



Validation of MODIS 3 km Land Aerosol Optical Depth from NASA's EOS Terra and Aqua Missions

Pawan Gupta^{1,2}, Lorraine A. Remer³, Robert C. Levy², Shana Mattoo^{2,4}

[1] {Goddard Earth Sciences Technology And Research (GESTAR), Universities Space Research Association (USRA), Columbia, MD, USA}

[2] {NASA Goddard Space Flight Center, Greenbelt, MD 20771, USA}

[3] {JCET, University of Maryland – Baltimore County, Baltimore, MD 21228, USA}

[4] {Science Systems and Applications, Inc, Lanham, MD 20709, USA}

Correspondence to: Pawan Gupta (pawan.gupta@nasa.gov)

Abstract

The two MODerate Resolution Imaging Spectroradiometer (MODIS) sensors, aboard Earth Observing Satellites (EOS) Terra and Aqua, have been making aerosol observations for more than 15 years. From these observations, the MODIS dark target (DT) aerosol retrieval algorithm provides aerosol optical depth (AOD) products, globally over both land and ocean. In addition to the standard resolution product (10x10 km²), the MODIS collection 6 (C006) data release included a higher resolution (3x3 km²). Other than accommodations for the two different resolutions, the 10 km and 3 km DT algorithms are basically the same. In this study, we perform global validation of the higher resolution AOD over global land by comparing against AERONET measurements. The MODIS-AERONET collocated data sets consist of 161,410 high-confidence AOD pairs from 2000 to 2015 for MODIS Terra and 2003 to 2015 for MODIS-Aqua. We find that 62.5% and 68.4 % of AODs retrieved from MODIS-Terra and MODIS-Aqua, respectively, fall within previously published expected error bounds of $\pm (0.05 + 0.2 \cdot \text{AOD})$, with a high correlation ($R=0.87$). The scatter is not random, but exhibits a mean positive bias of ~ 0.06 for Terra and ~ 0.03 for Aqua. These biases for the 3 km product are approximately 0.03 larger than the biases found in similar validations of the 10 km product. The validation results for the 3 km product did not have a relationship to aerosol loading (i.e. true AOD), but did exhibit dependence on quality flags, region, viewing geometry, and aerosol



spatial variability. Time series of global MODIS-AERONET differences show that validation is not static, but has changed over the course of both sensors' lifetimes, with MODIS-Terra showing more change over time. The likely cause of the change of validation over time is sensor degradation, but changes in the distribution of AERONET stations and differences in the global aerosol system itself could be contributing to the temporal variability of validation.

1. Introduction

The Moderate Resolution Imaging Spectroradiometer (MODIS) sensors, onboard the Earth Observing System (EOS) Terra and Aqua satellites, have been providing observations of Earth and the atmosphere for almost two decades (Salomonson et al., 1989). These data have been used to create a long-term set of atmospheric aerosol properties including aerosol optical depth (AOD – a measure of aerosol loading in the total atmospheric column) (Kaufman et al., 1997; Levy et al., 2013). In particular we are addressing products of the Dark Target (DT) algorithms that provide aerosol retrievals over both ocean and dark vegetated land surfaces (Kaufman et al., 1997; Remer et al., 2005; Levy et al., 2007a; 2007b; 2013). The DT products were designed with climate applications in mind and have been used to address a wide variety of geophysical science questions including the role of aerosols in climate-relevant processes (Kaufman et al., 2002; Christopher et al., 2002; Yu et al., 2006), cloud/precipitation modifications (Koren et al., 2009; 2012; Yuan et al., 2011; Oreopoulos et al., 2016), and long-range transport of aerosols (Kaufman et al., 2005; Yu et al., 2012). Users have even applied the DT aerosol product to address needs for monitoring, evaluating and forecasting air quality (al Saadi et al., 2005; Gupta et al., 2009; Van Donkelaar et al., 2015).

The MODIS DT algorithm produces an aerosol product, over land and ocean, at a nominal $10 \times 10 \text{ km}^2$ spatial resolution. This spatial resolution permits much selectivity in choosing which MODIS-measured reflectance pixels at $0.5 \times 0.5 \text{ km}^2$ resolution to include in the retrieval, and generally produces smooth and accurate fields of AOD and other aerosol parameters (Remer et al., 2012). By allowing the algorithm to discard up to 90% of the available pixels and still produce a high quality aerosol product, the algorithm avoids marginal situations unfavorable for an aerosol retrieval such as cloud fringes, fragments and shadows, as well as land surfaces that



do not agree with algorithm assumptions (Remer et al., 2012). The 10 km product has undergone lengthy evaluation and validation, updated after each major algorithm modification (Ichoku et al., 2002; Chu et al., 2002; Remer et al., 2002, 2005; Russell et al., 2007; Levy et al., 2005; 2010). Some of this evaluation was global in nature, while some local to a particular field experiment, but all concerned the 10 km MODIS DT aerosol product.

For climate studies, the initial intention of the algorithm, 10 km spatial resolution was sufficient to characterize global and regional aerosol loading. However, as the community expanded the use of MODIS AOD to a wide variety of purposes, need arose for a finer resolution product, and a nominal 3 x 3 km² resolution product was introduced as part of MODIS Collection 6 (Levy et al., 2013; Remer et al., 2013). The product is termed MYD04_3K for 3 km resolution aerosol parameters derived from the MODIS-Aqua sensor and MOD04_3K for those derived from MODIS-Terra. These products are produced operationally, over land and ocean, and the entire data records of Terra and Aqua have been reprocessed, creating a data record of almost two decades.

Before becoming operational, Remer et al., (2013) tested the algorithm by comparing six months of global 3 km retrievals from MODIS-Aqua against available ground truth, while other independent studies (Munchak et al., 2013; Nichol and Bilal, 2016; He et al., 2017 and others) have done subsequent evaluation of the product regionally and locally. These limited comparisons suggested that the new AOD product would be sufficiently accurate to provide useful information and new perspective to the aerosol community, but might introduce additional noise and/or bias that the original coarser resolution product successfully avoided. Now that the multi-decadal 3 km product is operational and available publicly, it is time to perform a comprehensive evaluation of this finer resolution MODIS DT aerosol product. We present here the results of an analysis of a comparison of the global long-term MODIS 3 km product with collocated AEROSOL RObotic NETwork (AERONET) (Holben et al., 1998) observations.

2. The MODIS dark target 3 km aerosol retrieval over land



1 The MODIS DT algorithm and products are described in detail in Levy et al., (2013) and also in
 2 the MODIS DT on-line Algorithm Theoretical Basis Document (ATBD, 2017). In summary, to
 3 retrieve aerosol parameters over land, the algorithm makes use of the reflectances measured in
 4 three of MODIS' 36 spectral channels, 0.47 μm , 0.65 μm and 2.1 μm (Levy et al., 2007a). These
 5 are provided in nominal spatial resolution of 0.5 x 0.5 km^2 (at nadir) other channels (some at 0.5
 6 x 0.5 km^2 , some at 1 x 1 km^2 resolution) are used for identifying appropriate surfaces for
 7 retrieval, and masking clouds, snow and ice. While the 10 km^2 (standard product – MxD04_L2,
 8 x=O for Terra, x=Y for Aqua) begins with 20x20 = 400 aggregation of 0.5 km^2 native-resolution
 9 pixels, the 3 km^2 aerosol retrieval box starts with 6 x 6 = 36 aggregation of such pixels. Native
 10 pixels are removed, as to retain only the ones most appropriate for a dark-target, over-land
 11 retrieval. Native pixels tagged as too bright for the dark target algorithm, or identified as
 12 containing cloud, or a water, snow or ice surface are removed from the aggregation. The
 13 remaining native pixels are sorted from darkest to brightest, and the darkest 20% and brightest
 14 50% of all remaining native pixels are removed as well. For the 3 km^2 retrieval this means at
 15 most 12 native pixels will remain, and likely fewer. For minimum statistical robustness, the 3 km^2
 16 algorithm requires at least 5 native pixels (out of the initial 36). If there are insufficient native
 17 pixels (e.g <5), output parameters are assigned fill values and no retrieval is attempted. Based on
 18 the aggregation and filtering, it is likely that there will be pixels native used by the 3 km^2
 19 retrieval that would have been discarded by the coarser (10 km^2) standard resolution product
 20 (Remer et al., 2013). For whichever product resolution, the remaining native pixels are
 21 averaged, leading to a single set of spectral reflectance values that drives the aerosol retrieval.
 22 Except for downstream decisions based on number of native pixels used, the 10 km^2 and 3 km^2
 23 retrievals proceed identically.

24

25 The retrieval uses a Look Up Table (LUT) procedure in which the LUT is constructed a priori of
 26 simulated top-of-atmosphere reflectances, calculated using assumptions of aerosol optical
 27 properties based on AERONET inversions (Dubovik and King, 2000) and radiative transfer. The
 28 surface reflectance is constrained by assuming a relationship between reflectance at 2.1 μm with
 29 the reflectances at 0.65 μm and 0.47 μm that is a function of a Normalized Difference Vegetation
 30 Index using Short Wave Infrared (SWIR) bands ($\text{NDVI}_{\text{SWIR}}$) and scattering angle (Levy et al.,
 31 2007a, 2007b; 2013). The algorithm finds the AOD that minimizes the differences between the



1 MODIS-observed mean spectral reflectances of the retrieval square and the simulated reflectance
 2 values of the LUT. The primary output of the retrieval is the AOD at 0.55 μm . Using a series of
 3 tests the algorithm assigns a quality assurance flag (QAF) of either 3, ‘good quality’ or 0, ‘bad
 4 quality’ to the retrieval. These values can be interpreted as “confidence” in the aerosol retrieval
 5 (whether the retrieval proceeded nominally, and whether there are enough native-resolution
 6 pixels). For quantitative use of the 10 km^2 product, the MODIS DT team has recommended
 7 limiting to QAF=3 retrievals (≥ 51 native pixels out of 120). The similar ratio for the 3 km^2
 8 product is ≥ 5 out of 12. Any fewer, and there is too little statistical information for any
 9 confidence in an aerosol retrieval. Therefore, for the 3 km^2 product, any fewer than 5 native
 10 pixels automatically receives QAF=0. QAF values assigned as 1 or 2 are based on other criteria.

11

12 **3. Data Sets**

13

14 **3.1. MODIS 3 km AOD**

15

16 The primary data set of this study is the Collection 6 MODIS dark target retrieved aerosol optical
 17 depth at 3 km spatial resolution, derived from Terra reflectances (MOD04_3K), or Aqua
 18 reflectances (MYD04_3K), as described in Section 2. These are publicly available and can be
 19 downloaded from <https://ladsweb.modaps.eosdis.nasa.gov/>. Of the products in the data sets, we
 20 analyze only the AOD at 0.55 μm .

21

22 Applying identical algorithms to two different sensors does not guarantee identical results (Levy
 23 et al., 2015). The two MODIS DT data sets, one from MODIS-Terra and one from MODIS-
 24 Aqua must be addressed separately as individual and independent products, even though they
 25 have been created from identical algorithms with no specific tuning of parameters for each
 26 sensor. While MODIS-Terra and MODIS-Aqua began as near-identical sensors, they have
 27 evolved over their lifetimes to develop their own instrumental characteristics. For example,
 28 some detectors in Aqua’s detector array at some wavelengths have died, resulting in fewer
 29 available reflectance pixels at those wavelengths. Terra’s detector array has not lost any
 30 detectors. At the same time, we have seen drift in some of Terra’s wavelengths, resulting in
 31 measureable artificial trends in MODIS-Terra’ aerosol products (Levy et al., 2013; Sayer et al.,



2015; Lyapustin et al., 2014). The most flagrant of those MODIS-Terra trends have been mitigated by aggressive radiometric calibration (Toller et al., 2013), which has been applied for creating the C006 DT products. Note that some projects (e.g. Lyapustin et al., 2014; Sayer et al., 2015) have since introduced additional mitigation, however, the DT retrieval has not applied these strategies. In this work, we will separately analyze the C006 aerosol products from the two MODIS sensors, (MOD04_3K and MYD04_3K), to provide users with clear information on the strengths and limitations of each one.

3.2. AERONET AOD

The Aerosol Robotic Network (AERONET) is NASA's global ground network of CIMEL sun-sky radiometers that make measurement of direct transmitted solar light and scattered sky light at several wavelengths during daylight hours (Holben et al., 1998). In this work, only the direct sun measurements will be used. AERONET processes these spectral measurements to derive AOD in their respective wavelengths. The AERONET spectral AOD product is a community standard for satellite-derived AOD validation, given that AERONET's AOD uncertainty of 0.01-0.02 (Eck et al., 1999) is sufficiently more accurate and precise than can be expected by any satellite retrieval. The typical temporal frequency of direct sun measurements is every 15 minutes. The network consists of hundreds of stations, located globally, across all continents and in a wide variety of aerosol, meteorological and surface type conditions. Only stations that sufficiently represent land areas will be used here, which means we are not comparing with observations taken on small islands, ocean platforms or mobile ships. The validation procedure requires calculating the spatio-temporal statistics of a collocated MODIS-retrieved and AERONET-measured AOD pair (Ichoku et al. 2002; Petrenko et al., 2012; Munchak et al., 2013; Remer et al., 2012). Figure 1 shows the location these stations and color-coding represents the number of collocated AERONET-MODIS AOD pairs over the station.

The configuration of the spectral bands varies, but typically is centered at 0.34, 0.38, 0.44, 0.50, 0.67, 0.87, and 1.02 μm . Here we use a quadratic log-log fit (Eck et al., 1999) to interpolate AERONET AOD to 0.55 μm to match the primary MODIS AOD product. AOD data from AERONET are reported for three different quality levels: unscreened (level 1.0), cloud screened



(level 1.5) and cloud screened and quality assured (level 2.0). We will only use level 2.0, version 2.0 AERONET AODs in this study.

3

4. Spatial and Temporal Collocation

5

We have created a collocated data set (CDS) of both MODIS-Terra and MODIS-Aqua with AERONET for nearly the entire mission (2003-2015 for Aqua and 2000-2015 for Terra). MODIS Collection 6, 3km AOD retrievals over land are collocated with AERONET version 2.0, level 2.0 AOD measurements using a modified spatio-temporal technique as outlined by (Ichoku et al., 2002; Petrenko et al., 2012; Munchak et al., 2013; Remer et al., 2013). From here on, we use the term “pixels” to refer to the MODIS retrieval product (e.g. 3 or 10 km² resolution); if referring to the native MODIS pixel resolution (e.g. 0.5 km²) we will denote as “native pixel”.

In previous validation studies of the standard 10 km product the spatio-temporal statistics were based on either 5 x 5 MODIS product pixels (~50x50 km² box) centered on the AERONET station (Ichoku et al., 2002; Levy et al., 2010) or all the MODIS product pixels within a 27.5 km radius around the AERONET station (Petrenko et al., 2012) matched with the temporal statistics of ±30 minutes of AERONET observations centered at satellite overpass time. These large spatial collocation boxes will not properly test the accuracy of finer resolution satellite products to represent small-scale aerosol gradients. Therefore, Remer et al., (2013) and Munchak et al. (2013) moved to a 7.5 km radius and ±30 minutes of overpass. The 7.5 km radius encompasses roughly 25 AOD pixels at nadir, which is analogous to the number of product pixels used with the coarser resolution product. In this study spatial statistics are calculated from all MODIS product pixels falling within a box of 0.15°x0.15° (latitude x longitude) centered over an AERONET location. Except for Polar Regions, this is similar to a 15 x 15 km² box or 7.5 km radius search at the nadir. Temporal statistics are calculated from all AERONET observations of AOD within ±30 minutes of satellite overpass.

As recommended by the MODIS DT science team (Levy et al., 2010), unless otherwise specified, only AOD pixels with quality assurance flag ‘very good’ (QAF=3) were included in averaging over the AERONET sites. To be consistent with previous validation exercises (Levy et



al., 2010), we have retained the collocated data sets only when there were at least 5 MODIS product pixels (out of a possible 25) and 2 AERONET measurements (out of a possible 2-4). The collocated data set (CDS) consists of 574 AERONET stations with 90,162 collocated pairs for MODIS-Terra and 71,248 collocated pairs for MODIS-Aqua. Figure 1 shows the number of CDS pairs at each AERONET station, for Terra and Aqua.

Thus, a data set (i.e. CDS) of collocated MODIS-AERONET pairs of AOD at 0.55 μm is created that can be organized and subsampled in any number of configurations. In any subsample, or for the entire data set, these ordered pairs can be plotted, one against the other to create a scatterplot, and collocation statistics calculated. We will use the following statistical parameters to quantify how well the MODIS retrievals match their collocated AERONET counterparts (Hyer et al, 2011):

- Correlation coefficient (R),
- Slope of the linear regression line,
- Root Mean Square Error (RMSE)

$$\text{Mean Bias} = \frac{1}{N} \sum (\text{MODIS AOD} - \text{AERONET AOD}) \dots\dots\dots (1)$$

Percentage of collocations falling within expected error,

$$\text{EE} = \pm(0.05 + 0.20 \times \text{AERONET AOD}) \dots\dots\dots (2)$$

Error Ratio (ER),

$$\text{ER} = (\text{MODIS AOD} - \text{AERONET AOD}) / \text{EE} \dots\dots\dots (3)$$

The coefficients in the EE equation were determined from evaluation of the 3 km^2 product over the six months of Aqua data analyzed by Remer et al., (2013). Those limited results suggested that expected error bounds should be broadened to the values seen in Eq. (2) from those derived for the 10 km^2 product ($\text{EE} = \pm 0.05 \pm 0.15 \times \text{AERONET AOD}$).

The number of collocations (N) is another parameter used to evaluate the 3 km^2 retrieval in the collocation data set.



5. Validation Results

5.1. Global Statistics

We first compare MODIS 3 km² AOD retrievals against collocated AERONET values, for both the recommended ‘high quality’ retrievals (QAF=3) and for all the retrievals, regardless of quality, keeping Terra and Aqua results separate. Results are plotted in Figure 2. Note that the 3 km² product only tags data as either ‘high quality’ or ‘low quality’. Table 1 presents the statistical parameters corresponding to this analysis while considering various combinations of QAFs.

Globally, there is strong correlation between MODIS 3 km² AOD and collocated AERONET equivalents. However, there is scatter and a positive bias to the retrievals, more so for Terra than Aqua, even though the correlation is similar between the satellites. Identified retrieval quality matters to product accuracy with QAF=3 showing stronger correlation, smaller RMSE and more retrievals falling within expected error than QAF=0, but the high quality data set loses about 20% of the retrievals. Figure 3 shows that the differences between Terra and Aqua in how they match AERONET values are much more apparent than the differences between QAF levels of the same satellite sensors. We note that only the high quality (QAF=3) Aqua 3 km² retrievals meet expectations in terms of falling within the standard expected error bars (Remer et al., 2012; and Eq. 2).

Table 1 also shows the corresponding validation statistics for the 10 km² product for QAF=3, distinguishing between Terra and Aqua. The 10 km² product, as expected, more closely matches AERONET values, having higher correlation, lower bias and RMSE, and producing more retrievals that fall within expected error bounds than does the 3 km² product. We note that even in the 10 km² validation statistics, mean bias for Terra is 0.03 higher than for Aqua, which is the same difference between sensors as found for the 3 km² product. The results in Table 1 confirm Remer et al. (2013)’s conclusion that the 3 km² product is less accurate than the standard 10 km² product. The remainder of the paper will be devoted to exclusively analyzing the differences



1 between the 3 km² product and AERONET, without further reference to the standard 10 km²
 2 product.

3

4 **5.2. Regional Statistics**

5

6 The accuracy of the 3 km AOD retrievals will be regionally and locally specific, depending on
 7 how well retrieval assumptions of surface and aerosol optical properties match actual conditions,
 8 local cloud conditions that may or may not introduce uncertainty into the retrieval and the
 9 spatial/temporal variability of the area that may create biases in the collocation methodology that
 10 depends on assumptions of aerosol homogeneity. Here we investigate how well the MODIS 3
 11 km² product matches AERONET over individual AERONET stations.

12

13 For the regional and local analyses, we will use only QAF=3 retrievals and calculate the same
 14 collocation statistics for each station, individually. Figure 4 plots the values for correlation
 15 coefficient, mean bias, percentage within expected error, and RMSE for each station that
 16 reported at least 100 collocations over the entire time series. In general, the MODIS 3 km²
 17 retrievals show high correlations over much of the northern mid-latitudes where there are
 18 AERONET stations in abundance. This relative abundance drives the global statistics of Table
 19 1. Correlation breaks down at some stations in California and the arid southwest of North
 20 America, in the Caribbean and Central America, Insular SE Asia and Australia, and especially in
 21 southern South America.

22

23 Not all stations with strong correlations exhibit small mean biases. For example, MODIS 3 km²
 24 retrievals severely under predict AOD in the stations of west Africa, falling well below expected
 25 error there, even though those stations report high correlations with AERONET. Stations in
 26 Australia, show relatively small mean biases and high percentages meeting expectations, despite
 27 poor correlations, suggesting that the poor correlations are the result of small dynamic range in
 28 the scatter plots that occur when AOD is consistently low.

29

30 In Figure 4, we see the local nature of the validation statistics. Stations in close proximity to
 31 each other can report very different statistics. For example, the stations clustered across northern



1 India, and those in an array across central South America (Brazil) range from strong positive to
 2 negative mean biases and RMSE error from 0.05 to 0.20, even though these groupings of stations
 3 will fall within the same region as defined in Figure 5. This is apparent in almost any region.
 4 Some of this variability may be due to differences in the temporal extent of the AERONET
 5 record at each individual station, so that even if stations are in close proximity in space, they may
 6 actually be making measurements in entirely different years or seasons. Other differences may
 7 be related to topography, urban surfaces, or other factors. Still, the variability seen in Figure 4
 8 shows how local conditions, and possibly the individual characteristics of the time series affect
 9 validation statistics.

10

11 The final point to note in Figure 4 is the difference between Terra and Aqua. For example, in the
 12 mean bias plots we see how the mean bias across the North American central plains fades from
 13 approximately positive 0.04-0.05 to slightly negative. For many of the stations, positive mean
 14 biases decrease from Terra to Aqua. This is in agreement with the global statistics presented in
 15 Table 1.

16

17 For the regional analysis, we will use only QAF=3 retrievals and calculate the same collocation
 18 statistics for all stations within each region. The 17 regions were defined following Hyer et al.,
 19 (2011), and as shown in Figure 5

20

21 Table 2 presents the regional validation statistics for each region defined in Figure 5. We know
 22 from previous analyses presented above that there are distinctive differences between Terra and
 23 Aqua mean biases; however, in calculating the regional statistics of Table 2, we combine Terra
 24 and Aqua collocations. Only QAF=3 retrievals are included.

25

26 The majority of collocations are found in the northern mid-latitude regions, with E. and W.
 27 CONUS (East and West Continental United States) representing 25% of the total collocations
 28 and Europe Mediterranean and Eurasian Boreal representing another 34% of the total. MODIS 3
 29 km retrievals from E. CONUS, Europe Mediterranean and Eurasian Boreal show very good
 30 overall agreement with AERONET, exhibiting $R \geq 0.78$, bias ≤ 0.05 and at least 2/3 of retrievals
 31 falling within expected error. W. CONUS retrievals agree with AERONET less well, exhibiting



some of the highest positive biases of any region on the globe. These four regions drive the global validation statistics, which reflect both the good agreement of E. CONUS and Europe, and the high bias of W. CONUS.

Regions where MODIS 3 km² retrievals exhibit especially good agreement with AERONET collocations include E. CONUS and Europe, as mentioned above, as well as north/central South America, equatorial and southern Africa, and Australia. Australia is particularly interesting, because even though its correlation is low, 75% of retrievals there fall within expected error bounds, suggesting small dynamic range in the scatter plots. Regions where retrievals exhibit especially poor agreement with AERONET include W. CONUS, as mentioned above, plus Central America, SW Asia, East Asia Mid-Latitudes, south South America and Insular Southeast Asia. East Asia is interesting because of its high correlation, but with insufficient retrievals falling within expected error bounds. It appears that the MODIS retrieval is sensitive to aerosol there, but is incorrectly modeling the aerosol and surface optical properties. Validation statistics are especially poor in Southwest Asia, where there are very few stations and collocations.

5.3 Error dependencies

We next explore the relationship between MODIS-AERONET 3 km² AOD differences and various parameters for the global collocation data set, keeping Terra and Aqua collocations separate. At each collocation, the AERONET AOD is subtracted from the MODIS AOD, so that a positive difference indicates a positive MODIS bias. The data is then sorted according a particular parameter in the database. Collocations are grouped into 87 bins for Terra and 67 bins for Aqua, each containing 1000 collocations. Thus, there are equal numbers of collocations in each bin, but the bins are not equally spaced along the x-axis, this way to calculated statistics free from the sample size. The mean, median and standard deviations of the MODIS-AERONET differences are calculated for each bin.

Figure 6 shows the results of this analysis as a function of AOD, both the true AOD, as measured by AERONET and by the MODIS-retrieved AOD. The AOD differences depend very little on the true AOD. There is some suggestion of a positive-negative-positive shift of differences at the



1 very lowest AOD (<0.1), but overall the differences are flat. Terra exhibits an overall positive
 2 mean bias against AERONET of about 0.06, with the bias in Aqua much less noticeable.
 3 Plotting these differences against the MODIS-retrieved AOD provides the user to evaluate an
 4 individual retrieval, given only a retrieved parameter. Here we see a distinctive pattern between
 5 MODIS AOD bias and MODIS AOD. The higher the retrieved AOD, the greater the positive
 6 difference between MODIS and AERONET. Significant biases of >0.10 are seen for MODIS
 7 AOD values > 0.40 . For retrieved AOD < 0.10 , the mean differences between MODIS and
 8 AERONET are negative. This indicates that a high value of retrieved AOD has greater
 9 probability of being too high than too low, and a low value of retrieved AOD has a greater
 10 probability of being too low than too high. These results are expected, as high AOD retrievals
 11 are more sensitive to true aerosol properties whereas true surface properties become more
 12 important in low AOD retrieval.

13
 14 Figure 7 shows how MODIS-AERONET AOD differences vary as function of AOD variability
 15 and availability of high-resolution (0.5 km^2) native pixels for retrieval to be performed at 3 km^2
 16 resolution. Standard deviation of the retrievals in the 5×5 collocation box is a measure of the
 17 homogeneity of the aerosol across the box. The collocation methodology assumes that MODIS
 18 spatial statistics will match AERONET temporal statistics, which holds best if the aerosol field is
 19 homogeneous in space and time at the AERONET station. As variability across the box
 20 increases (i.e. STD (AOD)), we expect differences between MODIS AOD and AERONET AOD
 21 to grow. However, we do not have a priori expectations of whether that growth will be positive
 22 or negative. We see from Figure 7 (top row), that differences are increasingly positive as
 23 variability increases. This is because the standard deviation is not normalized, and the
 24 differences increase simply because the AOD is increasing as it does in Figure 6.

25
 26 Another test of the collocation methodology assumptions is to look for error dependencies on the
 27 number of MODIS retrievals within the 5×5 collocation box. Note that the methodology requires
 28 at least 5 retrievals to represent the box and may have as many as 25. We see from Figure 7
 29 (middle row) that there are dependencies. Fewer numbers of retrievals are associated with
 30 positive differences, but having almost all of the 25 retrievals available is associated with
 31 negative differences. We understand, how collocation statistics might be skewed by having



fewer numbers of retrievals available to match AERONET, especially if the aerosols across the collocation box were not spatially homogeneous. Also fewer numbers of retrievals may be a result of marginal retrieval conditions caused by clouds and unfavorable surface conditions. It is less easy to understand the negative differences when the box is especially well represented with sufficient retrievals, and this require further investigation on individual retrieval.

The bottom row of Figure 7 shows the MODIS-AERONET AOD differences as a function of the average number of reflectance native pixels (0.5 km^2) used by the MODIS 3 km^2 retrieval in producing a value of AOD. The retrieval begins with 36 native pixels, and after masking, sorting and discarding; we left with between 5 and 12 pixels. The number of pixels used by the retrieval is indication of how much masking was required. If 12 pixels remain, then no masking was required and the situation is cloud-free and over favorable surfaces. If only 5 pixels remain, there are conditions in that retrieval that could raise concerns. In Figure 7, we see that the fewer the pixels used by the retrieval (i.e., more masking is needed), the higher the positive bias, especially for Terra. This suggests, in the Terra retrieval, that clouds or unfavorable surface conditions are contributing to the high bias we are seeing in the global data set. Interestingly, MODIS-AERONET differences are negative when masking is at a minimum, similar as to when the collocation box contains almost all possible retrievals. It seems that cloud-free situations with appropriate surface features are associated with MODIS under predicting AERONET AOD. The same functional relationship is apparent in the Aqua data set also, but the biases, both high and low, are less pronounced.

Figure 8 shows the MODIS – AERONET AOD differences as a function of geometry. The top row plots the differences against scattering angle, where we see positive bias increasing towards the extreme backscattering angles. The functional relationship is similar in both Terra and Aqua, but Terra's positive bias is more pronounced. The bottom row plots the differences against sensor view (i.e. zenith) angle, where the Aqua differences show little dependence on view angle, but the Terra differences increase positive biases in near nadir views. Geometrical dependencies in bias generally point to systematic inaccuracies in retrieval assumptions. These can be either in terms of surface angular functions or in aerosol optical properties. However, the difference between Terra and Aqua sensor zenith angle dependencies suggests an issue with



instrument characterization, which could include geometrical functionality due to the need to calibrate across the scan mirror.

5.4 Temporal Changes

Examining temporal changes of validation statistics across the entire time series of the collocation database further characterizes the accuracy of the 3 km² AOD product. Figure 9 plots monthly mean error ratios (Eq.3), and number of collocations for the time series of Terra (red) and Aqua (blue), separately. The error ratios (ER) compare the actual error (bias) to the expected errors (EE). The $-1 \leq ER \leq 1$ means the actual errors are smaller than EE whereas $|EE| > 1$, indicates a poor match. Even if the MODIS sensors and the algorithm were entirely consistent during the time series, AERONET stations go on and off line, causing global validation statistics to shift in local and regional emphasis, and introducing temporal variation in the global results. Therefore, we have selected 26 AERONET stations (Table 3, Figure 9) with long-term data records with consistent collocation over the entire time series for this analysis. The analysis over these selected stations allow us to examine the change in bias (and error ratios) over a longer time period without change in spatial and temporal distribution of AERONET stations. Only QAF=3 retrievals are used. During the 15 years of the collocation data set many factors have changed. For example, satellite sensor characterization is an ongoing process that employs several different measures of radiometric drift and then continuously adjusts calibration parameters to compensate for that drift. Thus, even though the algorithm remains consistent throughout the data record, the inputs to that algorithm may not be, despite the best efforts of the MODIS Characterization Team.

The time series of the monthly statistics shows strong seasonal variation of mean bias and number of collocations. Strong positive bias occurs in the April-August time period, followed by low or even negative bias in the October – February period

In addition to the seasonal variability, Figure 9 also exposes long-term temporal trends. There is a steady increase of the number of collocations per month, nearly doubling from the early years, up and through the beginning of 2012, as the AERONET network expands over time. The last



1 few years of the record show a decrease in collocations, in some part attributed to the lag in
 2 promoting AERONET records from Level 1.5 to Level 2.0. We only use AERONET Level 2.0
 3 for collocations.

4
 5 The temporal mean biases for the entire time series are 0.04 and 0.014 for Terra and Aqua,
 6 respectively, corresponding to Error Ratio (ER) in Figure 9 of 0.55 and 0.2. The mean biases also
 7 exhibit temporal trends with biases beginning to increase around 2008/2009. The bias for Terra
 8 increased from 0.038 to 0.048 whereas these numbers for Aqua are 0.014 and 0.016. The
 9 corresponding ER increase for Terra in 2008 is from 0.48 to 0.65. The increase in ER for Aqua
 10 is negligible.

11
 12 The systematic higher biases exhibited by Terra as compared with Aqua agree with the global
 13 analysis presented above. This offset in bias between the two MODIS sensors appears systematic
 14 from the beginning of the Aqua record to the end of the time series, although the magnitude of
 15 that offset increases over time as Terra's biases grow. The systematic greater number of
 16 collocations in the Terra data set than in Aqua's may result from diurnal cloud patterns that
 17 create cloudier conditions in the afternoon during Aqua overpass than during Terra's morning
 18 one. More clouds in the afternoon (King et al., 2013) may reduce the number of possible
 19 collocations. However, instrumental differences affecting available retrievals are another
 20 possibility.

21 22 **6. Discussion and Conclusions**

23
 24 To validate the MODIS 3 km² AOD products (MOD04_3K and MYD04_3K), which became
 25 publicly available in the MODIS Collection 6 release, we created a database of collocations
 26 between the product and AERONET observations for the extent of the MODIS record from 2000
 27 – 2015. Collocation criteria employed 5x5 MODIS retrievals centered at the AERONET station
 28 and all AERONET observations ±30 minutes of satellite overpass. Thus, the collocation box is
 29 approximately 15 km per side, for nadir views. Only Level 2, quality assured AERONET
 30 observations are included, and AERONET AOD is interpolated to 0.55 μm to match MODIS



1 values. Overall there are over 90,162 high quality collocations of Terra retrievals and over
 2 71,248 high quality collocations for Aqua.
 3
 4 The validation statistics examined include mean bias, regression slope, correlation coefficient
 5 and percentage falling within expected error bounds. In this validation exercise we hold the 3
 6 km² AOD product to expected error thresholds of $\pm 0.05 \pm 20\%$ (Remer et al., 2013). We find that
 7 the global 3 km² AOD product displays skill in matching AERONET observations with a
 8 correlation coefficient of 0.87, but there is RMSE of 0.15 and 0.13 for Terra and Aqua,
 9 respectively. The scatter is not random, but exhibits a mean positive bias of 0.06 for Terra and
 10 0.03 for Aqua. The coarse product error bounds capture 2/3 of the Aqua 3 km AOD retrieval, but
 11 less than 63% of the Terra retrievals. There is significant degradation of validation accuracy if
 12 MODIS retrievals of Poor data quality (QA<3) are included in the analysis. Approximately 20%
 13 of all MODIS retrievals were identified as Poor quality. Thus, on a global basis we recommend
 14 using only QAF=3 MODIS 3 km² retrievals for quantitative analysis. If doing so, then the
 15 expected error for the Aqua product is $\pm 0.05 \pm 0.20\text{AOD}$, on a global basis, but only $\pm 0.06 \pm$
 16 0.20AOD for Terra, where AOD is the true AOD. However, a more accurate representation of
 17 Terra's expected error is to account for the positive bias with asymmetrical error bounds: $-0.03 -$
 18 0.20AOD and $+0.13 + 0.20\text{AOD}$. The expected error bounds contain 2/3 of all AOD retrievals.
 19 To assess the mean bias of the retrieval based on the retrieved AOD, we find that the mean bias
 20 can be modeled as $0.19 + 0.17 \cdot \ln(\text{AOD_MODIS} + 0.25)$ for Terra and $0.15 +$
 21 $0.14 \cdot \ln(\text{AOD_MODIS} + 0.25)$ for Aqua. Note that mean bias itself is subject to uncertainty.
 22
 23 We find a wide range of accuracy in the 3 km² product locally and regionally, with spatially
 24 contiguous stations sometimes exhibiting significantly different validation statistics. The
 25 distribution of validation sites is highly skewed towards the northern mid-latitudes with over
 26 50% of all collocations in the database resulting from these areas. Within the northern mid-
 27 latitudes, eastern North America and Europe with boreal Eurasia show some of the best
 28 agreement with AERONET; however, western North America shows some of the poorest
 29 agreement. Regions outside of the northern mid-latitudes are less well represented in the
 30 database, but we find that north/central South America including the Amazon, equatorial and
 31 southern Africa, and Australia show good agreement with AERONET. Mexico and the



1 Caribbean, southern South America, SW Asia, East Asia, and the maritime continent of
 2 Southeast Asia generally show poor agreement. No attempt was made to isolate urban regions
 3 from rural ones, or to otherwise sort the data by surface type.
 4

5 The difference between MODIS-retrieved 3 km² AOD and AERONET observed values are
 6 mostly independent of true AOD. This is unexpected as error bounds are defined as a function of
 7 the percentage of AOD ($\pm 0.05 \pm 0.20 \cdot \text{AOD}$). However, the mean differences between MODIS
 8 3 km² AOD and AERONET are dependent on AOD variability and availability with the more
 9 variable the AOD, and the greater need for masking clouds and unfavorable surfaces in the
 10 original retrieval, the higher the positive offset between MODIS and AERONET. Some of this is
 11 due to the conditions of the original MODIS retrieval, and some is due to the difficulties of a
 12 spatio-temporal match-up in the collocation methodology. Interesting and unexplained is the
 13 tendency for the differences between MODIS and AERONET to go negative when conditions
 14 appear to be homogeneous and cloud-free. We also find error dependencies on geometry, with
 15 greater error in the far backscattering region and for Terra only, greater error in near-nadir views.
 16 Some of these geometrical errors are introduced by uncertainties in the assumptions of surface
 17 characteristics and aerosol optical properties in the MODIS retrieval, but the difference between
 18 Terra and Aqua suggests differences in the sensors themselves.
 19

20 We continue to see differences between the sensors in how validation statistics have evolved
 21 over time. By limiting our time series analysis to only 26 AERONET stations that span the entire
 22 time series, we eliminate changes in validation statistics due to changing AERONET station
 23 distribution. We find that both sensors exhibit time series with strong seasonal dependence and
 24 higher positive biases against AERONET in the northern spring and summer, than in northern
 25 fall and winter, with Terra's positive bias always greater than Aqua's. However, during the early
 26 years of the time series, both sensors were reporting similar number of retrievals falling within
 27 expected error. This changed during 2007-2009, when Terra's accuracy began to fall off and its
 28 positive biases increased. Aqua's bias against AERONET also increased during this time frame,
 29 but not as rapidly as Terra's. While, these drifts in validation accuracy suggest changes in
 30 characterization accuracy of the MODIS sensors themselves, there are other factors. The number
 31 of collocations has fallen off towards the end of the time series. We attribute this to a lag for



1 AERONET observations to be elevated to Level 2 status. Because of this lag, there may in fact
2 be a change in the distribution of AERONET stations in the temporal collocation database,
3 despite our best intentions, and this may introduce a temporal trend in the validation statistics.
4 Furthermore, the aerosol system itself has undergone significant changes since 2000, with the
5 U.S. and Europe drastically reducing their urban/industrial emissions and substituting wildfire
6 smoke as their primary source of aerosol. Likewise emissions and resulting AOD from other
7 regions experience both long-term trends and interannual variability. The combination of
8 variations in AERONET station distribution and the changing aerosol system over the time series
9 examined may be contributing to the trends seen in the validation statistics. However, the
10 differences between Terra and Aqua are difficult to explain without pointing to sensor
11 characterization stability.

12

13 The standard 10 km² product that meets expected error at 67% and 74% levels for Terra and
14 Aqua, respectively, on a global basis is measurably more accurate than the 3 km² product
15 examined here in detail. Similarly the global standard 10 km² AOD product exhibits half of the
16 mean bias with Terra and no bias at all for Aqua. These validation statistics for the 10 km²
17 standard product are preliminary and likely to change once a more comprehensive evaluation of
18 the Collection 6, 10 km² product is completed. The 10 km² product numbers are provided here
19 only to lend perspective to our results with the finer resolution product. Given this perspective,
20 we confirm the Remer et al. (2013) recommendation that users whose interests are global should
21 use the more robust and accurate 10 km² product, and leave the 3 km² product for specific
22 applications that require the finer resolution representation of the AOD field.

23

24 This validation study only addressed the 3 km² AOD product over land, and did not evaluate the
25 over water product. The study took a global and regional view, not a local one. Users of the
26 product on a local level are encouraged to consider particular biases that may occur due to local
27 conditions. For example, we know that the MODIS Collection 6 retrieval is systematically
28 biased over urban surfaces. This is true for both the 10 km² and 3 km² products. This problem
29 has been addressed and is substantially mitigated with the release of Collection 6.1 version of the
30 algorithm (Gupta et al., 2016). In the meantime, our results here show that overall the MODIS
31 Collection 6 algorithm is producing an AOD product at 3 km² resolution with sufficient accuracy



and with biases well-characterized so that it can be used quantitatively in a wide variety of science and practical applications.

7. Acknowledgement

This work was supported by the NASA ROSES program NNH13ZDA001N-TERAQEA: Terra and Aqua – Algorithms – Existing Data Products and NASA’s EOS program managed by Hal Maring. We thank MCST for their efforts to maintain and improve the radiometric quality of MODIS data, and LAADS/MODAPS for the continued processing of the MODIS products. The AERONET team (GSFC and site PIs) are thanked for the creation and continued stewardship of the sun photometer data record; which is available from <http://aeronet.gsfc.nasa.gov>.

8. Reference

Al-Saadi, J., Szykman, J., Pierce, R. B., Kittaka, C., Neil, D., Chu, D. A., Remer, L. A., Gumley, L., Prins, E., Weinstock, L., MacDonald, C., Wayland, R., Dimmick, F., and Fishman, J.: Improving national air quality forecasts with satellite aerosol observations, *Bull. Am. Meteorol. Soc.*, 86, 1249–1261, doi:10.1175/BAMS-86-9-1249, 2005.

Christopher, S. A., and Zhang, J.: Shortwave aerosol radiative forcing from MODIS and CERES observations over the oceans. *Geophysical Research Letters*, 29(18), 1859 - doi: 10.1029/2002GL014803, 2002.

Chu, D. A., Kaufman, Y. J., Ichoku, C., Remer, L. A., Tanre, D., and Holben, B. N.: Validation of MODIS aerosol optical depth retrieval over land, *Geophys. Res. Lett.*, 29, MOD2.1–MOD2.4, doi:10.1029/2001GL013205, 2002.

Dubovik, O., KING, M.D.: 2000. A flexible inversion algorithm for retrieval of aerosol optical properties from Sun and sky radiance measurements *J Geophys Res*, 105(D16): 20,673–20,696, doi:10.1029/2000JD900282, 2000.

Eck, T. F., Holben, B. N., Reid, J. S., Dubovik, O., Smirnov, A., O’Neill, N. T., Slutsker, I. and Kinne, S.: Wavelength dependence of the optical depth of biomass burning, urban, and desert dust aerosols, *J. Geophys. Res.*, 104, 31,333– 31,349, 1999.



- 1 Gupta, P., and Christopher, S. A.: Particulate matter air quality assessment using integrated
 2 surface, satellite, and meteorological products: Multiple regression approach, *J. Geophys.*
 3 *Res.*, 114, D14205, doi:10.1029/2008JD011496, 2009.
- 4 He, Qingqing & Zhang, Ming & Huang, B. T., Xuelian.: MODIS 3 km and 10 km aerosol optical
 5 depth for China: Evaluation and comparison. *Atmospheric Environment*. 153. 150-162.
 6 10.1016/j.atmosenv.2017.01.023, 2017.
- 7 Holben, B. N., Eck, T. F., Slutsker, I., Tanre, D., Buis, J. P., Setzer, A., Vermote, E., Reagan, J.
 8 A., Kaufman, Y. J., Nakajima, T., Lavenu, F., Jankowiak, I., and Smirnov, A.: AERONET – A
 9 federated instrument network and data archive for aerosol characterization, *Remote Sens.*
 10 *Environ.*, 66, 1–16, 1998.
- 11 Hsu, N.-Y. C., Jeong, M.-J. Bettenhausen, C.: Enhanced Deep Blue aerosol retrieval algorithm:
 12 The second generation, *J. Geophys. Res. Atmos.* 118 (16): 9296–9315, 10.1002/jgrd.50712,
 13 2013.
- 14 Hyer, E. J., Reid, J. S., and Zhang, J.: An over-land aerosol optical depth data set for data
 15 assimilation by filtering, correction, and aggregation of MODIS Collection 5 optical depth
 16 retrievals, *Atmos. Meas. Tech.*, 4, 379–408, doi:10.5194/amt-4-379-2011, 2011.
- 17 Ichoku, C., D. A. Chu, S. Mattoo, Y. J. Kaufman, L. A. Remer, D. Tanré, I. Slutsker and B. N.
 18 Holben, 2002: A spatio-temporal Approach for Global Validation and Analysis of MODIS
 19 Aerosol Products. *Geophys. Res. Lett.*, 29(12), 10.1029/2001GL013205, 2002.
- 20 Kaufman, Y. J., Gobron, N., Pinty, B., Widlowski, J., and Verstraete, M. M.: Relationship
 21 between surface reflectance in the visible and mid-IR used in MODIS aerosol algorithm –
 22 theory, *J. Geophys. Res.*, 29(23), 2116, doi:10.1029/2001GL014492, 2002.
- 23 Kaufman, Y. J., Koren, I., Remer, L. A., Tanre, D., Ginoux, P., and Fan, S.: Dust transport and
 24 deposition observed from the Terra- Moderate Resolution Imaging Spectroradiometer (MODIS)
 25 spacecraft over the Atlantic Ocean, *J. Geophys. Res.*, 110, D10S12, doi:10.1029/2003JD004436,
 26 2005.



- 1 Kaufman, Y. J., Tanre, D., Remer, L., Vermote, E., Chu, A., and Holben, B. N.: Operational
2 remote sensing of tropospheric aerosol over land from EOS moderate resolution imaging
3 spectroradiometer, *J. Geophys. Res.-Atmos.*, 102(D14), 17051– 17067, 1997.
- 4 King, M.D., Platnick, S.E., Menzel, W.P., Ackerman, S.A., Hubanks, P.A.: Spatial and temporal
5 distribution of clouds observed by MODIS onboard the Terra and Aqua satellites *IEEE*
6 *Transactions on Geoscience and Remote Sensing*, 51: 3826-3852, 2013.
- 7
8 Koren, I., Altaratz, O., Remer, L. A., Feingold, G., Martins, J. V. and Heiblum, R. H.: Aerosol-
9 induced intensification of rain from the tropics to the mid-latitudes, *Nat. Geosci.*, 5(2), 118–122,
10 doi:10.1038/ngeo1364, 2012.
- 11 Koren, I., Feingold, G., Jiang, H., and Altaratz, O.: Aerosol effects on the inter-cloud region of a
12 small cumulus cloud field, *Geophys. Res. Lett.*, 36, L14805, doi:10.1029/2009GL037424, 2009.
- 13 Levy, R. C., Mattoo, S., Munchak, L. A., Remer, L. A., Sayer, A. M., Patadia, F., and
14 Hsu, N. C.: The Collection 6 MODIS aerosol products over land and ocean, *Atmos. Meas. Tech.*,
15 6, 2989-3034, doi:10.5194/amt-6-2989-2013, 2013.
- 16 Levy, R. C., Munchak, L. A., Mattoo, S., Patadia, F., Remer, L. A., and Holz, R. E.: Towards a
17 long-term global aerosol optical depth record: applying a consistent aerosol retrieval algorithm to
18 MODIS and VIIRS-observed reflectance, *Atmos. Meas. Tech.*, 8, 4083-4110,
19 <https://doi.org/10.5194/amt-8-4083-2015>, 2015.
- 20 Levy, R. C., Remer, L. A., and Dubovik, O.: Global aerosol optical properties and application to
21 Moderate Resolution Imaging Spectroradiometer aerosol retrieval over land, *J. Geophys. Res.-*
22 *Atmos.*, 112, D13210, doi:10.1029/2006JD007815, 2007a.
- 23 Levy, R. C., Remer, L. A., Kleidman, R. G., Mattoo, S., Ichoku, C., Kahn, R., and Eck, T. F.:
24 Global evaluation of the Collection 5 MODIS dark-target aerosol products over land, *Atmos.*
25 *Chem. Phys.*, 10, 10399–10420, doi:10.5194/acp-10-10399-2010, 2010.
- 26 Levy, R. C., Remer, L. A., Mattoo, S., Vermote, E. F., and Kaufman, Y. J.: Second-generation
27 operational algorithm: retrieval of aerosol properties over land from inversion of Moderate



- 1 Resolution Imaging Spectroradiometer spectral reflectance, *J. Geophys. Res.-Atmos.*, 112,
2 D13211, doi:10.1029/2006JD007811, 2007b.
- 3 Levy, R.C., L.A. Remer, J.V. Martins, Y.J. Kaufman, A. Plana-Fattori, J. Redemann, P.B.
4 Russell and B. Wenny: Evaluation of the MODIS aerosol retrievals over ocean and land during
5 CLAMS. *J. Atmos. Sci.*, 62, 974-992, 2005.
- 6 Lyapustin, A., Wang, Y., Xiong, X., Meister, G., Platnick, S., Levy, R., Franz, B., Korkin, S.,
7 Hilker, T., Tucker, J., Hall, F., Sellers, P., Wu, A., and Angal, A.: Scientific impact of MODIS
8 C5 calibration degradation and C6+ improvements, *Atmos. Meas. Tech.*, 7, 4353–4365,
9 doi:10.5194/amt-7-4353-2014, 2014
- 10 Munchak, L. A., Levy, R. C., Mattoo, S., Remer, L. A., Holben, B. N., Schafer, J. S., Hostetler,
11 C. A., and Ferrare, R. A.: MODIS 3 km aerosol product: applications over land in an
12 urban/suburban region, *Atmos. Meas. Tech.*, 6, 1747-1759, [https://doi.org/10.5194/amt-6-1747-](https://doi.org/10.5194/amt-6-1747-2013)
13 2013, 2013.
- 14 Nichol, J. E., Bilal, M.: Validation of MODIS 3 km Resolution Aerosol Optical Depth Retrievals
15 Over Asia. *Remote Sens.* 8, 328, 2016.
- 16 Oreopoulos, L., Cho, N., Lee, D., Kato, S.: Radiative effects of global MODIS cloud regimes *J.*
17 *Geophys. Res. Atmos.*, 121(5): 2299–2317 [10.1002/2015jd024502, 2016.
- 18 Oreopoulos, L., N. Cho, N., and Lee, D.: Using MODIS cloud regimes to sort diagnostic signals
19 of aerosol-cloud-precipitation interactions, *J. Geophys. Res. Atmos.*, 122, 5416–5440,
20 doi:10.1002/2016JD026120, 2017.
- 21 Petrenko, M., Ichoku, C., and Leptoukh, G.: Multi-sensor Aerosol Products Sampling System
22 (MAPSS). *Atmos. Meas. Tech.*, 5, 913-926, DOI: 10.5194/amt-5-913-2012, 2012.
- 23 Remer, L. A., Kaufman, Y. J., Tanré, D., Mattoo, S., Chu, D. A., Martins, J. V., Li, R. R.,
24 Ichoku, C., Levy, R. C., Kleidman, R. G., Eck, T. F., Vermote, E., and Holben, B. N.: The
25 MODIS aerosol algorithm, products, and validation, *J. Atmos. Sci.*, 62, 947–973,
26 doi:10.1175/JAS3385.1, 2005.



- 1 Remer, L. A., Mattoo, S., Levy, R. C., and Munchak, L. A.: MODIS 3 km aerosol product:
2 algorithm and global perspective, Atmos. Meas. Tech., 6, 1829-1844,
3 <https://doi.org/10.5194/amt-6-1829-2013>, 2013.
- 4 Remer, L. A., Mattoo, S., Levy, R. C., Heidinger, A., Pierce, R. B., and Chin, M.: Retrieving
5 aerosol in a cloudy environment: aerosol product availability as a function of spatial resolution,
6 Atmos. Meas. Tech., 5, 1823–1840, doi:10.5194/amt-5-1823-2012, 2012.
- 7 Remer, L. A., Tanre, D., Kaufman, Y. J., Ichoku, C., Mattoo, S., Levy, R., Chu, D. A., Holben,
8 B., Dubovik, O., Smirnov, A., Martins, J. V., Li, R. R., and Ahmad, Z.: Validation of MODIS
9 aerosol retrieval over ocean, Geophys. Res. Lett., 29, MOD3.1– MOD3.4,
10 doi:10.1029/2001GL013204, 2002.
- 11 Russell, P. B., Livingston, J. M., Redemann, J., Schmid, B., Ramirez, S. A., Eilers, J., Kahn, R.,
12 Chu, A., Remer, L., Quinn, P. K., Rood, M. J., and Wang, W.: Multi-gridcell validation of
13 satellite aerosol property retrievals in INTEX/ITCT/ICARTT 2004, J. Geophys. Res., 112,
14 D12S09, doi:10.1029/2006JD007606, 2007.
- 15 Salomonson, V. V., Barnes, W. L., Maymon, P. W., Montgomery, H. E., and Ostrow, H.:
16 MODIS: advanced facility instrument for studies of the Earth as a system, IEEE T. Geosci.
17 Remote, 27, 145–153, 1989.
- 18 Sayer, A.M., Hsu, N.C., Bettenhausen, C., Jeong, M., Meister, G.: Effect of MODIS Terra
19 radiometric calibration improvements on Collection 6 Deep Blue aerosol products: validation
20 and Terra/Aqua consistency J. Geophys. Res, 120, 10.1002/2015JD023878, 2015.
- 21 Toller, G., Xiong, X., Sun, J., Wenny, B. N., Geng, X., Kuyper, J., Angal, A., Chen, H., Madhavan, S.
22 and Wu, A.: Terra and Aqua moderate-resolution imaging spectroradiometer collection 6
23 level 1B algorithm, J Appl Remote Sens, 7, 3557, doi:10.1117/1.JRS.7.073557, 2013.
- 24 van Donkelaar, A., Martin, R. V., Brauer, M., and Boys, B. L.: Use of Satellite Observations for
25 Long-Term Exposure Assessment of Global Concentrations of Fine Particulate Matter, Environ.
26 Health Perspect., 123, 135–143, doi:10.1289/ehp.1408646, 2015.



- 1 Yu, H., Kaufman, Y. J., Chin, M., Feingold, G., Remer, L. A., Anderson, T. L., Balkanski, Y.,
2 Bellouin, N., Boucher, O., Christopher, S., DeCola, P., Kahn, R., Koch, D., Loeb, N., Reddy, M.
3 S., Schulz, M., Takemura, T., and Zhou, M.: A review of measurement-based assessments of the
4 aerosol direct radiative effect and forcing, Atmos. Chem. Phys., 6, 613-666,
5 <https://doi.org/10.5194/acp-6-613-2006>, 2006.
- 6 Yu, H., Remer, L., Chin, M., Bian, H., Tan, Q., Yuan, T., and Zhang, Y.: Aerosols from
7 Overseas Rival Domestic Emissions over North America, Science, 337, 566–569, 2012.
- 8 Yuan, T., Remer, L. A., Pickering, K. E., and Yu, H.: Observational evidence of aerosol
9 enhancement of lightning activity and convective invigoration, Geophys. Res. Lett., 38, L04701,
10 [doi:10.1029/2010GL046052](https://doi.org/10.1029/2010GL046052), 2011.
- 11 Yuan, T., Remer, L., and Yu, H.: Microphysical, macrophysical and radiative signatures of
12 volcanic aerosols in trade wind cumulus observed by the A-Train, Atmos. Chem. Phys., 11,
13 7119-7132, DOI: 10.5194/acp-11-7119-2011, 2011.

14

15

16

17

18

19

20

21

22

23

24



Table 1. Global statistics of comparison between MODIS 3 km AOD at 0.55 μm retrievals and collocated AERONET observations for both Terra and Aqua, corresponding to three QA categories (QAF=0 for poor quality, QAF=0,3 for all quality and QAF=3 for high quality). The data used for the 10km validation does not represent the same time period as 3km. MODIS-Terra 10km data period is 03/2000 to 06/2013 where as MODIS-Aqua 10km data period is 01/2003 – 06/2013.

Sensor	MODIS-Terra				MODIS-Aqua			
Resolution	3 km			10 km	3 km			10 km
QAF	0	0, 3	3	3	0	0, 3	3	3
N	18055	112210	90162	82997	13935	89804	71248	66945
R	0.82	0.86	0.87	0.91	0.82	0.86	0.87	0.90
Bias	0.052	0.061	0.059	0.03	0.021	0.031	0.027	0.00
Slope	1.05	1.05	1.06	1.03	0.99	1.04	1.05	1.02
RMSE	0.18	0.15	0.15	0.11	0.16	0.14	0.13	0.10
Within EE%	52.56	59.62	62.47	68.68	58.55	66.08	68.36	74.38
Above EE%	35.03	33.50	31.33	24.27	25.17	23.51	21.47	15.42
Below EE%	12.42	6.88	6.2	7.04	16.28	10.42	10.18	10.21



1 Table 2. Regional statistics of inter-comparison between MODIS and AERONET. This is using
 2 join data sets of Terra and Aqua for QAF=3 only.

3
4

Region	N	Mean AOD	R	Bias	Slope	RMSE	Within EE%	Above EE%	Below EE%
N. American Boreal	4136	0.111	0.93	0.079	1.39	0.14	56.14	43.76	0.10
E. CONUS	22450	0.129	0.90	0.029	1.22	0.09	73.41	19.74	6.85
W. CONUS	17645	0.096	0.68	0.116	1.45	0.19	45.34	53.27	1.39
Central America	2499	0.203	0.87	0.084	1.25	0.18	51.62	40.42	7.96
South America	5577	0.276	0.96	-0.007	1.20	0.16	64.16	9.16	26.68
S. South America	5393	0.107	0.63	0.048	1.13	0.18	48.54	28.48	22.97
Africa south of equator	5849	0.184	0.81	-0.020	0.71	0.10	68.44	12.84	18.72
Equatorial Africa	270	0.203	0.90	0.002	1.03	0.08	77.78	11.85	10.37
Africa north of equator	9870	0.302	0.83	-0.039	0.63	0.18	61.00	18.50	20.50
SW Asia	405	0.220	0.78	0.164	1.25	0.21	33.58	66.17	0.25
Europe - Mediterranean	39792	0.162	0.79	0.043	1.06	0.11	70.62	24.63	4.75
Eurasian Boreal	13473	0.181	0.91	0.043	1.14	0.09	73.11	24.20	2.69
East Asia Mid- Latitudes	10009	0.370	0.91	0.110	1.09	0.20	56.03	41.22	2.75
Peninsular Southeast Asia	5259	0.501	0.91	0.039	1.05	0.18	68.09	22.02	9.89
Indian Subcontinent	8449	0.479	0.86	0.070	1.05	0.19	68.35	26.78	4.86
Insular Southeast Asia	853	0.243	0.85	0.118	1.03	0.20	50.41	48.30	1.29
Australian Continent	5965	0.087	0.59	-0.021	0.57	0.08	69.52	8.92	21.56

5
6
7



1 Table 3. List of selected AERONET stations for the long-term analysis as presented in figure
 2 9.
 3

Site Name	Latitude	Longitude
Canberra	-35.271	149.111
Skukuza	-24.992	31.587
Lake_Argyle	-16.108	128.749
CUIABA-MIRANDA	-15.729	-56.021
Mongu	-15.254	23.151
Jabiru	-12.661	132.893
Chiang_Mai_Met_Sta	18.771	98.972
Kanpur	26.513	80.232
Izana	28.309	-16.499
Saada	31.626	-8.156
Nes_Ziona	31.922	34.789
TABLE_MOUNTAIN_CA	34.380	-117.680
FORTH_CRETE	35.333	25.282
Blida	36.508	2.881
Cart_Site	36.607	-97.486
Fresno	36.782	-119.773
Evora	38.568	-7.912
GSFC	38.992	-76.840
KONZA_EDC	39.102	-96.610
XiangHe	39.754	116.962
BSRN_BAO_Boulder	40.045	-105.006
Lecce_University	40.335	18.111
Rome_Tor_Vergata	41.840	12.647
OHP_OBSERVATOIRE	43.935	5.710
Carpentras	44.083	5.058
Modena	44.632	10.945

4
 5
 6
 7
 8
 9
 10
 11
 12
 13
 14
 15
 16

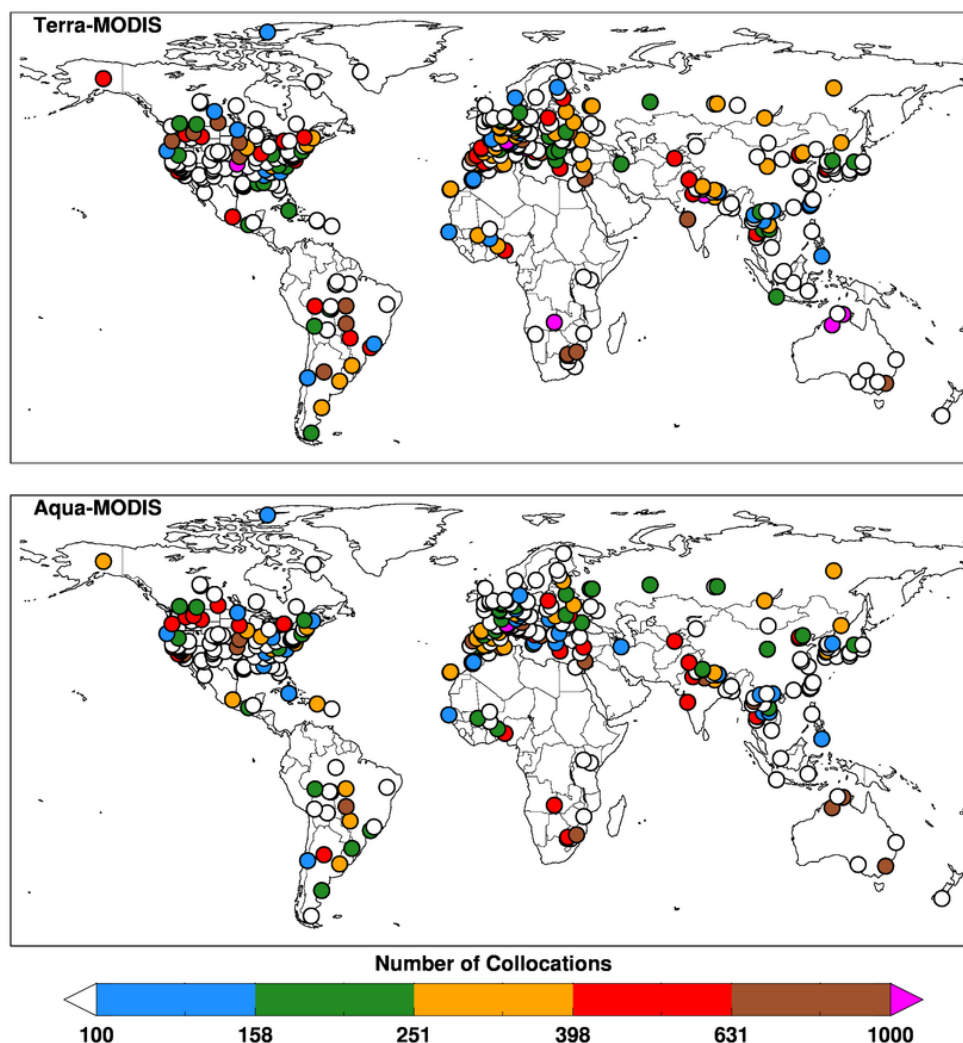


Figure 1. Locations of AERONET stations used in the validation study. The color scale shows the number of coincident MODIS-AERONET data points over each station for the entire period. Top panel is for MODIS-Terra and bottom panel for MODIS-Aqua. Most stations operated for only a subset of the 13 to 15 year record.

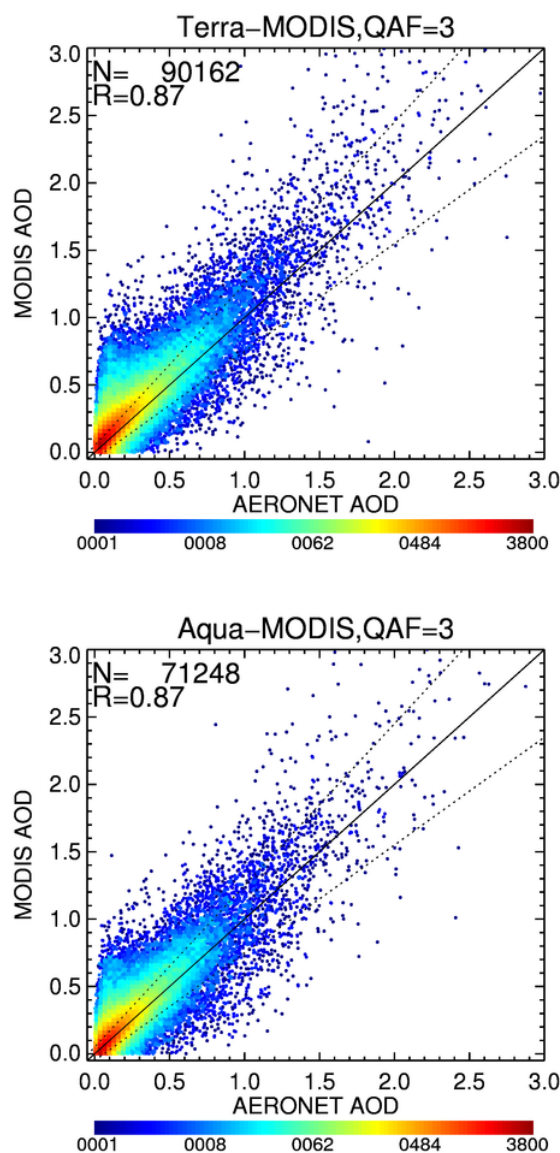


Figure 2. Two-dimensional density scatter plot of MODIS 3 km AOD versus AERONET observed AOD at 0.55 μm for the global collocation data set. The top panel is for MODIS-Terra for only the retrievals identified as ‘high quality’ (QAF=3), and the bottom panel is for MODIS-Aqua for QAF=3. The solid line denotes the 1:1 line, and the dashed lines denote the envelope of the expected error (EE), defined by Eq. 2.

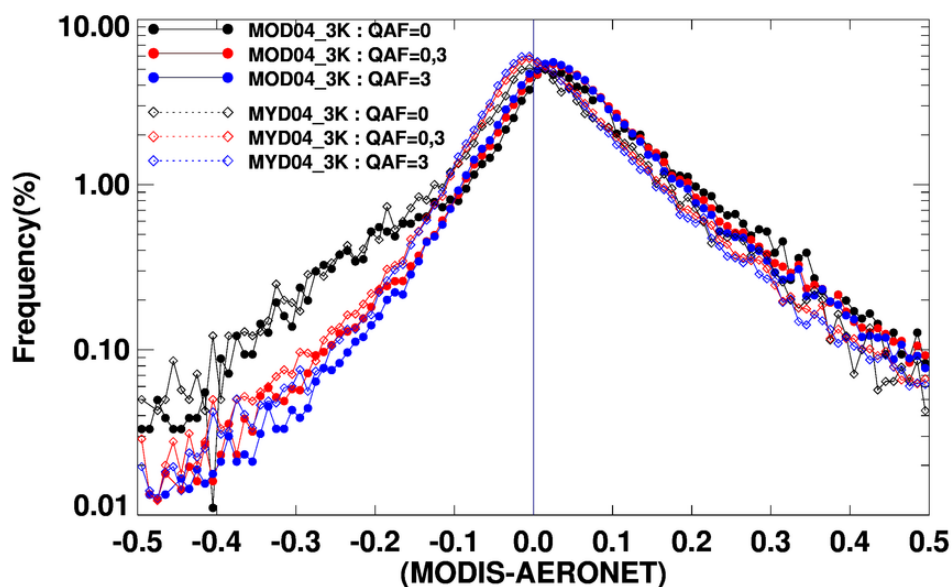


Figure 3. Global distribution of mean bias in MODIS 3 km² AOD retrievals with respect to collocated AERONET observations. The circular dots with solid lines are for Terra values and diamond dots with dotted lines are for Aqua values. The colors vary for the three quality levels (QAF=0, poor quality; QAF=3, high quality; and QAF=0&3, all quality).

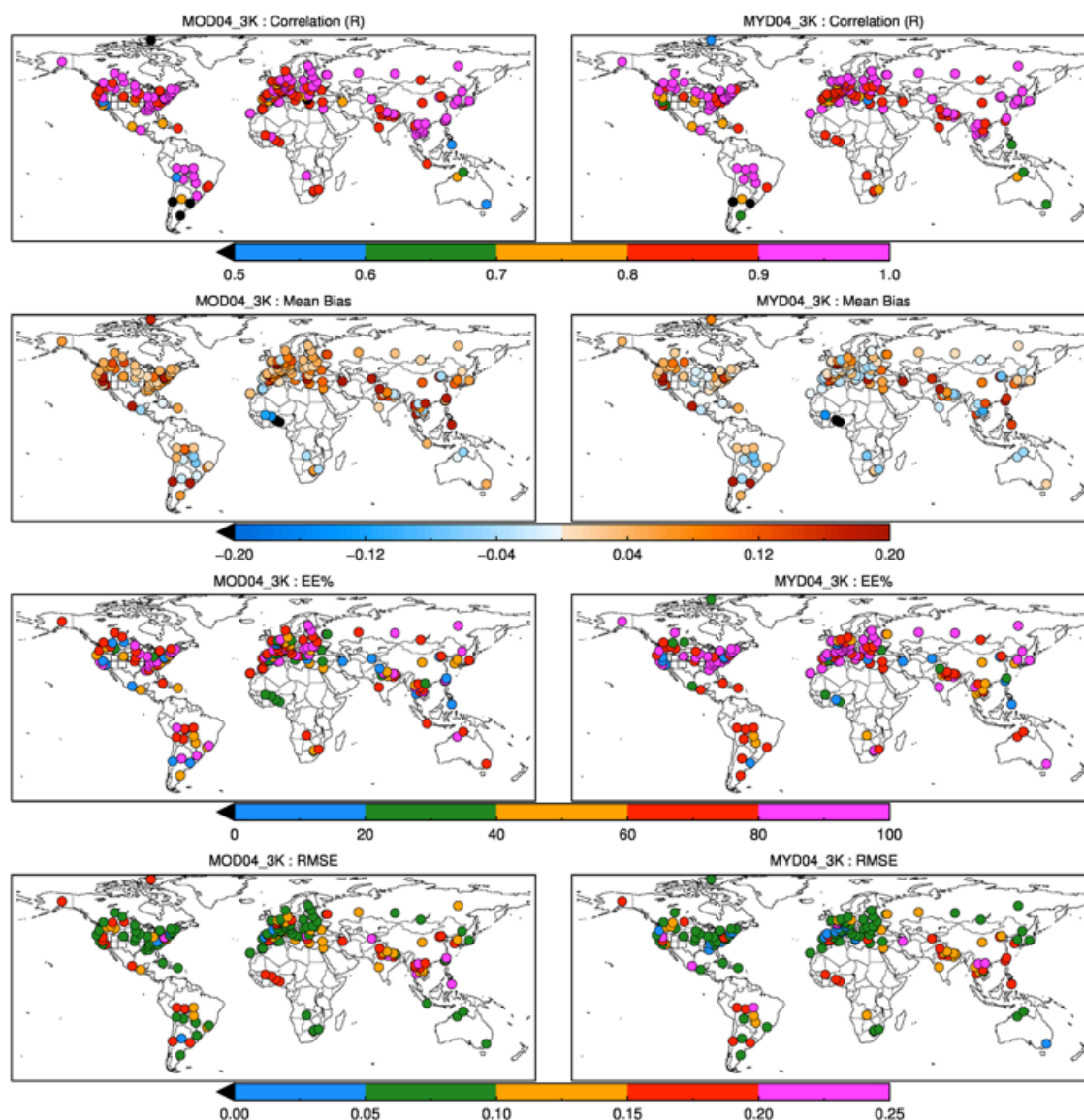


Figure 4. Statistics calculated from the collocation database at each AERONET station, individually, for Terra on the left and Aqua on the right. Shown are values for correlation coefficient (R), mean bias, percentage within expected error (EE%), and RMSE. Only stations with at least 100 collocations are plotted, which may differ between the two satellites, and only collocations with MODIS retrievals of QAF=3 are included.



1

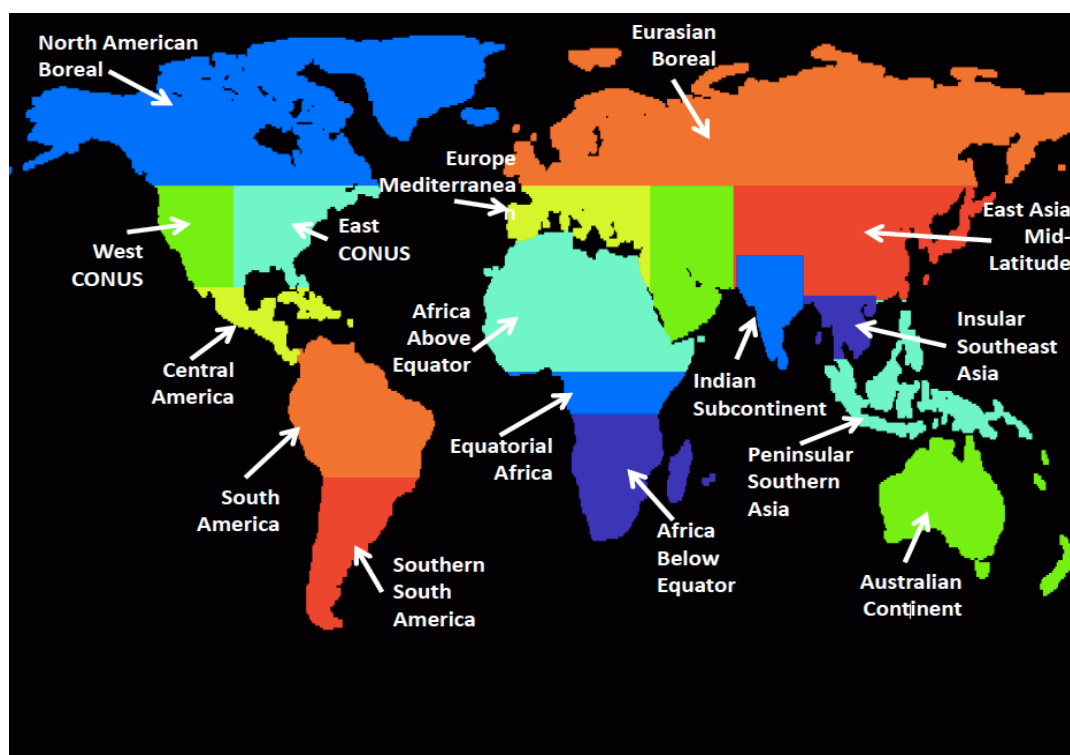


Figure 5. Map showing 17 selected parts of the world where regional analysis is performed.

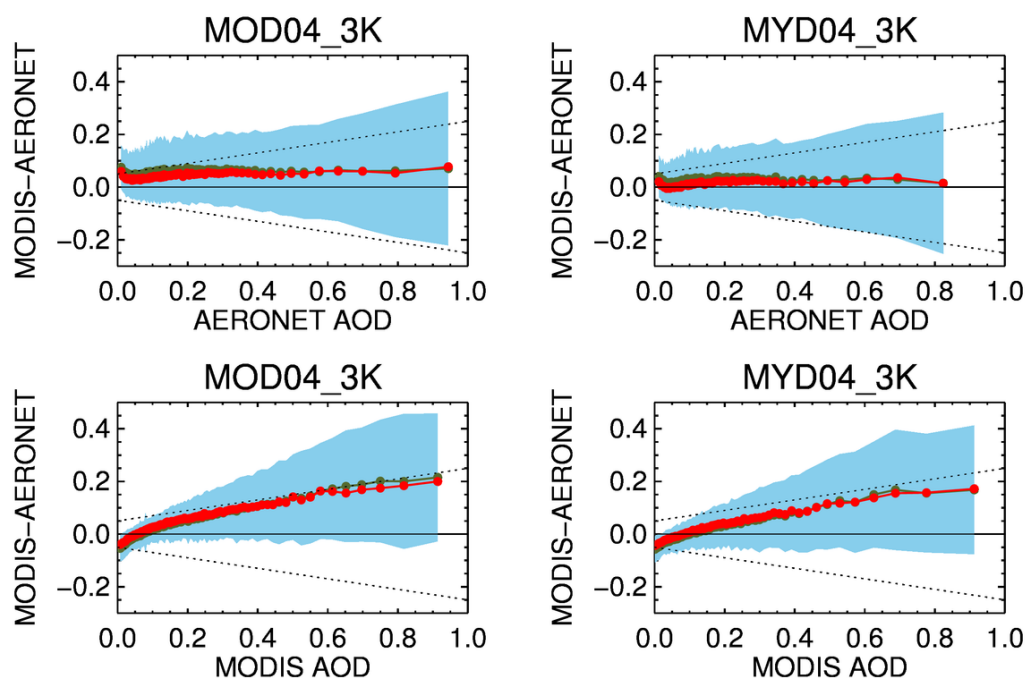


Figure 6. AOD differences between the MODIS 3 km product and AERONET for the global collocation data set, QAF=3, as a function of AERONET AOD (top), MODIS AOD (bottom). The left column shows Terra values and the right column shows Aqua. The global data set was sorted according to AOD, binned into bins with equal number of collocations, and then mean, median and standard deviation of each bin was calculated. Red dots and line show the mean. Black dots and line show the median. The blue cloud indicates one standard deviation of each bin. The horizontal black line denotes zero difference, and the dashed lines indicate EE envelopes. Positive values indicate that MODIS AOD is higher than AERONET.

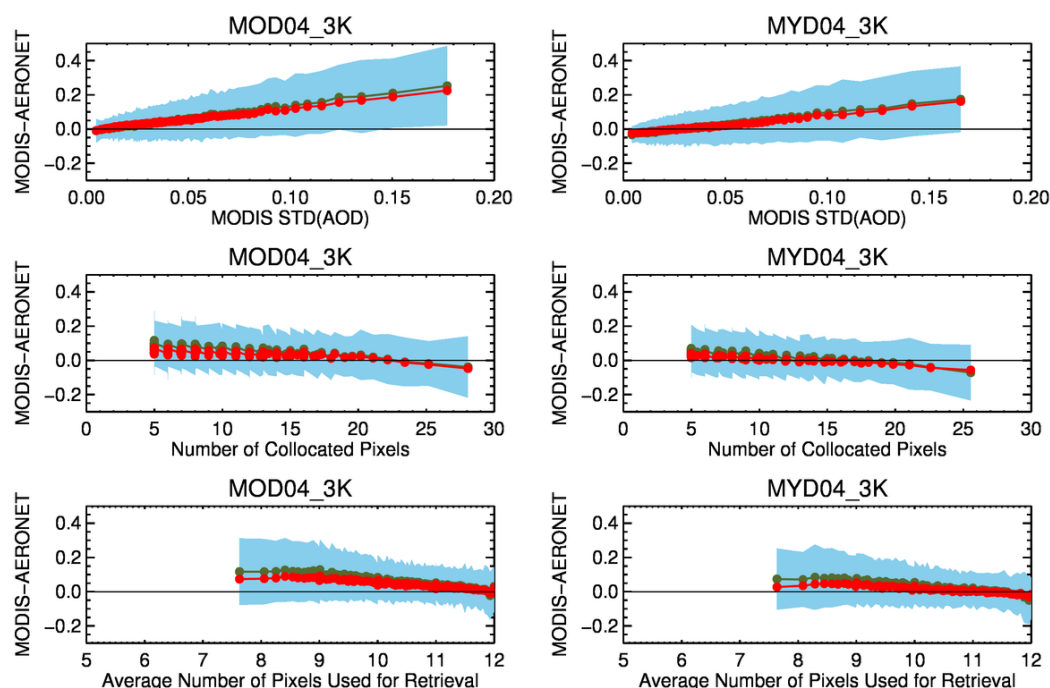


Figure 7. Same as Figure 6 except for standard deviation of MODIS AOD within the 5x5 collocation box (top row), number of MODIS retrievals with the 5x5 collocation box (middle row), and number of MODIS reflectance native pixels used by the retrieval, averaged for all retrievals made in the 5x5 collocation box (bottom row).

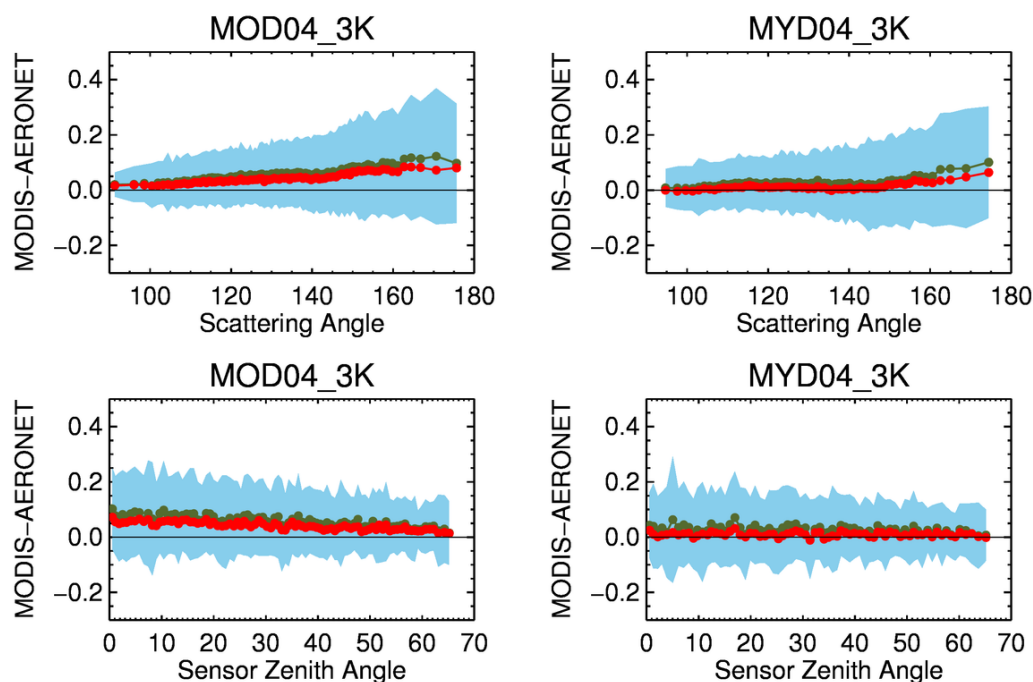
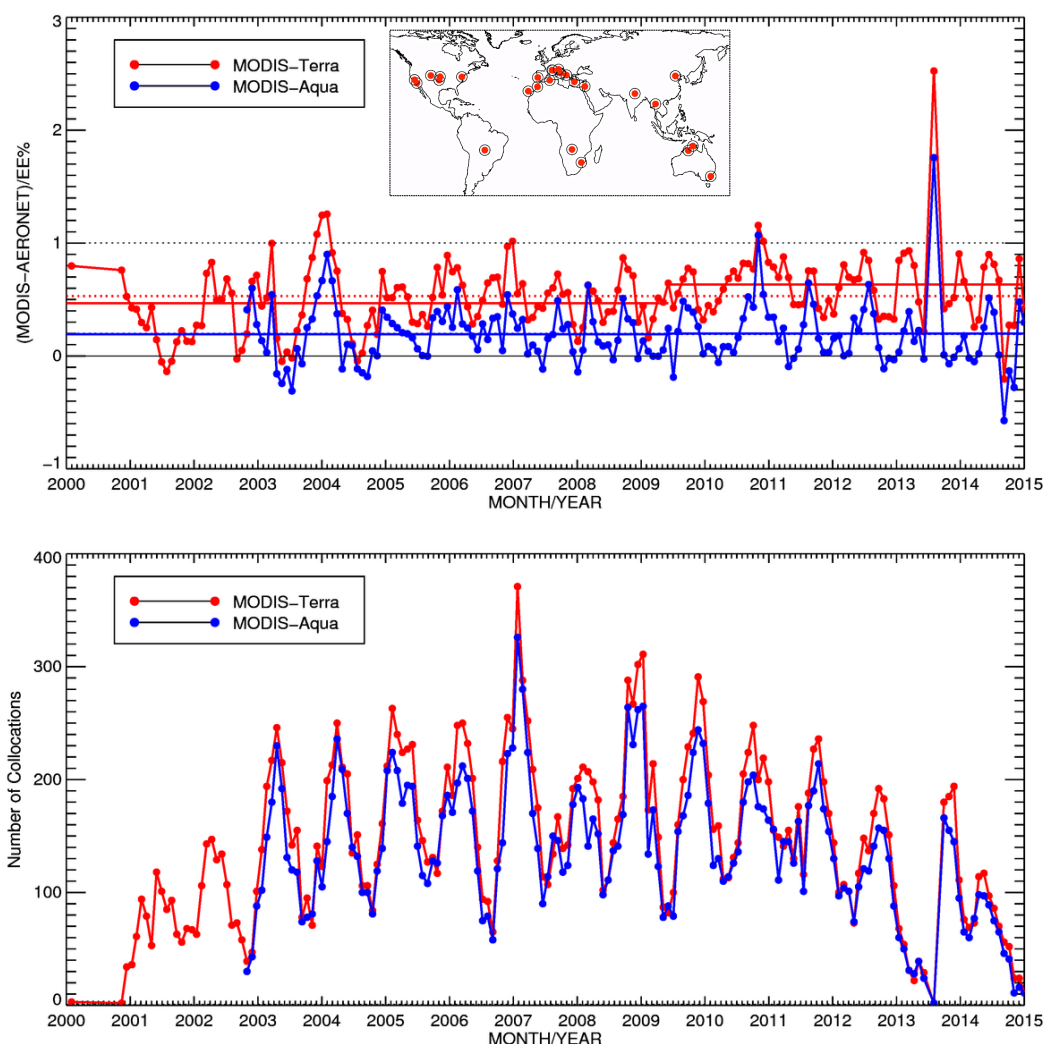


Figure 8. Same as Figure 6 except for scattering angle (top) and sensor zenith angle (bottom).



1
 2
 3 Figure 9. Time series of monthly mean error ratios (Eq. 3) (top), and number of collocations
 4 (bottom) for the global collocation data set from 26 selected long-term AERONET stations. The
 5 Terra record is in red, and Aqua in blue. Note Aqua's record begins two years after Terra, and
 6 the total number of collocations is temporally variant. Only MODIS 3 km retrievals with
 7 QAF=3 are included. The horizontal red and blue lines are temporal means of ER. Dotted red
 8 and blue horizontal lines indicate long-term temporal mean ER for each satellite. Solid red lines
 9 are temporal mean ER calculated for Terra for two periods (2000-2010) and (2011-2015). The
 10 map in the inset shows locations of AERONET stations used in this analysis with more details
 11 provided in table 3