

Response to Anonymous Referee #1 comments on “FRESCO-B: A fast cloud retrieval algorithm using oxygen B-band measurements from GOME-2” by Marine Desmons et al.

Authors: M. Desmons, P. Wang, P. Stammes, and L. G. Tilstra

The authors are grateful to the referee for the constructive evaluation and useful comments on the paper. In the following, a point-by-point reply is given, with the Referee's comments in italics.

This paper presents the new FRESCO-B algorithm, which retrieves cloud properties using measurements at the oxygen B-band regions. This new algorithm adopts the same infrastructure of the robust and well-tested FRESCO model. The paper gives the theoretical basis, radiative transfer simulations and performance evaluation. The work is helpful in further understand the cloud retrievals with oxygen A- and B-bands. The paper is well written and the topic is suitable for publication in AMT. I recommend publication after some minor revisions. Some comments are as follows:

1) Trying to understand the simulation results shown in Figure 6. Using Fig 6a as an example (it's easier to analyze since the surface albedo is the same for both FRESCO- B and FRESCO), why do you think the pressure difference is negative for the thin-and- low cloud case? There are three factors that contribute to the pressure difference: (a) the photon path differences inside clouds; (b) the photon path differences below clouds; and (c) the differences in the effective cloud fraction. The only one that can cause negative pressure difference is (c). Since FRESCO-B has a slightly larger effective cloud fraction, the total photon path for the clear part of the pixel should be smaller than FRESCO. I'm thinking if you force the effective cloud fraction to be the same, the negative cloud pressure difference will probably disappear for Fig6a.

This is a very interesting observation by the reviewer: for the case of a thin low cloud the FRESCO-B pressure is a bit lower than the FRESCO pressure. This holds for both the ocean and vegetation cases, except for the most oblique geometry for vegetation. In addition to the three factors mentioned by the reviewer there is also the factor of the light path above the cloud, which could be influenced by reflection by the cloud. Since the B-band is weaker than the A-band, multiple scattering between the cloud particles and the molecular Rayleigh scatterers above, inside, and below the cloud is stronger in the B-band than in the A-band. At 685 nm there is 50 % more Rayleigh scattering than at 760 nm. Most Rayleigh scattering is above 5 km, so by scattering above the cloud the B-band pressure would be lower than the A-band pressure.

The effective cloud fraction appears to be slightly larger in the B-band than in the A-band, which is well observed by the reviewer. This effect would counteract the above effect, since a larger cloud fraction means that more photons come from the cloud level instead of the clear sky. We followed the suggestion of the reviewer, and forced the effective cloud fraction in the B-band to be equal to that in the A-band. We obtain the same negative pressure difference. The values are resumed in the following table:

$c_{\text{eff}}=0.4679$	$P_{\text{FRESCO-B}}=685.285 \text{ hPa}$
	$P_{\text{FRESCO}}=689.318 \text{ hPa}$
$c_{\text{eff}}=0.4558$	$P_{\text{FRESCO-B}}=685.067 \text{ hPa}$
	$P_{\text{FRESCO}}=689.139 \text{ hPa}$

We changed the manuscript accordingly, section 4.3, p19:

...For the case of a thin low cloud, the FRESCO-B pressure is a bit lower than the FRESCO pressure. This holds for both the ocean and vegetation cases, except for the most oblique geometry for vegetation. This feature can be due to the Rayleigh scattering. Since the B-band is weaker than the A-band, multiple scattering between the cloud particles and the molecular Rayleigh scatterers above, inside, and below the cloud is stronger in the B-band than in the A-band. At 685 nm there is 50 % more Rayleigh scattering than at 760 nm. Most Rayleigh scattering is above 5 km, so the pressure retrieved in the B-band is lowered by the scattering happening above the cloud, leading to a smaller (negative) difference of pressures....

Specific comments:

P5 Fig 3: Please check the vertical axis of the figures. It doesn't look right.

Thanks for identifying this mistake. In Fig. 3 we plotted the reflectance at top-of-atmosphere, whereas we intended to plot the transmittance at the surface, to clearly show the absorption bands of oxygen. We corrected this. The atmospheric set-up and geometry is now given in the caption.

Here is the new figure:

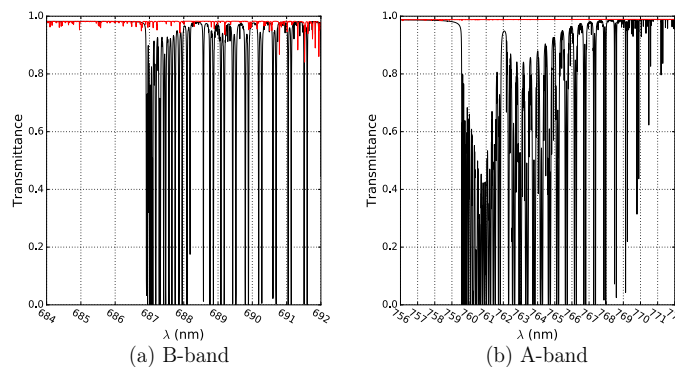


Figure 3: Line-by-line transmittances in the oxygen A (b) and B (a) bands (black lines). The transmittances of the overlapping water vapor lines are represented in red. The calculations are computed using the HITRAN 2016 database (Gordon, 2017) and for a solar zenith angle of 0° .

P18 Fig 9: Missing the (b) panel caption

Added. Thanks for noticing this oversight.

Response to the anonymous referee #2 comments on “FRESCO-B: A fast cloud retrieval algorithm using oxygen B-band measurements from GOME-2” by Marine Desmons et al.

Authors: M. Desmons, P. Wang, P. Stammes, and L. G. Tilstra

The authors are grateful to the referee for the constructive evaluation and useful comments. In the following, a point-by-point reply is given, with the Referee's comments in italics.

This paper presents a new FRESCO-B algorithm that was built up on the FRESCO algorithm but uses spectral measurements in the O₂ B-band, whereas FRESCO algorithm uses the O₂ A-band measurements. Both FRESCO-B and FRESCO algorithms retrieve cloud effective fraction and cloud top pressure from three 1-nm wide bands around the corresponding O₂ absorption bands. The development of FRESCO-B algorithm is motivated for its valuable application over vegetation surface that has much lower surface albedo in the O₂ B-band. This work demonstrated the promising retrievals using the FRESCO-B algorithm with both the synthetic data and the GOME- 2 measured data. The study is well designed and matches the scope of AMT. The manuscript is well written. I only have a few minor comments as below.

(1) For the algorithm, why select only three 1-nm wide windows instead of performing spectral fitting at all wavelengths of GOME-2 measurements around O₂ A- and B- band? Is it supposed to have more information if more wavelengths are used?

The use of a selected number of wavelengths in the O₂ A-band is one of the key features of the FRESCO method as devised by Koelemeijer et al. (2001). The basic idea is that the A-band (and B-band) contains much dependent information, since it consists of hundreds of lines. The reflectances in the selected three wavelength windows contain nearly all independent information that is available in the O₂ A band for instruments with the spectral resolution of GOME-2 (Kollewe et al., 1992, Fischer et al., 1992). Using the entire O₂ A-band (or B-band) would unnecessarily slow down the retrieval.

(2) Can authors include the definition for “effective cloud fraction” and explain why cloud effective fraction can exceed 1 (Figure 5)?

The definition of effective cloud fraction is now given in Sect. 4.2. p7:

...In Figure 5, we can see that the retrieved effective cloud fraction is sometimes higher than 1. This is due to the principle of the FRESCO algorithm. Indeed, the effective cloud fraction is the part of the pixel that the Lambertian cloud has to occupy to match the observed reflectance while the geometric cloud fraction is the part of the pixel that is covered by the true clouds. The effective cloud fraction is strongly coupled to the choice of the cloud albedo A_c . The choice of $A_c=0.8$ has been made in the FRESCO algorithm (Koelemeijer et al., 2001, Stammes et al., 2008) for different reasons (correction of trace gases for clouds, consistency of the Lambertian

model, ability to approach the measured reflectivities by simulations), and can lead to effective cloud fractions somewhat higher than 1....

(3) Figure 5 and 6. Need to label the left and right panels with (a) and (b), as these are indicated in the figure caption.

Done, we apologize for this omission

(4) Page 7 line 15: "Figure 4b" -> "Figure 5b".

Corrected.

(5) Table 2&3: As shown in the simulated retrievals in section 4, cloud effective fraction is larger from O2-B retrievals. Why the differences are negative for Land and Vegetation cases of GOME-2 retrievals? Ok, I found this is discussed on page 10.

Ok

(6) Figure 7a is not discussed in the text.

This figure is described in paragraph 5.1.1, p10

(7) It seems to me Figure 7b shows substantial land areas with negative cloud pressure difference. However, global average of this difference over Land is positive in Table 4, which may not consistent with Figure 7b. Please double check.

We have checked and both the table and the figure are correct. We understand your concern, but there are also many land areas with a positive difference, which balance the areas with a negative difference. We have added the histograms of the cloud effective fraction differences and of the cloud pressures differences over ocean and land in the paper in section 5.1 p13, and changed the text accordingly.

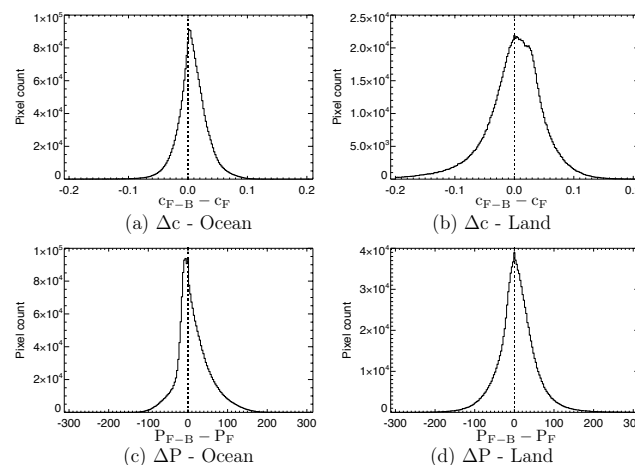
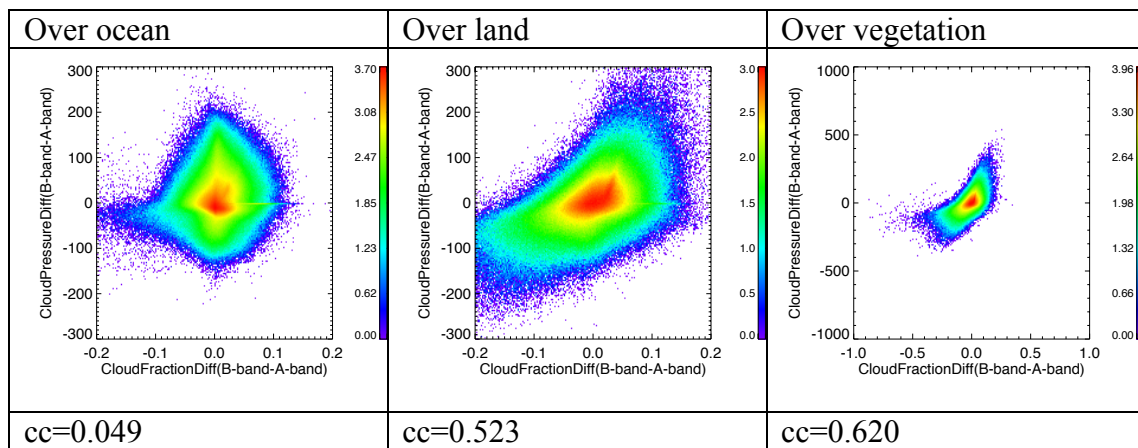


Figure 8: Histograms of the differences of effective cloud fractions (panels a and b) and of cloud pressures (panels c and d) between FRESCO-B and FRESCO for July 2014, over ocean (panels a and c) and over land (panels b and d).}

.... Over land, the cloud pressure difference is also positive on average with a mean value of 6.31 ± 49.1 hPa. However, as we can see on Figure 8d, the range of the values is larger than over ocean with a lot of pixels having a negative difference of pressure. This is again due to the high variability of the surface albedos for this type of surface and is coherent with the observations we have made on the simulations. Over land, the coefficient of correlation between the difference of pressures and the difference of cloud effective fractions is 0.523....

(8) In Figure 7, it appears some correlation may exist between the differences in effective cloud fraction and the differences in cloud top pressure. For instance, areas over land with negative cloud pressure difference tend to have negative cloud fraction difference. It seems a low bias in cloud fraction may lead to low bias in cloud pressure?

We have checked this point and plot the correlation between the difference of pressures and the difference of effective cloud fraction for different surfaces:



Consequently, the correlation between the difference of pressures and the difference of effective cloud fractions is inexistent for ocean pixels and moderate for land and vegetation cases. However, we think that the O₂ A- and B-bands react differently to cloud fraction and cloud albedo, which both influence effective cloud fraction, and to cloud vertical structure. These three cloud parameters represent different cloud properties that are not generally correlated. Thus, we think that the hints towards such a correlation are not strong enough.

We have added the value of those correlation coefficients in the paper, section 5.1.2:

.... Over ocean, the coefficient of correlation between the difference of pressures and the difference of cloud effective fractions is 0.0949....(p14)

.... Over land, the coefficient of correlation between the difference of pressures and the difference of cloud effective fractions is 0.523.... (p15)

(9) Figure 9: Need to add label “(b) Multi-layer clouds”

Done, thanks for noticing this oversight.

FRESCO-B: A fast cloud retrieval algorithm using oxygen B-band measurements from GOME-2

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Abstract. The FRESCO (Fast Retrieval Scheme for Clouds from the Oxygen A-band) algorithm is a simple, fast and robust algorithm used to retrieve cloud information in operational satellite data processing. It has been applied to GOME-1, SCIAMACHY, GOME-2 and more recently to TROPOMI. FRESCO retrieves effective cloud fraction and cloud pressure from measurements in the oxygen A-band around 761 nm. In this paper, we propose a new version of the algorithm, called
5 FRESCO-B, which is based on measurements in the oxygen B-band around 687 nm. Such a method is interesting for vegetated surfaces where the surface albedo is much lower in the B-band than in the A-band, which limits the ground contribution to the top-of-atmosphere reflectances. In this study we first perform retrieval simulations. These show that the retrieved cloud pressures from FRESCO-B and FRESCO differ only between -10 hPa and +10 hPa, except for high thin clouds over vegeta-
10 tion where the difference is larger, about +15 to +30 hPa, with FRESCO-B yielding higher pressures. Next, inter-comparison between FRESCO-B and FRESCO retrievals over one month of GOME-2B data reveals that the effective cloud fractions re-
trieved in the O₂ A and B bands are very similar (mean difference of 0.003) while the cloud pressures show a mean difference of 11.5 hPa, with FRESCO-B retrieving higher pressures than FRESCO. This agrees with the simulations and is partly due to deeper photons penetrations of O₂ B-band in clouds as compared to the O₂ A-band photons, and partly due to the surface albedo bias in FRESCO. Finally, validation with ground-based measurements shows that the FRESCO-B cloud pressure repre-
15 sents an altitude within the cloud boundaries for clouds that are not too far from the Lambertian reflector model, which occurs in about 50% of the cases.

Copyright statement. TEXT

1 Introduction

The Global Ozone Monitoring Experiment-2 (GOME-2) is a spectrometer flying on the Metop series of satellites: on Metop-
20 A since 2006, on Metop-B since 2012, and on Metop-C, launched in November 2018. The GOME-2 instruments sense the Earth back-scattered radiance and solar irradiance in the ultraviolet and visible part of the spectrum (240-790 nm) with a spectral resolution between 0.26 nm and 0.51 nm. The primary goal of GOME-2 measurements is the study of ozone as well as atmospheric trace gases and pollutants (e.g. nitrogen dioxide, sulfur dioxide, water vapor, bromine oxide). The trace

gas retrieval algorithms rely on information on cloud properties for each ground pixel. Indeed, clouds strongly affect trace gas retrievals from satellite measurements because of their shielding effect, their albedo effect and because of the in-cloud absorption effect (Stammes et al., 2008). Therefore, in order to ensure good quality trace gas retrievals, it is very important to retrieve cloud properties, and in particular, the cloud top height, cloud geometrical thickness, cloud fraction, cloud optical
5 thickness and the number of cloud layers (Boersma et al., 2004).

From a larger perspective, clouds are a key component of the Earth's climate system through their role in the Earth hydrological cycle and radiation balance. Global observation and description of clouds are necessary to understand and properly depict their overall and multiple effects. This is particularly true in the context of the climate change we are experiencing. Especially, the question whether the coverage of different cloud types will change, or the partition of low-level versus high-level clouds -
10 which have different and sometimes opposite radiative effects - will change is a crucial one. This is recognized as one of the major challenges in climate change predictions (Bony and Dufresne, 2005; Andrews et al., 2012; Vial et al., 2013).

The usage of the oxygen absorption for the retrieval of cloud pressure has already been studied for several decades. Indeed, O_2 is well mixed in the atmosphere and the degree of O_2 absorption can be related to a certain atmospheric path length. Above a bright surface, as a cloud acts in first approximation, O_2 absorption that affects solar radiation back-scattered towards a
15 space-borne sensor is mainly related to the scene height (the cloud height in our case). Such methods using reflected sunlight in oxygen absorption bands depend very weakly on the pressure/temperature vertical profiles. They do not suffer from a lack of sensitivity in the case of low clouds, and are not sensitive to temperature inversions, like retrievals with infrared measurements. The use of the oxygen A-band for remote sensing of cloud properties has been extensively studied (Wu, 1985; Fischer and Grassl, 1991; Kuze and Chance, 1994), and used in airborne campaigns (Fischer et al., 1991) and satellite missions
20 (Vanbauce et al., 1998; Koelemeijer et al., 2001; Fournier et al., 2006; Lindstrot et al., 2006; Preusker et al., 2007; Lelli et al., 2012), demonstrating the capabilities to retrieve an apparent cloud pressure using different sensors with narrow spectral bands centered on the oxygen A-band region. However, the use of the oxygen B-band for such retrievals remains quite limited, usually in association with measurements in the A-band (Kuze and Chance, 1994; Yang et al., 2013), or applied to the retrieval of aerosols or vegetation properties (Marshak and Knyazikhin, 2017; Xu et al., 2017). In this paper, we propose a new version
25 of the FRESCO algorithm, called FRESCO-B, which is based on measurements in the B-band. Such a method is interesting for vegetated surfaces. Indeed for this type of surface, the surface reflectance is much lower in the B-band than in the A-band (Tilstra et al., 2017), limiting the ground contribution to the top-of-atmosphere (TOA) measured reflectances.

This paper is ~~organised~~organized as follows. In Sect. 2, we present the FRESCO algorithm and our motivations to apply it in the oxygen B-band. In Sect. 3, we describe the oxygen A and B-bands, as well as the FRESCO-B retrieval method. In Sect.
30 4, we perform simulations of cloud retrievals in the oxygen A and B bands. In Sect. 5, we apply FRESCO-B to GOME-2 data, and validate the results by performing comparisons with FRESCO and with ground-based measurements. We conclude this study in Sect. 6.

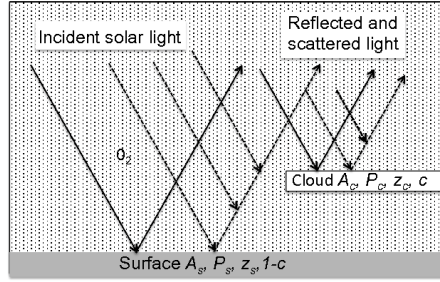


Figure 1. Atmospheric radiation model used in the FRESCO retrieval algorithm. Three light paths are considered: (1) from sun to surface to satellite, (2) from sun to cloud to satellite (1-2: solid lines), (3) from sun to atmosphere to satellite according to single Rayleigh scattering (dashed lines). The pixel is partly cloudy, with c the effective cloud fraction, and partly clear. A_s , P_s and z_s stand respectively for the surface albedo, pressure and altitude, while A_c , P_c and z_c indicate respectively the cloud albedo, pressure and altitude.

2 The FRESCO algorithm

In the FRESCO algorithm (Koelemeijer et al., 2001; Wang et al., 2008), information on cloud pressure and effective cloud fraction is derived from the reflectances in three 1 nm wide windows, situated in and around the O_2 A-band. The algorithm fits a simulated reflectance spectrum to the measured reflectance spectrum in the three windows, namely 758-759 nm, 760-761 nm and 765-766 nm. The atmospheric model used, shown in Figure 1, is based on the independent pixel approximation: the top-of-atmosphere simulated reflectances (R_{sim}) are computed as the weighted sum of the reflectances of the cloud-free and cloudy parts of the pixel, using the cloud fraction for the weight. The atmosphere above the ground surface or cloud is treated as an absorbing (due to oxygen) and purely Rayleigh scattering medium. The simulated reflectances can be written as (Wang et al., 2008):

$$10 \quad R_{sim} = (1 - c)T_s A_s + (1 - c)R_s + cT_c A_c + cR_c \quad (1)$$

where c is the effective cloud fraction while A_c and A_s stand for the cloud and surface albedo. R_c , T_c and R_s , T_s are the single Rayleigh scattering reflectances and transmittances of the cloudy and cloud-free part of the pixel, respectively. The transmission and Rayleigh scattering reflectances are pre-calculated as a function of the solar zenith angle (SZA), viewing zenith angle (VZA), azimuth difference, wavelength and pressure level (P_c , P_s). The transmission and reflectance spectra are calculated using a line-by-line method using the line parameters from HITRAN 2016 (Gordon et al., 2017) and then **convolved** convoluted using the instrument spectral response function at the measurement wavelength grid. In this model, reflection occurs only at the ground surface or cloud top, and both the ground surface and the cloud are assumed to be Lambertian reflectors. Consequently, the area below the cloud does not contribute in the radiative transfer calculation. The surface albedo is taken from a climatology (Tilstra et al., 2017) while the surface pressure is calculated from surface elevation, using the Mid-Latitude Summer atmospheric profile. The cloud albedo is assumed to be 0.8 or set to the reflectance at 758 nm if the reflectance is

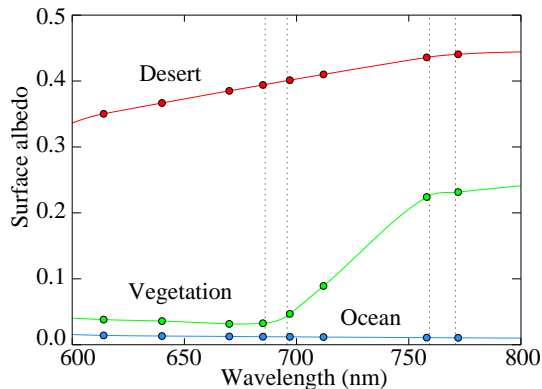


Figure 2. Mean surface albedo over one month (August) and for different surface types as a function of wavelength. “Desert” stands for the Sahara desert, “Ocean” for the Atlantic ocean, and “Vegetation” for the Amazonian tropical rain forest. The vertical dashed lines indicate the location of the oxygen A band (around 761 nm) and B band (around 687 nm). The values come from the SCIAMACHY surface albedo database described in Tilstra et al. (2017).

larger than 0.8. The retrieval method is based on minimizing the difference between the measured and simulated spectrum, using a Levenberg-Marquardt nonlinear least-squares method.

The primary aim of FRESKO is to serve the cloud correction in the trace gas retrievals, but cloud modellers are also interested in the FRESKO data, because the retrieval method uses oxygen and not temperature, and is therefore also sensitive to low, warm clouds. The FRESKO cloud algorithm is simple, fast and robust (Wang and Stammes, 2014) and therefore suitable for operational processing; it has been applied to GOME-1, SCIAMACHY, GOME-2 and more recently, to TROPOMI. Consequently, it is worth continuing to improve the method and to develop new applications.

A motivation to use the B-band to estimate the cloud height is the surface albedo at the B-band wavelengths. Figure 2 shows the mean surface albedo values from the SCIAMACHY surface Lambertian-Equivalent Reflectivity (LER) database (Tilstra et al., 2017). The values are taken for the month of August, for different regions (Atlantic ocean, Sahara desert and Amazonian Tropical rain forests), and is of the “MODE-LER” type. We can see that for vegetation, the surface albedo is significantly lower within the B-band than within the A-band. This means that the contribution of the ground in the top-of-atmosphere reflectances is lower in the oxygen B-band than in the A-band which may lead to more accurate cloud properties retrieval in the B-band over vegetation.

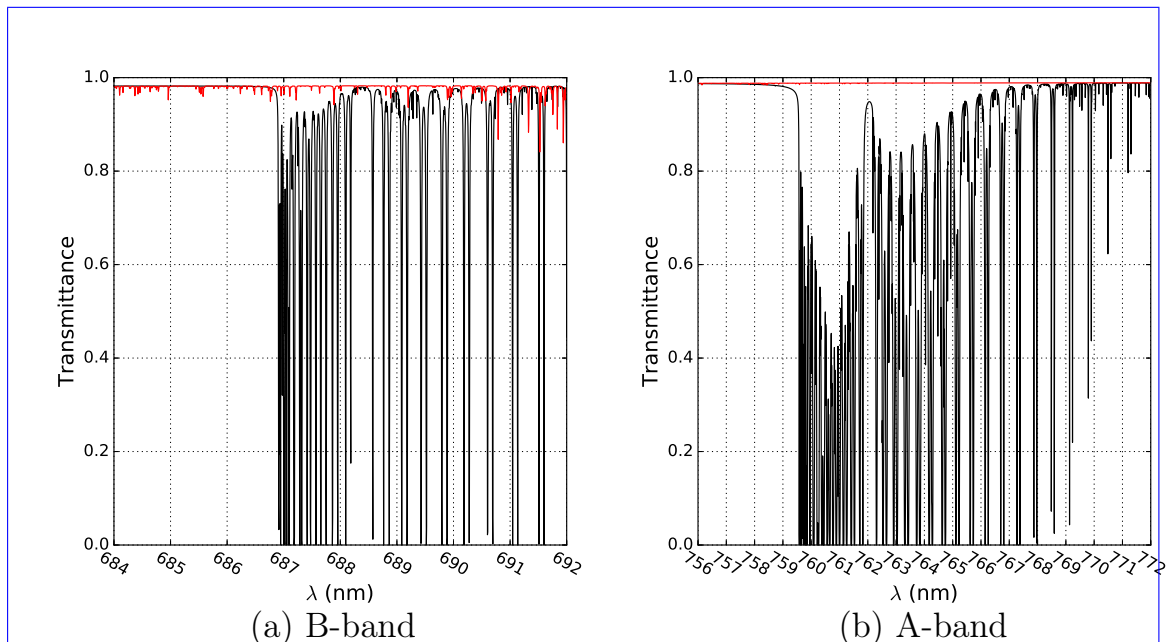


Figure 3. Line-by-line transmittances in the oxygen A (b) and B (a) bands (black lines). **On panel (a), the transmittance** The transmittances of the overlapping water vapor lines **is-are** represented in red. The transmittances-calculations are computed using the HITRAN 2016 database (Gordon et al., 2017) and for a solar zenith angle of 0°.

3 Radiative transfer in oxygen A and B-bands and FRESCO-B retrieval method

3.1 Oxygen A and B bands

The oxygen molecule has two strong absorption bands: the A-band around 761 nm and the B-band around 687 nm. Figure 3 shows the high-resolution transmittances, calculated using the 2016 HITRAN database (Gordon et al., 2017). We can clearly see that the B-band is less deep than the A-band, which means that less light is absorbed by oxygen in the B-band than in the A-band. With the hypothesis of a cloud acting like a high-albedo Lambertian reflector, the cloud height retrieved in the B-band is usually lower than the cloud height retrieved in the A-band (Yang et al., 2013). This is visible in Figure 4 which shows GOME-2B measurements in the oxygen A and B bands for two clouds at different altitudes over ocean. For the low cloud (black lines), FRESCO-B and FRESCO retrieves similar cloud pressures of 975 hPa and 980 hPa respectively. However, for the high cloud, the retrieved pressures are quite different as FRESCO-B cloud pressure is 461 hPa while FRESCO cloud pressure is 412 hPa. This large difference will be discussed later on but it is clear that retrieving cloud height from measurements in the oxygen B-band is very valuable. Indeed, it brings new information regarding the vertical location of the cloud, which can then be used together with the measurements in the oxygen A-band in order to have more information about the cloud layer.

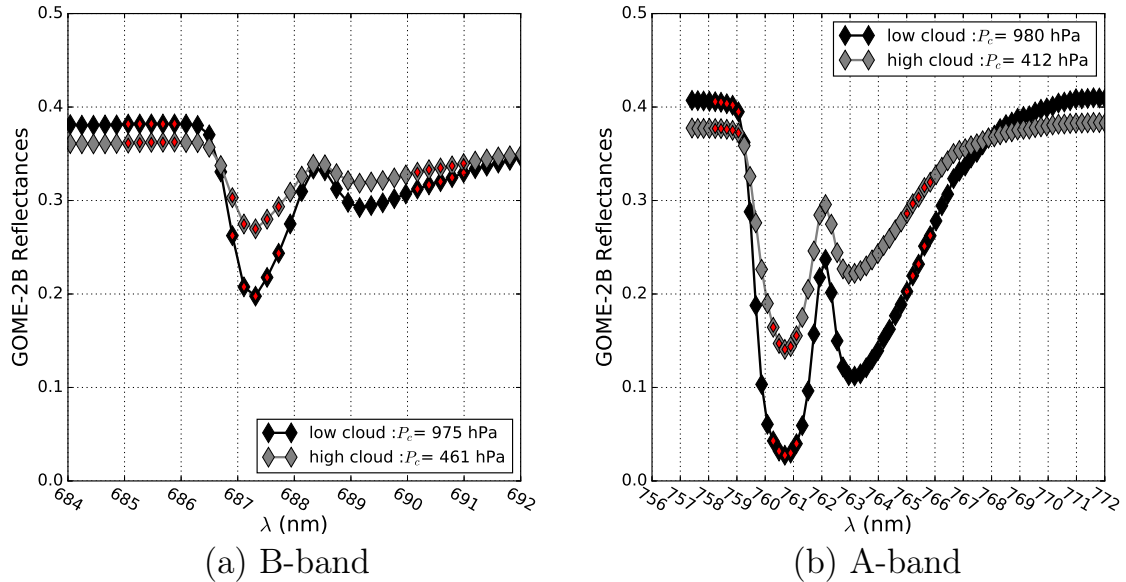


Figure 4. Example of GOME-2B spectral reflectance measurements in the B-band (a) and A-band (b) for low clouds (black line) and high clouds (grey line) for a pixel over ocean. The red colored symbols indicate the measurements that are used in the FRESCO-B and FRESCO algorithms, respectively. The cloud pressures retrieved with FRESCO-B and FRESCO are indicated.

3.2 FRESCO-B retrieval method

Similarly to FRESCO (Koelemeijer et al., 2001; Wang et al., 2008), the FRESCO-B retrieval method is based on minimizing the difference between a measured and a simulated spectrum using a Levenberg-Marquardt nonlinear least-squares method. FRESCO-B retrieves the effective cloud fraction as well as the cloud pressure from the top-of-atmosphere reflectances at three 1 nm wide windows, namely 685–686 nm, 686.8–687.8 nm, and 690–691 nm. The wavelengths are chosen in order to maximize the difference of absorption between the windows. Also, Figure 3 shows that the contamination by water vapor is small so we neglect it. Each of the three windows contains 5 reflectance measurements by GOME-2B as can be seen on Figure 4a. While the reflectances measured in the continuum (between 685 nm and 686 nm) are not impacted by the altitude of the cloud but only by its albedo, the amount of absorption in the two other windows varies according to the cloud altitude.

10 4 Simulations

4.1 Methodology

In order to understand FRESCO-B and its differences with FRESCO, we have simulated spectra for four different cloud cases. O_2 A and B-band reflectance spectra have been simulated with the DAK (Doubling-Adding KNMI) model (De Haan et al.,

1987; Stammes et al., 1989; Stammes, 2001), which is a line-by-line radiative transfer model. The spectral simulations have been made for a Mid-Latitude Summer atmosphere consisting of 33 plane-parallel homogeneous layers with multiple scattering and oxygen absorption. The O₂ absorption cross-sections were calculated using HITRAN 2016 line parameters (Gordon et al., 2017). In this atmosphere, homogeneous scattering cloud layers are inserted, with varying optical thickness. The cloud particle scattering phase function is a Henyey-Greenstein function with an asymmetry parameter of 0.85 and a single scattering albedo of 1. The cloud scenes are simulated with single-layer clouds which fully cover the pixels with a top altitude of either 5 km or 10 km, a geometrical thickness of 3 km, and for an optical thickness of 9 or 42. The four different cases, similar to the ones used in Sneep et al. (2008), are described in Table 1. The spectra were calculated from 684 nm to 692 nm for the B-band and from 756 nm to 772 nm for the A-band, both at 0.005 nm resolution. For each band, the simulations have been done for ocean and vegetation, with surface albedo values taken from the database described in Sect. 2. Then, we have ~~convolved~~ convoluted the obtained spectra with the GOME-2B slit functions and applied the FRESKO-B and FRESKO algorithms. Results are shown for 4 different viewing and solar geometries.

4.2 Cloud fraction

In Figure 5, we can see that the retrieved effective cloud fraction is sometimes higher than 1. This is due to the principle of the FRESKO algorithm. Indeed, the effective cloud fraction is the part of the pixel that the Lambertian cloud has to occupy to match the observed reflectance while the geometric cloud fraction is the part of the pixel that is covered by the true clouds. The effective cloud fraction is strongly coupled to the choice of the cloud albedo A_c . The choice of $A_c = 0.8$ has been made in the FRESKO algorithm (Koelemeijer et al., 2001, Stammes et al., 2008) for different reasons (correction of trace gases for clouds, consistency of the Lambertian model, ability to approach the measured reflectivities by simulations), and can lead to effective cloud fractions somewhat higher than 1.

In Figure 5a, we can see that for the cloud cases over ocean, the effective cloud fraction retrieved with FRESKO-B and FRESKO are very similar, as the difference ranges from 0 to 0.02. Figure 45b shows that over vegetation, the difference between the two effective cloud fractions is higher as it is comprised between 0.01 and 0.04. The effective cloud fraction retrieved with FRESKO-B is always higher than the one retrieved with FRESKO. This behaviour is as expected. Indeed, for wavelengths in the continuum (which are used in the effective cloud fraction retrievals), we may set $T = 1$ in Equation 1. Using $R_{sim}(\lambda) \approx R_{meas}(\lambda)$ leads to:

$$c = \frac{R_{meas} - A_s - R_s}{A_c - A_s + R_c - R_s} \quad (2)$$

Differentiation to A_s gives the change in the retrieved effective cloud fraction, Δc , due to a small change in the assumed surface albedo, ΔA_s (Koelemeijer et al., 2001):

$$\Delta c = -\frac{1 - c}{A_c - A_s + R_c - R_s} \Delta A_s \quad (3)$$

As the albedo of vegetation is lower in the B-band than in the A-band (see Fig. 2), we expect to retrieve a higher cloud fraction with FRESKO-B for this type of surface. For ocean, the surface albedo chosen for the simulations is the same in the two spectral regions, so we do not expect any impact on the effective cloud fraction.

	Simulated cloud cases			
	1	2	3	4
Description	thin and low	thick and low	thin and high	thick and high
Cloud top	5 km	5 km	10 km	10 km
Cloud bottom	2 km	2 km	7 km	7 km
Total cloud optical thickness	9	42	9	42
Surface albedo	Ocean	Vegetation		
Oxygen A-band	0.02	0.2		
Oxygen B-band	0.02	0.05		

Table 1. Parameters for the four cloud cases used in the retrieval simulations.

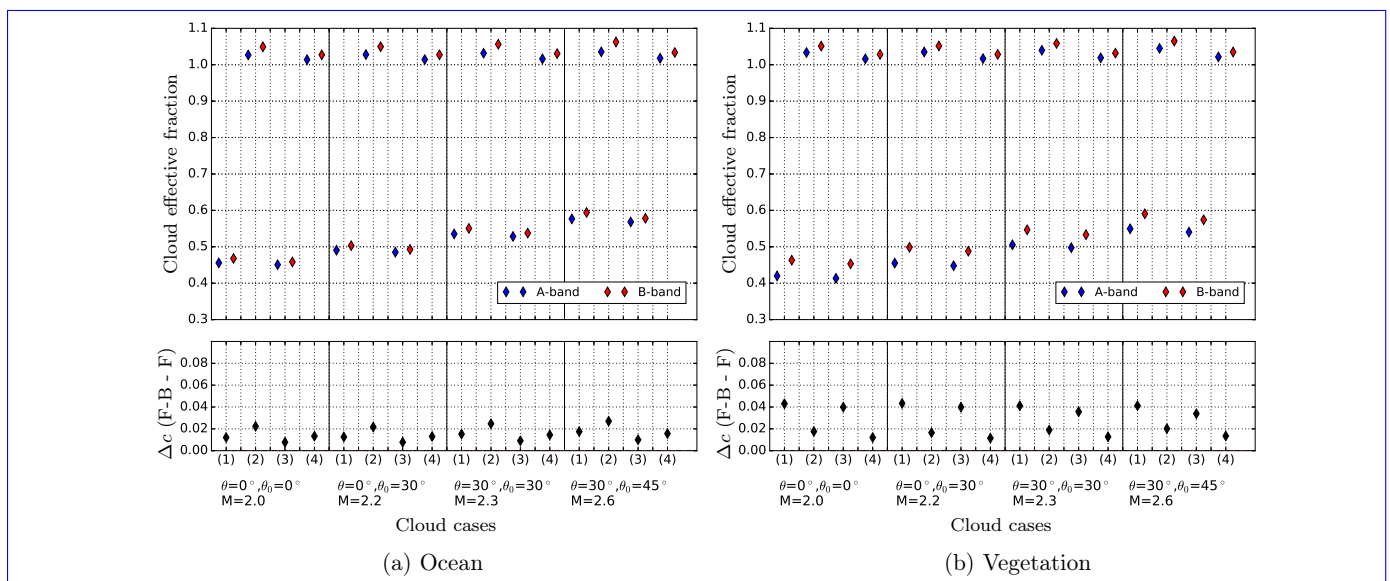


Figure 5. FRESKO-B (F-B) and FRESKO (F) effective cloud fraction retrievals (top panels) and effective cloud fraction differences (bottom panels) for the simulated cloud scenes over ocean (a) and vegetation (b). The cloud cases are described in Table 1. The simulations have been made with an azimuth angle difference of 90° .

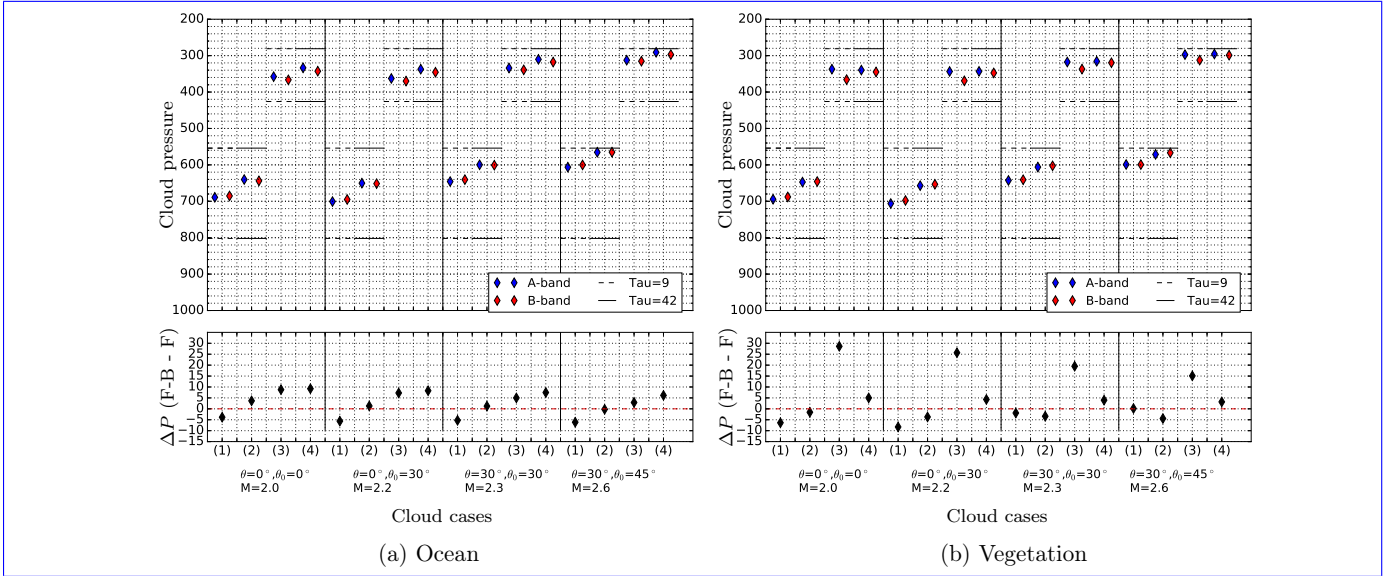


Figure 6. FRESKO-B (F-B) and FRESKO (F) cloud pressure retrievals (top panels) and cloud pressure differences (bottom panels) for the simulated cloud scenes over ocean (a) and vegetation (b). The cloud cases are described in Table 1. The simulations have been made with an azimuth angle difference of 90° .

4.3 Cloud pressure

In Figure 6, we see that the pressures retrieved by FRESKO-B and FRESKO in the simulations, indicate an altitude inside the cloud layer but well below the cloud top altitude. This feature is well known and common to algorithms which are based upon the hypothesis of a Lambertian cloud to retrieve the cloud top pressure (Saiedy et al., 1965; Vanbauce et al., 1998; Parol et al., 1999; Wang et al., 2008; Sneep et al., 2008). Indeed, real clouds do not act as perfect reflecting boundaries and the algorithm takes into account neither the photons reflected by the surface below the cloud, nor the photons penetration inside the cloud layer. In both cases, the photons path increases, as well as the oxygen absorption, which leads to a higher pressure than the cloud top. For optically thick clouds, the retrieved pressure is closer to the cloud top than for thinner clouds; the thick clouds are closer to the model of a Lambertian cloud. We also observe a decrease in the retrieved cloud pressures with increasing geometric air mass factor. Indeed, the larger the solar and viewing zenith angles are, the shallower the photon penetrates inside the cloud layer.

In Figure 6a, we can see that, over ocean, the difference in retrieved pressure between FRESKO-B and FRESKO is between about -5 hPa and +10 hPa, and increases with the cloud altitude. Indeed, the higher the cloud, the longer the photon's path under the cloud, where it undergoes oxygen absorption. As this path under the cloud is not taken into account in FRESKO and FRESKO-B, high cloud altitudes lead to larger pressure differences.

In Figure 6b, we can see that over vegetation, the difference of pressures between FRESKO-B and FRESKO is between about -10 hPa and +10 hPa, excepted for high (top altitude 10 km) optically thin clouds ($\tau = 9$) clouds where the difference

is much larger. The reason that FRESCO-B retrieves a (much) higher cloud pressure than FRESCO for high (thin) clouds is twofold. Firstly, the O₂ B-band is less strong than the O₂ A-band, so that radiation in the B-band penetrates deeper into clouds and the atmosphere than in the A-band, thus down to higher pressures. Secondly, for optically thin clouds there is a relatively large cloud-free part of the pixel in the FRESCO retrieval model (see Fig. 1). Since the surface albedo of land is higher in the O₂ A-band than in the O₂ B-band, there are more photons reflected by the surface in the A-band than in the B-band for the cloud-free part. These photons have experienced a high pressure and a large O₂ absorption. To compensate for this large O₂ absorption, the FRESCO retrieval places the cloud at a lower pressure than the FRESCO-B retrieval does. Both effects lead to a positive difference between FRESCO-B and FRESCO cloud pressures. The second effect explains the surface albedo bias of the FRESCO retrieval as compared to the FRESCO-B retrieval.

For the case of a thin low cloud, the FRESCO-B pressure is a bit lower than the FRESCO pressure. This holds for both the ocean and vegetation cases, except for the most oblique geometry for vegetation. This feature can be due to the Rayleigh scattering. Since the B-band is weaker than the A-band, multiple scattering between the cloud particles and the molecular Rayleigh scatterers above, inside, and below the cloud is stronger in the B-band than in the A-band. At 685 nm there is 50% more Rayleigh scattering than at 760 nm. Most Rayleigh scattering is above 5 km, so the pressure retrieved in the B-band is lowered by the scattering happening above the cloud, leading to a smaller (negative) difference of pressures.

The FRESCO algorithm works best for clouds over a dark surface, because then the major part of the radiation is coming from the cloud and not from the surface. When the surface albedo is increased in FRESCO, which happens when going from the O₂ B-band to the O₂ A-band, the cloud fraction decreases and the cloud pressure decreases as well (the cloud rises). This behaviour agrees with earlier studies, simulations and retrievals of FRESCO using the A-band when changing the surface albedo. In our simulations of clouds over vegetation the cloud fraction is indeed decreasing when going from the B-band to the A-band, and the cloud pressure is decreasing for high thin clouds by about 25-30 hPa when going from the B-band to the A-band.

5 Results

5.1 FRESCO in the oxygen A and B bands applied to GOME-2B data

In this section we compare the effective cloud fraction and cloud pressure retrieved by the two versions of FRESCO. We have run the two algorithms over one month of GOME-2B data in July 2014, excluding the snow and ice cases. The dataset contains 4 208 125 cases.

5.1.1 Cloud fraction

The differences of effective cloud fraction between FRESCO-B and FRESCO are shown of Figure 7a. ~~The distribution is centered around 0 and has a mean value of 0.003 and a standard deviation of 0.036.~~ We see the highest difference at high latitudes and over bright surfaces. The distributions of the differences between FRESCO-B and FRESCO retrievals over ocean

and land are shown on Figures 8a and b and the mean values are summarized in Table 2. ~~Table 2 shows the mean effective cloud fraction retrieved in the oxygen A and B bands for different kinds of surface. We can see that 2.~~ Both over ocean and land, the differences are almost zero, ~~which is with mean values of +0.0066 and -0.0052 respectively.~~ These very small differences are expected as the cloud fraction is mainly determined through the reflectance measurements performed in the non-absorbed part of the spectrum. Over land, the distribution of the cloud effective fraction differences is more spread than over ocean, which is due to the more important variability of the surface albedo over this type of surface.

However, over vegetation, both Figure 7a and Table 2 show that the effective cloud fraction retrieved in the B-band is lower than the one retrieved in the A-band while the simulations suggest the ~~opposite~~ opposite (cf. section 4.2). Indeed, although the surface reflectance has an anisotropic BRDF function, the surface is often assumed to be Lambertian, as in many situations the full BRDF is not available or the radiative transfer code used is not equipped to handle the BRDF properly. This is the case in FRESCO and ~~FRECO-B~~ FRESCO-B for which we use a MODE-LER surface albedo climatology established from SCHIAMACHY measurements (Tilstra et al., 2017). This anisotropy is stronger over vegetation which has non-isotropic elements like dense tree with heterogeneous leaves and shadow effects) and in the NIR (0.7-2.5 μm) where the atmosphere is more transparent. Recently, ~~(Lorente et al., 2018)~~ Lorente et al. (2018) has shown that this assumption of a Lambertian surface, leads to across-track biases on satellite retrievals that use those climatologies, like the effective cloud fraction, for the solar and viewing geometries of GOME-2. It is shown by Lorente et al. (2018), that the western part of the swath has the most biased FRESCO cloud fraction. Consequently, for the vegetation case, we have recompiled the mean effective cloud fractions for FRESCO and FRESCO-B, distinguishing the eastern, nadir and western pixels of the GOME-2B swath. The results are ~~summarised~~ summarized in Table 3. We see that for the eastern and nadir pixels of the swath, the effective cloud fraction retrieved by FRESCO-B is higher than the one retrieved by FRESCO, which agrees with our simulations, while for the western pixels, this is the opposite. Indeed, for the western part of the swath, the cloud effective fraction retrieved in the O₂ A-band is too high because the surface albedo ~~used~~ is high (red edge (Tilstra et al., 2017)) and the anisotropy is stronger, while in the O₂ B-band, the surface albedo is low and the anisotropy is smaller. Consequently, the error on A_s due to the assumption of a Lambertian surface has a smaller impact on effective cloud fraction in the B-band than in the A-band. This difference of behaviour of the retrievals according to the part of the swath corroborates the conclusions of Lorente et al. (2018).

5.1.2 Cloud pressure

In this subsection, we limit our study to the pixels for which the FRESCO effective cloud fraction is higher than or equal to 0.1, indeed, as mentioned in Wang et al. (2008), cloud pressures have big uncertainties when effective cloud fraction is lower than 0.1. This selection leaves us with 3 237 790 pixels.

The differences of cloud pressures retrieved with FRESCO-B and FRESCO are shown on ~~panel~~ Figure 7b. ~~The mean~~ We have also analyzed the differences separating the pixels according to the underlying surface type. The distributions of the differences between FRESCO-B and FRESCO are shown on Figures 8c and d while the mean values are summarized in Table 4. The mean cloud pressures are 736 ± 195 hPa in the A-band and 747 ± 176 hPa in the B-band. As already mentioned, this behaviour is expected as the B-band is less absorbing. ~~We also have analyzed the differences separating the pixels according~~

	Nr. of cases	Effective cloud fraction		Difference
		FRESCO-B	FRESCO	FRESCO-B - FRESCO
All cases	4 208 125	$\bar{c} = 0.389, SD=0.304$	$\bar{c} = 0.387, SD=0.297$	$\overline{\Delta c} = +0.0027, SD=0.036$
Ocean	2 582 287	$\bar{c} = 0.395, SD=0.298$	$\bar{c} = 0.388, SD=0.294$	$\overline{\Delta c} = +0.0066, SD=0.023$
Land	1 447 322	$\bar{c} = 0.374, SD=0.312$	$\bar{c} = 0.379, SD=0.302$	$\overline{\Delta c} = -0.0052, SD=0.051$
Vegetation	920 902	$\bar{c} = 0.380, SD=0.320$	$\bar{c} = 0.390, SD=0.310$	$\overline{\Delta c} = -0.0106, SD=0.064$

Table 2. Mean effective cloud fractions and standard deviations for FRESCO-B and FRESCO, as well as the mean differences, for GOME-2B measurements in July 2014 over different surfaces. The Vegetation category is a sub-part of the Land one.

Vegetation	Nr. of cases	Effective cloud fraction		Difference
		FRESCO-B	FRESCO	FRESCO-B - FRESCO
All swath	920 902	$\bar{c} = 0.380, SD=0.320$	$\bar{c} = 0.390, SD=0.310$	$\overline{\Delta c} = -0.0106, SD=0.064$
West	269 861	$\bar{c} = 0.396, SD=0.322$	$\bar{c} = 0.440, SD=0.290$	$\overline{\Delta c} = -0.0443, SD=0.078$
Nadir	286 768	$\bar{c} = 0.342, SD=0.311$	$\bar{c} = 0.338, SD=0.304$	$\overline{\Delta c} = +0.0042, SD=0.050$
East	364 273	$\bar{c} = 0.396, SD=0.322$	$\bar{c} = 0.394, SD=0.323$	$\overline{\Delta c} = +0.0026, SD=0.053$

Table 3. Mean effective cloud fractions and standard deviations for FRESCO-B and FRESCO, as well as the mean differences, for GOME-2B measurements in July 2014 over vegetation. We distinguish the eastern (pixels 1 to 8), nadir (pixels 9 to 16) and western (pixels 17 to 24) parts of the swath .

	Nr. of cases	Cloud pressure (hPa)		Difference (hPa)
		FRESCO-B	FRESCO	FRESCO-B - FRESCO
All cases	3 237 790	$\bar{P} = 747, SD=176$	$\bar{c} = 736$ $\bar{P} = 736, SD=195$	$\overline{\Delta P} = 11.5, SD=44.9$
Ocean	2 032 709	$\bar{P} = 763, SD=179$	$\bar{c} = 749$ $\bar{P} = 749, SD=199.$	$\overline{\Delta P} = 13.9, SD=42.1$
Land	1 063 650	$\bar{P} = 715, SD=165$	$\bar{c} = 709$ $\bar{P} = 709, SD=183$	$\overline{\Delta P} = 6.31, SD=49.1$
Vegetation	651 903	$\bar{P} = 689, SD=157$	$\bar{c} = 681$ $\bar{P} = 681, SD=174$	$\overline{\Delta P} = 8.52, SD=50.2$

Table 4. Mean cloud pressures and standard deviations for FRESCO-B and FRESCO, as well as the mean differences, for FRESCO-B and FRESCO for GOME-2B measurements in July 2014 over different surfaces. The Vegetation category is a sub-part of the Land one.

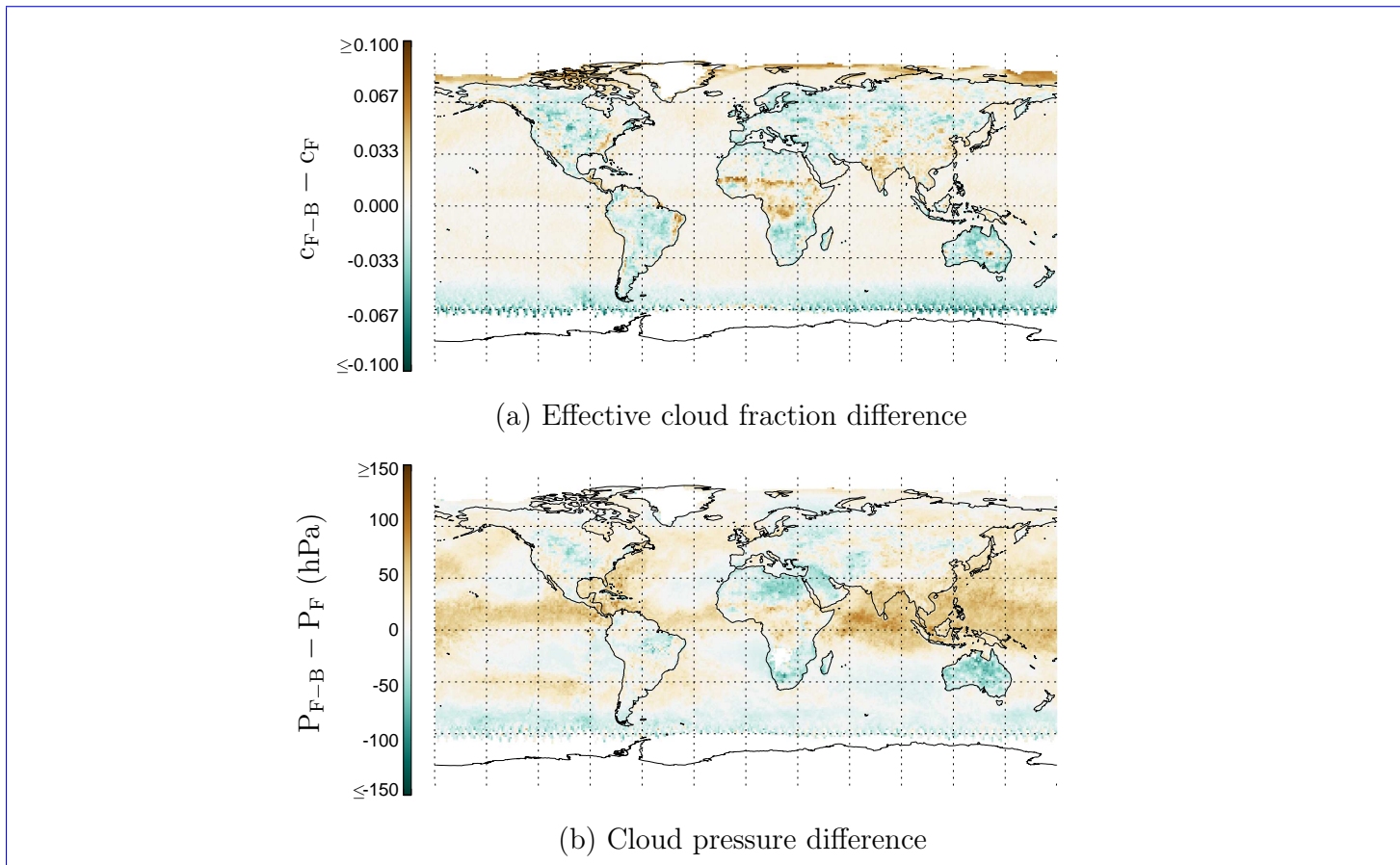


Figure 7. Differences of effective cloud fraction and cloud pressure between FRESCO-B and FRESCO for July 2014.

	Nr. of cases	Chi-squared		Difference
		FRESCO-B	FRESCO	FRESCO-B - FRESCO
All cases	3 237 790	$\bar{\chi}_2 = 10.3, SD=9.7$	$\bar{\chi}_2 = 5.8, SD=8.2$	$\overline{\Delta\chi_2} = 4.5, SD=9.1$
Ocean	2 032 709	$\bar{\chi}_2 = 10.1, SD=10.0$	$\bar{\chi}_2 = 5.5, SD=8.1$	$\overline{\Delta\chi_2} = 4.7, SD=9.2$
Land	1 063 650	$\bar{\chi}_2 = 10.3, SD=9.0$	$\bar{\chi}_2 = 6.4, SD=8.3$	$\overline{\Delta\chi_2} = 3.9, SD=8.8$
Vegetation	651 903	$\bar{\chi}_2 = 9.72, SD=9.0$	$\bar{\chi}_2 = 6.8, SD=8.4$	$\overline{\Delta\chi_2} = 2.9, SD=8.6$

Table 5. Mean Chi-squared and standard deviations for FRESCO-B and FRESCO, as well as the mean differences for FRESCO-B and FRESCO for GOME-2B measurements in July 2014 over different surfaces. The Vegetation category is a sub-part of the Land one.

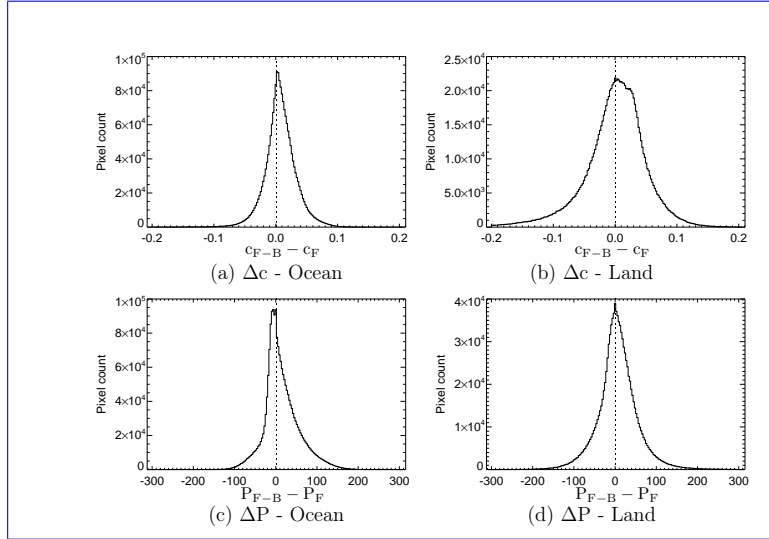


Figure 8. Histograms of the differences of effective cloud fractions (panels a and b) and of cloud pressures (panels c and d) between FRESCO-B and FRESCO for July 2014, over ocean (panels a and c) and over land (panels b and d).

Vegetation	Nr. of cases $c_{\text{eff}} > 0.1$	Cloud pressure (hPa)		Difference (hPa)
		FRESCO-B	FRESCO	FRESCO-B - FRESCO
All swath	651 903	$\bar{P} = 689$, SD=157	$\bar{P} = 681$, SD=174	$\overline{\Delta P} = 8.52$, SD=50.2
West	205 093	$\bar{P} = 699$, SD=154	$\bar{P} = 706$, SD=173	$\overline{\Delta P} = -7.19$, SD=54.1.
Nadir	185 336	$\bar{P} = 705$, SD=153	$\bar{P} = 689$, SD=171	$\overline{\Delta P} = +16.2$, SD=46.6
East	261 474	$\bar{P} = 671$, SD=161	$\bar{P} = 655$, SD=174	$\overline{\Delta P} = +15.4$, SD=46.6

Table 6. Mean cloud pressure for FRESCO-B and FRESCO, as well as differences, for GOME-2B measurements in July 2014 over vegetation. We distinguish the eastern (pixels 1 to 8), nadir (pixels 9 to 16) and western (pixels 17 to 24) parts of the swath .

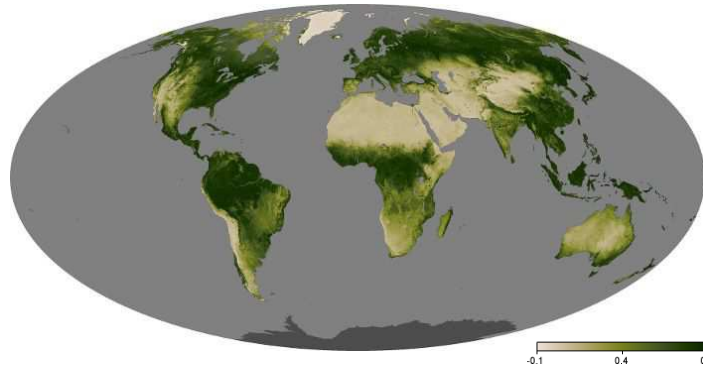


Figure 9. MODIS NDVI index for July 2014. Source: https://earthobservatory.nasa.gov/global-maps/TRMM3B43M/MODNDVI_M

~~to the underlying surface type. The mean cloud pressures as well the pressure differences between FRESKO-B and FRESKO for each surface are shown in Table 4.~~

Over ocean, the mean cloud pressure is 763 ± 179 hPa with FRESKO-B, while with FRESKO, this value is 749 ± 199 hPa and the mean difference is 13.9 ± 42.1 hPa. The surface albedo of ocean is very similar in oxygen A and B bands, therefore the differences we observe in cloud pressures are only due to a difference of absorption inside and under the cloud. For instance, on Figure 7b, we observe the larger pressure differences in the ITCZ, where there are a lot of high clouds (cumulonimbus, cirrus), which agrees with the simulations presented in section 4.3. Over ocean, the coefficient of correlation between the difference of pressures and the difference of cloud effective fractions is 0.0949.

Over land, the cloud pressure difference is also positive on average with a mean value of 6.31 ± 49.1 hPa. However, as we can see on Figure 8d, the range of the values is larger than over ocean with a lot of pixels having a negative difference of pressure. This is again due to the high variability of the surface albedos for this type of surface and is coherent with the observations we have made on the simulations. Over land, the coefficient of correlation between the difference of pressures and the difference of cloud effective fractions is 0.523.

Over vegetation, we observe on average small pressure differences, the mean cloud pressures being 689 ± 157 hPa with FRESKO-B and 681 ± 174 hPa with FRESKO. Figure 9 allows us to visualize the vegetated areas in July (the higher the index, the greener the area). We can see on Figure 7b that those areas (North-West of North America, North of South America, Northern part of Europe and Asia) have, on average, small cloud pressures differences, which also agrees with the simulations we have described in Section 4.3. Moreover, as seen on Figure 2, the surface albedo of vegetation surfaces is much lower in the oxygen B-band than in the oxygen A-band. Therefore, the contribution of the ground in the reflectances is lower in the oxygen B-band than in the A-band, so we expect more accurate retrievals in the B-band. Table 5 shows the value of χ^2 obtained with FRESKO and FRESKO-B for different types of surfaces. We can see that the χ^2 are always higher for FRESKO-B than FRESKO, this is due to the fits of the algorithms: the difference between the simulated reflectances and the measured ones is

always higher in the B-band. However, we can notice that while with FRESCO the χ^2 is the higher for vegetation, this is the opposite with FRESCO-B. FRESCO-B is the most suited over vegetated surfaces.

As we have previously mentioned the anisotropy of the surface albedo over vegetation and its consequences on effective cloud fraction retrieval, we have compiled the mean cloud pressure retrieved with FRESCO and FRESCO-B for the different part of the swath. The results are summarized in table 6. We see that the pressure retrieved with FRESCO-B is higher than the one retrieved with FRESCO for the eastern and nadir part of the swath, while it is the opposite for the western pixels. This last observation is due to the bias albedo on the western part of the swath: As mentioned in section 5.1.1, the use of a Lambertian albedo in FRESCO and FRESCO-B leads to artificially high values of FRESCO effective cloud fraction for those pixels and consequently to artificially low values of FRESCO cloud pressure (see equation 1). On the other side, the B-band is less impacted by the bias albedo and so is the FRESCO-B cloud pressure. This albedo bias impacting more the O₂ A-band measurements than the O₂ B-band leads to a higher cloud pressure difference than expected.

The mean difference of cloud pressure between FRESCO-B and FRESCO is 11.5 hPa, but the difference does not have the same signification according to the underlying surface: over ocean, the two pressures indicate different pieces of information about the vertical structure of the cloud layer, while over vegetation, the difference is partly due to the difference of surface albedo. Over vegetation, the FRESCO-B cloud pressure is more accurate than the FRESCO one. In the future we also recommend to use a surface albedo database which takes into account the anisotropy of the surface to avoid biases.

5.2 Comparison with ground-based measurements

5.2.1 Cloudnet target classification product

Cloudnet is a network of ground-based measurement facilities for the evaluation of clouds and aerosols in forecast models. Cloudnet started around 2001 with three observations sites (at Cabauw, Palaiseau and Chilbolton) and includes now five other permanent sites (at Juelich, Leipzig, Lindenberg, Mace Head and Potenza). In this study, we reject the data from the Mace Head site as this one is on the seaside, while the other sites can be considered as surrounded by vegetation. These sites are equipped with active sensors such as lidar and Doppler millimeter-wave radar that provide vertical profiles of cloud and aerosol properties as well as ice and liquid cloud water content at high temporal and spatial resolution. In this study, we use the Cloudnet Level 2 classification product (Illingworth et al., 2007) which is based on the combination of the vertically pointing Doppler cloud radar and backscatter lidar, and is available approximately every 30s. This product classifies each vertical layer as 1 of 11 classes, which distinguishes ice and water clouds, precipitation, aerosols, insects, clear sky and combination thereof. Indeed the radar detects large particles such as rain and drizzle drops, ice particles and insects, while the lidar is sensitive to smaller particles such as cloud droplets and aerosols. The target classification product also contains cloud top height and cloud base height. Cloud top and cloud base heights correspond respectively to the highest and lowest altitudes of the backscatter altitude grid boxes, that have clouds. Consequently, for ~~multilayer~~ multilayers cloud situations, the cloud top height refers to the top of the highest layer, while the cloud base height refers to the base of the lowest one.

5.2.2 Methodology

In this part, we perform comparisons between the two versions of FRESCO and the Cloudnet target classification product and cloud boundaries for the seven Cloudnet observations sites in July and August 2014. For every GOME-2B pixel collocated with a Cloudnet site, we select 1 hour (+/-30 min) of Cloudnet target classification. For every cloud height measurement from GOME-2B, there are about 120 (temporal) \times 495 (vertical) radar/lidar backscatter Cloudnet pixels, which are classified as one of the 11 categories mentioned earlier. In this study, we keep the cloudy cases, which consist of the classes "ice" and "cloud droplets only" as well as the precipitations ("drizzle or rain", "drizzle or rain and cloud droplets", "ice and supercooled droplets", "melting ice", "melting ice and cloud droplets"). We then determine the height distributions of the backscatter pixels, from 270 to 15000 m with a bin size of 270 m, following the method defined by Wang and Stammes (2014). If the distribution presents a unique mode, without an interruption by a clear-sky pixel, the cloud is considered as a monolayer case. Otherwise, we classify it as a ~~multilayer~~-multilayers case. The Cloudnet cloud top and base heights are calculated averaging the cloud top and base heights in the 1-hour period around GOME-2B overpass time.

5.2.3 Results

For the two considered months, the collocation process provides us 339 collocated cloud cases, that we further filter as follows:

- 15 – We keep only the cases for which the Cloudnet cloud fraction is higher than 0.05. This Cloudnet cloud fraction is calculated dividing the number of cloudy pixels by the number of pixels accumulated during the one hour period. With this filtering, we exclude the almost cloud-free scenes.
- Then we further filter the Cloudnet data excluding the cases for which the standard deviation of the cloud top height exceeds 1.5 km. This criterion used in Veefkind et al. (2016) allows to avoid the cases with a large temporal variability during the satellite overpass.
- 20 – Finally, we filter the cases according to the FRESCO-B cloud fraction, excluding the cases with $c < 0.1$. Indeed, it is known that the FRESCO cloud pressures are often too low when the cloud fraction is lower than 0.1 (Wang et al., 2008).

Those ~~eriteteria~~-criteria leave us with 138 cases: monolayer clouds are represented in panel (a) of Figure 10, while ~~multilayer~~-multilayers situations are shown in panel (b). On both panels, the clouds are ordered by increasing Cloudnet top height altitude.

25 We can distinguish three different regimes, for which the clouds properties are resumed in Table 7 :

1. For 68 cases (49% of the clouds), the FRESCO-B cloud pressure is inside the Cloudnet cloud boundaries. This population ~~correponds~~-corresponds to 46% of the monolayer clouds and to 59% of the multilayers clouds and concern vertically extended middle to high clouds (mean cloud height of 5.5 km and mean cloud thickness of 3.7 km). With FRESCO, the retrieved pressure is inside the Cloudnet boundaries for 64 cases (46% of the clouds) that are a bit higher (mean altitude of 6 km) and a bit thicker (mean thickness 4.1 km) than for FRESCO-B.
- 30

2. In 51 cases, the FRESCO-B cloud pressure indicate an altitude lower than the Cloudnet cloud boundaries. This expected behaviour concerns mainly middle to high (mean cloud height of 5.7 km) monolayer clouds with a limited vertical extent (mean geometrical thickness of 2 km). As already shown (Wang et al., 2008; Sneep et al., 2008), light can penetrate within the clouds where it is absorbed. As this phenomenon is not taken into account in FRESCO, the retrieved cloud height is lower than the cloud top height. This is the case in other retrieval algorithms based on oxygen absorption (Ferlay et al., 2010; Desmons et al., 2013). FRESCO cloud altitude is lower than Cloudnet for 53 cases. It concerns lower (mean altitude of 5.1 km) and thinner (mean thickness of 1.6 km) clouds than for FRESCO-B.
3. In the last 19 cases, the FRESCO-B cloud pressure indicate an altitude higher than the Cloudnet cloud boundaries. This concerns mainly monolayer low (mean top height of 1.9 km) thin (mean cloud geometrical thickness of 0.4 km) clouds. In those cases, many photons are transmitted and reach the surface before being reflected back to space. This situation concerns 21 cases for FRESCO, also quite low (mean top height of 2.3 km) and thin (mean cloud geometrical thickness of 0.5 km).

The results shown by FRESCO-B are very similar to the ones obtained with FRESCO, however FRESCO-B performs slightly better than FRESCO as with FRESCO-B, 49% of the retrieved cloud pressures are within the Cloudnet cloud boundaries, while this number is 46% with FRESCO. Also for monolayer clouds, FRESCO-B retrieves an altitude within Cloudnet range in 46% of the cases, while this value is 41% with FRESCO. For multilayers clouds, FRESCO performs ~~slightly~~ slightly better than FRESCO-B, indicating an altitude within Cloudnet range for 62% of the cases, while this value is 59% with FRESCO-B. Those performances are very promising and show that it would be very valuable to use both FRESCO-B and FRESCO retrievals, particularly in the case of multilayers clouds.

Like in the previous section, we have also filtered the pixels according to their position in the swath, keeping only the nadir and eastern pixels of the swath. Then, the FRESCO-B cloud pressure is inside the Cloudnet cloud boundaries in 49% of the cases, as with all the pixels, for FRESCO, this number stays stable too as it is now 47 % (46% for all the swath). Excluding the western pixels of the swath doesn't seem to change the performances of the two algorithms. However, the size of the database is too small to deduce strong conclusions. Also, we should keep in mind that the Cloudnet sites are situated in Europe, while the study made by Lorente et al. (2018) takes place in Amazonia. The vegetation is probably very different between those two regions (grass/forests) and so are their surface albedos.

6 Conclusions

In this paper, we present a new cloud retrieval algorithm called FRESCO-B, for GOME-2 measurements in the oxygen B-band. FRESCO-B is based on the algorithm FRESCO, which uses measurements in the oxygen A-band to retrieve cloud properties (effective cloud fraction and cloud pressure). However, while the surface albedo in the O₂ A-band is very low over ocean, the albedo takes larger values over vegetation, which can lead to biases in the retrievals. In this paper, we apply FRESCO-B to GOME-2B measurements in the O₂ B-band, which is a wavelength range where the surface albedo stays relatively low

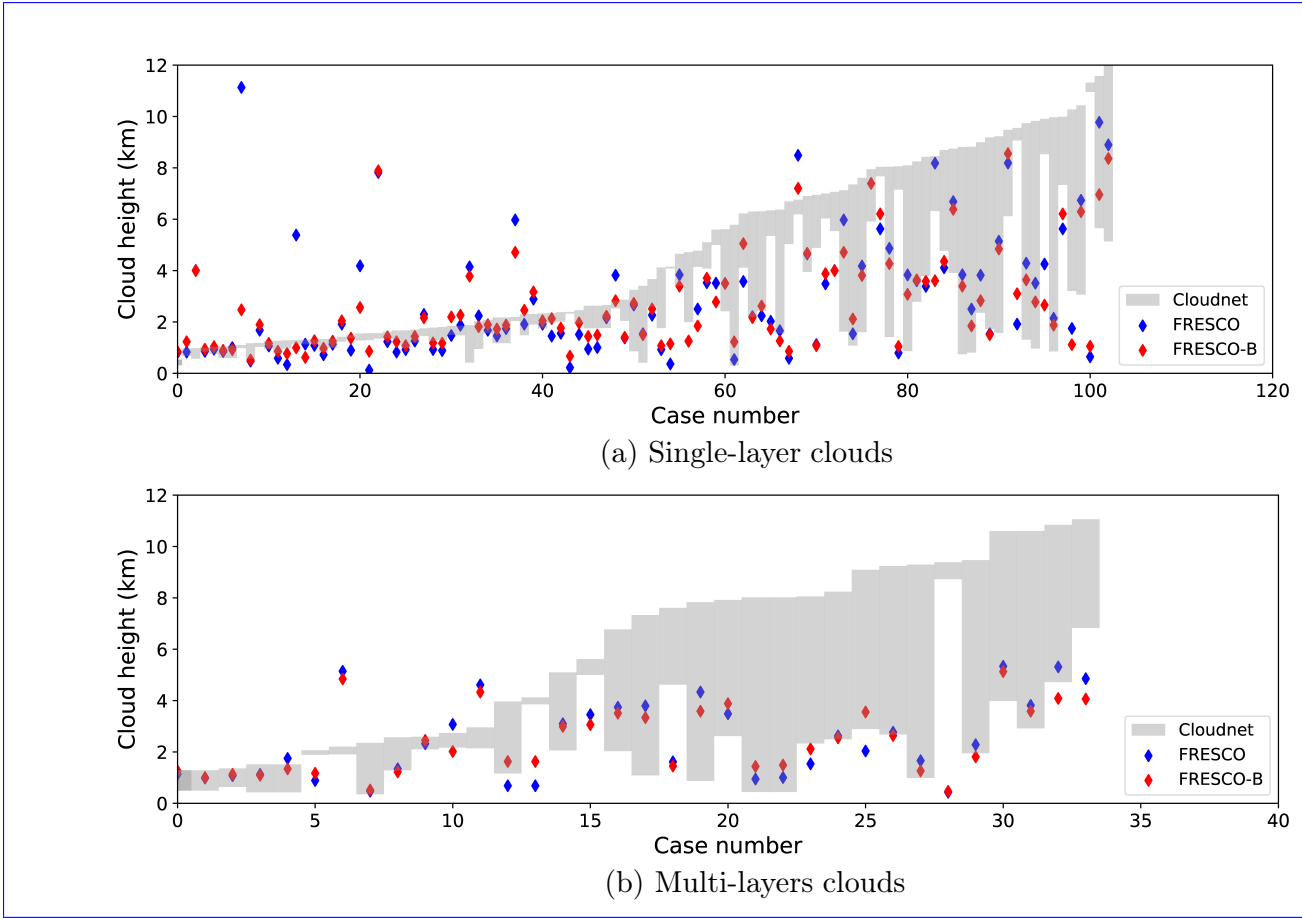


Figure 10. Comparison of the cloud heights retrieved by FRESKO-B, FRESKO and Cloudnet in July and August 2014, for six [Cloudnet](#) sites (Hyytiala, Juelich, Leipzig, Lindenberg, Palaiseau and Potenza).

	Collocated cases	Inside CN. range	Lower than CN. altitude range	Higher than CN. altitude range
FRESKO-B	138	68 (49%) $\overline{z_{CN}}=5.5$ km $\overline{h_{CN}}=3.7$ km	51 (35%) $\overline{z_{CN}}=5.7$ km $\overline{h_{CN}}=2.0$ km	19 (15%) $\overline{z_{CN}}=1.9$ km $\overline{h_{CN}}=0.4$ km
FRESKO	138	64 (46%) $\overline{z_{CN}}=6$ km $\overline{h_{CN}}=4.1$ km	53 (38%) $\overline{z_{CN}}=5.1$ km $\overline{h_{CN}}=1.6$ km	21 (16%) $\overline{z_{CN}}=2.3$ km $\overline{h_{CN}}=0.5$ km

Table 7. Number of collocated cloud cases for which FRESKO-B and FRESKO cloud pressures are inside, upper or higher than the Cloudnet cloud pressure range. For each case, the mean Cloudnet cloud altitude ($\overline{z_{CN}}$) and geometrical thickness ($\overline{h_{CN}}$) are indicated. Collocated Cloudnet/GOME-2B for July and August 2014 for the Cloudnet sites of Hyytiala, Juelich, Leipzig, Lindenberg, Palaiseau and Potenza.

whatever the underlying surface is.

First, we have simulated cloudy scenes which have shown that FRESCO-B and FRESCO retrievals indicate an altitude inside the cloud layer but well below the top, which was expected. We have noted that the difference between the pressures obtained with the two algorithms depends on the geometry of the scenes and ranges between -10 hPa ~~et~~ and +10 hPa, excepted for high thin clouds where it can reach +30 hPa. Then, inter-comparisons between FRESCO-B and FRESCO over one month of GOME-2B data have shown that the effective cloud fraction retrieved is very similar in the two bands. These comparisons have also revealed that FRESCO-B retrieves a higher cloud pressure than FRESCO (mean difference of 11.5 hPa), and is more accurate over vegetation (χ^2 is lower over vegetation than for other surfaces). Finally, we have validated FRESCO-B and FRESCO to in-situ data over vegetation obtained with the Cloudnet network of instruments. These comparisons have shown that FRESCO-B and FRESCO can retrieve a pressure which stands inside the cloud layers for cloud that are not too far from the Lambertian model. In the future, we would like to apply FRESCO-B and FRESCO to TROPOMI (Veefkind et al., 2012) measurements in the O₂ A and B-bands. Indeed its high spatial resolution (7 km × 3.5 km) compared to previous spectrometers (80 km × 40 km for GOME-2B) should provide much more detailed cloud structures.

In the future, the authors would like to merge the measurements in the O₂ A and B-bands in order to retrieve more information about the vertical structure of cloud layers. Excluding the western part of the swath of GOME-2B in the comparisons have improved the performances of FRESCO-B, which corroborates the conclusions of Lorente et al. (2018) on the surface albedo bias for these pixels. Since this bias leads to biases in cloud retrievals, the authors recommend to take into account the anisotropy of the surface reflectances for future surface albedo climatologies (on-going work within AC SAF).

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References

- Andrews, T., Gregory, J., Webb, M. J., and Taylor, K. E.: Forcing, feedbacks and climate sensitivity in CMIP5 coupled atmosphere-ocean climate models, *Geophys. Res. Lett.*, 39, L09712, <https://doi.org/10.1029/2012GL051607>, 2012.
- Boersma, K. F., Eskes, H. J., and Brinksma, E. J.: Error analysis for tropospheric NO₂ retrieval from space, *J. Geophys. Res.*, 109, <https://doi.org/10.1029/2003JD003962>, 2004.
- Bony, S. and Dufresne, J.-L.: Marine boundary layer clouds at the heart of tropical cloud feedback uncertainties in climate models, *Geophys. Res. Lett.*, 32, L20806, <https://doi.org/10.1029/2005GL023851>, 2005.
- De Haan, J. F., Bosma, P. B., and Hovenier, J. W.: The adding method for multiple scattering calculations of polarized light, *Astron. Astrophys.*, pp. 371–391, 1987.
- 10 Desmons, M., Ferlay, N., Parol, F., Mcharek, L., and Vanbauce, C.: Improved information about the vertical location and extent of monolayer clouds from POLDER3 measurements in the oxygen A-band, *Atmos. Meas. Tech.*, 6, 2221–2238, <https://doi.org/10.5194/amt-6-2221-2013>, 2013.
- Ferlay, N., Thieuleux, F., Cornet, C., Davis, A. B., Dubuisson, P., Ducos, F., Parol, F., Riédi, J., and Vanbauce, C.: Toward New Inferences about Cloud Structures from Multidirectional Measurements in the Oxygen A Band: Middle-of-Cloud Pressure and Cloud Geometrical Thickness from POLDER-3/PARASOL, *J. Appl. Meteor. Climatol.*, 49, 2492–2507, <https://doi.org/10.1175/2010JAMC2550.1>, 2010.
- 15 Fournier, N., Stammes, P., de Graaf, M., van der A, R., Pitters, A., Grzegorski, M., and Kokhanovsky, A.: Improving cloud information over deserts from SCIAMACHY Oxygen A-band measurements, *Atm. Chem. Phys.*, 6, 163–172, <https://doi.org/10.5194/acp-6-163-2006>, 2006.
- Gordon, I., Rothman, L., Hill, C., Kochanov, R., Tan, Y., Bernath, P., Birk, M., Boudon, V., Campargue, A., Chance, K., Drouin, B., Flaud, J.-M., Gamache, R., Hodges, J., Jacquemart, D., Perevalov, V., Perrin, A., Shine, K., Smith, M.-A., Tennyson, J., Toon, G., Tran, H., Tyuterev, V., Barbe, A., Császár, A. A., Devi, V., Furtenbacher, T., Harrison, J., Hartmann, J.-M., Jolly, A., Johnson, T., Karman, T., Kleiner, I., Kyuberis, A., Loos, J., Lyulin, O., Massie, S., Mikhailenko, S., Moazzen-Ahmadi, N., Müller, H., Naumenko, O., Nikitin, A., Polyansky, O., Rey, M., Rotger, M., Sharpe, S., Sung, K., Starikova, E., Tashkun, S., Auwera, J. V., Wagner, G., Wilzewski, J., Wcisło, P., Yu, S., and Zak, E.: The HITRAN 2016 molecular spectroscopic database, *J. Quant. Spectrosc. Radiat. Transfer*, 203, 3 – 69, <https://doi.org/10.1016/j.jqsrt.2017.06.038>, 2017.
- 20 Illingworth, A. J., Hogan, R. J., O'Connor, E. J., Bouniol, D., Delanoë, J., Pelon, J., Protat, A., Brooks, M. E., Gaussiat, N., Wilson, D. R., Donovan, D. P., Baltink, H. K., van Zadelhoff, G.-J., Eastment, J. D., Goddard, J. W. F., Wrench, C. L., Haeffelin, M., Krasnov, O. A., Russchenberg, H. W. J., Piriou, J.-M., Vinit, F., Seifert, A., Tompkins, A. M., and Willén, U.: CLOUDNET - Continuous evaluation of Cloud Profiles in Seven Operational Models using Ground-Based Observations, *Bull. Amer. Meteor. Soc.*, 88, 883–898, <https://doi.org/10.1175/BAMS-88-6-883>, 2007.
- 30 Koelemeijer, R. B. A., Stammes, P., Hovenier, J. W., and de Haan, J. F.: A fast method for retrieval of cloud parameters using Oxygen A band measurements from GOME, *J. Geophys. Res.*, 106, 3475–3490, <https://doi.org/10.1029/2000JD900657>, 2001.
- Kuze, A. and Chance, K. V.: Analysis of cloud top height and cloud coverage from satellites using the O₂ A and B bands, *J. Geophys. Res.*, 99, 14,481–14,492, <https://doi.org/10.1029/94JD01152>, 1994.
- 35 Lelli, L., Kokhanovsky, A. A., Rozanov, V. V., Vountas, M., Sayer, A. M., and Burrows, J. P.: Seven years of global retrieval of cloud properties using space-borne data of GOME, *Atmos. Meas. Tech.*, 5, 1551–1570, <https://doi.org/10.5194/amt-5-1551-2012>, 2012.

- Lindstrot, R., Preusker, R., Ruhtz, T., Heese, B., Wiegner, M., Lindemann, C., and Fischer, J.: Validation of MERIS cloud top pressure using airborne lidar measurements, *J. Appl. Meteor. Clim.*, 45, 1612–1621, <https://doi.org/10.1175/JAM2436.1>, 2006.
- Lorente, A., Boersma, K. F., Stammes, P., Tilstra, L. G., Richter, A., Yu, H., Kharbouche, S., and Muller, J.-P.: The importance of surface reflectance anisotropy for cloud and NO₂ retrievals from GOME-2 and OMI, *Atmos. Meas. Tech.*, 11, <https://doi.org/10.5194/amt-11-4509-2018>, 2018.
- 5 Marshak, A. and Knyazikhin, Y.: The spectral invariant approximation within canopy radiative transfer to support the use of the EPIC/DSCOVER oxygen B-band for monitoring vegetation, *J. Quant. Spectrosc. Radiat. Transfer.*, 191, 7 – 12, <https://doi.org/10.1016/j.jqsrt.2017.01.015>, 2017.
- Parol, F., Buriez, J.-C., Vanbauce, C., Couvert, P., Séze, G., Goloub, P., and Cheinet, S.: First results of the POLDER Earth Radiation Budget and Clouds operational algorithm, *IEEE Trans. Geosci. Remote Sens.*, 37, 3, 1597–1612, 1999.
- Preusker, R., Fischer, J., P., A., Bennartz, R., and Schüller, L.: Cloud-top pressure retrieval using the oxygen A-band in the IRS-3 MOS instrument, *Int. J. Remote. Sens.*, 28, 1957–1967, <https://doi.org/10.1080/01431160600641632>, 2007.
- Saiedy, F., Hilleary, D. T., and Morgan, W. A.: Cloud-top altitude measurements from satellites, *Applied Optics*, 4, 495–500, 1965.
- Sneep, M., de Haan, J. F., Stammes, P., Wang, P., Vanbauce, C., Vasilkov, A. P., and Levelt, P. F.: Three-way comparison between OMI and PARASOL cloud pressure products, *J. Geophys. Res.*, 113, <https://doi.org/10.1029/2007JD008694>, 2008.
- 15 Stammes, P.: Satellite Radiance Modelling in the UV-Visible Range, in: *Current Problems in Atmospheric Radiation*, pp. 385–388, A. Deepak, Hampton, VA., edited by Smith W. and Timofeyev Y., 2001.
- Stammes, P., de Haan, J. F., and Hovenier, J. W.: The polarized internal radiation field of a planetary atmosphere, *Astronomy and Astrophysics*, 225, 239–259, 1989.
- 20 Stammes, P., Sneep, M., de Haan, J. F., Veefkind, J. P., Wang, P., and Levelt, P. F.: Effective cloud fractions from the Ozone Monitoring Instrument: Theoretical framework and validation, *J. Geophys. Res.: Atmos.*, 113, <https://doi.org/10.1029/2007JD008820>, 2008.
- Tilstra, L. G., Tuinder, O. N. E., Wang, P., and Stammes, P.: Surface reflectivity climatologies from UV to NIR determined from Earth observations by GOME-2 and SCIAMACHY, *J. Geophys. Res.: Atmos.*, 122, 4084–4111, <https://doi.org/10.1002/2016JD025940>, 2017.
- Vanbauce, C., Buriez, J., Parol, F., Bonnel, B., Séze, G., and Couvert, P.: Apparent pressure derived from ADEOS-POLDER observations in the oxygen A-band over ocean, *Geophys. Res. Lett.*, 25, 3159–3162, 1998.
- 25 Veefkind, J. P., Aben, I., McMullan, K., Förster, H., de Vries, J., Otter, G., Claas, J., Eskes, H. J., de Haan, J. F., Kleipool, Q., van Weele, M., Hasekamp, O., Hoogeveen, R., Landgraf, J., Snel, R., Tol, P., Ingmann, P., Voors, R., Kruizinga, B., Vink, R., Visser, H., and Levelt, P. F.: TROPOMI on the ESA Sentinel-5 Precursor: A GMES mission for global observations of the atmospheric composition for climate, air quality and ozone layer applications, *Rem. Sens. Environ.*, 120, 70–83, <https://doi.org/10.1016/j.rse.2011.09.027>, 2012.
- 30 Veefkind, J. P., de Haan, J. F., Sneep, M., and Levelt, P. F.: Improvements to the OMI O₂–O₂ operational cloud algorithm and comparisons with ground-based radar–lidar observations, *Atmos. Meas. Tech.*, 9, 6035–6049, <https://doi.org/10.5194/amt-9-6035-2016>, 2016.
- Vial, J., Dufresne, J.-L., and Bony, S.: On the interpretation of inter-model spread in CMIP5 climate sensitivity estimates, *Clim. Dyn.*, 41, 3339–3362, <https://doi.org/10.1007/s00382-013-1725-9>, 2013.
- Wang, P. and Stammes, P.: Evaluation of SCIAMACHY Oxygen A band cloud heights using Cloudnet measurements, *Atmos. Meas. Tech.*, 7, 1331–1350, <https://doi.org/10.5194/amt-7-1331-2014>, 2014.
- 35 Wang, P., Stammes, P., van der A, R., Pinardi, G., and van Roozendaal, M.: FRESCO+: an improved O₂ A-band cloud retrieval algorithm for tropospheric trace gas retrievals, *Atmos. Chem. Phys.*, 8, 21, 6565–6576, <https://doi.org/10.5194/acp-8-6565-2008>, 2008.

Xu, X., Wang, J., , Wang, Y., Zeng, J., Torres, O., Yang, Y., Marshak, A., Reid, J., and Miller, S.: Passive remote sensing of altitude and optical depth of dust plumes using the oxygen A and B bands: First results from EPIC/DSCOVR at Lagrange-1 point, *Geophys. Res. Lett.*, 44, 7544–7554, <https://doi.org/10.1002/2017GL073939>, 2017.

5 Yang, Y., Marshak, A., Mao, J., Lyapustin, A., and Herman, J.: A method of retrieving cloud top height and cloud geometrical thickness with oxygen A and B bands for the Deep Space Climate Observatory (DSCOVR) mission: Radiative transfer simulations, *J. Quant. Spectrosc. Radiat. Transfer*, 122, 141 – 149, <https://doi.org/10.1016/j.jqsrt.2012.09.017>, 2013.