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Atmospheric Measurement Techniques Discussions

A new method to infer the size, number density, and charge of mesospheric dust from its in situ collection by the DUSTY probe.

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Abstract. We present a new extended method of analyzing measurements of mesospheric dust made with DUSTY rocket-borne Faraday cup probes. It yields the variation of fundamental dust parameters through a mesospheric cloud with an unrivalled altitude resolution down to 10 cm or less. A DUSTY probe was the first probe which unambiguously detected charged dust/aerosol particles in the Earth's mesosphere. DUSTY excluded the ambient plasma by various biased grids, which however allowed dust particles with radii above a few nanometer to enter, and it measured the flux of charged dust particles. The flux measurements directly yielded the total ambient dust charge density.

We extend the analysis of DUSTY data by using the impact currents on its main grid and the bottom plate as before, together with a dust charging model and a secondary charge production model, to allow the determination of fundamental parameters, such as dust radius, charge number and total dust density. We demonstrate the utility of the new analysis technique by considering observations made with the DUSTY probes during the MAXIDUSTY rocket campaign in June-July 2016 and comparing the results with those of other instruments (Lidar and photometer) also used in the campaign.

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1 Introduction.

The Earth's mesosphere has for a long time been the least known part of the Earth's atmosphere, and it probably still is. One reason for this is its inaccessibility to direct in situ observations – it being too high for balloons and planes, and too low for satellites. Its main cloud phenomena, the noctilucent clouds (NLC) which occurs in its polar regions, were first observed in 1885 (Jesse, 1885; Backhouse, 1885; Symons, 1888, Gadsden and Schröder, 1989). They are the highest altitude clouds in the Earth's atmosphere. It now appears that the NLC occurrence frequency is increasing with time and that the NLC spread further away from the poles with time (de Land et al., 2007), possibly due to changes in the composition of trace elements, like water vapor, in the mesosphere region. As such, one reason for the interest to understand the mesosphere is that it may be an indicator of climatic changes in the troposphere and stratosphere (Thomas, 1996). Another reason is that the mesosphere is the transition zone, between the outer space and the lower part of the atmosphere, where energetic particle precipitation, meteors and UV radiation normally deposits most of their energy. Disturbed magnetosphere conditions, with high energy particle precipitation, can create large amounts of reactive NO_x molecules which, when transported downwards, react with and reduce the ozone content (Reddman et al., 2013). Also, there is an influx of meteorites into the Earth's atmospheres, the total mass of which has been claimed to be from 4 to 300 t/day (Plane 2012; Asmus et al., 2015). Much of the meteorites evaporate as they are heated due to air friction when they enter the atmosphere, and the evaporated material re-condenses and creates nanometer sized particles, the meteoric smoke particles (MSP) (Rosinski and Snow, 1961; Hunten et al., 1980). The MSPs are thought to be crucial in creating NLC, where they probably act as condensation sites for water vapor to form the larger icy NLC particles, but homogeneous condensation may also be part of the cause of this (Turco et al., 1982; Rapp and Thomas, 2006). In the growth process the icy NLC particles, growing by water vapor condensing on them, also capture MSP, so that NLC particles will have MSPs embedded in them (Havnes and Naesheim, 2007; Havnes et al., 2009; Hervig et al., 2012, 2017). It also appears that the MSPs, when transported downwards, can influence on the cloud formation in the stratosphere and possibly also the troposphere (Ogurtsov and Raspopov, 2011).

In order to understand the mesosphere it is crucial to understand the evolution and role of various types of dust particles in it, such as the icy NLC and Polar Mesospheric Summer Echoes (PMSE) particles, and MSPs which probably also are present in the winter mesosphere to create the weak radar PMWE (Polar Mesospheric Winter Echoes) clouds (Czechovsky et al., 1979; Zeller et al., 2006; Latteck and Strelnikova, 2015). The progress in ground based instrumentation and observing techniques during the last few decades has been impressive. For example, lidars now routinely observe in full daylight to determine NLC particle sizes and densities (Baumgarten et al., 2007) and they also measure the metallic content in the mesosphere (Huang et al, 2015) and mesospheric temperatures (Höffner and Lautenbach, 2009). The powerful new MST radar MAARSY with its large increase in sensitivity has profoundly changed our knowledge of PMSE occurrence rates and the altitude ranges in which they can be found (Latteck and Strelnikova, 2015). Satellites have identified MSP cloud layers by observing along them (Hervig et al., 2009) and have also confirmed earlier predictions (Havnes and Næsheim, 2007; Havnes et al., 2009; Kassa et al., 2012) that MSPs are embedded in the icy NLC/PMSE particles with from 0.01 to 3% by volume (Hervig, 2012).

One of the obvious advantages of the ground based instrumentation and satellites, is that they can observe the mesospheric clouds continuously. However, they have a limited space resolution (ca. 100 m and upwards) and time resolution (seconds and upwards). Rocket instrumentation, on the other hand, although presenting only a snapshot of the conditions along its trajectory, observe with a time resolution typically of ~ 10^{-3} to 10^{-4} seconds, corresponding to a spatial resolution of ~ 0.1 to 1 m. Various rocket probes are developed to observe the plasma conditions (Friedrich and Rapp, 2009), the dust charge density (Havnes et al., 1996a), the total density of small dust (MSP) by a flashing technique (Rapp and Strelnikova, 2009) while MASS is a coarse dust mass spectrometer (Knappmiller et al., 2008; Amyx et al., 2008; Robertson et al., 2009, 2014). The MUDD (Multiple Dust Detector) mass analyze the collision fragments of the icy NLC particles and relate this to the mass distribution of embedded MSP (Havnes et al., 2014; Antonsen and Havnes, 2015; Antonsen et al., 2017).

In spite of the progress made with rocket instrumentation, there is a lack of high time/space resolution instruments to measure parameters as dust size, number density and charge. In the present paper we consider the principles of the much used DUSTY impact probe (Havnes et al., 1996a) and how its performance can be improved. The DUSTY probe, the principle of which is shown in Fig.1, is equipped with grids to prevent ambient plasma from reaching G2 and the bottom plate BP but allow dust particles to enter and collide with the grids and the BP. The potentials of the grids are given in Fig.1. The observed currents to the probe were originally used to find only the dust charge density of the ambient dust cloud, but in the present paper we will show how to extend the analysis of the DUSTY probe currents to allow it to also determine

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Discussion started: 3 September 2018

other dust parameters. The extension of the original method of analysis is based on earlier works, which have demonstrated the importance of secondary charge and secondary current production in glancing dust impacts on rocket probes and payload bodies (Havnes and Næsheim, 2007; Havnes et al., 2009; Kassa et al., 2012).

In Sec.2 we extend the earlier analysis method for the DUSTY impact probe and now use the currents to G2 and BP to find not only the dust charge density as before, but also the total dust density, the dust radius and the mean dust charge. In Sec. 3 we show the values for dust density and dust radius by this new method, used on the observations by the DUSTY probe on the payload MXD-1, which was launched on 30.06.2016 at 09:43:18 UT in the MAXIDUSTY rocket campaign. In Sec.4 we compare the DUSTY results with those from the RMR Lidar at Andøya (von Cossart et al, 1999; von Zahn et al, 2000; Baumgarten et al, 2007) and the on board MISU photometer (Gumbel et al., 2001; Hedin et al., 2008; Megner et al., 2009) and conclude the paper in Sec.5.

2 The extended analysis of dust observations made with DUSTY type Faraday cup probes.

The DUSTY probe (Havnes et al., 1996a; Havnes and Næsheim, 2007), the design of which is shown in Fig. 1, has grids G0, G1 and G2 and a solid bottom impact plate BP. The probe must point forward along the payload axis. The dust impact currents to G1, G2 and BP are all registered but not the current to G0, which is at the payload potential Φ_P. The registered currents are I_{G1}, I_{G2} and I_{BP}. The current I_{G1} will not be used in the analysis. It is the grid which is most influenced by effects like payload charging and the plasma environment and as such not directly connected to the measurements of dust. G0 and G1 are made of thin cylindrical wires and they each cover only 4.6% of the opening cross section of DUSTY. G2 is made of thicker wires to increase the secondary charging effect. It covers 23.5 % of the DUSTY cross section.

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Discussion started: 3 September 2018

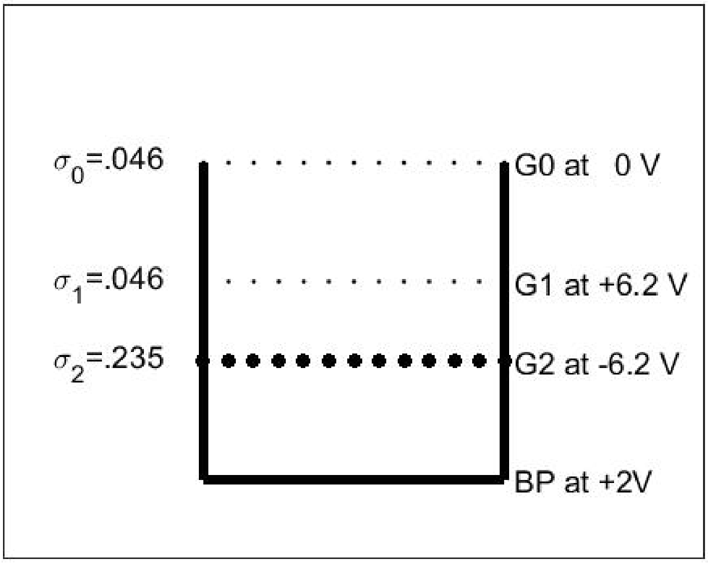

Figure 1. The design of the DUSTY probe used in the MAXIDUSTY campaign. The fractional coverage of the different grids, relative to the total probe cross section, are $\sigma_0 = \sigma_1 = 0.046$ and $\sigma_2 = 0.235$. The electric potentials of all the grids and the bottom plate are relative to the payload potential Φ_P. The currents are measured on G1, G2 and BP, but not on G0.

The dust current into the probe in front of G2, is designated I_D and is part of the expressions for the total current I_{G2} measured on G2

$$I_{G2} = \sigma_2 I_D + I_S \quad (1)$$

and for I_{BP} measured on the BP.

$$I_{BP} = (1 - \sigma_2) I_D - I_S \quad (2)$$

The current to G2 is made up of $\sigma_2 I_D$ which is the part of I_D which hits G2 and deposits its charge, plus the secondary current I_S which is produced by glancing dust impacts on G2 which rubs off electrons from it. If this last process is effective it can lead to that the total current I_{G2} can become positive even if the impacting dust particles are charged negatively. The current I_{BP}

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Discussion started: 3 September 2018

Atmospheric Measurement Techniques
Discussions
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to the bottom plate is made up of the direct hits on to BP by the dust which was not hitting G2, and minus the secondary current I_S. The electrons which are rubbed off from G2, producing a positive current I_S to G2, will be deposited on BP and create a negative current $-I_S$ there. We can eliminate I_S to find I_D by

$$I_D = I_{G2} + I_{BP} \quad . \tag{3}$$

The two upper grids G0 and G1 are made of thin wires and each cover only 4.6 % of the DUSTY cross section (Fig. 1). Much of the small negatively charged fragments produced on them by will be stopped by air friction and probe internal electric fields (Antonsen et al., 2017). We therefore neglected a possible contribution of their secondary production to the currents to G2 and BP. However, they will together stop ~9.2 % of the incoming dust current from passing G0 and G1. The current I_{Total} into the probe just above G0 can be expressed as $I_{Total} = I_D (1-\sigma_0)^{-2} = 1.1$ x I_D which gives us directly the observed ambient dust charge density $\Sigma\,(N_Z Z_D)$ from the relationship

$$I_{Total} = \pi R_p^2 V_R e \sum(N_Z Z_D) \quad . \tag{4}$$

Here R_p is the probe radius, and e = 1.6 x 10^{-19} C. The number density of dust particles with charge number Z_D is N_Z and the rocket velocity is V_R. We should note that the dust charge density $\sum(N_Z Z_D)$ which can be extracted from Eq. (4) is independent of the model for secondary production of charge since this cancels in Eq. (3).

Some information on the expected size of the dust particles, and the role of secondary charge production, can be found from examining the ratio

$$R = \frac{I_{G2}}{I_{BP}} = \frac{\sigma_{G2} I_D + I_S}{(1-\sigma_{G2}) I_D - I_S} \quad . \tag{5}$$

This ratio R should have values between $R = \frac{\sigma_{G2}}{1-\sigma_{G2}} = 0.31$ when the secondary charging current $I_S \rightarrow 0$, and R= -1 for $I_S >> I_D$. In Fig.2 we show R and I_D as function of altitude. It is reassuring that R, even though it varies significantly with altitude, stays so well within the above limits. This has been shown to be the case also in several earlier launches of the DUSTY probe (Havnes and Næsheim, 2007; Havnes et al., 2009).

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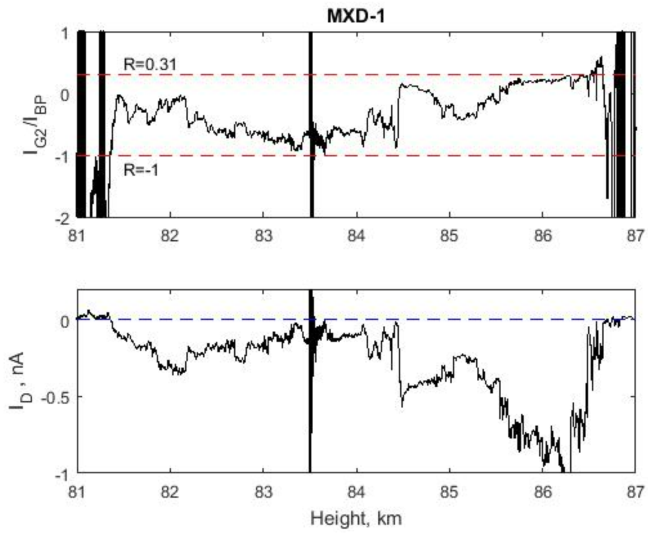

Figure 2. The ratio of the currents to G2 and BP in the upper panel, compared to the current I_D in the lower panel. The large disturbance at ~ 83.5 km altitude is caused by a squib being fired to open for another experiment on the payload. The values of *R*, at and outside the borders of the cloud are to be neglected since the dust density there is low or zero and *R* is therefore dominated by noise and uncertainties in their background level.

We see from Fig. 2 that the ratio *R* is dominated by secondary charging effects in the middle of the cloud system at ~82.5 to ~84.4 km, while at the upper edge around 86 km secondary charging is not very significant. This is in accordance with a scenario where small cloud particles normally can be expected to be found in the upper parts of the clouds (Robertson et al, 2009), from where they sink and grow, to reach maximum sizes in the middle regions of the clouds. In the lower parts, melting should lead to a reduction of dust sizes and release of embedded MSPs. Laboratory experiments show that the secondary production for fast impacts on metals by iron particles of radius above ~100 nm, is proportional to the volume of the impacting particle (Friichtenicht, 1964; Adams and Smith, 1971). Impacts of small ice particles below a radius ~100 nm, at impact velocities ~ 1400 m/s, indicate that the secondary production is proportional to the cross section of the impacting ice particle (Tomsic, 2001). Since the charge on a dust particle at given plasma conditions is roughly proportional to its radius, and since the cross section is proportional to the square of the radius, a significant secondary current (R<0) indicates large particles, while small secondary production (R>0) indicate small dust particles. We will later show that this is what we get for the dust size from the extended method.

The secondary charging, or the rubbing off effect by impacting dust on surfaces, is strongly dependent on the impact angle θ_i, the angle between the surface normal and the direction to the impacting particle. In the experiments with ice particles (Tomsic 2001) the maximum of the secondary production was at θ_i~ 86 degrees and it was reduced to 0 at 90 deg. Little secondary charge production took place below θ_i ~ 70 deg. This means that of the dust particles impacting on the cylindrical grid wires, only a fraction will rub off electrons from the grid. Havnes and Næsheim (2007) analyzed in detail the rotational effect on the currents to the grids of two DUSTY probes, launched in the summer of 1994 (Havnes et al., 1996a). They found that a substantial secondary charge production was needed to model the payload rotational effects on the grid impact currents. The effect of secondary charging has since been mapped in several other rocket flights (Havnes et al., 2009; Kassa et al., 2012; Havnes et al., 2014; Antonsen and Havnes, 2015; Antonsen et al., 2017). One result of the analysis of the secondary impact effects of NLC particles on the main grids of DUSTY type probes, was that it had to be very much more efficient than what has been found for impact of ice particles in laboratory experiments. A probable reason for this difference is most likely connected to that pure laboratory ice particles below ca 7 nm, have a tendency to stick to the impact surface and evaporate (Tomsic, 2001). On the other hand the NLC/PMSE icy particles, containing a substantial number of embedded MSPs (Hervig et al., 2012; Havnes and Næsheim, 2007) will partly fragment on impact and MSPs which are released will not evaporate but survive to carry away "rubbed off" electrons. With a MSP volume filling factor of 3% in a NLC/PMSE particle (Hervig et al., 2012), even a 7 nm NLC/PMSE icy particle will contain some 10 to 30 MSPs if their sizes are in the range 0.7 to 1 nm.

The secondary production, the number of charged fragments produced by one impacting NLC/PMSE particle of radius r_d, varies with the cross section of the impacting particle as

$$\eta_S(r_d) = \eta_{S,ref}\,(r_d/r_{d,ref})^2 \quad . \tag{6}$$

Havnes and Næsheim (2007) found that for a reference icy dust particle, of radius $r_{d,ref}$ = 50 nm a number of $\eta_{S,ref}$= 50 to 100 negative unit charges would be released. With 3% MSP volume filling factor (Hervig et al., 2012) this corresponds to that ~1% of the embedded MSPs become charged fragments, if we set the embedded MSP radius to 1 nm.

We can now express the secondary current I_S by a use of Eq. (6) and a knowledge of how large fraction of the grid wires which contribute to the secondary charge production. In the modeling by Havnes and Næsheim (2007) they found that secondary charges are produced on a fraction

Atmos. Meas. Tech. Discuss., https://doi.org/10.5194/amt-2018-222
Manuscript under review for journal Atmos. Meas. Tech.
Discussion started: 3 September 2018

$\sigma_{2,sec}$~0.28 of the G2 grid diameter, where the total area of G2 in MXD-1 covers a fraction σ_2 = 0.235 of the total probe cross section $\sigma_P = \pi R_P^2$. The probe radius is R_P = 0.04 m. From this we can express the secondary charge current as

$$I_S = eN_D V_R A_{sec} \eta_S(r_d) \quad . \qquad (7)$$

Here $N_D = \Sigma N_Z$, the total dust number density and $A_{sec} = \sigma_{2,sec}\sigma_2\sigma_p$ is the effective area of the probe for secondary charge production. This is only ~ 7% of the total probe cross section σ_p . The observed secondary charge current I_S is also found from Eqs. 1 and 2 as

$$I_S = (1-\sigma_2)I_{G2} - \sigma_2 I_{BP} \quad . \qquad (8)$$

Inserting Eq. (6) in Eq. (7) we can solve Eqs. (7) and (8) for the dust radius

$$\left(\frac{r_d}{r_{d,ref}}\right)^2 = \frac{(1-\sigma_2)I_{G2} - \sigma_2 I_{BP}}{A_{sec}\eta_{S,ref}\, e N_D V_R} \quad . \qquad (9)$$

Fixing the values for $\eta_{S,ref}$ and $r_{d,ref}$, the only unknown parameter on the right hand side is the total dust density N_D. If this is also known, we can find the dust radius from Eq. (9). However, the value of N_D is not directly available, but can be found in an iteration process which includes a charging model for the dust.

The charging model computes the equilibrium charge distribution of the ambient dust particles. The electron density n_e (Fig. 9) is measured by various probes on the payload. We require charge neutrality and find the ion density n_i from

$$n_i - n_e + \sum N_Z Z_D = 0 \quad . \qquad (10)$$

The plasma temperature is equal to the neutral temperature and we will use a temperature of 150 K. For our equilibrium charging model we require that the rate at which dust particles of charge Z are given the charge number $(Z\text{-}1)$ by an electron colliding with it and sticking to it, is equal to the rate by which dust with charge number $(Z\text{-}1)$ are given charge number Z by ions colliding and sticking to it

$$N_Z J_e(Z) = N_{Z-1} J_i(Z-1) \quad . \qquad (11)$$

Here $J_e(Z)$ and $J_i(Z)$ are the rates at which charged particles (electrons or ions) arrive at the surface of a dust particle with charge number Z, and stick to it. We have used the expressions for J_e and J_i from Draine and Sutin (1987) which include the short range polarization forces and refer to that paper for the full expressions.

Atmos. Meas. Tech. Discuss., https://doi.org/10.5194/amt-2018-222
Manuscript under review for journal Atmos. Meas. Tech.
Discussion started: 3 September 2018

The iteration procedure to extract values for dust radius r_d , dust total density N_D and also the dust charge distribution N_Z , together with other relevant parameters dependent on r_d and N_D, starts with a guess for the average dust charge number Z_{av}. A good guess is normally Z_{av}= -1. This will give an initial value for the total dust number density $N_D = \sum(N_Z Z_D)/Z_{av}$. Here $\sum(N_Z Z_D)$ is the observed dust charge density found from Eq. (4). From this value of N_D we calculate a value for the dust radius from Eq. (9). These approximations to N_D and r_d are now used in the charging model, together with known values for the plasma parameters, to calculate a new total dust density and a new average dust charge number which is used to find a new value for r_d . This process is repeatedly run through the charging code until it converges to a solution.

3 Measurements by the DUSTY probe on MAXIDUSTY-1, analyzed with the extended method.

We now use the observations by the DUSTY probe on MXD-1 and the new extended method to find the basic dust parameters: radius r_d, total density N_D and average dust charge number Z_{av} throughout the observed NLC/PMSE clouds. The electron data are taken from the results by the on board Faraday instrument (Friedrich and Rapp, 2009). In Fig. 3 we show smoothed raw currents I_{G2} and I_{BP} and the adopted background which will be subtracted from the raw currents to give the net currents . The curves show that the main cloud system extends from

Atmos. Meas. Tech. Discuss., https://doi.org/10.5194/amt-2018-222
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Discussion started: 3 September 2018

Atmospheric Measurement Techniques
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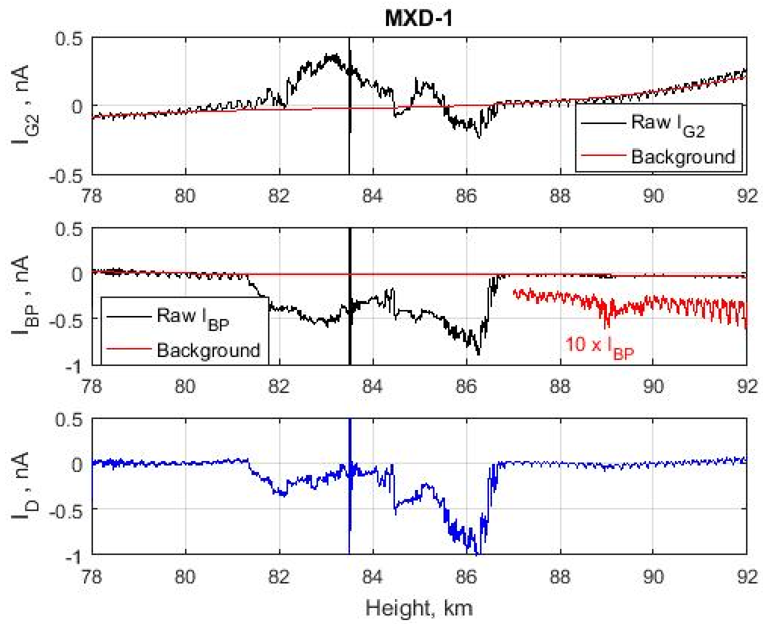

Figure 3. The smoothed currents I_{G2} and I_{BP} and the assumed background currents, are shown in the upper two panels. In the bottom panel we show the I_D current based on the currents I_{G2} and I_{BP} , corrected for background. The "event" at ~83.5 km is due to a squib being fired to open another instrument on the payload. In panel 2 we have also plotted in red a current *10 x* I_{BP} to empasize that there is a clear but weak dust structure at least spanning the altitude region from ~88.5 to ~89.9 km.

~81.3 to ~86.8 km with a clear but weak additional dust cloud structure between ~88.5 to ~89.9 km. We see indications that a weak structure also extends below 81.3 km, possibly down to ~ 80 km. This is apparent mainly in panel 1 where there is a weak I_{G2} in this interval and the payload rotation effect is different above and below 80 km, possibly indicating the presence of small MSP's in the size range up to several nm. They may have been released by melting of the larger icy particles and may be affected by the airstream around the payload and by the payload rotation.

In Fig. 4 we show the inferred values for dust radius r_d and N_D. The large noise signals around ~83.5 km in Figs .2 and 3, which were caused by a squib being fired, have been removed. The other 4 narrow and strong features in the middle of the cloud region (~83.3 to ~84.5 km) indicate the presence of dust layers, or "dust voids" with much larger dust sizes than just outside these

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Discussion started: 3 September 2018

Atmospheric Measurement Techniques
Open Access
Discussions

layers. The presence of dust of radius up to and even above 100 nm within the layers is indicated, compared to

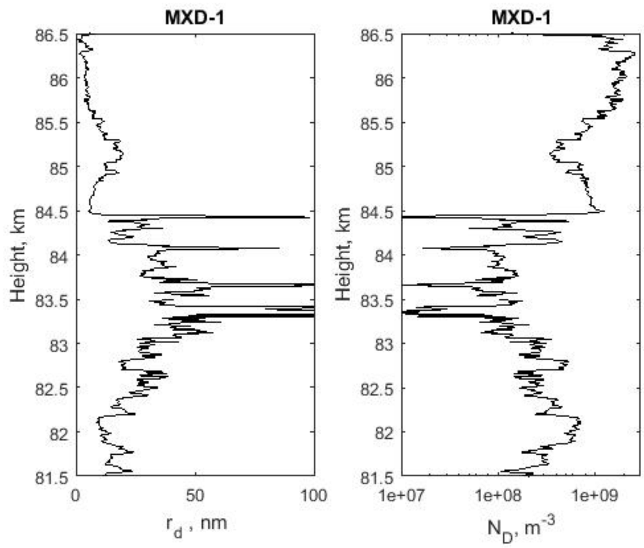

Figure 4. The inferred dust radius r_d and dust density N_D within the main cloud. We have applied a moderate sliding mean smoothing over 100 data points, changing the altitude resolution from 0.1 m in the observed data points, to 10 m. We have also removed the signals in the altitude region 83.5 to 83.55 km which are dominated by the strong noise from the squib firing, shown in Figs. 2 and 3.

dust sizes just outside the layers ranging from ~10 to ~ 40 nm. However, the values for r_d in these 4 narrow layers with large dust, are probably considerably more uncertain than in most other parts of the NLC/PMSE cloud. The reason for this is that these 4 layers (voids) have a very low dust density N_D, much lower than in the regions just outside the layers. We can see this from Figs. 2 and 3 where the current I_D is very low within the 4 layers and therefore the dust density N_D will also be low. This is directly evident from Fig .4, which show both r_d and N_D. The narrow layers with the large increase in dust sizes r_d also have low dust densities, where N_D can be down to ~ 10^7 m^{-3}. At such low values for the dust density, the dust radius r_d computed by Eq. (9), can be much affected by noise fluctuations in the signals, by payload rotational effects and uncertainties in the assumed background currents. This will lead to relatively large uncertainties in N_D and therefore also in r_d when computed with Eq. (9). The narrow layers or voids in NLC/PMSE clouds will probably still exist (see also Havnes et al., 1996b) and contain large dust particles but their peak values may be questionable.

4 Comparison of the extended DUSTY method results with lidar and photometer results.

As a test on the values of r_d and N_D found by the extended method we compare with corresponding values found from the ALOMAR RMR Lidar observations (von Zahn et al., 2000, Baumgarten et al., 2007) and the on board MISU photometer (Gumbel et al., 2001; Hedin et al., 2008; Megner et al., 2009).

The ALOMAR RMR Lidar is a twin-Lidar system with two power lasers simultaneously emitting at 1064, 532 and 355 nm wavelengths, and with two receiving telescopes each with a 1.8 m primary mirror. The Lidar can be operated all year and under daylight conditions. During the MAXDUSTY-1 launch one beam was pointed along the predicted payload trajectory at 85 km and one in the vertical direction. In Fig. 5 we show the RMR observations close to the payload trajectory where the separation of the lidar and rocket measurements was less than 2 km. The second lidar performed measurements above the lidar station about 18 km separated from MXD-1 measurements. At both locations a double layer was observed and both layers show up and downward motion indicating small scale perturbations of the atmosphere. The size of the particles is calculated from the signal of three wavelengths assuming a distribution of needle and plate like particles of multiple sizes (Baumgarten et al., 2007). The size values given here are radii of a volume equivalent sphere, and give the mode of a Gaussian distribution of particle sizes.

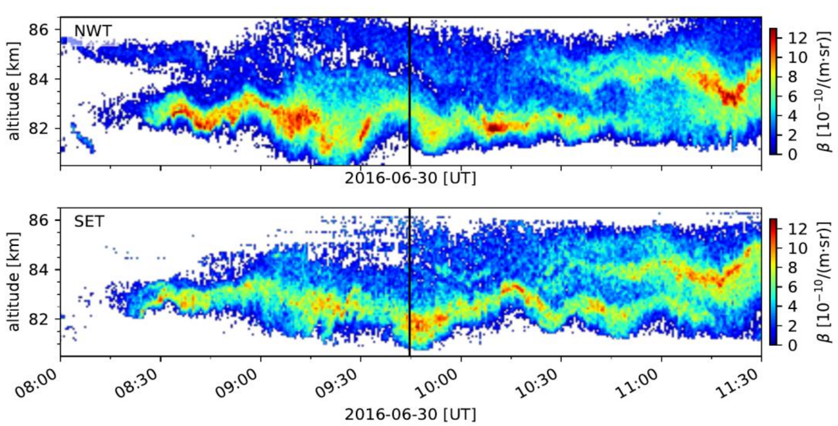

Figure 5. Backscatter coefficient (532 nm) measured by the RMR-Lidar along the payload trajectory of MXD-1 (upper panel) and about 18 km to the south-east of the trajectory (lower panel). The time of the rocket penetrating through the NLC layer is marked by the vertical black line.

The Side-looking MISU NLC photometer on board the payload also detected a two-layer NLC with an altitude profile very similar to the one in Fig. 5 at the time of the rocket measurement. Comparing the angle dependence of the scattering of sunlight on the NLC particles to theoretical Mie scattering phase functions, one can find an effective optical scattering radius, r_{Eff} of the particles in the NLC. This method is biased towards the largest particles due to the very strong dependence of scattering on dust radius. Below the layer, measuring the entire vertical extent of the NLC, the effective radius r_{Eff} = 46 (±4) nm. As we ascend through the NLC, the retrieved particle radius decreases with increasing altitude and the effective optical scattering radius in the top layer is 40 (±8) nm.

The two extended layers in Fig. 5, centered on ~ 83 and ~ 85 km also coincide with two layers at the same altitudes at which layers were detected with DUSTY. For DUSTY each of the two layers are characterized by containing large dust particles of low number density. This demonstrates again the strong dependence of scattering of light on the dust radius, increasing very rapidly with size so layers of low density but containing large dust can dominate the scattering. .

In Fig. 6 we show the DUSTY results, for one set of secondary charging parameters, for dust radius r_d, total dust number density N_D, and average dust charge number Z_{av}. We also show RMR Lidar results for 5 minutes centered on the MXD-1 measurements (09:44:36 UT) as well as the photometer measurements. The average sizes of the lidar measurements through the layer is 22 nm with standard deviation of 5 nm. The average width of the Gaussian size distribution is 8 nm. In the last panel we show the RMR Lidar observations of NLC brightness for 30 seconds around 09:44:36 UT compared with two model Lidar profiles computed for dust parameters inferred from the DUSTY observations and for the assumptions that the particles are pure ice or ice contaminated with 5% FeO which is the upper limit used by Hervig et al. (2012). We calculated the refractive index for mixture with FeO using the effective medium approximation (Garnett, 1904). We have excluded the data in the altitude region ~83.5 to ~83.7 km which were affected by the squib event.

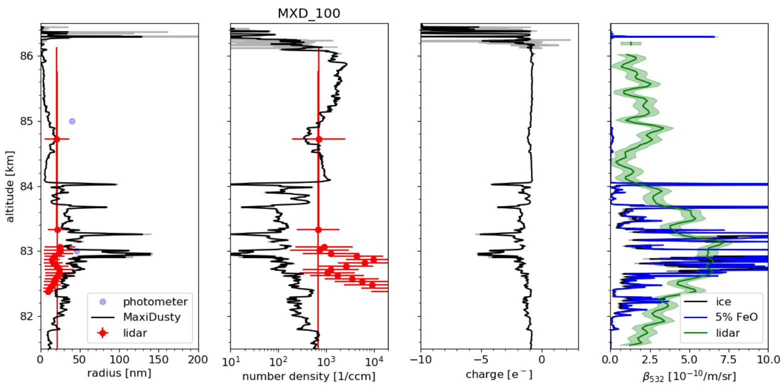

Figure 6 The first three panels show results for r_d , N_D and Z_{av} for an assumed value of $\eta_{S,ref}$ = 100. RMR Lidar results are marked by red dots while the two blue dots at 83 and 85 km are for the MISU photometer. The last panel shows the observed Lidar altitude profile where the black curve shows model results computed based on the MAXIDUSTY data of panel 1 and 2 and the assumption of pure ice particles, and the blue curve shows results based on the assumption that the ice particles contain 5% FeO. The green shaded area indicates the measurement uncertainty.

The variations of the DUSTY results for r_d, N_D and Z_{av} seem qualitatively reasonable. At the top of the cloud we find the smallest dust particles with sizes r_d well below 10 nm. These dust particles have presumably been created recently and now grow by deposition of water vapor which freezes out on their surface and contain embedded MSPs which become attached to them (Havnes and Næsheim, 2007; Hervig et al., 2012). The highest dust number density, close to $2x10^9$ m^{-3}, is found in this region. In the middle of the cloud the dust sizes outside the narrow dust voids have increased to a maximum value of around 40 nm and number density is around 10^8 m^{-3}. The dust radius becomes smaller further down into the bottom parts of the cloud with values of around ~ 20 nm and the number density increases to ~ $6x10^8$ m^{-3}. The average dust charge number is close to Z_{av} = -1 in the lower and upper parts of the cloud while in the middle part it is around Z_{av} ~ -2 to -3. That the comparatively large grains in the middle part do not have larger negative charge numbers is due to a paucity of electrons which is demonstrated by the electron bite out from ~ 82 to 84 km, shown in Fig. 7. In this figure we also show the dust charge density $\sum(N_Z Z_D)$ and note that the dust particles are the dominant negative charge carriers in practically the whole extent of the cloud.

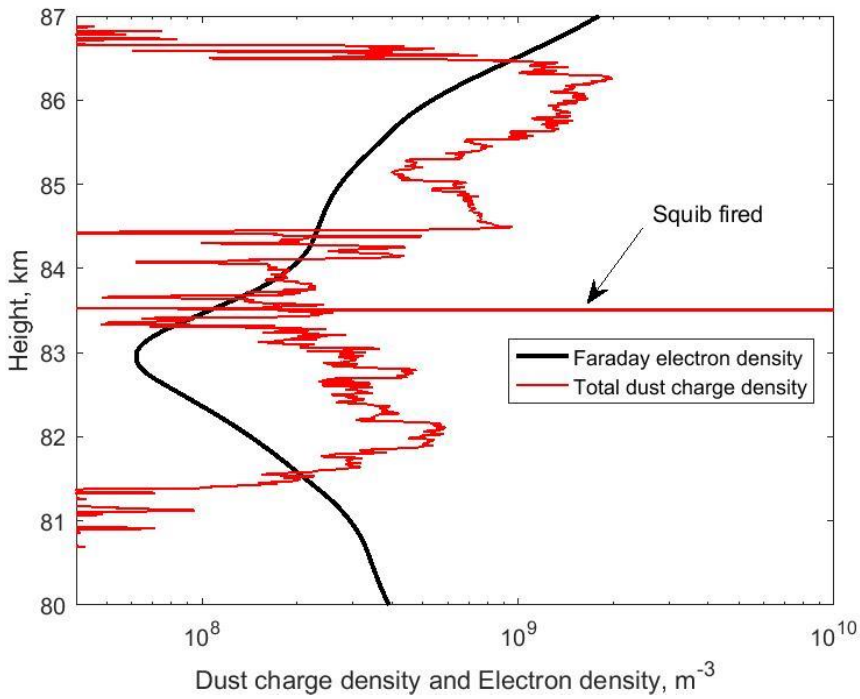

Figure 7. Electron density measured with the Faraday instrument, and the total dust charge density as observed by DUSTY, on MXD-1.

5 Discussion and conclusion. The extended method with its unsurpassed altitude resolution gives, in our opinion, reasonable results which compare well with the RMR Lidar and MISU photometer results (Fig. 6). It is noteworthy that the parameters for the secondary charging model in the present work have been taken from earlier modeling not aimed at finding r_d, N_D and Z_{av} but to demonstrate that secondary charging was essential in reproducing the currents to

Atmos. Meas. Tech. Discuss., https://doi.org/10.5194/amt-2018-222
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Discussion started: 3 September 2018

BP and G2 and their variation with payload rotation (Havnes and Næsheim, 2007; Havnes et al., 2009; Kassa et al., 2012).

If we compare the various results in Fig.6, where DUSTY results are based on $\eta_{S,ref} = 100$, there are some significant differences between DUSTY results and the RMR Lidar or MISU photometer results. The first is that the RMR Lidar in the region at and slightly below 83 km, finds particles of half or less the sizes that DUSTY finds. The MISU photometer is closer to the DUSTY values. Also, the Lidar total dust densities in the same altitude region are in general more than a magnitude larger than what DUSTY finds.

We should bear in mind that some of the differences may result from the Lidar and DUSTY probe sampling very different volumes. The sounding volumes are separated horizontally by about 2 km and differ in size. With an altitude resolution of 475 m and integration time of 300 sec the Lidar samples a volume of about 10^5 m^3 while DUSTY, with some smoothing of the data, samples 0.5 m^3 ($5x10^{-4}$ m^3 with unsmoothed data). These differences may be important taking into account small scale dynamics (Baumgarten and Fritts, 2014; Fritts et al., 2017). The time evolution shown in Fig. 5 indicates that such small scale variations were indeed likely during the time of the measurement.

For DUSTY we could lower the computed r_d and increase the N_D by increasing the secondary efficiency $\eta_{S,ref}$ in Eq. (9) from its “accepted” values between 50 and 100. This may require that the embedded MSPs occupy an exceptionally large volume of the icy NLC/PMSE particles. However, we see from Fig. 6d that the Lidar profile, computed on the basis of the DUSTY results for a $\eta_{S,ref} = 100$ compares reasonably with the observed Lidar profile while an increase of $\eta_{S,ref}$ to 150 will lead to the computed DUSTY Lidar profile becoming very weak compared the observed one. The best fit of the model DUSTY Lidar profile to the observed results is obtained for a value of $\eta_{S,ref}$ around 70 to 80.

The values of r_d, N_D and Z_{av} from the DUSTY data will also be affected by the electron density within the dust cloud. This can be critical if the dust density is large enough to create an electron bite-out with locally large reductions in the electron density. In such cases the dust charges can be reduced significantly compared to those that would occur if no bite-out were present. We see in Fig. 7 a significant electron bite-out with a minimum electron density of $6x10^7$ m^{-3} at an altitude of 83 km. At such low electron densities the Faraday method to determine the electron density is quite uncertain, which motivates us to examine the consequences of reducing the true electron density within the bite-out compared to that in Fig. 7. Reducing it by a factor of 10 will lead to a reduction of r_d by a factor of ~2 and an increase in N_D by a factor of ~3 within the bite-out.

The charge model we have used does not include the photodetachment effect (Havnes and Kassa, 2009; Rapp, 2009) and it does not include any photoelectric effect. Inclusion of a photodetachment effect will have some – but not serious - effect on dust particles less than ~5 nm. It will lead to a moderate increase in dust density and a decrease of the dust radius. In our model, using values of the photodetachment effect taken from Havnes and Kassa (2009), we get a moderate reduction of the dust radius r_d in the altitude region above ~ 85.5 km.

Another uncertainty, caused by the design of the DUSTY probe, is that small dust particles (less than ~ 2 nm at an altitude ~ 85 km), which may be carrying a non-negligible part of the charge density, will be swept away from the probe by the airstream around the payload and its probes (Horányi et al., 1999; Hedin et al., 2007). Observations by the MASS instrument (Robertson et al., 2009, 2014; Knappmiller, 2008) indicate that considerable amounts of small charged dust particles have a tendency to be present in the upper layers of NLC/PMSE clouds, together with larger NLC/PMSE cloud particles. We cannot exclude that this is also the case for the clouds observed by MXD-1. To evaluate the consequences of small charged particles potentially not being registered by DUSTY we will need a charging model with more than one dust size. Such models should also improve the comparison to lidar measurements, as these take the effect of different sizes into account and show that the ensemble of particles often has a width of the size distribution of about half the mode radius (Baumgarten et al., 2010).

We find that the development of the new extended method to analyze the DUSTY measurements, has given this probe a power which is astounding considering its simplicity. It can in principle be used to measure the dust radius, dust total density, dust charge density and dust charge – all with an unsurpassed altitude resolution down to 10 cm or smaller scales. This will also open up for a mapping of the distribution of dust size, dust density and dust charges within small scale dust structures (Havnes et al., 1996b). To achieve the best foundation for the extended method and future use of DUSTY-like probes, we plan to refine the analysis with a more complete charging model and to map the effects of changes in the various parameters involved in the method. A comparison with the RMR lidar and MISU photometer observations during the MXD-1 flight will continue to be essential in refining the method. This may also lead to a fine-tuning of the construction of the DUSTY probe for which the basic structure should be retained though modifications of G2 might be advantageous. For future campaigns we intend to improve the collocation of the measurement volumes and use the high resolution DUSTY measurements to derive the actual size distribution within the lidar sounding volume.

Author contribution.

OH, AB, TA and TWH extended the theory for analyzing the rocket data. OH and TA analyzed the rocket data. GB collected and analyzed the Lidar data. TA and ÅF tested rocket instruments. MF analyzed the Faraday data and provided the electron density data. JH collected the photometer data and analyzed them. OH prepared the manuscript with contributions from all co-authors.

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The authors declare that they have no conflict of interest.

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