

# Ultrasonic Nebulization for the Elemental Analysis of Microgram-Level Samples with Offline Aerosol Mass Spectrometry

Rachel E. O'Brien<sup>1\*</sup>, Kelsey J. Ridley<sup>2</sup>, Manjula R. Canagaratna<sup>3</sup>, John T. Jayne<sup>3</sup>, Philip L. Croteau<sup>3</sup>, Douglas R. Worsnop<sup>3</sup>, Sri Hapsari Budisulistiorini<sup>4†</sup>, Jason D. Surratt<sup>4</sup>, Christopher L. Follett<sup>5</sup>, Daniel J. Repeta<sup>5</sup>, Jesse H. Kroll<sup>2</sup>

[1] Department of Chemistry, College of William and Mary, Williamsburg, Virginia, 23185, USA

[2] Department of Civil and Environmental Engineering, Massachusetts Institute of Technology, Cambridge, Massachusetts 02139, USA

[3] Center for Aerosol and Cloud Chemistry, Aerodyne Research Inc., Billerica, Massachusetts 01821, USA

[4] Department of Environmental Sciences and Engineering, Gillings School of Global Public Health, University of North Carolina at Chapel Hill, Chapel Hill, North Carolina 27599, USA

[5] Department of Marine Chemistry and Geochemistry, Woods Hole Oceanographic Institution, Woods Hole, Massachusetts 02540, USA

† Now at: Earth Observatory of Singapore, Nanyang Technological University, Singapore 638789, Singapore

Corresponding author: Rachel O'Brien, reobrien@wm.edu

**Abstract.** The elemental composition of organic material in environmental samples – including atmospheric organic aerosol, dissolved organic matter, and other complex mixtures – provides insights into their sources and environmental processing. However, standard analytical techniques for measuring elemental ratios typically require large sample sizes (milligrams of material or more). Here we characterize a method for measuring elemental ratios in environmental samples, requiring only micrograms of material, using a Small Volume Nebulizer (SVN). The technique uses ultrasonic nebulization of samples to generate aerosol particles (100-300 nm diameter), which are then analyzed using an aerosol mass spectrometer (AMS). We demonstrate that the technique generates aerosol from complex organic mixtures with minimal changes to the elemental composition of the organic material and that quantification is possible using internal standards (e.g.,  $\text{NH}_4^{15}\text{NO}_3$ ). Sample volumes of 2-4  $\mu\text{L}$  with total solution concentrations of at least 0.2 g/L form sufficient particle mass for elemental ratio measurement by the AMS, despite only a small fraction (~0.1%) of the sample forming fine particles after nebulization (with the remainder ending up as larger droplets). The method was applied to aerosol filter extracts from the field and laboratory, as well as to the polysaccharide fraction of dissolved organic matter (DOM) from the North Pacific Ocean. In the case of aerosol particles, the mass spectra and elemental ratios from the SVN-AMS agree with those from online AMS sampling; similarly, for DOM, the elemental ratios determined from the SVN-AMS agree with those determined using combustion analysis. The SVN-AMS provides a platform for the rapid quantitative analysis of the elemental composition of complex organic mixtures and non-

38 refractory inorganic salts from microgram samples with applications that include analysis of aerosol extracts, and  
39 terrestrial, aquatic, and atmospheric dissolved organic matter.

## 40 **1 Introduction**

41 A large number of environmental systems, including the atmosphere, natural waters, and terrestrial systems,  
42 contain complex organic mixtures composed of hundreds to thousands of molecular species. Our ability to understand  
43 and model such complex chemical systems is often greatly improved when we characterize them in terms of simple  
44 chemical frameworks. On the simplest level, the analysis of average elemental ratios can provide important  
45 information on potential sources of organic matter samples, as well as the chemical and/or biological transformation  
46 processes that modify their composition. For example, the elemental ratios of atmospheric organic aerosol – e.g.,  
47 oxygen/carbon ratio (O:C), hydrogen/carbon ratio (H:C), and nitrogen/carbon ratio (N:C) – provide information on  
48 aerosol sources and aging (Aiken et al., 2008; Canagaratna et al., 2015; Chen et al., 2015; Daumit et al., 2013; Heald  
49 et al., 2010; Jimenez et al., 2009; Kroll et al., 2011). Similarly, in water and soil samples, the elemental ratios of  
50 carbon, nitrogen, and phosphorous reveal insights into sources and processing of dissolved and particulate organic  
51 matter (Becker et al., 2014; Hansman et al., 2015; Koch et al., 2005; Lu et al., 2015).

52 The most widespread technique for elemental analysis is high-temperature combustion followed by elemental  
53 (carbon, hydrogen, nitrogen, and sulfur (CHNS)) analysis, which is highly accurate but can require milligrams of  
54 material (Skoog et al., 1998). For many trace environmental samples, like atmospheric aerosol, this can require  
55 extremely long collection times which lead to low time resolution, limiting the amount of information provided for  
56 systems that exhibit high temporal variability, such as air masses in major urban regions. An alternative approach for  
57 measuring the elemental ratios of aerosol is online (real-time) techniques. The most widely-used instrument for such  
58 measurements is the Aerodyne High-Resolution Time-of-Flight Aerosol Mass Spectrometer (HR-ToF-AMS) (Decarlo  
59 et al., 2006), which can measure elemental ratios of ambient aerosol using just nanograms of material. Over the last  
60 decade, in-situ analysis of aerosol particles with the AMS has enabled rapid, sensitive characterization of aerosol  
61 concentrations, sources, and atmospheric aging, improving our ability to model atmospheric aerosol and consequently  
62 its climate and health effects (Kroll et al. 2015; Ng et al. 2011; Jimenez et al. 2009; Canagaratna et al. 2007).

63 Recently, a number of researchers have used the AMS in an “offline mode,” in which atmospheric samples  
64 are collected on filters, extracted, and then atomized into the AMS. Examples include the analysis of sources and  
65 aging of atmospheric organic material from aerosol filter extracts (Bozzetti et al., 2017; Huang et al., 2014; Sun et al.,  
66 2011; Xu et al., 2015; Ye et al., 2017), cloud/fog water samples (Kaul et al., 2014; Lee et al., 2012), and organic  
67 material in glaciers (Xu et al., 2013). Offline AMS has proven especially useful for the analysis of aerosol particles  
68 larger than 1  $\mu\text{m}$  (Bozzetti et al., 2016; Daellenbach et al., 2016; Ge et al., 2017). Offline AMS has also proven useful  
69 in investigating fractionation and solubility of atmospheric organic material in water and organic solvents  
70 (Daellenbach et al., 2016; Mihara and Mochida, 2011; Xu et al., 2016). These studies used both custom-made and  
71 commercial atomizers with solvent volumes of at least 5-15 mL. To generate aerosol particles in the size range needed  
72 for the AMS, this corresponds to necessary sample masses on the order of 50  $\mu\text{g}$ . While this represents a substantial  
73 improvement over the sample mass requirements of conventional CHNS analysis, it is still sufficiently large to limit

74 the applicability of the approach since it can require relatively large organic samples collected with high-volume  
75 samplers, often over 24 hours or more.

76 In this work, we characterize a new technique for the elemental analysis of very small sample masses, using  
77 ultrasonic nebulization. Aerosol generation with a small volume nebulizer (SVN) expands the range of environmental  
78 samples that can be measured, where either sample size is limited or solvent contamination is a concern. The SVN  
79 generates aerosol suitable for analysis with aerosol instrumentation, including not only the AMS but also Scanning  
80 Mobility Particle Sizers (SMPS); single particle mass spectrometers (e.g. Particle Analysis by Laser Mass  
81 Spectrometry (PALMS) (Murphy et al., 1998)); soft ionization sources (e.g. Extractive Electrospray Ionization (EESI)  
82 (Gallimore and Kalberer, 2013)); and thermal desorption chemical ionization mass spectrometers (e.g. Filter Inlet for  
83 Gases and AEROSols, (FIGAERO-CIMS) (Lopez-Hilfiker et al., 2014)). Here, we present results characterizing the  
84 SVN using an HR-ToF-AMS and an SMPS and demonstrate production and elemental analysis of aerosol using 2-4  
85  $\mu\text{L}$  of liquid samples, with masses of organic material as low as  $\sim 0.4 \mu\text{g}$ . In some cases, depending on the sample,  
86 pre-concentration is required to generate suitable solutions for analysis. The concentration ranges needed (described  
87 below) are comparable to the concentrations used for other offline characterizations including soft ionization with  
88 electrospray ionization into mass spectrometers. Thus, this technique provides a platform for direct comparison  
89 between offline-AMS samples and other analytical techniques. Quantification of total organic concentrations is  
90 demonstrated using internal standards. We examine the effects of aerosol collection, extraction, and nebulization on  
91 the mass spectra and elemental ratios observed for offline and online AMS. The aim of this work is to demonstrate  
92 that offline analysis of organic mixtures with the SVN-AMS can provide quantitative characteristic elemental ratios  
93 for trace environmental and biological samples using just micrograms of sample.

## 94 **2 Experimental**

### 95 **2.1 Small Volume Nebulizer**

96 The SVN, illustrated in Figure 1, creates an aerosol by ultrasonically nebulizing a small droplet placed on a  
97 thin film stretched across a water reservoir. The aerosol is then carried by a gentle flow of either house air (zero air,  
98 Aadco Instruments) or argon (Airgas, 99.999% purity) into the AMS. The three main components of the SVN,  
99 described in detail below, are (1) a bottom cylinder with an ultrasonic transducer and water bath, (2) a thin film that  
100 is press-fit onto the top of the water bath by an upper cylinder with a slightly larger ID, and (3) a vertical glass tube  
101 that connects to the AMS. The connections between all components are airtight, but the apparatus is easily  
102 disassembled to inject samples onto the film, as well as to clean the thin film and change the water bath.

103 In the bottom section of the SVN, the 2.4 MHz ultrasonic transducer (Sonaer, Inc., Model 241VM) is located  
104 just under the liquid reservoir, with a thin film stretched across the top of the reservoir to provide a clean nebulization  
105 surface for the sample. We use a 0.001" thick Kapton film or Teflon film, as these two were found to have the lowest  
106 background signal and the best performance in terms of the amount of aerosol generated compared to other materials  
107 tested. Press-fit onto the bottom piece is another PVC cylinder that has two side ports with carrier gas inlets, and a  
108 larger hole in the top into which a 15 cm glass tube is seated. The distance from the thin film to the bottom of the glass

109 tube is ~1.5 cm. During experiments, the nebulized aerosol is carried up through the vertical glass tube, into the  
110 stainless steel tubing that leads to the AMS. Additional components such as Nafion™ (Perma Pure LLC) driers can  
111 be placed inline if desired, but such modifications were not investigated in the present work.

112 Samples can be introduced into the SVN using two different approaches: discrete injections of individual  
113 samples (for individual “one-shot” measurements) or continuous addition of a sample flow (for continual analysis,  
114 enabling signal averaging). For most studies, MilliQ water was used as the solvent; in some cases we used HPLC-  
115 grade methanol, though the organic background signal is higher in that case, likely due to a combination of increased  
116 organic background in organic solvents and incomplete evaporation of methanol prior to measurement. For most of  
117 the work described here, we used discrete injections of 2-5  $\mu\text{L}$  of aqueous solutions manually deposited directly onto  
118 the center of the Kapton film. For continuous injections, solutions made with MilliQ or organic solvents were  
119 introduced via a syringe pump (Harvard Apparatus Model 22), which sends liquid flow (20-40  $\mu\text{L}/\text{min}$ ) through a  
120 borosilicate capillary entering the SVN via a small downward-facing hole in the upper PVC piece (Figure 1). In the  
121 future, such a port could also be used to provide automated discrete sample introduction using an autosampler.

122 For aqueous samples containing salts and small organic molecules, only 1-2% of the original sample mass  
123 was observed to remain on the thin film after a discrete injection (Figure S2). To ensure a clean surface between  
124 different samples, the surface was cleaned by nebulizing 2-8  $\mu\text{L}$  of MilliQ water off the surface 5-10 times over  
125 approximately one minute. The cleanliness of the surface was then verified by nebulizing a salt solution (at least 0.5  
126 g/L) between each sample. The salt solution is necessary to ensure that any contaminants can be seen, since pure water  
127 risks generating aerosol particles that are too small to be measured in the AMS. For samples in which carryover was  
128 observed (for example, the dissolved organic matter solutions discussed in section 3.1), additional cleaning of the film  
129 was undertaken with sonication in a deionized water bath followed by rinsing with HPLC-grade methanol for > 30  
130 seconds. Careful maintenance of the surface ensures uncontaminated mass spectra and accurate quantification of the  
131 solution components.

## 132 **2.2 AMS Data Collection and Analysis**

133 While a number of different aerosols instruments could be used with the SVN, here we focus primarily on  
134 elemental analysis by the HR-ToF-AMS. The AMS has previously been described in detail (Canagaratna et al., 2007;  
135 Decarlo et al., 2006) and provides quantitative measurements of non-refractory material (organics, ammonium sulfate,  
136 ammonium nitrate, etc.) for aerosol particles between approximately 40 and 1,000 nm. The mass spectrometer used  
137 in the AMS is a high resolution time-of-flight mass spectrometer (HToF-MS, ToFwerk AG), run under “V mode” for  
138 a mass resolution of 2,000-3,000  $m/\Delta m$ . This mass-resolving power enables peak fitting and identification of all  
139 organic fragment ions observed here ( $< 130 m/z$ ), which enables the calculation of quantitative elemental ratios for  
140 the organic mixture, after correcting for fragmentation bias during electron ionization (Aiken et al., 2007, 2008;  
141 Canagaratna et al., 2015). For AMS data collected using indoor or outdoor air, the intensities of  $\text{CO}^+$  and  $\text{H}_2\text{O}^+$  are  
142 complicated by gas-phase interferences ( $\text{N}_2^+$  and gas-phase  $\text{H}_2\text{O}^+$ ). For samples compared to chamber or ambient  
143 online-AMS data sets, house air was the carrier gas, standard empirical estimates were used, and the improved-ambient  
144 method for elemental ratios was applied (Canagaratna et al., 2015). With the SVN, inert carrier gases such as argon

145 can also be used, allowing for the direct measurement of the CO<sup>+</sup> ion intensity (as demonstrated below for dissolved  
146 organic matter, the majority of the other samples were run with zero air).

147 For discrete sampling, “fast MS” mode (Kimmel et al., 2010) was used because the pulse length of a single  
148 injection is ~30-60 seconds long. Fast MS mode generates mass spectra every 0.5-2 seconds and the instrument cycles  
149 between the “closed” state, in which the aerosol beam is blocked, and the “open” state, in which the aerosol beam can  
150 reach the vaporization/ionization region for detection. For the work shown here, mass spectra were collected every  
151 0.5 seconds for ~15-18 seconds in the “open” state, followed by 3 seconds in the closed state. The closed spectrum  
152 provides information on the instrument background, including contributions from gas phase species, and is subtracted  
153 from the open spectrum in data processing. For the high resolution peak fitting and the analysis of the mass spectrum  
154 and the elemental ratios, the average mass spectrum across all injections is used. For quantification, the total signal  
155 under each injection pulse is used. For continuous injections, the standard AMS operating mode (“GenAlt mode”)  
156 was used. This provides an average mass spectrum (by subtracting the closed signal from the open signal), as well as  
157 particle time-of-flight (PToF) data (providing aerosol size distributions for all aerosol components), once per minute.  
158 All AMS data were analyzed using software packages SQUIRREL (v1.57I) and PIKA (v1.16I), available at  
159 <http://cires1.colorado.edu/jimenez-group/ToFAMSResources/ToFSoftware/>.

160 The aerodynamic lens on the AMS has a transmission efficiency of nearly 100% for particles with  
161 aerodynamic diameters of 70-500 nm; for somewhat smaller particles (40-70 nm), this transmission is lower but not  
162 negligible (Jimenez et al., 2003). Thus, high enough solution concentrations are used such that the dried particles  
163 formed in the nebulizer are larger than ~100 nm aerodynamic diameter. Collection efficiencies (CE) in the AMS can  
164 vary depending on the extent to which aerosol particles bounce off the thermal element prior to vaporization (Docherty  
165 et al., 2013). This can impact the absolute concentrations observed, but for internally mixed samples, the relative  
166 concentrations of different aerosol components are independent of CE. In this work, most measurements (including  
167 elemental ratios) are reported as relative measurements, and thus no CE correction is applied. Some biases may arise  
168 if the aerosol is not internally mixed, but for all systems examined so far in PToF, no size-dependence in composition  
169 was observed (Figure S1).

### 170 **2.3 Sample Collection and Solution Preparation**

171 As described below, samples were prepared from a number of sources, including commercially available  
172 standards, the extracts of chamber and ambient aerosol particles collected on filters, and dissolved organic matter from  
173 the Pacific Ocean. For all solutions, either ultrapure water (18.2 M Ω cm, MilliQ) or HPLC-grade methanol was used.  
174 Prior to use, all glassware was cleaned with a methanol solvent wash and baked at 450°C for 6 hours.

175 Chamber aerosol (enabling offline vs. online comparisons) was generated in the MIT 7.5-m<sup>3</sup> Teflon  
176 environmental chamber, run in continuous-volume, “semi-batch” mode. Details on the facility are given elsewhere  
177 (Hunter et al., 2014). Experiments were run at 20 °C, < 5% RH, in the dark, and under low-NO<sub>x</sub> (< 10 ppb) conditions  
178 using ozone as the oxidant. Ammonium sulfate seeds were added for an initial concentration of ~60 μg/m<sup>3</sup>. The  
179 VOC, α-pinene, had an initial mixing ratio of 100 ppb; a penray lamp (Jelight model 600) was used to add an initial  
180 ozone concentration of ~700 ppb ozone. The ozone concentration decreased due to consumption and dilution to 400

181 ppb by the end of the experiment. The initial organic loading was  $\sim 70 \mu\text{g}/\text{m}^3$  and decayed due to dilution, sampling,  
182 and wall loss to a final value of  $\sim 18 \mu\text{g}/\text{m}^3$ . Filter samples were collected on Zeflour® PTFE Membrane Filters (0.5  
183  $\mu\text{m}$  pore size) at flow rates of  $\sim 5 \text{ L}/\text{min}$  for 10 hr. Laboratory blank filters were prepared by placing separate filters  
184 in the filter holder for 10 minutes before the start of the experiments. All filters were stored in baked aluminum foil  
185 packets, sealed in plastic bags, and placed in a freezer at  $-20 \text{ }^\circ\text{C}$  until extraction. Filters were extracted with  $\sim 4 \text{ mL}$  of  
186 HPLC-grade methanol. In order to avoid oxidation of the organic species in the extract, no sonication was used;  
187 instead, the vials were gently agitated by hand intermittently over 3 hours. Solutions were concentrated by evaporating  
188 to dryness under a gentle stream of ultra-high purity  $\text{N}_2$ . Dried samples were stored in the freezer at  $-20 \text{ }^\circ\text{C}$  until  
189 reconstitution with MilliQ water and analysis by the SVN-AMS. Blank subtraction was carried out with a scaling of  
190 the filter blank to 12% of the sample signal, as determined from the internal standard in each sample.

191 Field samples from the Southern Oxidant and Aerosol Study (SOAS) in 2013 were collected on pre-baked  
192 Tissuquartz™ Filters (Pall Life Science, 8 x 10 in) at Look Rock, TN starting on 06/16/2013 using a high-volume  
193 aerosol filter sampler with a  $\text{PM}_{2.5}$  cyclone (Tisch Environmental, Inc.) as described by Budisulistiorini et al. (2015).  
194 For filter extraction, a 37 mm punch was extracted in a pre-cleaned scintillation vials with 20 mL high-purity methanol  
195 (LC-MS Chromasolv-grade®, Sigma Aldrich) by sonication for 45 min. Filter extract was filtered through 0.2  $\mu\text{m}$   
196 syringe filter (Acrodisc® PTFE membrane, Pall Life Sciences) to remove suspended filter fibers. The filtered extract  
197 was then blown down to dryness under a gentle  $\text{N}_2(\text{g})$  stream at room temperature. An aerosol chemical speciation  
198 monitor (ACSM) (Ng et al., 2011a) was deployed at the same field site (Budisulistiorini et al., 2015); the average mass  
199 spectrum for the length of the filter sample was used for comparison with the present SVN-AMS measurements. For  
200 all analyses presented here (chamber and ambient) sufficient mass was extracted to enable the analysis of individual  
201 filter samples, with no combination of extracts from different samples required.

202 Standard solutions were prepared from commercially available compounds dissolved in MilliQ water.  
203 Reagents used included ammonium sulfate, ammonium nitrate, isotopically-labelled ammonium nitrate ( $\text{NH}_4^{15}\text{NO}_3$ ),  
204 citric acid, mannitol, PEG-400, 4-hydroxy-3-methoxy-DL-mandelic acid (HMMA), and HPLC grade methanol, all  
205 from Sigma-Aldrich, all at  $\geq 99\%$  purity.

206 The DOM polysaccharide sample was collected at the Natural Energy Laboratory Hawaii Authority facility  
207 in Kona, Hawaii. Seawater from a depth of 20 m was pumped through a 0.2  $\mu\text{m}$  filter to remove particles and the high  
208 molecular weight fraction of organic matter in the filtrate was concentrated by ultrafiltration using a membrane with  
209 a 1 nm pore size and a nominal 1,000 Da molecular weight cut off. This fraction was desalted by serial  
210 dilution/concentration with MilliQ water and then freeze-dried. Low-molecular weight humic substances and residual  
211 salts were removed by stirring with anion (hydroxide form) and cation exchange resins (hydrogen form). The final  
212 product was freeze-dried to yield a fluffy white powder. Conventional CHNS analysis was carried out using a CE-440  
213 Elemental Analyzer (Exeter Analytical). This powder was dissolved in MilliQ water at approximately 1 g/L to prepare  
214 solutions for analysis.

## 215 **3 Results and Discussion**

### 216 **3.1 Nebulization and Aerosol Size**

217 Figure 2a shows a time series of measured aerosol mass concentrations of a typical nebulized aerosol pulse  
218 from a 4  $\mu\text{L}$  solution containing approximately 0.33 g/L each of mannitol, ammonium sulfate, and ammonium nitrate.  
219 The nebulizer is turned on at  $t = 0$  and shortly afterwards ( $t = \sim 10$  s) the aerosol packet is observed in the AMS. The  
220 start of the nebulization is timed so that a closed (background) measurement occurs during the downslope of the signal  
221 ( $t = \sim 16$ -21 s, gaps). This background is subtracted from the aerosol particle signal during data processing.  
222 Measurements are collected until the signal returns to the baseline ( $t = \sim 44$  s).

223 Figure 2b shows the size distribution of the particles generated by nebulizing an aqueous solution of citric  
224 acid with continuous injection via syringe pump and a total concentration of  $\sim 1$  g/L into an SMPS (TSI). The particles  
225 have size distributions centered at 150-200 nm (electrical mobility diameter). These particles were sampled into the  
226 SMPS without passing through a drier. The SVN was approximately three meters further away from the inlet of the  
227 SMPS so the particles are likely to be somewhat smaller than those entering the AMS, due to water evaporation in the  
228 dry carrier gas. We find injections of solutions with total concentrations above 0.2 g/L provide sufficient aerosol mass  
229 for analysis (Figure S1). These measurements compare well with calculations based on the size of droplets reported  
230 by the manufacturer (Sonaer inc.) of approximately 1.7  $\mu\text{m}$  using water solutions. Assuming that the density of the  
231 dried particle is 1.3 g/cm<sup>3</sup> (Nakao et al., 2013), the minimum sample concentration that will form a 100 nm dried  
232 particle is approximately 0.3 g/L. More dilute solutions do not generate signal in the AMS because the majority of the  
233 aerosol particles that are formed are too small for transmission through the aerodynamic lens of the AMS (Figure 2b).  
234 To generate large enough aerosol particles from more dilute solutions, larger initial droplets could be formed by  
235 changing the transducer to one that vibrates at a lower frequency. However, for these larger droplets, drying will  
236 require the loss of a greater amount of solvent, so that any impurities in the solvent will make up a larger (and possibly  
237 even dominant) fraction of the resulting fine particles. Thus the use of ultrasonic nebulization at lower frequencies  
238 was not investigated here.

239

### 240 **3.2 Quantification**

#### 241 **3.2.1 Nebulization Efficiency**

242 A key quantity describing the potential sensitivity of the SVN-AMS is the SVN nebulization efficiency, the  
243 ratio of the mass measured in the AMS compared to the mass of analyte placed on the thin film. This was determined  
244 by loading 4  $\mu\text{L}$  of a known solution onto the film and measuring the mass of each component in the AMS integrated  
245 over the injection pulse, determined by:

$$246 \quad M_{AMS} = \int_{t_1}^{t_2} f(t) dt \times v_{AMS}$$

247 where  $M_{AMS}$  is the mass measured by the AMS in  $\mu\text{g}$ ,  $f(t)$  is the instantaneous mass concentration measured in the  
248 AMS ( $\mu\text{g}/\text{m}^3$ ), and  $v_{ams}$ , is the gas flow rate into the AMS in  $\text{m}^3/\text{s}$ . For each injection, the background-subtracted AMS  
249 signal is calculated (Figure 2a). The gaps due to closed cycles are bridged by interpolation and the area under the

250 injection curve is calculated via trapezoidal integration from time points before and after the pulse ( $t_1$  and  $t_2$ ,  
251 respectively) with the time steps ( $dt$ ) corresponding to the MS cycle time (here 0.5 s). The mass measured in the AMS  
252 is affected by three factors: the amount of aerosol formed and transported out of the SVN, the fraction of the gas flow  
253 from the SVN that is sampled by the AMS (typically  $\sim 1/2$ ), and the fraction of aerosol that bounces off the heater  
254 element before vaporizing (the AMS CE).

255 Figure 3 shows the mass measured in the AMS compared to the mass deposited on the nebulizer for replicate  
256 injections of four different aqueous solutions of citric acid, ammonium nitrate, ammonium sulfate, and isotopically-  
257 labeled ammonium nitrate ( $\text{NH}_4^{15}\text{NO}_3$ , used later as an internal standard) with concentrations ranging between  
258 approximately 0.1 and 0.2 g/L for each of the components (but with the same total concentration, 0.75 g/L). Six  
259 replicate injections of 4  $\mu\text{L}$  drops of the solutions from one of the calibration curves (section 3.2.2 below) were  
260 atomized, and the total mass observed in the AMS was calculated as described above. (Details on the concentrations  
261 of analytes in these calibration solutions for Figures 3 and 4 are provided in the supplemental.) There are variations in  
262 the efficiency from sample to sample and run to run, thus the trends shown in Figure 3 are illustrative only. The key  
263 trait observed is that the measured nebulization efficiencies are on the order of 0.02-0.06%, indicating that the aerosol  
264 mass detected with the AMS is approximately three orders of magnitude lower than the mass originally deposited on  
265 the thin film.

266 The majority of the sample mass loss likely occurs during the nebulization process itself. For aqueous  
267 solutions in the SVN, large droplets are observed to be ejected off the surface of the film at the same time as the  
268 aerosol is generated. These ejected droplets are then lost to the walls of the SVN. The ejection of these droplets  
269 appears to be a necessary part of the nebulization mechanism for water samples as smaller volumes ( $< 1 \mu\text{L}$ ) of water  
270 do not generate such droplets and also do not appear to form aerosol. This observed mechanism is in agreement with  
271 previous studies of aerosol generation for ultrasonic nebulization, in which cavitation within the droplet (Lang, 1962)  
272 and boiling and/or jetting from a droplet chain (Simon et al., 2015) have been observed.

273 The size distribution and number of aerosol particles from ultrasonic nebulization have been shown to be  
274 affected by the frequency of the ultrasonic vibration, the properties of the liquid including surface tension, density,  
275 and viscosity, and the concentration of the solution (Donnelly et al., 2005; Lang, 1962; Simon et al., 2015). The present  
276 application involves relatively dilute solution, so the only parameter that is likely to vary is the surface tension, by use  
277 of different solvents. Nebulization of solvents with lower surface tension, such as methanol, led to the ejection of  
278 much smaller droplets, and consequently substantially higher nebulization efficiencies ( $\sim 10\%$ ). However, methanol  
279 (and other HPLC-grade organic solvents) was found to give higher background signals in the AMS than MilliQ water,  
280 likely due to higher levels of low-volatility contaminants. This difference was also observed by Daellenbach *et al.*  
281 (2016); therefore, MilliQ water appears to be the ideal solvent to use for most environmental samples. However, with  
282 adequate solvent background characterization, organic solvents may be optimal for environmental samples with more  
283 non-polar components (e.g. petroleum or fresh tail pipe emissions).



### 284 3.2.2 Internal Standards and Calibration Curves

285 In Figure 3, the vertical spread of data points illustrates the variation in nebulization efficiency from one  
286 injection to the next. This is likely the result of small differences in the droplet shape or position on the film, leading  
287 to differences in how the droplets are ejected from the surface during aerosol formation. This run-to-run variability  
288 in nebulization efficiency, as well as the lack of a linear correlation between the mass placed on the film and the mass  
289 observed, complicates quantification, and necessitates the use of an internal standard to quantify the concentration of  
290 organic species within the original sample. In some cases, an inorganic ion that is independently quantified, such as  
291 sulfate, can serve as this internal standard (Daellenbach et al., 2016). However, in many cases such an independent  
292 measurement is not available; additionally, some environmental samples may not contain appreciable levels of  
293 measurable inorganic species, or else such species may not be soluble in the solvent of choice (e.g. ionic species in  
294 organic solvents). In these cases, an internal standard needs to be added to the solution prior to nebulization.

295 For use with the AMS, the internal standard must meet a number of requirements: it must be non-refractory,  
296 soluble, unreactive with the other sample components, not already present in the solution, and easily distinguishable  
297 from other species in the sample. For nebulization of samples dissolved in organic solvents, organic internal standards  
298 (e.g., phthalic acid (Chen et al., 2016; Han et al., 2016)) meet these requirements. In the present work, which focuses  
299 on aqueous samples only, we use an inorganic internal standard of isotopically-labelled ammonium nitrate  
300 ( $\text{NH}_4^{15}\text{NO}_3$ ). An example mass spectrum for an internal standard solution is shown in Figure 4a. The background  
301 signal from other components (organic material, sulfate, and nitrate) is very low. Another tested option is ammonium  
302 iodide ( $\text{NH}_4\text{I}$ ). Both of these salts work well as internal standards for both laboratory and ambient samples, since  
303 neither  $^{15}\text{NO}_3$  nor iodide are present in appreciable amounts in the atmosphere and there is usually a very small  
304 contribution of organic fragments at the fragment masses observed for those salts. Typically, the internal standards  
305 are added at the same order of magnitude concentration as the sample. For all tests of background signals and blanks,  
306 the internal standard is added to the solutions at concentrations between 0.5-1 g/L in order to generate aerosols of  
307 sufficient size for the AMS. This allows an analysis of any trace material present in the blank by creating an aerosol  
308 population to transfer the trace material into the AMS and allows for a background subtraction using the internal  
309 standard.

310 Figure 4b shows calibration curves with linear responses for three different organic compounds (citric acid,  
311 4-hydroxy-3-methoxy-DL-mandelic acid (HMMA), and polyethylene glycol 400 (PEG-400)) at four concentrations  
312 using  $\text{NH}_4^{15}\text{NO}_3$  as the internal standard. For the calibration curve, the ratios of the AMS signals for the analyte over  
313 the internal standard are compared to the ratios for known solution concentrations, thus correcting any variations in  
314 the mass of analyte nebulized. For quantification of unknowns, known concentrations of the internal standard are  
315 added to the samples at ratios comparable to what is used for the calibration curve. The ratio of the measured AMS  
316 signals can then be used to calculate the unknown analyte concentration from the calibration curve.

317 For quantification of complex organic mixtures using this technique, the most accurate organic calibration  
318 standards will have chemical structures similar to the average structure of the mixture. The slope of each line is related  
319 to the relative ionization efficiency (RIE) of the organic compound in the AMS (Jimenez et al., 2003). The RIE values  
320 in Figure 4b for HMMA and citric acid (1.01 and 1.95, respectively) bracket the range of RIE values for different

321 types of organics measured using standard AMS calibration techniques (Jimenez et al., 2016). This range likely arises  
322 from differences in how the organic compounds dissociate during volatilization on the heater. The heater in the AMS  
323 is typically set at 600°C, and so most organic molecules found in organic aerosol thermally decompose prior to electron  
324 impact ionization (Canagaratna et al., 2015) leading to RIEs in the range of 1.0-2.0. In contrast, the slope of 2.62 for  
325 PEG-400 is substantially outside of the range of values. However, with the AMS, complex mixtures are less likely to  
326 show large variations in RIE than different individual compounds, such as those used in Figure 4. For extracts of  
327 atmospheric aerosol or other smaller organic mixtures, the RIE of 1.4, which is typically used for AMS measurements  
328 (Canagaratna et al., 2007; Jimenez et al., 2016; Xu et al., 2018), is likely the best value to use as an initial calibration  
329 slope. For extracts of other types of organic mixtures, compounds that have a structure similar to the average organic  
330 composition should be used to calibrate the samples.

### 331 3.3 Mass Spectral Analysis

332 The primary goal of the SVN-AMS is to measure quantitative chemical information, specifically elemental  
333 ratios, from complex organic mixtures. We have characterized these for a number of different chemical systems,  
334 described below. Results are summarized in Figure 5 (comparing SVN-AMS and online AMS mass spectra) and Table  
335 1 (comparing elemental ratios measured with SVN-AMS with those measured by either online AMS or CHNS  
336 analysis).

337 One concern with using ultrasonic nebulization to generate aerosol particles is the possibility that the high  
338 temperatures possibly reached by the solution during nebulization may degrade the organic compounds, affecting their  
339 mass spectra (and hence measured elemental composition). Figure 5a shows a comparison of a solution containing 1  
340 g/L citric acid aerosolized with a TSI atomizer (TSI 3076, 40 psi gas) (black) and the SVN (gold), with the inset  
341 showing a direct comparison between the intensities measured for each ion in the mass spectrum. The degree of  
342 agreement can be described by the dot product of the intensities for matching peaks in the two spectra, as well as the  
343 log of the intensities before taking the dot product (log-dot product), which gives the lower intensity peaks greater  
344 weight. Very good overlap between the two mass spectra is observed, with a dot product of 0.99 and a log-dot product  
345 of 0.96. This indicates minimal degradation of the citric acid by ultrasonic nebulization.

346 A high degree of similarity is also observed between offline and online aerosol measurements for more  
347 complex mixtures. Figure 5b shows mass spectra for a comparison of offline (gold) vs. online (black) SOA sample,  
348 generated from the dark ozonolysis of  $\alpha$ -pinene. The online mass spectra is the average real-time AMS mass spectrum  
349 averaged over the 10 hours of filter collection. For all filter samples, spectra from the SVN are background subtracted  
350 using spectra collected from blank filter samples. These blanks provide the background for any trace organic material  
351 on the filters before collection as well as any background organic material introduced during sample preparation. The  
352 overlap in Figure 5b between the mass spectra is very good, with a dot product of 0.98 and a log-dot product of 0.98.  
353 The elemental ratios are also very similar between the two samples with an H:C of 1.6 for both and O:C of 0.48 for  
354 the chamber and 0.49 for the SVN samples (Table 1). The largest difference is observed at  $m/z$  44 ( $\text{CO}_2^+$ ) and  $m/z$  43  
355 ( $\text{C}_2\text{H}_3\text{O}^+$ ) with a larger fraction of  $\text{CO}_2^+$  in the offline sample. The intensity of  $\text{CO}^+$  ( $m/z$  28) is also different, but only  
356 because it is set equal to the intensity of the  $\text{CO}_2^+$  ion, as is commonly done for ambient sampling with the AMS

357 (given that the  $\text{CO}^+$  ion generally cannot be distinguished from the much more abundant  $\text{N}_2^+$  ion (Aiken et al., 2007)).  
358 The organic contribution from  $\text{H}_2\text{O}^+$ ,  $\text{OH}^+$ , and  $\text{O}^+$  is also constrained by the  $\text{CO}_2^+$  signal so any differences in  $\text{CO}_2^+$   
359 intensity will also show up in those ions (Aiken et al., 2008). The observed difference in  $\text{CO}_2^+$  and  $\text{C}_2\text{H}_3\text{O}^+$  ion intensity  
360 is likely a result of the extraction step prior to nebulization, which may preferentially dissolve the most water-soluble  
361 (oxidized) SOA components. Additionally, the online measurement is for  $\text{PM}_1$  while offline measurement is collecting  
362 the full range of particle sizes on the filter. However, based on the agreement in H:C and O:C in the online and offline  
363 cases, these factors do not appear to bias elemental ratio measurements substantially.

364 Figure 5c shows a comparison of online and offline measurements of ambient organic aerosol, specifically  
365 ACSM measurements and SVN-AMS measurements of a filter extract collected simultaneously during the 2013  
366 SOAS field campaign in Look Rock, TN (8 pm July 4 to 7am July 5, 2013; EST). Since the ACSM is a unit-mass-  
367 resolution instrument, the HR-AMS data are degraded to unit mass resolution, and ions that are determined from the  
368  $m/z$  44 signal ( $m/z=15, 16, 17, 18,$  and  $28$ ) are excluded from the analysis. Additionally, ions at  $m/z$  30 and 31 were  
369 removed from comparison because of interferences from the internal standard ( $m/z$  31) and nitrate in the sample ( $m/z$   
370 30).

371 The two mass spectra in Figure 5c have a high degree of agreement between the major ions (dot product of  
372 0.98). However, there is substantially more variation between the two techniques than in the chamber study, especially  
373 in the lower-abundance peaks ( $m/z>45$ ; see inset), as reflected in the lower log-dot product of only 0.90. Possible  
374 reasons for this lower correlation include fractionation from the extraction step, the different sizes measured ( $\text{PM}_{2.5}$   
375 for the filter vs.  $\text{PM}_1$  for the ACSM) (Daellenbach et al., 2016), the uncertainty in ACSM signals at higher masses due  
376 to uncertainty in the relative ion transmission curve (Ng et al., 2011a), and the losses of more volatile compounds  
377 during collection, extraction, and handling. Additional work is necessary to quantify the importance of these effects  
378 and care should be taken when comparing the full mass spectra for on-line compared to off-line SVN-AMS analysis.  
379 The high degree of overlap in the intensities of the dominant ions between the online (AMS/ACSM) measurements  
380 and offline (SVN-AMS) results indicates that the ensemble organic composition for these aerosol samples is generally  
381 well-represented by the SVN-AMS measurements (Table 1). However, the estimated elemental ratios from a lower  
382 resolution AMS are more uncertain than from the HR-ToF-AMS. Thus, the ratios for these samples in Table 1 are  
383 provided only as a demonstration of the overall agreement between the two techniques.

384 For the SVN, the small sample volume requirements can make it attractive for the analysis of other  
385 environmental samples that are soluble in water (or organic solvents), and that have low enough vapor pressures to  
386 remain in the condensed phase after sample preparation and nebulization. Here we demonstrate the analysis of the  
387 high molecular weight fraction of the polysaccharide fraction of dissolved organic matter (DOM) with the SVN-AMS.  
388 The DOM sample was prepared using a standard protocol for the isolation of this fraction of the organic material (see  
389 section 2.3). This preparation removes the lower molecular weight compounds so chemicals such as methane sulfonic  
390 acid are not expected to be observed. Figure 5d shows an example AMS mass spectrum from DOM collected from  
391 the Pacific Ocean. The mass spectrum is dominated by oxidized fragments containing one or more oxygen atoms  
392 with smaller amounts of nitrogen-containing fragments. The sample preparation for the DOM removed all salts, thus  
393 the ammonium fragments were assumed to be from organonitrogen species and were assigned to the organic fraction.

394 The measured N:C and H:C values of 0.081 and 1.7, respectively, matches those measured by CHNS analysis (0.080  
395 and 1.74, respectively). This demonstrates that with the SVN, microgram quantities of dissolved environmental  
396 mixtures can be nebulized and sampled into the AMS providing a rapid, quantitative method to determine elemental  
397 ratios in these complex organic mixtures.

#### 398 **4 Conclusions**

399 A new ultrasonic nebulizer has been described and characterized for generation of aerosol from very small  
400 sample masses. We demonstrate the application of this technique to offline AMS analysis of complex organic  
401 mixtures from aerosol filter extracts and DOM. Data sets that include quantitative organic mass, characteristic mass  
402 spectra, and quantitative elemental ratios can be generated from only 0.4-1.2  $\mu\text{g}$  of material. For these samples, pre-  
403 concentration was required to prepare a suitable solution concentration for analysis. This will be required for some  
404 types of environmental samples and care should be taken to minimize artifacts during solution preparation. A direct  
405 comparison between the mass spectra generated by commercial spray atomizers or by particles sampled directly from  
406 the atmosphere showed high degrees of agreement, indicating minimal composition changes during sample  
407 preparation and nebulization. Nebulization of aqueous samples generated measurable aerosol from 0.1% of the sample  
408 mass. Higher nebulization efficiencies (and smaller ejected droplets) were observed for methanol, likely due to its  
409 lower surface tension. The SVN, combined with offline-AMS, provides rapid analysis of non-refractory organic and  
410 inorganic compounds. For other types of characterization, including analysis of refractory material or organic  
411 molecular composition, the SVN can also be coupled with other aerosol instrumentation such as PALMS or CIMS  
412 instruments.

413 Future improvements in the nebulization and collection efficiency of the SVN-AMS will enable analysis with  
414 even lower sample mass requirements. The use of organic internal standards is one method to potentially improve  
415 collection efficiency in the AMS as the higher organic content may decrease the bounce of particles off the vaporizer.  
416 Additionally, the use of solvents with lower surface tension than water shows promise for improved nebulization  
417 efficiencies. Finally, in contrast to atomizers (where the carrier gas generates the aerosol), ultrasonic nebulizers  
418 decouple the aerosol formation from the carrier gas flow rate, enabling potential concentration of the aerosol prior to  
419 sampling. A useful future direction for this technique will be to characterize the background signal in different organic  
420 solvents and optimize the continuous flow configuration to minimize the return of large ejected droplets back onto the  
421 film. Continuous flow with organic solutions will also enable the analysis of more hydrophobic organic samples such  
422 as fresh vehicle emissions, cooking oils, and petrochemical samples. In the future, the SVN can be used to generate  
423 aerosol for quantitative and qualitative analysis of other environmental samples to investigate sources or  
424 processing/aging of these organic mixtures. The SVN, combined with aerosol measurement techniques such as the  
425 AMS, provides a rapid, quantitative method to characterize the chemical and elemental properties of complex organic  
426 mixtures, producing rich data sets for the exploration of exceptionally trace environmental samples.

427

428 **Supporting Information**

429 The supporting information is available free of charge at DOI:xxx. The document contains additional information on  
430 particle sizes and memory effects between runs, (file type, PDF).

431 **Corresponding Author**

432 Rachel E. O'Brien, College of William and Mary, reobrien@wm.edu.

433 **Author Contributions**

434 MRC, JTJ, PLC, DRW, JHK, and KJR designed and built the SVN. SHB, JDS, CLF, DJR provided ambient aerosol  
435 samples and DOM. REO and JHK designed experiments and REO carried them out. REO prepared the manuscript  
436 with contributions from all authors.

437 **Data availability**

438 All data sets including mass spectra and SMPS data are available on request from REO, reobrien@wm.edu.

439 **Acknowledgement**

440 This work was supported by National Oceanic and Atmospheric Administration Grants No. NA13OAR4310072 and  
441 NA140AR4310132. KJB acknowledges support from the National Science Foundation. SHB and JDS acknowledges  
442 support from the US Environmental Protection Agency Award No. 835404, Electric Power Research Institute (EPRI),  
443 and National Oceanic and Atmospheric Administration Grant No. NA13OAR4310064. Special thanks to Dr. David  
444 Karl and Mr. Eric Grabowski, University of Hawaii, for the CHNS elemental analysis of DOM. DJR acknowledges  
445 support from the Gordan and Betty Moore Foundation award 6000 and the Simons Foundation SCOPE award 329108.

446 **References**

447 Aiken, A. C., DeCarlo, P. F. and Jimenez, J. L.: Elemental analysis of organic species with electron ionization high-  
448 resolution mass spectrometry, *Anal. Chem.*, 79(21), 8350–8358, doi:10.1021/ac071150w, 2007.

449 Aiken, A. C., Decarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S., Ulbrich, I. M., Mohr, C.,  
450 Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M., Ziemann, P. J., Canagaratna, M. R.,  
451 Onasch, T. B., Alfarra, M. R., Prevot, A. S. H., Dommen, J., Duplissy, J., Metzger, A., Balensperger, U. and Jimenez,  
452 J. L.: O/C and OM/OC Ratios of Primary, Secondary, and Ambient Organic Aerosols with High-Resolution Time-of-  
453 Flight Aerosol Mass Spectrometry, *Environ. Sci. Technol*, 42, 4478–4485, doi:10.1021/es703009q, 2008.

454 Becker, J. W., Berube, P. M., Follett, C. L., Waterbury, J. B., Chisholm, S. W., Delong, E. F., Repeta, D. J. and Metz,

455 T.: Closely related phytoplankton species produce similar suites of dissolved organic matter, *Front. Microbiol.*, 5(111),  
456 1–14, doi:10.3389/fmicb.2014.00111, 2014.

457 Bozzetti, C., Daellenbach, K. R., Hueglin, C., Fermo, P., Sciare, J., Kasper-Giebl, A., Mazar, Y., Abbaszade, L., El  
458 Kazzi, M., Gonzalez, R., Shuster-Meiseles, T., Flasch, M., Wolf, R., Kr, A., Francesco Canonaco, E., Schnelle-Kreis,  
459 R., Slowik, J. G., Zimmermann, R., Rudich, Y., Baltensperger, U., El Haddad, I. and Preo, A. H.: Size-Resolved  
460 Identification, Characterization, and Quantification of Primary Biological Organic Aerosol at a European Rural Site,  
461 *Environ. Sci. Technol.*, 50, 3425–3434, doi:10.1021/acs.est.5b05960, 2016.

462 Bozzetti, C., Sosedova, Y., Xiao, M., Daellenbach, K. R., Ulevicius, V., Dudoitis, V., Mordas, G., Byčenkienė E, S.  
463 E., Plauškaitė E, K., Vlachou, A., Golly, B., Chazeau, B., Besombes, J.-L., Baltensperger, U., Jaffrezo, J.-L., Slowik,  
464 J. G., Haddad, I. El and Prévôt, A. S. H.: Argon offline-AMS source apportionment of organic aerosol over yearly  
465 cycles for an urban, rural, and marine site in northern Europe, *Atmos. Chem. Phys*, 17, 117–141, doi:10.5194/acp-17-  
466 117-2017, 2017.

467 Budisulistiorini, S. H., Li, X., Bairai, S. T., Renfro, J., Liu, Y., Liu, Y. J., Mckinney, K. A., Martin, S. T., McNeill, V.  
468 F., Pye, H. O. T., Nenes, A., Neff, M. E., Stone, E. A., Mueller, S., Knote, C., Shaw, S. L., Zhang, Z., Gold, A. and  
469 Surratt, J. D.: Examining the effects of anthropogenic emissions on isoprene-derived secondary organic aerosol  
470 formation during the 2013 Southern Oxidant and Aerosol Study (SOAS) at the Look Rock, Tennessee ground site,  
471 *Atmos. Chem. Phys*, 15, 8871–8888, doi:10.5194/acp-15-8871-2015, 2015.

472 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q., Onasch, T. B., Drewnick, F.,  
473 Coe, H., Middlebrook, A., Delia, A., Williams, L. R., Trimborn, A. M., Northway, M. J., Decarlo, P. F., Kolb, C. E.,  
474 Davidovits, P. and Worsnop, D. R.: Chemical and Microphysical Characterization of Ambient Aerosols with the  
475 Aerodyne Aerosol Mass Spectrometer, *Mass Spec Rev*, 26, 185–222, doi:10.1002/mas.20115, 2007.

476 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P., Hildebrandt Ruiz, L., Fortner,  
477 E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue, N. M., Jayne, J. T. and Worsnop, D. R.: Elemental ratio  
478 measurements of organic compounds using aerosol mass spectrometry: characterization, improved calibration, and  
479 implications, *Atmos. Chem. Phys*, 15, 253–272, doi:10.5194/acp-15-253-2015, 2015.

480 Chen, Q., Heald, C. L., Jimenez, J. L., Canagaratna, M. R., Zhang, Q., He, L.-Y., Huang, X.-F., Campuzano-Jost, P.,  
481 Palm, B. B., Poulain, L., Kuwata, M., Martin, S. T., Abbatt, J. P. D., Lee, A. K. Y. and Liggio, J.: Elemental  
482 composition of organic aerosol: The gap between ambient and laboratory measurements, *Geophys. Res. Lett.*, 42(10),  
483 4182–4189, doi:10.1002/2015GL063693, 2015.

484 Chen, Q., Ikemori, F. and Mochida, M.: Light Absorption and Excitation–Emission Fluorescence of Urban Organic  
485 Aerosol Components and Their Relationship to Chemical Structure, *Environ. Sci. Technol.*, 50, 10859–10868,  
486 doi:10.1021/acs.est.6b02541, 2016.

487 Daellenbach, K. R., Bozzetti, C., Křepelová, A., Canonaco, F., Wolf, R., Zotter, P., Fermo, P., Crippa, M., Slowik, J.  
488 G., Sosedova, Y., Zhang, Y., Huang, R. J., Poulain, L., Szidat, S., Baltensperger, U., El Haddad, I. and Prévôt, A. S.  
489 H.: Characterization and source apportionment of organic aerosol using offline aerosol mass spectrometry, *Atmos.*  
490 *Meas. Tech.*, 9(1), 23–39, doi:10.5194/amt-9-23-2016, 2016.

491 Daumit, K. E., Kessler, S. H. and Kroll, J. H.: Average chemical properties and potential formation pathways of highly  
492 oxidized organic aerosol, *Faraday Discuss.*, 165, 181–201, doi:10.1039/c3fd00045a, 2013.

493 Decarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C., Gonin, M., Fuhrer, K.,  
494 Horvath, T., Docherty, K. S., Worsnop, D. R. and Jimenez, J. L.: Field-Depolyable, High-Resolution, Time-of-Flight  
495 Aerosol Mass Spectrometer, *Anal. Chem.*, 78(24), 8281–8289, doi:8410.1029/2001JD001213.Analytical, 2006.

496 Docherty, K. S., Jaoui, M., Corse, E., Jimenez, J. L., Offenberg, J. H., Lewandowski, M. and Kleindienst, T. E.:  
497 Collection Efficiency of the Aerosol Mass Spectrometer for Chamber-Generated Secondary Organic Aerosols,  
498 *Aerosol Sci. Technol.*, 47, 294–309, doi:10.1080/02786826.2012.752572, 2013.

499 Donnelly, T. D., Hogan, J., Mugler, A., Schubmehl, M., Schommer, N., Bernoff, A. J., Dasnurkar, S. and Ditmire, T.:  
500 Using ultrasonic atomization to produce an aerosol of micron-scale particles, *Rev. Sci. Instrum.*, 76(11), 1–10,  
501 doi:10.1063/1.2130336, 2005.

502 Gallimore, P. J. and Kalberer, M.: Characterizing an Extractive Electrospray Ionization (EESI) Source for the Online  
503 Mass Spectrometry Analysis of Organic Aerosols, *Environ. Sci. Technol.*, (47), 734–7331, doi:10.1021/es305199h,  
504 2013.

505 Ge, X., Li, L., Chen, Y., Chen, H., Wu, D., Wang, J., Xie, X., Ge, S., Ye, Z., Xu, J. and Chen, M.: Aerosol  
506 characteristics and sources in Yangzhou, China resolved by offline aerosol mass spectrometry and other techniques,  
507 *Environ. Pollut.*, 225, 74–85, doi:10.1016/j.envpol.2017.03.044, 2017.

508 Han, Y., Kawamura, K., Chen, Q. and Mochida, M.: Formation of high-molecular-weight compounds via the  
509 heterogeneous reactions of gaseous C8–C10 n-aldehydes in the presence of atmospheric aerosol components, *Atmos.*  
510 *Environ.*, 126, 290–297, doi:10.1016/j.atmosenv.2015.11.050, 2016.

511 Hansman, R. L., Dittmar, T. and Herndl, G. J.: Conservation of dissolved organic matter molecular composition during  
512 mixing of the deep water masses of the northeast Atlantic Ocean, *Mar. Chem.*, 177, 288–297,  
513 doi:10.1016/j.marchem.2015.06.001, 2015.

514 Heald, C. L., Kroll, J. H., Jimenez, J. L., Docherty, K. S., Decarlo, P. F., Aiken, A. C., Chen, Q., Martin, S. T., Farmer,  
515 D. K. and Artaxo, P.: A simplified description of the evolution of organic aerosol composition in the atmosphere,  
516 *Geophys. Res. Lett.*, 37(L08803), 1–5, doi:10.1029/2010GL042737, 2010.

517 Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S. M.,  
518 Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski,  
519 M., Abbazade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., Haddad, I. El and  
520 Prévôt, A. S. H.: High secondary aerosol contribution to particulate pollution during haze events in China, *Nature*,  
521 514, doi:10.1038/nature13774, 2014.

522 Hunter, J. F., Carrasquillo, A. J., Daumit, K. E. and Kroll, J. H.: Secondary Organic Aerosol Formation from Acyclic,  
523 Monocyclic, and Polycyclic Alkanes, *Environ. Sci. Technol.*, 48, 10227–10234, doi:10.1021/es502674s, 2014.

524 Jimenez, J. L., Jayne, J. T., Shi, Q., Kolb, C. E., Worsnop, D. R., Yourshaw, I., Seinfeld, J. H., Flagan, R. C., Zhang,  
525 X., Smith, K. A., Morris, J. W. and Davidovits, P.: Ambient aerosol sampling using the Aerodyne Aerosol Mass  
526 Spectrometer, *J. Geophys. Res.*, 108, 8425, doi:10.1029/2001JD001213, 2003.

527 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., Decarlo, P. F., Allan, J.  
528 D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J.,  
529 Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen,  
530 J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman,  
531 J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S.,  
532 Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimo,  
533 A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M.,  
534 Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U. and Worsnop, D. R.: Evolution of  
535 Organic Aerosols in the Atmosphere, *Science* (80- ), 326, 1525–1529, doi:10.1126/science.1179518, 2009.

536 Jimenez, J. L., Canagaratna, M. R., Drewnick, F., Allan, J. D., Alfarra, M. R., Middlebrook, A. M., Slowik, J. G.,  
537 Zhang, Q., Coe, H., Jayne, J. T. and Worsnop, D. R.: Comment on “The effects of molecular weight and thermal  
538 decomposition on the sensitivity of a thermal desorption aerosol mass spectrometer,” *Aerosol Sci. Technol.*, 50, i–xv,  
539 doi:10.1080/02786826.2016.1205728, 2016.

540 Kaul, D. S., Gupta, T. and Tripathi, S. N.: Source Apportionment for Water Soluble Organic Matter of Submicron  
541 Aerosol: A Comparison between Foggy and Nonfoggy Episodes, *Aerosol Air Qual. Res.*, 14, 1527–1533,  
542 doi:10.4209/aaqr.2013.10.0319, 2014.

543 Kimmel, J. R., Farmer, D. K., Cubison, M. J., Sueper, D., Tanner, C., Nemitz, E., Worsnop, D. R., Gonin, M. and  
544 Jimenez, J. L.: Real-time aerosol mass spectrometry with millisecond resolution, *Int. J. Mass Spectrom.*, 303, 15–26,  
545 doi:10.1016/j.ijms.2010.12.004, 2010.

546 Koch, B. P., Witt, M., Engbrodt, R., Dittmar, T. and Kattner, G.: Molecular formulae of marine and terrigenous  
547 dissolved organic matter detected by electrospray ionization Fourier transform ion cyclotron resonance mass  
548 spectrometry, *Geochim. Cosmochim. Acta*, 69(13), 3299–3308, doi:10.1016/j.gca.2005.02.027, 2005.



549 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R., Altieri, K. E.,  
550 Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D. and Kolb, C. E.: Carbon oxidation state as a  
551 metric for describing the chemistry of atmospheric organic aerosol, *Nat. Chem.*, 3, 133–139,  
552 doi:10.1038/NCHEM.948, 2011.

553 Kroll, J. H., Lim, C. Y., Kessler, S. H. and Wilson, K. R.: Heterogeneous Oxidation of Atmospheric Organic Aerosol:  
554 Kinetics of Changes to the Amount and Oxidation State of Particle-Phase Organic Carbon, *J. Phys. Chem. A*, 119,  
555 10767–10783, doi:10.1021/acs.jpca.5b06946, 2015.

556 Lang, R. J.: Ultrasonic Atomization of Liquids, *J. Acoust. Soc. Am.*, 34(1), 6–8, 1962.

557 Lee, A. K. Y., Hayden, K. L., Herckes, P., Leitch, W. R., Liggio, J., Macdonald, A. M. and Abbatt, J. P. D.:  
558 Characterization of aerosol and cloud water at a mountain site during WACS 2010: secondary organic aerosol  
559 formation through oxidative cloud processing, *Atmos. Chem. Phys. Atmos. Chem. Phys.*, 12, 7103–7116,  
560 doi:10.5194/acp-12-7103-2012, 2012.

561 Lopez-Hilfiker, F. D., Mohr, C., Ehn, M., Rubach, F., Kleist, E., Wildt, J., Mentel, T. F., Lutz, A., Hallquist, M.,  
562 Worsnop, D. and Thornton, J. A.: A novel method for online analysis of gas and particle composition: description and  
563 evaluation of a Filter Inlet for Gases and AEROSols (FIGAERO), *Atmos. Meas. Tech.*, 7, 983–1001, doi:10.5194/amt-  
564 7-983-2014, 2014.

565 Lu, Y., Li, X., Mesfioui, R., Bauer, J. E., Chambers, R. M., Canuel, E. A. and Hatcher, P. G.: Use of ESI-FTICR-MS  
566 to Characterize Dissolved Organic Matter in Headwater Streams Draining Forest-Dominated and Pasture-Dominated  
567 Watersheds, *PLoS One*, 1–21, doi:10.1371/journal.pone.0145639, 2015.

568 Mihara, T. and Mochida, M.: Characterization of Solvent-Extractable Organics in Urban Aerosols Based on Mass  
569 Spectrum Analysis and Hygroscopic Growth Measurement, *Environ. Sci. Technol.*, 45, 9168–9174,  
570 doi:10.1021/es201271w, 2011.

571 Murphy, D. M., Thomson, D. S. and Mahoney, M. J.: In Situ Measurements of Organics, Meteoritic Material,  
572 Mercury, and Other Elements in Aerosols at 5 to 19 Kilometers, *Science (80-. )*, 282, 1664–1669, 1998.

573 Nakao, S., Tang, P., Tang, X., Clark, C. H., Qi, L., Seo, E., Asa-Awuku, A. and Iii, D. C.: Density and elemental ratios  
574 of secondary organic aerosol: Application of a density prediction method, *Atmos. Environ.*, 68, 273–277,  
575 doi:10.1016/j.atmosenv.2012.11.006, 2013.

576 Ng, N. L., Herndon, S. C., Trimborn, a., Canagaratna, M. R., Croteau, P. L., Onasch, T. B., Sueper, D., Worsnop, D.  
577 R., Zhang, Q., Sun, Y. L. and Jayne, J. T.: An Aerosol Chemical Speciation Monitor (ACSM) for Routine Monitoring  
578 of the Composition and Mass Concentrations of Ambient Aerosol, *Aerosol Sci. Technol.*, 45(7), 780–794,  
579 doi:10.1080/02786826.2011.560211, 2011a.

580 Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Chhabra, P. S., Seinfeld, J. H. and Worsnop, D. R.: Changes in organic  
581 aerosol composition with aging inferred from aerosol mass spectra, *Atmos. Chem. Phys.* *Atmos. Chem. Phys.*, 11,  
582 6465–6474, doi:10.5194/acp-11-6465-2011, 2011b.

583 Simon, J. C., Sapozhnikov, O. A., Khokhlova, V. A., Crum, L. A. and Bailey, M. R.: Ultrasonic atomization of liquids  
584 in drop-chain acoustic fountains, *J Fluid Mech.*, 766, 129–146, doi:10.1017/jfm.2015.11.Ultrasonic, 2015.

585 Skoog, D. F., Holler, F. J. and Nieman, T. A.: *Principles of Instrumental Analysis*, Saunders College Pub,  
586 Philadelphia., 1998.

587 Sun, Y., Zhang, Q., Zheng, M., Ding, X., Edgerton, E. S. and Wang, X.: Characterization and Source Apportionment  
588 of Water-Soluble Organic Matter in Atmospheric Fine Particles (PM 2.5) with High-Resolution Aerosol Mass  
589 Spectrometry and GC-MS, *Environ. Sci. Technol.*, 45, 4854–4861, doi:10.1021/es200162h, 2011.

590 Xu, J., Zhang, Q., Li, X., Ge, X., Xiao, C., Ren, J. and Qin, D.: Dissolved Organic Matter and Inorganic Ions in a  
591 Central Himalayan Glacier-Insights into Chemical Composition and Atmospheric Sources, *Environ. Sci. Technol.*,  
592 47, 6181–6188, doi:10.1021/es4009882, 2013.

593 Xu, J. Z., Zhang, Q., Wang, Z. B., Yu, G. M., Ge, X. L. and Qin, X.: Chemical composition and size distribution of  
594 summertime PM<sub>2.5</sub> at a high altitude remote location in the northeast of the Qinghai-Xizang (Tibet) Plateau: Insights  
595 into aerosol sources and processing in free troposphere, *Atmos. Chem. Phys.*, 15(9), 5069–5081, doi:10.5194/acp-15-  
596 5069-2015, 2015.

597 Xu, L., Guo, H., Weber, R. J. and Ng, N. L.: Chemical Characterization of Water-Soluble Organic Aerosol in  
598 Contrasting Rural and Urban Environments in the Southeastern United States, *Environ. Sci. Technol.*, 51, 78–88,  
599 doi:10.1021/acs.est.6b05002, 2016.

600 Xu, W., Lambe, A., Silva, P., Hu, W., Onasch, T., Williams, L., Croteau, P., Zhang, X., Renbaum-Wolff, L., Fortner,  
601 E., Jimenez, J. L., Jayne, J., Worsnop, D. and Canagaratna, M.: Laboratory evaluation of species-dependent relative  
602 ionization efficiencies in the Aerodyne Aerosol Mass Spectrometer Laboratory evaluation of species-dependent  
603 relative ionization efficiencies in the Aerodyne Laboratory eva, *Aerosol Sci. Technol.*, 52(6), 626–641,  
604 doi:10.1080/02786826.2018.1439570, 2018.

605 Ye, Z., Liu, J., Gu, A., Feng, F., Liu, Y., Bi, C., Xu, J., Li, L., Chen, H., Chen, Y., Dai, L., Zhou, Q. and Ge, X.:  
606 Chemical characterization of fine particulate matter in Changzhou, China, and source apportionment with offline  
607 aerosol mass spectrometry, *Atmos. Chem. Phys.*, 17, 2573–2592, doi:10.5194/acp-17-2573-2017, 2017.

608

609

610 **Figures and Tables**

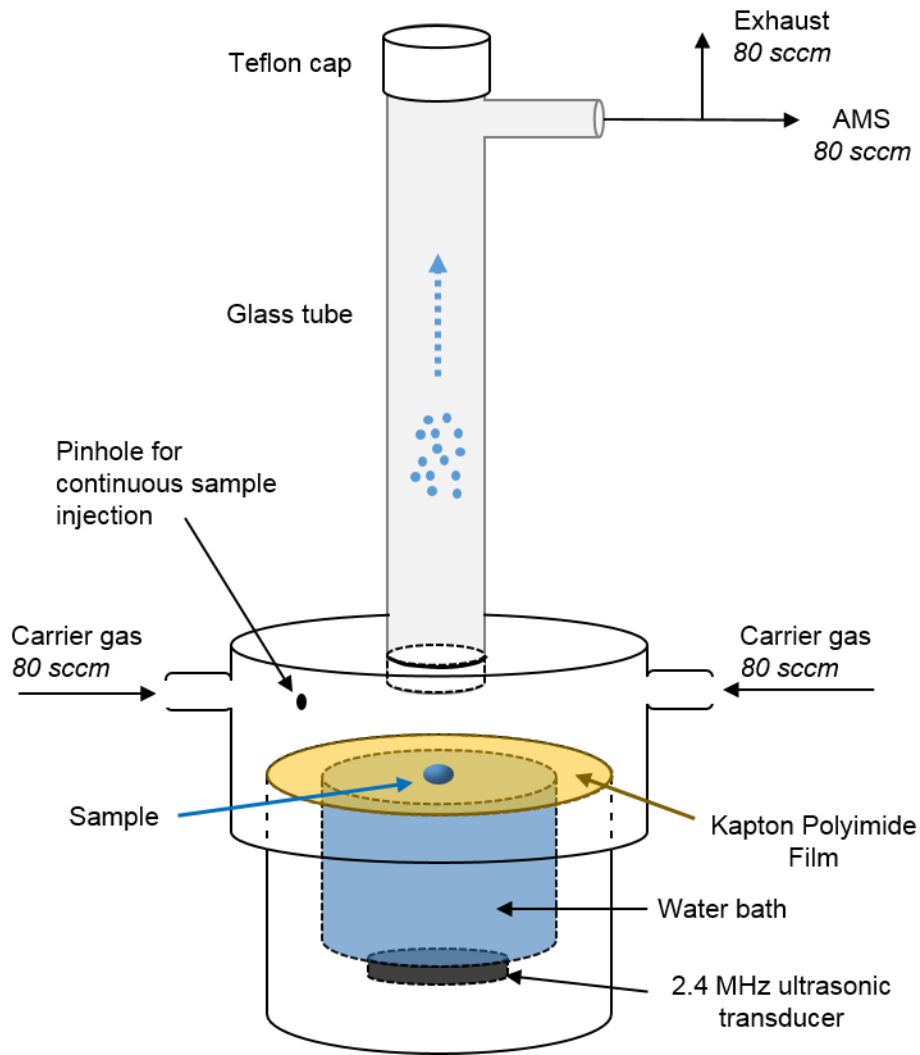
611 Table 1. Elemental ratios measured by SVN-AMS vs. other techniques for the various mixtures examined in this work.

Sample		O:C	H:C	N:C
Citric acid	Atomizer-AMS	1.0	1.4	--
	SVN-AMS	1.1	1.3	--
	Actual	1.2	1.3	--
$\alpha$ -pinene SOA	Online-AMS	0.48	1.6	< 0.002
	SVN-AMS	0.50	1.6	< 0.002
Look Rock	Online-ACSM <sup>a</sup>	0.13 ( $f_{44}=0.19$ )	1.3 ( $f_{43}=0.062$ )	-- <sup>b</sup>
	SVN-AMS <sup>a</sup>	0.13 ( $f_{44}=0.16$ )	1.3 ( $f_{43}=0.051$ )	-- <sup>b</sup>
DOM	CHNS analyzer	ND	1.74	0.080
	SVN-AMS	0.77	1.7	0.081

612

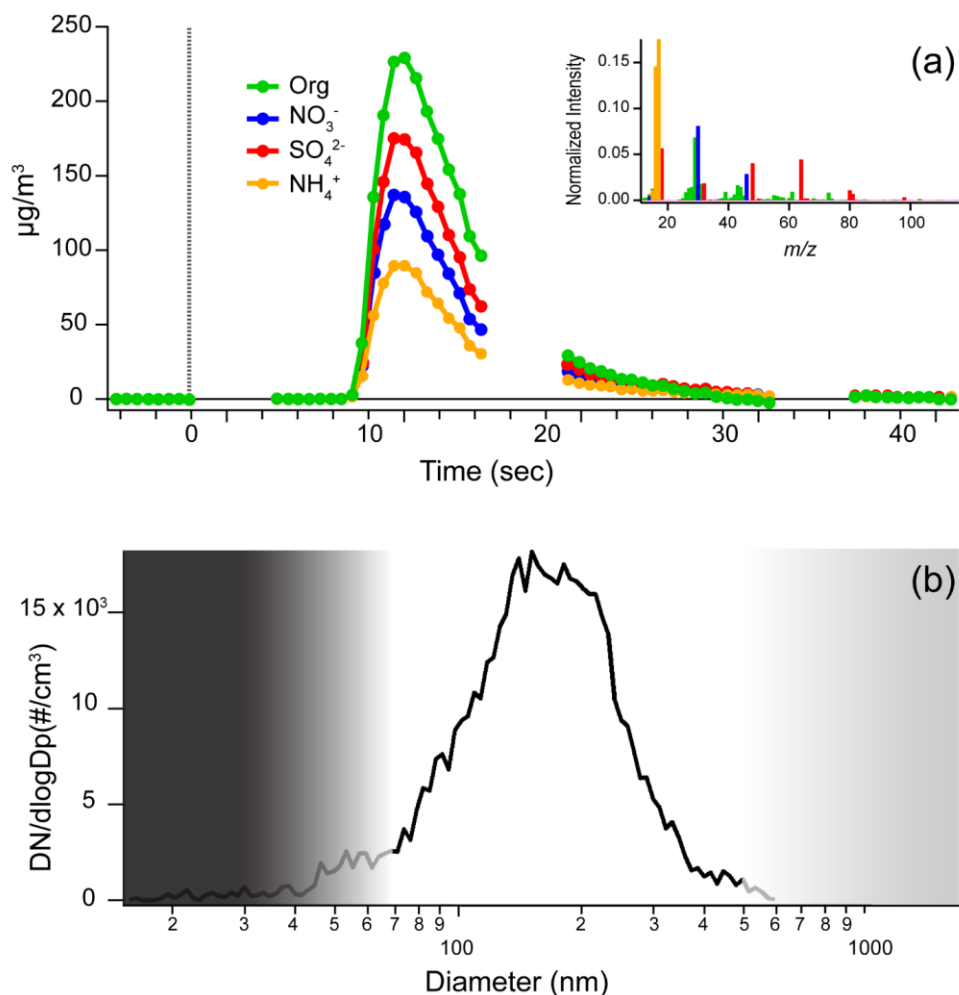
613 a. Elemental ratios are estimated from parameterizations for  $f_{44}$  and  $f_{43}$  (Aiken et al., 2008; Ng et al., 2011b).

614 b. There is no established method for determining N/C from UMR data.

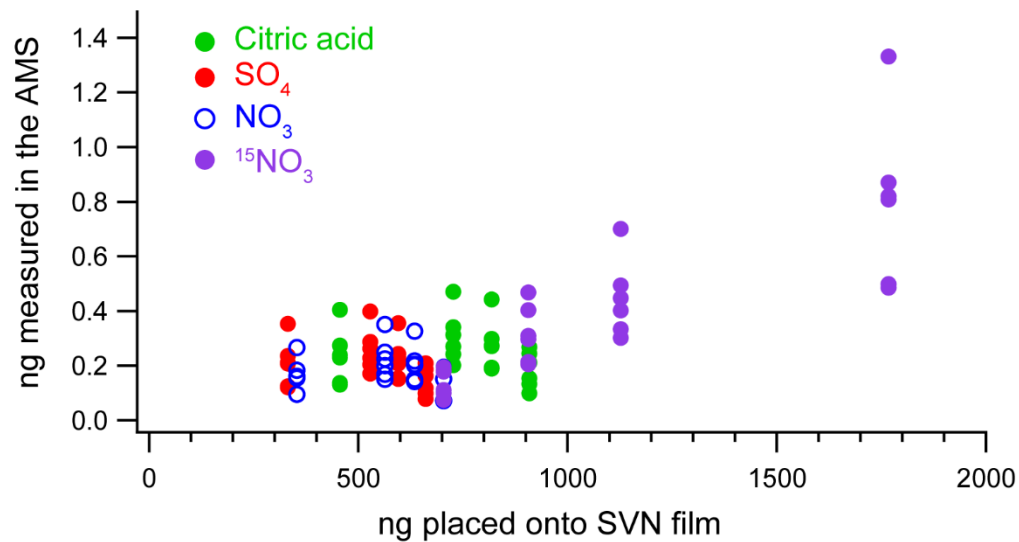


615

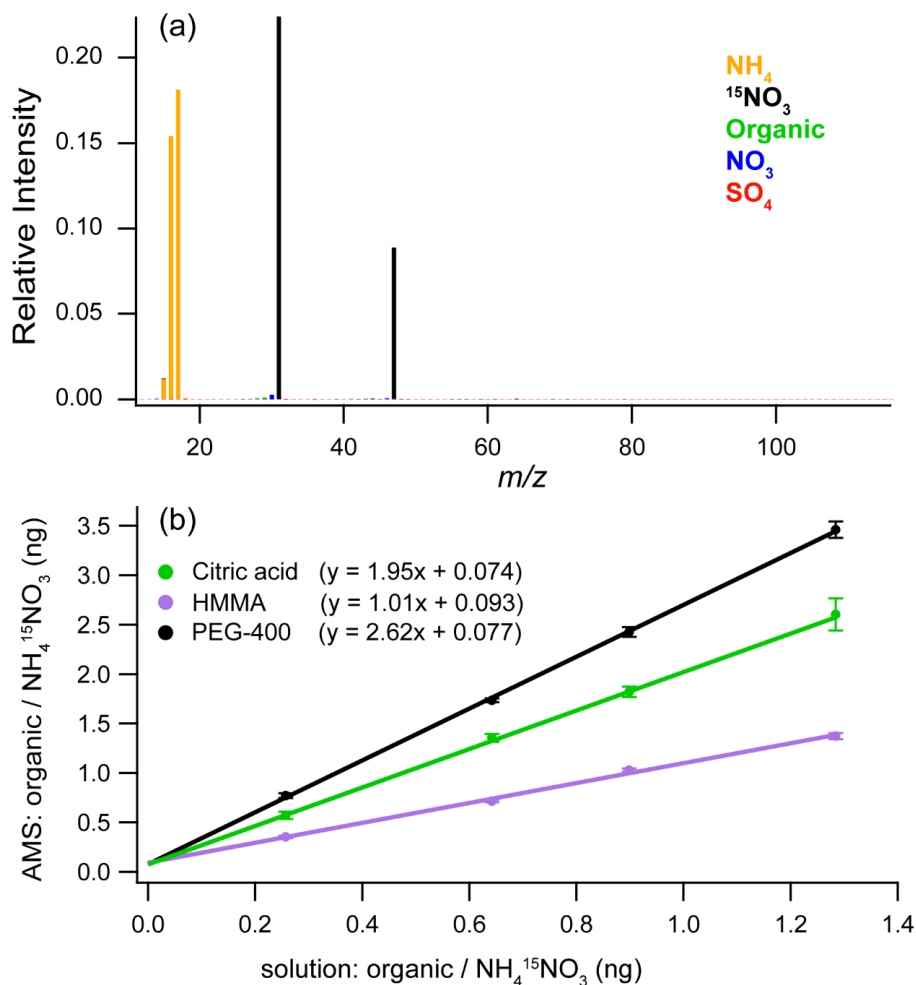
616 **Figure 1.** Schematic diagram of the small volume ultrasonic nebulizer (SVN). Samples (2-4 $\mu$ L) are loaded on the Kapton  
 617 (or Teflon) film through either the hole in which the glass tube is seated (for discrete injections) or through the pinhole (for  
 618 continuous injections). After the transducer is turned on, the aerosol is carried up through the glass tube and into the  
 619 instrument by a ~160 sccm flow of zero air or argon carrier gas. The water bath between the transducer and the Kapton  
 620 film carries ultrasonic waves up to the film and serves to cool the ultrasonic transducer.



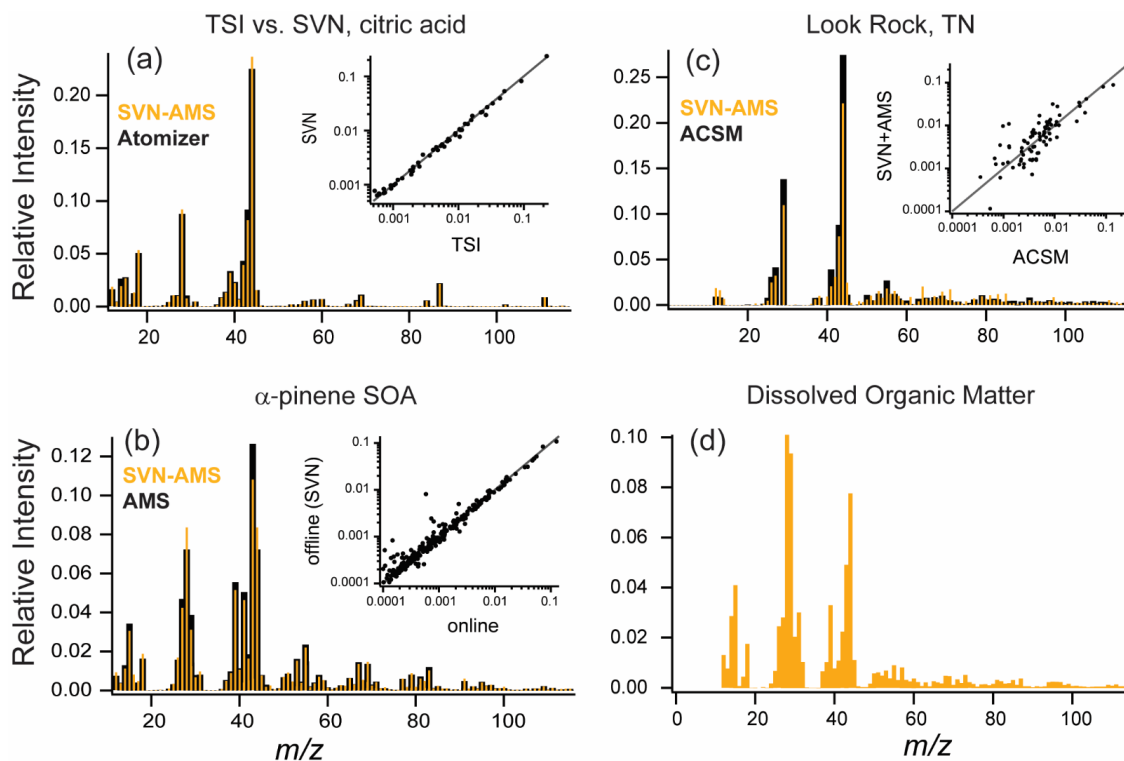
621 **Figure 2. Measurements of the composition and size of nebulized samples from the SVN. (a) Time series of aerosol**  
 622 **composition from a single 4  $\mu\text{L}$  nebulization of an aqueous solution (mannitol, ammonium nitrate, and ammonium sulfate).**  
 623 **Data were recorded using fast-mode MS for the AMS-open scans, with a mass spectrum collected every 0.5 s (filled circles).**  
 624 **The gaps in the trace correspond to closed cycles where the aerosol beam was blocked to provide a background subtraction**  
 625 **(gas-phase and instrument background) that was applied during data processing. Measured concentrations are not**  
 626 **corrected for collection efficiency (CE) in the AMS, which affects the absolute values but not the relative concentrations.**  
 627 **The inset shows the average mass spectrum acquired across the injection, normalized to total ion signal. (b) Aerosol size**  
 628 **distribution from a  $\sim 1\text{g/L}$  citric acid solution measured with an SMPS (black line). The gradient represents the**  
 629 **transmission efficiency for particles into the AMS with nearly 100% between 70-500 nm and decreased but substantial**  
 630 **transmission for spherical particles 30-70 nm and 500 nm to 2.5  $\mu\text{m}$  (Jimenez et al., 2003); thus, the smallest particles in the**  
 631 **distribution will not be efficiently detected by the AMS.**



632 **Figure 3. Mass of each component placed on the thin film vs. the mass measured by the AMS for 4 different solutions with**  
 633 **varying concentrations of citric acid, ammonium sulfate, ammonium nitrate, and the internal standard (NH<sub>4</sub><sup>15</sup>NO<sub>3</sub>), all**  
 634 **with a total solution concentration of 0.75 g/L. Each sample had 5 replicate injections, with the vertical spread in the**  
 635 **measured masses indicating substantial run-to-run variability (up to a factor of 3) between injections.**



636 **Figure 4. (a) Blank of the Kapton film using 1 g/L internal standard solution ( $^{15}\text{N}$ - ammonium nitrate). (b) Calibration**  
 637 **curves made using an internal standard for solutions with three different organic compounds: citric acid, 4-hydroxy-3-**  
 638 **methoxy-DL-mandelic acid (HMMA), and polyethylene glycol 400 (PEG-400). The error bars are  $\pm 1\sigma$  for five replicate**  
 639 **injections.**



641 **Figure 5.** Online (or TSI atomizer) (black) vs. SVN nebulizer (orange) mass spectra for (a) an aqueous solution of citric  
 642 acid at 1 g/L; (b)  $\alpha$ -pinene +  $O_3$  chamber SOA; (c) a SOAS campaign sample from Look Rock, TN with online data collected  
 643 on an ACSM. Smaller insets in a, b, and c show direct comparison of intensities for each mass spectrum on a log scale. (d)  
 644 AMS mass spectra from North Pacific Ocean dissolved organic matter (polysaccharide fraction) nebulized with the SVN  
 645 (since this sample was not from aerosol particles, no online samples are available).