



1 **Revisiting particle sizing using grayscale optical array probes:**  
2 **evaluation using laboratory experiments and synthetic data**

3

4 **Sebastian J. O’Shea<sup>1</sup>, Jonathan Crosier<sup>1,2</sup>, James Dorsey<sup>1,2</sup>, Waldemar Schledewitz<sup>1</sup>, Ian**  
5 **Crawford<sup>1</sup>, Stephan Borrmann<sup>3,4</sup> Richard Cotton<sup>5</sup> and Aaron Bansemer<sup>6</sup>**

6

7 [1]{School of Earth and Environmental Sciences, University of Manchester, Manchester, UK}

8 [2]{National Centre for Atmospheric Science, University of Manchester, Manchester, UK}

9 [3]{Particle Chemistry Department, Max Planck Institute for Chemistry, Mainz, Germany}

10 [4]{Institute for Atmospheric Physics, Johannes Gutenberg University, Mainz, Germany}

11 [5]{Met Office, Exeter, UK}

12 [6]{National Center for Atmospheric Research, Boulder CO, USA}

13 Correspondence to: S. J. O’Shea ([sebastian.oshea@manchester.ac.uk](mailto:sebastian.oshea@manchester.ac.uk))

14

15 **Abstract**

16 In-situ observations from research aircraft and instrumented ground sites are important  
17 contributions to developing our collective understanding of clouds, and are used to inform and  
18 validate numerical weather and climate models. Unfortunately, biases in these datasets may be  
19 present, which can limit their value. In this paper, we discuss artefacts which may bias data  
20 from a widely used family of instrumentation in the field of cloud physics, Optical Array Probes  
21 (OAPs). Using laboratory and synthetic datasets, we demonstrate how grayscale analysis can  
22 be used to filter data, constraining the sample volume of the OAP, and improving data quality  
23 particularly at small sizes where OAP data are considered unreliable. We apply the new  
24 methodology to ambient data from two contrasting case studies: one warm cloud and one cirrus  
25 cloud. In both cases the new methodology reduces the concentration of small particles (< 60  
26  $\mu\text{m}$ ) by approximately an order of magnitude. This significantly improves agreement with a  
27 Mie scattering spectrometer for the liquid case and with a holographic imaging probe for the



1 cirrus case. Based on these results, we make specific recommendations to instrument  
2 manufacturers, instrument operators, and data processors about the optimal use of greyscale  
3 OAP's. We also raise the issue of bias in OAP's which have no greyscale capability.

4

## 5 **1 Introduction**

6 Optical array probes (OAPs) are widely used to provide in situ measurements of cloud particle  
7 size, habit and concentration (Wendisch and Brenguier, 2013). These measurements provide  
8 insights into key cloud microphysical processes such as ice nucleation, particle growth and  
9 precipitation (Field, 1999; Lawson et al., 2015). In situ measurements are an important means  
10 to validate remote sensing instrumentation, which are routinely used to initialise operational  
11 weather forecast models (Fox et al., 2018; Mace & Benson, 2017).

12 OAPs consist of a laser illuminating a linear array of photo-diodes. A particle crossing the laser  
13 beam is detected if the laser intensity at any of the elements of the array drops below a threshold  
14 value. A shadow image is constructed by appending consecutive slices from the detectors as  
15 the particle moves perpendicular to the laser beam.

16 Monoscale OAPs use a 50% decrease in signal intensity as their threshold for detection (Lawson  
17 et al., 2006), resulting in 1-bit binary images with pixels either in an active state, or an inactive  
18 state. Greyscale OAPs are also available, which detect particles at multiple intensity thresholds,  
19 resulting in 2-bit images with pixels having three different active states, and one in-active state.  
20 For example, a greyscale probe could be configured to record images with pixels off (inactive),  
21 or triggered at shadow intensity levels of 25%, 50% and 75%. We use the abbreviations  $A_{25-50}$ ,  
22  $A_{50-75}$  and  $A_{75-100}$  for the number of pixels associated with a decrease in detector signal of 25 –  
23 50%, 50 – 75% and 75 – 100%, respectively. Similarly we use the abbreviations  $D_{25}$ ,  $D_{50}$ , and  
24  $D_{75}$  for the diameter of images with decreases in detector signal greater than 25, 50 and 75%,  
25 respectively. Korolev et al. (1991) describes a hybrid mono-grey system for the closely related  
26 1-D type of probe; this used a similar array as an OAP to size particles using a 50% shadow  
27 intensity, but also had on-board signal processing to provide additional filtering and the  
28 requirement for at least 1 pixel in any measured image to have a shadow intensity  $>67\%$ . This  
29 resulted in the reduction of artefacts due to poorly imaged particles near the edges of the depth  
30 of field.



1 Particles which are imaged by an OAP are fully in focus at the object plane, with image quality  
2 deteriorating as the particle location moves away from this object plane. The distance from the  
3 object plane that a particle can be observed is known as the depth of field (DoF) and is used to  
4 determine the instruments sample volume (together with the air speed and effective optical  
5 array width). Previous studies have found that the depth of field (at the 50% intensity level)  
6 follows a relationship of the form (Knollenberg, 1970)

$$7 \quad DoF = \pm \frac{cD_0^2}{4\lambda}$$

8 Equation 1

9 where  $D_0$  is the particle diameter and  $\lambda$  is the laser wavelength.  $c$  is a dimensionless constant,  
10 typically between 3 and 8 (Lawson et al., 2006; Gurganus & Lawson, 2018).

11 The size of the measured image depends on the particle's distance from the object plane.  
12 However, this dependence has been shown to be non-monotonic (Joe and List, 1987). Korolev  
13 et al. (1991) show that OAP images of transparent spheres (e.g. liquid drops) can be accurately  
14 approximated by the Fresnel diffraction from an opaque disc. The ratio of the detected image  
15 diameter to the actual physical diameter  $D_0$  is purely a function of the normalised, dimensionless  
16 distance from the object plane  $Z_d$ ,

$$17 \quad Z_d = \frac{4\lambda Z}{D_0^2}$$

18 Equation 2

19 where  $Z$  is the distance from the object plane. The spatial intensity distributions from  
20 transparent spheres are independent of particle size. A distinct feature of these distributions is  
21 a bright spot at the centre of the image known as the Poisson spot. Korolev et al. (2007, hereafter  
22 K07) describes a method for determining a spherical particle's distance from the object plane  
23 and size using the size of the Poisson spot.

24 Joe and List (1986) suggested significantly reducing the depth of field so that the image size  
25 could be assumed to be equal to the particle size. Particles outside the new depth of field were  
26 identified using the ratio  $A_{75-100} / (A_{25-50} + A_{50-75} + A_{75-100})$ . The disadvantage of reducing the  
27 depth of field in this way is that it can lead to poor sampling statistics. Similarly, by requiring  
28 at least one pixel to have a >67% decrease in detector signal Korolev et al. (1991) removes the  
29 most severely miss sized particles. Reuter & Bakan (1998, hereafter RB98) suggested an



1 alternative approach by assuming a linear relationship between image size and the greyscale  
2 ratio  $A_{25-50} / (A_{25-50} + A_{50-75} + A_{75-100})$ , which they then use to determine the particle size. This  
3 relationship was determined using laboratory experiments with a rotating disk with printed  
4 circular spots.

5 This paper describes tests on a Droplet Measurement Technologies Inc. (DMT) greyscale cloud  
6 imaging probe (CIP-15) using a droplet generation system. Liquid drops were injected into the  
7 probe at measured distances from the object plane to examine how this impacts the ability of  
8 the probe to accurately size particles. Results from these experiments are compared to synthetic  
9 images calculated assuming Fresnel diffraction (Korolev et al. 1991). Section 3.1 evaluates the  
10 efficacy of the K07 and RB98 size correction algorithms. Section 3.2 uses greyscale intensity  
11 ratios to determine a particle's distance from the object plane near the edge of the probe's depth  
12 of field. This allows significantly fragmented images to be removed and a revised depth of field  
13 to be used to determine particle concentrations. Section 3.4 examines how these results impact  
14 field measurements of particle size distributions using two research flights: one in a warm liquid  
15 cloud and one in cirrus.

16

## 17 **2 Methods**

### 18 **2.1 Cloud imaging probe (CIP-15)**

19 The CIP-15 is a commercially available greyscale OAP (DMT Inc., USA; Baumgardner et al.,  
20 2001). It has a 64 element photo-diode array with an effective pixel size of 15  $\mu\text{m}$ , giving the  
21 probe a nominal size range of 15 to 960  $\mu\text{m}$ . Images are recorded at three greyscale thresholds,  
22 which can be varied in the probe's data acquisition software. For the drop generator experiments  
23 these thresholds were set to 25, 50 and 75%; and 25, 50 and 67%. In Sect. 3.4.1 the thresholds  
24 were 40, 50 and 70%; and in Sect. 3.4.2 they were 25, 50 and 75%. The probe is fitted with  
25 anti-shatter tips to minimise ice shattering on the leading edge of the probe during field  
26 measurements. The measurements presented in this paper are from two CIP-15 systems. The  
27 major difference between the two probes is that one CIP-15 has an arm separation of 7 cm, and  
28 the other has an arm separation of 4 cm. The laboratory experiments and warm cloud results in  
29 Sect. 3.4.1 used the CIP with 7 cm arm separation, whereas the cirrus results in Sect. 3.4.2 use  
30 the CIP with 4 cm arm separation.

31



## 1    **2.2    Supporting measurements**

2    Section 3.4 compares measurements from the CIP-15 on board the FAAM Bae-146 research  
3    aircraft to those from a DMT Inc. Cloud Droplet Probe (CDP) and a holographic imaging probe.  
4    The CDP sizes particles in the range 3 to 50  $\mu\text{m}$  using the scattered light intensity from a diode  
5    laser and assuming Mie scattering theory (Lance et al., 2010). The probe was calibrated during  
6    the campaign using glass calibration beads.

7    HALOHolo is a holographic imaging probe from the Institute for Atmospheric Physics at the  
8    University of Mainz and Max Planck Institute for Chemistry Mainz. It has a  $6576 \times 4384$  pixel  
9    CCD detector with an effective pixel size of 2.95  $\mu\text{m}$ . This equates to a sample volume of  
10    approximately  $19 \times 13 \times 155$  mm ( $\sim 38$   $\text{cm}^3$ ). At 6 frames per second and an average airspeed of  
11    about  $100$   $\text{ms}^{-1}$ , this equates to a volume sample rate of  $\sim 230$   $\text{cm}^3$   $\text{s}^{-1}$ . Particles between 6  $\mu\text{m}$   
12    (2 pixels) and 1 cm (half the detector width) are resolvable in the hologram reconstructions.  
13    However, the detection of small particles is limited by noise in the background image. Therefore  
14    a minimum size threshold of 35  $\mu\text{m}$  is applied, above which it is estimated that the probe's  
15    detection rate is greater than 90% (Schlenczek, 2017). Shattered particles were minimised by  
16    removing all particles with inter particle distances less than 10 mm (Fugal & Shaw, 2009;  
17    O'Shea et al., 2016)

18

## 19    **2.3    Drop generator**

20    A monodisperse stream of droplets was generated using a commercially available droplet  
21    generator. The generator is similar to that described by Lance et al. (2010) and uses  
22    piezoelectrically actuated printheads. Three MicroFab, Inc (USA) printheads with 60, 90 and  
23    120  $\mu\text{m}$  orifices were used during these experiments. Each printhead was in turn vertically  
24    mounted on two perpendicularly positioned translation stages (MTS50/M-Z8 - 50 mm,  
25    Thorlabs), each have 50 mm travel range and a 0.05  $\mu\text{m}$  minimum increment. The printheads  
26    were connected to a fluid reservoir and the fluid pressure was adjusted using a pneumatic  
27    pressure controller so that the meniscus was at the end of the printhead. The printheads were  
28    actuated using a JetDrive III electronics module (MicroFab, Inc) to give a droplet production  
29    rate of 50 Hz.

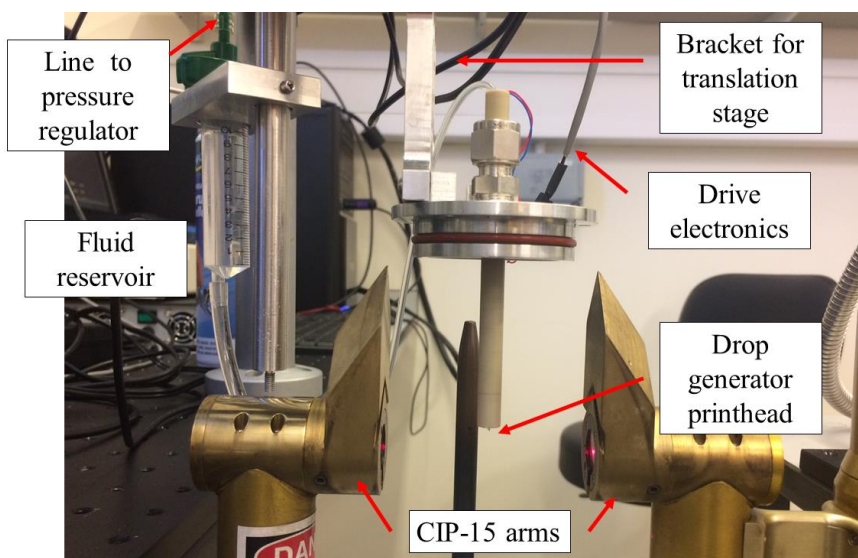
30    The generation of a stable stream of monodisperse droplets depends on several factors,  
31    including the drive electronics and the physical properties of the fluid used. These factors



1 together with size of the printhead orifice control the size of the generated droplets. Previous  
2 work has shown that more stable outputs can be achieved using mixtures of water and ethylene  
3 glycol compared to using pure water (Jang et al., 2009; Liu, 2016). For these experiments, with  
4 the 60  $\mu\text{m}$  printhead a 50% water and 50% ethylene glycol solution was used, while 100%  
5 ethylene glycol was used with the 90  $\mu\text{m}$  and 120  $\mu\text{m}$  printheads.

6 Figure 1 shows the setup for experiments that were performed on the CIP-15. The CIP-15 was  
7 vertically mounted below the drop generator. For each printhead the drop generator's position  
8 was stepped in 0.5 mm increments between the CIP-15's arms, dwelling for 3s after each  
9 movement.

10



11

12 *Figure 1. Photograph of the droplet generator and CIP-15.*

13

14 To test the stability of the droplet generator separate experiments were performed where the  
15 droplets were monitored using a high speed camera (FASTCAM Mini AX, Photron, UK) with  
16 a zoom lens (Navitar, USA). The pixel size of the image for a given magnification was  
17 calibrated using a stage graticule (50 x 2  $\mu\text{m}$ , Graticules Ltd, London, UK). The droplets were  
18 monitored for a 1 hour period at a droplet production rate of both 10 droplets  $\text{s}^{-1}$  and 250 droplets  
19  $\text{s}^{-1}$ . The interquartile range of the drop diameters, as measured by the camera, were 2  $\mu\text{m}$  and 3  
20  $\mu\text{m}$ , respectively.



1

## 2 **2.4 Synthetic data**

3 Modelled images were generated using optical, electronic, and diode thresholding simulations.  
4 The optical simulation considers only Fresnel diffraction by a round object, following the  
5 methods of Korolev et al. (1991) and K07. An electronic time delay simulation was also  
6 performed (Baumgardner and Korolev, 1997), but the effects are negligible due to the fast  
7 response of the electronics ( $\tau = 51$  ns) and the slow speed of the simulation (air speed of 10  
8  $\text{m s}^{-1}$ ). The photodiodes are assumed to have a rectangular shape with a 5:4 aspect ratio and a  
9 20% spacing between them, as in Korolev (2007). Similar to the drop generator experiments  
10 particles were injected at known distances from the object plane and the probe's response was  
11 simulated. This was done for particles with diameters 50 to 150  $\mu\text{m}$  at intervals of 5  $\mu\text{m}$ . These  
12 were positioned at 1 mm intervals over the range -5 to +5 cm from the object plane. Images  
13 were simulated using four different combinations of greyscale thresholds: 25, 50 and 75%; 30,  
14 50 and 70%; 40, 50 and 70%; and 25, 50 and, 67%.

15

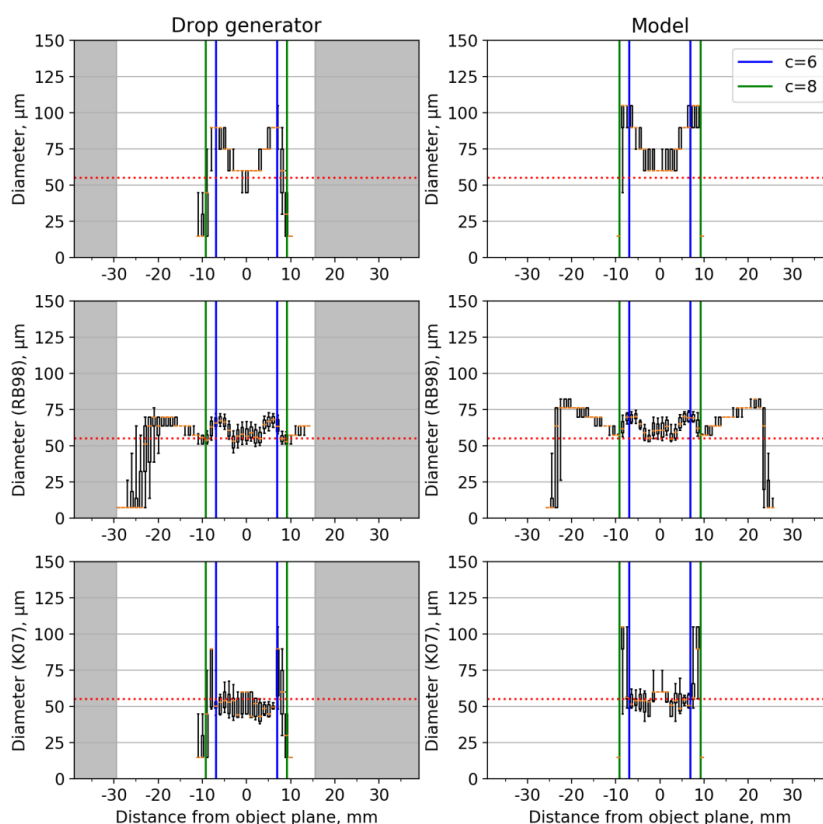
## 16 **3 Results**

### 17 **3.1 Particle size correction**

18 Figures 2-4 show the image diameter as a function of distance from the object plane. The  
19 diameters in these plots have been calculated along the axis of the optical array. Left panels  
20 show results from the laboratory experiments using the 60  $\mu\text{m}$ , 90  $\mu\text{m}$  and 120  $\mu\text{m}$  drop  
21 generator printheads. Example images from the 90  $\mu\text{m}$  printhead at three distances from the  
22 object plane are shown in Figure 5. Theoretically the size of the image at the centre of the object  
23 plane should be the closest approximation to the drop size. The median size of these are 60  $\mu\text{m}$ ,  
24 90  $\mu\text{m}$  and 90  $\mu\text{m}$  for the 60  $\mu\text{m}$ , 90  $\mu\text{m}$  and 120  $\mu\text{m}$  printheads, respectively. These  
25 measurements are subject to a 15  $\mu\text{m}$  uncertainty due to the pixel resolution of the CIP-15. The  
26 right panels show similar plots for the synthetic data for 55  $\mu\text{m}$ , 80  $\mu\text{m}$  and 90  $\mu\text{m}$  particles  
27 (dashed red lines). These sizes were chosen as they were the closest match to the droplet  
28 calibrations. The position of the depth of field as calculated using Eq.1 with a  $c$  value of 6 (blue)  
29 and 8 (green) are shown as vertical lines. If calculated correctly particles should not be visible  
30 outside the depth of field. If not removed, such particles would bias the measured



1 concentrations. Figures 2-4 show that a  $c$  value of 8 effectively bounds the region where  
2 particles are visible using a 50% intensity threshold. The drop velocity will have an impact on  
3 this due to the probe's electronic time delay (Baumgardner and Korolev, 1997). These tests  
4 were performed with relatively slow droplet velocities ( $< 10 \text{ m s}^{-1}$ ), especially when compared  
5 to aircraft measurements (approximately  $100 \text{ m s}^{-1}$ ). However, this effect is minimised by the  
6 fast time response of modern probes such as the CIP-15 ( $\tau = 51 \text{ ns}$ ).



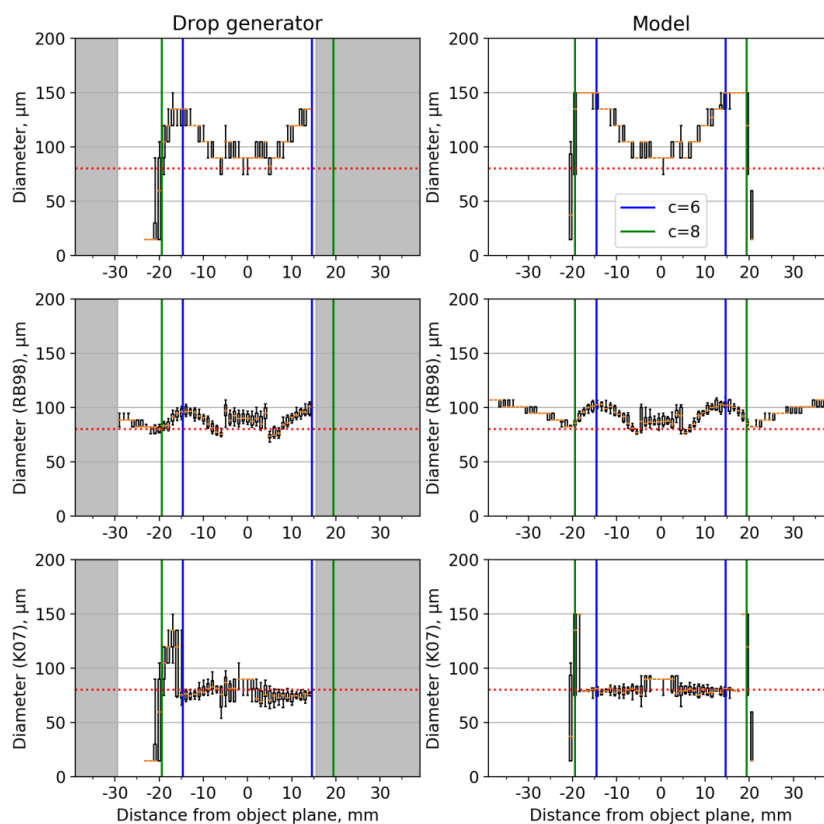
7  
8 *Figure 2. Box and whisker plots showing image diameter as a function of distance from the*  
9 *object plane. Left panels show results from the laboratory experiments using a 60 μm printhead.*  
10 *The grey shaded regions were not sampled using the drop generator. Right panels show the*  
11 *model image diameter from a 55 μm particle. Top panels show the image diameter using a 50%*





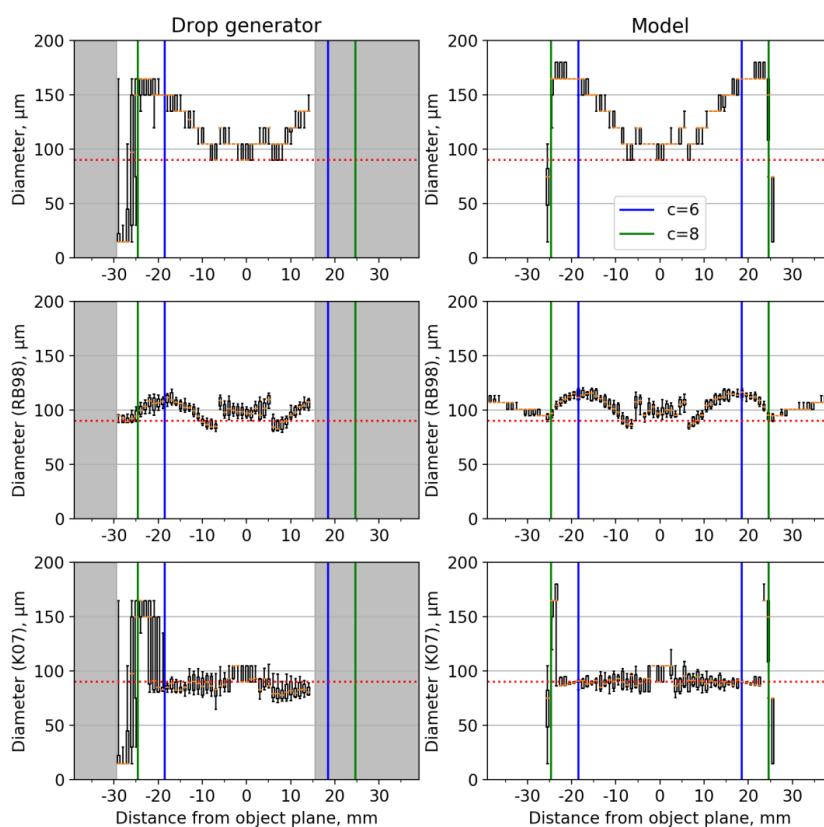
1 decrease in intensity threshold ( $D_{50}$ ). Middle panels show the diameter after the Reuter &  
2 Bakan (1998) size correction has been applied. Bottom panels show the diameter corrected  
3 using the Korolev et al. (2007) algorithm. Dashed red line shows the estimated droplet  
4 diameter. The position of the depth of field calculated using Eq. 1 with a  $c$  value of 6 (blue) and  
5 8 (green) are shown as vertical lines.

6



7

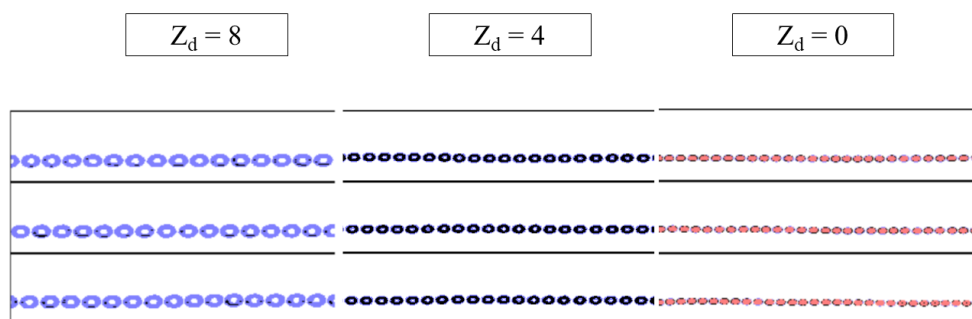
8 Figure 3. Same as Fig. 2 but left panels show results from the laboratory experiments using a  
9 90  $\mu\text{m}$  printhead and right panels show the modelled image diameter from an 80  $\mu\text{m}$  particle.



1

2 *Figure 4. Same as Fig. 2 but left panels show results from the laboratory experiments using a*  
3 *120 µm printhead and right panels show the modelled image diameter from a 90 µm particle.*

4



1

2 *Figure 5. Example images of droplets from the droplet generator using a 90  $\mu\text{m}$  printhead at*  
 3  *$Z_d = 8$  (left),  $Z_d = 4$  (middle) and  $Z_d = 0$  (right). Decreases in detector intensity of 25 to 50%,*  
 4 *50 to 75% and >75% are shown as light blue, dark blue and orange pixels, respectively.*

5

6 The middle panels show the image diameter as function of distance from the object plane once  
 7 the RB98 size correction algorithm has been applied. This algorithm assumes a linear  
 8 relationship between  $D_{25}$  and the greyscale ratio  $A_{25} / (A_{25} + A_{50} + A_{75})$ . In reality this  
 9 relationship is not completely linear, as a result the corrected diameter is not independent of  
 10 position. Additionally there is a bias in the corrected size when compared to the particle model.  
 11 To quantify this for the synthetic data we calculate the median of each position bin, Table 1  
 12 shows statistics for  $Z_d < 6$ . It is clear that the RB98 algorithm has a bias of the order 10  $\mu\text{m}$ . A  
 13 similar bias is seen in RB98 when compared to K07 if both algorithms are applied to results  
 14 from the drop generator (Table 1).

15

Drop Generator				Model				
Printhead , $\mu\text{m}$	$D_{50}$ ( $Z=0$ ) )	Image diameter , $\mu\text{m}$	Reuter & Bakan	Korole v et al. (2007),	Particle diameter , $\mu\text{m}$	Image diameter , $\mu\text{m}$	Reuter & Bakan	Korole v et al. (2007),



			(1997) , $\mu\text{m}$				(1997) , $\mu\text{m}$	
60	60	68 (15)	60 (10)	51 (10)	55	75 (26)	62 (9)	54 (7)
90	90	105 (30)	90 (11)	76 (8)	80	105 (32)	92 (13)	80 (4)
120	90	105 (23)	99 (7)	86 (7)	90	120 (43)	104 (15)	91 (5)

1 *Table 1. Median (inter-quartile range) image diameter for  $Z_d < 6$  from the drop generator*  
 2 *experiments and the model images.*

3

4 The lower panels in Figures 2-4 show the image diameter after the K07 algorithm has been  
 5 applied. Across much of the depth of field this algorithm removes the image diameter's position  
 6 dependence. For the synthetic data the median diameter across the depth of field is now within  
 7  $1 \mu\text{m}$  of the true particle diameter and the inter-quartile range is reduced (Table 1). At the edge  
 8 of the depth of field the Poisson spots become sufficiently large that the outer ring fragments at  
 9 the 50% threshold (Fig. 5, left panel). Once this happens the K07 algorithm is not able to correct  
 10 the size of these severely misshapen images. As shown in Figures 2-4 these fragmented images  
 11 have large variability in their size and can be either much larger or smaller than the true particle  
 12 size.

13

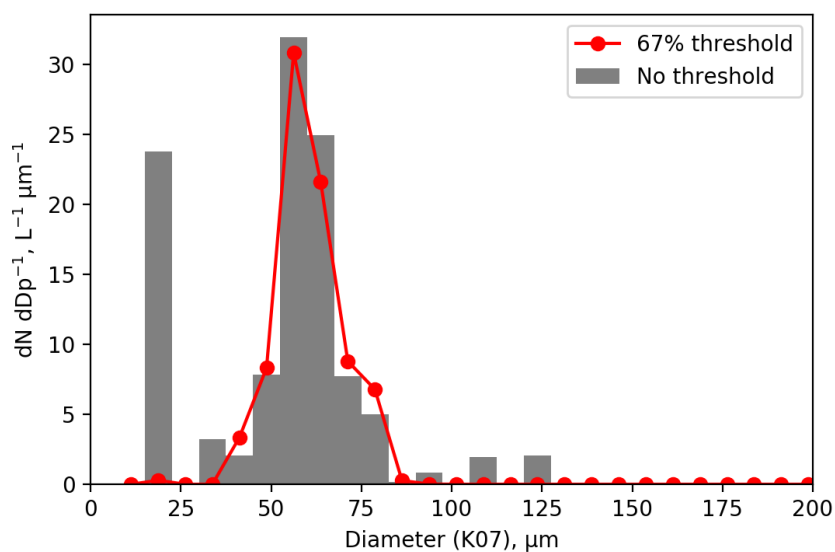
### 14 **3.2 Identifying fragmented images**

15 The K07 algorithm effectively corrects the diameter of imaged spherical particles across much  
 16 of the depth of field for binary images at the 50% threshold. However for  $Z_d$  greater than  
 17 approximately 7 the images are too fragmented and the correction no longer works effectively.  
 18 These fragmented images need to be removed from further analysis, otherwise they will bias  
 19 the measured size distributions.

20 The 1-D probes described Korolev et al. (1991) have an element of greyscale filtering. They do  
 21 not record particle images, rather they just measure the diameter of particles using a 50%  
 22 threshold. Only particles that have at least one detector with a  $>67\%$  drop in detector intensity



1 are recorded. To test the efficacy of this approach to remove fragmented particles drop  
2 generator scans were performed with the CIP-15 thresholds set to 25, 50 and 67%. Figure 6  
3 shows a size distribution of the K07 corrected diameter for a scan using the 60  $\mu\text{m}$  printhead.  
4 The grey bars show data from all droplets, while the red markers show only images with at least  
5 one pixel above the 67% threshold. By performing this filtering the fragmented images are  
6 minimised and the depth of field is constrained to  $Z_d < 4.8$ .  
7



8

9 *Figure 6. Size distribution of the K07 corrected diameter for a drop generator scan using the*  
10 *60  $\mu\text{m}$  printhead. The grey bars show data from all droplets, while the red markers show only*  
11 *images with at least one pixel with a >67% decrease in detector intensity. This reduces the*  
12 *depth of field to  $Z_d < 4.8$ .*

13

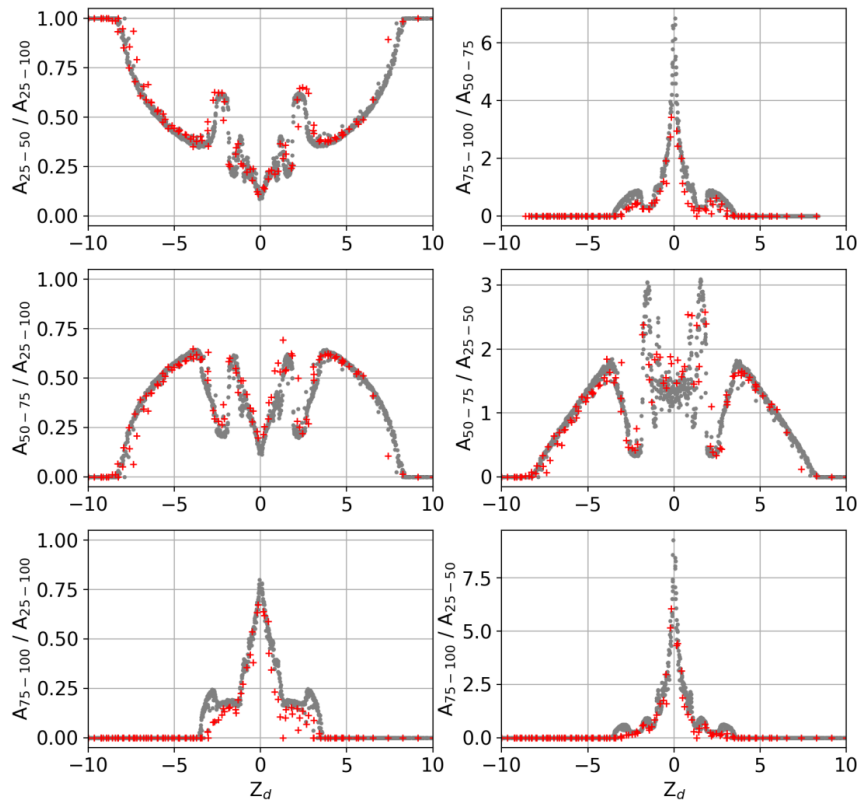
14 Ideally greyscale information could be used to uniquely determine a particle's  $Z_d$ , which could  
15 either be used to correct the image size or exclude fragmented images from further analysis.  
16 Figure 7 shows various combinations of greyscale ratios as a function of  $Z_d$ . Results from the  
17 model for the particle sizes 50 to 150  $\mu\text{m}$  are shown in grey, while results from the drop  
18 generator for the three printhead sizes (60, 90 and 120  $\mu\text{m}$ ) are shown in red. None of these



1 ratios are monotonic and most exhibit very complex behaviour. As a consequence they can't  
2 easily be used to determine a particle's position across the whole depth of field. However,  
3 within certain regions some of the ratios are monotonic.

4 The ratio  $A_{50-75} / A_{25-50}$  (middle right panel) is near linear for the approximate range  $3.5 < |Z_d|$   
5  $< 8.5$ . This is an important region since it is where the images begin to fragment and the true  
6 particle size can no longer be accurately retrieved using the K07 algorithm. Before this ratio  
7 can be used to determine a particle's position we need to check that  $|Z_d|$  is within the linear  $A_{50-}$   
8  $75 / A_{25-50}$  region. If an image has  $A_{75-100} > 0$  then  $|Z_d|$  can be limited to less than approximately  
9 3.5. Similarly, if  $A_{50-75}$  is equal to zero then  $|Z_d|$  will be greater than approximately 8.5 and  
10 likely too fragmented for accurate sizing.

11



1

2 *Figure 7. The ratios of the number of pixels between different thresholds from the drop*  
 3 *generator experiments (red) and model simulations of particles in the size range 50 to 150  $\mu\text{m}$*   
 4 *(grey) as a function of normalised distance from the object plane ( $Z_d$ ).*

5

6 Figure 8 shows results from the model (positive  $Z_d$ ) that meet the criteria  $A_{75-100} = 0$  and  $A_{50-75}$   
 7  $> 0$  (blue markers). The following equation can be fit to the data with an  $R^2$  of 0.98,

8

9

$$|Z_d| = \frac{\frac{A_{50-75}}{A_{25-50}} - 3.2048}{-0.3772}$$



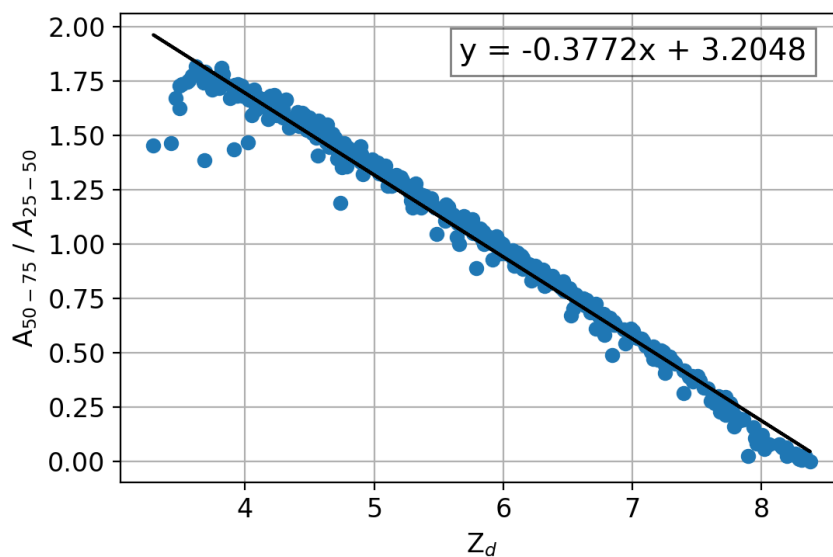
1

Equation 3

2 This equation allows the particle position to be retrieved over the approximate range  $3.5 < |Z_d|$   
3  $< 8.5$ . It should be noted that the uncertainty is greater for particles in the region  $3.5 < |Z_d| < 5$   
4 due to the larger number of outliers.

5 Figure 9 shows size distributions of the K07 corrected diameter for drop generator scans using  
6 the  $60 \mu\text{m}$  (top panel),  $90 \mu\text{m}$  (middle panel) and  $120 \mu\text{m}$  (bottom panel) printheads. The grey  
7 bars show data from all droplets, while the coloured lines show size distributions that have been  
8 filtered using different  $Z_d$  thresholds. The  $Z_d$  of each droplet was determined using Eq. 3. By  
9 applying a  $Z_d$  threshold both the very large and very small outliers are removed from the size  
10 distribution.

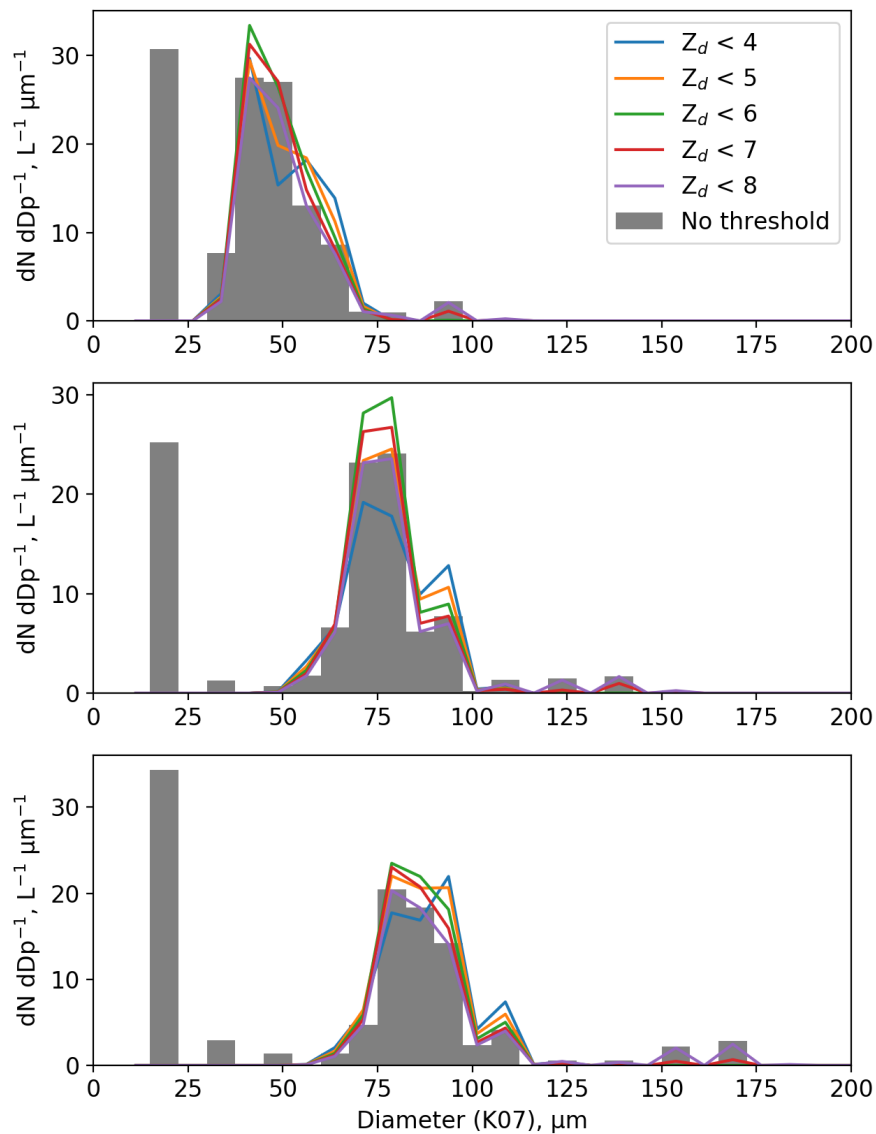
11



12

13 *Figure 8. The ratio of the number of pixels between grayscale thresholds 50 – 75% and 25 –*  
14 *50% (positive  $Z_d$ ) that meet the criteria  $A_{75-100} = 0$  and  $A_{50-75} > 0$  (blue markers). This data is*  
15 *from model simulations of particles in the size range 50 to 150  $\mu\text{m}$ .*







1 *Figure 9. Size distributions of the K07 corrected diameter for drop generator scans using the*  
 2 *60 μm (top panel), 90 μm (middle panel) and 120 μm (bottom panel) printheads. The grey bars*  
 3 *show data from all droplets, while the coloured lines show size distributions that have been*  
 4 *filtered using different  $Z_d$  thresholds. The  $Z_d$  of each droplet was determined using Eq. 3.*

5

6 Similar relationships can be derived using different greyscale thresholds. We tested several  
 7 different combinations of thresholds that could be used: first 40, 50 and 70%; second 30, 50  
 8 and 70%; and finally 25, 50 and 67%. A similar procedure was employed of first removing  
 9 particles with no pixels above the highest threshold and ones without at least one pixel above  
 10 the middle threshold. For each particle the number of pixels with greyscale intensity between  
 11 the middle and upper threshold ( $A_{Mid}$ ) was divided by the number of pixels between the lower  
 12 and middle thresholds ( $A_{Low}$ ). A linear equation of the following form was then fit to this ratio  
 13 versus  $Z_d$ ,

$$14 \quad |Z_d| = \frac{A_{Mid}}{A_{Low}} + i$$

15

Equation 4

16 Table 2 shows the fit coefficients for the four different combinations of greyscale thresholds  
 17 from the model data. Also shown is the approximate  $Z_d$  range where each relationship is  
 18 applicable.

19

Greyscale thresholds			Coefficients		$R^2$	$Z_d$ range
High	Mid	Low	i	j		
75	50	25	-3.2048	-0.3772	0.98	3.5 to 8.5
70	50	40	-7.8282	-0.9507	0.99	4.1 to 8.2
70	50	30	-4.0856	-0.4885	0.99	4.1 to 8.4
67	50	25	-3.3619	-0.4009	0.99	4.8 to 8.4

20 *Table 2. Fit coefficients  $i$  and  $j$  for Eq. 4 for different combinations of greyscale thresholds.*

21



### 1 3.3 Sample volume

2 The previous section described how  $Z_d$  can be determined for images as they begin to fragment.  
3 This allows a threshold  $Z_d$  to be employed to remove these images from further analysis. To  
4 correctly determine the particle concentration the sample volume needs to be adjusted to take  
5 account of the  $Z_d$  threshold. The revised depth of field is calculated by setting  $c$  in Eq. 1 equal  
6 to the chosen  $Z_d$  threshold. For  $Z_d$  to be correctly calculated using Eq. 4 and the K07 size  
7 correction to be applicable the entire particle needs to be imaged. Images that have pixels  
8 greater than the low greyscale threshold in contact with the edge of the optical array should not  
9 be used to calculate the concentration. The probe sample volume (SVol) for a given  $D_0$  can then  
10 be calculated using

11

$$12 \quad SVol = TAS \cdot \int_{-DoF}^{+DoF} (NR - D_{Low}(Z)) dZ$$

13

Equation 5

14 where  $D_{Low}$  is the image diameter using all pixels greater the low greyscale threshold, TAS is  
15 the true air speed, N is the number of array elements and R is the resolution of the probe. This  
16 equation has been modified from Korolev et al. (1991) so that it uses  $D_{Low}$  rather than  $D_{50}$ .  
17 Across much of the probe's size range the  $D_{25}$  sample area is less than 10% smaller than the  $D_{50}$   
18 sample area, however this increases for larger particles. The integration of the effective array  
19 width ( $NR - D_{Low}(Z)$ ) is performed over whichever is smaller out of the depth of field or the  
20 probe arm width.

21

### 22 3.4 Airborne measurements

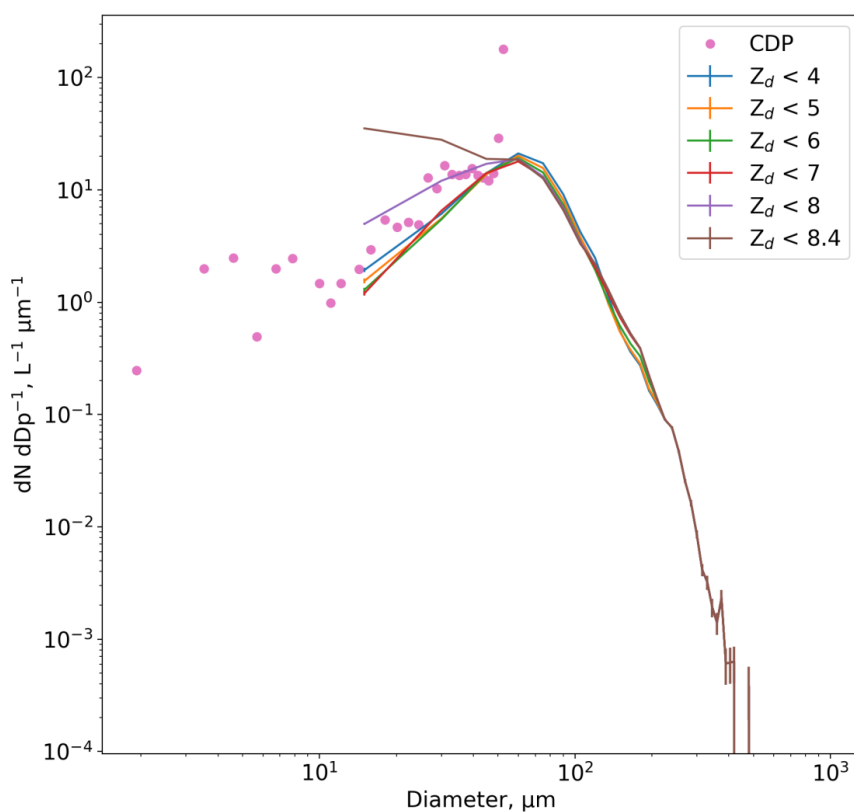
23 The following section applies the results from Sections 3.1 to 3.3 to field measurements from  
24 two research flights.

#### 25 3.4.1 Liquid cloud

26 As part of the CLOUDS and Aerosol Radiative Impacts and Forcing (CLARIFY) project the  
27 FAAM Bae-146 Research Aircraft performed sorties out of Ascension Island. On 5 September  
28 2017, Pockets of Open Cells were sampled. This flight was characterised by a clean marine



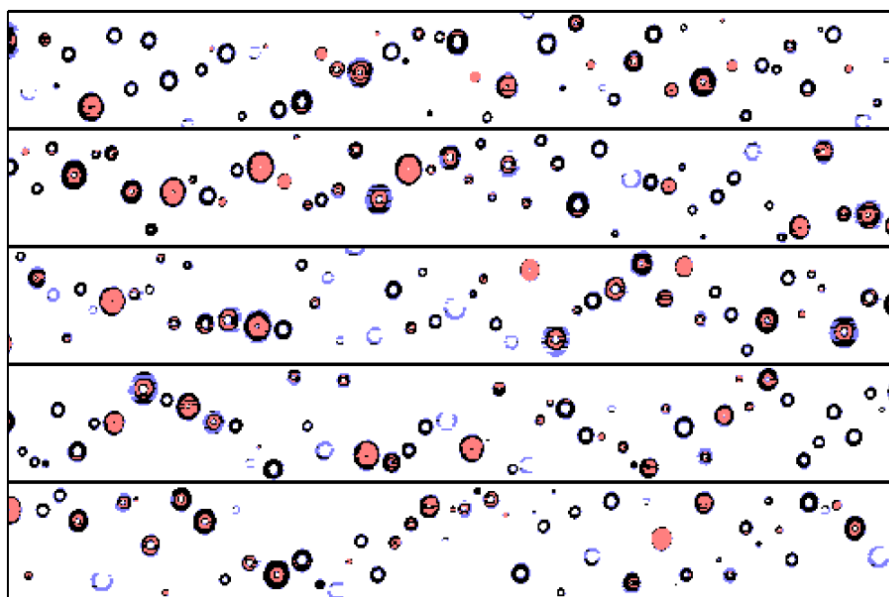
1 boundary layer and large cloud droplets/drizzle. Figure 10a shows size distributions from the  
2 CIP-15 that have been averaged over a straight and level run at 14°C (16:42:10 to 16:43:15  
3 GMT). Pink markers show the CDP size distribution averaged over the same period. The other  
4 coloured lines have been calculated using K07 using various  $Z_d$  thresholds. Figure 10b shows  
5 example CIP-15 images from this period. The particle diameters in this section have been  
6 calculated as the mean of the particle size along the axis of the optical array and the particle  
7 trajectory. Using a  $Z_d$  threshold less than 7 significantly reduces the concentration of drops  
8 smaller than 60  $\mu\text{m}$ , which is in better agreement with the CDP. The decrease is over an order  
9 of magnitude for the smallest CIP-15 bin.





1 *Figure 10a. Comparison in liquid cloud between the CDP and CIP-15 using different  $Z_d$*   
2 *thresholds. Figure 10b) shows example CIP-15 images from this period. Decreases in detector*  
3 *intensity of 25 to 50%, 50 to 75% and >75% are shown as light blue, dark blue and orange*  
4 *pixels, respectively.*

5



6

7 *Figure 10b.*

8

### 9 3.4.2 Cirrus

10 A number of studies have found a persistent small ice mode in their OAP measurements of  
11 cirrus clouds (Cotton et al., 2013; Jackson et al., 2015; O'Shea et al. 2016). O'Shea et al. (2016)  
12 hypothesised that in their measurements this was largely due to out of focus larger crystals. Due  
13 to the size dependence of the sample volume only a relatively small proportion of miss sized  
14 large particles are needed to cause a significant number concentration of small particles. The  
15 relationships between greyscale ratios and  $Z_d$  determined in this paper have been developed  
16 using spherical droplets. Similarly, K07 is strictly only applicable to spherical droplets.  
17 However, ice crystals can be a variety of complex shapes.

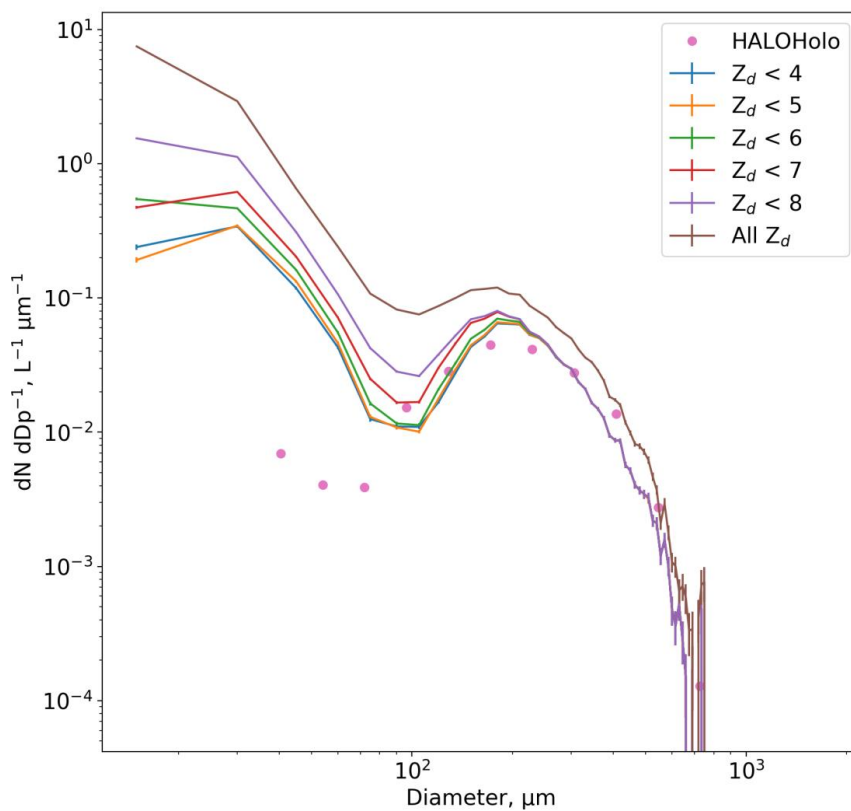


1 To examine whether the greyscale relationships in this paper can be applied to glaciated clouds  
2 we use measurements from the PICASSO project (Parameterizing Ice Clouds using Airborne  
3 obServationS and triple-frequency dOppler radar). On 7 February 2018, the FAAM BAe-146  
4 sampled cirrus over southern UK. Figure 11a shows size distributions for a straight and level  
5 run at  $-42\text{ }^{\circ}\text{C}$  (16:02:00 to 16:10:00 GMT). Crystals were predominantly rosettes and columns,  
6 with a smaller proportion of aggregates. Example CIP-15 images from this period are shown in  
7 Fig. 11b. Particles associated with inlet shattering were minimised by filtering particles with  
8 inter-arrival times less than  $1 \times 10^{-5}\text{ s}$  (Field et al., 2006). The particles in Fig. 10a have not been  
9 corrected using K07.

10 Similar to O'Shea et al. (2016), if no  $Z_d$  filter is applied the CIP-15 cirrus size distribution is  
11 bimodal with one mode at approximately  $200\text{ }\mu\text{m}$  and another at the smallest measured sizes.  
12 As a more restrictive  $Z_d$  threshold is applied the small particle mode (less than  $100\text{ }\mu\text{m}$ )  
13 decreases. Similar to the liquid case (Section 3.4.1) the concentration of small particles  
14 decreases by an order of magnitude for  $Z_d < 6$  compared to when no filtering is applied.  
15 However, this algorithm doesn't completely remove the small particle mode. There are a  
16 number of possible explanations for this: first, the mode may be real and due to ice nucleation.  
17 However, coincident holographic measurements do not show the small particle mode (pink  
18 markers, Fig. 11a), suggesting it is an artefact associated with the OAP measurement technique.  
19 Second, it may be due to shattering on the inlet of the probe. However as mentioned previously,  
20 shattering events should be associated with short inter-arrival times and a stringent inter-arrival  
21 threshold has been applied to this dataset. Third, noise in the CIP-15 images will degrade the  
22 accuracy of the  $Z_d$  retrieval. Finally, the non-spherical shape of ice crystals will mean that the  
23 greyscale relationships aren't directly applicable. Further work is needed to examine greyscale  
24  $Z_d$  relationships for specific particle habits and whether a spherical approximation is applicable.

25 At large sizes the sample volume decreases with size till it is zero for particles larger than  $960$   
26  $\mu\text{m}$ . The Poisson counting uncertainty in the size distribution is shown as error bars in Fig. 10.  
27 As shown in Fig. 11 the number of counts becomes small and the counting uncertainty increases  
28 significantly for particles larger than approximately  $700\text{ }\mu\text{m}$ .

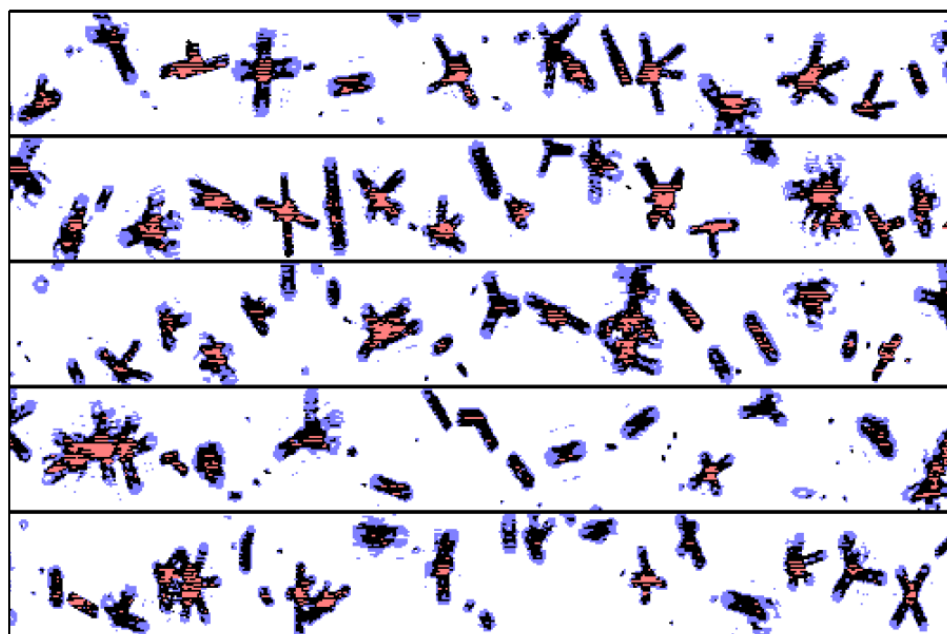
29



1

2 *Figure 11a. CIP-15 size distribution in cirrus using different  $Z_d$  thresholds. Figure 10b shows*  
3 *example CIP-15 images from this period. Decreases in detector intensity of 25 to 50%, 50 to*  
4 *75% and >75% are shown as light blue, dark blue and orange pixels, respectively.*

5



1

2 *Figure 11b.*

3

#### 4 **4 Conclusions**

5 This paper has described tests on a grayscale OAP using a droplet generator, results from which  
6 have been compared to synthetic data. Despite recent advances in holographic instruments for  
7 cloud microphysical measurements (Fugal & Shaw, 2009) work is still needed to better  
8 characterise the uncertainties associated with this technique. Additionally holographic probes  
9 require high performance computers to post-process the significant amounts of data they  
10 generate (e.g. HALOHolo generated several terabytes per 2-5 hour flight during PICASSO).  
11 This makes it challenging to routinely deploy such instruments. Therefore it is likely that OAPs  
12 will be still widely used for at least another 10 years. We make the following recommendations  
13 for their use:

14

- 15 • K07 should be used to correct the image size of spherical particles. This algorithm is  
16 found to perform better than RB98 across much of the depth of field ( $Z_d < 6$ ). However,





- 1 K07 is not able to correct the size of the severely fragmented images of particles near  
2 the edge of the probe's depth of field ( $Z_d > 6$ ).  
3
- 4 • Fragmented images from particles near the edge of the depth of field need to be removed  
5 to avoid significant bias to the derived particle size distributions. This is particularly a  
6 problem for diameters less than approximately  $100 \mu\text{m}$  due to the relatively small depth  
7 of field at these sizes.  
8
  - 9 • Greyscale information should be used to filter fragmented images and the probe's  
10 sample volume should be adjusted. The following four combinations of greyscale  
11 thresholds were tested: 25, 50 and 75%; 40, 50 and 70%; 30, 50 and 70%; and 25, 50  
12 and 67%. Using these thresholds and the relationships presented in this paper it is  
13 possible to determine a particle's position near the edge of the depth of field. This  
14 methodology was tested on measurements from two research flights. In both cases this  
15 reduced the concentration of small particles ( $< 60 \mu\text{m}$ ) by approximately an order of  
16 magnitude, significantly improving agreement with a Mie scattering spectrometer for  
17 the liquid case and with a holographic imaging probe for the cirrus case.  
18
  - 19 • The data from monoscale OAPs is unreliable below approximately  $100 \mu\text{m}$  due to  
20 fragmented larger particles. A small number of monoscale probes exist that reject  
21 particles that do not have at least one detector with a  $>67\%$  decrease in intensity. If this  
22 filtering is performed it would greatly minimise the impact of out-of-focus particles.  
23 However, this feature is not available on commonly used modern probes such as the  
24 2DS (SPEC Inc., Lawson et al., 2006).  
25
  - 26 • Reintroducing a 67% intensity rejection criteria on monoscale probes should be a high  
27 priority if possible. If this requires hardware modifications, it may be more appropriate  
28 to upgrade to full greyscale capability.  
29
  - 30 • Past datasets from greyscale OAPs should be re-examined. The filtering and sample  
31 volume adjustments presented in this paper should be applied.  
32



1 **Data availability**

2 The data presented here can be provided on request to the contact author.

3

4

5 **Acknowledgements**

6 The authors would like to thank Brian Derby, Andy Wallwork and Rachel Saunders for their  
7 assistance with the drop generator setup. We would like to thank Chris Westbrook and the  
8 PICASSO and CLARIFY teams for provision of the airborne data. Airborne data were obtained  
9 using the BAe-146-301 Atmospheric Research Aircraft (ARA) flown by Directflight Ltd and  
10 managed by the Facility for Airborne Atmospheric Measurements (FAAM), which is a joint  
11 entity of the Natural Environment Research Council (NERC) and the Met Office. The CIP-15s  
12 were provided by the National Centre for Atmospheric Science and FAAM. The National  
13 Centre for Atmospheric Science provided support for the droplet generator experiments. This  
14 work was supported by the NERC grants NE/P012426/1 and NE/L013584/1.

15



## 1 **References**

- 2 Baumgardner, D., Jonsson, H., Dawson, W., O'Connor, D., and Newton, R.: The cloud, aerosol  
3 and precipitation spectrometer: a new instrument for cloud investigations, *Atmos. Res.*, 59–60,  
4 251–264, [https://doi.org/10.1016/S0169-8095\(01\)00119-3](https://doi.org/10.1016/S0169-8095(01)00119-3), 2001.
- 5 Baumgardner, D. and Korolev, A.: Airspeed Corrections for Optical Array Probe Sample  
6 Volumes, *J. Atmos. Ocean. Tech.*, 14, 1224–1229, [https://doi.org/10.1175/1520-0426\(1997\)014<1224:ACFOAP>2.0.CO;2](https://doi.org/10.1175/1520-0426(1997)014<1224:ACFOAP>2.0.CO;2), 1997.
- 8 Cotton, R. J., Field, P. R., Ulanowski, Z., Kaye, P. H., Hirst, E., Greenaway, R. S., Crawford,  
9 I., Crosier, J., and Dorsey, J.: The effective density of small ice particles obtained from in situ  
10 aircraft observations of mid-latitude cirrus, *Q. J. Royal Meteor. Soc.*, 139, 1923–1934, 2013.
- 11 Field, P. R.: Aircraft Observations of Ice Crystal Evolution in an Altostratus Cloud, *Journal of  
12 the Atmospheric Sciences*, 56, 1925–1941, [https://doi.org/10.1175/1520-0469\(1999\)056<1925:AOOICE>2.0.CO;2](https://doi.org/10.1175/1520-0469(1999)056<1925:AOOICE>2.0.CO;2), 1999.
- 14 Field, P. R., Heymsfield, A. J., and Bansemer, A.: Shattering and Particle Interarrival Times  
15 Measured by Optical Array Probes in Ice Clouds, *J. Atmos. Oceanic Technol.*, 23, 1357–1371,  
16 2006.
- 17 Fox, S., Mendrok, J., Eriksson, P., Ekelund, R., O'Shea, S. J., Bower, K. N., Harlow, R. C., and  
18 Pickering, J. C.: Airborne validation of radiative transfer modelling of ice clouds at millimetre  
19 and sub-millimetre wavelengths, *Atmos. Meas. Tech. Discuss.*, <https://doi.org/10.5194/amt-2018-308>, in review, 2018.
- 21 Fugal, J. P. and Shaw, R. A.: Cloud particle size distributions measured with an airborne digital  
22 in-line holographic instrument, *Atmos. Meas. Tech.*, 2, 259–271, <https://doi.org/10.5194/amt-2-259-2009>, 2009.
- 24 Gurganus, C., and Lawson, P.: Laboratory and flight tests of 2D imaging probes: Toward a  
25 better understanding of instrument performance and the impact on archived data, *J. Atmos.  
26 Oceanic Technol.* 35, 7, 1533–1553, doi: 10.1175/JTECH-D-17-0202.1, 2018.
- 27 Jackson, R. C., McFarquhar, G. M., Fridlind, A. M., and Atlas, R.: The dependence of cirrus  
28 gamma size distributions expressed as volumes in  $N_0\text{-}\lambda\text{-}\mu$  phase space and bulk cloud properties  
29 on environmental conditions: Results from the Small Ice Particles in Cirrus Experiment  
30 (SPARTICUS), *J. Geophys. Res.-Atmos.*, 120, 10351–10377, doi:10.1002/2015JD023492,



- 1 2015.
- 2 Jang, D., Kim, D., Moon, J.: Influence of fluid physical properties on ink-jet printability.
- 3 *Langmuir*, 25, 2629–2635, 2009.
- 4 Joe, P., and List, R.: Testing and performance of two dimensional optical array spectrometers
- 5 with greyscale, *J. Atmos. Oceanic Technol.*, 4, 139–150, 1987.
- 6 Knollenberg, R. G.: The optical array: An alternative to scattering or extinction for airborne
- 7 particle size determination, *J. Appl. Meteorol.*, 9, 86–103, doi:10.1175/1520-
- 8 0450(1970)009<0086:TOAAAT>2.0.CO;2, 1970.
- 9 Korolev, A.: Reconstruction of the sizes of spherical particles from their shadow images. Part
- 10 I: Theoretical considerations, *J. Atmos. Ocean. Tech.*, 24, 376–389, 2007.
- 11 Korolev, A. V., Emery, E. F., Strapp, J. W., Cober, S. G., Isaac, G. A., Wasey, M. and Marcotte,
- 12 D.: Small ice particles in tropospheric clouds: Fact or artifact? Airborne icing instrumentation
- 13 evaluation experiment, *Bull. Am. Meteorol. Soc.*, 92(8), 967–973,
- 14 doi:10.1175/2010BAMS3141.1, 2011.
- 15 Korolev, A. V., Kuznetsov, S. V., Makarov, Yu. E., and Novikov, V. S.: Evaluation of
- 16 measurements of particle size and sample area from optical array probes. *J. Atmos. Oceanic*
- 17 *Technol.*, 8, 514–522, 1991.
- 18 Lance, S., Brock, C. A., Rogers, D. and Gordon, J. A.: Water droplet calibration of the Cloud
- 19 Droplet Probe (CDP) and in-flight performance in liquid, ice and mixed-phase clouds during
- 20 ARCPAC, *Atmos. Meas. Tech.*, 3(6), 1683–1706, doi:10.5194/amt-3-1683-2010, 2010.
- 21 Lawson, R. P., O’Connor, D., Zmarzly, P., Weaver, K., Baker, B., Mo, Q. and Jonsson, H.: The
- 22 2D-S (stereo) probe: Design and preliminary tests of a new airborne, high-speed, high-
- 23 resolution particle imaging probe, *J. Atmos. Ocean. Technol.*, 23(11), 1462–1477,
- 24 doi:10.1175/JTECH1927.1, 2006.
- 25 Lawson, R. P., Woods, S., and Morrison, H.: The Microphysics of Ice and Precipitation
- 26 Development in Tropical Cumulus Clouds, *J. Atmos. Sci.*, 72, 2429–2445, doi:10.1175/JAS-
- 27 D-14-0274.1, 2015.
- 28 Liu, Y.: Inkjet printed drops and three-dimensional ceramix structures, PhD Thesis, University
- 29 of Manchester, 2016.
- 30 Mace, G., G., and Benson, S.: Diagnosing cloud microphysical process information from



- 1 remote sensing measurements: A feasibility study using aircraft data. Part I: Tropical anvils
- 2 measured during TC4. *J. Appl. Meteor. Climatol.*, 56, 633–649, <https://doi.org/10.1175/JAMC->
- 3 D-16-0083.1, 2017.
- 4 O’Shea, S. J., Choullarton, T. W., Lloyd, G., Crosier, J., Bower, K. N., Gallagher, M., Abel, S.
- 5 J., Cotton, R. J., Brown, P. R. A., Fugal, J. P., Schlenczek, O., Borrmann, S., and Pickering, J.
- 6 C.: Airborne observations of the microphysical structure of two contrasting cirrus clouds, *J.*
- 7 *Geophys. Res.*, 121, 13, <https://doi.org/10.1002/2016JD025278>, 2016.
- 8 Reuter, A. and Bakan, S.: Improvements of cloud particle sizing with a 2-D-grey probe, *J.*
- 9 *Atmos. Ocean. Tech.*, 15, 1196–1203, 1998.
- 10 Schlenczek, O.: Airborne and ground-based holographic measurement of hydrometeors in
- 11 liquid-phase, mixed-phase and ice clouds, PhD Thesis, University of Mainz, 2017.
- 12 Wendisch, M. and Brenguier, J.-L. (Eds.): *Airborne Measurements for Environmental*
- 13 *Research: Methods and Instruments.*, vol. ISBN: 978-3-527-40996-9, Wiley-VCH Verlag
- 14 GmbH & Co. KGaA, Weinheim, Germany, 2013.