

Intercomparison of lidar, aircraft, and surface ozone measurements in the San Joaquin Valley during the California Baseline Ozone Transport Study (CABOTS)

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Abstract. The California Baseline Ozone Transport Study (CABOTS) was conducted in the late spring and summer of 2016 to investigate the influence of long-range transport and stratospheric intrusions on surface ozone (O₃) concentrations in California with emphasis on the San Joaquin Valley (SJV), one of two “extreme” ozone non-attainment areas in the U.S. One of the major objectives of CABOTS was to characterize the vertical distribution of O₃ and aerosols above the SJV to aid in the identification of elevated transport layers and assess their surface impacts. To this end, the NOAA Earth System Research Laboratory (ESRL) deployed the Tunable Optical Profiler for Aerosol and oZone (TOPAZ) mobile lidar to the Visalia Municipal Airport (36.315°N, -119.392°E) in the central SJV between 27 May and 7 August 2016. Here we compare the TOPAZ ozone retrievals with co-located *in-situ* surface measurements and nearby regulatory monitors, and to airborne *in-situ* measurements from the University of California at Davis/Scientific Aviation (SciAv) Mooney and NASA Alpha Jet Atmospheric eXperiment (AJAX) research aircraft. Our analysis shows that the lidar and aircraft measurements agree, on average, to within 5 ppbv, the sum of their stated uncertainties of 3 and 2 ppbv, respectively.

1 **1 Introduction**

2 The San Joaquin Valley (SJV) of California is one of only two “extreme” ozone (O₃) non-attainment areas
3 remaining in the United States with a 2016 ozone Design Value, i.e. the metric used by the U.S. EPA to determine
4 air quality compliance that is calculated as the 3-yr average of the 4th highest measured maximum daily 8-h average
5 mixing ratio (MDA8), that is more than 20 parts-per-billion by volume (ppbv) greater than the primary National
6 Ambient Air Quality Standard (NAAQS) of 70 ppbv (<https://www3.epa.gov/airquality/greenbook/hdtdc.html>). Such
7 high O₃ concentrations are harmful to human health (U.S. Environmental Protection Agency, 2014) and impair plant
8 growth and productivity (Avnery et al., 2011a, b), adversely affecting both the \$15 billion agricultural industry in
9 the SJV and the iconic forests of the nearby Sequoia and Kings Canyon National Parks (Panek et al., 2013).

10
11 The need to better understand the causes for the high surface O₃ in the San Joaquin Valley has motivated several
12 major air quality studies over the years including the San Joaquin Valley Air Quality Study (SJVAQS) in 1990
13 (Lagarias and Sylte, 1991), the Central California Ozone Study (CCOS) in 2000, (Reynolds et al., 2010) and the
14 California Research at the Nexus of Air Quality and Climate Change (CalNex) field campaign in 2010 (Ryerson et
15 al., 2013; Brune et al., 2016). More recently, this issue was addressed by the 2016 California Baseline Ozone
16 Transport Study (CABOTS) organized and supported by the California Air Resources Board (CARB)
17 (<https://www.arb.ca.gov/research/cabots/cabots.htm>). CABOTS was designed to investigate the contributions of
18 background O₃ (Jaffe et al., 2018) and the influence of stratospheric intrusions (Lin et al., 2012a) and long-range
19 transport from Asia (Lin et al., 2012b) on surface O₃ concentrations in the SJV during late spring and summer.
20 Characterization of the vertical distribution of O₃ in the lower and middle free troposphere above the SJV and
21 upwind regions with an accuracy of at least 10%, the nominal accuracy of ECC ozonesondes in the troposphere
22 (Smit, et al., 2014), was a key objective of the campaign, and O₃ profiles were measured using three different
23 techniques (lidar, aircraft, and ozonesondes) in various parts of California. Integration of these datasets requires that
24 these measurements be intercompared (Ancellet and Ravetta, 2005; Beekmann et al., 1995; Kempfer et al., 1994;
25 Schäfer et al., 2002) and any differences between the various techniques understood and characterized. For pollution
26 studies, it is important that this validation includes the lowest 100 m, which is inaccessible to most ozone lidars
27 (Wang et al. 2017). In this paper, we compare O₃ measurements from the NOAA ESRL multi-angle Tunable Optical
28 Profiler for Aerosol and oZone (TOPAZ) lidar with *in-situ* measurements from nearby regulatory and research
29 surface monitors, and from instruments flown aboard the UC Davis/Scientific Aviation Mooney (Trousdel et al.,
30 2016) and Alpha Jet research aircraft based at NASA’s Ames Research Center (Hamill et al., 2016; Yates et al.,
31 2015). These comparisons, together with those from the multi-lidar (including TOPAZ) and ozonesonde Southern
32 California Ozone Observation Project (SCOOP) intercomparison conducted by the NASA-sponsored Tropospheric
33 Ozone Lidar Network (TOLNet) immediately after CABOTS (Leblanc et al., 2018), provide this validation.

34 35 **2 California Baseline Ozone Transport Study (CABOTS)**

36 The CABOTS field campaign was conducted between mid-May and mid-August of 2016. The primary
37 measurements (cf. **Figure 1a**) included electrochemical cell (ECC) ozonesondes (Johnson et al., 2002) launched

1 daily from Bodega Bay (38.319°N, -123.075°E, 12 m asl) (6 May-17 August) and Half Moon Bay (37.505°N, -
2 122.483°E, 9 m asl) (15 July-17 August) by the San Jose State University (SJSU), *in-situ* aircraft sampling of O₃ and
3 other compounds above central California by the University of California, Davis (UC Davis)/Scientific Aviation
4 (Trousdel et al., 2016) and the NASA Alpha Jet Atmospheric eXperiment (AJAX) (Yates et al., 2015), and ozone
5 and backscatter lidar measurements by the truck-based NOAA ESRL TOPAZ lidar system (Alvarez et al., 2011) at
6 the Visalia Municipal Airport (VMA, 36.315°N, -119.392°E, 88 m above mean sea level, asl) (27 May-18 June and
7 18 July-7 August) (**Figure 2**). Surface O₃ measurements were also made at the ozonesonde and lidar sites, and at the
8 UC Davis monitoring station at the Chews Ridge Observatory (36.306°N, -121.567°E, 1520 m asl) (Asher et al.,
9 2018) in the Santa Lucia Mountains west of Visalia, as well as the extensive networks of regulatory surface monitors
10 maintained by the California Air Resources Board and the San Joaquin Valley Unified Air Pollution Control District
11 (SJVAPCD).

12
13 The Bodega Bay and Half Moon Bay sites were located on the coast to sample the Pacific inflow, and the VMA was
14 chosen for the TOPAZ operations because of its central location in the SJV, the availability of the runway and
15 airspace for low approaches and aircraft profiles, and the presence of the co-located SJVAPCD wind profiler and
16 Radio Acoustic Sounding System (RASS) (Bao et al., 2008). The TOPAZ truck was parked on the west side of the
17 VMA between the airport runway and the heavily-trafficked multi-lane CA-99 and adjacent San Joaquin Valley
18 Railroad (SJVR) (**Figure 2**). The VMA is located about 10 km west of downtown Visalia (pop. 130,000) and lies
19 about one-third (60 km) of the way from Fresno to Bakersfield (**Figure 1a,b**). Visalia is located about 400 km from
20 Bodega Bay, and 300 km from Half Moon Bay, which limited the usefulness of comparisons between the lidar and
21 ozonesondes.

22

23 **3 Ozone Measurement Platforms**

24

25 **3.1 NOAA/ESRL TOPAZ lidar**

26 The TOPAZ differential absorption lidar (DIAL) system was originally developed for the profiling of O₃ and
27 particulate backscatter in the planetary boundary layer and lower free troposphere from NOAA Twin Otter aircraft
28 (Alvarez et al., 2011;Langford et al., 2011;Senff et al., 2010;Langford et al., 2012;Langford et al., 2010). The lidar
29 was reconfigured for mobile ground-based measurements in 2012 and deployed in this configuration to several field
30 campaigns including the 2013 Las Vegas Ozone Study (LVOS) (Langford et al., 2015) prior to CABOTS. The lidar
31 is installed in the back of a medium box truck (cf. **Figure 2**) equipped with a commercial UV absorption monitor for
32 *in-situ* O₃ measurements (2B Technologies Model 205) that samples air 5 m above the surface and an Airmar
33 150WX weather station to measure temperature, pressure, relative humidity, and wind speed and direction. The 2B
34 Model 205 has been approved by the EPA as a Federal Equivalent Method (FEM) for surface O₃ monitoring and has
35 a nominal (1 σ) precision and accuracy that is the greater of 1 ppbv or 2% for 10-s averages. Modified versions of
36 the same instrument were flown on both the Scientific Aviation Mooney and NASA Alpha Jet. Comparisons

1 between the NOAA 2B at the VMA and a mobile calibration source operated by CARB revealed a 3% low bias in
2 the recorded 2B measurements that has been corrected in the data used here.

3
4 The eye safe TOPAZ lidar is built around a low pulse energy ($\sim 100 \mu\text{J}$), high repetition rate (1 kHz) quadrupled
5 Nd:YLF pumped Ce:LiCAF laser that is re-tuned between each pulse to generate light at three different wavelengths
6 from 286 to 294 nm with an effective repetition rate of 333 Hz for each wavelength (Alvarez et al., 2011). The laser
7 pulses are transmitted and the lidar return signals collected by a coaxial transmitter/receiver equipped with a
8 commercial (Licel) photomultiplier-based dual analog/photon counting system. This hybrid data acquisition system
9 was installed in 2016 and replaced the original fast analog data acquisition system that was optimized for aircraft
10 operations (Alvarez et al., 2011; Wang et al., 2017). This modification increased the maximum useful range to ~ 6 km
11 during the day and to more than 8 km at night, depending on the laser power, atmospheric extinction, and solar
12 background light.

13
14 The truck-mounted version of TOPAZ incorporates a large scannable turning mirror above the vertically pointing
15 transmitter/receiver to allow profile measurements at different slant angles. These slant profiles can be combined to
16 create vertical profiles that start much closer to the ground (25-30 m) than conventional vertically staring lidar
17 systems (Proffitt and Langford, 1997). During CABOTS, the scanning mirror was moved sequentially between
18 elevation angles of 90, 20, 6, and 2° with a 225-s averaging time at 90° and 75-s averaging times at the other 3
19 angles. The cycle was repeated approximately every 8 minutes and the vertical projections combined to create a
20 single vertical profile starting at 27.5 ± 5 m above ground level (agl). This approach assumes a fair degree of
21 horizontal homogeneity and the lidar slant paths were oriented parallel to the VMA runway (135°) over open
22 farmland to avoid populated neighborhoods and minimize the effects of NO emissions from the often heavy traffic
23 on CA-99 (cf. **Figure 2**), which could locally titrate ozone and create strong horizontal concentration gradients near
24 the surface.

25
26 The O_3 profiles shown here were retrieved using two wavelengths (~ 287 and 294 nm) with 30-m range gates and a
27 smoothing filter that increased from 270 m wide at the minimum range (815 ± 15 m) to 1400 m wide at the maximum
28 range (8 km). The effective vertical resolution increased from ~ 10 m near the surface to ~ 150 m above 500 m agl
29 and 900 m at 6 km. Profiles of the backscatter from aerosols, smoke, and dust were retrieved with a constant 7.5 m
30 resolution at 294 nm. The ozone profiles were computed using the O_3 absorption cross-sections from Malicet *et al.*
31 (1995) and an iterative technique to correct for differential aerosol backscatter and extinction that assumes a
32 backscatter-to-extinction ratio of 40 and fixed Ångstrom coefficients of 0 for backscatter and -0.5 for extinction
33 (Alvarez et al., 2011). These values offer a good compromise for a wide variety of particulate types (Völger et al.,
34 1996). The actual aerosol composition in the SJV was not measured during CABOTS, but measurements during the
35 2010 Carbonaceous Aerosols and Radiative Effects Study (CARES) typically found a mix of organics, sulfate,
36 nitrate, ammonium, and soil dust in the northern part of the valley (Zaveri et al., 2012). Smoke from the Soberanes
37 Fire near Big Sur dominated the aerosol mix in the SJV during the second IOP. We varied the aerosol backscatter

1 Angstrom coefficient between -1 and 1 and the aerosol extinction Angstrom coefficient between 0 and -1 for a
2 “worst case scenario” of a thin smoke layer with very high aerosol backscatter embedded in an otherwise clean
3 atmosphere to estimate the error in the ozone retrieval introduced by using these fixed parameters. The sharp aerosol
4 gradients at the smoke layer edges tend to magnify errors in the ozone retrieval if the aerosol correction is not
5 properly implemented. Temperature and pressure profiles interpolated from the 3-h National Centers for
6 Environmental Prediction (NCEP) North American Regional Reanalysis (NARR) using the grid point closest to the
7 TOPAZ lidar location were used to account for the temperature dependence of the O₃ cross-sections and to convert
8 O₃ number densities to mixing ratios. The total uncertainties in the 8-min ozone retrievals in the absence of strong
9 aerosol gradients are estimated to increase from ±3 ppbv below 4 km to ±10 ppbv at the top of the profile. When
10 strong backscatter gradients are present, the O₃ uncertainty can potentially increase by another ±3 ppbv.

11

12 **3.2 UC Davis/Scientific Aviation Mooney**

13 The University of California at Davis and Scientific Aviation, Inc. (<http://www.scientificaviation.com>), conducted a
14 series of research flights above the SJV during the summer of 2016 using a Scientific Aviation single-engine
15 Mooney TLS or Ovation aircraft as part of the CARB-supported Residual Layer Ozone Study (RLO)
16 (<https://www.arb.ca.gov/research/apr/past/14-308.pdf>). Several of these flights overlapped with the TOPAZ
17 operations during CABOTS, as did some of the 12 additional flights (EPA) funded by the U.S. EPA and the Bay
18 Area Air Quality Management District (BAAQMD). The Mooney carried a 2B Technologies Model 205 O₃
19 monitor, an Eco Physics Model CLD 88 (NO) with a photolytic converter to measure NO and NO₂, and a Picarro
20 2301f Cavity Ring-Down Spectrometer (CRDS) to measure CO₂, CH₄, and H₂O (Trousdel et al., 2016). The 2B
21 model 205 was used with the minimum integration time of 2 s, which corresponds to a mean distance of 150 m at
22 the typical level flight speed (the data stream was sampled at 1-s intervals). As noted above, the 2B has a nominal
23 accuracy of 2% for concentrations above 5 ppbv, and a precision of 2% for concentrations above 5 ppbv if 10-s
24 averages are used. If the limiting noise is randomly distributed, this implies a precision of 5 ppbv for 2-s averages.
25 Calibrations of the Scientific Aviation 2B using an external ozone source (2B, Model 306) found the instrument to
26 have offsets and slopes less than 1.5 ppb and within 4% of unity, respectively.

27

28 **3.3 NASA Alpha Jet Atmospheric eXperiment (AJAX)**

29 The NASA Ames Alpha Jet Atmospheric eXperiment (AJAX) (Hamill et al., 2016) sampled O₃ and other
30 tropospheric constituents above California during CABOTS using a two-person jet based at Moffett Field, CA (MF,
31 37.415° N, -122.050° E). The Alpha Jet carried an external wing pod with a modified commercial UV absorption
32 monitor (2B Technologies Inc., model 205) to measure O₃ (Ryoo et al., 2017; Yates et al., 2015; Yates et al., 2013)
33 and a (Picarro model 2301-m) cavity ringdown analyzer to measure CO₂, CH₄, and H₂O (Tanaka et al., 2016). A
34 second wing pod carried a non-resonant laser-induced fluorescence instrument to measure formaldehyde (CH₂O)
35 (St. Clair et al., 2017). The pod mounting kept the residence times of the sample inlets to less than 2 s. The aircraft is
36 also equipped with GPS and inertial navigation systems to provide altitude and position information, and the NASA

1 Ames-developed Meteorological Measurement Systems (MMS) to provide highly accurate pressure, temperature,
2 and 3-D wind data. The 2B O₃ data, recorded every 2 s, are averaged over 10 s to increase the signal-to-noise ratio.
3 Ozone calibrations were performed before/after each flight using an external ozone source (2B Technologies Inc.,
4 model 306 referenced to the NIST scale, certified annually). Raw flight O₃ data were corrected using the linearity
5 correction factor and zero offset from the calibration closest in time to the flight. Overall accuracy of the O₃
6 instrument is determined to be 3 ppbv or better at 10-s resolution, with uncertainty improving at lower altitudes, as
7 determined from pressure chamber tests; see Yates et al., (2013) for a more detailed error analysis.

8 9 **4 Results and Comparisons**

10 The TOPAZ measurements were conducted over two 3-week intensive operating periods (IOPs) in the late spring
11 (27 May to 18 June) and summer (18 July to 7 August) of 2016. A total of 440 hours of lidar data were recorded
12 during the first (1654 profiles over 22 days) and second (1686 profiles over 21 days) IOPs with an average of more
13 than 10 hours of nearly continuous measurements per day. The skies above Visalia were mostly cloud free during
14 the study, with only a few profiles truncated by high clouds during IOP1. However, during IOP2 the SJV was
15 fumigated by smoke from the Soberanes Fire that started on 22 July about 200 km west of Visalia near Big Sur.

16 17 **4.1 Comparisons between lidar and surface measurements**

18 The NOAA 2B ozone monitor operated continuously at the VMA throughout the TOPAZ deployment with the
19 system response checked during each IOP by an external mobile calibration source operated by CARB. These
20 calibration checks revealed a 3% low bias in the NOAA 2B instrument that has been corrected in the data shown
21 here. **Figure 3** plots time series (Pacific Daylight Time, PDT, or UT-7 h) of the 1-min averaged *in-situ* surface
22 mixing ratios (gray dots) measured 5 m above the ground from each IOP together with the TOPAZ mixing ratios
23 retrieved from a height of 27.5±5 m (black line) and a range of 815±15 m along the slant path above the agricultural
24 fields to the southeast (cf. **Figure 2**). **Figure 4a** is an enlarged view of the VMA surface measurements (gray line)
25 from 9-13 June together with the mixing ratios from the 27.5 m TOPAZ measurements (filled black circles). Also
26 plotted are the 1-h average ozone mixing ratios measured 6.7 m agl by the CARB regulatory API/Teledyne 400
27 monitor located on N. Church Street in Visalia (102 m asl) about 10 km to the east of VMA (solid black line), and
28 measured 5 m agl by the SJVAPCD API/Teledyne 400 monitor in Hanford (82 m asl) about 22 km to the west of
29 VMA (dotted black line). The four sets of measurements agreed fairly well during the day but diverged markedly at
30 night and in the early morning when O₃ was removed by surface deposition and titration by NO_x within the surface
31 layer. The losses were greatest at the VMA monitor which was located in the TOPAZ truck next to the heavily-
32 trafficked CA-99 and SJVR railroad line. Titration by NO was undoubtedly much greater here, but there were no
33 NO_x measurements available to confirm this hypothesis. Much smaller losses were measured by the rural Hanford
34 monitor and intermediate losses were measured by the Visalia monitor which is located on a downtown rooftop. A
35 scatter plot of all of the coincident TOPAZ and *in-situ* measurements from CABOTS (**Figure 4b**, filled gray circles)
36 shows that the *in-situ* concentrations measured at VMA were often much smaller than the concentrations measured
37 815±15 m away by the lidar, and even titrated to zero under some conditions. The data converge (filled black

1 circles) when the comparison is restricted to conditions when the two measurements are expected to sample a
2 common airmass, i.e. during the day after the nocturnal inversion has dissipated (0900 to 1830 PDT) and the winds
3 were southeasterly (125 to 145°) and greater than 2.5 m s⁻¹. The results of Orthogonal Distance Regression (ODR)
4 fits of these data are shown both in the figure and in **Table 1**. We use ODR fits, which assume that both variables
5 can have uncertainties, for our analyses instead of simple linear regressions which assume that all of the
6 uncertainties lie in the dependent variable. Fits of the filtered data give a slope of 1.00±0.03 and an intercept of -
7 2.6±1.5 ppbv where the errors represent the 95% confidence limits of the ODR fits.

8
9 **Figure 5** compares the 27.5 m TOPAZ O₃ measurements to the regulatory O₃ surface measurements from the
10 monitors at Visalia (8.5 km) and Hanford (24 km) described above, and from the more distant SJVAPCD monitors
11 at Parlier (34 km) and Porterville (43 km). The TOPAZ mixing ratios were slightly higher than those at Visalia and
12 Hanford, but lower than those at Parlier and Porterville, which are closer to the Sierra foothills and measure some of
13 the highest O₃ concentrations found in the SJVAB. The degree of correlation decreased with distance as expected,
14 yet remained quite good more than 40 km from the VMA at Porterville. This suggests that the O₃ measurements
15 acquired at the VMA during CABOTS can be considered representative of the central San Joaquin Valley.

16 17 **4.2 Comparisons between lidar and aircraft measurements**

18 Comparisons between the ground-based lidar and aircraft measurements are subject to much larger uncertainties
19 arising from spatial and temporal sampling differences compared to the comparison with nearby surface monitors.
20 During CABOTS, the fixed wing aircraft conducted both low approaches above the VMA runway (cf. **Figure 2**) and
21 spiral profiles around the airport, but never directly sampled the vertical column probed by the lidar. The
22 comparisons were also conducted as brief elements of multi-hour sampling flights with other objectives, and time
23 constraints and air traffic considerations sometimes contributed to the spatial and temporal mismatches. The piston-
24 engine Mooney took about 25 minutes to execute an ascending profile from the surface to 3 km, while the Alpha Jet
25 took about 9 minutes (similar to the 8-min TOPAZ integration time) to conduct a descending profile from 3 km to
26 the surface. Spatial mismatches were also created by the vertically smoothing of the DIAL retrieval, which can both
27 smooth and displace sharp vertical concentration gradients seen by the aircraft. Similar considerations apply to
28 comparisons between lidars and ozonesondes since balloons have a finite rise time and can be carried many
29 kilometers downwind from the launch site (Leblanc et al. 2018). Despite these caveats, we show that the lidar and
30 aircraft measurements usually agreed to within ±10%, the nominal accuracy of ECC ozonesondes in the troposphere
31 (Smit, et al., 2014), which is the generally accepted reference standard for ozone profile measurements.

32 33 **4.2.1 UC Davis/Scientific Aviation Mooney**

34 The RLO flights were executed as a series of 2 to 3-day deployments with as many as 4 flights per day lasting 2 to 3
35 hours each between Fresno and Bakersfield. Two of these deployments, RLO2 (2-4 June), and RLO4 (24-26 July),
36 overlapped with the first and second TOPAZ IOPs, respectively, and included low approaches at VMA on most of
37 the flights with spiral profiles near VMA on several. Both deployments occurred as warm temperatures (>40°C) and

1 weak anticyclonic winds associated with synoptic high-pressure systems resulted in the buildup of surface ozone
2 across the South Coast and San Joaquin Valley Air Basins. The highest measured MDA8 O₃ in the SJVAB during
3 the first IOP was recorded on 4 June at Clovis (91 ppbv), which lies about 65 km northwest of VMA (cf. **Figure 1b**).
4 The highest reported MDA8 O₃ during the second IOP (and the year) was recorded on 27 July at Parlier (101 ppbv),
5 which lies midway between Clovis and the VMA. The monitors at Visalia and Hanford reported MDA8
6 concentrations of 72 and 88 ppbv, respectively, on 4 June, and 83 and 85 ppbv on 27 July. **Figure 3** shows that the
7 highest O₃ mixing ratios measured by the VMA surface monitor and TOPAZ (27.5 m agl) were also recorded on
8 these two days.

9
10 The flight tracks from all of the Mooney sorties during the RLO2 and RLO4 deployments are plotted in **Figure 6a**.
11 FLT29 (RLO4) was a transit flight from the Scientific Aviation home base near Sacramento to Fresno. The
12 remaining RLO flights were between Fresno and Bakersfield as noted above. The two EPA deployments (27-29 July
13 and 4-6 August) were of longer duration than the RLO flights with morning and afternoon sorties that placed more
14 emphasis on cross-valley measurements and transects to the coast (**Figure 6b**) including profiles above the South
15 Bay (EPA1) and Chews Ridge (EPA2). The afternoon flights during both series included legs to Visalia.

16
17 **Figure 7** shows the sections of the RLO and EPA flight tracks that passed within 5 km of TOPAZ (dashed black
18 circles). Most of these flights included low (<10 m) passes along the VMA runway that approached to within ~350
19 m horizontally of the TOPAZ truck and within 1000 m of the center of the 27.5 m agl TOPAZ slant path
20 measurements (cf. **Figure 2**). **Figures 8a-8d** show time series of the 27.5 m TOPAZ and 5 m *in-situ* measurements
21 during all of the RLO and EPA low approaches together with the ozone measured by the aircraft between the surface
22 and 25 m agl. All of the aircraft measurements lie within 10% of the O₃ retrieved by TOPAZ with the exception of
23 the much higher values (>100 ppbv) measured by the Mooney around 1400 PDT on 3 June (**Figure 8a**, see below).
24 The scatter plots in **Figures 8e** and **8f** show that the aircraft also measured much higher concentrations than the *in-*
25 *situ* surface monitor during the night and early morning, in agreement with the lidar measurements in **Figure 4**. The
26 differences were smaller on 27 July than on 3 June, and also less pronounced than those in **Figure 4**. Closer
27 agreement between the aircraft and surface measurements might be expected since some of the aircraft
28 measurements were made within 200 m of the lidar truck (cf. **Figure 2**). The dark blue points show that the low bias
29 in the surface measurements decreased during the day after the surface inversion had dissipated (there were too few
30 measurements to effectively filter them by windspeed or direction). The mean ODR fit parameters based on the
31 measurements from both RLO2 and RLO4 listed in **Table 1** are very similar to those found for the lidar which
32 suggests that the filtered surface measurements still have low bias that could be either instrumental or sampling
33 related.

34
35 **Figure 9** compares the aircraft and lidar O₃ measurements made during 5 of the ascending profiles conducted by the
36 Mooney near the VMA. FLT19 was conducted in the early afternoon of 3 June, and FLT33, FLT35, FLT36, and
37 FLT37 were conducted over the 24-hour period beginning just after local midnight on 25 July. The four consecutive

1 TOPAZ profiles acquired during the time required for the Mooney to reach the top of each profile (~15-30 minutes
2 at a climb rate of $\sim 2.2 \text{ m s}^{-1}$) are plotted in each panel. The gray envelopes show the lidar mean profile $\pm 10\%$. The
3 differences between consecutive profiles reflect the combined effects of atmospheric variability and the precision of
4 the lidar measurements.

5
6 Overall, the agreement between the TOPAZ and Mooney profiles in **Figure 9** is within $\pm 10\%$, but there are some
7 notable discrepancies. Most of these arise from the coarser vertical resolution of the lidar retrievals, which smooth
8 out abrupt concentration changes such as those seen at the top of the boundary layer ($\sim 0.8 \text{ km agl}$) in **Figure 9a**, and
9 between 2 and 3 km in **Figure 9e** where several narrower layers are smoothed into one broad layer in the lidar
10 profile. **Figure 9e** also shows that the agreement between the lidar and aircraft measurements is better at low
11 altitudes where the addition of the slant path measurements significantly improves the effective vertical resolution of
12 the lidar. Fine-scale variability in O_3 also contributes to some of the observed differences, particularly on 3 June
13 where the aircraft-measured O_3 concentrations varied by as much 25 ppbv during the low approach over the VMA
14 runway. This unusually large variability is also seen in the large and rapid changes in the lidar measurements near
15 the top of the boundary layer (**Figure 9a**) and challenges the assumptions about horizontal homogeneity used in the
16 calculation of the TOPAZ vertical profiles near the surface.

17
18 The lidar profiles from 26 July (**Figure 9e**) also show large profile-to-profile changes in the narrow high O_3 layer
19 lying just above the top of the nocturnal boundary layer ($\sim 0.3 \text{ km asl}$). The 25 and 26 July measurements (**Figures**
20 **9b-9e**) were made several days after the Soberanes Fire started and the low altitude “layer” near 400 m in **Figure 9e**
21 is actually a short-lived puff of smoke and elevated O_3 from the fire. This is more obvious in the expanded view of
22 the profiles shown in **Figure 10a**. Only two of the four lidar profiles from **Figure 9e** are plotted: the first profile
23 coinciding with the aircraft measurements (solid trace, $\pm 10\%$) and the profile acquired 16-24 minutes later when the
24 puff had mostly disappeared (dashed trace). The corresponding lidar backscatter measurements are plotted in **Figure**
25 **10b**, and **Figure 10c** shows the NO_2 and H_2O profiles measured by the aircraft. The backscatter measurements show
26 that the TOPAZ retrievals are unaffected by strong backscatter gradients, which can create second-derivative like
27 inflection points in the DIAL O_3 profiles (Kovalev and McElroy, 1994). The absence of a corresponding structure in
28 the aircraft NO_2 and H_2O profiles confirms that the high O_3 layer seen in the lidar and aircraft measurements was not
29 an artifact caused by interferences from these species, which weakly absorb between 280 and 300 nm (Proffitt and
30 Langford, 1997).

31 32 **4.2.2 NASA Alpha Jet Atmospheric eXperiment (AJAX)**

33 AJAX conducted 4 research flights over the SJV while TOPAZ was operational, with 2 additional flights (21 June
34 and 7 July) between the two IOPs. The Alpha Jet executed descending spiral profiles from 4 to 5 km down to the
35 surface that ended in low approaches on three of these flights: AJX190 on 3 June, AJX191 on 15 June, and AJX195
36 on 21 July. The aircraft also conducted a very low approach ($\sim 5 \text{ m}$) at VMA on 28 July (AJX196) but did not
37 execute a full profile. These low approach measurements are represented by the filled yellow circles in Figures 8a

1 and 8c. The first and last flights (AJX190 and AJX196) coincided with the high ozone episodes mentioned earlier
2 and the third flight (AJX195) also occurred during a period of high pressure. The second flight (AJX191) was
3 conducted as a deep closed low moved into the Pacific Northwest, however, bringing unseasonably cool
4 temperatures (26 °C) and strong surface winds to the SJV. This cyclonic system advected a large Asian pollution
5 plume across the valley in the middle troposphere, but surface ozone remained low with the peak MDA8 O₃
6 concentration in the SJVAB only reaching 59 ppbv at the Sequoia-Kings Canyon monitor.

7
8 **Figures 11 and 12** are similar to **Figures 6 and 7**, but instead show the AJAX flight tracks. The first AJAX flight
9 (AJX190) on 3 June during IOP1 overlapped with the UC Davis/Scientific Aviation RLO2 deployment. AJAX191
10 took place about two weeks later in IOP1, and AJX195 occurred several days prior to the RLO4 deployment in
11 IOP2. AJAX also executed profiles (not shown here) above and upwind of Chews Ridge on AJX190 and AJAX191
12 and near Bodega Bay on AJX191 and 195 and sampled the Soberanes Fire plume on AJX196.

13
14 **Figure 13** displays coincident AJAX and TOPAZ profiles in plots similar to those shown for the Mooney in **Figure**
15 **9**, but with an extended vertical axis to reflect the higher range of these profiles. The points in **Figure 13** are sparser
16 than those in **Figure 9** in part because of the 10-s averaging time, and in part because the Alpha Jet executed its
17 descending profiles with an airspeed of about 110 m s⁻¹ compared to about 60 m s⁻¹ for the ascending Mooney
18 profiles.

19
20 The agreement between the Alpha Jet and TOPAZ measurements is within ±10% on all 3 days except for 3 June
21 (**Figure 13a**) when the measured aircraft and retrieved lidar concentrations differ by as much as 12 ppbv (20%) at
22 2.5 km asl and 20 ppbv (~50%) at 5.2 km asl. The disparities between the inbound and outbound measurements in
23 **Figure 13a** show that the Alpha Jet encountered strong horizontal gradients below 800 m in the boundary layer
24 when it arrived at the VMA about 3 hours after the Mooney found similar horizontal variability (cf. **Figures 8a** and
25 **9a**). The Google Earth map and latitude-altitude and longitude-altitude plots in **Figure 14** better illustrate the extent
26 of the horizontal variability in the boundary layer. These figures also show weaker horizontal gradients above 3 km
27 where the disagreement between the lidar and aircraft is most pronounced.

28 29 **5 Discussion**

30 The results of the different O₃ comparisons are summarized in **Table 1**. As was noted above, comparisons between
31 the lidar and aircraft profiles are subject to uncertainties arising from sampling differences introduced by the
32 intrinsic vertical smoothing of the lidar retrievals and horizontal displacements between the aircraft and lidar. The
33 potential impact of horizontal displacements on the comparisons when the O₃ spatial variability is large is illustrated
34 by **Figure 14**, and a good example of the differences created by the lidar smoothing is seen near the top of the
35 boundary layer around 0.8 km in **Figure 9a**. These uncertainties can be reduced by averaging the measurements to
36 be compared over larger volumes. **Figure 15** compares the lidar and aircraft measurements from the profiles plotted
37 in **Figures 9 and 13**, and from several other RLO and EPA flights not shown, with each individual profile averaged

1 over 1 km segments (0 to 1 km, 1 to 2 km, etc.). This averaging decreases the influence of O₃ spatial variability, and
2 also reduces the statistical uncertainties in both the lidar retrievals and aircraft measurements, with the effective
3 temporal averaging of the AJAX and SciAv measurements increasing to about 2 and 4 minutes, respectively. Each
4 point in the scatter plots of **Figure 15a** and **15b** represents the mean mixing ratio from one of these 1 km segments,
5 with the error bars showing the standard deviation of the mean. The intercepts and slopes derived from orthogonal
6 distance regressions of both datasets overlap with zero and unity, respectively, within the 95% confidence limits of
7 the ODR fits. The lower panels (**Figures 15c** and **15d**) plot the same data as differences which show that the TOPAZ
8 and SciAv measurements (**Figure 15c**) agree to within 1 ppbv on average, and the TOPAZ and AJAX
9 measurements (**Figure 15d**) to within 4.2 ppbv. Neither plot shows evidence of a systematic altitude dependence in
10 the differences.

11
12 Both lidar/aircraft comparisons are limited by the small number of common measurements with only 3 profiles
13 available for the AJAX comparisons. The SciAv comparisons include data from 7 flights, but only the 5 profiles shown
14 in **Figure 9** extend above 2 km and only 3 of those reach 3 km. These limited datasets make the comparisons more
15 sensitive to the influence of individual points. For example, the point surrounded by the dashed circle in **Figure 15d**
16 includes the measurements from within the dashed oval in **Figure 13b** where the lidar retrieval is clearly smoothing
17 out vertical gradient compared to the aircraft measurements. If this measurement point is excluded, the mean TOPAZ-
18 AJAX difference decreases to 3.9 ± 2.6 . In either case, the differences between the TOPAZ lidar retrievals and the in-
19 situ surface and aircraft measurements lie within the combined uncertainties of the different measurements and well
20 within the 10% accuracy standard set by the ECC ozonesonde.

21 22 **6 Summary and Conclusions**

23 The lidar, aircraft, and ozonesonde profiles acquired during the 2016 CABOTS field campaign provide an
24 unprecedented look at the vertical distribution of lower tropospheric O₃ above California during late spring and
25 summer. The good agreement between the low elevation TOPAZ measurements and the collocated and regional
26 (<45 km) surface monitors suggests that the measurements made at the VMA during CABOTS can be considered
27 representative of the central San Joaquin Valley. Comparisons between the NOAA TOPAZ lidar profiles and the
28 surface and aircraft measurements agree within the stated uncertainties, and we conclude that all of these O₃
29 measurements may be used with confidence.

30
31 The coordinated lidar and aircraft sampling of O₃ above the central San Joaquin Valley during CABOTS also
32 illustrates the synergy between the two types of measurements. Lidar can provide long time series of the O₃ (and
33 backscatter) vertical distributions above a fixed location while the aircraft can place the lidar measurements within a
34 larger spatial context and measure other important parameters. This synergy is illustrated by the two time-height
35 curtain plots displayed in **Figure 16**. **Figure 16a** shows the continuous TOPAZ measurements from a 14-hour time
36 span on 25-26 July with the data from SciAv FLT 35, 36, and 37 superimposed. The aircraft measurements made

1 within 5 km of VMA are highlighted by colored squares outlined in white. **Figure 16b** is similar, but shows 10-
2 hours of continuous TOPAZ measurements from 15 June with the AJAX measurements (AJX191) superimposed.

3
4 The CABOTS ozonesondes were launched too far away (>300 km) from the VMA to allow quantitative
5 comparisons with the lidar. However, TOPAZ was relocated to the NASA Jet Propulsion Laboratory (JPL) Table
6 Mountain Facility (TMF) in the San Gabriel Mountains immediately after CABOTS for the Southern California
7 Ozone Observation Project (SCOOP), a multiple lidar and ozonesonde intercomparison organized by the NASA-
8 sponsored Tropospheric Ozone Lidar Network or TOLNet (<https://www-air.larc.nasa.gov/missions/TOLNet/>) at the
9 NASA Jet Propulsion Laboratory (JPL) Table Mountain Facility (TMF) (Leblanc et al., 2018). The results from the
10 SCOOP intercomparison and those presented here complete the inter-validation of the CABOTS lidar, aircraft, and
11 ozonesonde profile measurements.

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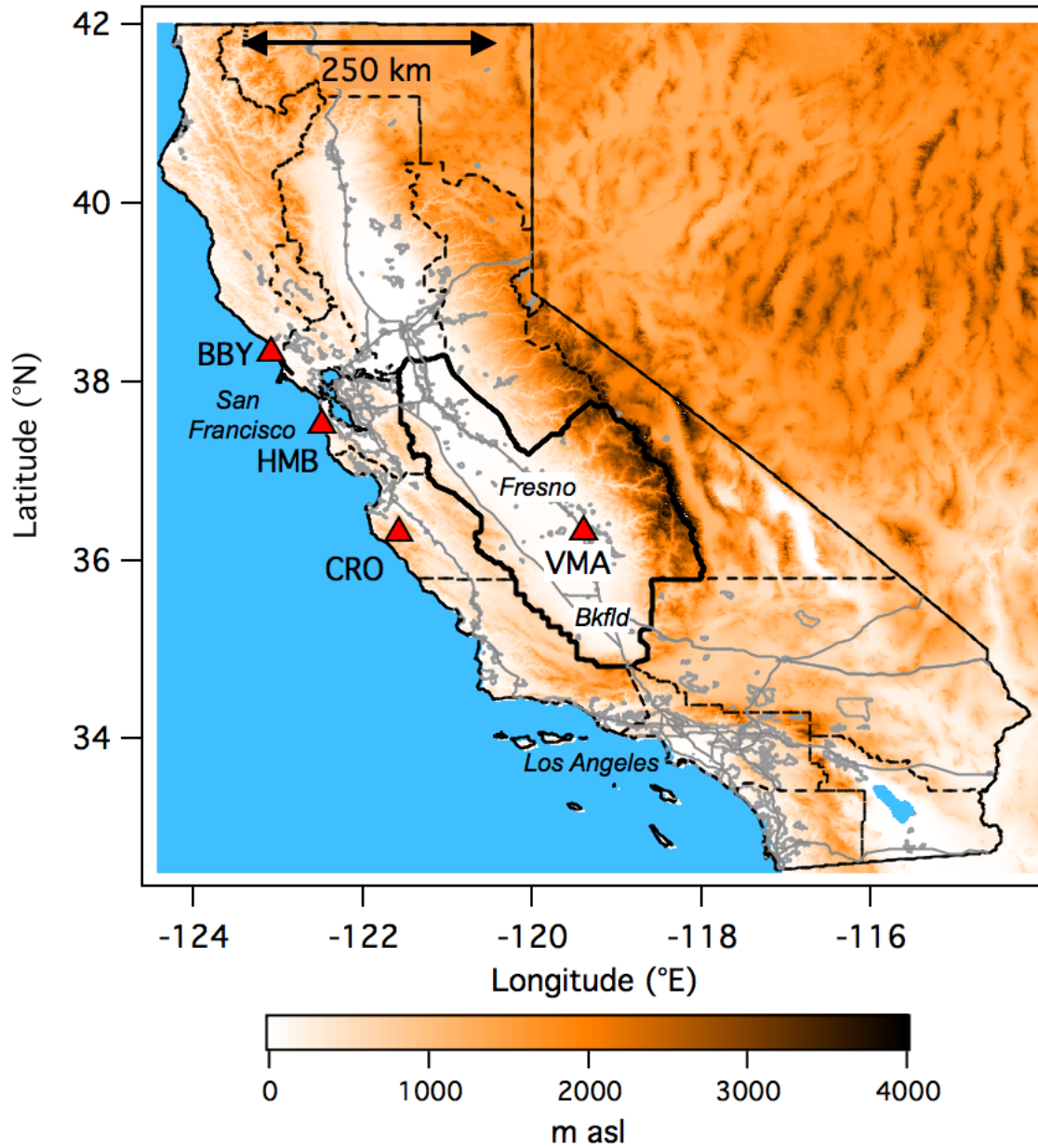
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Table 1. Summary of the lidar, surface, and aircraft comparisons

A	B	Ratio $\pm 1\sigma$ (A/B)	Diff. $\pm 1\sigma$ (A-B)	Slope* (A vs B)	Int.* (A vs B)
TOPAZ	VMA	1.06 \pm 0.08	2.9 \pm 3.7 ppbv	1.00 \pm 0.03	-2.6 \pm 1.5 ppbv
SciAv	VMA	1.07 \pm 0.10	5.0 \pm 5.0 ppbv	1.01 \pm 0.01	-4.5 \pm 1.1 ppbv
TOPAZ	SciAv	1.01 \pm 0.04	0.8 \pm 2.8 ppbv	1.00 \pm 0.13	1.0 \pm 9.0 ppbv
TOPAZ	AJAX	1.08 \pm 0.06	4.2 \pm 0.8 ppbv	1.07 \pm 0.13	1.8 \pm 3.4 ppbv

*from Orthogonal Distance Regression (ODR) fits. Uncertainties are 95% confidence limits.

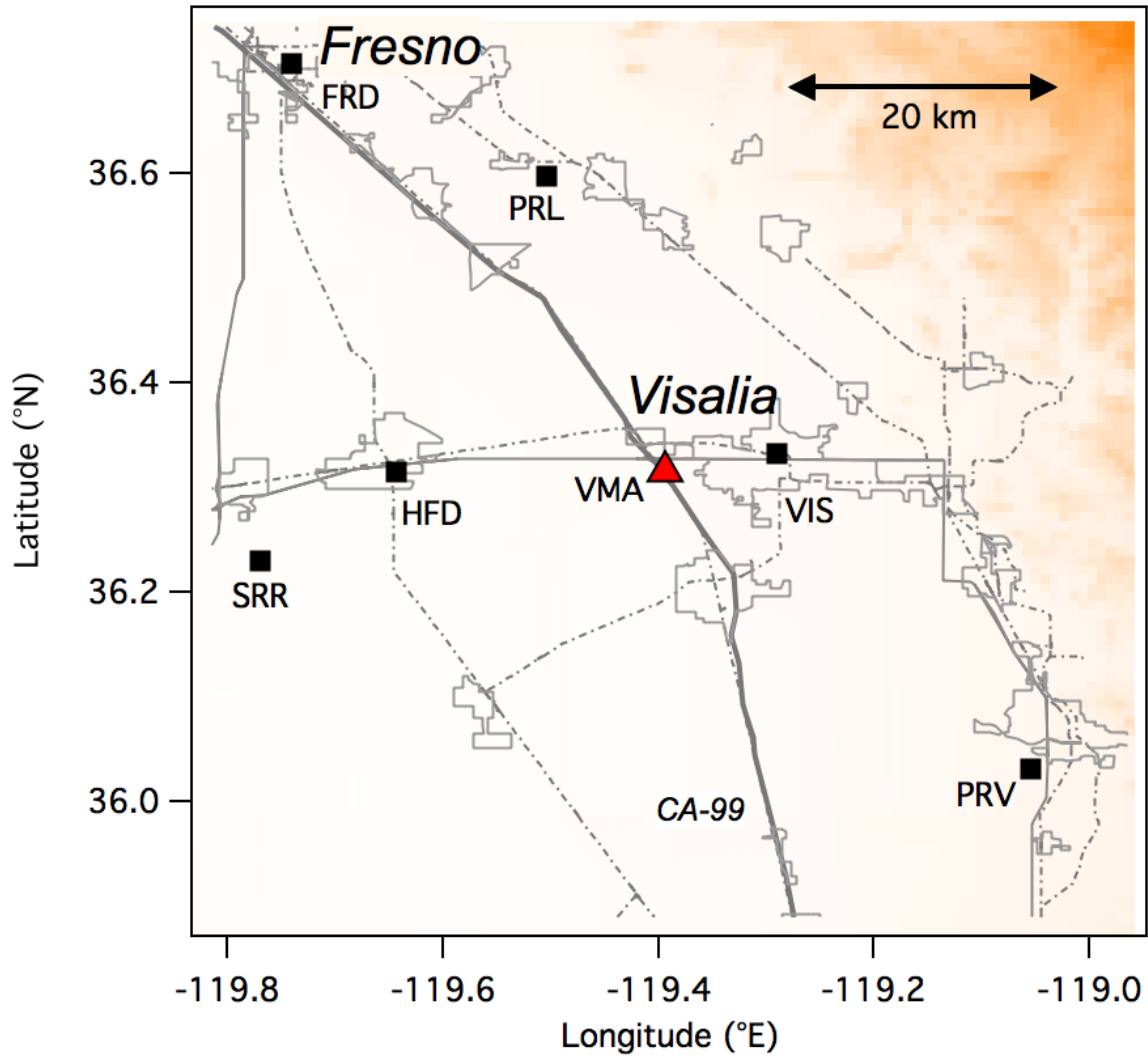
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Figure 1. (a) Topographic map showing the air basins of California (dashed black lines); the San Joaquin Valley Air Basin (SJVAB) is outlined in heavy solid black. Interstate highways and urban areas are shown in gray. The filled red triangles show the CABOTS measurement sites at Bodega Bay (BBY), Half Moon Bay (HMB), Visalia Municipal Airport (VMA), and Chews Ridge Observatory (CRO).

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Figure 1. (b) Same as (a), but showing an enlarged view of the area surrounding the VMA. The solid and dot-dash gray lines represent the major highways and railroads, respectively, with the heavier solid line showing CA-99 (see text). The filled black squares show the 6 closest regulatory O₃ monitors active during the CABOTS campaign: Visalia (VIS), Hanford (HFD), Santa Rosa Rancheria (SRR), Fresno-Drummond St. (FRD), Parlier (PRL), and Porterville (PRV). The elevation scale is the same as in (a).

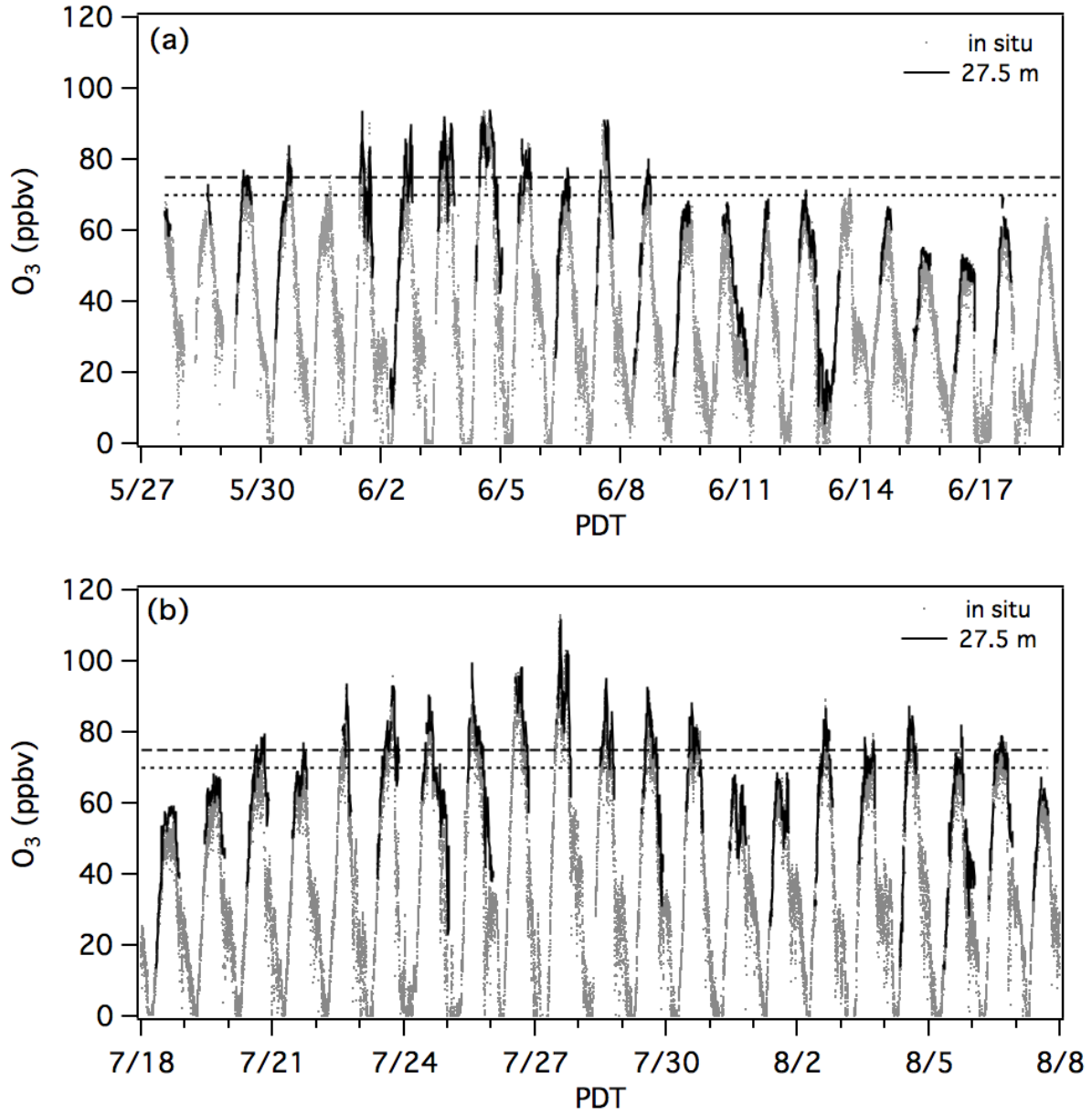
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Figure 2. Aerial view of the Visalia Municipal Airport (VMA) showing the 1 km lidar slant path line of sight as a yellow arrow with the TOPAZ truck located at the base. The Scientific Aviation Mooney and AJAX Alpha Jet are shown flanking the NOAA ESRL TOPAZ truck below the Google Earth image. Mooney and TOPAZ photos by A. Langford. Alpha Jet photo by W. von Dauster.

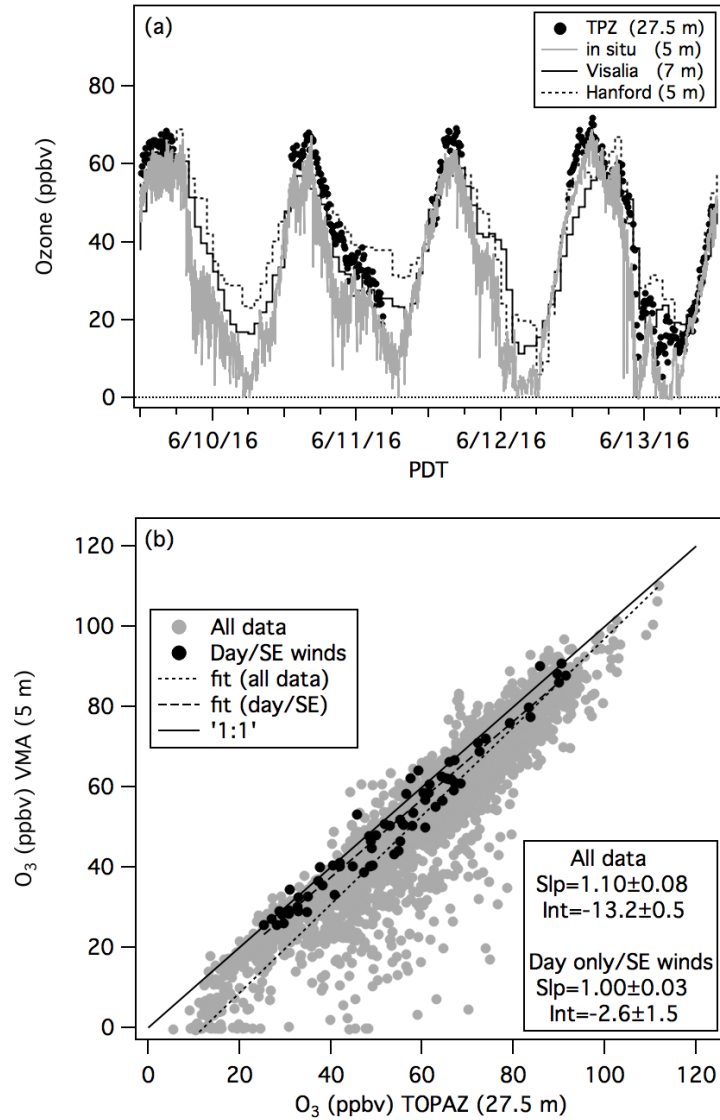
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Figure 3. Time series plots (local Pacific Daylight Time, PDT) of the O₃ concentrations retrieved 815±15 m downrange and 27.5 m above the surface by TOPAZ (black line) with the measurements from the *in-situ* 2B monitor sampling 5 m agl at the TOPAZ location (gray dots) during the first (a) and second (b) IOPs. The dashed and dotted lines respectively show the 2008 (75 ppbv) and 2015 (70 ppbv) O₃ NAAQS.

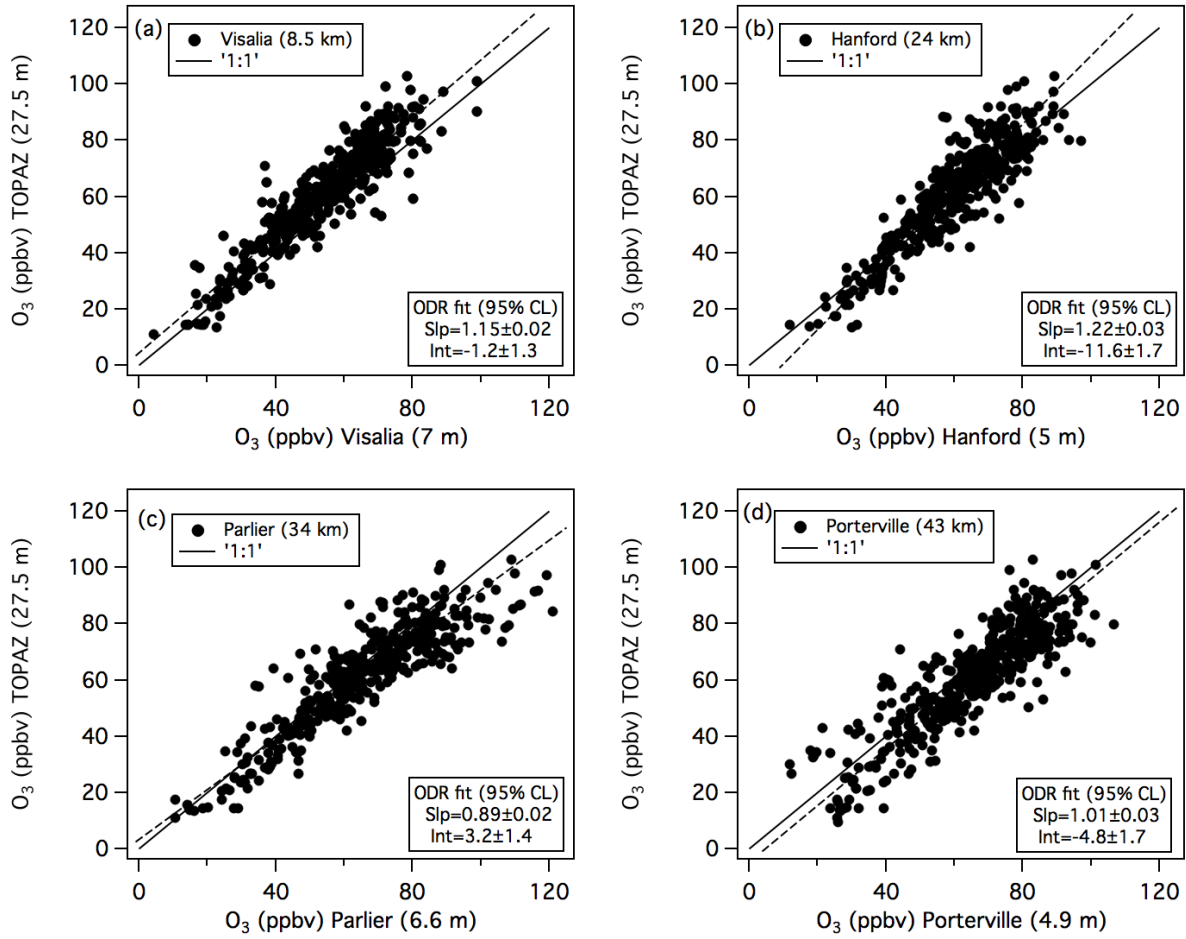
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Figure 4. (a) Four-day time series (9-13 June) showing the O₃ concentrations in air sampled 5 m agl above the TOPAZ truck at the VMA (gray line) and the O₃ mixing ratios at a height of 27.5±5 m and distance of 815±15 m retrieved from the TOPAZ measurements (filled black circles). The solid black and dotted staircase lines show the 1-h measurements from the Visalia and Hanford regulatory monitors. (b) Scatter plot comparing the 27.5 m TOPAZ measurements to the interpolated 5 m *in-situ* measurements. The filled gray circles (with dotted ODR fit) show the entire CABOTS data set from Figure 3, and the filled black circles (with dashed ODR fit) show only those measurements made during the day (0900 to 1830 PDT) when the winds were southeasterly (125 to 145°) and greater than 2.5 m s⁻¹. The solid line shows the 1:1 correspondence.

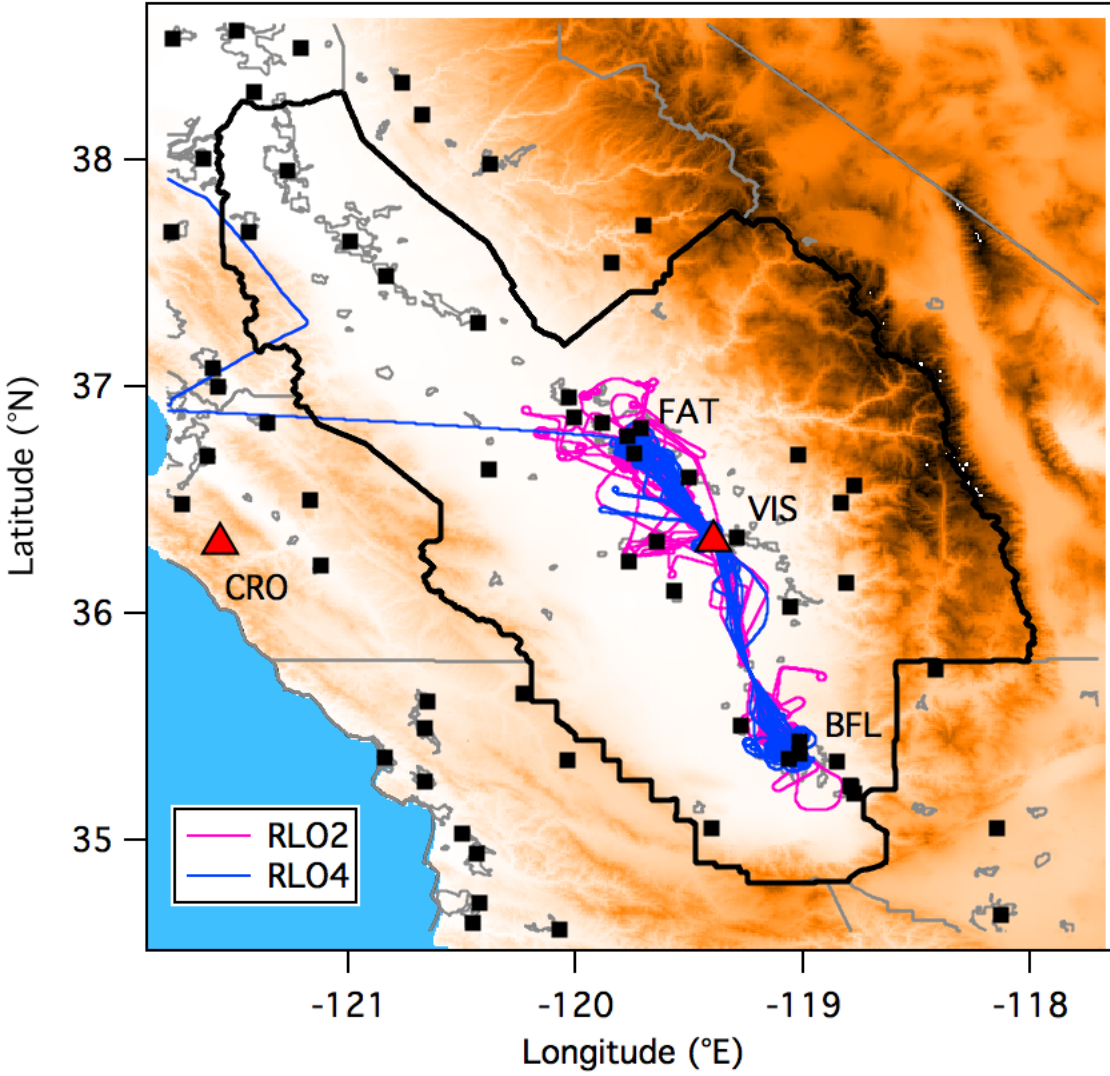
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Figure 5. Scatter plots with ODR fits comparing the 27.5 m TOPAZ measurements with the 1-h measurements from the regulatory monitors at (a) Visalia-N. Church Street, (b) Hanford, (c) Parlier, and (d) Porterville. The measurements in the upper box and x-axis label refer to the distance from the VMA and sampling height above ground, respectively. The Visalia monitor is operated by the California Resources Board. The remaining three are operated by CARB and the SJVAPCD. The TOPAZ measurements are interpolated to the 1-h time base of the regulatory measurements for the comparison.

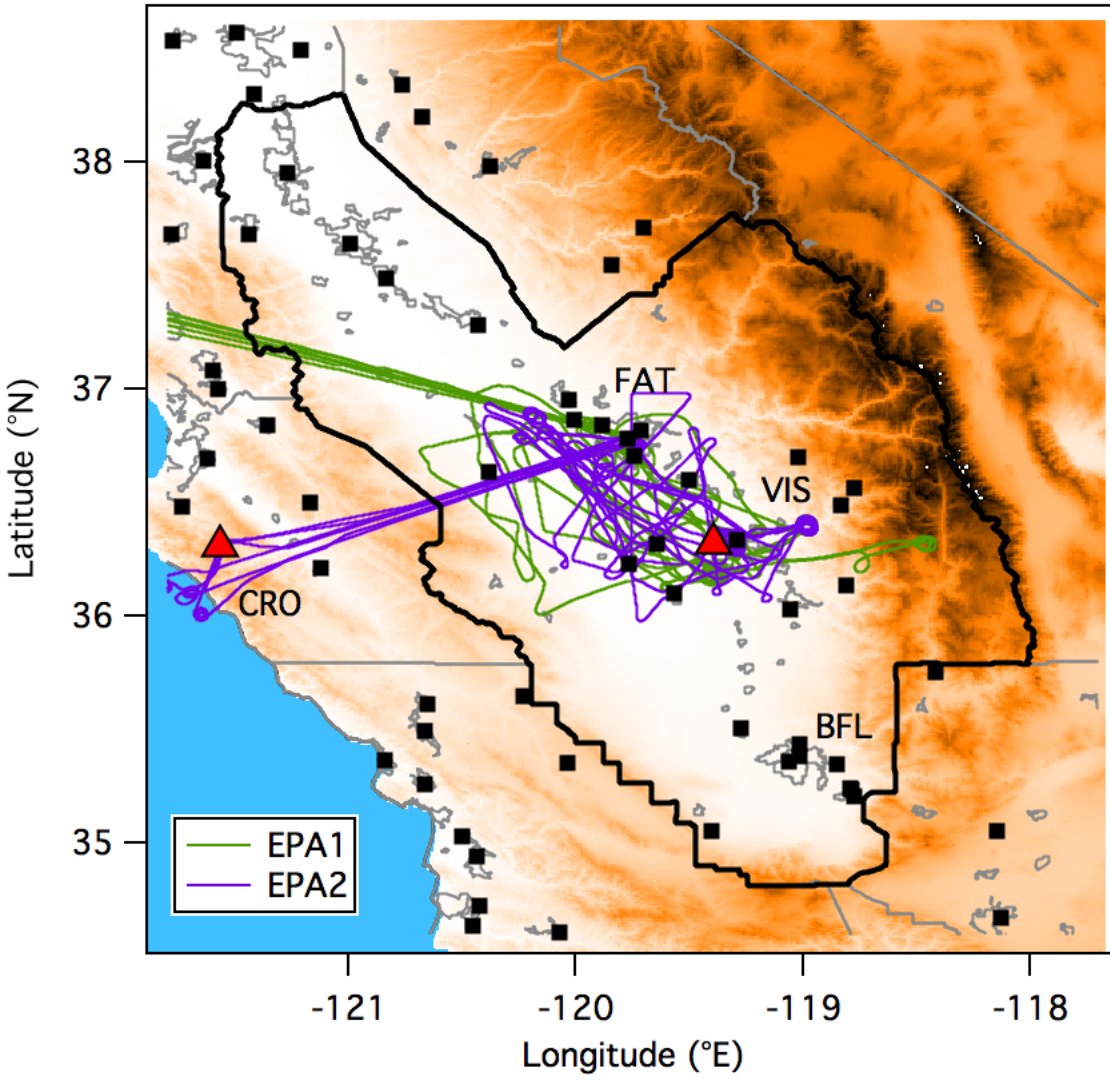
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Figure 6. (a) Map of the San Joaquin Valley showing the RLO flight tracks coincident with the TOPAZ measurements (RLO2 and RLO4). The filled black squares show the regulatory surface monitors. The CABOTS sampling sites at CRO and VMA are marked by red triangles. The other abbreviations are the Fresno (FAT), Visalia (VIS), and Bakersfield (BFL) airport codes. Note that VMA and VIS refer to the same airport.

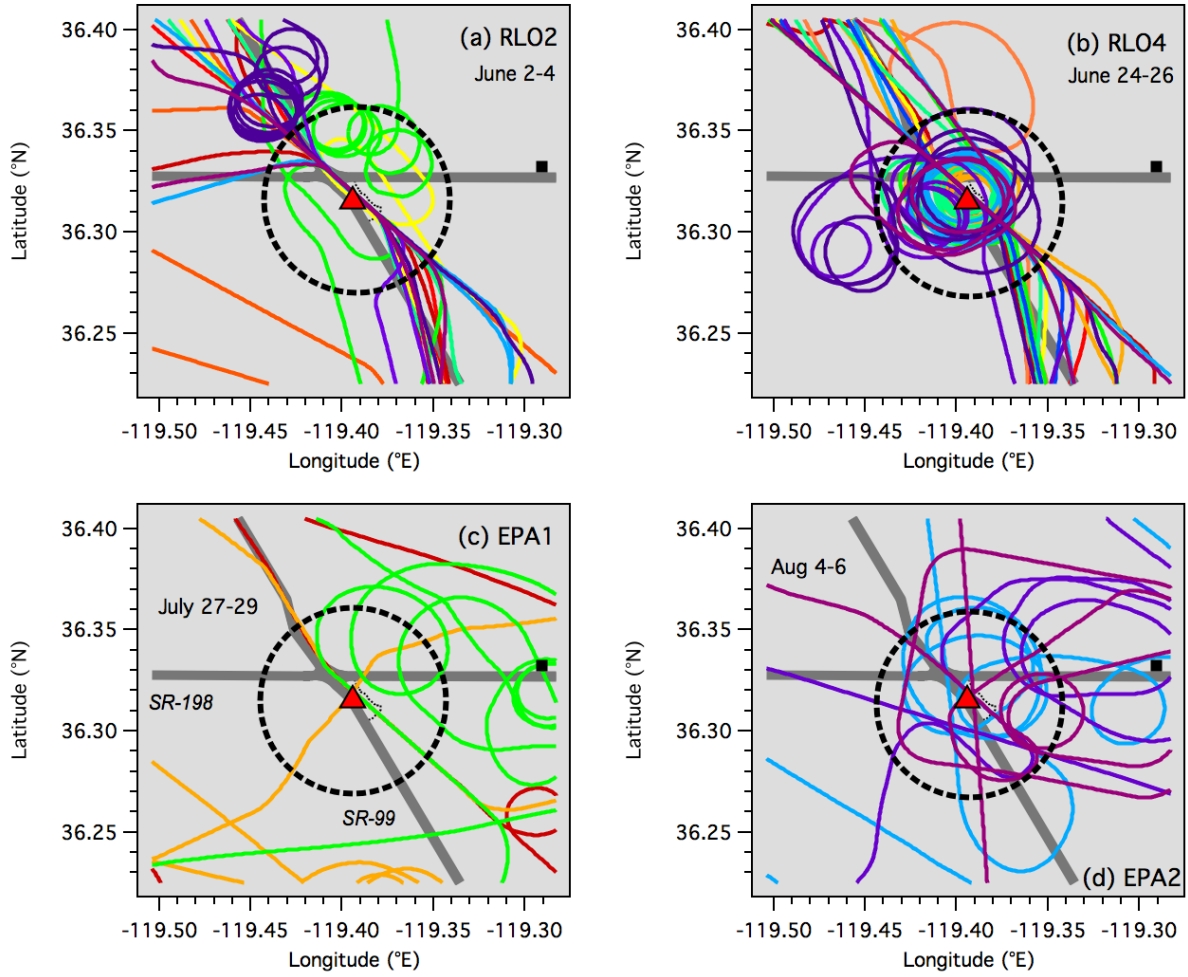
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Figure 6. (b) Same as (a), but with the EPA flight tracks (EPA1 and EPA2).

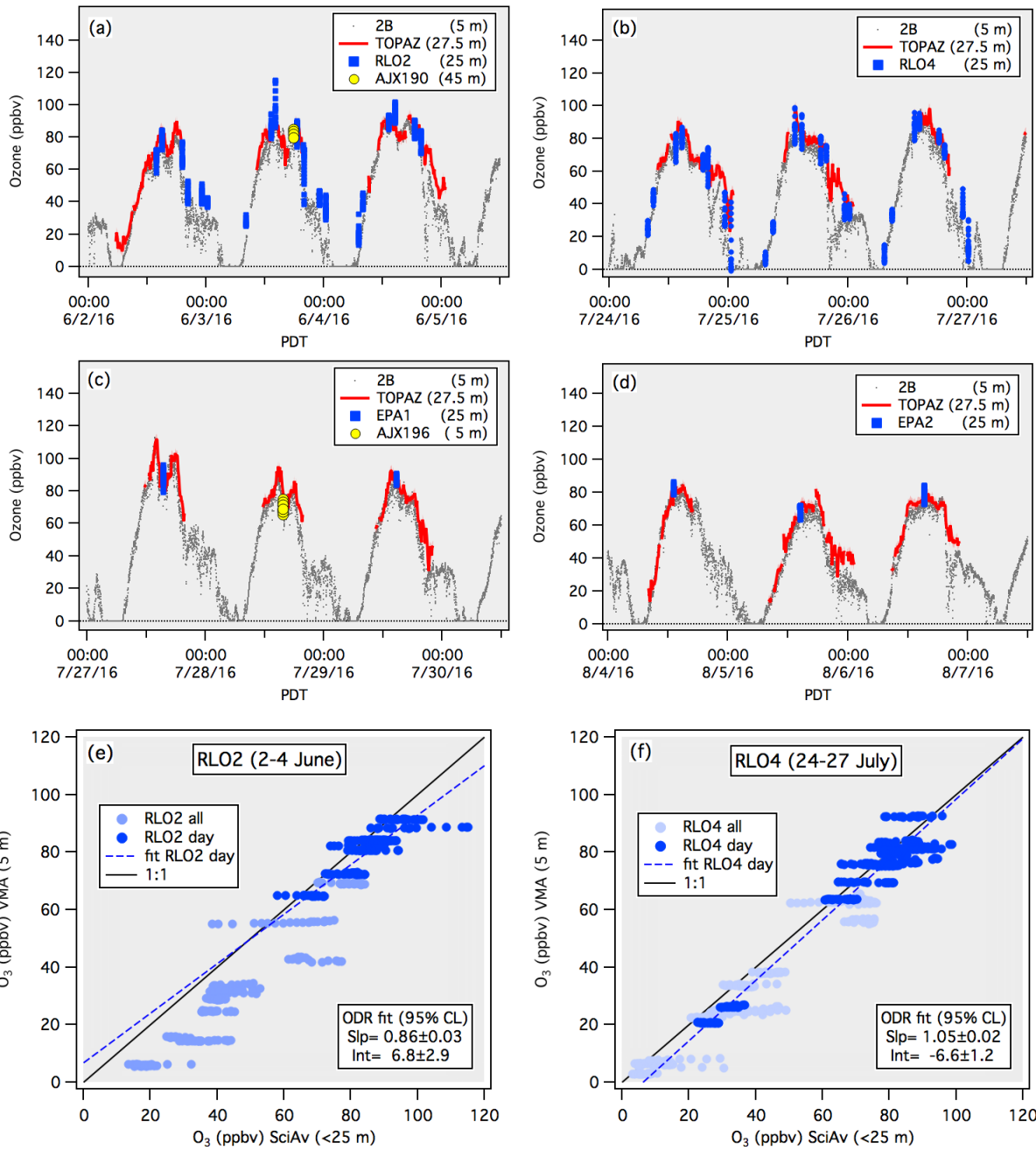
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Figure 7. RLO and EPA flight tracks in the vicinity of TOPAZ. (a) RLO2 (2-4 June), (b) RLO4 (24-26 July), (c) EPA1 (27-29 July), and (d) EPA2 (4-6 August). Each color represents a different flight. The red triangle marks the location of TOPAZ at the VMA and the dashed black circles show the 5 km radius used for the profile comparisons. The black square represents the Visalia-N. Church St. O₃ monitor.

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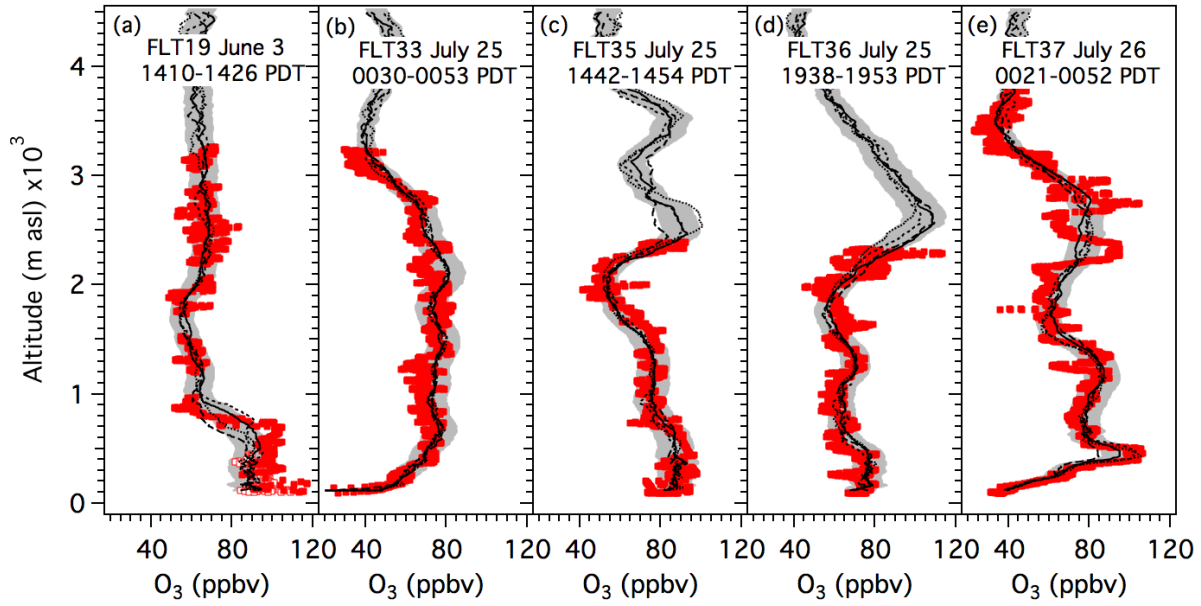


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Figure 8. (a)-(d) Time series of the surface in-situ O₃ (gray dots) and 27.5 m TOPAZ O₃ (red line) measured during the RLO and EPA low approaches on (a) 2-5 June, (b) 24-27 July, (c) 27-30 July, and (d) 4-7 August 2016. The red envelope shows the the TOPAZ data ±3 ppbv, the nominal accuracy of the lidar retrievals. The blue squares represent the 1-s sampled (2-s recorded) Scientific Aviation measurements made between the surface and 25 m agl. The filled yellow circles in (a) and (c) show 2-s measurements from AJAX low approaches (see text). Panels (e) and (f) show scatter plots of the in-situ surface measurements and the Scientific Aviation data from the RLO flights in panels (a) and (b), respectively. The ODR fit parameters refer to the dark blue points which represent the measurements from daytime (0830-1830 PDT) flights.

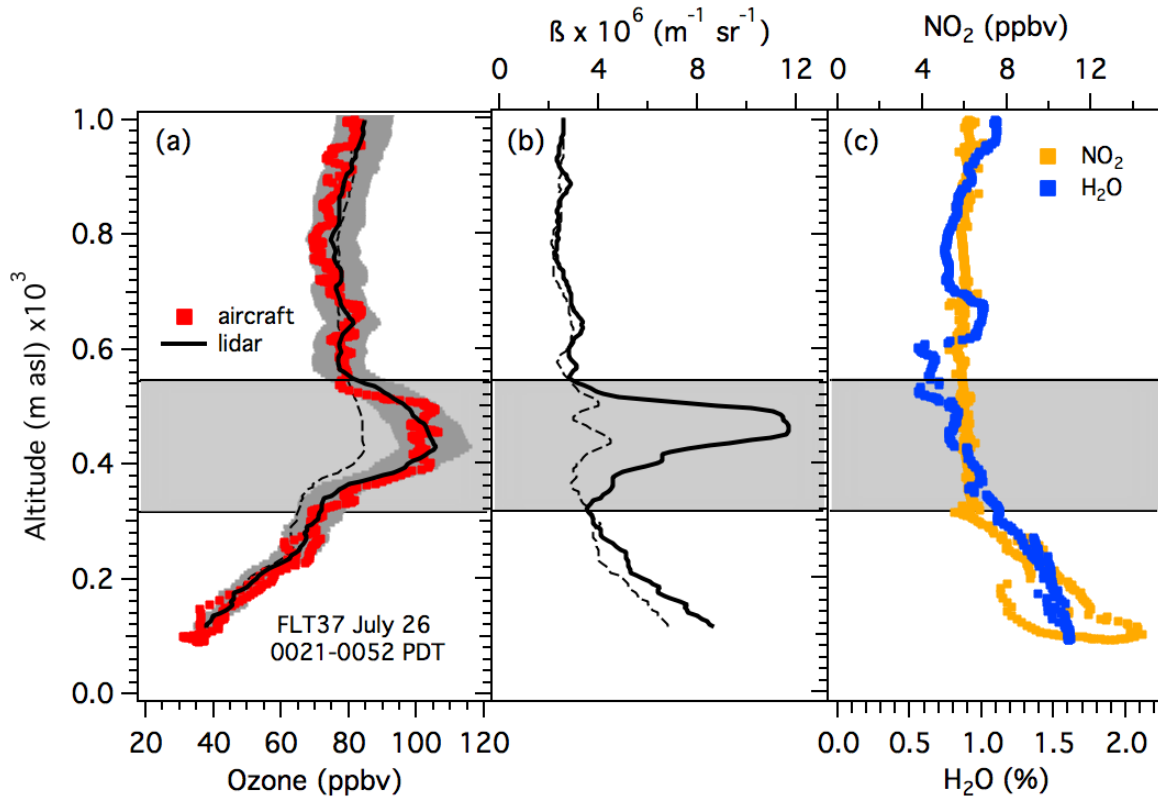
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Figure 9. Profile plots comparing the TOPAZ (black lines) and Scientific Aviation (red squares) O₃ measurements on (a) FLT19, 3 June, (b) FLT33, 25 July, (c) FLT35 25 July, (d) FLT36, 25 July and (e) FLT 37, 26 July. The dotted, short dash, solid, and long dash lines show the four consecutive 8-min lidar profiles acquired during the aircraft profiles. The gray envelopes show the mean lidar profile $\pm 10\%$ as reference. Note the large variability near the surface and sharp transition at 800 m in the 3 June aircraft measurements (cf. Figure 3a).

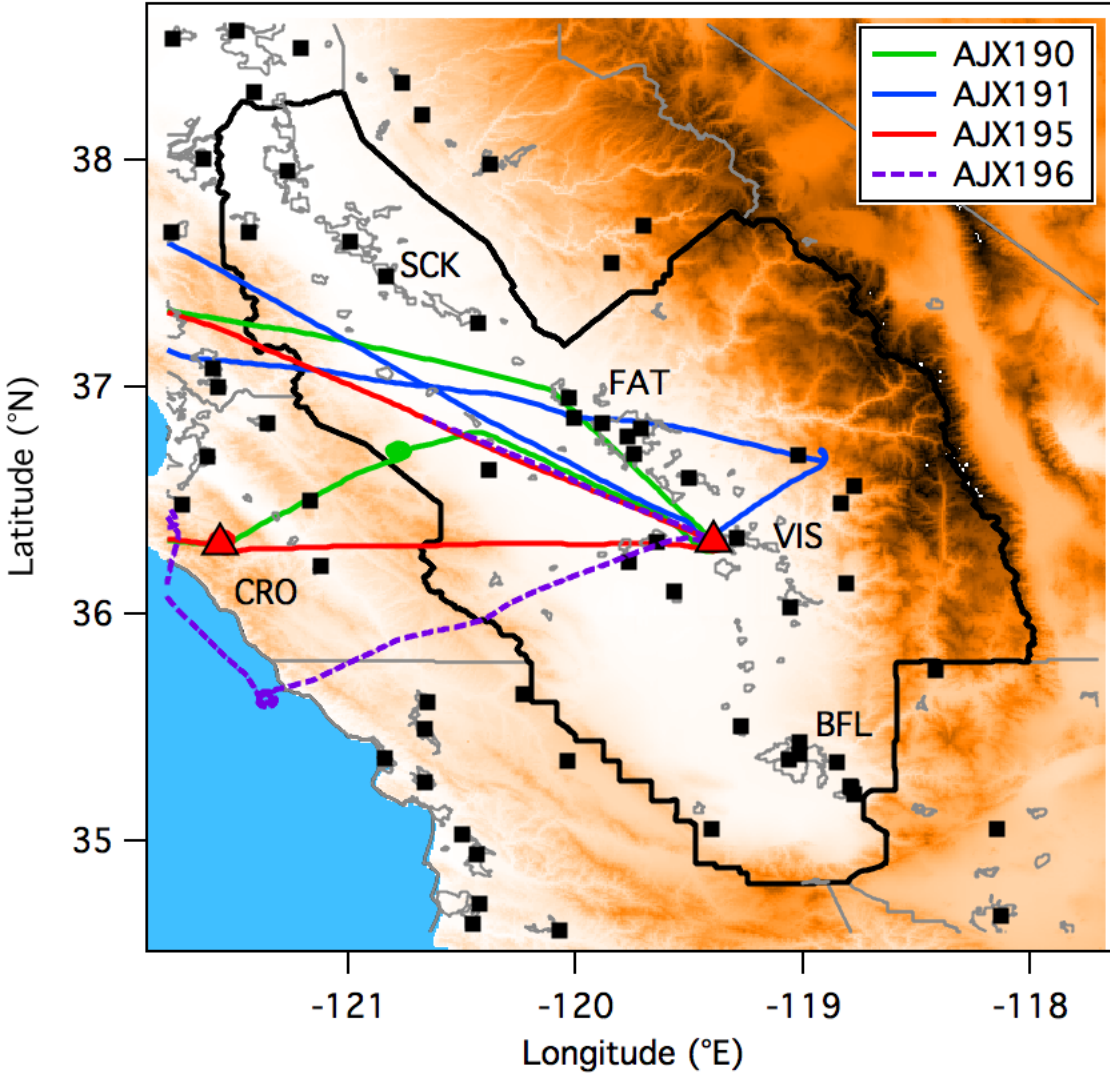
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Figure 10. (a) Expanded view of the lidar and aircraft O_3 profiles from Figure 9e plotted with coincident: (b) lidar backscatter, and (c) aircraft NO_2 and H_2O profiles. The solid black profile ($\pm 10\%$ in gray) in (a) shows the lidar profile coinciding with the aircraft measurements below 1 km; the dashed black line shows the profile measured 16-24 minutes later. Likewise, for the backscatter profiles in (b). The horizontal gray band highlights the smoke puff from the Soberanes fire.

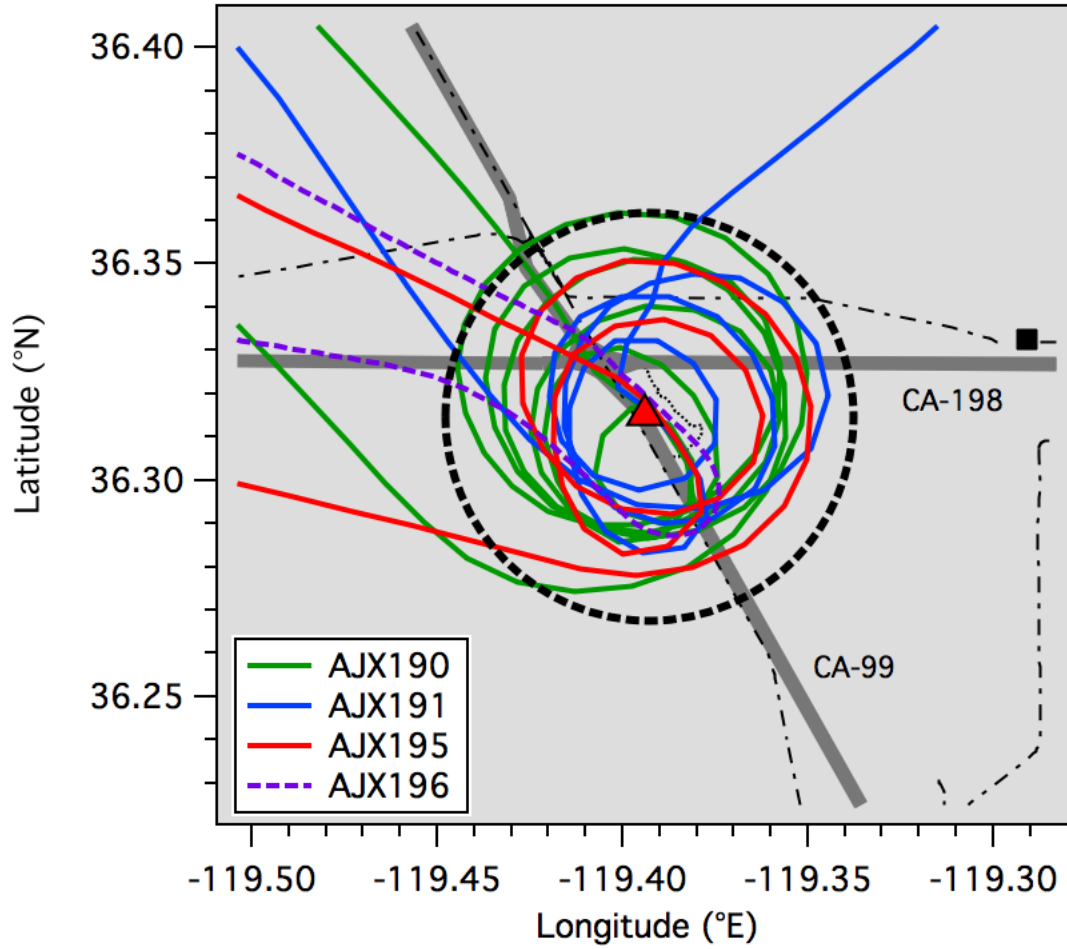
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Figure 11. Map of the San Joaquin Valley showing the AJAX flight tracks on 3 June (AJX190), 15 June (AJX191), 21 July (AJX195), and 28 July (AJX196). The abbreviations and symbols are the same as in Figure 6.

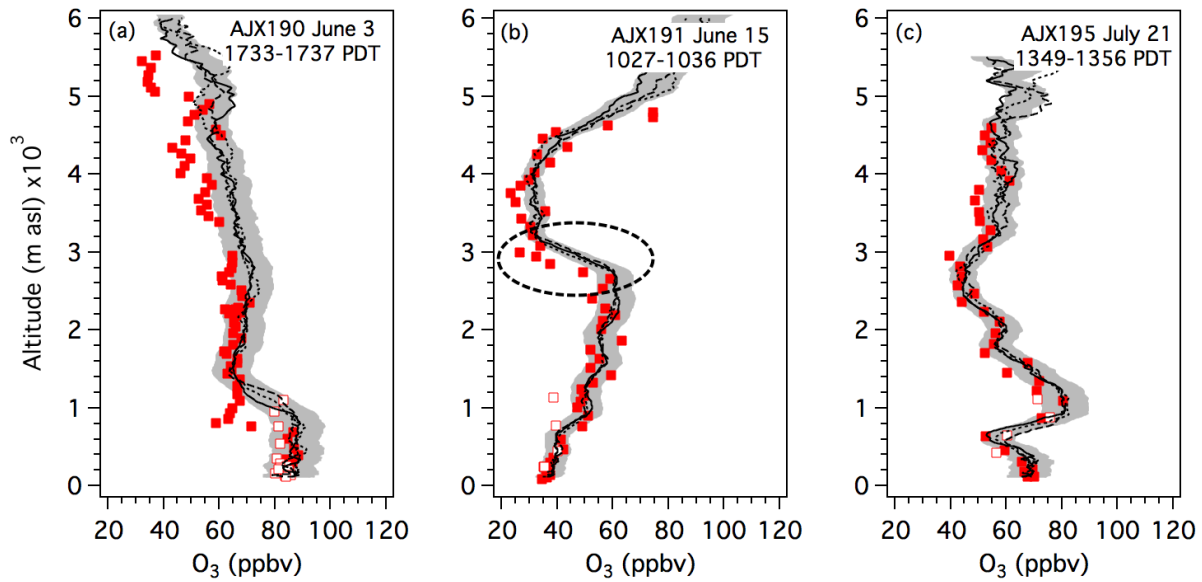
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Figure 12. AJAX flight tracks in the vicinity of the VMA (red triangle). The black square represents the Visalia-N. Church St. O₃ monitor and the dashed black circle marks the 5 km radius window used for the profile comparisons. The heavy gray lines show the major highways and the black dot-dash lines the railroads.

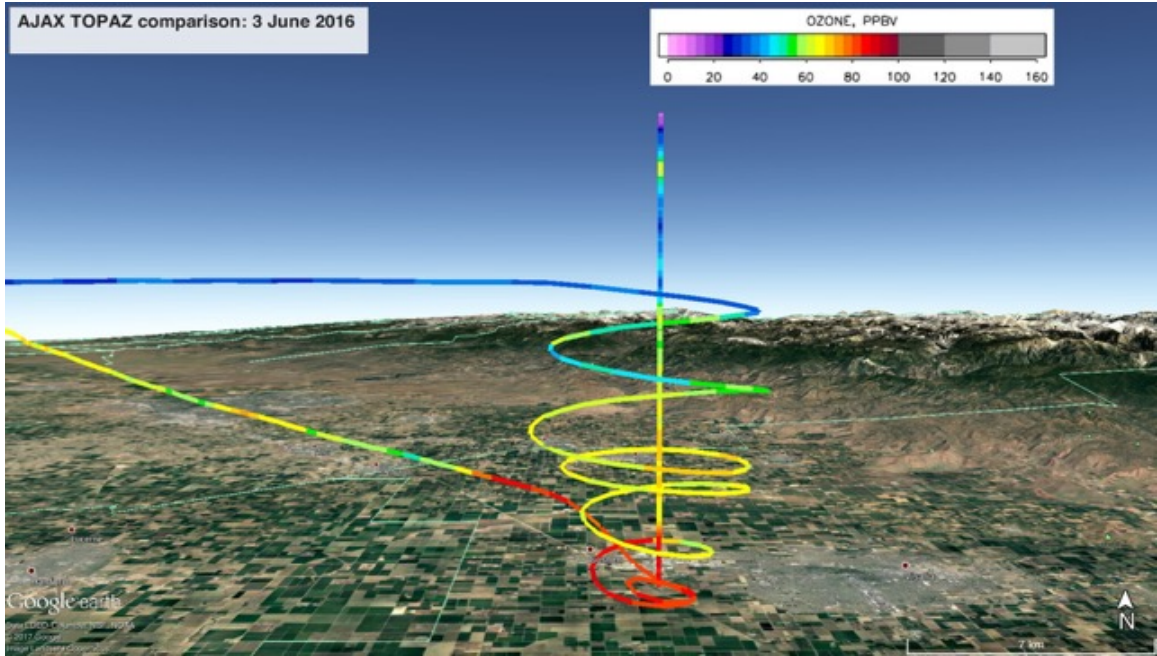
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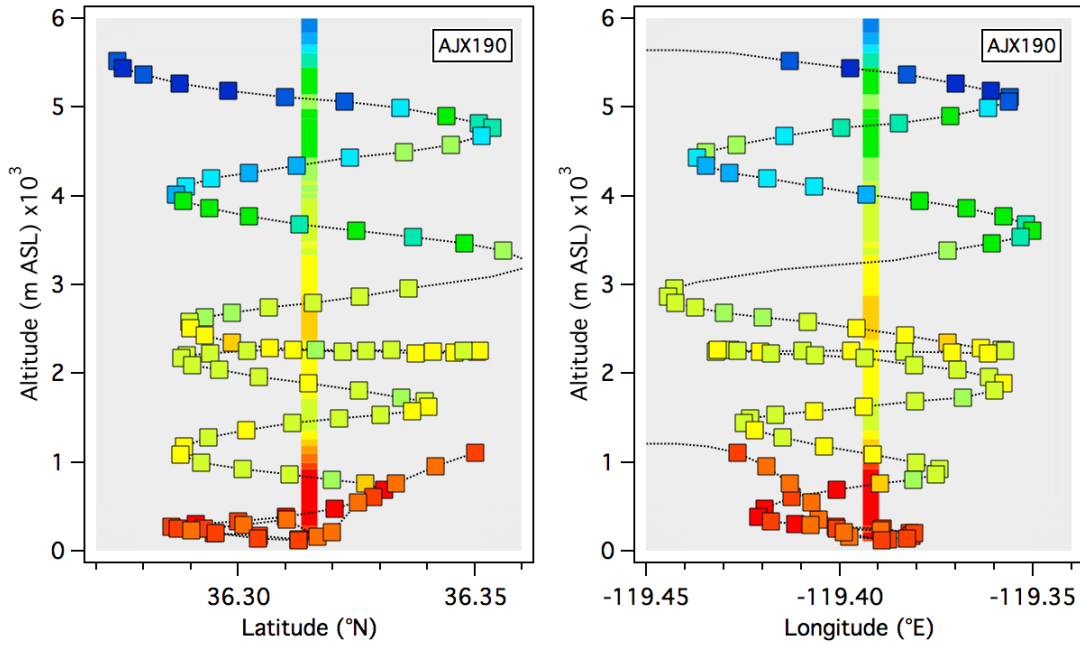
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Figure 13. Profile plots comparing the TOPAZ (black lines) and 10-s AJAX (red squares) measurements on (a) AJX190, 3 June, (b) AJX191, 15 June, and (c) AJX195, 21 July. The closed squares correspond to the Alpha Jet descent and the open squares the subsequent climb out. Note the differences between these measurements. The dotted, dashed, and solid lines show the order of the three 8-min lidar profiles that bracket the AJAX profile. The gray envelopes show the mean lidar profile $\pm 10\%$ as reference. The significance of the dashed oval in (b) is discussed in the text.

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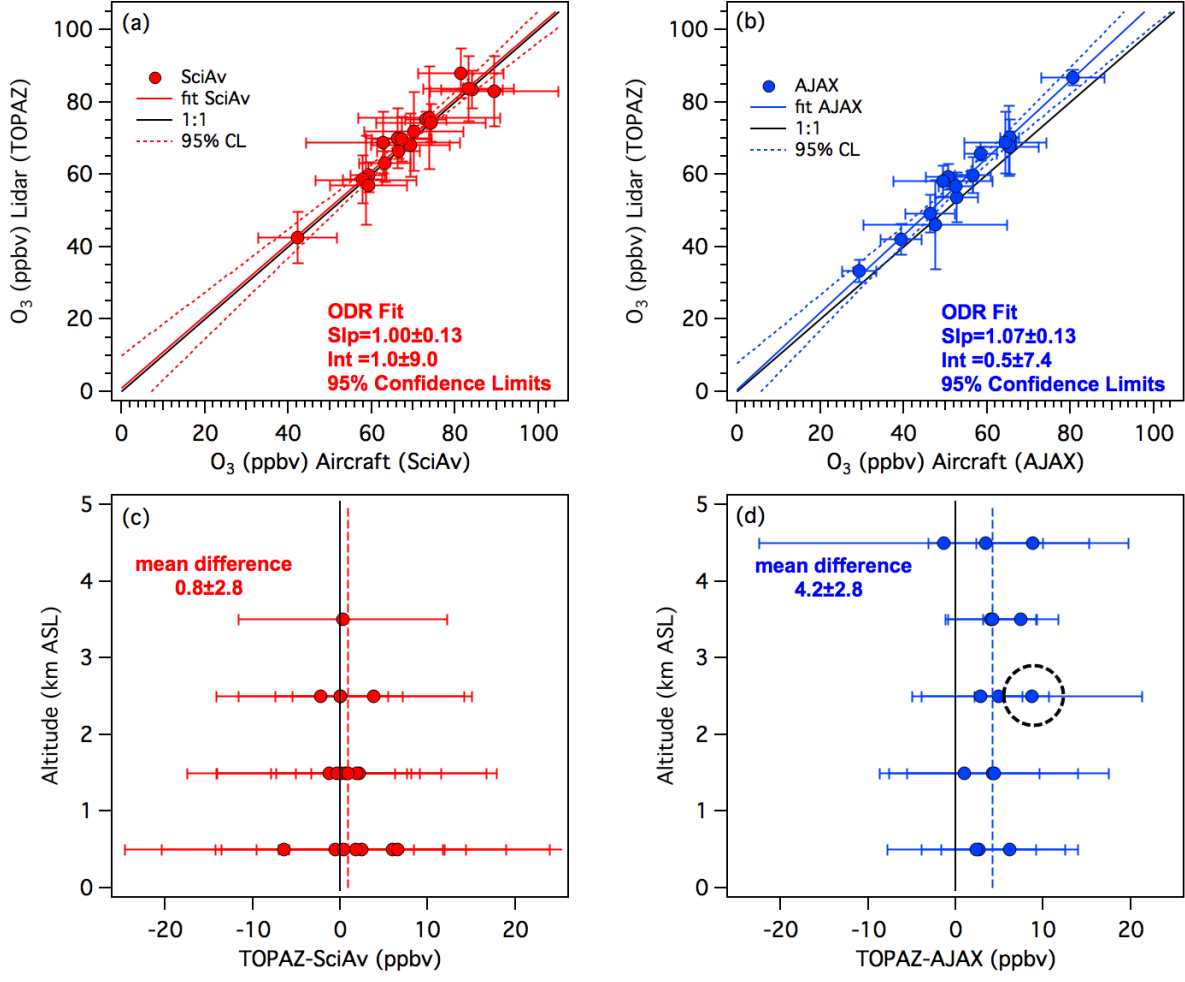
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Figure 14. (top) Google Earth image of the TOPAZ and AJAX profiles from 3 June 2016 showing the spatial variations across the ~8 km diameter spiral profile by the Alpha Jet during its descent and climb out over the VMA. (bottom) AJAX and TOPAZ profiles from Figure 13a plotted as a function of latitude (left) and longitude (right). Both plots are 10 km wide. Note the strong horizontal gradients below 1.2 km.

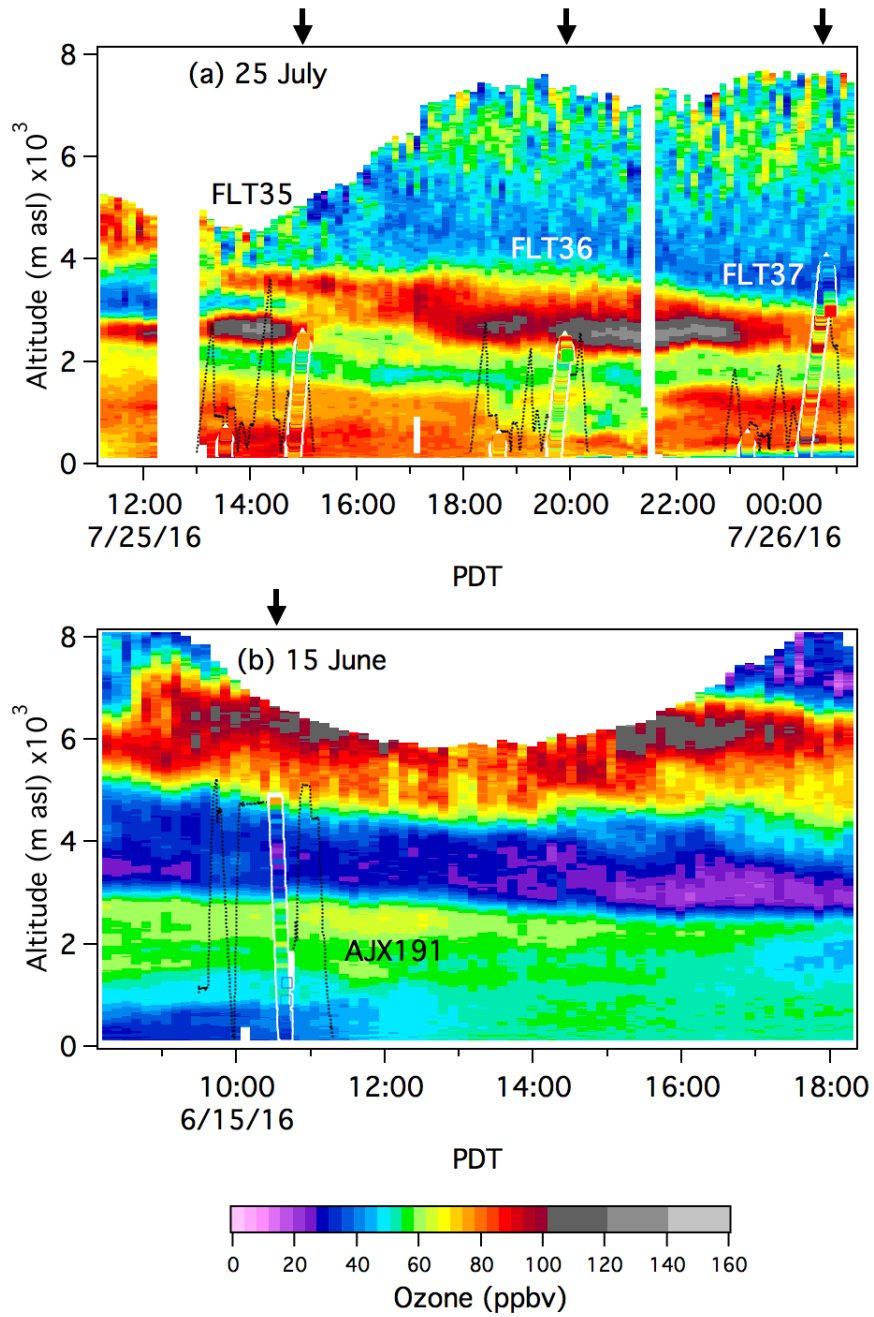
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Figure 15. (a) and (b), Scatter plots comparing the TOPAZ lidar retrievals to *in-situ* O₃ measurements from 7 SciAv Mooney and 3 NASA Alpha Jet flights, respectively, averaged over 1 km vertical bins. The error bars show the standard deviations of the 1 km column means. (c) and (d), Differences between the 1 km mean TOPAZ and aircraft measurements from (a) and (b) plotted as a function of altitude. The vertical dashed lines show the mean differences. The dashed circle in (d) corresponds to the dashed oval in Figure 13b (see text).

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Figure 16. Time-height curtain plots of the TOPAZ ozone measurements from (a) 25-26 July with the Scientific Aviation profiles from FLT35, 36, and 37 superimposed, and (b) 15 June with the coincident AJAX profile superimposed. The aircraft measurements made within 5 km of VMA (arrows) are highlighted by squares and colored using the same scale as the TOPAZ data. The high O_3 layers around 3 km asl in (a) are related to the Soberanes Fire; the measurements plotted in the lower right corner of (a) correspond to the data shown in Figure 10.