1 Instrument Artifacts Lead to Uncertainties in Parameterizations of Cloud

2 Condensation Nucleation

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6 Abstract

8	The concentrations of cloud condensation nuclei (CCN) modulate cloud properties, rainfall
9	location and intensity, and climate forcings. This work assesses uncertainties in CCN
10	measurements and the apparent hygroscopicity parameter (κ_{app}) which is widely used to represent
11	CCN populations in climate models. CCN measurements require accurate operation of three
12	instruments: the CCN instrument, the differential mobility analyzer (DMA), and the condensation
13	particle counter (CPC). Assessment of DMA operation showed that varying the ratio of aerosol to
14	sheath flow from 0.05 to 0.30 resulted in discrepancies between the κ_{app} values calculated from
15	CCN measurements and the literature value. Discrepancies were found to increase from effectively
16	zero to 0.18 for sodium chloride, and from effectively zero to 0.08 for ammonium sulfate. The
17	ratio of excess to sheath flow was also varied, which shifted the downstream aerosol distribution
18	towards smaller particle diameters (for excess flow < sheath flow) or larger particle diameters (for
19	excess flow > sheath flow) than predicted. For the CPC instrument, undercounting occurred at
20	high concentrations, resulting in calculated κ_{app} lower than the literature values. Lastly,
21	undercounting by CCN instruments at high concentration was also assessed, taking the effect of
22	supersaturation on counting efficiency into account. Under recommended operating conditions,
23	the combined DMA, CPC, and CCN uncertainties in κ_{app} are 1.1 % or less for 25 to 200 nm
24	diameter aerosols.

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27 **1. Introduction**

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29 Aerosol-cloud interactions represent a major uncertainty in current predictions of the Earth's climate 30 (IPCC, 2013). According to well-known Köhler theory, an aerosol's potential to catalyze cloud droplet 31 formation by activating as a cloud condensation nucleus (CCN) depends on its physical and chemical 32 properties. For any given composition, the CCN activation potential of an aerosol increases as its 33 diameter decreases. While the relationship between aerosol diameter and CCN activation is straightforward, the effect of composition on an aerosol's ability to participate in cloud formation is 34 35 more complex (Petters and Kreidenweis, 2013; Ovadnevaite et al., 2011). Predicting the cloud forming capacity of various air masses based on the properties of the aerosol they contain is essential for 36 37 evaluating relative contributions from pollution, continental background and marine aerosol sources (Brooks and Thornton, 2018; Carslaw et al., 2013). Long-term CCN measurements are available from 38 numerous locations globally (Schmale et al., 2018). However, understanding regional and temporal 39 40 variability in CCN populations requires the ability to assess whether observed differences reflect true 41 physical differences or simply variations in CCN sampling strategies.

42

43 Parameterizations of CCN activity which accurately prescribe CCN measurements are needed for 44 climate models, cloud resolving models, and air quality predictions (Betancourt and Nenes, 2014; 45 Betancourt et al., 2013; Chang et al., 2017; Crosbie et al., 2015; Karydis et al., 2012; Kawecki and Steiner, 2018). One parameterization was designed to represent the cloud droplet activation potential 46 47 ambient aerosol masses of unknown composition with a single variable, kappa (κ) based on the dry 48 aerosol's hygroscopicity, or ability to uptake water and form a solution droplet (Petters and 49 Kreidenweis, 2007). Various names and abbreviations have been given to κ throughout the literature: "hygroscopicity parameter", "single hygroscopicity parameter", κ (Petters and Kreidenweis, 2007; 50

51	Carrico et al., 2008; Asa-Awuku et al., 2010; Moore et al., 2012b); "CCN-derived κ ", κ_{CCN} (Carrico et
52	al., 2008; Petters and Kreidenweis, 2007); and the "apparent hygroscopicity parameter" κ_{app} (Sullivan
53	et al., 2009; Collins et al., 2016; Petters and Kreidenweis, 2013). The term apparent hygroscopicity is
54	favored by many because it emphasizes that fact that while CCN activation can often be predicted
55	accurately by hygroscopic water uptake, they are different physical processes. It is possible for a
56	compound to have high intrinsic hygroscopicity and low apparent hygroscopicity if it is poorly soluble
57	in water (Sullivan et al., 2009).
58	
59	Parameterizations of hygroscopicity that pre-date Petters and Kreidenweiss 2007 exist as well. Winkler
60	1973 developed an equation for approximating the growth of an aerosol particle with relative humidity,
61	based on the quantity and physical characteristics of the soluble species in the particle. Another
62	approximation for the relationship between the equilibrium size of a particle and relative humidity was
63	derived by Fitzgerald in 1975, in which the soluble fraction and composition of the soluble
64	<u>component(s) are taken into account.</u>
65	
66	Fitzgerald et al., 1982 derived a particle composition parameter using the mass fraction and physical
67	properties of soluble material in a particle. Kreidenweis et al., 2005 determined that the critical
68	activation diameter of dry aerosol particles can be calculated from simplified Köhler theory using the
69	physical properties of water and the solute in a solution droplet. This parameterization has been used in
70	CCN closure studies (Bougiatioti et al., 2009; Moore et al., 2011; Moore et al., 2012a). The earliest
71	prediction of CCN concentrations for specific particle diameters and hygroscopicity used this
72	parameterization as well (Mochida et al., 2006).
73	

74	Once calculated, hygroscopicity parameters are useful tools for comparing CCN field measurements
75	conducted in various regions and seasons and for making predictions about cloud formation, aerosol-
76	cloud interactions in weather, and climate models. Values of κ_{app} can be used to compare the CCN
77	results in field and laboratory studies, including sea spray aerosol. For example, aggregation of results
78	from several mesocosm experiments and marine field studies found submicron (30-100 nm) κ_{app} for sea
79	spray aerosol as low as 0.4 and as high 1.3 (Collins et al., 2016). Another study, which included a
80	survey of observational CCN data, reported that marine and continental aerosols could be described by
81	κ_{app} values of 0.7 \pm 0.2 and 0.3 \pm 0.1 respectively (Andreae and Rosenfeld, 2008).
82	
83	Several studies have examined the sensitivity of models to κ values derived from HTDMA

84 measurements. An analysis of the NASA Global Modeling Initiative Chemical Transport Model and the 85 GEOS-Chem CTM (Karydis et al., 2012) found that cloud droplet number concentration is sensitive to κ 86 in Arctic and remote regions, where background aerosol loadings are low. Another study (Betancourt 87 and Nenes, 2014) found that a ± 50 % uncertainty range in the κ of secondary organic aerosols and 88 particulate organic matter resulted in a cloud droplet number concentration uncertainty of up to 15 % 89 and 16 %, respectively. Updating precipitation models with lab-derived κ values for specific inorganic 90 and organic species may increase the accuracy of storm forecasts by providing better predictions of 91 intense precipitation (Kawecki and Steiner, 2018). In terms of climate, (Liu and Wang, 2010) found that 92 increasing the κ of primary organic aerosols from 0 to 0.1, and decreasing the κ of secondary organics 93 aerosols from 0.14 to 0.07, resulted in an uncertainty in global secondary aerosol indirect forcing of 0.4 Wm⁻² from pre-industrial times to present day. 94

95

96 The sensitivity of weather and climate models to hygroscopicity parameters demonstrates the need for 97 accurate measurements. In this study, we examine experimental uncertainties in CCN measurements and

- 98 the resulting uncertainties in determination of κ_{app} . Differences in report κ_{app} values may result from
- 99 experimental artifacts rather any actual differences in aerosol's ability to facilitate cloud formation. By
- 100 systematically quantifying sources of experimental error, this study provides a framework for
- 101 determining the significance of variations in CCN properties reported in multiple studies and defining
- 102 the operating conditions which minimize instrumental artifacts.

103 2. Background

104

The Köhler equation relates water vapor saturation <u>ratio</u> at the surface of a wet droplet, *s*, to its radius at
equilibrium (Rogers and Yau, 1989):

107

108
$$s = \left(1 - \frac{b}{r^3}\right) \exp\left(\frac{a}{r}\right) \tag{1a}$$

109

110
$$a = \frac{2\sigma_w M_w}{\rho_w RT}$$
(1b)

111

$$b = \frac{3im_s M_w}{4\pi\rho_w M_s} \tag{1c}$$

113

114 where *s* is the equilibrium saturation ratio of a solution droplet with radius r, σ_w is the surface tension of 115 water, M_w is the molecular weight of water, *R* is the ideal gas constant, *T* is temperature in Kelvin, ρ_w is 116 the density of water, and M_s is the molecular weight of the solute. The minimum saturation ratio that is 117 required for spontaneous droplet growth, s_{act} , is therefore:

118

119
$$s_{crit} = 1 + \sqrt{\frac{4a^3}{27b}}$$

(2)

120

121 Petters and Kreidenweis [2007] reformulated the Köhler equation as κ-Köhler theory:

122

123
$$s_{crit} = exp\left(\sqrt{\frac{4A^3}{27D_{act}^3\kappa_{app}}}\right)$$
(3a)

124 and

126
$$A = \frac{4\sigma_{lv}M_w}{_{RT}\rho_w} \tag{3b}$$

Where s_{crit} is the critical water vapor saturation <u>ratio</u>, D_{act} is the dry particle activation diameter and κ_{app} is the apparent hygroscopicity parameter. Solving for κ_{app} yields:

130

131
$$\kappa_{app} = \frac{4A^3 \sigma_{lv}^3}{27T^3 D_{act}^3 ln^2(s_{crit})}$$
(4)

132

The apparent hygroscopicity parameter can be calculated from experimental CCN results, where the dry diameter and water vapor saturation <u>ratio</u> are known. For a chosen aerosol diameter, the activated fraction is the ratio of the concentration aerosols that activate as CCN to the total aerosol concentration:

137
$$Activated \ fraction = \frac{CCN \ Concentration}{Aerosol \ Concentration}$$
(5)

138

Activated fraction data is fit with a sigmoid error function to determine the percent supersaturation at which 50 % of the particles have activated as CCN(activated fraction = 0.50), which is considered the operationally defined critical percent supersaturation SS_{crit} (Rose et al., 2008). The critical saturation ratio s_{crit} can then be determined and entered into Eq. (4) in order to calculate κ_{app} for the nearmonodisperse aerosol:

144

$$s_{crit} = 1 + \frac{SS_{crit}}{100} \tag{6}$$

146

147 Reporting κ_{app} as a function of diameter allows for the comparison of the cloud condensation nucleation 148 abilities of multimodal aerosol populations, without overlooking differences which arise due to aerosol 149 composition.

- 150
- 151 The apparent hygroscopicity parameter is related to chemical composition; therefore, the calculated κ_{app}
- 152 of a pure substance should be constant across CCN experiments. However, discrepancies between κ_{app}
- for a single chemical species have been observed. Experimental results for ammonium nitrate are
- 154 <u>inconsistent with reported values ranging from 0.577 $\leq \kappa_{app} \leq 0.753$ (Svenningsson et al., 2006).</u>
- Also, large ranges are often observed for organic compounds, such as glutaric acid ($0.054 \le \kappa_{app} \le$
- 156 0.16) and malonic acid (0.199 $\leq \kappa_{app} \leq 0.255$) (Koehler et al., 2006; Kumar et al., 2003; Hartz et al.,
- 157 <u>2006</u>). Below we evaluate potential sources of uncertainties in CCN measurements and the resulting
- 158 uncertainties in κ_{app} .

3. Artifacts derived from sized CCN measurements

101	
162	CCN measurements used for calculating apparent hygroscopicity from monodisperse aerosol require
163	accurate operation of three instruments: the CCN, the differential mobility analyzer (DMA), and the
164	condensational particle counter (CPC). The setup for laboratory CCN experiments is shown in Fig. 1.
165	First, a polydisperse population of aerosols is generated by an atomizer and dried using a desiccant tube
166	packed with silica gel. A near-monodisperse flow is obtained through size-selection in the DMA. The
167	flow is then split between a CPC (which measures aerosol concentration) and a CCN counter (which
168	measures the concentration of particles that activate as cloud condensation nuclei at a given percent
169	supersaturation). Instrument artifacts will first be assessed separately for the DMA, CPC, and CCN
170	counter. In the concluding section of the paper (and Fig. 10), the overall uncertainty due to the
171	combination of these is presented and discussed.
172	
173	We note that this study considers sized CCN measurements which may be used for the determination of
174	κ_{app} . In contrast, a number of earlier CCN studies conducted on the full ambient aerosol population
175	without sizing the aerosol (Jennings et al., 1996; Hudson and Xie, 1998; Modini et al., 2015; Duan et al.,
176	2017; Schmale et al., 2018; Leng et al., 2013). While useful, such studies do not produce the data
177	required for accurate determination of κ_{app} from the CCN measurements.
178	
 179	3.1 Artifacts derived from differential mobility analyzers
180	
181	3.1.1 DMA operation and electrical mobility
182	

183 Differential mobility analyzers used in atmospheric science include commercially available instruments 184 from Grimm Aerosol Technik, TSI Incorporated, and MSP Corporation. They have also been custom built by a number of research groups (Mei et al., 2011;Barmpounis et al., 2016;Jokinen and Makela, 185 1997;Seol et al., 2000). All models allow for the selection of particles through electrical mobility, the 186 ability of a particle to move through a medium (such as air) while acted upon by an electrical field. The 187 188 DMA size-selects near-monodisperse aerosol from a polydisperse aerosol source, as shown in Fig. 2 189 (modeled after the Vienna-type long Differential Mobility Analyzer from Grimm Technologies). The electrical mobility Z_p of a particle with mobility diameter d_m can be calculated according to: 190

191

192
$$Z_p = \frac{neC_C(d_m)}{3\pi\eta d_m} \tag{7}$$

193

where *n* is the number of charges on the particle (assumed to be one in this study), *e* is the elementary unit of charge, η is the gas dynamic viscosity, and $C_C(d_m)$ is the Cunningham slip correction factor:

197
$$C_{C}(d_{m}) = 1 + \frac{2\lambda}{d_{m}} \left(\alpha_{CC} + \beta_{CC} \exp\left[-\frac{\gamma_{CC}}{2\lambda/d_{m}} \right] \right)$$
(8)

198

where λ is the mean free path (DeCarlo et al., 2004). For the Vienna-type long Differential Mobility Analyzer from Grimm Technologies, Inc. considered here, $\alpha_{CC} = 1.246$, $\beta_{CC} = 0.42$, and $\gamma_{CC} = 0.86$ (Grimm Aerosol Technik, 2009).

202

203 Particle-laden flow enters the differential mobility analyzer through the aerosol inlet (flow Q_a), and travels 204 down the DMA column (inner radius r_1 , outer radius r_2) with the clean air sheath flow Q_{sh} . Positively-205 charged particles are attracted by the negatively-charged inner electrode, to which voltage V_0 has been applied. Ideally, selection of a voltage allows only particles of a specific mobility diameter to exit the DMA through the sample flow Q_s . All particles with larger diameter (lower Z_p) or smaller diameter (higher Z_p) will exit the DMA through the excess flow Q_e . In other words, Q_s would ideally <u>consist only</u> of aerosols with diameters equal to, or very nearly equal to, the selected diameter.

In reality, the aerosol flow that leaves the DMA through Q_s is polydisperse with a mobility distribution determined by instrumental parameters. A triangular approximation has been chosen as a model for this distribution, as particle inertia is negligible for the diameters considered in this study (Stratmann et al., 1997; Mamakos et al., 2007). The probability that a particle at the aerosol inlet will exit with the sampling flow is defined by transfer function $f(Z_p, Z_{p,mid})$:

216

217
$$f(Z_p, Z_{p,mid}) = \frac{\alpha_{TF}}{2\beta_{TF}} \left(\left| \frac{Z_p}{Z_{p,mid}} - (1 + \beta_{TF}) \right| + \left| \frac{Z_p}{Z_{p,mid}} - (1 - \beta_{TF}) \right| - 2 \left| \frac{Z_p}{Z_{p,mid}} - 1 \right| \right)$$
(9)

where $Z_{p,mid}$ is the midpoint mobility of the transfer function, and α_{TF} and β_{TF} are flow-derived constants, defined as:

220

$$\alpha_{TF} = \frac{Q_s + Q_a}{2Q_a} \tag{10a}$$

222 and

$$\beta_{TF} = \frac{Q_s}{Q_{sh}} \tag{10b}$$

224

The midpoint and half-width of the transfer function are respectively calculated according to: (Knutsonand Whitby, 1975)

227
$$Z_{p,mid} = \frac{Q_e + Q_{sh}}{4\pi L V_0} \ln\left(\frac{r_2}{r_1}\right)$$
(11a)

228 and

- 229
- 230

231

$$\Delta Z_p = \frac{Q_a}{2\pi L V_0} \ln\left(\frac{r_2}{r_1}\right) \tag{11b}$$

where L is the distance between the DMA inlet and outlet.

233

234 **3.1.2** κ_{app} artifacts arising from DMA flow ratios

235

Next we assess the ramifications of the DMA transfer function for the derived κ_{app} . A lognormal 236 237 theoretical aerosol number distribution was used to represent a polydisperse ambient aerosol population 238 (Fig. 3a). This distribution was converted to an electrical mobility distribution using Eq. (7) and Eq. (8), 239 assuming that the aerosols in the distribution were spherical and singly charged. From the distribution, a 240 series of single aerosol sizes were selected (25, 50, 100, and 200 nm diameter). For each aerosol size, 241 the resulting DMA transfer functions were calculated for seven cases using Eq. (9) and the various 242 parameters for DMA sheath, excess, aerosol, and sample flow listed in Table 1. These seven cases were 243 chosen to represent possible measurements scenarios that may be encountered in a CCN experiment. 244 The aerosol/sheath ratio is varied in Cases 1-4 in order to study the effects of chosen experimental 245 parameters. Sheath flow is predetermined in some DMAs (for example, the Grimm Vienna DMA 246 considered in this study), but can be varied in other instruments. The aerosol flow rate may also be 247 selected in an experiment. Cases 5-7 vary the excess/sheath ratio in order to take proper instrument 248 operation into account. The excess and sheath flow should be identical, but small discrepancies may 249 occur. 250

For example, the resulting DMA transfer functions for a 100 nm aerosol conditions constrained by Cases 1-4 are shown in Fig. 3b, where an increase in Q_a/Q_{sh} from 0.1 (black line) to 0.3 (green line) tripled the width of the number distribution, and decreasing Q_a/Q_{sh} to 0.05 (blue line) from 0.10 halved the width of the number distribution. The result of applying the transfer functions shown in Fig. 3b to the distribution in Fig. 3a is shown in Fig. 3c.

256

257 All downstream distributions for all seven DMA cases and all aerosol sizes are shown in Fig. S1 in the 258 Supplement. DMA Cases 1-4 represent experimental conditions in which the sheath and excess air flows are equal and the aerosol/sheath flow ratio is varied. As Q_a/Q_{sh} increases, the width of the 259 260 number distribution measured downstream of the DMA increases, while the midpoint diameter remains 261 constant. It was found that doubling the aerosol to sheath ratio doubled the width of the downstream 262 number distribution for 25, 50, 100, and 200 nm particles. For example, when selecting 200 nm 263 particles, increasing Q_a/Q_{sh} from 0.10 to 0.20 increased the downstream diameter range from 181-222 264 nm (a spread of 41 nm) to 167-250 nm (a spread of 87 nm). The particle diameter ranges that would be observed downstream of the DMA are summarized in Table 2. 265

266

To assess the variations in CCN properties resulting from DMA uncertainties the critical percent supersaturation were calculated for representative atmospheric aerosols. The value of SS_{crit} was calculated for each particle diameter using Eq. (3a), using literature values for apparent hygroscopicity of 0.61 for ammonium sulfate and 1.28 for sodium chloride (Clegg et al., 1998). It should be noted that this analysis considers two homogeneous aerosol distributions of hygroscopic salts. Real aerosol distributions tend to be mixtures of many species, and the shape of the number distribution can vary between species.

275 Note that in the absence of DMA diameter uncertainty, this single component aerosol population should 276 be characterized by a single κ_{ann} regardless of diameter. To test how uncertainties in DMA diameter translate to uncertainties in κ_{app} , the true critical saturation ratio s_{crit} was then put into Eq. (4) in order 277 to calculate the "perceived" κ_{app} for each diameter given the chosen transfer function from Cases 1-7. 278 279 For example, if 100 nm particles were selected from the DMA by the user, the transfer functions would allow larger and smaller particles to pass into the sample flow, as shown in Table 2. Particles with 280 diameter > 100 nm would be "perceived" to have higher apparent hygroscopicity than particles with 281 282 diameter < 100 nm, since the equilibrium vapor pressure over the surface of a particle decreases as its diameter increases (and as curvature decreases). Using Eq. (12), these diameter-specific "perceived" 283 284 κ_{app} values were volume-weighted, resulting in $\kappa_{app,theory}$:

- 285
- 286

$$\kappa_{app,theory} = \sum_{i} \epsilon_{i} \kappa_{i} \tag{12}$$

287

where ϵ_i is the volume fraction of aerosol of each diameter *i*, and κ_i is the perceived κ_{app} for each diameter (adapted from Petters and Kreidenweis [2007]). Results for ammonium sulfate and sodium chloride are shown in Fig. 4a. The critical saturation ratio was calculated from $\kappa_{app,theory}$ using Eq. (3a) for each case and converted to critical <u>percent</u> supersaturation. The results are compared to theoretical κ -Köhler theory curves for ammonium sulfate and sodium chloride generated using the literature κ_{app} for each compound (Fig. 4b).

- 294
- 295 Discrepancies between $\kappa_{app,theory}$ calculated in this study and literature values (hereon referred to as 296 " κ_{app} artifacts") are shown for both compounds in Fig. 4c-d.

The greatest κ_{ann} artifacts were found in DMA case 4 (where the aerosol/sheath ratio was the highest) 298 299 for both ammonium sulfate and sodium chloride aerosols. The artifacts for ammonium sulfate in DMA case 4 were 0.05-0.08, or 8-13 % of the literature value used for $\kappa_{app}^{(NH_4)_2SO_4}$, while the sodium chloride 300 301 artifacts in DMA case 4 were 0.11-0.18, or 9-14 % of the literature value used for κ_{app}^{NaCl} . Artifacts were also high for DMA case 6 ($-0.025 \le \kappa_{app,artifact}^{(NH_4)_2SO_4} \le -0.018$) and DMA case 7 ($0.016 \le$ 802 $\kappa_{app,artifact}^{(NH_4)_2SO_4} \leq 0.017$), where sheath and excess flow were unequal. This result demonstrates that 803 804 artifacts may still occur when low aerosol/sheath flow ratios are chosen (0.15 and 0.08 for DMA cases 6 805 and 7, respectively) due to small differences between sheath and excess flow rates (5% and 2% for 306 DMA cases 6 and 7, respectively). 307 308 κ_{app} artifacts were larger for sodium chloride ($-0.05 \le \kappa_{app,artifact}^{NaCl} \le 0.18, 4 - 14\%$ of κ_{app}^{NaCl}) than 309 for ammonium sulfate $(-0.03 \le \kappa_{app,artifact}^{(NH_4)_2SO_4} \le 0.08, 5 - 13\%$ of $\kappa_{app}^{(NH_4)_2SO_4}$) across the DMA cases. 310 311 As our results show, when two or more compounds are compared, the more hygroscopic compound will 312 have larger κ_{app} artifacts. 313 314 **3.1.3** Effect of double and triple charges on particles 815 816 During normal operation, the Grimm DMA employs a bipolar charger (also known as a neutralizer) to 317 charge aerosol particles through the capture of gaseous ions. The analysis in Section 3.1.2 assumes that 318 each particle carries a single (+1) charge. In reality, the methods used to charge particles prior to 819 entering a DMA may impart two, three, or more charges to individual particles (Fuchs, 1963). The 820 charge distribution resulting from a bipolar charger is roughly approximated using the Boltzmann law

321	(Keefe et al., 1959). However, the Boltzmann law assumes symmetric aerosol particle charging (equal		
322	concentrations of negatively and positively charged particles). Deviation from symmetric charging is		
323	observed in regions of high ionizations, and this deviation becomes more pronounced as particle size		
324	increases (Hoppel and Frick, 1990).		
325			
326	A more accurate estimation of stationary charge distribution has been calculated using an approximation		
327	formula for the charge distribution produced by a bipolar charger:		
328			
329	$f(k) = 10^{\left[\sum_{i=0}^{i=5} a_i(k)(\log_{10} D_{nm})^i\right]} $ (13)		
330			
331	where $f(k)$ is the fraction of particles carrying k charges, $a_i(k)$ are approximation coefficients		
332	determined using a least-squares regression analysis, and D_{nm} is the particle diameter in nanometers		
333	(Wiedensohler, 1988). The approximation coefficients only apply to particles with 0, ± 1 , and ± 2		
334	charges. In a separate study, Maricq et al., 2008 determined approximation coefficients for poly (α_{-}		
335	olefin) oligomer oil droplets with ± 1 , ± 2 , and ± 3 charges. The approximation coefficients reported by		
336	these two studies were in excellent agreement for particles with ± 1 and in weak agreement for ± 2		
337	charges (+2 and -2 charging efficiencies were overestimated by 50% and 100%, respectively).		
338	Therefore, this analysis will use the approximation coefficients from Wiedensolher, 1988 for particles		
339	with +1 and +2 charges, and the approximation coefficient for particles with +3 charge from Maricq et		
340	<u>al., 2008.</u>		
341			
342	In order to assess the impact of multiple charges on κ_{app} , Eq. (13) and the approximation coefficients		
343	from Wiedensolher, 1988 and Maricq et al., 2008 were used to calculate the charge distribution of the		
344	representative aerosol population shown in Fig. 3a. The resulting charge distribution is shown in Fig.		
I	18		

345	S2a. An increase in multiple charging is observed as particle diameter increases, though this is offset		
346	somewhat by the decrease in concentration with particle size above 50 nm.		
347			
348	It follows that aerosols incorrectly sized due to double and triple changing will be passed from the DMA		
349	to the CCN and result in an additional uncertainty in the CCN measurements. To illustrate this, activated		
350	fraction curves, were generated for 25, 50, 100, and 200 nm sodium chloride particle selection by the		
351	DMA (Fig. 5). The activation of sodium chloride is represented by sigmoid curves, where the midpoint		
352	of each activation curve is the κ-Köhler-derived critical supersaturation of sodium chloride, and the		
353	standard deviation of each curve is one-tenth of this value (consistent with the standard		
354	deviation/midpoint ratio observed from our instrument's ammonium sulfate CCN calibration data). For		
355	each particle diameter, D , the observed activated fraction, $AF_{D,weighted}^{SS}$, for each percent supersaturation		
356	SS was determined by weighting the activated fraction $AF_{D,i}^{SS}$ of each particle diameter/charge at that		
357	percent supersaturation, by the fraction of particles of that diameter:		
358			
359	$AF_{D,weighted}^{SS} = \sum_{i=1}^{i=3} \frac{\text{concentration of particles with charge i and diameter D}}{\text{concentration of particles with charge+1,+2,+3,and diameter D}} AF_{D,i}^{SS} $ (14)		
360			
361	The raw data shown in Fig. 5 (green curves) can be corrected for multiple charging by determining the		
362	fraction of particles with $> +1$ charge from the lower plateau in each plot (dashed lines). The adjusted		
363	<u>activated fraction for each percent supersaturation, $AF_{adjusted}$, is calculated using the equation:</u>		
364			
365	$AF_{adjusted} = \frac{AF_{raw} - AF_{plateau}}{1 - AF_{plateau}} $ (15)		
	$\frac{11}{1 - AF_{plateau}} $ (13)		
366			

367	where AF_{raw} is the raw activated fraction at that percent supersaturation, and $AF_{plateau}$ is the activated
368	fraction corresponding to the lower plateau (Rose, 2008). The adjusted activated fraction curves are
369	shown in Fig. 5 (blue curves). These are in good agreement with the theoretical κ-Köhler-derived
370	activation curves for sodium chloride (not shown).
371	
372	Critical supersaturation was determined for each diameter by calculating the percent supersaturation at
373	which the raw $AF_{D,weighted}^{SS} = 0.5$. These critical supersaturations are shown in Fig. 6a, and the
374	theoretical critical supersaturations calculated from κ -Kohler theory are shown for comparison. Eq. 4
375	was used to calculate apparent hygroscopicity for each particle diameter, shown in Fig. 6b. A dashed
376	line in Fig. 6b indicates the literature value for κ_{app}^{NaCl} . It is apparent that failing to account for multiply-
377	charged particle in the activated fraction curves shown in Fig. 5 leads to an overestimation of κ_{app} .
378	<u>Artifacts in κ_{app} are shown in Fig. 6c.</u>
379	
380	For the theoretical aerosol distribution used in this analysis (Fig. 3a), small, positive deviations from κ -
381	<u>Köhler theory and the literature value for κ_{app}^{NaCl} were observed (0.01 $\leq \kappa_{app,artifact}^{NaCl} \leq 0.04, 1 - 1$)</u>
382	3 % of κ_{app}^{NaCl}). As shown in the figure, κ_{app} artifacts resulting from unaccounted-for multiple charges
383	decrease with particle diameter for this theoretical aerosol population. Greater κ_{app} artifacts would be
384	expected for aerosol populations with more prevalent accumulation modes.
385	
386	The aerosol/sheath ratio within the DMA also modulates the effect of multiple charges on κ_{app} . As the
387	aerosol/sheath ratio increases, the transfer function broadens, allowing particles that are both larger and
388	smaller than the selected diameter to exit the DMA. This in turn broadens the CCN activated fraction
389	curve (Rose et al., 2008). The larger particles will activate as CCN at lower supersaturations than

390	particles with the selected diameter, resulting in an increase in the activated fraction plateau due to
391	multiple-charged particles and a further decrease in the determined SS _{crit} . Petters et al. 2007b showed
392	that CCN activated fraction curves are significantly skewed by multiply-charged particles when the
393	mode diameter of the aerosol population upstream of the DMA exceeds the critical diameter of the size-
394	selected particles. In an example CCN activated fraction curve, Rose et al. 2008 demonstrated that a 1:6
395	ratio of doubly-to-singly charged particles resulted in an underestimation of the critical activation
396	diameter by 2%. Zhao-Ze and Liang, 2014 also showed that multiply-charged particles can introduce
397	significant uncertainty into hygroscopicity calculations.
 398	
399	3.1.4 Additional artifacts resulting from DMA measurements
400	
401	Several additional factors that may impact experimental κ_{app} are beyond the scope of this study, but are
402	worth mentioning as they represent additional potential sources of error in some cases. First, volatile
403	aerosols may partially evaporate inside the DMA, resulting in a decrease in particle size exiting the
404	DMA. DMA sizing error due to aerosol volatility (defined as the ratio of sampled diameter to the
405	selected diameter) increases with volatility, though sizing error can be decreased by increasing the
406	sheath flow rate in the DMA. Conversely, hygroscopic aerosols may grow inside the DMA, resulting in
407	larger particles existing the DMA. Operationally, errors in DMA sizing due to hygroscopic growth can
408	be mitigated if aerosols entering the DMA inlet are in wet metastable states (higher aerosol RH at DMA
409	inlet), and if DMA sheath flow rates are kept low (Khlystov, 2014).
410	
411	Voltage shifts within the DMA (differences between the selected voltage and the actual voltage inside
412	the DMA) can lead to discrepancies between selected and sampled particle diameters. Voltage shifts
413	may result from a space-charge field generated by the motion of charges within the DMA. Particles

- 414 <u>charged by the bi-polar neutralizer will either be attracted towards or repelled away from the inner</u>
- 415 <u>column of the DMA, depending on whether they are positively or negatively charged</u>. This charge
- 416 <u>separation creates a space-charge field which shifts the actual voltage within the DMA from the selected</u>
- 417 voltage. The impact of the space-charge field on the midpoint and spread of the DMA transfer function
- 418 increases as particle mobility increases (as particle size decreases), and as particle concentration
- 419 increases (Alonso and Kousaka, 1996; Alonso et al., 2000; Alonso et al., 2001).

- 420 **4. Artifacts derived from condensation particle counters**
- 421

422 **4.1 CPC operation at low concentration**

423

The second instrument which must function accurately during CCN experiments is the condensation 424 425 particle counter. CPC performance is characterized by the maximum counting efficiency (which may be influenced by the working fluid in the instrument) and the 50 %-cut-off diameter (d_{50}), the particle 426 diameter at which 50 % counting efficiency is observed, both of which can vary between commercially 427 428 available models and even between individual CPCs (Heim et al., 2004). One study found that nbutanol CPCs (TSI, Inc. Models 3772, 3775, and 3776) exhibited smaller d_{50} for silver particles than 429 sodium chloride (3.3 $nm \le d_{50}^{Ag} \le 7.8 nm$ and 4.1 $nm \le d_{50}^{NaCl} \le 14.7 nm$), due to the more effective 430 431 condensation of n-butanol on silver particles (Hermann et al., 2007).

432

Maximum counting efficiencies in that study varied from 88.9 % to 100.3 %. Another comparison of nbutanol CPCs (TSI Inc. Models 3010 and 3022, Grimm Tech. Inc. Model 5.403) found $3.1 nm \le d_{50} \le$ 11.9 nm for sodium chloride aerosols (Heim et al., 2004). In another study, the counting efficiencies observed in measurements of tungsten oxide particles by different instruments of the same model (TSI 3025) varied from 88.9 % to 138.9 %, while $d_{50}^{WO_x}$ varied from 3.2 nm to 11.0 nm (Hameri et al., 2002).

While some issues can cause undercounting at all concentrations, the additional issue of uncounted
particles due to the arrival of more than one particle in the detector's field of view at any time arises
only at higher concentrations. The cut-off between "low" and "high" concentration is not exact and
varies between instruments. <u>CPC undercounting issues which arise even at relatively low concentrations</u>
(which one would expect to encounter under standard experimental conditions) will be discussed in this

444 445 446 section. Concentration-dependent effects encountered at higher concentrations will be explored in Sect. <u>4.2.</u>

447	Six counting efficiency curves were generated using sigmoidal distributions and the 50 % cut-off
448	diameters and maximum counting efficiencies listed in Table 3. Chosen values represent d_{50} values and
449	maximum counting efficiencies reported in the literature under relatively low concentrations of 1000-
450	4000 cm ⁻³ (Hermann et al., 2007). The resulting sigmoidal distributions (Fig. 7a) were used to
451	determine the counting efficiency of 25, 50, 100, and 200 nm particles.
452	
453	Next, κ_{app} was calculated from theoretical critical <u>percent</u> supersaturations for each chosen diameter. To
l 454	do so, four sigmoid curves representing sodium chloride CCN activation (hereon referred to as
455	"activation curves") for 25, 50, 100, and 200 nm were generated. The κ -Köhler-SS _{crit} of sodium
456	chloride was used as the midpoint of each activation curve, and one-tenth of this value was used as the
457	standard deviation (100 % CE, Fig. 7b-e). These values are consistent with the standard
458	deviation/midpoint ratio observed from our instrument's ammonium sulfate CCN calibration data.
459	
460	Activation curves were then generated for CPC Cases 1-6 by dividing the activated fraction for each dry
461	particle diameter by the counting efficiency for that diameter. SS_{crit} was determined for each CPC case
462	by finding the percent supersaturation at which activated fraction $= 0.50$. Results are summarized in
463	Fig. 7f. Next, critical supersaturation was converted to saturation, and $\kappa_{app,theory}$ was calculated for
464	each diameter in each CPC Case using Eq. (4) (see Fig. 7g). As above, κ_{app} artifacts were calculated by
465	finding the difference between these results and the literature value of κ_{app} for sodium chloride (see Fig.
466	7h).
467	

468	For the diameters studied	l, the effect of maxir	num counting efficiency	on CPC concentration (and
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- 469 activated fraction) is greater than the effect of 50 %-cutoff diameter. However, neither characteristic
- 470 resulted in large κ_{app} artifacts. The largest κ_{app} artifact observed at "low" concentrations was 0.035 for
- 471 CPC Case 4, 2.4 % of the literature value for the apparent hygroscopicity factor for sodium chloride.

4.2 CPC operation at high concentration

Λ	7	2
4	1	J

474	Operation at high concentrations introduces an additional source of undercounting through particle
475	coincidence at the CPC optical counter. For the TSI 3010 CPC, undercounting is observed is for particle
476	concentrations above $1 \times 10^4 \ cm^{-3}$. At $5 \times 10^4 \ cm^{-3}$, the detector saturates and cannot detect higher
477	concentrations. By comparison, the TSI 3025 is effective at counting higher particle concentrations, of
478	up to $2.5 \times 10^4 \ cm^{-3}$ (Hameri et al., 2002;Sem, 2002).
479	
480	To model undercounting due to particle coincidence, four CPC counting curves (Fig. 8a) were generated
481	using the equations in Table 4. Case 7 represents a CPC where counting efficiency decreases with
482	particle concentration, without reaching saturation. Cases 8-10 represent CPCs where saturation is
483	reached at 4 \times 10 ⁴ cm ⁻³ , 2 \times 10 ⁴ cm ⁻³ , and 1 \times 10 ⁴ cm ⁻³ , respectively. These saturation
484	concentrations are of similar magnitude to those observed from TSI 3010 concentration data. It should
485	be noted that the CPC concentration in Cases 7-10 levels off at the saturation concentration for each
486	case.
487	
488	In order to assess the importance of undercounting in CPC Cases 7-10, four theoretical aerosol
489	distributions with a peak concentration at 50 nm were employed (Table 5, Fig. 8b). CPC Distribution 1
490	represents a worst-case scenario of similar magnitude to the highest particle concentrations measured
491	during a coastal nucleation event (Hameri et al., 2002; Sem, 2002), while CPC Distributions 2, 3, and 4
492	are lower in concentration (due to the lack of undercounting in CPC Distributions 2, 3, and 4 as
493	demonstrated in Figure 6b, the remaining analysis for CPC operation at high concentration considers
494	only CPC Distribution 1.) CPC Cases 8-10 were applied to- CPC Distribution 1 in order to determine
I	

the concentration measured by the CPC for 25, 50, 100 and 200 nm aerosols. The counting efficiency
 was then calculated for each case and aerosol diameter in CPC Distribution 1.

- 498 Sigmoidal activated fraction curves were generated for 25, 50, 100 and 200 nm sodium chloride 499 aerosols. As in the low concentration cases, the midpoint of each 100 % CE curve was chosen to be 500 equal to the κ -Köhler-derived SS_{crit} of sodium chloride at each dry diameter, and the standard deviation 501 of each curve is equal to one-tenth of the SS_{crit} . These activated fraction curves were adjusted using the 502 counting efficiencies calculated in the previous step. In cases where the activated fraction has increased 503 due to undercounting by the CPC, the theoretical sigmoidal curve shifts to the left relative to the 100 % 504 CE case (Fig. 8c-f). Thus, undercounting by the CPC effectively increases the reported activated 505 fraction. As before, SS_{crit} was determined from each of these curves, and $\kappa_{app,theory}$ was subsequently 506 calculated using Eq. (4) (Fig. 8g-h).
- 507

 $\kappa_{app,theory}$ fell over a much wider range for 25, 50, and 100 nm particles (1.30-1.56, 1.32-1.70, and 1.30-1.55, respectively) than for 200 nm particles (1.28-1.29) due to the lower concentration of 200 nm particles in the chosen aerosol distribution, which resulted in a higher counting efficiency for these aerosols. In comparison, the largest range in $\kappa_{app,theory}$ was observed for 50 nm aerosols, the peak diameter in this aerosol distribution.

513

514 A wider range in $\kappa_{app,theory}$ was observed for the high-concentration CPC Cases (7-10) compared to the 515 low-concentration CPC Cases (1-6). The lowest counting efficiency observed across the low-516 concentration cases was 89.9 % for 25 nm aerosol in Case 4, while the lowest counting efficiency

517 observed in the high-concentration cases was 18.0 % for 50 nm aerosol in Case 10.

518

- 519 Artifacts in the apparent hygroscopicity parameter are shown in Fig. 8i. κ_{app} artifacts were the greatest
- 520 for a CPC that becomes saturated at 20,000 particles/cm³ (0.0131 $\leq \kappa_{app} \leq 0.4206$). The lower the
- 521 concentration at which a CPC becomes saturated, the more quickly its counting efficiency will drop as
- 522 concentration increases, resulting in increased activated fraction and increased apparent hygroscopicity.
- 523 The magnitude of artifacts due to CPC undercounting depends on the saturation concentration of the
- 524 CPC and the distribution of the aerosol population being studied.

525 5. Artifacts derived from cloud condensation nuclei instruments

526

Finally, the third instrument whose performance accuracy contributes to the overall certainty in CCN 527 528 assessment in the CCN instrument itself. Several instruments have been implemented for measuring CCN concentrations over the last few decades. Older models include the Continuous Flow Parallel Plate 529 530 Diffusion Chamber (Sinnarwalla, 1973) and the Hudson CCN spectrometer (Hudson, 1989) which both 531 employ an applied temperature gradient perpendicular to the aerosol flow. Newer models, such as the widely-used Droplet Measurement Technology Cloud Condensation Nuclei Counter (DMT CCN-100), 532 533 operate with a streamwise temperature gradient and continuous, laminar flow (Lance et al., 2006). The following analysis considers the DMT CCN-100. According to the CCN-100 manual, the counting 534 535 efficiency for this CCN instrument depends on concentration and supersaturation (Fig. 9a). The counting efficiency decreases rapidly with concentration at < 0.2 % SS due to rapid water vapor 536 537 depletion at these low supersaturations, and falls off more slowly for > 0.2 % SS (DMT CCN-100 538 manual).

539

The counting efficiency of the DMT CCN-100 was tested for four lognormal aerosol distributions with
peak concentrations at 50 nm and varying total concentrations (Table 5, Fig. 9b). Note that CCN Cases
<u>1-4 are identical to the aerosol distributions CPC Distributions 1-4 used for the high-concentration CPC</u>
<u>cases.</u>

544

The counting efficiencies for each case were applied to theoretical sodium chloride sigmoidal activated fraction curves to produce normalized activated fraction curves (Fig. 9c-f). As above, the midpoint is set to the SS_{crit} of sodium chloride at each dry diameter, and the standard deviation is assumed to be one-tenth of SS_{crit} . CCN undercounting effectively decreases activated fraction, therefore shifting the activated fraction curve downwards and towards higher percent supersaturations. The opposite effect is observed when CPC undercounting occurs. Critical supersaturation was determined for each CCN case, as above (Fig. 9g). Values of SS_{crit} were then converted to saturation, and $\kappa_{app,theory}$ was calculated using Eq. (4) (Fig. 9h).

553

Significant deviations from κ -Köhler theory were only observed in CCN Case 1, with total aerosol concentration 5 × 10⁶ particles/cm³ (Fig. 9g-i). The largest deviation for CCN Case 1 was observed in 100 nm particles ($\kappa_{app,artifact} = -0.57$), due to the higher concentration of 100 nm particles compared to 25 and 200 nm particles, and the lower <u>percent</u> supersaturation necessary for activation. The largest artifacts across CCN Cases 2 and 3 were also observed for 100 nm particles, though no artifacts were observed for any particle diameter in CCN Case 4 due to the much lower concentrations.

560

561 Sodium chloride is very hygroscopic. It should be noted that aerosols consisting of less hygroscopic 562 compounds will activate at higher percent supersaturations (> 0.2 % SS regime) which will lead to smaller κ_{app} artifacts when the same aerosol distribution and total aerosol concentration is considered. 563 564 If a mixture was considered (for example, sodium chloride with a non-hygroscopic species such as soot) 565 the results may also be different. The shape of the aerosol distribution must also be taken into account. 566 A distribution with a narrower peak than the one generated for this analysis would be at risk for larger κ_{app} artifacts for any total aerosol concentration, and these artifacts would be greater at the peak 567 diameter, while a broader distribution would result in less variation in κ_{app} artifacts for each particle 568 569 diameter.

- 570 6. <u>Counting statistics in CCN and CPC measurements</u>
- 571
- 572 Though it is beyond the scope of this analysis, it should be mentioned that sampling at very low particle
- 573 <u>concentrations (< 200 cm⁻³ total particle concentration) can introduce additional error into CCN and</u>
- 574 <u>CPC measurements. This error can be mitigated by increasing scan times (Moore et al., 2010). For</u>
- 575 example, Moore et al., 2010 averaged CCN and particle concentrations over 5-second intervals for
- 576 monodisperse particle concentrations $< 10 \text{ cm}^{-3}$, and increased averaging time to 20-second intervals
- 577 when the monodisperse particle concentration reached $< 6 \text{ cm}^{-3}$.
- 578

7. Discussion

581	A comparison of several instrument sources of error in CCN-derived κ_{app} is shown in Fig. 10. In					
582	addition, the best and worst case combination of errors, determined by additive error propagation, are					
583	also shown. DMA Case 4, CPC Case 4, CPC Case 10, and CCN Case 1 represent the operating					
584	conditions that resulted in the largest κ_{app} artifacts in this study. In DMA Case 4, the aerosol/sheath					
585	ratio of 0.30 resulted in a broadened aerosol distribution downstream of the DMA. Compared to DMA					
586	Case 1, where $Q_a/Q_{sh} = 0.10$, the downstream diameter range in DMA Case 4 was 300 % higher for 25					
587	nm particles, resulting in a spread of 20-36 nm. Similarly, the diameter ranges for 50, 100, and 200 nm					
588	diameter were 220 %, 230 %, and 220 % wider than in Case 1, respectively. Compared to the most ideal					
589	DMA case presented in this study (DMA Case 2), where $Q_a/Q_{sh} = 0.05$, the downstream diameter					
590	range in DMA Case 4 was 700 % higher for 25 nm particles; the diameter ranges for 50, 100, and 200					
591	nm diameter were 540 %, 560 %, and 520 % wider than in Case 2, respectively. <u>The results</u>					
592	demonstrate that limiting Q_a/Q_{sh} to ≤ 0.10 will result in a narrow particle size distribution downstream					
593	of the DMA. Other studies have recommended employing DMA sample/sheath ratios of 0.2 (Petters et					
594	al., 2007; Carrico et al., 2008; Moore et al., 2010) or 0.1 (Moore et al., 2010; Zhao-Ze and Liang, 2014)					
595	in order to minimize measurement aerosols due to transfer function broadening.					
596						
597	The effects of multiply-charged particles on κ_{app} calculations were also quantified, as shown in Fig. 10.					
598	<u>Small, positive κ_{app} artifacts (1 – 3 % of κ_{app}^{NaCl}) were observed when particles with +2 and +3 charges</u>					
599	were not accounted for. This analysis considered a theoretical aerosol distribution in which most of the					

500 particles measure less than 100 nm in diameter. Actual aerosol distributions vary temporally and

501 spatially, and often include accumulation and coarse modes that would result in larger κ_{app} artifacts.

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CPC Cases 8 and 10, respectively). It should be noted that undercounting was only observed for one of					
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<u>ultiple</u> 50,					
a s					

- artifact cases (DMA Case 2, CPC Case 3, and CCN Case 4) are < 0.001 except for 200 nm particles,
- 625 where $\kappa_{app,artifact} = 0.0013$.
- 626

- 627 We note that Fig. 4c-d demonstrated that κ_{app} error may result from instrument artifacts for ammonium
- 528 sulfate and sodium chloride, two standard compositions used in calibration of CCN instruments (Rose et
- 629 <u>al., 2008</u>). Therefore, the κ_{app} error encountered while calibrating the CCN instrument may compensate
- 630 for the CCN measurement bias of aerosol samples. However, as also demonstrated in Fig. 4c-d, the
- 631 <u>magnitude of this instrumentally-derived bias varies by compound.</u>

632 Conclusions

633

The sensitivity of weather and climate models to accuracy in CCN activation predictions has been
demonstrated in other works. Possible sources of apparent hygroscopicity artifacts calculated from CCN
measurements have been presented in this study. This analysis has focused on sodium chloride and
ammonium sulfate aerosols, but it can be extended to other aerosol populations, including mixtures and
field samples.

639

640	The greatest combine	ed artifacts (0.22	$< \kappa_{app.artifact}$	< 0.34, NaCl)	occurred as a result of the
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641 combined issues of the highest DMA aerosol/sheath ratio, uncorrected multiple particle charging, and

undercounting by both CPC and CCN instrument. The lowest combined artifacts (0.00018 <

643 $\kappa_{app,artifact} < 0.0013$, NaCl) occurred as a result of ideal operating conditions :lowest DMA/sheath

ratio, corrected multiple particle charging, and little to no undercounting.

645

The largest single-instrument artifacts ($-0.57 < \kappa_{app,artifact} < 0.42$ for sodium chloride) in this study arise from undercounting by either the CPC or CCN counter at high concentration. This problem arises during attempts to measure aerosol concentrations of $\sim 10^4$ cm⁻³ which is much higher than the recommended concentration ranges for either instrument, (CPC Cases 7-10 and CCN Case 4). Corrective action should be taken to dilute aerosol samples in order to avoid undercounting. It should be noted that these artifacts are for individual instruments and do not take combined operation of the CPC and CCN into account; when both instruments undercount, artifacts in $\kappa_{app,artifact}$ are reduced.

654 Smaller single-instrument artifacts ($\kappa_{app,artifact} < 0.04$) were observed for the CPC cases where 50 %-655 cut-off diameter and maximum counting efficiency were varied. Given the chosen particle diameters 656 (25, 50, 100, 200 nm), κ_{app} artifacts due to d_{50} were minimal. The largest κ_{app} artifacts for a CPC 657 counting at low concentration (0.031-0.035) were observed where the maximum counting efficiency was 658 equal 0.90. This may represent a compositional mismatch between n-butanol as the working fluid and 659 sodium chloride as the aerosol, due to the poor solubility of the latter in the former. Individual n-butanol 660 CPCs may exhibit higher maximum counting efficiencies for sodium chloride.

661

Uncertainty arising from the DMA depended greatly on the chosen aerosol and sheath settings. One set of DMA cases (Cases 2-4) examined the effect of aerosol/sheath ratio. By decreasing this ratio, a narrower near-monodisperse flow can be produced, which increases the accuracy of calculated κ_{app} . The κ_{app} artifacts for an aerosol/sheath ratio of 0.10 were 1.1 % of $\kappa_{literature}$ for 25 nm sodium chloride aerosols, 0.93 % for 50 nm, 0.81 % for 100 nm, and 0.66 % for 200 nm. Decreasing the aerosol/sheath ratio to 0.05 resulted in $\kappa_{app} \approx 0.01$ % of $\kappa_{literature}$ for NaCl. When a ratio of 0.30 was employed, the resulting artifacts that were ~10 % of $\kappa_{literature}$ for sodium chloride.

The second set of DMA cases (5-7) were designed to evaluate the effects of holding the sheath flow constant while varying the excess air flow by -2 %, +2 %, and +5 %. These resulted in shifts of \leq 2 nm for 25 nm and 50 nm particles, \leq 4 nm for 100 nm particles, and \leq 7 nm for 200 nm particles. The downstream aerosol distribution was shifted towards larger particle diameters when sheath flow exceeded excess flow, and towards smaller particle diameter when sheath flow was less than excess flow. When taking field measurements, the composition of the sample may vary with particle diameter, thereby introducing another source of error from a broader DMA distribution.

678 Under optimal operating conditions, where the DMA sample/sheath ratio is 0.10 and excess/sheath ratio 679 is 1.0, and in the absence of undercounting by the CPC or CCN, uncertainties in κ_{app} are within ±1.1 % for 25 to 200 nm particles. When the DMA sample/sheath ratio drops to 0.05, κ_{app} uncertainties decrease to ±0.01 %. Additionally, errors in activated fraction (and therefore κ_{app}) resulting from the bipolar charge distribution can be corrected by determining the fraction of particles with multiple charges.

684

685 Apparent hygroscopicity parameter artifacts were calculated for two pure, inorganic species in this study. This analysis could be used to estimate κ_{app} artifacts for ambient aerosol populations, which 686 687 may result in a better understanding of the "real' differences between these populations. As discussed in the introduction, Collins et al. 2016 aggregated κ_{app} from several mesocosm and field studies for 30-688 100 nm sea spray aerosol ($0.4 < \kappa_{app}^{SSA} < 1.3$). The wide range of κ_{app} in these studies may be 689 690 attributed to differences in composition, experimental artifacts, or a combination of the two. Quantification of experimental artifacts would facilitate interpretation of κ_{app} in aerosol populations and 691 692 constrain the importance of composition in CCN activation. There has been a recent proliferation of 693 CCN data availability from multiple researchers and multiple experimental setups. To maximize the 694 utility of these studies and to compare cloud-activating properties of various ambient aerosol masses, it 695 is essential that artifacts are considered in both CCN data collection and in reporting of the data.

696	Supplement Link
697	Will be included by Copernicus
698	
699	Author Contribution
700	Sarah D. Brooks provided the conceptual framework and contributed to the writing of the manuscript.
701	Jessica A. Mirrielees performed the analysis and lead the writing of the manuscript.
702	
703	Competing Interests
704	The authors declare that they have no conflict of interest.
705	
706	Disclaimer
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708	
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Notation			
$\alpha_{cc}, \beta_{cc}, \gamma_{cc}$	Empirically-determined constants used to		
	calculate Cunningham slip correction factor		
Z_p	Aerosol particle electrical mobility		
Cc	Cunningham slip correction factor		
d_m	Electrical mobility diameter		
n	Number of charges on particle		
е	Elementary unit of charge		
η	η Gas dynamic viscosity		
λ	Mean free path		
Q_{sh}	Sheath flow		
${oldsymbol Q}_e$	Excess air flow		
\boldsymbol{Q}_{a}	Aerosol flow		
Q_s	Sample flow		
κ_{app}	Apparent hygroscopicity parameter		
κ _{app,artifact}	Apparent hygroscopicity parameter artifact		
S	Equilibrium water vapor saturation		
S _{crit}	Critical saturation (50 % of aerosols active as		
	cloud condensation nuclei)		
A	Constant used in calculating κ_{app}		
σ_{lv}	Surface tension of water		
Т	Temperature		
D _{act}	Activation diameter		
SS _{crit}	Critical percent supersaturation		
α_{TF}	Height of DMA transfer function		
β_{TF}	Half-width of DMA transfer function		
Z'_p	Mobility of particle at DMA inlet		
$Z_{p,mid}$	Midpoint of transfer function		
ΔZ_p	Half-width of transfer function		
V ₀	Voltage selected at DMA		
r_1			
r_2	DMA outer radius		
Ĺ	DMA length		
<i>d</i> ₅₀	50 %-cut-off diameter		

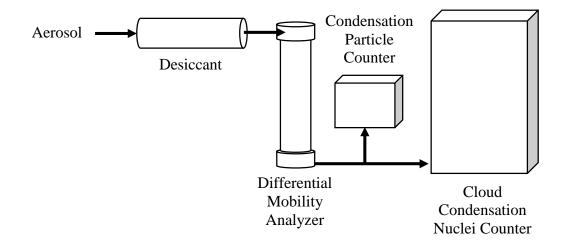
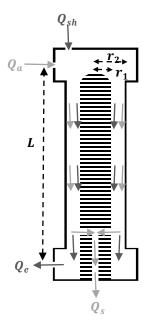


Figure 1 Experimental setup <u>used for obtaining sized</u> CCN and particle concentration measurements

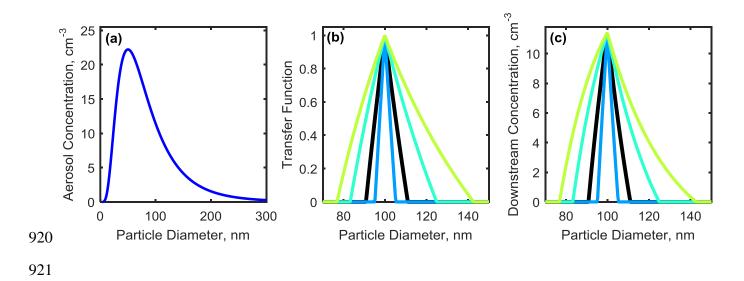
915 <u>from an aerosol sample</u>.



917 **Figure 2** Simplified flow diagram of a DMA with an inner electrode radius r_1 , outer electrode radius r_2 ,

918 distance between aerosol inlet and sample outlet L, clean sheath air flow Q_{sh} , aerosol flow Q_a , excess air

919 <u>flow Q_e , and sample air flow Q_s .</u>



922 Figure 3 (a) A theoretical aerosol distribution generated using a lognormal function centered at 50 nm.
923 (b) The transfer function calculated using Eq. (7). (c) <u>Multiplying the distribution by the transfer function</u>
924 gives the downstream aerosol concentration (cm⁻³).

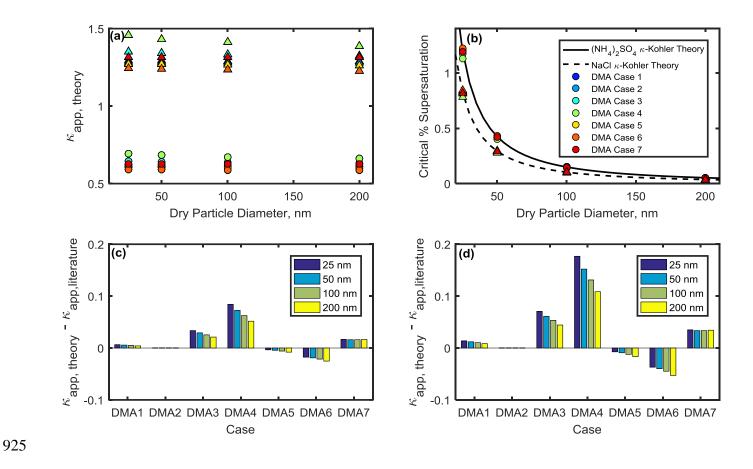


Figure 4 (a) Apparent hygroscopicity κ_{app} for <u>DMA cases 1-7 for sodium chloride (triangles) and</u> ammonium sulfate (circles) [see legend in (b)]. (b) Critical supersaturation of ammonium sulfate and sodium chloride particles calculated using κ_{app} values derived in (a). Ammonium sulfate and sodium chloride curves from κ -Köhler theory are shown for comparison. Legend colors apply to both salts. (c,d) DMA-flow-derived artifacts in κ_{app} are shown for each DMA case and both salts.

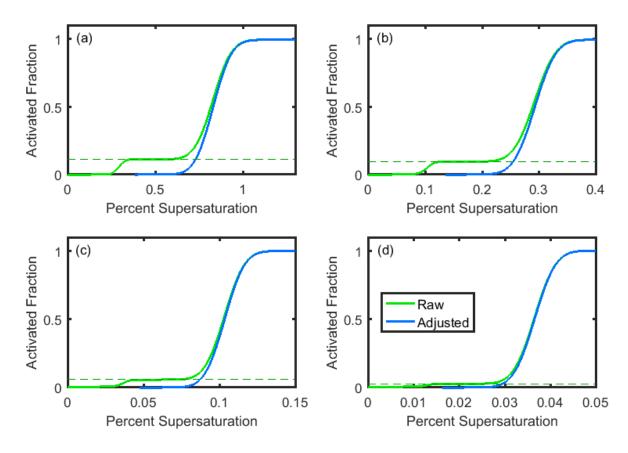
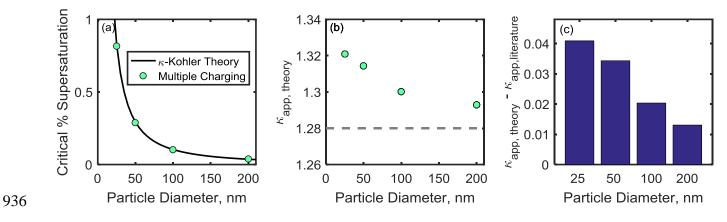
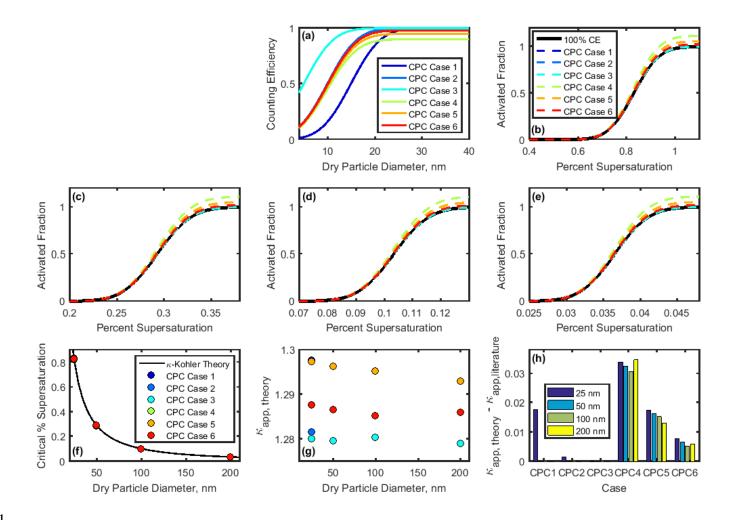
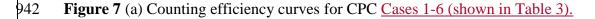


Figure 5 Theoretical raw (green) and adjusted (blue) activated fraction curves for (a) 25 nm (+1), 50 nm
(+2), and 75 nm (+3) particles; (b) 50 nm (+1), 100 nm (+2), and 150 nm (+3) particles; (c) 100 nm
(+1), 200 nm (+2), and 300 nm (+3) particles; (d) 200 nm (+1), 400 nm (+2), and 600 nm (+3) particles.
All particles are pure sodium chloride.

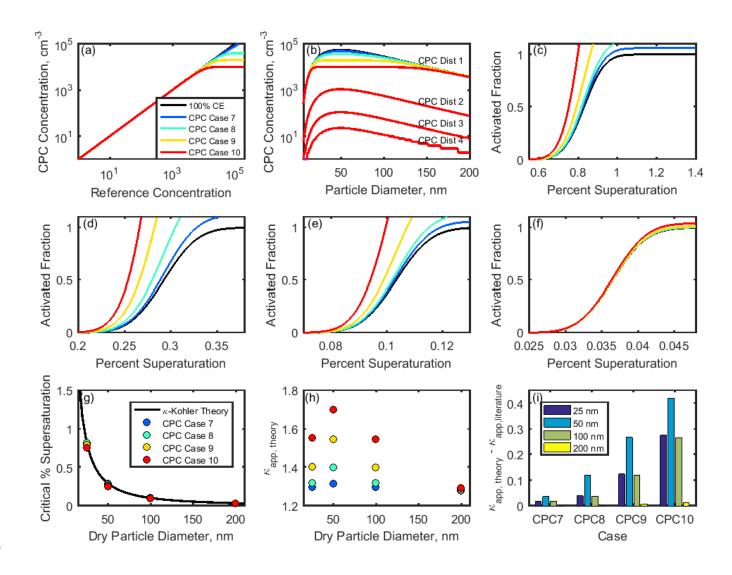


937 **Figure 6** (a) Critical percent supersaturation of sodium chloride particles determined from activated 938 fraction curves shown in Fig. 5. A κ -Köhler curve for sodium chloride is shown for comparison. (b) 939 Theoretical κ_{app} for each particle diameter (gray dashed line indicates literature value for κ_{app}^{NaCl}). (c) 940 Artifacts in κ_{app} resulting from multiple particle charges.



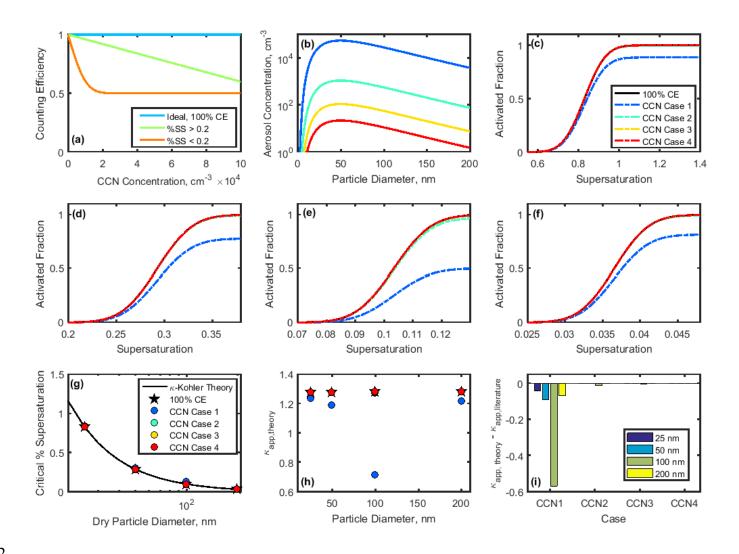


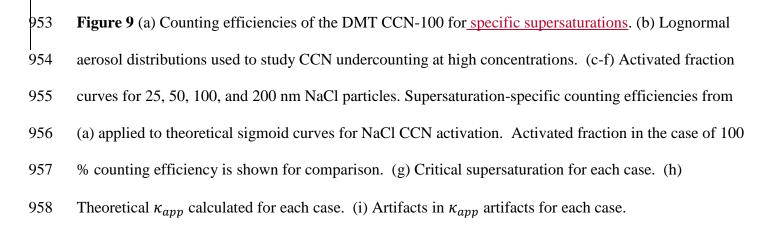
943 (b-e) CCN activated fraction curves for 25, 50, 100, and 200 nm NaCl, respectively. (f) Critical 944 supersaturation calculated for each <u>particle diameter</u>. (g) Theoretical κ_{app} for each CPC case <u>and</u> 945 <u>particle diameter</u>. (h) Artifacts in κ_{app} for each CPC case <u>and particle diameter</u>.



946

Figure 8 (a) Theoretical relationships between the reference aerosol concentration and CPC concentration. (b) Concentration-dependent counting efficiencies from (a) were applied to four theoretical aerosol distributions. (c-f) Activated fraction curves for <u>CPC Distribution 1 and particle</u> diameters 25, 50, 100, and 200 nm NaCl aerosol, respectively. (g,h) Critical supersaturation and κ_{app} for each case. (i) Artifacts in κ_{app} for each case.





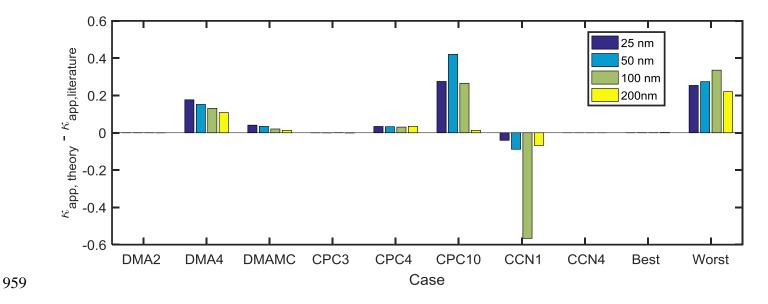


Figure 10 Comparison of κ_{app} artifacts derived from best and worst case scenarios for instrumental measurements for sodium chloride. Combined artifacts for the lowest-artifact cases (Best: DMA Case 2, CPC Case 3, and CCN Case 4) and the highest-artifact cases (Worst: DMA Case 4, <u>multiple charging</u>, CPC Case 4, CPC Case 8, and CCN Case 1).

Table 1 Theoretical DMA Flow Test Cases						
Case	Q_{sh} (L min ⁻¹)	Q_e (L min ⁻¹)	Q_a (L min ⁻¹)	Q_s (L min ⁻¹)	Q_a/Q_{sh}	Q_e/Q_{sh}
DMA 1	3.00	3.00	0.30	0.30	0.10	1.00
DMA 2	3.00	3.00	0.15	0.15	0.05	1.00
DMA 3	3.00	3.00	0.60	0.60	0.20	1.00
DMA 4	3.00	3.00	0.90	0.90	0.30	1.00
DMA 5	3.00	3.06	0.36	0.30	0.12	1.02
DMA 6	3.00	3.15	0.45	0.30	0.15	1.05
DMA 7	3.00	2.94	0.24	0.30	0.08	0.98

Table 2	Table 2 Predicted downstream particle diameter range for each DMA case.				
Case	25 nm	50 nm	100 nm	200 nm	
DMA 1	23-27	46-56	91-111	181-222	
DMA 2	24-26	48-53	95-105	190-211	
DMA 3	21-31	42-62	83-125	167-250	
DMA 4	20-36	39-71	77-143	154-285	
DMA 5	23-27	45-55	90-110	181-220	
DMA 6	22-27	45-54	89-107	178-215	
DMA 7	23-28	46-56	92-112	183-225	

Table 3 Values of 50%-cutoff diameter and maximum counting efficiency used in investigating κ_{app} artifacts for low particle concentrations measured by a CPC.

Case	<i>d</i> ₅₀ , nm	Maximum Counting Efficiency
CPC 1	15	100 %
CPC 2	10	100 %
CPC 3	5	100 %
CPC 4	10	90 %
CPC 5	10	95 %
CPC 6	10	98 %

Table 4 Equations used to model the relationship between a reference or "true" aerosol concentration x (particles cm⁻³), and the concentration measured by a condensation particle counter y (particles cm⁻³).

Case	Equation
CPC 7	$y = x - 2 \times 10^{-6} x^2$
CPC 8	$y = 40000 \operatorname{erf}\left(\frac{x}{32000\sqrt{2}}\right)$
CPC 9	$y = 20000 \operatorname{erf}\left(\frac{x}{16000\sqrt{2}}\right)$
CPC 10	$y = 10000 \operatorname{erf}\left(\frac{x}{8000\sqrt{2}}\right)$

Table 5 Total concentrations used in theoretical aerosol distribution for CPC operation at hig	;h
concentration and CCN-derived κ_{app} artifacts.	

CPC Distribution	CCN Case	Total Concentration (particles cm ⁻³)
CPC Distribution 1	CCN 1	5×10^{6}
CPC Distribution 2	CCN 2	1×10^{5}
CPC Distribution 3	CCN 3	1×10^{4}
CPC Distribution 4	CCN 4	2×10^{3}