4.34
<i>LMT Column O_3 (0-2 km)</i>				
Correlation (R)	-0.68	0.03	-0.19	0.83
Bias $\pm 1\sigma$ (ppb)	2.9 ± 9.7	-2.9 ± 8.5	-4.9 ± 8.0	-1.3 ± 4.4
RMSE (ppb)	9.75	8.65	9.06	4.39
TROPOZ	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-10 km)</i>				
Correlation (R)	-0.09	0.26	0.38	0.82
Bias $\pm 1\sigma$ (ppb)	2.2 ± 9.7	3.3 ± 10.0	-4.6 ± 9.1	2.4 ± 6.0
RMSE (ppb)	9.73	10.33	9.99	6.30
<i>LMT Column O_3 (0-2 km)</i>				
Correlation (R)	-0.15	-0.09	-0.18	0.47
Bias $\pm 1\sigma$ (ppb)	-11.1 ± 7.5	-4.4 ± 7.3	-7.4 ± 7.4	-6.7 ± 6.2
RMSE (ppb)	13.23	8.43	10.33	8.93
JPL TMF	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-10 km)</i>				
Correlation (R)	-0.35	0.76	0.80	0.72
Bias $\pm 1\sigma$ (ppb)	0.3 ± 18.7	-5.0 ± 13.8	-10.6 ± 13.4	-0.5 ± 14.6
RMSE (ppb)	18.38	14.41	16.86	14.29
<i>LMT Column O_3 (0-2 km)</i>				
Correlation (R)	-0.53	-0.21	0.22	0.49
Bias $\pm 1\sigma$ (ppb)	3.3 ± 13.6	-2.4 ± 12.7	-4.0 ± 11.7	0.9 ± 10.4
RMSE (ppb)	13.72	12.68	12.14	10.24



Table 3. Time-series evaluation of the TB-Clim, GEOS-5 FP, MERRA2, and GEOS-Chem hourly-averaged tropospheric and LMT column O_3 with the RO3QET, TROPOZ and JPL TMF lidars. The statistics include correlation (R), mean bias, bias standard deviation (1σ), and root mean squared error (RMSE).

RO3QET	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-5 km)</i>				
Correlation (R)	-0.54	-0.55	-0.51	0.68
Bias $\pm 1\sigma$ (ppb)	3.5 ± 1.4	2.6 ± 1.6	-1.2 ± 1.5	2.1 ± 1.1
RMSE (ppb)	3.77	2.98	1.86	2.37
<i>LMT Column O_3 (0-2 km)</i>				
Correlation (R)	0.20	0.55	-0.43	0.76
Bias $\pm 1\sigma$ (ppb)	1.9 ± 3.9	-3.3 ± 3.6	-5.9 ± 4.0	0.3 ± 2.6
RMSE (ppb)	4.20	4.73	7.04	2.45
TROPOZ	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-10 km)</i>				
Correlation (R)	-0.07	-0.38	-0.56	0.78
Bias $\pm 1\sigma$ (ppb)	2.6 ± 2.5	3.3 ± 2.6	-5.1 ± 3.2	2.2 ± 1.7
RMSE (ppb)	3.57	4.17	6.00	2.74
<i>LMT Column O_3 (0-2 km)</i>				
Correlation (R)	0.26	0.76	0.67	0.92
Bias $\pm 1\sigma$ (ppb)	-12.6 ± 6.9	-7.5 ± 6.6	-9.6 ± 6.9	-7.7 ± 4.8
RMSE (ppb)	14.25	9.91	11.70	9.01



Table 4. Time-series evaluation of daily-averaged X_r predictions using the TB-Clim, GEOS-5 FP, MERRA2, and GEOS-Chem data as a priori information in theoretical TEMPO retrievals of tropospheric and LMT column O_3 values with RO3QET, TROPOZ and JPL TMF lidars. The statistics include mean bias, bias standard deviation (1σ), root mean squared error (RMSE), and the number of occurrences where error exceeds 10 ppb.

RO3QET	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-5 km)</i>				
Bias $\pm 1\sigma$ (ppb)	1.4 ± 2.3	1.3 ± 2.7	-0.2 ± 2.5	1.0 ± 2.0
RMSE (ppb)	2.66	2.91	2.43	2.17
10 ppb error exceedance	0	0	0	0
<i>LMT Column O_3 (0-2 km)</i>				
Bias $\pm 1\sigma$ (ppb)	0.2 ± 6.1	-3.8 ± 5.5	-3.4 ± 5.1	-2.2 ± 2.5
RMSE (ppb)	5.88	6.44	5.97	3.26
10 ppb error exceedance	1	3	2	0
TROPOZ	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-10 km)</i>				
Bias $\pm 1\sigma$ (ppb)	-0.9 ± 4.2	-0.6 ± 4.8	-2.2 ± 4.4	-0.5 ± 2.7
RMSE (ppb)	4.21	4.72	4.85	2.66
10 ppb error exceedance	1	1	2	0
<i>LMT Column O_3 (0-2 km)</i>				
Bias $\pm 1\sigma$ (ppb)	-11.4 ± 6.2	-6.4 ± 6.3	-5.1 ± 5.9	-4.8 ± 4.8
RMSE (ppb)	12.95	8.85	7.67	6.71
10 ppb error exceedance	10	6	4	3
JPL TMF	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-10 km)</i>				
Bias $\pm 1\sigma$ (ppb)	-0.2 ± 4.0	-0.8 ± 3.1	-1.7 ± 3.0	-0.3 ± 3.3
RMSE (ppb)	3.97	3.14	3.42	3.29
10 ppb error exceedance	1	1	1	1
<i>LMT Column O_3 (0-2 km)</i>				
Bias $\pm 1\sigma$ (ppb)	3.1 ± 14.8	1.9 ± 13.7	4.8 ± 12.6	1.0 ± 12.7
RMSE (ppb)	14.87	13.57	13.27	12.54
10 ppb error exceedance	9	8	10	6

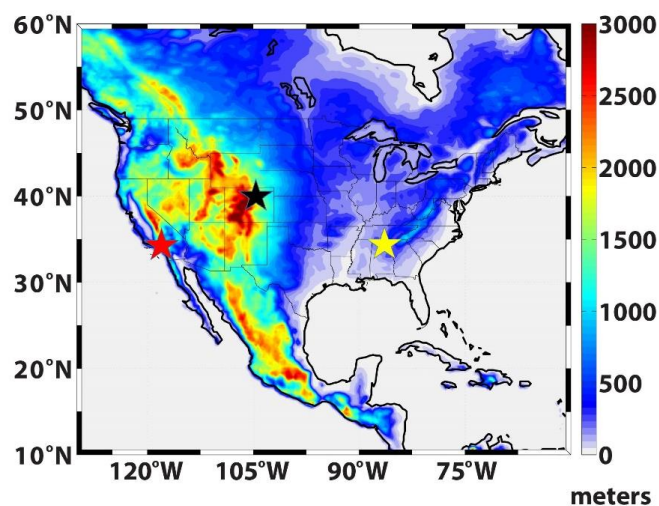


Table 5. Time-series evaluation of hourly-averaged TOLNet observations and X_r predictions using the TB-Clim, GEOS-5 FP, MERRA2, and GEOS-Chem data as a priori information in theoretical TEMPO retrievals of tropospheric and LMT column O_3 values at the location of RO3QET (07 August, 2014) and TROPOZ (22 July, 2014). The statistics include mean, min/max, and standard deviation (1σ) from observations and theoretical TEMPO retrievals.

RO3QET 07 August, 2014	TOLNet	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-5 km)</i>					
Mean (ppb)	60.7	59.8	59.5	59.0	59.5
Max/Min (ppb)	67.5/56.4	64.7/56.8	64.1/56.9	63.8/56.1	65.1/55.5
Std. Dev. (ppb)	3.62	2.63	2.35	2.55	3.18
<i>LMT Column O_3 (0-2 km)</i>					
Mean (ppb)	65.2	56.5	53.4	53.1	62.1
Max/Min (ppb)	79.4/54.3	62.6/52.5	59.4/49.8	59.4/48.8	70.6/54.6
Std. Dev. (ppb)	9.27	3.41	3.33	3.67	5.38
TROPOZ 22 July, 2014	TOLNet	TB-Clim	GEOS-5 FP	MERRA2	GEOS-Chem
<i>Tropospheric Column O_3 (0-10 km)</i>					
Mean (ppb)	50.5	52.4	52.2	50.7	50.3
Max/Min (ppb)	55.8/46.3	55.7/49.2	55.5/49.0	53.3/47.7	53.3/47.3
Std. Dev. (ppb)	3.25	2.60	2.52	2.06	2.40
<i>LMT Column O_3 (0-2 km)</i>					
Mean (ppb)	75.0	44.3	49.9	51.2	56.3
Max/Min (ppb)	97.0/58.6	47.5/41.3	54.3/45.6	54.9/47.3	66.4/47.8
Std. Dev. (ppb)	12.77	2.27	2.96	2.81	5.93



701 Figures



702

703 **Figure 1.** Location of the GSFC TROPOZ (black star), JPL TMF (red star), and the UAH RO3QET (yellow
 704 star) TOLNet systems during the summer of 2014. The locations are overlaid on the topographic heights
 705 (meters) from the GEOS-5 model.

706

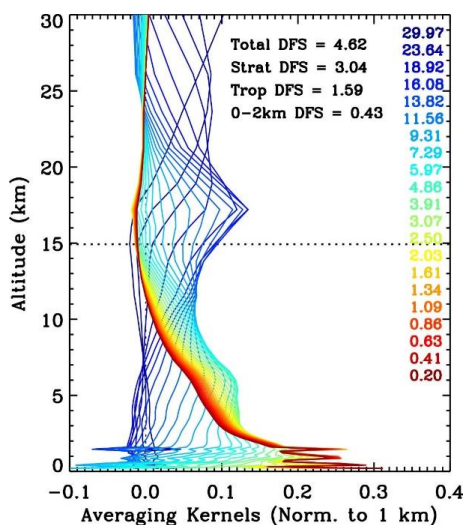
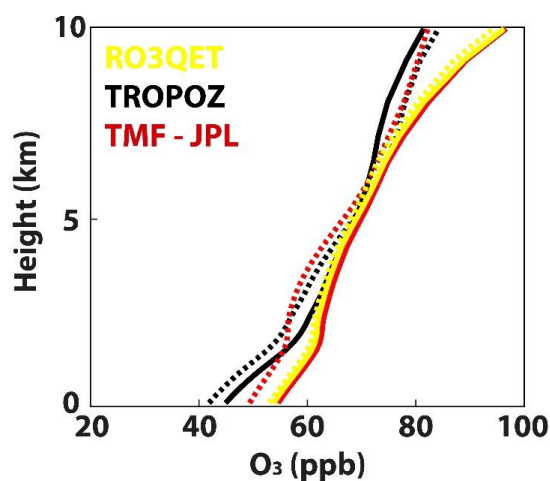


Figure 2. Simulated TEMPO O₃ retrieval AK matrix (normalized to 1 km layer) from joint UV+VIS measurements (290-345 nm, 540-650 nm) from the surface to 30 km above ground level used at the UAH TOLNet site during August at 20 UTC. The AK lines are for individual vertical levels (km above ground level), with the colors ranging from red to blue representing vertical levels from surface air to ~30 km. The legend presents the DFS for the total (Total), stratosphere (Strat), troposphere (Trop), and 0-2 km columns.



714
 715 **Figure 3. Monthly-averaged vertical profiles of O_3 (ppb) from TB-Clim data at the location of the RO3QET**
 716 **(yellow lines), TROPOZ (black lines), and JPL TMF (red lines) TOLNet systems for July (solid lines) and**
 717 **August (dashed lines). The monthly-averages are derived using the hourly TB-Clim data during the hours/days**
 718 **of observations obtained at each location.**
 719

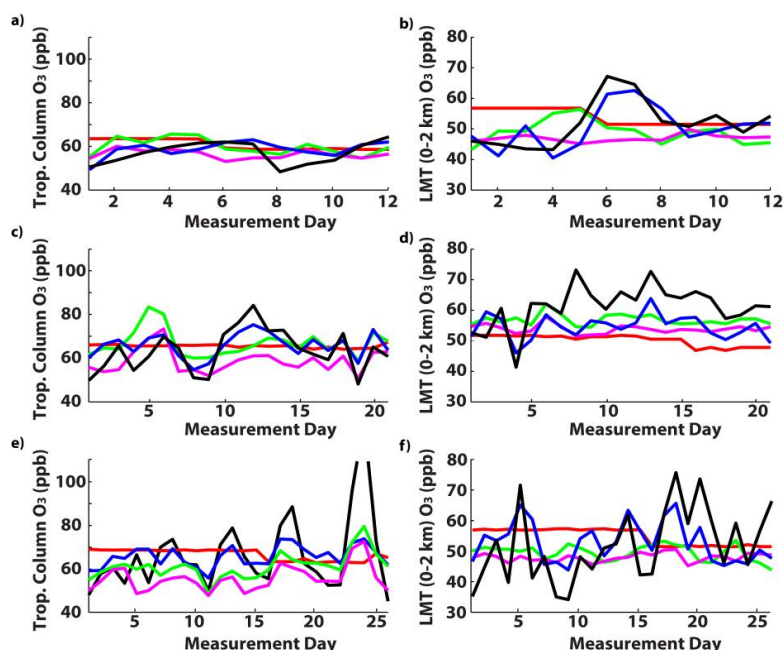


Figure 4. Time-series of daily-averaged tropospheric column (0-10 km) O₃ (ppb) from TB-Clim (red line), GEOS-5 FP (green line), MERRA2 (magenta line), and GEOS-Chem (blue line) compared to TOLNet (black line) at the locations of a) RO3QET, c) TROPOZ, and e) JPL TMF. Panels b), d), and f) are similar but for the comparison of LMT column (0-2 km) O₃.

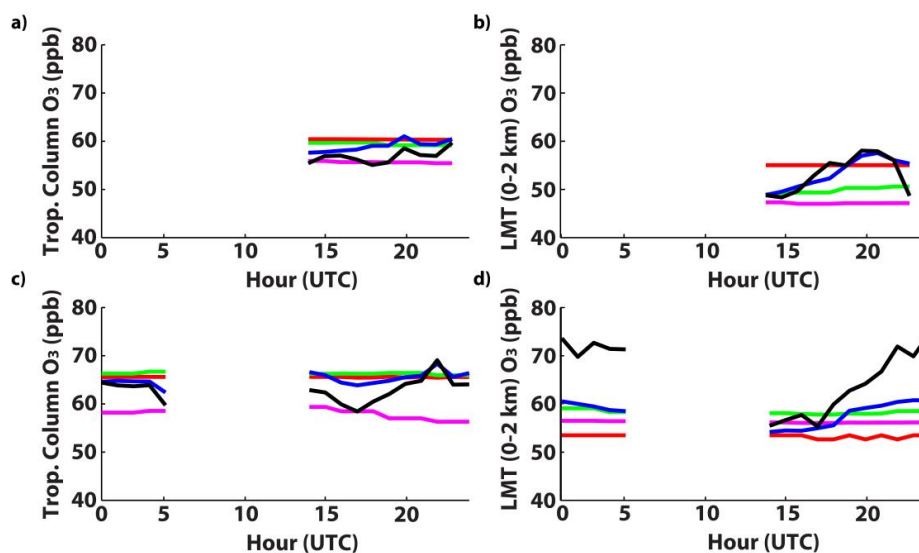
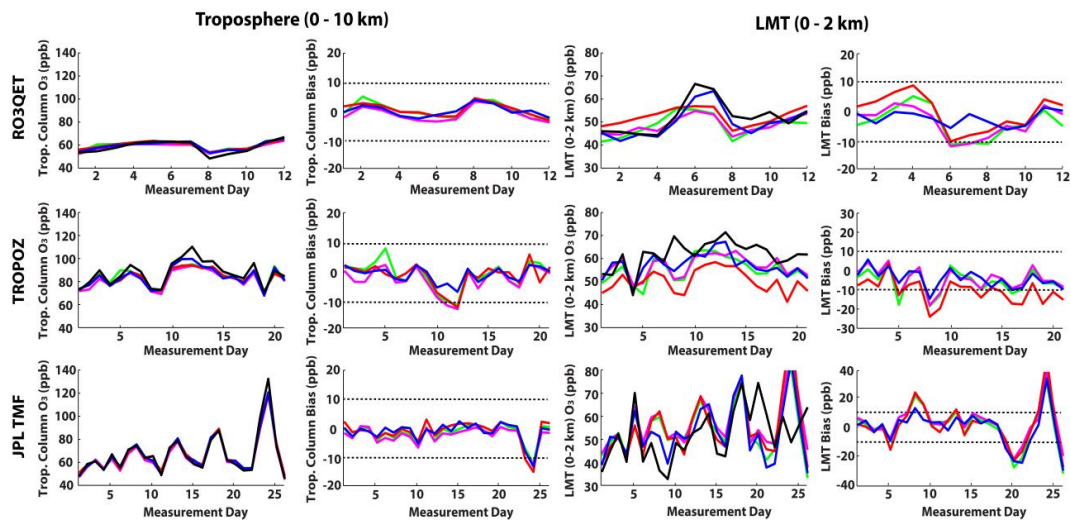


Figure 5. Diurnal time-series of hourly-averaged tropospheric column (0-10 km) O₃ (ppb) from TB-Clim (red line), GEOS-5 FP (green line), MERRA2 (magenta line), and GEOS-Chem (blue line) compared to TOLNet (black line) at the locations of a) RO3QET and c) TROPOZ. Panels b) and d) are similar but for the comparison of LMT column (0-2 km) O₃. The times of missing data are hours where no TOLNet observations were taken during the summer of 2014.



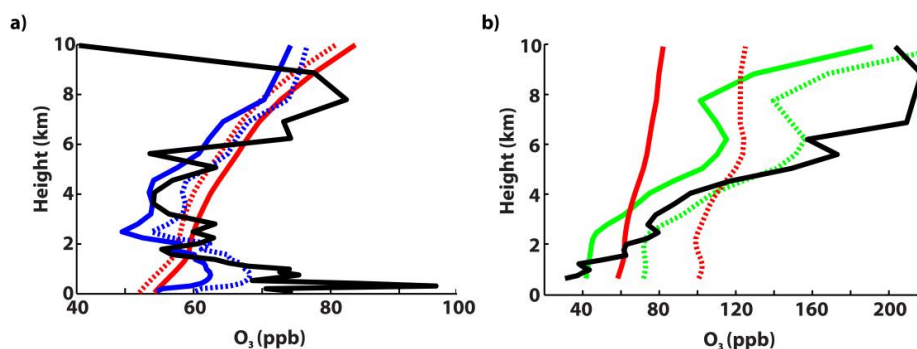
733



734

735 **Figure 6.** Time-series of daily-averaged tropospheric and LMT column X_r and bias values (ppb) when using
736 TB-Clim (red line), GEOS-5 FP (green line), MERRA2 (magenta line), and GEOS-Chem (blue line) as the a
737 priori when compared to observed O_3 by TOLNet (black line) at the locations of RO3QET (top row), TROPOZ
738 (middle row), and JPL TMF (bottom row). The dashed black lines represent the 10 ppb precision/accuracy
739 requirement for TEMPO O_3 retrievals.

740



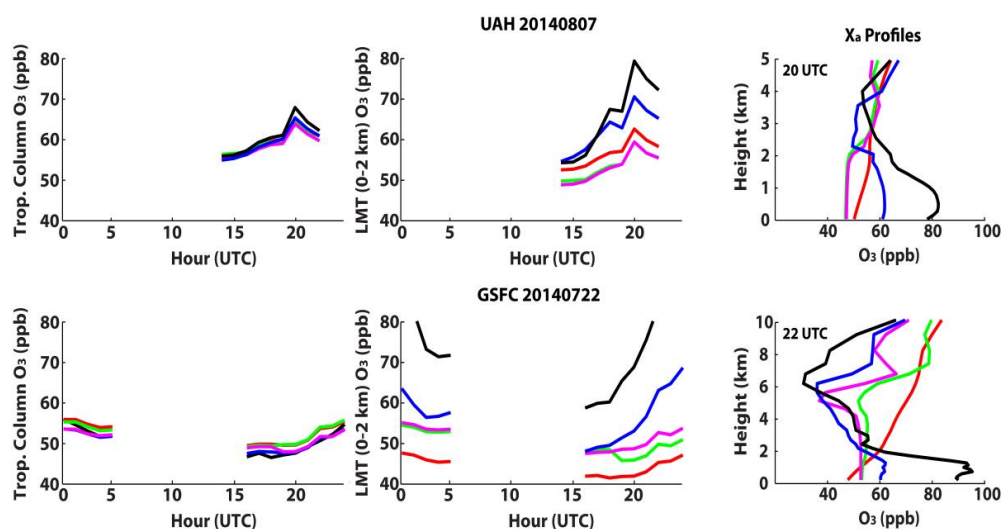
741

742 **Figure 7. Vertical profiles of a) daily-averaged X_a (solid line) and X_r (dashed line) O_3 values when applying**
 743 **TB-Clim (red line) and GEOS-Chem (blue line) as a priori information in TEMPO retrievals compared to**
 744 **TOLNet (black line) at the locations of the JPL TMF lidar on 08 July, 2014. Panel b) shows daily-averaged X_a**
 745 **and X_r O_3 values when applying TB-Clim (red line) and GEOS-5 FP (green line) as a priori information in**
 746 **TEMPO retrievals compared to TOLNet (black line) at the locations of the JPL TMF lidar on 21 August, 2014.**

747



748



749

750 **Figure 8.** Diurnal time-series of hourly-averaged tropospheric (0-10 km) and LMT (0-2 km) column X_r , O_3
 751 (ppb) values with a priori from TB-Clim (red line), GEOS-5 FP (green line), MERRA2 (magenta line), and
 752 GEOS-Chem (blue line) compared to TOLNet (black line) at the locations of RO3QET location on 07 August
 753 2014 (top row) and TROPOZ on 22 July 2014 (bottom row). The hourly-averaged a priori vertical profiles are
 754 also presented (right column) along with TOLNet (black line) for the hour of largest LMT O_3 observed by
 755 TOLNet in the time-series.

756