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Evaluating the performance of five different chemical ionization techniques for detecting gaseous oxygenated organic species

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- 5 Matthieu Riva^{1,2}, Pekka Rantala¹, Jordan E. Krechmer³, Otso Peräkylä¹, Yanjun Zhang¹, Liine
- 6 Heikkinen¹, Olga Garmash¹, Chao Yan¹, Markku Kulmala^{1,4}, Douglas Worsnop^{1,3}, Mikael Ehn¹
- 7 Institute for Atmospheric and Earth System Research / Physics, Faculty of Science, University of
- 8 Helsinki, Helsinki, 00140, Finland
- ⁹ Univ Lyon, Université Claude Bernard Lyon 1, CNRS, IRCELYON, F-69626, Villeurbanne,
- 10 France.
- ³ Aerodyne Research Inc., Billerica, MA, USA.
- ⁴ Aerosol and Haze Laboratory, Beijing Advanced Innovation Center for Soft Matter Science and
- 13 Engineering, Beijing University of Chemical Technology (BUCT), Beijing, China

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- 15 Correspondence to:
- Matthieu Riva (matthieu.riva@ircelyon.univ-lyon1.fr) & Mikael Ehn (mikael.ehn@helsinki.fi)

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Abstract

21 The impact of aerosols on climate and air quality remains poorly understood due to multiple factors. One of 22 the current limitations is the incomplete understanding of the contribution of oxygenated products, generated 23 from the gas-phase oxidation of volatile organic compounds (VOC), to aerosol formation. Indeed, atmospheric 24 gaseous chemical processes yield thousands of (highly) oxygenated species, spanning a wide range of chemical 25 formulas, functional groups and, consequently, volatilities. While recent mass spectrometric developments 26 have allowed extensive on-line detection of a myriad of oxygenated organic species, playing a central role in 27 atmospheric chemistry, the detailed quantification and characterization of this diverse group of compounds remains extremely challenging. To address this challenge, we evaluated the capability of current state-of-the-28 29 art mass spectrometers equipped with different chemical ionization sources to detect the oxidation products 30 formed from α-pinene ozonolysis under various conditions. Five different mass spectrometers were deployed 31 simultaneously for a chamber study. Two chemical ionization atmospheric pressure interface time-of-flight 32 mass spectrometers (CI-APi-TOF) with nitrate and amine reagent ion chemistries and an iodide chemical 33 ionization time-of-flight mass spectrometer (TOF-CIMS). Additionally, a proton transfer reaction time-offlight mass spectrometer (PTR-TOF 8000) and a new "Vocus" PTR-TOF were also deployed. In the current 34 study, we compared around 1000 different compounds between each of the five instruments, with the aim to 35 determine which oxygenated VOC (OVOC) the different methods were sensitive to, and identifying regions 36 37 where two or more instruments were able to detect species with similar molecular formulae. We utilized a 38 large variability in conditions (including different VOC, ozone, NO_x and OH scavenger concentrations) in our 39 newly constructed atmospheric simulation chamber for a comprehensive correlation analysis between all 40 instruments. This analysis, combined with estimated concentrations for identified molecules in each instrument, yielded both expected and surprising results. As anticipated based on earlier studies: the PTR 41 42 instruments were the only ones able to measure the precursor VOC; the iodide TOF-CIMS efficiently detected many semi-volatile organic compounds (SVOC) with 3 to 5 oxygen atoms; and the nitrate CI-APi-TOF was 43 mainly sensitive to highly-oxygenated organic (O > 5) molecules (HOM). In addition, the Vocus showed good 44 agreement with the iodide TOF-CIMS for the SVOC, including also a range of organonitrates. The amine CI-45 APi-TOF agreed well with the nitrate CI-APi-TOF for HOM dimers. However, the loadings in our experiments 46 47 caused the amine reagent ion to be considerably depleted, causing non-linear responses for monomers. This

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study explores and highlights both benefits and limitations of currently available chemical ionization mass spectrometry instrumentation for characterizing the wide variety of OVOC in the atmosphere. While specifically shown for the case of α -pinene ozonolysis, we expect our general findings to be valid also for a wide range of other VOC-oxidant systems. As discussed in this study, no single instrument configuration can be deemed better or worse than the others, as the optimal instrument for a particular study ultimately depends on the specific target of the study.

1. Introduction

Atmospheric aerosols, a mixture of solid and liquid particles consisting of organic and inorganic substances suspended in the air, have significant impact on climate (Albrecht, 1989; Hallquist et al., 2009; Intergovernmental Panel on Climate Change, 2014; Twomey, 1977). They are also recognized to adversely impact air quality and human health, representing nowadays the fifth-ranking human health risk factor, globally (Gakidou et al., 2017). Depending on the regions, organic aerosol on average contributes 20-90% to the submicron aerosol mass (Jimenez et al., 2009), with secondary organic aerosol (SOA) as the largest source of atmospheric organic aerosol (Hallquist et al., 2009; Jimenez et al., 2009). SOA is predominantly formed through the gas-phase oxidation of volatile organic compounds (VOC), producing oxygenated VOC (OVOC), which can subsequently condense onto pre-existing aerosol particles. Generally, the more oxidized OVOC, the lower its volatility is, and the greater is the probability of this compound to partition to the particle phase. Recently, studies have provided new insights on how highly-oxygenated organic molecules (HOM) can form faster than previously expected, and at high enough yields to make them a major source of condensing, or even nucleating, compounds (Ehn et al., 2014; Jokinen et al., 2015; Kirkby et al., 2016; Wang et al., 2018).

The quantitative assessment of the impact of aerosol on climate remains poorly understood due to a number of factors, including an incomplete understanding of how VOC oxidation processes contribute to new particle and SOA formation (Glasius and Goldstein, 2016). Indeed, atmospheric oxidation processes can lead to the formation of thousands of oxidized products from a single precursor (Glasius and Goldstein, 2016; Goldstein and Galbally, 2007). As a result of these complex oxidation processes, atmospheric organic species span an extremely wide range of chemical formulas, structures, and physicochemical properties. Volatilities range from volatile species present only in the gas phase, via low- and semi-volatile organic compounds

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(LVOC and SVOC), to extremely low volatility organic compounds (ELVOC) present mainly in particle phase (Donahue et al., 2012). The chemical complexity of OVOC poses a major challenge in detecting, quantifying, and characterizing such a large number and wide variety of organic compounds.

Mass spectrometric techniques, which can detect a large range of species simultaneously, are well-suited to tackle these challenges. This is underlined by the major role of the mass spectrometers in improving our understanding of the atmospheric chemical composition over the last twenty years (Breitenlechner et al., 2017; Ehn et al., 2014; Jokinen et al., 2012; Krechmer et al., 2018; Lindinger et al., 1998; Yuan et al., 2017). Proton transfer reaction (PTR) has been one of the most used low-pressure ionization techniques since the mid-1990s (Lindinger et al., 1998). Since then, the PTR-MS technique has been greatly improved in terms of sensitivity, detection limit and mass resolution by introducing the PTR-TOF-MS (Yuan et al., 2017). The latest version has detection limits as low as 10⁷ molecules cm⁻³. While such techniques can characterize VOC, the PTR-MS technique has not been able to measure more oxygenated organic species. This is mostly due to losses of these low volatile compounds in the sampling lines and on the walls of the inlet (caused e.g. by very low flow rates), as the instrument was designed to primarily measure volatile compounds.

Several different chemical ionization mass spectrometry (CIMS) methods have been developed, including low-pressure systems like CF₃O⁻-CIMS for specific detection of oxygenated VOC and SVOC including hydroperoxides (Crounse et al., 2006), acetate-CIMS for selective detection of organic acids (Bertram et al., 2011), and the iodide adduct ionization CIMS for the detection of wider range of OVOC, including alcohols, hydroperoxides and peroxy acids (Lee et al., 2014; Riva et al., 2017). These instruments, based on negative ion chemistry, can detect oxygenated gas phase compounds at concentrations as low as ~10⁶ molecules cm⁻³. Finally, the discovery of the HOM was possible due to the development of nitrate chemical ionization source connected to an atmospheric pressure interface time-of-flight mass spectrometer (CI-APi-TOF) (Ehn et al., 2014; Jokinen et al., 2012). The selectivity and high sensitivity for molecules containing many functional groups (detection limit below 10⁵ molecules cm⁻³) of the nitrate CI-APi-TOF makes this instrument perfect for detecting HOM and even certain radicals (e.g., peroxy radicals). As part of the rapid development in gas-phase mass spectrometry, several new reagent ion chemistries have been tested over the last few years. With improvements in sensitivity and/or selectivity, new methods are now able to detect a wide

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variety of oxygenated species, including radicals and stabilized Criegee intermediates (Berndt et al., 2015, 2017, 2018; Breitenlechner et al., 2017; Hansel et al., 2018; Krechmer et al., 2018).

The selectivity and sensitivity of the different ionization chemistries makes it impossible for one mass spectrometer to be able to measure the full range of VOC and OVOC present in the atmosphere. Hence, only a simultaneous deployment of several mass spectrometry techniques can provide a comprehensive chemical characterization of the gaseous composition. While such a multi-instrument approach maximizes the fraction of organic species measured (Isaacman-VanWertz et al., 2017, 2018), a number of questions and limitations can arise in both laboratory and field measurements. For instance, the extent to which instruments (i) can measure species with identical molecular composition, (ii) can cover the entire range of oxygenated species, (iii) can provide constant sensitivity across different conditions, have to be determined. Most studies are typically limited to one, or perhaps two, mass spectrometers, and then it is also important to know which fraction of the OVOC distribution these instruments are sensitive to. To our knowledge, systematic comparisons of the most commonly used, or recently developed, gas-phase mass spectrometers are not yet available. In this work, we compared the suitability of five different chemical ionization methods (including iodide TOF-CIMS, nitrate and amine CI-APi-TOFs, a PTR-TOF and the newly developed Vocus PTR-TOF) for the detection of OVOC formed from α-pinene ozonolysis during a comprehensive chamber study with varying VOC, O₃, and NO₃ concentrations. We characterized the time evolution of around 1000 compounds and explored the capability of these instruments to measure OVOC of different oxygenation level within different compound groups.

2. Experimental Section

2.1. Chamber experiments

Experiments were performed at the University of Helsinki in a 2-m³ atmospheric simulation Teflon (FEP) chamber. The "COALA" chamber (named after the project in which it was constructed: Comprehensive molecular characterization of secondary Organic AerosoL formation in the Atmosphere) was operated under steady-state conditions, meaning that a constant flow of reactants and oxidants were continuously added to the chamber, while chamber air was sampled by the instruments. Under the conditions used in this study, the average residence time in the chamber was ~ 30 min, and the majority of conditions were kept constant for 6

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to 12 hours before changing to new conditions. These experiments focused on the characterization of the oxidation products arising from the α -pinene ($C_{10}H_{16}$) ozonolysis. α -pinene was used for the generation of oxidation products because it is the most abundant monoterpene emitted by the boreal forests and is one of the most important SOA precursors on a global scale (Jokinen et al., 2015; Kelly et al., 2018).

The experiments were conducted at room temperature ($27 \pm 2^{\circ}$ C) and under dry conditions (RH < 1%). An overview of the measurements, as well as the experimental conditions, are presented Figure 1. α -pinene was introduced to the chamber from a gas cylinder, and steady-state concentrations of α -pinene were varied from 20 to 100 ppb. As alkene ozonolysis yields OH radicals (Atkinson et al., 1997), in some experiments, ~ 1500 ppm of carbon monoxide (CO) was injected to serve as OH scavenger. 10 to 50 ppb of O₃ was generated by injecting purified air through an ozone generator (Dasibi 1008-PC) and monitored over the process of the campaign using a UV photometric analyzer (Model 49P, Thermo-Environmental). In the experiments performed in the presence of NO_x, 400 nm LED lights were used to generate NO in the chamber from the photolysis of the injected NO₂. The purified air ([O₃/NO_x] and [VOC] reduced to less than 1 ppb and 5 ppb, respectively), generated by an air purification system (AADCO, 737 Series, Ohio, USA) running on compressed air, was used as a bath gas. Temperature, relative humidity (RH) and pressure were monitored by a Vaisala Humidity and Temperature Probe (INTERCAP® HMP60) and a differential pressure sensor (Sensirion SDP1000-L025).

2.2. Mass spectrometers

We deployed five chemical ionization schemes to the COALA chamber in order to characterize the chemical composition of the gas-phase oxidation products for formed from α -pinene ozonolysis. In this section, we briefly present each instrument, summarized in Table 1. As each mass spectrometer has slightly different working principles, references to more detailed descriptions are provided. Specific benefits and limitations, which were not often discussed in earlier studies, are reviewed in section 2.4. Each of the mass spectrometers were equipped with a mass analyzer manufactured by Tofwerk AG, either an HTOF (mass resolving power \sim 5000) or Long-TOF (LTOF, mass resolving power \sim 10000) version.

In the analysis, we focused primarily on the relative behavior of the ions measured by the different mass spectrometers. An absolute comparison was also performed, but this approach has larger uncertainty, as

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the sensitivity towards every molecule is different in each of the mass spectrometers, depending on molecular size, functionality, proton affinity, polarizability, etc. We attempted a rough estimate of absolute concentrations for each instrument, despite the fact that with around a thousand of ions analyzed, it is evident that we make no claim to be accurate for them all. As will be shown, the concentrations of gas-phase VOC and OVOC vary up to 7 orders of magnitude, and therefore useful information can still be obtained even in cases where concentration estimates could be off by an order of magnitude. Details about instruments used in this study as well as calibrations and instrumental limitations are discussed in the following sections.

2.2.1. PTR-TOF

 α -pinene concentration was measured in the COALA chamber by a proton transfer reaction time-of-flight mass spectrometer (PTR-TOF 8000, Ionicon Analytik Gmbh) – later referred to as "PTR-TOF". The technical details have been described in detail elsewhere (Graus et al., 2010; Jordan et al., 2009). The sample air from the COALA chamber was drawn to the instrument using 2-m long PTFE tubing (6 mm o.d, 4 mm i.d.) and a piece of 20-cm capillary PEEK tubing (1.6 mm o.d., 1 mm i.d.), with the sampling flow of 0.8 LPM (liters per minute). The instrument was operated using a drift tube at the pressure of around 2 mbar and a drift tube at the temperature of 60 degrees (°C). Drift tube voltage was kept at 600 V leading to E/N = 145 Td where E is the electrical field strength and N the gas number density. With these settings, the primary ion isotope (H₃¹⁸O⁺, 21.0221 Th) level stayed at 4500 cps (counts per second), and the mass resolving power of the HTOF mass analyzer was ~ 4500. Data were recorded using a time resolution of 10 s. The background of the instrument was measured approximately every day with VOC-free air generated using a custom-made catalytic converter heated to 350 °C (Schallhart et al., 2016).

2.2.2. Vocus

The Vocus PTR-TOF (proton transfer reaction time-of-flight mass spectrometer, Tofwerk AG/Aerodyne Research, Inc.), later referred to as "Vocus", is based on a new PTR-inlet design (i.e., focusing ion-molecule reactor, FIMR) with sub-ppt detection limits (Krechmer et al., 2018). Sample air was drawn to the instrument using 1-m long PTFE tubing (6 mm o.d, 4 mm i.d.), with a flow rate of 4.5 LPM. Most of the sample air was directed to the exhaust while the actual flow to the Vocus was around 0.15 LPM. The instrument was operated with 1.0 mbar drift tube pressure, the voltages being 350 (axial) and 400 (radial) and E/N = 120 Td. The signal

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level of the instrument had some instability during the campaign, thus the primary ion signal (H_3O^+ , 19.0178 Th) varied from a few hundred to few thousand cps and the isotope of the second water cluster ($H_2^{18}OH_2OH_3O^+$, 57.0432 Th) was around $10^4 - 10^5$ cps. The much lower signal at H_3O^+ was due to a high-pass band filter, that removes most of the ions < 35 Th (Krechmer et al., 2018). The mass resolving power of the LTOF mass analyzer was 12000-13000 for the whole campaign. Data were recorded using a time resolution of 10 s. Zero air was produced with a built-in active carbon filter and background was measured hourly except during December 15 – 17 due the malfunctioning of the zero air pump.

2.2.3. Iodide

Another deployed instrument was a time-of-flight chemical ionization mass spectrometer (TOF-CIMS, Tofwerk AG/Aerodyne Research, Inc.), equipped with iodide (Γ) reagent ion chemistry– later referred to as "Iodide". While the molecules could be detected as deprotonated species or as adducts with Γ , we restricted the analysis in this work to ions containing only an iodide adduct, which guarantees detection of the parent organic compounds without substantial fragmentation. Iodide TOF-CIMS has been described previously and has high sensitivity towards (multifunctional) oxygenated organic compounds (Iyer et al., 2017; Lee et al., 2014). The instrument was operated at 1 LPM reagent flow rate into the Ion-Molecule Reaction (IMR) chamber of the instrument. Iodide ions were generated from methyl iodide (CH₃I) using a polonium (Po-210) source. Sample air was drawn to the instrument using 1 m long PTFE tubing (6 mm o.d, 4 mm i.d.) with the flow rate of 2 LPM. The IMR was temperature controlled at 40°C and operated at a nominal pressure of 100 mbar. The instrument, equipped with an HTOF mass analyzer, was configured to measure singularly charged ions from 1 to 1000 Th with a mass resolving power and time resolution of 4000 – 5000 and 10 s, respectively.

2.2.4. Amine/Nitrate

Two chemical ionization atmospheric pressure interface time-of-flight mass spectrometers (CI-APi-TOF, Tofwerk AG/Aerodyne Research, Inc.) were also deployed (Ehn et al., 2014; Jokinen et al., 2012). The inlet was designed to minimize wall losses through the use of coaxial sample (10 LPM) and sheath flows (~ 30 LPM), in order to sample (extremely) low-volatile species which are easily lost to the walls. Two types of ionization schemes were utilized: the promising new amine reagent ion chemistry (Berndt et al., 2017, 2018) and the more commonly used nitrate chemistry - later referred to as Amine and Nitrate, respectively. The

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Amine has been shown to be sensitive towards a very wide range of OVOC, both closed shell species and peroxy radicals, from molecules with a few oxygen atoms all the way to HOM (Berndt et al., 2018). The Nitrate, on the other hand, has mainly been used for detection of HOM (Ehn et al., 2014).

Sample air was drawn to the instruments using a common 1 m long PTFE inlet line (19.05 mm o.d, 16 mm i.d.) with the flow rate being ~ 20 LPM (~ 10 LPM for each mass spectrometer). Nitrate (NO₃⁻) ions were formed from nitric acid (HNO₃) using an X-ray source while protonated butylamine (C₄H₁₂N⁺) ions were produced using butylamine with a 7.5 MBq Am-241 source. NO₃⁻ or C₄H₁₂N⁺ ions enter the ion reaction zone together with a clean sheath air flow, concentric with the sample flow, and the two do not mix turbulently. The ions are then guided into the sample flow by an electrical field. The residence time in the IMR was ~ 200 ms. The main reagent ions were NO₃⁻ (m/z 62), HNO₃NO₃⁻ (m/z 125) and (HNO₃)₂NO₃⁻ (m/z 188) for the Nitrate and C₄H₁₂N⁺ (m/z 74) for the Amine. Both instruments were equipped with LTOF mass analyzers providing a mass resolving power of 9000 – 10000.

2.3. Calibration of the mass spectrometers

In order to estimate absolute concentrations of all detected molecules, each instrument's signals, using an averaging period of 15 min, were normalized to the reagent ion signals (to eliminate the influence of changes affecting all signals in the instruments, e.g., due to degrading response of the detector) followed by multiplication with a scaling factor. The reagent ion quantity used for normalization is described below, separately for each instrument. Normalized ion count rates are reported as normalized cps, ncps. The scaling factors were derived differently for each instrument (details provided below). For Iodide, Nitrate and Amine, the same factor was used for all ions in the spectrum, while for the PTR instruments the factors were different depending on the type of molecules (e.g., VOC or OVOC). For the PTR instruments and the Iodide, a duty cycle correction was applied to compensate for mass-dependent transmission due to the orthogonal extraction of the mass analyzers. Finally, we emphasize that the scaling factors should not be compared between instruments as a measure of sensitivity, since multiple factors impact these values, including e.g., the specific normalization approach and the chosen extraction frequency of the mass analyzers.

The PTR-TOF was calibrated twice using a calibration unit consisting of a calibration gas mixture of 16 different VOC (Apel-Riemer Environmental Inc., USA) that was diluted with clean air purified by a catalytic converter (1.2 LPM of zero air, and 8 sccm of standard gas), producing VOC mixing ratios of around

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7 ppb (parts per billion) (Schallhart et al., 2016). Sensitivities were calculated to be 12.31 ncps ppb⁻¹, 27.92 ncps ppb⁻¹ and 30.51 ncps ppb⁻¹ based on the concentration of the monoterpenes, MVK (methyl vinyl ketone) and m-/o-xylenes, respectively. PTR-TOF signals were normalized using the sum of the first primary ion isotope at m/z 21.0221 and the first water cluster isotope at m/z 39.0327 (e.g., Schallhart et al., 2016). According to common practice, the sensitivities above were scaled to correspond to a situation where the total reagent ion signal equaled 10^6 cps.

The Vocus was calibrated four times during the campaign using the same calibration gas mixture as used for the PTR-TOF. There was variability in the sensitivity during the campaign and therefore the uncertainty in the Vocus results are slightly larger than normal. Sensitivities were highest for acetone, maximum around 1800 cps ppb⁻¹ and around 650 cps ppb⁻¹ for monoterpenes. α -pinene concentration was retrieved using the authentic standard while the concentrations of the OVOC and $C_{10}H_{14}H^+$ were estimated using the calibration factor of the MVK and sum of m-/o-xylenes, respectively. MVK and m-/o-xylene sensitivity was around 1700 cps ppb⁻¹ and 700 cps ppb⁻¹, respectively. Vocus signals were normalized using the primary ion signal at m/z 19.0178 only, as the water clusters have a negligible effect on the ion chemistry inside the FIMR (Krechmer et al., 2018). Due to the high-pass filter that removes almost all the signal at m/z 19.0178, we do not report the normalized sensitivities (i.e., in ncps ppb⁻¹) for the Vocus, in order to avoid direct comparisons with the PTR-TOF. Instead, the sensitivities above are given without normalization, although a normalization was used for the final data.

The Iodide was calibrated twice during the campaign (December 15 and 23) by injecting known amounts of formic acid to the instrument. Due to unknown reasons, the response of the Iodide decayed throughout the campaign, and therefore only data measured before December 17 was included. While normalization should compensate for this type of behavior, this particular instrument utilized a time-to-digital converter (TDC) acquisition card, which meant the primary ion peak was heavily saturated. Lacking any isotopic signatures for Γ , we found that utilizing a region of the rising edge of the Γ signal (126.5-126.65 Th) provided a reasonable correction to our data. The sensitivity without normalization was 1.0 cps ppt⁻¹ for formic acid, and following the normalization, this sensitivity was applied for all ions throughout the period where Iodide data was included in the analysis. We acknowledge that this brings with it a large uncertainty, as the

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Iodide has sensitivities ranging over a few orders of magnitude depending on the specific molecule (Lee et al., 2014).

Standards for OVOC compounds measurable by the Nitrate are still lacking, and this instrument was therefore not directly calibrated during the campaign. However, to be able to roughly estimate concentrations, a calibration was inferred by assuming that the molar yield of HOM, i.e., molecules with six or more oxygen atoms, during α-pinene ozonolysis experiments was 5%. Different values have been reported for the HOM yield in this system, ranging from slightly above to slightly below 5% (Ehn et al., 2014; Jokinen et al., 2014, 2015). Clearly such an approach yields large uncertainties, and we estimated it here to roughly \pm 70%. Earlier work with more direct calibrations reported an uncertainty of \pm 50% (Ehn et al., 2014), and the added 20 p.p. in this work reflects the increased uncertainty in scaling the sensitivity based on expected HOM yields. This method requires knowledge on the wall loss rate of HOM in the COALA chamber, which was estimated to be $1/300 \text{ s}^{-1}$ in our study. The resulting calibration coefficient was $2 \cdot 10^{10}$ molecules cm⁻³ ncps⁻¹, which is similar as in previous studies (Ehn et al., 2014; Jokinen et al., 2012). As for Nitrate, the Amine was also not calibrated directly, and in order to achieve an estimate of the concentrations measured by this instrument, we scaled the sensitivity of the Amine to match that of the Nitrate for specific HOM dimers (C₁₉H_{28/30}O₁₂₋₁₇ and C₂₀H_{30/32}O₁₂₋₁₇ ₁₇), which were found to correlate very well between the two instruments (as described in more detail in the Results section). This approach gave a calibration factor of 6 · 10⁸ molecules cm⁻³ ncps⁻¹, with similar uncertainty estimates as for the Nitrate. In the CI-APi-TOFs, the calibration factor is generally close to 1010 molecules cm⁻³ ncps⁻¹, but as discussed later in section 3.1, the Amine reagent ion was considerably depleted during the experiments, which led to the relatively low calibration factor. As mentioned earlier, the scaling factors should not be compared directly between instruments. The lower value for the Amine is a result of the normalization rather than an indication of higher sensitivity. This reagent ion depletion also means that the most abundant species were most likely no longer responding linearly to concentration changes, and therefore their concentrations can be off by an order of magnitude or more.

2.4. Instrumental limitations and considerations

In this section we aim at highlighting some of the limitations involved when characterizing and quantifying OVOC measured by online mass spectrometers. The list below is not exhaustive, but addresses several issues that are relevant for the interpretation of our results.

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2.4.1. Mass resolving power

One major limitation for all of the mass spectrometers described above, is the mass resolving power, ranging from 4000 to 14000. Even though the new generation of LTOF mass analyzers with higher resolving power can enhance the separation of measured ions, it remains challenging to accurately identify and deconvolve the elemental composition of many ions. Indeed, it is common for one CIMS mass spectrum to include more than 1000 different ions. For high-resolution (HR) peak identification and separation, firstly one needs to generate a list of ions, i.e., "peak list". Its construction can be time-consuming even if only based on one single spectrum, and once conditions change, different ions may appear. For measurements lasting weeks or months, it is nearly impossible to assure that all ions are correctly identified and fitted. If the peak list contains too few ions compared to reality, signals from non-fitted ions will assign the adjacent ions with artificially high signals. On the contrary, if too many closely lying ions are included in the peak list, even small errors in the mass axis determination can cause signal to be fitted to specific ions even though their signals are non-existent. In such extreme cases, with closely overlapping ions, traditional HR analysis becomes impossible.

While less selective detection techniques can sound more useful to monitor and characterize OVOC, spectra acquired using such ionization techniques (e.g., PTR, Iodide or Amine), pose a significant challenge for data analysis and may ultimately provide even less useful information. Statistical analysis techniques can be used in order to better constrain the uncertainties associated with peak fitting, as recently proposed (Cubison and Jimenez, 2015; Stark et al., 2015). These previous studies pointed out that the uncertainties related to the peak fitting can become significant if the overlapping peaks are separated by less than a full-width at half-maximum (Cubison and Jimenez, 2015). This is very often the case for CIMS instruments, and the more the ions overlap, the larger the uncertainty is. Peak fitting becomes increasingly problematic as molecular masses increase, since the number of potential ions increases dramatically with mass.

2.4.2. Ionization, declustering and fragmentation

The response of a mass spectrometer to a certain compound is to first approximation a result of two factors: the ionization probability of the neutral molecule, and the detection probability of the formed ion. The ionization process is largely controlled by the stability of the products compared to the primary ions, whether a question of adduct formation or (de)protonation processes. Different reagent ion chemistries have been

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321 Hyttinen et al., 2015, 2018; Iyer et al., 2016). While a neutral molecule can bind to a reagent ion at the collision 322 limit, the adduct can undergo collision-induced dissociation (i.e., declustering) during transport through the 323 interface to the high vacuum in the mass analyzer. Ultimately, the binding strength of the adduct and the energy 324 of the collisions in the mass spectrometer will define the survival probability of the ions. To address this issue, 325 procedures have been proposed, e.g., to probe the response of adducts to different collision energies (Isaacman-326 VanWertz et al., 2018; Lopez-Hilfiker et al., 2016), providing critical information on the sensitivity of the 327 instrument. 328 Similarly, to declustering, (de)protonated compounds can undergo fragmentation reactions where 329 molecular bonds are broken. For example, the detection of monoterpenes (C₁₀H₁₆) using PTR instruments often 330 shows equally large signals at the parent ion $(C_{10}H_{17}^{+})$ and at a fragment ion $(C_{6}H_{9}^{+})$. Also, Iodide adducts 331 have been shown to cause molecules to fragment, as in the case of peroxy acids decomposing to carboxylate 332 anions (Lee et al. 2014). Both declustering and fragmentation processes are associated with the optimization 333 of the voltages of each instrument, which is performed by the instrument operator (Breitenlechner et al., 2017; 334 Krechmer et al., 2018; Lopez-Hilfiker et al., 2016). While using voltage scans to probe such processes is 335 possible, and even desirable, performing, interpreting and utilizing the results, across the mass spectrum and across different conditions, remains challenging and has only been utilized in a few studies to date (Isaacman-336 337 VanWertz et al., 2018; Lopez-Hilfiker et al., 2016). 338 2.4.3. Quantification 339 For quantification, the instrument sensitivity is generally determined via calibration standards, while a 340 background level was measured by zero air. The challenges involved in these procedures are highly dependent 341 on the type of compounds to be quantified. As an example, we discuss three kinds of molecules with different 342 volatility: VOC, SVOC and ELVOC. 343 (a) VOC: volatile species are relatively easy to quantify since they can be contained in gas bottles or easily 344 evaporated from standard samples in known quantities. Their responses are also fast due to negligible 345 adsorption/evaporation from walls. 346 (b) SVOC: Many semi-volatile organic compounds (SVOC) are commercially available and can be 347 evaporated in known amounts from liquid standards into the gas phase. However, the nature of SVOC

studied computationally in recent years, successfully reproducing several observations (Berndt et al., 2017;

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results in both condensed and gas phases for these species, meaning that once clean air is introduced, the signal of SVOC will often show a gradual decay over minutes, or even hours, due to evaporation of the 'leftover' from surfaces in the inlet lines and the inlet itself (Pagonis et al., 2017). The procedure to determine the "correct" blank is not trivial, and the blank will look different depending on whether it is done at the entrance of the instrument or at the sampling inlet, and depending on the duration of the blank measurement itself. Another related challenge for SVOC quantification is that temperature fluctuations of a few degrees may cause net evaporation (temperature increasing) or condensation (temperature decreasing) of SVOC from sampling lines and the inlet. (c) ELVOC: For ELVOC, finding standard compounds for calibration remains extremely difficult. Most organic compounds, including hydroperoxide or acid, with such low volatility are likely to decompose before evaporating. Thus, their quantification is often inferred from other similar compounds. For the nitrate CI-APi-TOF, sulfuric acid is often used for calibration, by forming it in-situ from SO₂ (Kürten et al., 2012). This is, to some extent, a similar approach as we took for the Nitrate in this work, scaling it to the estimated HOM yield, as both methods require knowledge of formation rates from the initial precursors and loss rates of the formed compound of interest. Other studies have used permeation sources of perfluorinated carboxylic acids, which are semi-volatile, yet found to bind strongly to nitrate ions. However, while the calibration is complicated, the blank measurements are often not even needed, for exactly the same reasons. Whatever contaminants might be present in the system, most are irreversibly lost to instrument surfaces and unable to evaporate into the gas phase due to the extremely low vapor pressures. Potential oxidation processes occurring inside the mass spectrometer may be an exception, but to our knowledge, this has not been reported as a large concern for ELVOC. In addition to the list above, the response of an instrument to specific molecules may vary according to the conditions at which they were sampled. Temperature (change) was listed as one consideration, and water vapor, or relative humidity (RH), is another important limitation for several mass spectrometers described above (Breitenlechner et al., 2017; Krechmer et al., 2018; Kürten et al., 2012; Lee et al., 2014; Li et al., 2018). For chemical ionization techniques, the water vapor can either compete with the OVOC ionization, leading to decrease of the sensitivity, or stabilize the adduct resulting in an increase of the sensitivity. Alternatively, if a

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compound forms a very stable complex, it may have an adduct formation efficiency that is independent of water vapor. If the sensitivity is RH-dependent, calibrations and blanks should optimally be performed at the same RH as the sampling in order to be representative. This, in turn, may cause considerable practical challenges for both RH control and calibration and blank cleanliness.

In summary, recent computational and experimental work has shown that many approaches exist for optimizing the ability of CIMS instruments to quantify OVOC, including different blanks, calibration methods, voltage scans, etc. However, all these approaches are very rarely utilized in a single study, simply due to the immense time and effort required, both during the experiments and during the data analysis, where the results of all steps need to be incorporated. Ultimately, each study needs to prioritize between producing larger amounts of data (i.e., performing more measurements) with less capability for detailed quantification, or to produce a smaller amount of data with more accurate quantification.

3. Results and discussion

We applied our five CIMS instruments at the COALA chamber over a period of nearly one month, where we tried to provide different types of atmospherically relevant oxidation conditions for α -pinene. With such a high variability in conditions, we compared signals between the mass spectrometers more robustly, even though certain limitations were inevitable. For example, it is often the case that mass spectra will show some signal at almost every mass, which can be due to multiple reasons, and it is important to separate when the signal is truly from the sampled air and not from some internal background or contamination. Similarly, one needs to assess whether the instrument is measuring the majority of the species with the same elemental composition, or only detecting a small subset of those compounds due to specific selectivity for one isomer. In addition, an instrument may be able to detect a certain molecule, but the resulting signal remains unreliable. This may be the case if the sensitivity is extremely low for the molecule, or if the peak is close to a much larger unrelated signal, which will create large interferences when performing HR fitting. In both cases the signal is likely to be influenced by different types of noise.

First, we performed correlation analyses in order to identify signals which were physically meaningful. We conducted the analysis with the whole dataset (a total of ~ 1000 ions in each instrument) rather than selectively focusing on individual ions. This comprehensive approach utilized more data, but also resulted

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in larger uncertainties as not all fitted ions could be validated for all CIMS. From the correlation analysis we identified when two instruments agree, i.e., observing identical elemental compositions and having a similar temporal behavior, concerning some group of compounds. From a subsequent absolute comparison, we estimated which chemical ionization method was likely to be detecting a certain group of compounds more efficiently, in cases where the correlation between instrument was high, or which CIMS instruments were detecting a larger number, or the more abundant, isomers at a certain composition, in cases where the correlation between instrument was low.

3.1. Instrument comparisons: correlations

3.1.1. Low pressure ionization mass spectrometers

Peak fitting was performed utilizing the Igor-based Tofware or Matlab-based Toftools software (Junninen et al., 2010) for ion mass-to-charge up to \sim 600 Th, depending on the mass spectrometers. To select which ions to fit (i.e., include in the peak lists), both the exact masses and the isotopic distributions were used as criteria. A Pearson correlation coefficient R, was calculated between molecules having the same elemental composition measured by the different instruments. As a practical example, the time series of $C_{10}H_{16}O_3$ measured by Vocus and Iodide are shown in Figure 1C, and the time series correlation for this compound between the two instruments was R = 0.85. For later comparisons we will use R squared, and in this case $R^2 = 0.73$. For Iodide, the data set covered only the first half of the campaign, but the other instruments covered nearly the whole period. This includes a wide variety of conditions, with and without NO_x , and therefore high correlations are very suggestive of two instruments measuring the exact same compound(s) at that specific elemental composition. However, as an increase in α -pinene is likely to increase almost all measured OVOC signals to some extent, low positive correlations can arise artificially and should not be over-interpreted. Due to the selectivity and the sensitivity of the ionization methods, all ions were not observed in all the different instruments, and thus only a certain fraction of the identified compounds can be compared between mass spectrometer.

Figure 2 shows the correlation analysis for the low-pressure chemical ionization mass spectrometers, with marker size scaled by R^2 . In those figures, the abscissa represents the measured m/z of the compounds and the y-axis their mass defect, which is calculated as the exact mass of the compound minus the mass rounded

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to the closest integer (Schobesberger et al., 2013). For example, the mass of $C_{10}H_{16}O_3$ is 184.110 Da and the mass defect is + 0.110 Da. The contribution of the reagent ions has been removed in the different Figures. A mass defect diagram helps to separate the molecules into two dimensions and allows some degree of identification of the plotted markers.

As expected, the PTR-TOF and the Vocus are strongly correlated for compounds with low (0-3) oxygen number (Figure 2A). Contrariwise, only a few compounds were identified by the PTR-TOF and the Iodide with a fairly good correlation (i.e., $R^2 > 0.5$). The correlating compounds included small acids such as formic and acetic acid. As discussed earlier, the inlet of the PTR-TOF is not well designed to sample OVOC having low volatility, which explained the lack of correlations for larger and more oxidized products between the PTR-TOF and the Nitrate CI-APi-TOF. The molecules with the lowest correlations ($R^2 < 0.2$) were not included in the plots, as the intention is to show regions where instruments agree. If an ion is included in a peak list, it will always be fit, and thereby a value of $R^2 > 0$ is always expected, filling markers throughout the MD-mass space.

In addition to VOC, the Vocus was able to measure a large range of OVOC (m/z 150-300 Th) as revealed in Figure 2B, displaying a very good correlation with species identified by the Iodide. Indeed, most of the identified compounds have $R^2 > 0.7$. As noted earlier, several different experimental conditions were tested (Figure 1), and these high correlations indicate that both instruments were likely sensitive to the same compounds. In other words, as good correlation was seen in this mass range for nearly all compositions, the Iodide and the Vocus did not seem to be strongly impacted by the exact chemical conformation of the organic compounds. Interestingly no dimers (m/z > 300 Th) were observed with the Vocus, which suggests some potential limitation of the instrument. As a result, very limited correlation was observed between compounds measured by the Vocus and the Amine or Nitrate CI-APi-TOFs. The two main exceptions being $C_5H_6O_7$ (178.011 Da) and $C_7H_9NO_8$ (235.033 Da). Note that the latter is less clear, as the correlation is nearly identical between three instruments (Nitrate, Vocus and Iodide). The lack of correlation was not only due to lack of ion transmission at higher masses in the Vocus, since the instrument was able to detect some ions up to 400 Th, including $C_{10}H_{30}O_5Si_5H^+$ and $C_{19}H_{29}O_6NH^+$. One possibility was that since the compounds above ~ 300 Th were likely to contain hydroperoxides, or in the case of dimers, organic peroxides, the ions might may have fragmented before detection in the Vocus, either during the protonation or due to the strong electric fields in

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the Vocus FIMR. In the case of HOM monomers with more than 7 oxygen atoms, an additional limitation comes from more abundant and closely overlapping ions in the spectra, impacting the accurate fitting of these ion signals in the Vocus. From our data set, it was not possible to determine the exact cause(s) for this lack of sensitivity for larger molecules in the Vocus, but it is possible that changes in instrument operating conditions can extend the range of molecules detectable using the Vocus in future studies.

As shown in Figure 2C, the Iodide was capable of measuring ions with larger masses (i.e., m/z > 300Th) indicating the detection of more complex (e.g., dimers) and oxygenated compounds than the Vocus. This was the case in spite of the lower flow rate for the Iodide than the Vocus, and thus less optimal for sampling of low-volatile species (Table 1). The Iodide seemed to have the widest detection range of the mass spectrometers deployed in this study, showing high correlation with other instruments for organic molecules from C_1 (like formic acid) to C_{20} , as long as the molecules had at least two oxygen atoms. This is in line with earlier findings that the Iodide is sensitive to most species that are polar or have polarizable functional groups (Iyer et al., 2017; Lee et al., 2014). However, the correlation with the CI-APi-TOFs was still somewhat limited $(R^2 < 0.7)$ for HOM monomers and dimers. One reason may have been that these HOM contain peroxy acid functionalities, which have been shown to undergo reactions in the Iodide TOF-CIMS (Lee et al., 2014). In this work, we only analyzed the ions containing Γ as these were believed to be the ones where the parent molecule remained intact. Another reason for lower correlation was the fact that Γ is less selective than other ionization methods resulting in many overlapping peaks at the same m/z and ambiguous peak fitting (Lee et al., 2014; Stark et al., 2015, 2017), similar to the case in the Vocus. This means that although the Iodide and/or the Vocus might be able to charge a specific molecule, and it would not fragment before detection, the ion may remain unquantifiable due to highly ambiguous peak fitting as a result of multiple overlapping signals.

3.1.2. Atmospheric pressure interface mass spectrometers

Figure 3 shows similar comparisons as in Figure 2, now for the Nitrate (Figure 3A) and the Amine (Figure 3B). Interestingly, these two instruments show excellent correlation ($R^2 > 0.9$) for dimeric products (molecules within 350-500 Th), but showed mostly low correlations ($R^2 < 0.6$) with other instruments in the monomer range. The Nitrate had some agreement with the Iodide for certain monomer compounds, but in the HOM monomer range where the Nitrate generally saw its largest signals (C_{10} molecules with 7 to 11 oxygen atoms; Ehn et al., 2014), none of the other instruments showed strongly correlating signatures.

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Despite showing signals at almost all OVOC, the Amine presented low correlations for all OVOC except the dimers. In the Amine the reagent ion was greatly depleted due to the relatively high signals (Figure 4), likely leading to a non-linear response for most of the OVOC, apparently with the exception of the HOM dimers. It may be that the amine reagent ion formed extremely stable clusters with these dimers, and thus any collision involving these dimers with the reagent ion (regardless of whether already clustered with an OVOC) in the IMR lead to an amine-dimer cluster. While the Amine showed very low correlation with the other instruments for most molecules, it has been demonstrated to be an extremely useful detector of both radicals and closed shell OVOC under very clean, low-loading flow tube experiments (Berndt et al., 2017, 2018). In other words, it can provide information on a wide variety of OVOC, but to obtain quantitative information, the amine CI-APi-TOF has to be used in very diluted system (with very clean air) and at low loadings. Determining more explicitly the limitations requires further studies, but as a rough approximation, the typical CI-APi-TOF sensitivity of ~10¹⁰ molecules cm⁻³ ncps⁻¹ means that when sampling detectable molecules at 10¹⁰ molecules cm⁻³ (~0.4 ppb), these molecules will have ion signals of equal abundance as the reagent ions. Consequently, once the concentration of measurable molecules exceeds roughly 100 ppt, the CI-APi-TOF may no longer be an optimal choice. For the Nitrate CI-APi-TOF, which mainly detects HOM with short lifetimes due to their low volatilities, this has rarely been a limitation, but for less selective reagent ions, like amines, this can be an important consideration.

3.2. Instrument comparisons: concentration estimates

Concentrations of the identified compounds were estimated for all the different instruments, as described in section 2.6. It should be noted that no separate inlet loss corrections were applied. The estimations for the results of PTR-TOF and the Vocus are the most reliable as both instruments were calibrated using authentic standards with a proven method, while larger uncertainties in the total measured concentrations are expected for the Iodide and the CI-APi-TOFs.

With around 1000 identified ions for each instrument, except for the PTR-TOF, we decided to focus our attention in this section on a few particular compound groups: the most abundant C_{10} -monomers (i.e., $C_{10}H_{14/16}O_n$), C_{10} -organonitrates ($C_{10}H_{15}NO_n$) and dimers ($C_{20}H_{32}O_n$). For the non-nitrate compounds, the concentrations were measured during a steady-state conditions of December 9 from 15:30 to 23:00 with $[O_3]$ = 25 ppb and $[\alpha$ -pinene] = 100 ppb) during period I (Figure 1 in blue). The organonitrate concentrations were

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compared using steady-state conditions of December 20 from 02:45 to 07:45 with $[O_3] = 35$ ppb, $[\alpha$ -pinene] = 100 ppb and NO = 0.5 ppb, during period IV (Figure 1 in purple). Figures 5A-D show the concentrations of the selected species as a function of oxygen number in the molecules. While we again emphasize that all the concentrations were only rough estimates, these plots painted a similar picture to the correlation analysis, as described in more detail in the next paragraphs.

Focusing first on the non-nitrate monomers (Figures 5A-B), for compounds with zero or one oxygen atoms, the PTR-TOF agreed well with the concentration estimated by the Vocus, while molecules with more than two oxygen atoms were already close to, or below, the noise level of the PTR-TOF. In contrast, as the number of oxygen atoms in the molecule reached two or more, the Iodide signal increased, and for most compounds showed concentrations similar to the Vocus. These two instruments agreed on concentration estimates fairly well all the way up to an oxygen content of around 9 oxygen atoms, where the measured signals were close to the instruments' noise levels. However, when comparing to the Nitrate, which is assumed to have good sensitivity for HOM with 7 or more oxygen atoms, the concentrations suggested by the Vocus and Iodide for the O_7 and O_8 monomers were very high. We preliminarily attributed this to an over-estimation of the concentrations of HOM by these two instruments, possibly due to higher sensitivities towards these molecules as compared to the compounds used for calibration (i.e., MVK). We also did not correct for potential backgrounds using the blanks for the Iodide, although measured, since the variability in the blank concentrations (see also discussion in section 2.4) was large enough to cause artificially high fluctuations in the final signals. Therefore, we opted to not include such a correction, but also note that even if half the signal at a given ion was attributable to background in the Iodide, then it would only have a small impact on the logarithmic scales used in Figure 5. Other possible reasons for this discrepancy was that the Iodide and Vocus were able to detect isomers that the Nitrate was not, or that the Nitrate sensitivity was under-estimated. However, considering that the Nitrate HOM signal was scaled to match a 5% molar HOM yield, it was unlikely that the HOM concentrations can be considerably higher than this. Other estimated parameters involved in the formation and loss rates of HOM also had uncertainties, but we did not expect any of them to be off by more than 50%. This concentration discrepancy thus remained unresolved, and will require more dedicated future studies.

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Finally, the quantities estimated using the Amine are significantly lower (1-2 orders of magnitude) for all monomers when comparing to the other instruments. This was presumably related to the titration of the reagent ion, which meant that the majority of charged OVOC will undergo multiple subsequent collisions with other OVOC, potentially losing their charge in the process. The Nitrate had, as expected, very low sensitivity towards less oxygenated compounds, and its highest detection efficiency for HOM (i.e., molecules with at least 6 oxygen atoms).

The organonitrate comparison in Figure 5C suggested that both the Vocus and the Iodide were efficient at detecting these compounds, as both instruments agreed well ($R^2 > 0.7$) for C₁₀-organonitrates with 5 to 10 oxygen atoms. While organonitrates have been detected before using the Iodide (Lee et al., 2016), this was the first observation that also the Vocus can detect such compounds efficiently. For larger oxygen content, the Nitrate again seemed to be most sensitive, showing clear signals above 10 oxygen atoms, where the previous instruments were already close to noise levels. The Amine seemed worse at detecting organonitrates compared to non-nitrate monomers.

Neither of the PTR instruments were able to detect any dimers in this study. The Amine and the Nitrate were able to quantify the widest range of HOM dimers, while the Iodide was able to detect less oxidized dimers (Figure 5D). Based on the concentration estimates, the Amine's detection range also extended to less oxidized dimers than the Nitrate, as has already been shown by Berndt et al. (2018). Dimers measured by the Iodide were more abundant than the ones detected by the Amine, but already from the monomer comparisons we speculated that the Amine might be underestimating concentrations while the Iodide might be overestimating them. With the data available to us, we can only speculate on the relative sensitivities of the instruments able to detect dimers, especially with the Vocus providing no support to the comparison.

One aspect lending credibility to the Amine dimer data, in addition to the good time series correlation with the Nitrate, was the odd-even oxygen atom patterns visible both in the Amine and Nitrate data. Such a pattern is to be expected, since the 32 hydrogen atoms in the selected dimers indicates that they have been formed from RO₂ radicals where one had 15 hydrogen atoms (which is what ozonolysis will yield, following OH loss) (Docherty et al., 2005; Lee et al., 2006; Ziemann and Atkinson, 2012), while the second RO₂ had 17 hydrogen atoms (which is the number expected from OH oxidation of an alkene where OH adds to the double bond). The first RO₂ from ozonolysis had 4 oxygen atoms, and further autoxidation will keep an even number

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of oxygen atoms, while the opposite was true for the OH-derived RO_2 which started from 3 oxygen atoms. In other words, the major dimers from this pathway should contain an odd number of oxygen atoms after combination. In the case of $C_{20}H_{30}O_n$ dimers, mainly formed from two ozonolysis RO_2 , the pattern was expected to show peaks at even n, which is also the case (not shown).

Such odd-even patterns for the oxygen content was not visible in the Iodide, but the reason remained unknown. It was possible that the dimers detected by the Iodide might be formed via other pathways, where such a selectivity did not occur. This topic should be explored further in future studies, since dimers formed form the oxidation of biogenic compounds are important for new-particle formation, and it is therefore critical to accurately identify and quantify the formation and evolution of different types of dimers. To date, both dimers measured by Iodide (Mohr et al., 2017) and Nitrate (Tröstl et al., 2016) have been found to be important for particle formation from monoterpenes.

3.3. Performance in detecting oxygenated species

Figure 6 summarizes our results and depicts the performance of each mass spectrometer in detecting monomer and dimer monoterpene oxidation products. Molecules of $C_{10}H_{16}O_n$, $C_{10}H_{15}NO_n$ and $C_{20}H_{30}O_n$ were provided as examples. We emphasized that the oxygen content alone was not the determining factor for whether a certain type of mass spectrometer will detect a compound, but we utilized this simplified representation in order to provide an overview of the performances of the different chemical ionization schemes. The results were primarily based on the correlation analysis from section 3.1, and as apparent from the y-axis, this comparison was only qualitative. However, our aim was to provide an easy-to-interpret starting point, especially for new CIMS users wanting to compare different available techniques.

For monomer compounds without N-atoms, shown in Figure 6A, the PTR-TOF was limited to the detection of VOC, while the Vocus was additionally able to measure a large range of OVOC, up to at least 5-6 oxygen atoms. The Iodide detected OVOC with oxygen content starting from \sim 3 atoms, but did not seem to efficiently observe HOM monomers (i.e., $C_{10}H_xO_{>7}$). While being a very promising instrument for a broad detection of OVOC, the performance of the Amine was limited in our study due to a significant drop of the reagent ion to \sim 40% of the total signal. Therefore, the Amine was marked with a shaded region rather than a line, with the lower limit based roughly on its usefulness under the conditions we probed, while the upper limit was an estimate based on findings in a cleaner system with low loadings (Berndt et al., 2018). Finally, the

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Nitrate was mainly selective towards HOM. The detection and quantification of monomeric OVOC containing 5 to 8 oxygen atoms remained as the most uncertain, since there were inconsistencies in both concentration and correlation between the Nitrate, measuring the more oxygenated species, and the Vocus/Iodide, which detected the less oxidized compounds.

In Figure 6B, the suitability for the different instruments was plotted for organonitrate monomers. The Vocus efficiently detected the less oxidized organonitrates, while the Iodide displayed a good sensitivity for the same compounds, with the exception of the least oxygenated ones. For larger number of oxygens, the Nitrate again seemed the most suitable method. For dimers (Figure 6C), neither of the PTR techniques showed any ability to detect these compounds in our study. We did not extend the lines all the way down to n = 0 for the compounds, as it was still possible that these methods can be able to detect the least oxidized and most volatile C₂₀ compounds, which might not have been present during our experiments. The Iodide showed some correlation with the Nitrate, but had good signals mainly in the range of dimers with 4 to 8 oxygen atoms. The Amine and Nitrate correlated well for the most oxidized dimers, suggesting good suitability for dimer detection of HOM dimers. The Amine concentrations stayed high, with the expected odd-even pattern in oxygen number, even at lower oxygen content than the Nitrate, and therefore the suitability extended further towards lower Oatom contents. Again, the shaded area was based on a combination of our findings and those of Berndt et al. (2018).

As a final test for each instrument, we estimated how much of the reacted carbon (in ppbC) the different mass spectrometers can explain. As shown in Figure 7, both the Iodide and Vocus seemed to capture most of the reacted carbon, within uncertainties. The concentration determined using the Vocus was overestimated, explaining more carbon than was reacted. Out of the largest contributors to the reacted carbon, pinonaldehyde ($C_{10}H_{16}O_2$) was not efficiently detected by Iodide, but otherwise most of the abundant molecules were quantified by both Vocus and Iodide. Any carbon lost by condensation to walls or particles would not have been quantifiable by any of the instruments in this study. While the Nitrate was calibrated with an assumption that it can measure 5% of the reacted α -pinene, it only detected less than a tenth of that amount. The reason was that the HOM it can detect were quickly lost to walls (or particles), and thus the gas-phase concentration was not equivalent to the branching ratio of the VOC oxidation reaction. In fact, and as revealed by the slow changes in the times series in Figure 7C, most of the carbon ultimately measured by the Nitrate

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was semi-volatile, as such compounds accumulated and reached higher concentration in the chamber, unlike HOM. Thus, while the Nitrate was able to detect a critical group of OVOC from an aerosol formation perspective, i.e., HOM, for carbon closure studies (Isaacman-VanWertz et al., 2017, 2018) it will be of limited use. This again highlights the need to first determine the target of a study before deciding which CIMS technique is the most useful. For the closure comparison in our study, the overestimations emphasized the need to perform calibration with an extensive set of OVOC, ideally with monoterpene-oxidation products, in order to better constrain the sensitivity of the products of interest. The study by Isaacman-VanWertz et al. (2018), as the only study to achieve full carbon closure during chamber oxidation of α -pinene by OH, also successfully utilized voltage scanning to determine sensitivities of each compound.

4. Conclusions

The primary goal of this work was to evaluate the performance of 5 chemical ionization mass spectrometers (PTR-TOF, Vocus PTR, Iodide TOF-CIMS, Amine CI-APi-TOF and Nitrate CI-APi-TOF) in the identification and quantification of a wide variety of products formed in the ozonolysis of α -pinene. In addition, we wanted to estimate the capabilities of the newly developed Vocus PTR in measuring OVOC species. By comparing the regions of coverage of the instruments across multiple experimental conditions (i.e., in different O₃, VOC, NO and OH radical concentrations) we demonstrated that current instrumentation captures nearly the entire range of OVOC, spanning from VOC to ELVOC. The PTR-TOF was only able to measure the most volatile compounds, while the Vocus appeared to be able to measure both VOC and most of the OVOC up to 5 to 6 oxygen atoms. In combination with the Iodide and Nitrate, most of the OVOC range can be measured. The Iodide showed good overlap with the Vocus for most SVOC with 3 to 5 oxygen atoms, while the Nitrate detected mainly products with 6 or more oxygen atoms. No dimer species were observed with either of the PTR instruments, which might be due to wall losses (likely at least for the PTR-TOF) and/or potential fragmentation in the instruments. The Amine CI-APi-TOF is a promising technique, as shown in earlier studies, but it likely requires low loadings in order to not titrate the reagent ion, limiting its utility for many chamber experiments and, potentially, atmospheric observations. The large uncertainties in attempting a quantification of the wide variety of species measurable with these mass spectrometers underline the urgent need of developing robust, simple and complete calibration methods in order to obtain a better estimation of

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performed by the user will ultimately impact the sensitivity, the selectivity, and the interpretability of the 652 653 results attainable from each instrument. 654 655 Data availability All data are available by contacting the corresponding authors. 656 657 Acknowledgments 658 659 This work was supported by the European Research Council (ERC-StG COALA, grant nr 638703). We 660 gratefully acknowledge Pasi Aalto, Petri Keronen, Frans Korhonen, and Erkki Siivola for technical support. 661 O.G. thanks Doctoral Programme in Atmospheric Sciences (ATM-DP) at the University of Helsinki for 662 financial support. O.P. would like to thank the Vilho, Yrjö and Kalle Väisälä Foundation. We thank the 663 tofTools team for providing tools for mass spectrometry data analysis. 664 Author contributions 665 666 M.R. and M.E. designed the experiments. Instrument deployment, operation, and data analysis were carried out by: M.R., P.R. J.E.K., O.P., Y.Z., L.H., O.G. C.Y., and M.E.; M.R., P.R., O.P., and M.E., interpreted the 667 668 compiled data set. M.R., P. R. and M.E. wrote the paper. All co-authors discussed the results and commented the manuscript. The authors declare that they have no conflict of interest. 669

the concentrations. Finally, it is important to underline that the experimental and analytical procedures

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Table. 1 Overview and characteristics of the mass spectrometers deployed during the campaign at the COALAchamber.

Instrument ^a	Ionization ^b	Resolving power ^c	Sampling flow rate	T in IMR ^d (°C)	Residence time in IMR	IMR pressure	Inlet length
	D .		(LPM)	-	HVIK	(mbar)	(m)
PTR-TOF	Proton	4500	0.8	60	100 μs	2.0	2
	transfer						
Vocus	Proton	12000	4.5	30	82 μs	1.0	1
	transfer						
Iodide	I- adduct	4500	2	40	94 ms	100	1
- 01- 01-0	C II NI ⁺		_	-			•
Amine	$C_4H_{12}N^+$	10000	10	Ambient	200 ms	Ambient	1
	adduct						
Nitrate	NO_3^-	9000	10	Ambient	200 ms	Ambient	1
	adduct						

instrument; ^c corresponds to the mass resolution of the instruments under the conditions used in this study. ^d
IMR = Ion-molecule reaction chamber, i.e. the region where sample molecules are mixed with reagent ions.
The IMR has a different design in each of the instruments, except for the Nitrate and Amine, which are identical.

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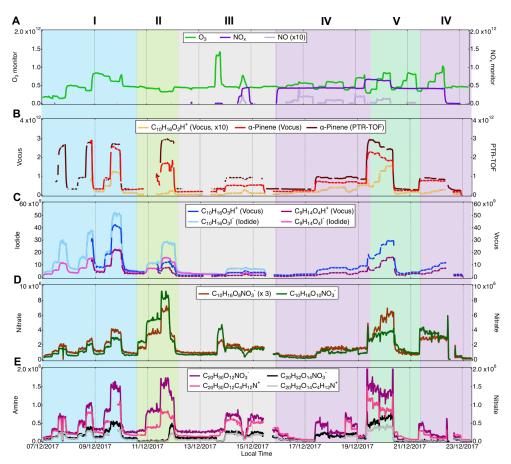


Figure 1: Campaign overview, including the concentration of O_3 , NO, NO_x (A) as well as α-pinene measured by the Vocus and the PTR-TOF and pinonaldehyde measured using the Vocus (B). Concentrations of pinonic and pinic acids (Vocus & Iodide) are presented in C, example of HOM monomers from Nitrate (D) and example of HOM dimers from Amine and Nitrate (E). Concentrations for all the gaseous species are in molecules cm⁻³, see text for details on quantification. The experiments were separated in 5 types: I: α-pinene $+ O_3$; II: α-pinene $+ O_3 + CO$ (as an OH scavenger); III: tests (NO₂ injection, H₂O₂ injection to generate HO₂); IV: α-pinene $+ O_3 + NO$ and V: α-pinene $+ O_3 + NO$ and V: α-pinene $+ O_3 + NO$ and C₁₀H₁₆O₈NO₃⁻ are scaled for clarity.

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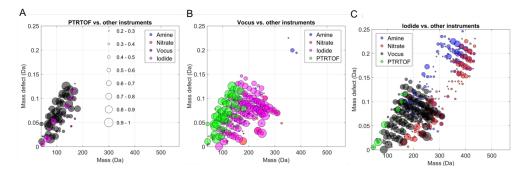


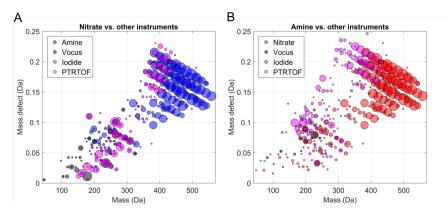
Figure 2: Mass defect plots showing the compounds for which time series correlation ($R^2 > 0.2$) was observed by the low-pressure chemical ionization mass spectrometers (A) PTR-TOF, (B) Vocus and (C) Iodide. Each circle represents a distinct molecular composition and the marker area represents the correlation (R^2 , legend shown in A) of the time series of that molecule between two different CIMS instruments. The color of each marker depicts the instrument against which the correlation is calculated.

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Figure 3: Mass defect plots showing the compounds for which time series correlation ($R^2 > 0.2$) was observed by the atmospheric-pressure chemical ionization mass spectrometers (A) Nitrate and (B) Amine. Each circle represents a distinct molecular composition and the marker area represents the correlation (R^2 , legend shown in Figure 2A) of the time series of that molecule between two different CIMS instruments. The color of each marker depicts the instrument against which the correlation is calculated.

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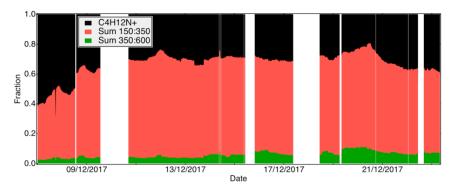


Figure 4: Contribution of the reagent ion, sum of ions from m/z 150 to 350 Th and sum of ions from m/z 350 to 600 Th to total ion count throughout the campaign for the Amine CI-APi-TOF. Only a negligible fraction of the signal was found below *m/z* 150 Th (excluding C₄H₁₂N⁺).

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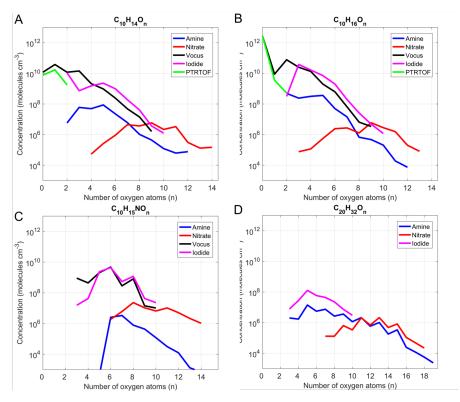


Figure 5: Estimated concentrations of the main α-pinene C_{10} -monomer oxidation products (A & B), C_{10} -monomer organonitrates (C) and α-pinene dimers (D) by the different mass spectrometers deployed in this study. The average concentrations were estimated when the system reached steady state in two experiments: without NO (panels A, B & D), December 9 (15:30-23:00) and with NO (panel C), December 20 (02:45 to 07:45). See text for more details. Data is plotted only for ions for which the average concentrations were higher than 3 times the standard deviation during the campaign.

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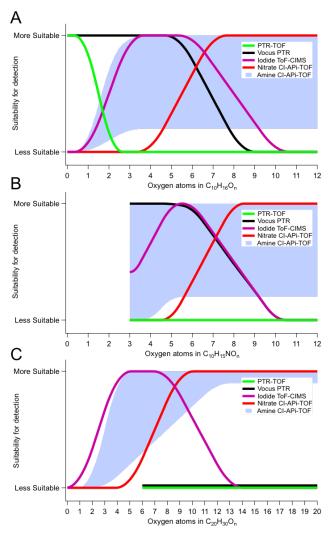


Figure 6: Estimated detection suitability of the different CIMS techniques for α -pinene and its oxidation products, plotted as a function of the number of oxygen atoms. Each panel symbolizes a compound group: monomers (A), organonitrate monomers (B) and dimers (C). The figures are indicative only, as none of the reagent ion chemistries are direct functions of the oxygen atom content in the molecules. See text for more details.

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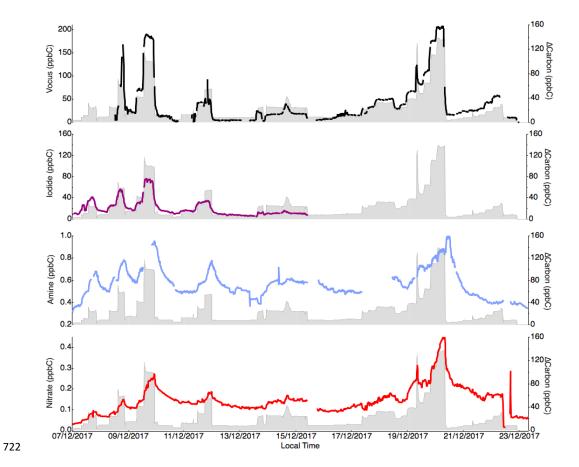


Figure 7: Concentration (in ppbC) of the sum of the compounds measured by each instrument (Vocus, Iodide, Amine and Nitrate) throughout the campaign, compared to the amount of reacted carbon through α -pinene oxidation. Large uncertainties remain in the quantification of the OVOC for all instruments, but it is clear that the Iodide and Vocus are able to measure a large fraction of the reacted carbon in the gas phase.

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728 REFERENCES

- 729 Albrecht, B. A.: Aerosols, Cloud Microphysics, and Fractional Cloudiness, Science, 245(4923), 1227-1230,
- 730 doi:10.1126/science.245.4923.1227, 1989.
- 731 Atkinson, R., Baulch, D. L., Cox, R. A., Hampson, R. F., Kerr, J. A., Rossi, M. J. and Troe, J.: Evaluated
- 732 Kinetic and Photochemical Data for Atmospheric Chemistry: Supplement VI. IUPAC Subcommittee on Gas
- 733 Kinetic Data Evaluation for Atmospheric Chemistry, Journal of Physical and Chemical Reference Data, 26(6),
- 734 1329–1499, doi:10.1063/1.556010, 1997.
- 735 Berndt, T., Richters, S., Kaethner, R., Voigtländer, J., Stratmann, F., Sipilä, M., Kulmala, M. and Herrmann,
- 736 H.: Gas-Phase Ozonolysis of Cycloalkenes: Formation of Highly Oxidized RO 2 Radicals and Their Reactions
- 737 with NO, NO₂, SO₂, and Other RO₂ Radicals, The Journal of Physical Chemistry A, 119(41), 10336–10348,
- 738 doi:10.1021/acs.jpca.5b07295, 2015.
- 739 Berndt, T., Herrmann, H. and Kurtén, T.: Direct Probing of Criegee Intermediates from Gas-Phase Ozonolysis
- 740 Using Chemical Ionization Mass Spectrometry, Journal of the American Chemical Society, 139(38), 13387–
- 741 13392, doi:10.1021/jacs.7b05849, 2017.
- 742 Berndt, T., Scholz, W., Mentler, B., Fischer, L., Herrmann, H., Kulmala, M. and Hansel, A.: Accretion Product
- 743 Formation from Self- and Cross-Reactions of RO 2 Radicals in the Atmosphere, Angewandte Chemie
- 744 International Edition, 57(14), 3820–3824, doi:10.1002/anie.201710989, 2018.
- 745 Bertram, T. H., Kimmel, J. R., Crisp, T. A., Ryder, O. S., Yatavelli, R. L. N., Thornton, J. A., Cubison, M. J.,
- 746 Gonin, M. and Worsnop, D. R.: A field-deployable, chemical ionization time-of-flight mass spectrometer,
- 747 Atmospheric Measurement Techniques, 4(7), 1471–1479, doi:10.5194/amt-4-1471-2011, 2011.
- 748 Breitenlechner, M., Fischer, L., Hainer, M., Heinritzi, M., Curtius, J. and Hansel, A.: PTR3: An Instrument for
- 749 Studying the Lifecycle of Reactive Organic Carbon in the Atmosphere, Analytical Chemistry, 89(11), 5824–
- 750 5831, doi:10.1021/acs.analchem.6b05110, 2017.
- 751 Crounse, J. D., McKinney, K. A., Kwan, A. J. and Wennberg, P. O.: Measurement of Gas-Phase
- 752 Hydroperoxides by Chemical Ionization Mass Spectrometry, Analytical Chemistry, 78(19), 6726–6732,
- 753 doi:10.1021/ac0604235, 2006.
- 754 Cubison, M. J. and Jimenez, J. L.: Statistical precision of the intensities retrieved from constrained fitting of
- 755 overlapping peaks in high-resolution mass spectra, Atmospheric Measurement Techniques, 8(6), 2333–2345,
- 756 doi:10.5194/amt-8-2333-2015, 2015.
- 757 Docherty, K. S., Wu, W., Lim, Y. B. and Ziemann, P. J.: Contributions of Organic Peroxides to Secondary
- 758 Aerosol Formed from Reactions of Monoterpenes with O₃, Environmental Science & Technology, 39(11),
- 759 4049–4059, doi:10.1021/es050228s, 2005.
- 760 Donahue, N. M., Robinson, A. L., Trump, E. R., Riipinen, I. and Kroll, J. H.: Volatility and Aging of
- 761 Atmospheric Organic Aerosol, in Atmospheric and Aerosol Chemistry, vol. 339, edited by V. F. McNeill and
- P. A. Ariya, pp. 97–143, Springer Berlin Heidelberg, Berlin, Heidelberg., 2012.
- 763 Ehn, M., Thornton, J. A., Kleist, E., Sipilä, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann,
- 764 R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S.,
- 765 Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B., Jørgensen, S., Kjaergaard, H. G.,
- 766 Canagaratna, M., Maso, M. D., Berndt, T., Petäjä, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop,
- 767 D. R., Wildt, J. and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature,
- 768 506(7489), 476–479, doi:10.1038/nature13032, 2014.
- 769 Gakidou, E., Afshin, A., Abajobir, A. A., Abate, K. H., Abbafati, C., Abbas, K. M., Abd-Allah, F., Abdulle,
- 770 A. M., Abera, S. F., Aboyans, V., Abu-Raddad, L. J., Abu-Rmeileh, N. M. E., Abyu, G. Y., Adedeji, I. A.,

Manuscript under review for journal Atmos. Meas. Tech.

Discussion started: 10 December 2018 © Author(s) 2018. CC BY 4.0 License.





- 771 Adetokunboh, O., Afarideh, M., Agrawal, A., Agrawal, S., Ahmadieh, H., Ahmed, M. B., Aichour, M. T. E.,
- 772 Aichour, A. N., Aichour, I., Akinyemi, R. O., Akseer, N., Alahdab, F., Al-Aly, Z., Alam, K., Alam, N., Alam,
- 773 T., Alasfoor, D., Alene, K. A., Ali, K., Alizadeh-Navaei, R., Alkerwi, A., Alla, F., Allebeck, P., Al-Raddadi,
- 774 R., Alsharif, U., Altirkawi, K. A., Alvis-Guzman, N., Amare, A. T., Amini, E., Ammar, W., Amoako, Y. A.,
- 775 Ansari, H., Antó, J. M., Antonio, C. A. T., Anwari, P., Arian, N., Ärnlöv, J., Artaman, A., Aryal, K. K.,
- 776 Asayesh, H., Asgedom, S. W., Atey, T. M., Avila-Burgos, L., Avokpaho, E. F. G. A., Awasthi, A., Azzopardi,
- P., Bacha, U., Badawi, A., Balakrishnan, K., Ballew, S. H., Barac, A., Barber, R. M., Barker-Collo, S. L.,
- 778 Bärnighausen, T., Barquera, S., Barregard, L., Barrero, L. H., Batis, C., Battle, K. E., Baumgarner, B. R.,
- Baune, B. T., Beardsley, J., Bedi, N., Beghi, E., Bell, M. L., Bennett, D. A., Bennett, J. R., Bensenor, I. M.,
- 780 Berhane, A., Berhe, D. F., Bernabé, E., Betsu, B. D., Beuran, M., Beyene, A. S., Bhansali, A., Bhutta, Z. A.,
- 781 Bicer, B. K., Bikbov, B., Birungi, C., Biryukov, S., Blosser, C. D., Boneya, D. J., Bou-Orm, I. R., Brauer, M.,
- 782 Breitborde, N. J. K., et al.: Global, regional, and national comparative risk assessment of 84 behavioural,
- 783 environmental and occupational, and metabolic risks or clusters of risks, 1990–2016: a systematic analysis for
- 784 the Global Burden of Disease Study 2016, The Lancet, 390(10100), 1345-1422, doi:10.1016/S0140-
- 785 6736(17)32366-8, 2017.
- 786 Glasius, M. and Goldstein, A. H.: Recent Discoveries and Future Challenges in Atmospheric Organic
- 787 Chemistry, Environmental Science & Technology, 50(6), 2754–2764, doi:10.1021/acs.est.5b05105, 2016.
- 788 Goldstein, A. H. and Galbally, I. E.: Known and Unexplored Organic Constituents in the Earth's Atmosphere,
- 789 Environmental Science & Technology, 41(5), 1514–1521, doi:10.1021/es072476p, 2007.
- 790 Graus, M., Müller, M. and Hansel, A.: High Resolution PTR-TOF: Quantification and Formula Confirmation
- 791 of VOC in Real Time, Journal of the American Society for Mass Spectrometry, 21(6), 1037–1044,
- 792 doi:10.1016/j.jasms.2010.02.006, 2010.
- 793 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue,
- 794 N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M.,
- 795 Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A.,
- 796 Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R. and Wildt, J.: The formation, properties and
- 797 impact of secondary organic aerosol: current and emerging issues, Atmospheric Chemistry and Physics, 9(14),
- 798 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.
- 799 Hansel, A., Scholz, W., Mentler, B., Fischer, L. and Berndt, T.: Detection of RO 2 radicals and other products
- 800 from cyclohexene ozonolysis with NH 4 + and acetate chemical ionization mass spectrometry, Atmospheric
- 801 Environment, 186, 248–255, doi:10.1016/j.atmosenv.2018.04.023, 2018.
- 802 Hyttinen, N., Kupiainen-Määttä, O., Rissanen, M. P., Muuronen, M., Ehn, M. and Kurtén, T.: Modeling the
- 803 Charging of Highly Oxidized Cyclohexene Ozonolysis Products Using Nitrate-Based Chemical Ionization,
- The Journal of Physical Chemistry A, 119(24), 6339–6345, doi:10.1021/acs.jpca.5b01818, 2015.
- 805 Hyttinen, N., Otkjær, R. V., Iyer, S., Kjaergaard, H. G., Rissanen, M. P., Wennberg, P. O. and Kurtén, T.:
- 806 Computational Comparison of Different Reagent Ions in the Chemical Ionization of Oxidized Multifunctional
- 807 Compounds, The Journal of Physical Chemistry A, 122(1), 269–279, doi:10.1021/acs.jpca.7b10015, 2018.
- 808 Intergovernmental Panel on Climate Change, Ed.: Climate Change 2013 The Physical Science Basis:
- 809 Working Group I Contribution to the Fifth Assessment Report of the Intergovernmental Panel on Climate
- 810 Change, Cambridge University Press, Cambridge., 2014.
- 811 Isaacman-VanWertz, G., Massoli, P., O'Brien, R. E., Nowak, J. B., Canagaratna, M. R., Jayne, J. T., Worsnop,
- 812 D. R., Su, L., Knopf, D. A., Misztal, P. K., Arata, C., Goldstein, A. H. and Kroll, J. H.: Using advanced mass
- 813 spectrometry techniques to fully characterize atmospheric organic carbon: current capabilities and remaining
- gaps, Faraday Discussions, 200, 579–598, doi:10.1039/C7FD00021A, 2017.
- 815 Isaacman-VanWertz, G., Massoli, P., O'Brien, R., Lim, C., Franklin, J. P., Moss, J. A., Hunter, J. F., Nowak,

Manuscript under review for journal Atmos. Meas. Tech.

Discussion started: 10 December 2018 © Author(s) 2018. CC BY 4.0 License.





- 816 J. B., Canagaratna, M. R., Misztal, P. K., Arata, C., Roscioli, J. R., Herndon, S. T., Onasch, T. B., Lambe, A.
- 817 T., Jayne, J. T., Su, L., Knopf, D. A., Goldstein, A. H., Worsnop, D. R. and Kroll, J. H.: Chemical evolution
- 818 of atmospheric organic carbon over multiple generations of oxidation, Nature Chemistry, 10(4), 462–468,
- 819 doi:10.1038/s41557-018-0002-2, 2018.
- 820 Iyer, S., Lopez-Hilfiker, F., Lee, B. H., Thornton, J. A. and Kurtén, T.: Modeling the Detection of Organic and
- 821 Inorganic Compounds Using Iodide-Based Chemical Ionization, The Journal of Physical Chemistry A, 120(4),
- 822 576–587, doi:10.1021/acs.jpca.5b09837, 2016.
- 823 Iyer, S., He, X., Hyttinen, N., Kurtén, T. and Rissanen, M. P.: Computational and Experimental Investigation
- of the Detection of HO₂ Radical and the Products of Its Reaction with Cyclohexene Ozonolysis Derived RO
- 825 2 Radicals by an Iodide-Based Chemical Ionization Mass Spectrometer, The Journal of Physical Chemistry A,
- 826 121(36), 6778–6789, doi:10.1021/acs.jpca.7b01588, 2017.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F.,
- 828 Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A.
- 829 L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A.,
- 830 Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R.,
- Cubison, M. J., E., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo,
- Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin,
- 833 R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J.
- 834 R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A.
- 835 M., Kolb, C. E., Baltensperger, U. and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere,
- 836 Science, 326(5959), 1525–1529, doi:10.1126/science.1180353, 2009.
- Jokinen, T., Sipilä, M., Junninen, H., Ehn, M., Lönn, G., Hakala, J., Petäjä, T., Mauldin, R. L., Kulmala, M.
- 838 and Worsnop, D. R.: Atmospheric sulphuric acid and neutral cluster measurements using CI-APi-TOF,
- 839 Atmospheric Chemistry and Physics, 12(9), 4117–4125, doi:10.5194/acp-12-4117-2012, 2012.
- Jokinen, T., Sipilä, M., Richters, S., Kerminen, V.-M., Paasonen, P., Stratmann, F., Worsnop, D., Kulmala,
- 841 M., Ehn, M., Herrmann, H. and Berndt, T.: Rapid Autoxidation Forms Highly Oxidized RO 2 Radicals in the
- 842 Atmosphere, Angewandte Chemie International Edition, 53(52), 14596–14600, doi:10.1002/anie.201408566,
- 843 2014.
- Jokinen, T., Berndt, T., Makkonen, R., Kerminen, V.-M., Junninen, H., Paasonen, P., Stratmann, F., Herrmann,
- 845 H., Guenther, A. B., Worsnop, D. R., Kulmala, M., Ehn, M. and Sipilä, M.: Production of extremely low
- 846 volatile organic compounds from biogenic emissions: Measured yields and atmospheric implications,
- 847 Proceedings of the National Academy of Sciences, 112(23), 7123–7128, doi:10.1073/pnas.1423977112, 2015.
- 848 Jordan, A., Haidacher, S., Hanel, G., Hartungen, E., Märk, L., Seehauser, H., Schottkowsky, R., Sulzer, P. and
- 849 Märk, T. D.: A high resolution and high sensitivity proton-transfer-reaction time-of-flight mass spectrometer
- 850 (PTR-TOF-MS), International Journal of Mass Spectrometry, 286(2), 122–128,
- 851 doi:10.1016/j.ijms.2009.07.005, 2009.
- 852 Junninen, H., Ehn, M., Petäjä, T., Luosujärvi, L., Kotiaho, T., Kostiainen, R., Rohner, U., Gonin, M., Fuhrer,
- 853 K., Kulmala, M. and Worsnop, D. R.: A high-resolution mass spectrometer to measure atmospheric ion
- 854 composition, Atmospheric Measurement Techniques, 3(4), 1039–1053, doi:10.5194/amt-3-1039-2010, 2010.
- 855 Kelly, J. M., Doherty, R. M., O& apos; Connor, F. M. and Mann, G. W.: The impact of biogenic,
- 856 anthropogenic, and biomass burning volatile organic compound emissions on regional and seasonal variations
- in secondary organic aerosol, Atmospheric Chemistry and Physics, 18(10), 7393-7422, doi:10.5194/acp-18-
- 858 7393-2018, 2018.
- 859 Kirkby, J., Duplissy, J., Sengupta, K., Frege, C., Gordon, H., Williamson, C., Heinritzi, M., Simon, M., Yan,
- 860 C., Almeida, J., Tröstl, J., Nieminen, T., Ortega, I. K., Wagner, R., Adamov, A., Amorim, A., Bernhammer,

Manuscript under review for journal Atmos. Meas. Tech.

Discussion started: 10 December 2018 © Author(s) 2018. CC BY 4.0 License.





- 861 A.-K., Bianchi, F., Breitenlechner, M., Brilke, S., Chen, X., Craven, J., Dias, A., Ehrhart, S., Flagan, R. C.,
- 862 Franchin, A., Fuchs, C., Guida, R., Hakala, J., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Kim,
- 863 J., Krapf, M., Kürten, A., Laaksonen, A., Lehtipalo, K., Makhmutov, V., Mathot, S., Molteni, U., Onnela, A.,
- Peräkylä, O., Piel, F., Petäjä, T., Praplan, A. P., Pringle, K., Rap, A., Richards, N. A. D., Riipinen, I., Rissanen,
- 865 M. P., Rondo, L., Sarnela, N., Schobesberger, S., Scott, C. E., Seinfeld, J. H., Sipilä, M., Steiner, G., Stozhkov,
- 866 Y., Stratmann, F., Tomé, A., Virtanen, A., Vogel, A. L., Wagner, A. C., Wagner, P. E., Weingartner, E.,
- Wimmer, D., Winkler, P. M., Ye, P., Zhang, X., Hansel, A., Dommen, J., Donahue, N. M., Worsnop, D. R.,
- 868 Baltensperger, U., Kulmala, M., Carslaw, K. S. and Curtius, J.: Ion-induced nucleation of pure biogenic
- particles, Nature, 533(7604), 521–526, doi:10.1038/nature17953, 2016.
- 870 Krechmer, J., Lopez-Hilfiker, F., Koss, A., Hutterli, M., Stoermer, C., Deming, B., Kimmel, J., Warneke, C.,
- 871 Holzinger, R., Jayne, J. T., Worsnop, D. R., Fuhrer, K., Gonin, M. and de Gouw, J. A.: Evaluation of a New
- 872 Reagent-Ion Source and Focusing Ion-Molecule Reactor for use in Proton-Transfer-Reaction Mass
- 873 Spectrometry, Analytical Chemistry, doi:10.1021/acs.analchem.8b02641, 2018.
- 874 Kürten, A., Rondo, L., Ehrhart, S. and Curtius, J.: Calibration of a Chemical Ionization Mass Spectrometer for
- 875 the Measurement of Gaseous Sulfuric Acid, The Journal of Physical Chemistry A, 116(24), 6375–6386,
- 876 doi:10.1021/jp212123n, 2012.
- 877 Lee, A., Goldstein, A. H., Keywood, M. D., Gao, S., Varutbangkul, V., Bahreini, R., Ng, N. L., Flagan, R. C.
- 878 and Seinfeld, J. H.: Gas-phase products and secondary aerosol yields from the ozonolysis of ten different
- 879 terpenes, Journal of Geophysical Research, 111(D7), doi:10.1029/2005JD006437, 2006.
- 880 Lee, B. H., Lopez-Hilfiker, F. D., Mohr, C., Kurtén, T., Worsnop, D. R. and Thornton, J. A.: An Iodide-Adduct
- 881 High-Resolution Time-of-Flight Chemical-Ionization Mass Spectrometer: Application to Atmospheric
- 882 Inorganic and Organic Compounds, Environmental Science & Technology, 48(11), 6309-6317,
- 883 doi:10.1021/es500362a, 2014.
- Lee, B. H., Mohr, C., Lopez-Hilfiker, F. D., Lutz, A., Hallquist, M., Lee, L., Romer, P., Cohen, R. C., Iyer, S.,
- 885 Kurtén, T., Hu, W., Day, D. A., Campuzano-Jost, P., Jimenez, J. L., Xu, L., Ng, N. L., Guo, H., Weber, R. J.,
- 886 Wild, R. J., Brown, S. S., Koss, A., de Gouw, J., Olson, K., Goldstein, A. H., Seco, R., Kim, S., McAvey, K.,
- 887 Shepson, P. B., Starn, T., Baumann, K., Edgerton, E. S., Liu, J., Shilling, J. E., Miller, D. O., Brune, W.,
- 888 Schobesberger, S., D'Ambro, E. L. and Thornton, J. A.: Highly functionalized organic nitrates in the southeast
- 889 United States: Contribution to secondary organic aerosol and reactive nitrogen budgets, Proceedings of the
- 890 National Academy of Sciences, 113(6), 1516–1521, doi:10.1073/pnas.1508108113, 2016.
- 891 Li, X., Chee, S., Hao, J., Abbatt, J. P. D., Jiang, J. and Smith, J. N.: Relative Humidity Effect on the Formation
- 892 of Highly Oxidized Molecules and New Particles during Monoterpene Oxidation, Atmospheric Chemistry and
- 893 Physics Discussions, 1–26, doi:10.5194/acp-2018-898, 2018.
- 894 Lindinger, W., Hansel, A. and Jordan, A.: On-line monitoring of volatile organic compounds at pptv levels by
- 895 means of proton-transfer-reaction mass spectrometry (PTR-MS) medical applications, food control and
- environmental research, International Journal of Mass Spectrometry and Ion Processes, 173(3), 191–241,
- 897 doi:10.1016/S0168-1176(97)00281-4, 1998.
- 898 Lopez-Hilfiker, F. D., Iyer, S., Mohr, C., Lee, B. H., D' Ambro, E. L., Kurtén, T. and Thornton, J. A.:
- 899 Constraining the sensitivity of iodide adduct chemical ionization mass spectrometry to multifunctional organic
- 900 molecules using the collision limit and thermodynamic stability of iodide ion adducts, Atmospheric
- 901 Measurement Techniques, 9(4), 1505–1512, doi:10.5194/amt-9-1505-2016, 2016.
- 902 Mohr, C., Lopez-Hilfiker, F. D., Yli-Juuti, T., Heitto, A., Lutz, A., Hallquist, M., D'Ambro, E. L., Rissanen,
- 903 M. P., Hao, L., Schobesberger, S., Kulmala, M., Mauldin, R. L., Makkonen, U., Sipilä, M., Petäjä, T. and
- 904 Thornton, J. A.: Ambient observations of dimers from terpene oxidation in the gas phase: Implications for new
- 905 particle formation and growth: Ambient Observations of Gas-Phase Dimers, Geophysical Research Letters,
- 906 44(6), 2958–2966, doi:10.1002/2017GL072718, 2017.

Manuscript under review for journal Atmos. Meas. Tech.

Discussion started: 10 December 2018 © Author(s) 2018. CC BY 4.0 License.





- 907 Pagonis, D., Krechmer, J. E., de Gouw, J., Jimenez, J. L. and Ziemann, P. J.: Effects of gas-wall partitioning
- 908 in Teflon tubing and instrumentation on time-resolved measurements of gas-phase organic compounds,
- 909 Atmospheric Measurement Techniques, 10(12), 4687–4696, doi:10.5194/amt-10-4687-2017, 2017.
- 910 Riva, M., Budisulistiorini, S. H., Zhang, Z., Gold, A., Thornton, J. A., Turpin, B. J. and Surratt, J. D.:
- 911 Multiphase reactivity of gaseous hydroperoxide oligomers produced from isoprene ozonolysis in the presence
- 912 of acidified aerosols, Atmospheric Environment, 152, 314–322, doi:10.1016/j.atmosenv.2016.12.040, 2017.
- 913 Schallhart, S., Rantala, P., Nemitz, E., Taipale, D., Tillmann, R., Mentel, T. F., Loubet, B., Gerosa, G., Finco,
- 914 A., Rinne, J. and Ruuskanen, T. M.: Characterization of total ecosystem-scale biogenic VOC exchange at a
- 915 Mediterranean oak-hornbeam forest, Atmospheric Chemistry and Physics, 16(11), 7171-7194,
- 916 doi:10.5194/acp-16-7171-2016, 2016.
- 917 Schobesberger, S., Junninen, H., Bianchi, F., Lonn, G., Ehn, M., Lehtipalo, K., Dommen, J., Ehrhart, S.,
- 918 Ortega, I. K., Franchin, A., Nieminen, T., Riccobono, F., Hutterli, M., Duplissy, J., Almeida, J., Amorim, A.,
- 919 Breitenlechner, M., Downard, A. J., Dunne, E. M., Flagan, R. C., Kajos, M., Keskinen, H., Kirkby, J., Kupc,
- 920 A., Kurten, A., Kurten, T., Laaksonen, A., Mathot, S., Onnela, A., Praplan, A. P., Rondo, L., Santos, F. D.,
- 921 Schallhart, S., Schnitzhofer, R., Sipila, M., Tome, A., Tsagkogeorgas, G., Vehkamaki, H., Wimmer, D.,
- 922 Baltensperger, U., Carslaw, K. S., Curtius, J., Hansel, A., Petaja, T., Kulmala, M., Donahue, N. M. and
- 923 Worsnop, D. R.: Molecular understanding of atmospheric particle formation from sulfuric acid and large
- 924 oxidized organic molecules, Proceedings of the National Academy of Sciences, 110(43), 17223–17228,
- 925 doi:10.1073/pnas.1306973110, 2013.
- 926 Stark, H., Yatavelli, R. L. N., Thompson, S. L., Kimmel, J. R., Cubison, M. J., Chhabra, P. S., Canagaratna,
- 927 M. R., Jayne, J. T., Worsnop, D. R. and Jimenez, J. L.: Methods to extract molecular and bulk chemical
- 928 information from series of complex mass spectra with limited mass resolution, International Journal of Mass
- 929 Spectrometry, 389, 26–38, doi:10.1016/j.ijms.2015.08.011, 2015.
- 930 Stark, H., Yatavelli, R. L. N., Thompson, S. L., Kang, H., Krechmer, J. E., Kimmel, J. R., Palm, B. B., Hu,
- 931 W., Hayes, P. L., Day, D. A., Campuzano-Jost, P., Canagaratna, M. R., Jayne, J. T., Worsnop, D. R. and
- 932 Jimenez, J. L.: Impact of Thermal Decomposition on Thermal Desorption Instruments: Advantage of
- 933 Thermogram Analysis for Quantifying Volatility Distributions of Organic Species, Environmental Science &
- 934 Technology, 51(15), 8491–8500, doi:10.1021/acs.est.7b00160, 2017.
- 935 Tröstl, J., Chuang, W. K., Gordon, H., Heinritzi, M., Yan, C., Molteni, U., Ahlm, L., Frege, C., Bianchi, F.,
- 936 Wagner, R., Simon, M., Lehtipalo, K., Williamson, C., Craven, J. S., Duplissy, J., Adamov, A., Almeida, J.,
- 937 Bernhammer, A.-K., Breitenlechner, M., Brilke, S., Dias, A., Ehrhart, S., Flagan, R. C., Franchin, A., Fuchs,
- 938 C., Guida, R., Gysel, M., Hansel, A., Hoyle, C. R., Jokinen, T., Junninen, H., Kangasluoma, J., Keskinen, H.,
- 939 Kim, J., Krapf, M., Kürten, A., Laaksonen, A., Lawler, M., Leiminger, M., Mathot, S., Möhler, O., Nieminen,
- 940 T., Onnela, A., Petäjä, T., Piel, F. M., Miettinen, P., Rissanen, M. P., Rondo, L., Sarnela, N., Schobesberger,
- 941 S., Sengupta, K., Sipilä, M., Smith, J. N., Steiner, G., Tomè, A., Virtanen, A., Wagner, A. C., Weingartner, E.,
- 942 Wimmer, D., Winkler, P. M., Ye, P., Carslaw, K. S., Curtius, J., Dommen, J., Kirkby, J., Kulmala, M.,
- 943 Riipinen, I., Worsnop, D. R., Donahue, N. M. and Baltensperger, U.: The role of low-volatility organic
- 944 compounds in initial particle growth in the atmosphere, Nature, 533(7604), 527-531,
- 945 doi:10.1038/nature18271, 2016.
- 946 Twomey, S.: The Influence of Pollution on the Shortwave Albedo of Clouds, Journal of the Atmospheric
- 947 Sciences, 34(7), 1149–1152, doi:10.1175/1520-0469(1977)034<1149:TIOPOT>2.0.CO;2, 1977.
- 948 Wang, S., Riva, M., Yan, C., Ehn, M. and Wang, L.: Primary Formation of Highly Oxidized Multifunctional
- 949 Products in the OH-Initiated Oxidation of Isoprene. A Combined Theoretical and Experimental Study,
- 950 Environmental Science & Technology, doi:10.1021/acs.est.8b02783, 2018.
- 951 Yuan, B., Koss, A. R., Warneke, C., Coggon, M., Sekimoto, K. and de Gouw, J. A.: Proton-Transfer-Reaction
- 952 Mass Spectrometry: Applications in Atmospheric Sciences, Chemical Reviews, 117(21), 13187–13229,

Manuscript under review for journal Atmos. Meas. Tech.

Discussion started: 10 December 2018 © Author(s) 2018. CC BY 4.0 License.





- 953 doi:10.1021/acs.chemrev.7b00325, 2017.
- 954 Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol formation,
- 955 Chemical Society Reviews, 41(19), 6582, doi:10.1039/c2cs35122f, 2012.