



1 **Instrument Artifacts Lead to Uncertainties in Parameterizations of Cloud**

2 **Condensation Nucleation**

3 **Jessica A. Mirrielees and Sarah D. Brooks**

4 Texas A&M University, College Station, TX 77843

5 Correspondence E-mail: [jmirrieles@tamu.edu](mailto:jmirrieles@tamu.edu)



6 **Abstract**

7

8 The concentrations of cloud condensation nuclei (CCN) modulate cloud properties, rainfall  
9 location and intensity, and climate forcings. This work assesses uncertainties in CCN  
10 measurements and the apparent hygroscopicity parameter ( $\kappa_{app}$ ) which is widely used to represent  
11 CCN populations in climate models. CCN measurements require accurate operation of three  
12 instruments: the CCN instrument, the differential mobility analyzer (DMA), and the condensation  
13 particle counter (CPC). Assessment of DMA operation showed that varying the ratio of aerosol to  
14 sheath flow from 0.05 to 0.30 resulted in discrepancies between the  $\kappa_{app}$  values calculated from  
15 CCN measurements and the literature value. Discrepancies were found to increase from effectively  
16 zero to 0.18 for sodium chloride, and from effectively zero to 0.08 for ammonium sulfate. The  
17 ratio of excess to sheath flow was also varied, which shifted the downstream aerosol distribution  
18 towards smaller particle diameters (for excess flow < sheath flow) or larger particle diameters (for  
19 excess flow > sheath flow) than predicted. For the CPC instrument, undercounting occurred at  
20 high concentrations, resulting in calculated  $\kappa_{app}$  lower than the literature values. Lastly,  
21 undercounting by CCN instruments at high concentration was also assessed, taking the effect of  
22 supersaturation on counting efficiency into account. Under recommended operating conditions,  
23 the combined DMA, CPC, and CCN uncertainties in  $\kappa_{app}$  are 1.1 % or less for 25 to 200 nm  
24 diameter aerosols.



25 **Copyright Statement**

26 Will be provided by Copernicus



## 27 **Introduction**

28

29 Aerosol-cloud interactions represent a major uncertainty in current predictions of the Earth's climate  
30 (IPCC, 2013). According to well-known Köhler theory, an aerosol's potential to catalyze cloud droplet  
31 formation by activating as a cloud condensation nucleus (CCN) depends on its physical and chemical  
32 properties. For any given composition, the CCN activation potential of an aerosol increases as its  
33 diameter decreases. While the relationship between aerosol diameter and CCN activation is  
34 straightforward, the effect of composition on an aerosol's ability to participate in cloud formation is  
35 more complex (Petters and Kreidenweis, 2013; Ovadnevaite et al., 2011). Predicting the cloud forming  
36 capacity of various air masses based on the properties of the aerosol they contain is essential for  
37 evaluating relative contributions from pollution, continental background and marine aerosol sources  
38 (Brooks and Thornton, 2018; Carslaw et al., 2013). Long-term CCN measurements are available from  
39 numerous locations globally (Schmale et al., 2018). However, understanding regional and temporal  
40 variability in CCN populations requires the ability to assess whether observed differences reflect true  
41 physical differences or simply variations in CCN sampling strategies.

42

43 Parameterizations of CCN activity which accurately prescribe CCN measurements are needed for  
44 climate models, cloud resolving models, and air quality predictions (Betancourt and Nenes, 2014;  
45 Betancourt et al., 2013; Chang et al., 2017; Crosbie et al., 2015; Karydis et al., 2012; Kawecki and  
46 Steiner, 2018). One parameterization was designed to represent the cloud droplet activation potential  
47 ambient aerosol masses of unknown composition with a single variable, kappa ( $\kappa$ ) based on the dry  
48 aerosol's hygroscopicity, or ability to uptake water and form a solution droplet (Petters and  
49 Kreidenweis, 2007). Various names and abbreviations have been given to  $\kappa$  throughout the literature:  
50 "hygroscopicity parameter", "single hygroscopicity parameter",  $\kappa$  (Petters and Kreidenweis, 2007;



51 Carrico et al., 2008;Asa-Awuku et al., 2010;Moore et al., 2012); “CCN-derived  $\kappa$ ”,  $\kappa_{CCN}$  (Carrico et al.,  
52 2008; Petters and Kreidenweis, 2007); and the “apparent hygroscopicity parameter”  $\kappa_{app}$  (Sullivan et  
53 al., 2009; Collins et al., 2016; Petters and Kreidenweis, 2013). The term *apparent* hygroscopicity is  
54 favored by many because it emphasizes that fact that while CCN activation can often be predicted  
55 accurately by hygroscopic water uptake, they are different physical processes. It is possible for a  
56 compound to have high intrinsic hygroscopicity and low apparent hygroscopicity if it is poorly soluble  
57 in water (Sullivan et al., 2009).

58

59 Once calculated, hygroscopicity parameters are useful tools for comparing CCN field measurements  
60 conducted in various regions and seasons and for making predictions about cloud formation, aerosol-  
61 cloud interactions in weather, and climate models. Values of  $\kappa_{app}$  can be used to compare the CCN  
62 results in field and laboratory studies, including sea spray aerosol. For example, aggregation of results  
63 from several mesocosm experiments and marine field studies found submicron (30-100 nm)  $\kappa_{app}$  for sea  
64 spray aerosol as low as 0.4 and as high 1.3 (Collins et al., 2016).

65

66 Several studies have examined the sensitivity of models to  $\kappa$  values derived from HTDMA  
67 measurements. An analysis of the NASA Global Modeling Initiative Chemical Transport Model and the  
68 GEOS-Chem CTM (Karydis et al., 2012) found that cloud droplet number concentration is sensitive to  $\kappa$   
69 in Arctic and remote regions, where background aerosol loadings are low. Another study (Betancourt  
70 and Nenes, 2014) found that a  $\pm 50$  % uncertainty range in the  $\kappa$  of secondary organic aerosols and  
71 particulate organic matter resulted in a cloud droplet number concentration uncertainty of up to 15 %  
72 and 16 %, respectively. Updating precipitation models with lab-derived  $\kappa$  values for specific inorganic  
73 and organic species may increase the accuracy of storm forecasts by providing better predictions of  
74 intense precipitation (Kawecki and Steiner, 2018). In terms of climate, (Liu and Wang, 2010) found that



75 increasing the  $\kappa$  of primary organic aerosols from 0 to 0.1, and decreasing the  $\kappa$  of secondary organics  
76 aerosols from 0.14 to 0.07, resulted in an uncertainty in global secondary aerosol indirect forcing of 0.4  
77  $\text{Wm}^{-2}$  from pre-industrial times to present day.

78

79 The sensitivity of weather and climate models to hygroscopicity parameters demonstrates the need for  
80 accurate measurements. In this study, we examine experimental uncertainties in CCN measurements and  
81 the resulting uncertainties in determination of  $\kappa_{app}$ . Differences in report  $\kappa_{app}$  values may result from  
82 experimental artifacts rather than actual differences in aerosol's ability to facilitate cloud formation. By  
83 systematically quantifying sources of experimental error, this study provides a framework for  
84 determining the significance of variations in CCN properties reported in multiple studies and defining  
85 the operating conditions which minimize instrumental artifacts.

86



87 **1. Background**

88

89 The Köhler equation relates water vapor saturation at the surface of a wet droplet,  $s$ , to its radius at  
90 equilibrium (Rogers and Yau, 1989):

91

92 
$$s = \left(1 - \frac{b}{r^3}\right) \exp\left(\frac{a}{r}\right) \quad (1a)$$

93

94 
$$a = \frac{2\sigma_w M_w}{\rho_w R T} \quad (1b)$$

95

96 
$$b = \frac{3im_s M_w}{4\pi\rho_w M_s} \quad (1c)$$

97

98 where  $s$  is the equilibrium saturation ratio of a solution droplet with radius  $r$ ,  $\sigma_w$  is the surface tension of  
99 water,  $M_w$  is the molecular weight of water,  $R$  is the ideal gas constant,  $T$  is temperature in Kelvin,  $\rho_w$  is  
100 the density of water, and  $M_s$  is the molecular weight of the solute. The minimum saturation ratio that is  
101 required for spontaneous droplet growth,  $s_{act}$ , is therefore:

102

103 
$$s_{crit} = 1 + \sqrt{\frac{4a^3}{27b}} \quad (2)$$

104

105 Petters and Kreidenweis [2007] reformulated the Köhler equation as  $\kappa$ -Köhler theory:

106

107 
$$s_{crit} = \exp\left(\sqrt{\frac{4A^3}{27D_{act}^3 \kappa_{app}}}\right) \quad (3a)$$

108 and



109

110 
$$A = \frac{4\sigma_{lv}M_w}{RT\rho_w} \quad (3b)$$

111

112 Where  $s_{crit}$  is the critical water vapor saturation,  $D_{act}$  is the dry particle activation diameter and  $\kappa_{app}$  is  
113 the apparent hygroscopicity parameter. Solving for  $\kappa_{app}$  yields:

114

115 
$$\kappa_{app} = \frac{4A^3\sigma_{lv}^3}{27T^3D_{act}^3\ln^2(s_{crit})} \quad (4)$$

116

117 The apparent hygroscopicity parameter can be calculated from experimental CCN results, where the dry  
118 diameter and water vapor saturation are known. For a chosen aerosol diameter, the activated fraction is  
119 the ratio of the concentration aerosols that activate as CCN to the total aerosol concentration:

120

121 
$$\text{Activated fraction} = \frac{\text{CCN Concentration}}{\text{Aerosol Concentration}} \quad (5)$$

122

123 Activated fraction data is fit with a sigmoid error function to determine the supersaturation at which 50  
124 % of the particles have activated as CCN, which is considered the operationally defined critical  
125 supersaturation  $SS_{crit}$  (Rose et al., 2008). The critical saturation  $s_{crit}$  can then be determined and  
126 entered into Eq. (4) in order to calculate  $\kappa_{app}$  for the near-monodisperse aerosol:

127

128 
$$s_{crit} = 1 + \frac{SS_{crit}}{100} \quad (6)$$

129



130 Reporting  $\kappa_{app}$  as a function of diameter allows for the comparison of the cloud condensation nucleation  
131 abilities of multimodal aerosol populations, without overlooking differences which arise due to aerosol  
132 composition.

133

134 The apparent hygroscopicity parameter is related to chemical composition; therefore, the calculated  $\kappa_{app}$   
135 of a pure substance should be constant across CCN experiments. However, discrepancies between  $\kappa_{app}$   
136 for a single chemical species have been observed. Experimental results for ammonium nitrate are  
137 inconsistent  $0.577 \leq \kappa_{app} \leq 0.753$ , and large ranges are often observed for organic compounds, such  
138 as glutaric acid ( $0.054 \leq \kappa_{app} \leq 0.16$ ) and malonic acid ( $0.199 \leq \kappa_{app} \leq 0.255$ ) (Koehler et al.,  
139 2006; Kumar et al., 2003; Hartz et al., 2006; Svenningsson et al., 2006). Below we evaluate potential  
140 sources of uncertainties in CCN measurements and the resulting uncertainties in  $\kappa_{app}$ .

141



## 142 2. Artifacts derived from CCN measurements

143

144 CCN measurements require accurate operation of three instruments: the CCN, the differential mobility  
145 analyzer (DMA), and the condensational particle counter (CPC). The setup for laboratory CCN  
146 experiments is shown in Fig. 1. First, a polydisperse population of aerosols is generated by an atomizer  
147 and dried using a desiccant tube packed with silica gel. A near-monodisperse flow is obtained through  
148 size-selection in the DMA. The flow is then split between a CPC (which measures aerosol  
149 concentration) and a CCN counter (which measures the concentration of particles that activate as cloud  
150 condensation nuclei at a given percent water vapor supersaturation).

151

152

153

### 154 2.1 Artifacts derived from differential mobility analyzers

155

#### 156 2.1.1 DMA operation and electrical mobility

157

158 Differential mobility analyzers used in atmospheric science include commercially available instruments  
159 from Grimm Aerosol Technik, TSI Incorporated, and MSP Corporation. They have also been custom  
160 built by a number of research groups (Mei et al., 2011; Barmounis et al., 2016; Jokinen and Makela,  
161 1997; Seol et al., 2000). All models allow for the selection of particles through electrical mobility, the  
162 ability of a particle to move through a medium (such as air) while acted upon by an electrical field. The  
163 DMA size-selects near-monodisperse aerosol from a polydisperse aerosol source, as shown in Fig. 2  
164 (modeled after the Vienna-type long Differential Mobility Analyzer from Grimm Technologies). The  
165 electrical mobility  $Z_p$  of a particle with mobility diameter  $d_m$  can be calculated according to:



166

167 
$$Z_p = \frac{neC_C(d_m)}{3\pi\eta d_m} \quad (7)$$

168

169 where  $n$  is the number of charges on the particle (assumed to be one in this study),  $e$  is the elementary  
170 unit of charge,  $\eta$  is the gas dynamic viscosity, and  $C_C(d_m)$  is the Cunningham slip correction factor:

171

172 
$$C_C(d_m) = 1 + \frac{2\lambda}{d_m} \left( \alpha_{CC} + \beta_{CC} \exp \left[ -\frac{\gamma_{CC}}{2\lambda/d_m} \right] \right) \quad (8)$$

173

174 where  $\lambda$  is the mean free path (DeCarlo et al., 2004). For the Vienna-type long Differential Mobility  
175 Analyzer from Grimm Technologies, Inc. considered here,  $\alpha_{CC} = 1.246$ ,  $\beta_{CC} = 0.42$ , and  $\gamma_{CC} = 0.86$   
176 (Grimm Aerosol Technik, 2009).

177

178 Particle-laden flow enters the differential mobility analyzer through the aerosol inlet (flow  $Q_a$ ), and travels  
179 down the DMA column (inner radius  $r_1$ , outer radius  $r_2$ ) with the clean air sheath flow  $Q_{sh}$ . Positively-  
180 charged particles are attracted by the negatively-charged inner electrode, to which voltage  $V_0$  has been  
181 applied. Ideally, selection of a voltage allows only particles of a specific mobility diameter to exit the  
182 DMA through the sample flow  $Q_s$ . All particles with larger diameter (lower  $Z_p$ ) or smaller diameter  
183 (higher  $Z_p$ ) will exit the DMA through the excess flow  $Q_e$ . In other words,  $Q_s$  would ideally be a truly  
184 monodisperse flow.

185

186 In reality, the aerosol flow that leaves the DMA through  $Q_s$  is polydisperse with a mobility distribution  
187 determined by instrumental parameters. A triangular approximation has been chosen as a model for this  
188 distribution, as particle inertia is negligible for the diameters considered in this study (Stratmann et al.,



189 1997;Mamakos et al., 2007). The probability that a particle at the aerosol inlet will exit with the sampling  
 190 flow is defined by transfer function  $f(Z_p, Z_{p,mid})$ :

191

$$192 \quad f(Z_p, Z_{p,mid}) = \frac{\alpha_{TF}}{2\beta_{TF}} \left( \left| \frac{Z_p}{Z_{p,mid}} - (1 + \beta_{TF}) \right| + \left| \frac{Z_p}{Z_{p,mid}} - (1 - \beta_{TF}) \right| - 2 \left| \frac{Z_p}{Z_{p,mid}} - 1 \right| \right) \quad (9)$$

193 where  $Z_{p,mid}$  is the midpoint mobility of the transfer function, and  $\alpha_{TF}$  and  $\beta_{TF}$  are flow-derived  
 194 constants, defined as:

195

$$196 \quad \alpha_{TF} = \frac{Q_s + Q_a}{2Q_a} \quad (10a)$$

197 and

$$198 \quad \beta_{TF} = \frac{Q_s}{Q_{sh}} \quad (10b)$$

199

200 The midpoint and half-width of the transfer function are respectively calculated according to: (Knutson  
 201 and Whitby, 1975)

$$202 \quad Z_{p,mid} = \frac{Q_e + Q_{sh}}{4\pi LV_0} \ln \left( \frac{r_2}{r_1} \right) \quad (11a)$$

203 and

204

$$205 \quad \Delta Z_p = \frac{Q_a}{2\pi LV_0} \ln \left( \frac{r_2}{r_1} \right) \quad (11b)$$

206

207

208



## 209 2.1.2 $\kappa_{app}$ artifact analysis and results

210

211 Next we assess the ramifications of the DMA transfer function for the derived  $\kappa_{app}$ . A lognormal  
212 theoretical aerosol number distribution was used to represent a polydisperse ambient aerosol population  
213 (Fig. 3a). This distribution was converted to an electrical mobility distribution using Eq. (7) and Eq. (8),  
214 assuming that the aerosols in the distribution were spherical and singly charged. From the distribution, a  
215 series of single aerosol sizes were selected (25, 50, 100, and 200 nm diameter). For each aerosol size,  
216 the resulting DMA transfer functions were calculated for 7 cases using Eq. (9) and the various  
217 parameters for DMA sheath, excess, aerosol, and sample flow listed in Table 1. For example, the  
218 resulting DMA transfer functions for a 100 nm aerosol conditions constrained by Cases 1-4 are shown in  
219 Fig. 3b, where an increase in  $Q_a/Q_{sh}$  from 0.1 (black line) to 0.3 (green line) tripled the width of the  
220 number distribution, and decreasing  $Q_a/Q_{sh}$  to 0.05 (blue line) from 0.10 halved the width of the  
221 number distribution. The result of applying the transfer functions shown in Fig. 3b to the distribution in  
222 Fig. 3a is shown in Fig. 3c.

223

224 All downstream distributions for all seven DMA cases and all aerosol sizes are shown in Fig. S1 in the  
225 Supplement. DMA Cases 1-4 represent experimental conditions in which the sheath and excess air  
226 flows are equal and the aerosol/sheath flow ratio is varied. As  $Q_a/Q_{sh}$  increases, the width of the  
227 number distribution measured downstream of the DMA increases, while the midpoint diameter remains  
228 constant. It was found that doubling the aerosol to sheath ratio doubled the width of the downstream  
229 number distribution for 25, 50, 100, and 200 nm particles. For example, when selecting 200 nm  
230 particles, increasing  $Q_a/Q_{sh}$  from 0.10 to 0.20 increased the downstream diameter range from 181-222  
231 nm (a spread of 41 nm) to 167-250 nm (a spread of 87 nm). The particle diameter ranges that would be  
232 observed downstream of the DMA are summarized in Table 2.



233

234 To assess the variations in CCN properties resulting from DMA uncertainties the critical supersaturation  
235 were calculated for representative atmospheric aerosols. The value of  $SS_{crit}$  was calculated for each  
236 particle diameter using Eq. (3a), using literature values for apparent hygroscopicity of 0.61 for  
237 ammonium sulfate and 1.28 for sodium chloride (Clegg et al., 1998). It should be noted that this  
238 analysis considers two homogeneous aerosol distributions of hygroscopic salts. Real aerosol  
239 distributions tend to be mixtures of many species, and the shape of the number distribution can vary  
240 between species.

241

242 Note that in the absence of DMA diameter uncertainty, this single component aerosol population should  
243 be characterized by a single  $\kappa_{app}$  regardless of diameter. To test how uncertainties in DMA diameter  
244 translate to uncertainties in  $\kappa_{app}$ , the true critical saturation ratio  $s_{crit}$  was then put into Eq. (4) in order  
245 to calculate the “perceived”  $\kappa_{app}$  for each diameter given the chosen transfer function from Cases 1-7.  
246 For example, if 100 nm particles were selected from the DMA by the user, the transfer functions would  
247 allow larger and smaller particles to pass into the sample flow, as shown in Table 2. Particles with  
248 diameter  $> 100$  nm would be “perceived” to have higher apparent hygroscopicity than particles with  
249 diameter  $< 100$  nm, since the equilibrium vapor pressure over the surface of a particle decreases as its  
250 diameter increases (and as curvature decreases). Using Eq. (12), these diameter-specific “perceived”  
251  $\kappa_{app}$  values were volume-weighted, resulting in  $\kappa_{app,theory}$ :

252

$$253 \quad \kappa_{app,theory} = \sum_i \epsilon_i \kappa_i \quad (12)$$

254

255 Results for ammonium sulfate and sodium chloride are shown in Fig. 4a. The critical saturation ratio  
256 was calculated from  $\kappa_{app,theory}$  using Eq. (3a) for each case and converted to critical supersaturation.



257 The results are compared to theoretical  $\kappa$ -Köhler theory curves for ammonium sulfate and sodium  
258 chloride generated using the literature  $\kappa_{app}$  for each compound (Fig. 4b).  
259  
260 Discrepancies between  $\kappa_{app,theory}$  calculated in this study and literature values (hereon referred to as  
261 “ $\kappa_{app}$  artifacts”) are shown for both compounds in Fig. 4c-d.  
262  
263 The greatest  $\kappa_{app}$  artifacts were found in DMA case 4 (where the aerosol/sheath ratio was the highest)  
264 for both ammonium sulfate and sodium chloride aerosols. The artifacts for ammonium sulfate in DMA  
265 case 4 were 0.05-0.08, or 8-13 % of the literature value used for  $\kappa_{app}^{(NH_4)_2SO_4}$ , while the sodium chloride  
266 artifacts in DMA case 4 were 0.11-0.18, or 9-14 % of the literature value used for  $\kappa_{app}^{NaCl}$ . Artifacts were  
267 also high for DMA case 6 ( $-0.025 \leq \kappa_{app,artifact}^{(NH_4)_2SO_4} \leq -0.018$ ) and DMA case 7 ( $0.016 \leq$   
268  $\kappa_{app,artifact}^{(NH_4)_2SO_4} \leq 0.017$ ), where  $Q_e > Q_{sh}$ .  
269  
270  $\kappa_{app}$  artifacts were larger for sodium chloride ( $-0.05 \leq \kappa_{app,artifact}^{NaCl} \leq 0.18$ , 4 – 14 % of  $\kappa_{app}^{NaCl}$ ) than  
271 for ammonium sulfate ( $-0.03 \leq \kappa_{app,artifact}^{(NH_4)_2SO_4} \leq 0.08$ , 5 – 13 % of  $\kappa_{app}^{(NH_4)_2SO_4}$ ) across the DMA cases.  
272 As our results show, when two or more compounds are compared, the more hygroscopic compound will  
273 have larger  $\kappa_{app}$  artifacts.



### 274 3. Artifacts derived from condensation particle counters

275

#### 276 3.1. CPC operation at low concentration

277

278 The second instrument which must function accurately during CCN experiments is the condensation  
279 particle counter. CPC performance is characterized by the maximum counting efficiency (which may be  
280 influenced by the working fluid in the instrument) and the 50 %-cut-off diameter ( $d_{50}$ ), the particle  
281 diameter at which 50 % counting efficiency is observed, both of which can vary between commercially  
282 available models and even between individual CPCs (Heim et al., 2004). One study found that n-  
283 butanol CPCs (TSI, Inc. Models 3772, 3775, and 3776) exhibited smaller  $d_{50}$  for silver particles than  
284 sodium chloride ( $3.3 \text{ nm} \leq d_{50}^{\text{Ag}} \leq 7.8 \text{ nm}$  and  $4.1 \text{ nm} \leq d_{50}^{\text{NaCl}} \leq 14.7 \text{ nm}$ ), due to the more effective  
285 condensation of n-butanol on silver particles (Hermann et al., 2007).

286

287 Maximum counting efficiencies in that study varied from 88.9 % to 100.3 %. Another comparison of n-  
288 butanol CPCs (TSI Inc. Models 3010 and 3022, Grimm Tech. Inc. Model 5.403) found  $3.1 \text{ nm} \leq d_{50} \leq$   
289  $11.9 \text{ nm}$  for sodium chloride aerosols (Heim et al., 2004). In another study, the counting efficiencies  
290 observed in measurements of tungsten oxide particles by different instruments of the same model (TSI  
291 3025) varied from 88.9 % to 138.9 %, while  $d_{50}^{\text{WO}_x}$  varied from 3.2 nm to 11.0 nm (Hameri et al., 2002).

292

293 While some issues can cause undercounting at all concentrations, the additional issue of uncounted  
294 particles due to the arrival of more than one particle in the detector's field of view at any time arises  
295 only at higher concentrations. The cut-off between "low" and "high" concentration is not exact and  
296 varies between instruments. CPC undercounting issues which arise even at relatively low concentrations  
297 will be discussed in this section, and concentration-dependent effects will be explored in Sect. 3.2.



298

299 Six counting efficiency curves were generated using sigmoidal distributions and the 50 % cut-off  
300 diameters and maximum counting efficiencies listed in Table 3. Chosen values represent  $d_{50}$  values and  
301 maximum counting efficiencies reported in the literature under relatively low concentrations of 1000-  
302  $4000 \text{ cm}^{-3}$  (Hermann et al., 2007). The resulting sigmoidal distributions (Fig. 5a) were used to  
303 determine the counting efficiency of 25, 50, 100, and 200 nm particles.

304

305 Next,  $\kappa_{app}$  was calculated from theoretical critical supersaturations for each chosen diameter. To do so,  
306 four sigmoid curves representing sodium chloride CCN activation (hereon referred to as “activation  
307 curves”) for 25, 50, 100, and 200 nm were generated. The  $\kappa$ -Köhler-derived critical supersaturation of  
308 sodium chloride was used as the midpoint of each activation curve, and one-tenth of this value was used  
309 as the standard deviation (100 % CE, Fig. 5b-e). These values are consistent with the standard  
310 deviation/midpoint ratio observed from our instrument’s ammonium sulfate CCN calibration data.

311

312 Activation curves were then generated for CPC Cases 1-6 by dividing the activated fraction for each dry  
313 particle diameter by the counting efficiency for that diameter. Critical supersaturation was determined  
314 for each CPC case. Results are summarized in Fig. 5f. Next, critical supersaturation was converted to  
315 saturation, and  $\kappa_{app,theory}$  was calculated for each diameter in each CPC Case using Eq. (4) (see Fig.  
316 5g). As above,  $\kappa_{app}$  artifacts were calculated by finding the difference between these results and the  
317 literature value of  $\kappa_{app}$  for sodium chloride (see Fig. 5h).

318

319 For the diameters studied, the effect of maximum counting efficiency on CPC concentration (and  
320 activated fraction) is greater than the effect of 50 %-cutoff diameter. However, neither characteristic



321 resulted in large  $\kappa_{app}$  artifacts. The largest  $\kappa_{app}$  artifact observed at “low” concentrations was 0.035 for  
322 CPC Case 4, 2.4 % of the literature value for the apparent hygroscopicity factor for sodium chloride.



### 323 **3.2. CPC operation at high concentration**

324

325 Operation at high concentrations introduces an additional source of undercounting through particle  
326 coincidence at the CPC optical counter. For the TSI 3010 CPC, undercounting is observed is for particle  
327 concentrations above  $1 \times 10^4 \text{ cm}^{-3}$ . At  $5 \times 10^4 \text{ cm}^{-3}$ , the detector saturates and cannot detect higher  
328 concentrations. By comparison, the TSI 3025 is effective at counting higher particle concentrations, of  
329 up to  $2.5 \times 10^4 \text{ cm}^{-3}$  (Hameri et al., 2002; Sem, 2002).

330

331 To model undercounting due to particle coincidence, four CPC counting curves (Fig. 6a) were generated  
332 using the equations in Table 4. Case 7 represents a CPC where counting efficiency decreases with  
333 particle concentration, without reaching saturation. Cases 8-10 represent CPCs where saturation is  
334 reached at  $4 \times 10^4 \text{ cm}^{-3}$ ,  $2 \times 10^4 \text{ cm}^{-3}$ , and  $1 \times 10^4 \text{ cm}^{-3}$ , respectively. These saturation  
335 concentrations are of similar magnitude to those observed from TSI 3010 concentration data.

336

337 In all cases, the aerosol population was assumed to be a distribution consisting of  $5 \times 10^6 \text{ particles/cm}^3$   
338 with a peak concentration at 50 nm (Fig. 6b). Each case was applied to this theoretical distribution in  
339 order to determine the concentration measured by the CPC for 25, 50, 100 and 200 nm aerosols. Then,  
340 the counting efficiency was calculated for each case and aerosol diameter.

341

342 Sigmoidal activated fraction curves were generated for 25, 50, 100 and 200 nm sodium chloride  
343 aerosols. As in the low concentration cases, the midpoint of each 100 % CE curve was chosen to be  
344 equal to the  $\kappa$ -Köhler-derived critical supersaturation of sodium chloride at each dry diameter, and the  
345 standard deviation of each curve is equal to one-tenth of the critical supersaturation. These activated  
346 fraction curves were adjusted using the counting efficiencies calculated in the previous step. In cases



347 where the activated fraction has increased due to undercounting by the CPC, the theoretical sigmoidal  
348 curve shifts to the left relative to the 100 % CE case (Fig. 7c-f). Thus, undercounting by the CPC  
349 effectively increases the reported activated fraction. As above, critical supersaturation was determined  
350 from each of these curves, and  $\kappa_{app,theory}$  was subsequently calculated using Eq. (4) (Fig. 7g-h).

351

352  $\kappa_{app,theory}$  fell over a much wider range for 25, 50, and 100 nm particles (1.30-1.56, 1.32-1.70, and  
353 1.30-1.55, respectively) than for 200 nm particles (1.28-1.29) due to the lower concentration of 200 nm  
354 particles in the chosen aerosol distribution, which resulted in a higher counting efficiency for these  
355 aerosols. In comparison, the largest range in  $\kappa_{app,theory}$  was observed for 50 nm aerosols, the peak  
356 diameter in this aerosol distribution.

357

358 A wider range in  $\kappa_{app,theory}$  was observed for the high-concentration CPC Cases (7-10) compared to the  
359 low-concentration CPC Cases (1-6). The lowest counting efficiency observed across the low-  
360 concentration cases was 89.9 % for 25 nm aerosol in Case 4, while the lowest counting efficiency  
361 observed in the high-concentration cases was 18.0 % for 50 nm aerosol in Case 10.

362

363 Artifacts in the apparent hygroscopicity parameter are shown in Fig. 6i.  $\kappa_{app}$  artifacts were the greatest  
364 for a CPC that becomes saturated at 20,000 particles/cm<sup>3</sup> ( $0.0131 \leq \kappa_{app} \leq 0.4206$ ). The lower the  
365 concentration at which a CPC becomes saturated, the more quickly its counting efficiency will drop as  
366 concentration increases, resulting in increased activated fraction and increased apparent hygroscopicity.  
367 The magnitude of artifacts due to CPC undercounting depends on the saturation concentration of the  
368 CPC and the distribution of the aerosol population being studied.

369

370



371 **4. Artifacts derived from cloud condensation nuclei instruments**

372

373 Finally, the third instrument whose performance accuracy contributes to the overall certainty in CCN  
374 assessment in the CCN instrument itself. Several instruments have been implemented for measuring  
375 CCN concentrations over the last few decades. Older models include the Continuous Flow Parallel Plate  
376 Diffusion Chamber (Sinnarwalla, 1973) and the Hudson CCN spectrometer (Hudson, 1989) which both  
377 employ an applied temperature gradient perpendicular to the aerosol flow. Newer models, such as the  
378 widely-used Droplet Measurement Technology Cloud Condensation Nuclei Counter (DMT CCN-100),  
379 operate with a streamwise temperature gradient and continuous, laminar flow (Lance et al., 2006). The  
380 following analysis considers the DMT CCN-100. According to the CCN-100 manual, the counting  
381 efficiency for this CCN instrument depends on concentration and supersaturation (Fig. 6a). The  
382 counting efficiency decreases rapidly with concentration at  $< 0.2$  % SS due to rapid water vapor  
383 depletion at these low supersaturations, and falls off more slowly for  $> 0.2$  % SS (DMT CCN-100  
384 manual).

385

386 The counting efficiency of the DMT CCN-100 was tested for four lognormal aerosol distributions with  
387 peak concentrations at 50 nm and varying total concentrations (Table 5, Fig. 7b). Note that for  
388 comparison, CCN Case 1 was chosen to be identical to the aerosol distribution used for the high-  
389 concentration CPC cases. CCN Cases 2-4 follow the same distribution shape, with decreased total  
390 particle concentration.

391

392 The counting efficiencies for each case were applied to theoretical sodium chloride sigmoidal activated  
393 fraction curves to produce normalized activated fraction curves (Fig. 8c-f). As above, the midpoint is  
394 set to the critical supersaturation of sodium chloride at each dry diameter, and the standard deviation is



395 assumed to be one-tenth of the critical supersaturation. CCN undercounting effectively decreases  
396 activated fraction, therefore shifting the activated fraction curve downwards and towards higher percent  
397 supersaturations. The opposite effect is observed when CPC undercounting occurs. Critical  
398 supersaturation was determined for each CCN case, as above (Fig. 7g). Values of  $SS_{crit}$  were then  
399 converted to saturation, and  $\kappa_{app,theory}$  was calculated using Eq. (4) (Fig. 7h).  
400  
401 Significant deviations from  $\kappa$ -Köhler theory were only observed in CCN Case 1, with total aerosol  
402 concentration  $5 \times 10^6$  particles/cm<sup>3</sup> (Fig. 8g-i). The largest deviation for CCN Case 1 was observed in  
403 100 nm particles ( $\kappa_{app,artifact} = -0.57$ ), due to the higher concentration of 100 nm particles compared  
404 to 25 and 200 nm particles, and the lower supersaturation necessary for activation. The largest artifacts  
405 across CCN Cases 2 and 3 were also observed for 100 nm particles, though no artifacts were observed  
406 for any particle diameter in CCN Case 4 due to the much lower concentrations.  
407  
408 Sodium chloride is very hygroscopic. It should be noted that aerosols consisting of less hygroscopic  
409 compounds will activate at higher supersaturations (> 0.2 % SS regime) which will lead to smaller  $\kappa_{app}$   
410 artifacts when the same aerosol distribution and total aerosol concentration is considered. If a mixture  
411 was considered (for example, sodium chloride with a non-hygroscopic species such as soot) the results  
412 may also be different. The shape of the aerosol distribution must also be taken into account. A  
413 distribution with a narrower peak than the one generated for this analysis would be at risk for larger  
414  $\kappa_{app}$  artifacts for any total aerosol concentration, and these artifacts would be greater at the peak  
415 diameter, while a broader distribution would result in less variation in  $\kappa_{app}$  artifacts for each particle  
416 diameter.

417 **5. Discussion**

418

419 A comparison of several instrument sources of error in CCN-derived  $\kappa_{app}$  is shown in Fig. 8. DMA  
420 Case 4, CPC Case 4, CPC Case 10, and CCN Case 1 represent the operating conditions that resulted in  
421 the largest  $\kappa_{app}$  artifacts in this study. In DMA Case 4, the aerosol/sheath ratio of 0.30 resulted in a  
422 broadened aerosol distribution downstream of the DMA. Compared to DMA Case 1, where  
423  $Q_a/Q_{sh} = 0.10$ , the downstream diameter range in DMA Case 4 was 300 % higher for 25 nm particles,  
424 resulting in a spread of 20-36 nm. Similarly, the diameter ranges for 50, 100, and 200 nm diameter  
425 were 220 %, 230 %, and 220 % wider than in Case 1, respectively. Compared to the most ideal DMA  
426 case presented in this study (DMA Case 2), where  $Q_a/Q_{sh} = 0.05$ , the downstream diameter range in  
427 DMA Case 4 was 700 % higher for 25 nm particles; the diameter ranges for 50, 100, and 200 nm  
428 diameter were 540 %, 560 %, and 520 % wider than in Case 2, respectively.

429

430 CPC Case 4 represents  $\kappa_{app}$  artifacts (0.031-0.035) due to undercounting that arises from poor  
431 maximum CPC counting efficiency (90 %), which may be observed when using butanol as a working  
432 fluid while measuring the concentration of inorganic aerosols. In contrast,  $\kappa_{app}$  artifacts are negligible  
433 ( $< 0.10$  % of  $\kappa_{app}^{NaCl}$ ) in CPC Case 3, where maximum counting efficiency = 100 %. CPC Cases 8 and 10  
434 represent undercounting at high concentration with CPCs where saturation is observed at  $4 \times 10^4 cm^{-3}$   
435 and  $1 \times 10^4 cm^{-3}$ , respectively. Counting efficiency drops off more rapidly with concentration in the  
436 latter case, resulting in  $\kappa_{app}$  artifacts that are highest at the peak of the aerosol distribution (0.1190 and  
437 0.4206 for 50 nm aerosols in CPC Cases 8 and 10, respectively).

438

439 CCN Case 1 represents CCN undercounting at high concentration (total aerosol concentration =  $5 \times$   
440  $10^6 cm^{-3}$ ). CCN undercounting is greatest for low supersaturation ( $< 0.2$  %) and high concentration,



441 resulting in the lowest counting efficiency and highest  $\kappa_{app}$  artifacts (- 0.57) for 100 nm aerosols in  
442 CCN Case 1. The largest CCN-derived  $\kappa_{app}$  artifact observed outside of CCN Case 1 was - 0.01 for  
443 100 nm aerosols in CCN Case 2.  
444  
445 The combined artifacts for the cases where the highest artifacts were observed (DMA Case 4, CPC Case  
446 4, CPC Case 10, CCN Case 1) are 0.21, 0.24, 0.32, and 0.21 for 25, 50, 100, and 200 nm particles  
447 respectively, as shown in Fig. 8. The combined artifacts for the lowest-artifact cases (DMA Case 2,  
448 CPC Case 3, and CCN Case 4) are < 0.001 except for 200 nm particles, where  $\kappa_{app,artifact} = 0.0013$ .



449 **Conclusions**

450

451 The sensitivity of weather and climate models to accuracy in CCN activation predictions has been  
452 demonstrated in other works. Possible sources of apparent hygroscopicity artifacts calculated from CCN  
453 measurements have been presented in this study. This analysis has focused on sodium chloride and  
454 ammonium sulfate aerosols, but it can be extended to other aerosol populations, including mixtures and  
455 field samples.

456

457 The largest artifacts ( $-0.57 < \kappa_{app,artifact} < 0.42$ ) in this study arise from undercounting by  
458 condensation particle counters and cloud condensation nuclei counters at high concentration. This  
459 problem arises in cases which represent attempts to measure aerosol concentrations much higher than  
460 are recommended for these instruments,  $\sim 10^4 \text{ cm}^{-3}$  (CPC Cases 7-10 and CCN Case 4). Corrective  
461 action should be taken to dilute aerosol samples in order to avoid undercounting.

462

463 Smaller artifacts ( $\kappa_{app,artifact} < 0.04$ ) were observed for the CPC cases where 50 %-cut-off diameter  
464 and maximum counting efficiency were varied. Given the chosen particle diameters (25, 50, 100, 200  
465 nm),  $\kappa_{app}$  artifacts due to  $d_{50}$  were minimal. The largest  $\kappa_{app}$  artifacts for a CPC counting at low  
466 concentration (0.031-0.035) were observed where the maximum counting efficiency was equal 0.90.  
467 This may represent a compositional mismatch between n-butanol as the working fluid and sodium  
468 chloride as the aerosol, due to the poor solubility of the latter in the former. Individual n-butanol CPCs  
469 may exhibit higher maximum counting efficiencies for sodium chloride.

470

471 Uncertainty arising from the DMA depended greatly on the chosen aerosol and sheath settings. One set  
472 of DMA cases (Cases 2-4) examined the effect of aerosol/sheath ratio. By decreasing this ratio, a



473 narrower near-monodisperse flow can be produced, which increases the accuracy of calculated  $\kappa_{app}$ .

474 The  $\kappa_{app}$  artifacts for an aerosol/sheath ratio of 0.10 were 1.1 % of  $\kappa_{literature}$  for 25 nm sodium  
475 chloride aerosols, 0.93 % for 50 nm, 0.81 % for 100 nm, and 0.66 % for 200 nm. Decreasing the  
476 aerosol/sheath ratio to 0.05 resulted in  $\kappa_{app} \approx 0.01$  % of  $\kappa_{literature}$  for NaCl. When a ratio of 0.30 was  
477 employed, the resulting artifacts that were  $\sim 10$  % of  $\kappa_{literature}$  for sodium chloride.

478

479 The second set of DMA cases (5-7) were designed to evaluate the effects of holding the sheath flow  
480 constant while varying the excess air flow by -2 %, +2 %, and +5 %. These resulted in shifts of  $\leq 2$  nm  
481 for 25 nm and 50 nm particles,  $\leq 4$  nm for 100 nm particles, and  $\leq 7$  nm for 200 nm particles. The  
482 downstream aerosol distribution was shifted towards larger particle diameters when sheath flow  
483 exceeded excess flow, and towards smaller particle diameter when sheath flow was less than excess  
484 flow. When taking field measurements, the composition of the sample may vary with particle diameter,  
485 thereby introducing another source of error from a broader DMA distribution.

486

487 Under optimal operating conditions, where the DMA sample/sheath ratio is 0.10 and excess/sheath ratio  
488 is 1.0, and in the absence of undercounting by the CPC or CCN, uncertainties in  $\kappa_{app}$  are within  $\pm 1.1$  %  
489 for 25 to 200 nm aerosols. When the DMA sample/sheath ratio drops to 0.05,  $\kappa_{app}$  uncertainties  
490 decrease to  $\pm 0.01$  %.

491

492 Apparent hygroscopicity parameter artifacts were calculated for two pure, inorganic species in this  
493 study. This analysis could be used to estimate  $\kappa_{app}$  artifacts for ambient aerosol populations, which  
494 may result in a better understanding of the “real” differences between these populations. As discussed in  
495 the introduction, Collins et al. 2016 aggregated  $\kappa_{app}$  from several mesocosm and field studies for 30-  
496 100 nm sea spray aerosol ( $0.4 < \kappa_{app}^{SSA} < 1.3$ ). The wide range of  $\kappa_{app}$  in these studies may be



497 attributed to differences in composition, experimental artifacts, or a combination of the two.  
498 Quantification of experimental artifacts would facilitate interpretation of  $\kappa_{app}$  in aerosol populations and  
499 constrain the importance of composition in CCN activation. There has been a recent proliferation of  
500 CCN data availability from multiple researchers and multiple experimental setups. To maximize the  
501 utility of these studies and to compare cloud-activating properties of various ambient aerosol masses, it  
502 is essential that artifacts are considered in both CCN data collection and in reporting of the data.



503 **Supplement Link**

504 Will be included by Copernicus

505

506 **Author Contribution**

507 Sarah D. Brooks provided the conceptual framework and reviewed the manuscript. Jessica A. Mirrielees  
508 performed the analysis and wrote the manuscript.

509

510 **Competing Interests**

511 The authors declare that they have no conflict of interest.

512

513 **Disclaimer**

514 Will be included by Copernicus

515

516 **Acknowledgements**

517 This project was supported by the National Science Foundation of the United States (Award  
518 #15398810). In addition, Mirrielees thanks Texas A&M University for support through Institute for and  
519 Advanced Studies HEEP PhD Fellowship and a Lechner Scholarship.



## 520 References

- 521 Asa-Awuku, A., Nenes, A., Gao, S., Flagan, R. C., and Seinfeld, J. H.: Water-soluble SOA from Alkene  
522 ozonolysis: composition and droplet activation kinetics inferences from analysis of CCN activity,  
523 *Atmos. Chem. Phys.*, 10, 1585-1597, 10.5194/acp-10-1585-2010, 2010.
- 524 Barmounis, K., Maissner, A., Schmidt-Ott, A., and Biskos, G.: Lightweight differential mobility  
525 analyzers: Toward new and inexpensive manufacturing methods, *Aerosol Sci. Technol.*, 50, 4,  
526 10.1080/02786826.2015.1130216, 2016.
- 527 Betancourt, R. M., Nenes, A., and Liu, X. H.: Relative Contributions of Aerosol Properties to Cloud  
528 Droplet Number: Adjoint Sensitivity Approach in a GCM, in: AIP Conference Proceedings, 19th  
529 International Conference on Nucleation and Atmospheric Aerosols (ICNAA), Colorado State Univ, Ctr  
530 Arts, Fort Collins, CO, 2013, WOS:000319766400170, 679-682, 2013.
- 531 Betancourt, R. M., and Nenes, A.: Understanding the contributions of aerosol properties and  
532 parameterization discrepancies to droplet number variability in a global climate model, *Atmos. Chem.*  
533 *Phys.*, 14, 4809-4826, 10.5194/acp-14-4809-2014, 2014.
- 534 Brooks, S. D., and Thornton, D. C. O.: Marine Aerosols and Clouds, in: Annual Review of Marine  
535 Sciences, Vol 10, edited by: Carlson, C. A., and Giovannoni, S. J., Annual Review of Marine Science,  
536 289-313, 2018.
- 537 Carrico, C. M., Petters, M. D., Kreidenweis, S. M., Collett, J. L., Engling, G., and Malm, W. C.: Aerosol  
538 hygroscopicity and cloud droplet activation of extracts of filters from biomass burning experiments, *J.*  
539 *Geophys. Res.-Atmos.*, 113, 9, 10.1029/2007jd009274, 2008.
- 540 Carslaw, K. S., Lee, L. A., Reddington, C. L., Pringle, K. J., Rap, A., Forster, P. M., Mann, G. W.,  
541 Spracklen, D. V., Woodhouse, M. T., Regayre, L. A., and Pierce, J. R.: Large contribution of natural  
542 aerosols to uncertainty in indirect forcing, *Nature*, 503, 67-+, 10.1038/nature12674, 2013.
- 543 Chang, D. Y., Lelieveld, J., Tost, H., Steil, B., Pozzer, A., and Yoon, J.: Aerosol physicochemical  
544 effects on CCN activation simulated with the chemistry-climate model EMAC, *Atmospheric*  
545 *Environment*, 162, 127-140, 10.1016/j.atmosenv.2017.03.036, 2017.
- 546 Clegg, S. L., Brimblecombe, P., and Wexler, A. S.: Thermodynamic model of the system H<sup>+</sup>-NH<sub>4</sub><sup>+</sup>-  
547 SO<sub>4</sub><sup>2-</sup>-NO<sub>3</sub><sup>-</sup>-H<sub>2</sub>O at tropospheric temperatures, *J. Phys. Chem. A*, 102, 2137-2154, 10.1021/jp973042r,  
548 1998.
- 549 Collins, D. B., Bertram, T. H., Sultana, C. M., Lee, C., Axson, J. L., and Prather, K. A.: Phytoplankton  
550 blooms weakly influence the cloud forming ability of sea spray aerosol, *Geophys. Res. Lett.*, 43, 9975-  
551 9983, 10.1002/2016gl069922, 2016.
- 552 Crosbie, E., Youn, J. S., Balch, B., Wonaschutz, A., Shingler, T., Wang, Z., Conant, W. C., Betterton, E.  
553 A., and Sorooshian, A.: On the competition among aerosol number, size and composition in predicting  
554 CCN variability: a multi-annual field study in an urbanized desert, *Atmos. Chem. Phys.*, 15, 6943-6958,  
555 10.5194/acp-15-6943-2015, 2015.
- 556 DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle morphology  
557 and density characterization by combined mobility and aerodynamic diameter measurements. Part 1:  
558 Theory, *Aerosol Sci. Technol.*, 38, 1185-1205, 10.1080/027868290903907, 2004.
- 559 Hameri, K., O'Dowd, C. D., and Hoell, C.: Evaluating measurements of new particle concentrations,  
560 source rates, and spatial scales during coastal nucleation events using condensation particle counters, *J.*  
561 *Geophys. Res.-Atmos.*, 107, 11, 10.1029/2001jd000411, 2002.
- 562 Hartz, K. E. H., Tischuk, J. E., Chan, M. N., Chan, C. K., Donahue, N. M., and Pandis, S. N.: Cloud  
563 condensation nuclei activation of limited solubility organic aerosol, *Atmospheric Environment*, 40, 605-  
564 617, 10.1016/j.atmosenv.2005.09.076, 2006.
- 565 Heim, M., Kasper, G., Reischl, G., and Gerhart, C.: Performance of a New Commercial Electrical  
566 Mobility Spectrometer, *Aerosol Sci. Technol.*, 38, 3-14, 10.1080/02786820490519252, 2004.



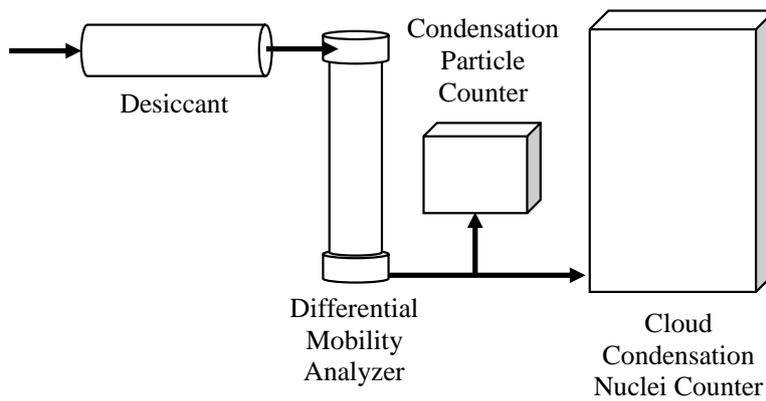
- 567 Hermann, M., Wehner, B., Bischof, O., Han, H. S., Krinke, T., Liu, W., Zerrath, A., and Wiedensohler,  
568 A.: Particle counting efficiencies of new TSI condensation particle counters, *J. Aerosol. Sci.*, 38, 674-  
569 682, 10.1016/j.jaerosci.2007.05.001, 2007.
- 570 Hudson, J. G.: An Instantaneous CCN Spectrometer, *J. Atmos. Ocean. Technol.*, 6, 1055-1065,  
571 10.1175/1520-0426(1989)006<1055:aics>2.0.co;2, 1989.
- 572 IPCC, 2013: Summary for Policymakers, Intergovernmental Panel on Climate Change, Cambridge  
573 University Press, 14, 2013.
- 574 Jokinen, V., and Makela, J. M.: Closed-loop arrangement with critical orifice for DMA sheath excess  
575 flow system, *J. Aerosol. Sci.*, 28, 643-648, 10.1016/s0021-8502(96)00457-0, 1997.
- 576 Karydis, V. A., Capps, S. L., Russell, A. G., and Nenes, A.: Adjoint sensitivity of global cloud droplet  
577 number to aerosol and dynamical parameters, *Atmos. Chem. Phys.*, 12, 9041-9055, 10.5194/acp-12-  
578 9041-2012, 2012.
- 579 Kawecki, S., and Steiner, A. L.: The Influence of Aerosol Hygroscopicity on Precipitation Intensity  
580 During a Mesoscale Convective Event, *J. Geophys. Res.-Atmos.*, 123, 424-442, 10.1002/2017jd026535,  
581 2018.
- 582 Knutson, E. O., and Whitby, K. T.: Aerosol classification by electric mobility: apparatus, theory, and  
583 applications, *J. Aerosol. Sci.*, 6, 443-451, [https://doi.org/10.1016/0021-8502\(75\)90060-9](https://doi.org/10.1016/0021-8502(75)90060-9), 1975.
- 584 Koehler, K. A., Kreidenweis, S. M., DeMott, P. J., Prenni, A. J., Carrico, C. M., Ervens, B., and  
585 Feingold, G.: Water activity and activation diameters from hygroscopicity data - Part II: Application to  
586 organic species, *Atmos. Chem. Phys.*, 6, 795-809, 10.5194/acp-6-795-2006, 2006.
- 587 Kumar, P. P., Broekhuizen, K., and Abbatt, J. P. D.: Organic acids as cloud condensation nuclei:  
588 Laboratory studies of highly soluble and insoluble species, *Atmos. Chem. Phys.*, 3, 509-520, 2003.
- 589 Lance, S., Medina, J., Smith, J. N., and Nenes, A.: Mapping the operation of the DMT Continuous Flow  
590 CCN counter, *Aerosol Sci. Technol.*, 40, 242-254, 10.1080/02786820500543290, 2006.
- 591 Liu, X. H., and Wang, J. A.: How important is organic aerosol hygroscopicity to aerosol indirect  
592 forcing?, *Environ. Res. Lett.*, 5, 10, 10.1088/1748-9326/5/4/044010, 2010.
- 593 Mamakos, A., Ntziachristos, L., and Sarnaras, Z.: Diffusion broadening of DMA transfer functions.  
594 Numerical validation of Stolzenburg model, *J. Aerosol. Sci.*, 38, 747-763,  
595 10.1016/j.jaerosci.2007.05.004, 2007.
- 596 Mei, F., Fu, H. J., and Chen, D. R.: A cost-effective differential mobility analyzer (cDMA) for multiple  
597 DMA column applications, *J. Aerosol. Sci.*, 42, 462-473, 10.1016/j.jaerosci.2011.04.001, 2011.
- 598 Moore, R. H., Raatikainen, T., Langridge, J. M., Bahreini, R., Brock, C. A., Holloway, J. S., Lack, D.  
599 A., Middlebrook, A. M., Perring, A. E., Schwarz, J. P., Spackman, J. R., and Nenes, A.: CCN Spectra,  
600 Hygroscopicity, and Droplet Activation Kinetics of Secondary Organic Aerosol Resulting from the 2010  
601 Deepwater Horizon Oil Spill, *Environ. Sci. Technol.*, 46, 3093-3100, 10.1021/es203362w, 2012.
- 602 Ovadnevaite, J., Ceburnis, D., Martucci, G., Bialek, J., Monahan, C., Rinaldi, M., Facchini, M. C.,  
603 Berresheim, H., Worsnop, D. R., and O'Dowd, C.: Primary marine organic aerosol: A dichotomy of low  
604 hygroscopicity and high CCN activity, *Geophys. Res. Lett.*, 38, 5, 10.1029/2011gl048869, 2011.
- 605 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and  
606 cloud condensation nucleus activity, *Atmos. Chem. Phys.*, 7, 1961-1971, 10.5194/acp-7-1961-2007,  
607 2007.
- 608 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and  
609 cloud condensation nucleus activity - Part 3: Including surfactant partitioning, *Atmos. Chem. Phys.*, 13,  
610 1081-1091, 10.5194/acp-13-1081-2013, 2013.
- 611 Rogers, R. R., and Yau, M. K.: in: *A Short Course In Cloud Physics*, Third ed., edited by: Haar, D. T.,  
612 Elsevier Science Inc., Tarrytown, New York, USA, 87-89, 1989.
- 613 Rose, D., Gunthe, S. S., Mikhailov, E., Frank, G. P., Dusek, U., Andreae, M. O., and Poschl, U.:  
614 Calibration and measurement uncertainties of a continuous-flow cloud condensation nuclei counter



615 (DMT-CCNC): CCN activation of ammonium sulfate and sodium chloride aerosol particles in theory  
616 and experiment, *Atmos. Chem. Phys.*, 8, 1153-1179, 10.5194/acp-8-1153-2008, 2008.  
617 Scanning Mobility Particle Sizer Series 5.400 and 5.500, Grimm Aerosol Technik, Ainring, Germany,  
618 2009.  
619 Schmale, J., Henning, S., Decesari, S., Henzing, B., Keskinen, H., Sellegri, K., Ovadnevaite, J., Pohlker,  
620 M. L., Brito, J., Bougiatioti, A., Kristensson, A., Kalivitis, N., Stavroulas, I., Carbone, S., Jefferson, A.,  
621 Park, M., Schlag, P., Iwamoto, Y., Aalto, P., Aijala, M., Bukowiecki, N., Ehn, M., Frank, G., Frohlich,  
622 R., Frumau, A., Herrmann, E., Herrmann, H., Holzinger, R., Kos, G., Kulmala, M., Mihalopoulos, N.,  
623 Nenes, A., O'Dowd, C., Petaja, T., Picard, D., Pohlker, C., Poschl, U., Poulain, L., Prevot, A. S. H.,  
624 Swietlicki, E., Andreae, M. O., Artaxo, P., Wiedensohler, A., Ogren, J., Matsuki, A., Yum, S. S.,  
625 Stratmann, F., Baltensperger, U., and Gysel, M.: Long-term cloud condensation nuclei number  
626 concentration, particle number size distribution and chemical composition measurements at regionally  
627 representative observatories, *Atmos. Chem. Phys.*, 18, 2853-2881, 10.5194/acp-18-2853-2018, 2018.  
628 Sem, G. J.: Design and performance characteristics of three continuous-flow condensation particle  
629 counters: a summary, *Atmos. Res.*, 62, 267-294, 10.1016/s0169-8095(02)00014-5, 2002.  
630 Seol, K. S., Tsutani, Y., Camata, R. P., Yabumoto, J., Isomura, S., Okada, Y., Okuyama, K., and  
631 Takeuchi, K.: A differential mobility analyzer and a Faraday cup electrometer for operation at 200-930  
632 Pa pressure, *J. Aerosol. Sci.*, 31, 1389-1395, 10.1016/s0021-8502(00)00037-9, 2000.  
633 Sinarwalla, A. M. a. A., D.J.: A Cloud Nucleus Counter with Long Available Growth Time, *Journal of*  
634 *Applied Meteorology*, 831-835, 1973.  
635 Stratmann, F., Kauffeldt, T., Hummes, D., and Fissan, H.: Differential electrical mobility analysis: A  
636 theoretical study, *Aerosol Sci. Technol.*, 26, 368-383, 10.1080/02786829708965437, 1997.  
637 Sullivan, R. C., Moore, M. J. K., Petters, M. D., Kreidenweis, S. M., Roberts, G. C., and Prather, K. A.:  
638 Timescale for hygroscopic conversion of calcite mineral particles through heterogeneous reaction with  
639 nitric acid, *Phys. Chem. Chem. Phys.*, 11, 7826-7837, 10.1039/b904217b, 2009.  
640 Svenningsson, B., Rissler, J., Swietlicki, E., Mircea, M., Bilde, M., Facchini, M. C., Decesari, S., Fuzzi,  
641 S., Zhou, J., Monster, J., and Rosenorn, T.: Hygroscopic growth and critical supersaturations for mixed  
642 aerosol particles of inorganic and organic compounds of atmospheric relevance, *Atmos. Chem. Phys.*, 6,  
643 1937-1952, 10.5194/acp-6-1937-2006, 2006.  
644

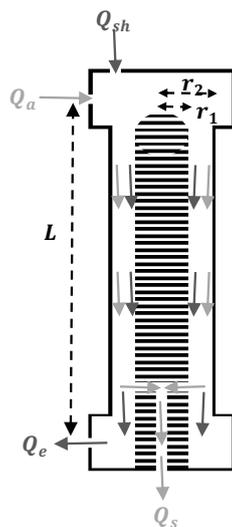


Notation	
$\alpha_{cc}, \beta_{cc}, \gamma_{cc}$	Empirically-determined constants used to calculate Cunningham slip correction factor
$Z_p$	Aerosol particle electrical mobility
$C_c$	Cunningham slip correction factor
$d_m$	Electrical mobility diameter
$n$	Number of charges on particle
$e$	Elementary unit of charge
$\eta$	Gas dynamic viscosity
$\lambda$	Mean free path
$Q_{sh}$	Sheath flow
$Q_e$	Excess air flow
$Q_a$	Aerosol flow
$Q_s$	Sample flow
$\kappa_{app}$	Apparent hygroscopicity parameter
$\kappa_{app,artifact}$	Apparent hygroscopicity parameter artifact
$s$	Equilibrium water vapor saturation
$s_{crit}$	Critical saturation (50 % of aerosols active as cloud condensation nuclei)
$A$	Constant used in calculating $\kappa_{app}$
$\sigma_{lv}$	Surface tension of water
$T$	Temperature
$D_{act}$	Activation diameter
$SS_{crit}$	Critical percent supersaturation
$\alpha_{TF}$	Height of DMA transfer function
$\beta_{TF}$	Half-width of DMA transfer function
$Z'_p$	Mobility of particle at DMA inlet
$Z_{p,mid}$	Midpoint of transfer function
$\Delta Z_p$	Half-width of transfer function
$V_0$	Voltage selected at DMA
$r_1$	DMA inner radius
$r_2$	DMA outer radius
$L$	DMA length
$d_{50}$	50 %-cut-off diameter



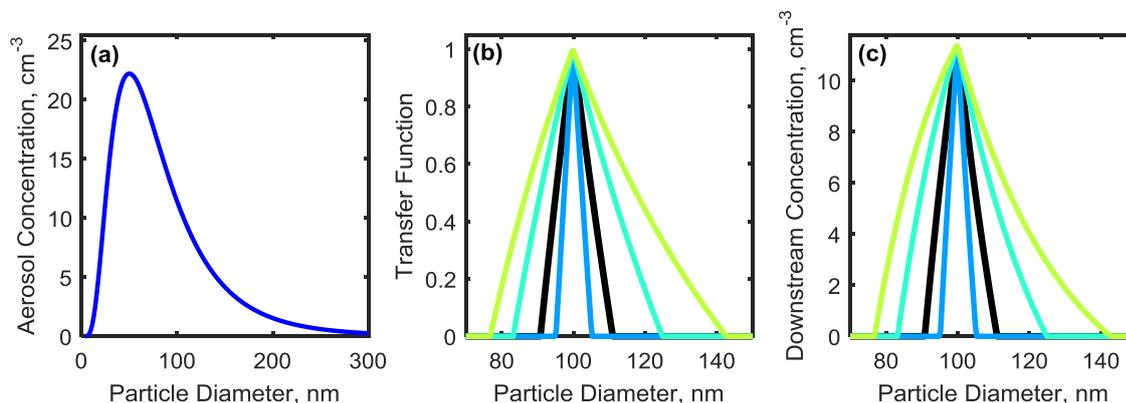
646

647 **Figure 1** Experimental setup for CCN measurements.



648

649 **Figure 2** Simplified flow diagram of a DMA with an inner electrode radius  $r_1$ , outer electrode radius  $r_2$ ,  
650 and distance between aerosol inlet and sample outlet  $L$ .  $Q_{sh}$  represents the clean sheath air flow,  $Q_a$   
651 represent the aerosol flow,  $Q_e$  represents the excess air flow, and  $Q_s$  represents the sample air flow.



652

653

654 **Figure 3** (a) A theoretical aerosol distribution generated using a lognormal function centered at 50 nm.

655 The total concentration for this distribution is 2000 particles/cm<sup>3</sup>. (b) The transfer function calculated

656 using Eq. (7). (c) Downstream aerosol concentration, cm<sup>-3</sup> (Distribution x transfer function = downstream

657 concentration).



**Table 1 Theoretical DMA Flow Test Cases**

Case	$Q_{sh}$ (L min <sup>-1</sup> )	$Q_e$ (L min <sup>-1</sup> )	$Q_a$ (L min <sup>-1</sup> )	$Q_s$ (L min <sup>-1</sup> )	$Q_a/Q_{sh}$	$Q_e/Q_{sh}$
<b>DMA 1</b>	3	3	0.3	0.3	0.1	1
<b>DMA 2</b>	3	3	0.15	0.15	0.05	1
<b>DMA 3</b>	3	3	0.6	0.6	0.2	1
<b>DMA 4</b>	3	3	0.9	0.9	0.3	1
<b>DMA 5</b>	3	3.06	0.36	0.3	0.12	1.02
<b>DMA 6</b>	3	3.15	0.45	0.3	0.15	1.05
<b>DMA 7</b>	3	2.94	0.24	0.3	0.08	0.98

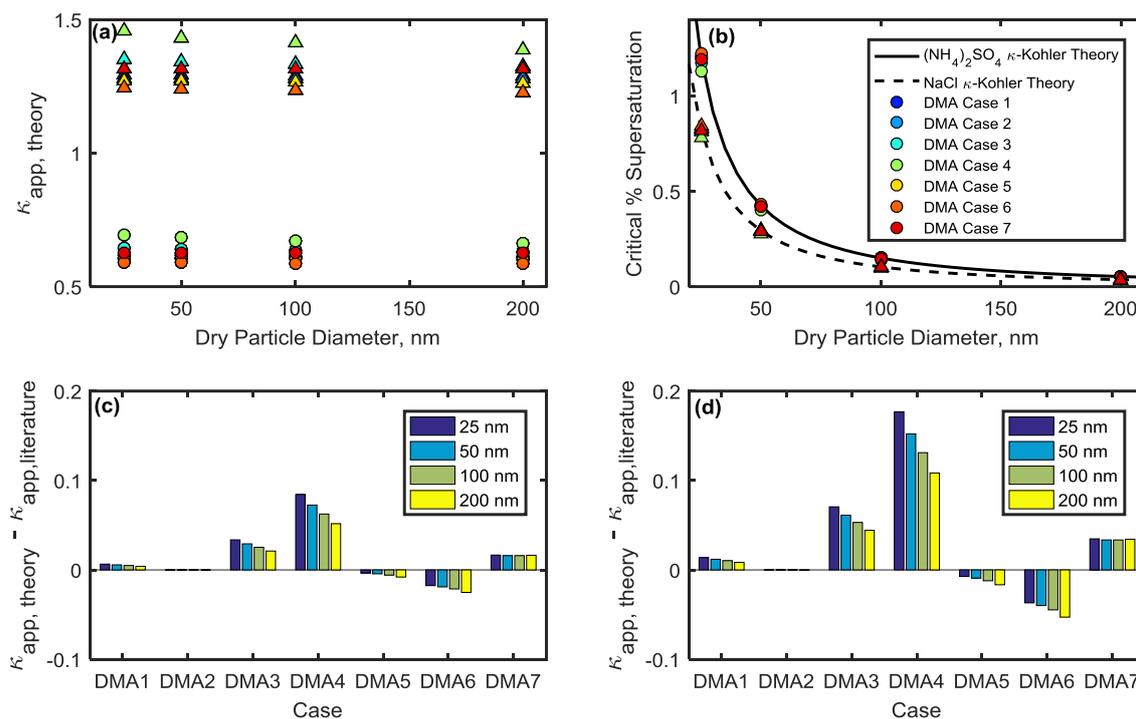
658



**Table 2** Predicted downstream particle diameter range for each DMA case.

Case	25 nm	50 nm	100 nm	200 nm
<b>DMA 1</b>	23-27	46-56	91-111	181-222
<b>DMA 2</b>	24-26	48-53	95-105	190-211
<b>DMA 3</b>	21-31	42-62	83-125	167-250
<b>DMA 4</b>	20-36	39-71	77-143	154-285
<b>DMA 5</b>	23-27	45-55	90-110	181-220
<b>DMA 6</b>	22-27	45-54	89-107	178-215
<b>DMA 7</b>	23-28	46-56	92-112	183-225

659



660

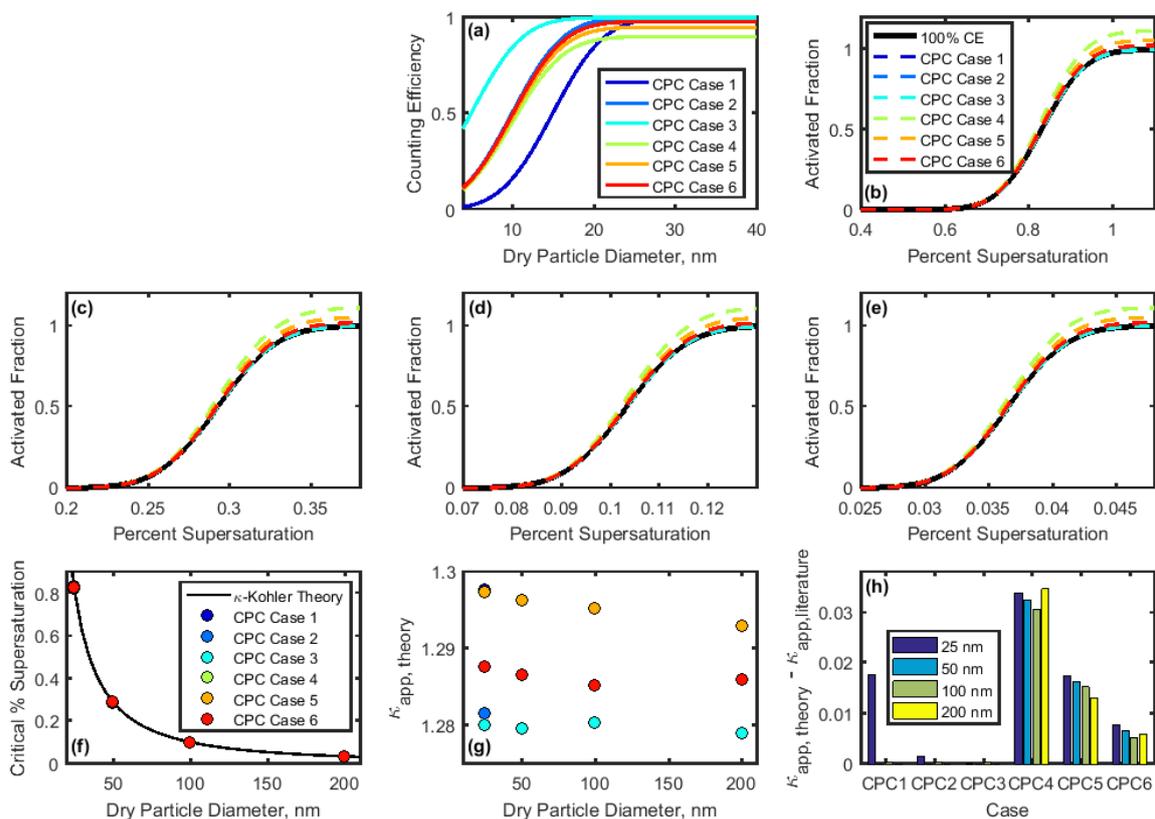
661 **Figure 4** (a) Apparent hygroscopicity  $\kappa_{app}$  for each DMA case. Triangles represent the results for  
 662 sodium chloride, and circles represent the results for ammonium sulfate; only one is shown in (b) to  
 663 conserve space. (b) The critical supersaturation of ammonium sulfate and sodium chloride particles  
 664 calculated for Cases 1-7 using  $\kappa_{app}$  values derived in (a). Ammonium sulfate and sodium chloride  
 665 curves from  $\kappa$ -Köhler theory are shown for comparison. (c,d) DMA-flow-derived artifacts in  $\kappa_{app}$  for  
 666 ammonium sulfate are shown for each DMA case for ammonium sulfate aerosols and sodium chloride  
 667 aerosols.



**Table 3** Parameters used in investigating  $\kappa_{app}$  artifacts for low concentrations measured by a condensation particle counter.

Case	$d_{50}$ , nm	Maximum Counting Efficiency
CPC 1	15	100 %
CPC 2	10	100 %
CPC 3	5	100 %
CPC 4	10	90 %
CPC 5	10	95 %
CPC 6	10	98 %

668



669

670 **Figure 5** (a) Counting efficiency curves based on CPC measurements.

671 (b-e) CCN activation for 25, 50, 100, and 200 nm NaCl, respectively. (f) Critical supersaturation

672 calculated for each activated fraction curve. (g) Apparent hygroscopicity for each case. (h)  $\kappa_{app}$

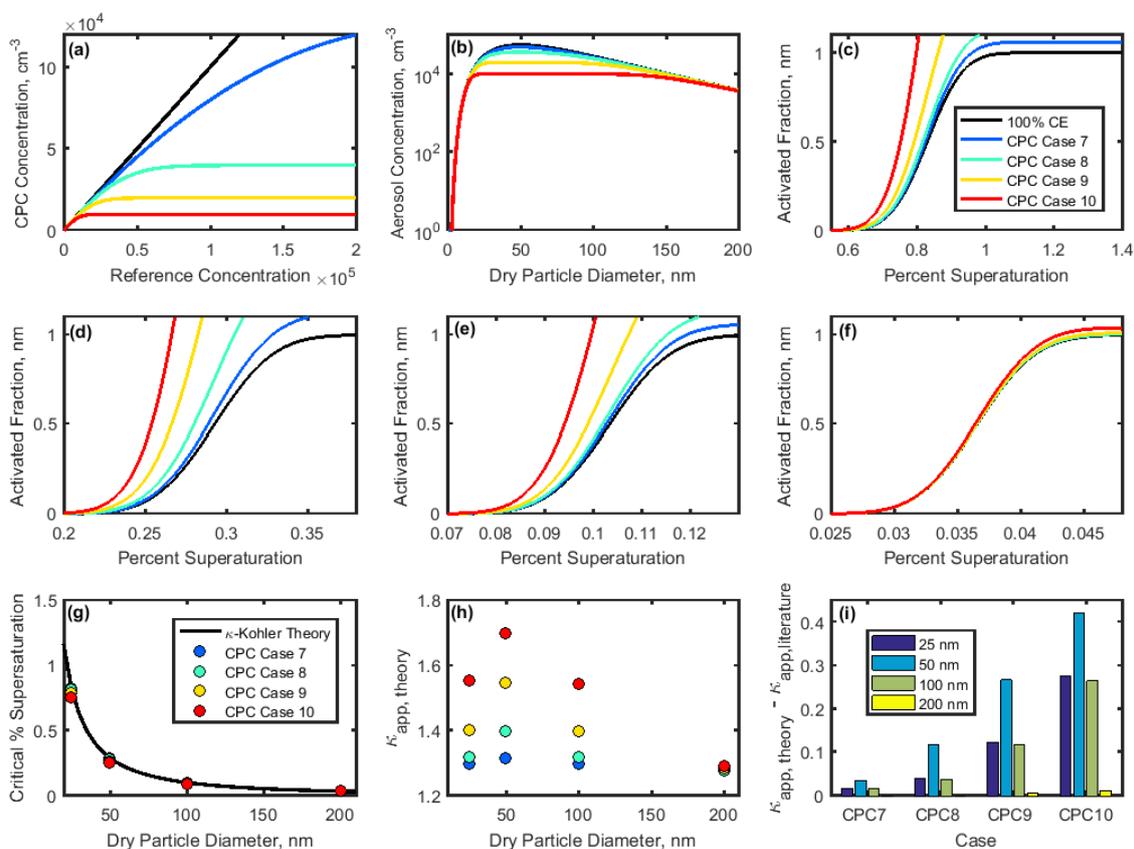
673 artifacts for each case.



**Table 4** Equations used to model the relationship between a reference or “true” aerosol concentration  $x$  (particles  $\text{cm}^{-3}$ ), and the concentration measured by a condensation particle counter  $y$  (particles  $\text{cm}^{-3}$ ).

Case	Equation
CPC 7	$y = x - 2 \times 10^{-6}x^2$
CPC 8	$y = 40000 \operatorname{erf}\left(\frac{x}{32000\sqrt{2}}\right)$
CPC 9	$y = 20000 \operatorname{erf}\left(\frac{x}{16000\sqrt{2}}\right)$
CPC 10	$y = 10000 \operatorname{erf}\left(\frac{x}{8000\sqrt{2}}\right)$

674



675

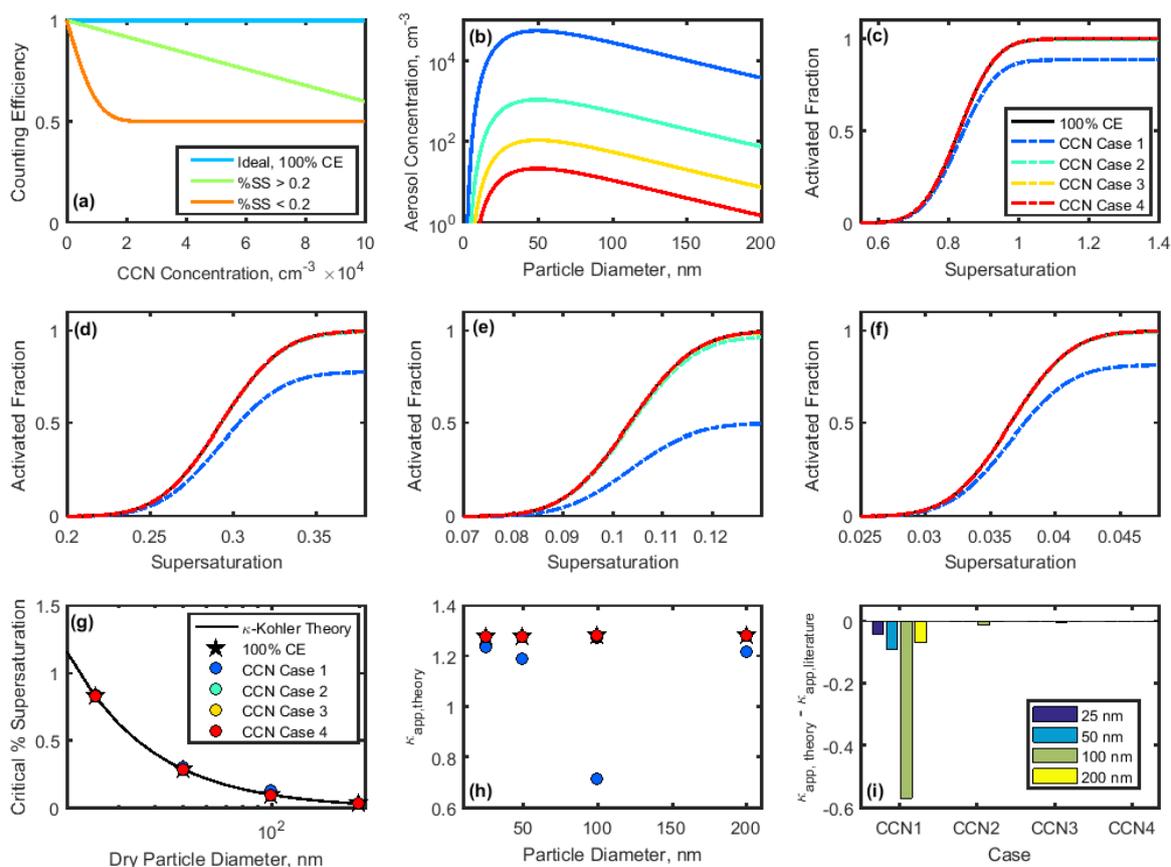
676 **Figure 6** (a) Theoretical relationships between the reference aerosol concentration and CPC  
 677 concentration. (b) Concentration-dependent counting efficiencies were applied to a theoretical aerosol  
 678 distribution. The ideal case (100 % counting efficiency), shown in black, is a lognormal distribution  
 679 centered at 50 nm, consisting of 5 × 10<sup>6</sup> total particles. (c-f) Activated fraction for 25, 50, 100, and 200  
 680 nm NaCl aerosol, respectively. (g,h) Critical supersaturation and apparent hygroscopicity for each  
 681 concentration-dependent case. (i)  $\kappa_{app}$  artifacts for each case.



**Table 5** Total concentrations used in theoretical aerosol distribution for CCN-derived  $\kappa_{app}$  artifacts.

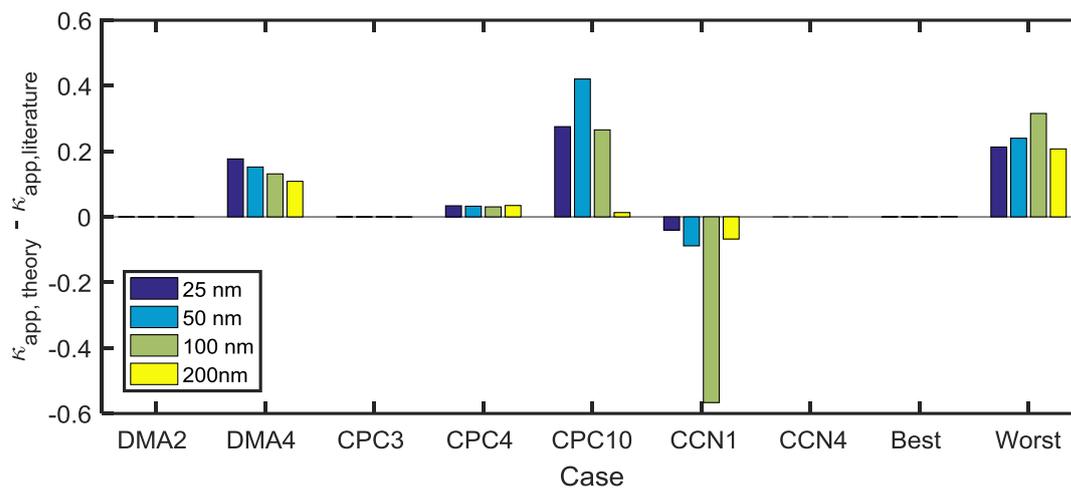
Case	Total Concentration (particles cm <sup>-3</sup> )
CCN 1	$5 \times 10^6$
CCN 2	$1 \times 10^5$
CCN 3	$1 \times 10^4$
CCN 4	$2 \times 10^3$

682



683

684 **Figure 7** (a) Counting efficiency of the DMT CCN-100. (b) Lognormal distributions used to study CCN  
 685 undercounting at high concentrations. (c-f) Activation fractions for 25-200 nm NaCl particles,  
 686 respectively. Supersaturation-specific counting efficiencies from (a) were applied to theoretical sigmoid  
 687 curves for NaCl CCN activation. Activated fraction in the case of 100 % counting efficiency is shown  
 688 for comparison. (g) Critical supersaturation for each case. (h)  $\kappa_{\text{app}}$  calculated for each case. (i)  $\kappa_{\text{app}}$   
 689 artifacts for each case.



690

691 **Figure 8** Comparison of kappa artifacts derived from best and worst case scenarios for instrumental  
 692 measurements for sodium chloride. Combined artifacts for the lowest-artifact cases (Best: DMA Case  
 693 2, CPC Case 3, and CCN Case 4) and the highest-artifact cases (Worst: DMA Case 4, CPC  
 694 Case 8, and CCN Case 1).

695