

Dear Editors,

Thanks very much for your valuable time and clear instructions on our manuscript. We have revised our manuscript carefully based on the valuable comments and suggestions from the three reviewers. Please kindly find our point-to-point responses below.

Best regards,

The authors

Anonymous Referee #1

Received and published: 2 June 2018

see my comments in the attached PDF file

Please also note the supplement to this comment:

<https://www.atmos-meas-tech-discuss.net/amt-2018-83/amt-2018-83-RC1-supplement.pdf>

Ps and Tm are required during the conversion process from zenith wet delay (ZWD) to precipitable water vapour (PWV) from GNSS measurements. An alternative method for accurate determination of PWV was proposed for near-real-time applications using GNSS data and nearby synoptic observations, and a method to construct PWV maps with the use of GNSS network was presented. The demand for assimilating zenith total delay (ZTD) into numerical weather prediction (NWP) models is very high in meteorological practice, in particular its role in weather now-casting. The research conducted is of importance to improve weather predictions over the region investigated in HuNan China. Overall I have no serious concerns with respect to the manuscript. Your analysis is sound and you show plenty of results upon which you draw your conclusions. Also the English is very good and generally your statements are clear. However, there are a few minor points I would like to raise (in no particular order):

R. Thanks for your positive conclusions towards our manuscript.

1. Page 3 LINE 2, You stated that “None of them is however collocated with meteorological sensors”? Actually, more than 70 stations have meteorological observations collected by HuNan Meteorological Bureau. You need either remove this statement or make a different statement. In fact the HuNan CORS network does have significant amount of collocated radiosonde measurements since 2015. They are typically easily accessible.

R. Thanks for your useful information. At present, we are not able get the meteorological observations collected by Hunan Meteorological Bureau. In the future, we will try to collaborate with the Hunan Meteorological Bureau to get those data. We thus revised the sentence ‘None of them is however collocated with meteorological sensors’ to ‘However, some stations in the Hunan GNSS network are not collocated with meteorological sensors, thus they cannot be directly used for water vapor monitoring. Except for the Hunan GNSS network, there are many GNSS stations without meteorological observations distributed across the province, which could be included for enhancing the quality of constructed PWV maps in the future’.

2. In Figure 1, it seems RSCZ radiosonde station collocated with one of the CORS stations. Pls compare the CORS derived PWV with RS-derived PWV.

R. Thanks for this comment. Yes, the distance between the RSCZ radiosonde station and the GNSS station is only 266 m, and their height difference is 8 m. In section 4.2, the RSCZ radiosonde-derived PWVs were compared with those interpolated from surrounding 4 GNSS stations including the collocated station using the following equation:

$$y_G = \frac{\sum_{i=1}^n \exp(-d_i^2) \cdot y_i}{\sum_{i=1}^n \exp(-d_i^2)}$$

We can get that the larger the distance, the smaller the weight. Except the collocated GNSS stations, the distances of the other 3 stations to the radiosonde stations all exceed 50 km. Therefore, the weight for the collocated GNSS station is greater than 0.99 in the PWV calculation. For this reason, the Figure 4(b) can be seen the direct PWV comparison between radiosonde and corresponding GNSS station. The RMS error of their PWV differences is 2.94 mm. As you suggested, we also compare the radiosonde PWVs with the collocated GNSS PWVs, the statistical results are basically the same.

Thank you very much again.

3. In page4, line 14, “The blind model global pressure and temperature (GPT)”, this sentence should be changed in a better order.

R. Thanks for your correction. We have revised the sentence ‘The blind model global pressure and temperature (GPT)’ to ‘The global pressure and temperature (GPT) model’.

4. In page5, line 25, where do the formula (5-6) come from? Please show the references and the unit of parameters.

R. These two formulas are from the following journal paper:

Zhang, H., Yuan, Y., Li, W., Ou, J., Li, Y. and Zhang, B.: GPS PPP-derived

precipitable water vapor retrieval based on Tm/Ps from multiple sources of meteorological data sets in China, J. Geophys. Res. Atmospheres, doi:10.1002/2016JD026000, 2017.

Note that the original formula for the P_s in the reference is as follows:

$$P_s = P_r e^{\frac{-gM(H_s - H_r)}{RT_v}}$$

We slightly revise the above formula to $P_s = P_r e^{\mu(H_s - H_r)}$ to refine the modeling of the P_s . In addition, we cited the above paper when describing the formula (5-6) and gave the unit of parameters.

Thank you very much.

5. In section 4.1 please show the formula (4-8) coefficients estimated locally at each synoptic site using reanalysis products.

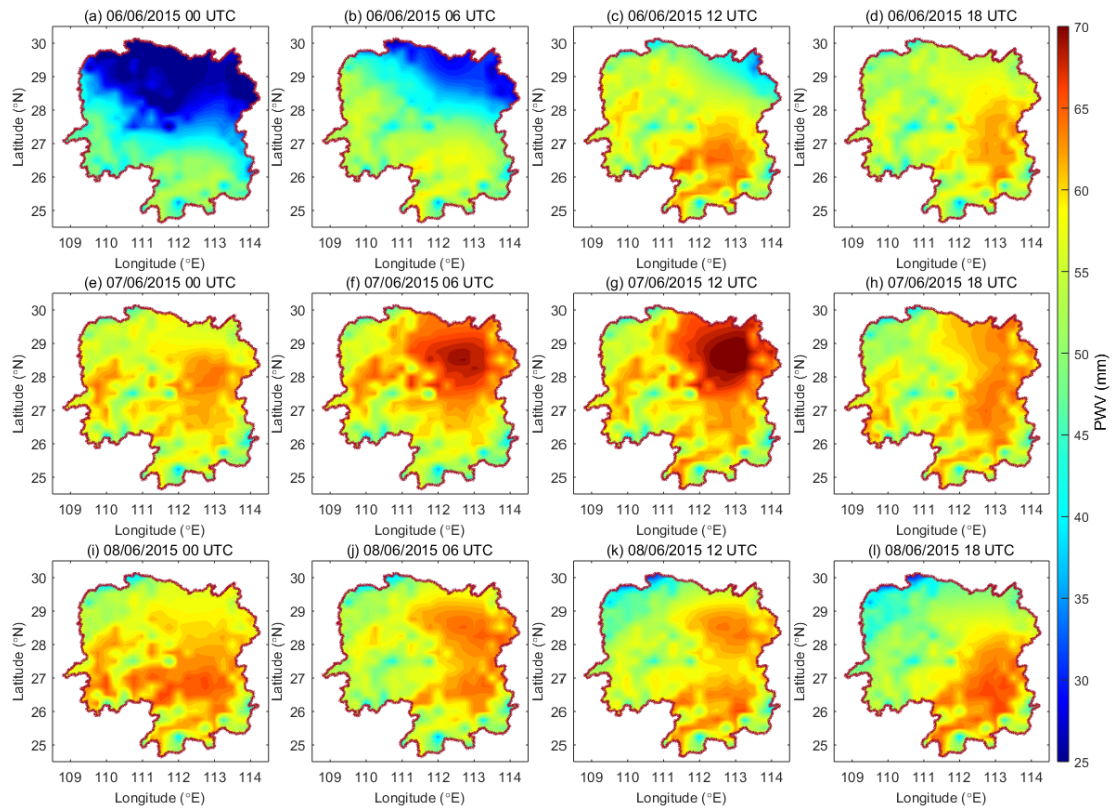
R. Thanks for your comment. The estimated parameters using the ERA-I reanalysis products over the year of 2014 are shown in the following table. We have added this table in section 4.1.

Table 1 Estimated values of a , b , μ and α for the 20 synoptic sites using ERA-I atmospheric profiles over the whole year of 2014

Station	Parameters			
	a	b	μ	α
S01	264.72	0.82	-0.1110	-4.47
S02	264.40	0.83	-0.1112	-4.48
S03	264.90	0.82	-0.1106	-4.25
S04	267.08	0.75	-0.1102	-3.76
S05	265.67	0.79	-0.1111	-4.05
S06	266.46	0.78	-0.1104	-3.90
S07	265.68	0.79	-0.1103	-4.16
S08	266.49	0.77	-0.1108	-3.79
S09	267.32	0.73	-0.1101	-3.88
S10	267.23	0.73	-0.1102	-4.09
S11	269.07	0.67	-0.1097	-3.66
S12	267.99	0.72	-0.1105	-3.66
S13	268.40	0.70	-0.1105	-3.64
S14	268.74	0.65	-0.1074	-4.06
S15	269.02	0.68	-0.1103	-3.69
S16	269.56	0.66	-0.1099	-3.78
S17	269.43	0.66	-0.1102	-3.70
S18	269.52	0.65	-0.1099	-3.96
S19	270.27	0.63	-0.1096	-4.04

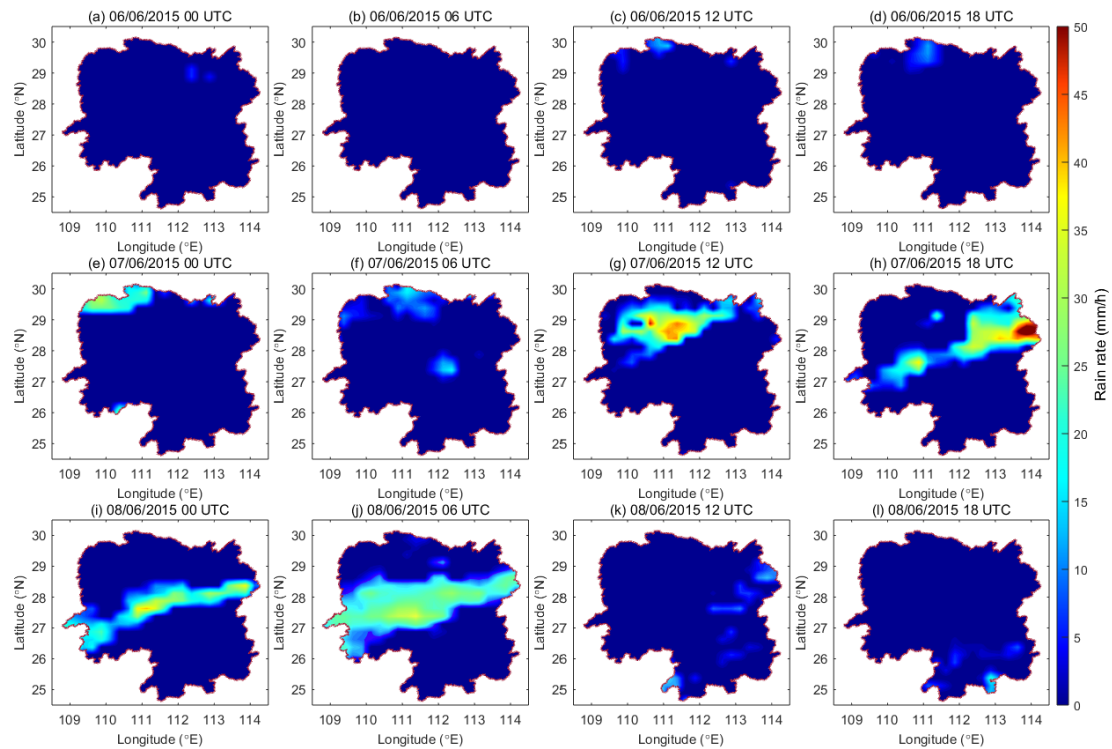
6. All the PWV maps show minor color variations, which is difficult to show the large moisture variations during a rainstorm. I would recommend add 2 more colors between blue and yellow, such as red and purple.

R. Thanks very much for your useful comment. We changed the contour color for the PWV maps. Below shows the revised PWV maps. As you mentioned, it's more clear to show the moisture variations.



7. In Figure 9, the PWV maps should be compared to precipitation every 6 hours. The temporal span of 24h in Figure 10 is too large to explain the situation effectively.

R. Thanks very much for your constructive comments. According to your suggestion, we also show the rain rates every 6 h in the revised manuscript. However, we cannot observe close correlations between the PWV and the rain rate. Larger moisture convergence is not necessarily linked with higher rain rate occurrence. This is because the moisture convergence is not the only cause of precipitation, whilst also controlled by many other factors such as wind, temperature and terrain. However, the GNSS-derived PWV maps are able to reveal the moisture advection, transportation and convergence during the heavy precipitation event.



Evolution of rain rate maps for the Hunan province every 6 h from UTC 00, 6 June 2015 to UTC 18, 8 June 2015.

The rain rate data were retrieved from the TRMM with a spatial resolution of $0.25^{\circ} \times 0.25^{\circ}$.

8. Two recent journal publications to do the same region (and perhaps similar data and data sources are used) in China should be consulted. Their publication details are listed below.

- Li L, Wu S, Wang X, Tian Y, He C, and Zhang K (2018) Modelling of weighted-mean temperature using regional radiosonde observations in Hunan China, *Terr. Atmos. Ocean. Sci.*, Vol.29, No.2, 187-199, doi: 10.3319/TAO.2017.05.26.01.
- Li LI, Suqin Wu, Xiaoming Wang, Ying Tian, Changyong He and Kefei Zhang (2017) Seasonal Multi-Factor Modelling of Weighted-Mean Temperature for Ground-Based GNSS Meteorology in Hunan, China, *Advances in Meteorology*, volume 17, <https://doi.org/10.1155/2017/3782687>.

R. Thanks very much for your information. We have carefully read the above two references. They are very informative and useful. We thus cited them in our revised manuscript.

Thank you again for your invaluable comments. Your constructive comments and suggestions have improved the quality of our manuscript greatly.

Anonymous Referee #2

Received and published: 14 June 2018

Dear Authors, This is a very interesting manuscript which applies well-used methods to the GNSS-derived ZTD from Hunan, China. I have no major comments on the scientific contents. It is well written with minor grammatical and spelling mistakes. However, they are a few of which some I have pointed out in the annotated manuscript I have uploaded (amt-2018-083-supplement.pdf).

R. Thanks for your positive conclusions towards our manuscript.

General comments:

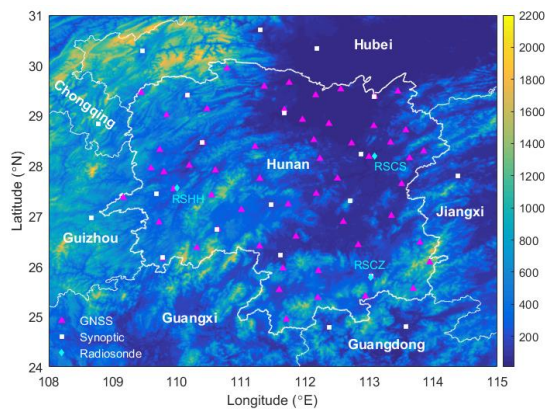
1) You do not mention how the GNSS data have been processed or where the solution is from. Clearly GNSS-derived ZTD are not raw observations and a couple of sentences on this step or a reference pointing to details of the GNSS processing strategy are required.

R. Thanks for your comments. We added several sentences ‘In this study, the ZTDs are estimated using GNSS precise point positioning (PPP) technique with the Bernese 5.2 software (Dach et al., 2015). To examine their performance in real-time applications, IGS (International GNSS Service) ultra-rapid satellite orbit data and clock corrections are adopted in PPP processing. The ZTDs are estimated with an interval of 30 min, whilst the horizontal gradients are estimated every 12 h. The global mapping function (GMF) is used (Boehm et al., 2006) in the estimation, and GNSS observations with elevation angles below 5° are rejected.’ in section 2.1 to introduce the method of estimating ZTD used in this study.

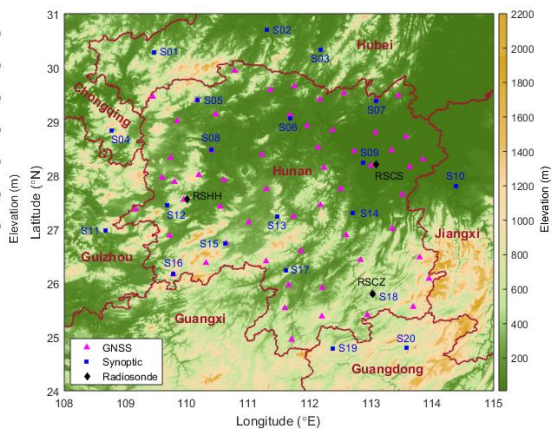
2) Figure 1. Change the colour scheme of this figure to something more commonly used for the presentation of topography. For example, low lying areas should be in green and high areas in brown (green-orange-yellow-red-brown). Also, I did not notice the legend at first. It might be useful for other readers if you box it in and make the background of the legend white.

R. Thanks very much for your invaluable comments. We modified the figure according to your suggestions. Below show the original and modified figures:

Original



Modified



3) You have only employed the GPT2 model. It is well known that this model can only reflect the annual variation in p and t and not the daily fluctuations. For this the values from the VMF1 model (although derived from ECMWF) would be more adequate for the comparison. As there is a VMF1 model for forecasts, you could also employ this in near-real-time.

R. Thanks for your comment. At present, three different types of parameters/models can be provided by the Institute of Geodesy and Geophysics at the Vienna University of Technology. They are 1) VMF1; 2) Empirical Models (like GMF or GPT); and 3) Other Parameters (z_{200} (Height of the 200 hPa pressure level), T_{mean} (Mean temperature), LHG (Linear horizontal gradients), GRAD (Horizontal Gradients GRAD)) (refer to <http://ggosatm.hg.tuwien.ac.at/DELAY/readme.txt>). However, VMF1 does not provide the pressure and weighted mean temperature. Although they provide T_{mean} data, but have a latency of more than 24 h. Therefore, we chose to use GPT2w model in this study. I have also contacted Prof. Boehm (leader of the Institute of Geodesy and Geophysics at the Vienna University of Technology) to consult the availability of forecasted pressure and mean temperature. He replied me that those parameters are not yet provided on global forecast grids (see the screenshot below).

Thanks again!



2018/6/23 (周六) 19:49

Böhm, Johannes <Johannes.Boehm@geo.tuwien.ac.at>

RE: Forecast pressure and Tm

收件人 yeary124@csu.edu.cn

Dear Prof. Chen,

thanks for your email. These parameters are not (yet) provided on global forecast grids.

Best regards, Johannes

From: yeary124@csu.edu.cn <yeary124@csu.edu.cn>

Sent: Freitag, 22. Juni 2018 16:20

To: Böhm, Johannes <Johannes.Boehm@geo.tuwien.ac.at>

Subject: Forecast pressure and Tm

Dear Prof. Boehm,

I am studying the PWV retrieval in near real time.

As we know, the pressure and the weighted mean temperature Tm are required.

Previously, I used the GPT2w model to get those parameters.

I am wondering, besides the ZHD and ZWD, does the VMF1 model provide forecasted surface pressure and Tm?

Looking forward for your reply. Thanks very much.

4) Section 4.2. You mention the wet/dry biases between the PWVs from GNSS and radiosonde data. There are references out there and you should mention that other authors have found similar biases, linking your work more to previously published work. Where do the biases come from?

R. Thanks very much for your insightful comments. When PWV values are less than 10 mm, there is an obvious wet bias relative to the radiosonde. This is probably related to the dry bias of radiosonde sensors caused by solar heating (Moradi et al., 2013). Whereas, an obvious dry bias can be observed for PWV values larger than 65 mm. The dry bias is likely due to the overestimation of water vapor by radiosonde as the humidity sensors suffer contamination from rain and clouds during radiosonde ascents (Bock et al., 2005). As you suggested, we added the above explanations in the modified manuscript.

Bock, O., Keil, C., Richard, E., Flamant, C. and Bouin, M.: Validation of precipitable water from ECMWF model analyses with GPS and radiosonde data during the MAP SOP, Q. J. R. Meteorol. Soc., 131(612), 3013–3036, doi:10.1256/qj.05.27, 2005.

Moradi, I., Soden, B., Ferraro, R., Arkin, P. and Vömel, H.: Assessing the quality of humidity measurements from global operational radiosonde sensors, J. Geophys. Res. Atmospheres, 118(14), 8040–8053, doi:10.1002/jgrd.50589, 2013.

5) Additional reference, place as indicated.

Guerova, G., J. Jones, J. Dousa, G. Dick, S. de Haan, E. Pottiaux, O. Bock, R. Pacione, G. Elgered, H. Vedel and M. Bender (2016).

"Review of the state-of-the-art and future prospects of the ground-based GNSS meteorology in Europe." Atmos. Meas. Tech., 9, 5385-5406, <https://doi.org/10.5194/amt-9-5385-2016>, 2016.

R. Thanks very much for your information. We have carefully read the above paper. It is very informative and useful. We thus cited it in our revised manuscript.

6) minor comment: often when you talk about ECMWF you actually mean the ERA-Interim reanalysis, hence ERA-I would be a better abbreviation.

R. As you suggested, we modified the ECMWF to ERA-I in our manuscript. Thanks very much for your useful comment.

Please also note the supplement to this comment:

<https://www.atmos-meas-tech-discuss.net/amt-2018-83/amt-2018-83-RC2-supplement.pdf>

R. Thanks very much for your detailed grammar corrections. We have revised them accordingly.

Thank you again for your invaluable comments. Your constructive comments and suggestions have improved the quality of our manuscript greatly.

Anonymous Referee #3

Received and published: 30 June 2018

General Comments.

The manuscript describes a method to convert GNSS-derived Zenith Total Delay into Precipitable Water Vapour using, for each GNSS stations, surface pressure and weighted mean temperature of the atmosphere obtained interpolating nearby synoptic observations. The analysis is well presented and the results sound reasonable. However, I would raise the following issues which has to be clarified prior to the publication.

R. Thanks for your positive conclusions towards our manuscript.

1. In the paper, there is no indication on how GNSS data are analyzed: which strategy is applied for estimating ZTD? Which global products are used? What is the ZTD sampling rate? What is the accuracy of the GNSS ZTD estimates? What is the latency of GNSS ZTD estimates?

R. Thanks very much for your constructive comments. In our study, the ZTDs are estimated using GNSS precise point positioning (PPP) technique with the Bernese 5.2 software. To examine their performance in real-time applications, IGS (International GNSS Service) ultra-rapid satellite orbit data and clock corrections are adopted in PPP processing. The ZTDs are estimated with an interval of 30 min, whilst the horizontal gradients are estimated every 12 h. The global mapping function (GMF) is used in the estimation, and GNSS observations with elevation angles below 5° are rejected. Evaluation results show that our estimated ZTDs have an accuracy of 8.8 mm in the comparison by radiosonde measured ones. The presented work is a first-step study to investigate the PWV map construction in near real-time. Therefore, currently the system has not been operated in real-time mode. However, based on our preliminary estimation (58 GNSS stations; data sampling rate of 30s; a common PC (Intel(R) Core(TM) i7-7700 CPU @ 3.60 GHz; RAM 16 GB)), the GNSS ZTD estimates could be provided with a latency of less than 5 min.

We have added some contents in section 2.1 to describe the ZTD estimation strategies used in this study. Thanks again!

2. The authors claim they are presenting a method inferring ‘accurate’ Ps and Tm and for the construction of ‘high-quality’ PWV maps. Both ‘accurate’ and ‘high-quality’ has to be quantified with respect to the target application the authors are interested in. This because the observational requirements are different according to the different target application. I would suggest reviewing the title by adding in it the target application of this research.

R. Thanks for your insightful comment. The ultimate goal of constructing high quality

PWV maps is to enhance the precipitation forecasts and analysis for the Hunan province, China. As you suggested, we revised the title from ‘Constructing Precipitable Water Vapor Map from Regional GNSS Network Observations without Collocated Meteorological Data’ to ‘Constructing Precipitable Water Vapor Map from Regional GNSS Network Observations without Collocated Meteorological Data for Weather Forecasting’.

Below specific comments.

Line 13 pag.1. My suggestion is to replace ‘(GNSS) data’ with ‘(GNSS) Zenith Total Delay (ZTD) estimates’?

Line 25 pag.1 replace ‘ERA reanalysis’ with ‘ERA-Interim reanalysis’?

R. Thanks for your corrections. We have revised them accordingly.

Line 4 pag.2 Suggested reference: Guerova, G., Jones, J., Douša, J., Dick, G., de Haan, S., Pottiaux, E., Bock, O., Pacione, R., Elgered, G., Vedel, H., and Bender, M.: Review of the state of the art and future prospects of the ground-based GNSS meteorology in Europe, Atmos. Meas. Tech., 9, 5385-5406, doi:10.5194/amt-9-5385-2016, 2016.

R. Thanks very much for your information. We have carefully read the above paper. It is very informative and useful. We thus cited it in our revised manuscript as you suggested.

Line 7 pag.2 ‘for a better performance’, please clarify this statement.

R. Thanks for your comment. We modified the ‘for a better performance’ to ‘for a better GNSS positioning performance’. We think it is clear now.

Line 8 pag.2 See general comment. The accuracy of the GNSS ZTD estimates depend on how the data are processed and on the global products used in the processing. For example, the agreement of reprocessed ZTD estimates is at 2 mm level (reference Pacione, R., Araszkievicz, A., Brockmann, E., and Dousa, J.: EPN Repro2: A reference GNSS tropospheric dataset over Europe, Atmos. Meas. Tech., 10, 1689–1705, doi: 10.5194/amt-2016-369, 2017).

R. Thanks very much for your recommended paper. It is very interesting, and we thus cited it in our revised manuscript. We totally agree with that the accuracy of the GNSS ZTD estimates depend on how the data are processed and on the global products used in the processing. We added the sentence ‘The accuracy of the GNSS ZTD estimates depends on the data processing strategies and on the global products

used in the processing' in line 8 of page 2.

Line 24 pag.3 what is the average distance between a synoptic station and a GNSS station?

R. The average distance between a synoptic station and a GNSS station is about 41 km. We have added such statement in the revised manuscript.

Thank you!

Line 26 pag.3 I guess the ZTD sampling rate is higher than 6h, right? If so, how do you interpolate in time pressure and temperature data measured at the synoptic station? What is the error of this interpolation? Such error has to be added in the error analysis done in section 4.1.

R. Thanks very much for your insightful comments. Yes, the ZTD estimates have a time resolution of 30 min in our study. The currently used synoptic data have a time resolution of 6 h. Actually, in the accuracy analysis presented in section 4.1, considering the real-time accessibility the Ps and Tm are not real data measured at the synoptic stations. They are extrapolated by using the empirical models established based on the past 20 days (an optimal number of days after various tests). The empirical models are established using 4-order Fourier function.

Here, we show the time series of extrapolated and measured surface pressure and temperature at a sampling synoptic station in Figure 1. We can see the extrapolated synoptic data in general agree well with the measured ones. Table 1 further gives the statistics of the comparison results. RMS errors of the extrapolated surface pressure and temperature are 2.43 hPa and 2.31 °C, respectively.

Since in section 4.1, the Ps and Tm are derived from the extrapolated data at the synoptic stations, the evaluated errors by radiosonde already include the interpolation errors in time. In section 2.1, we have added more contents to clarify the synoptic data used in the evaluation. In the near future, we will be able to access the hourly synoptic data from the China Meteorological Agency. At that time, more accurate extrapolated surface pressure and temperature can be obtained for our applications.

Thanks again!

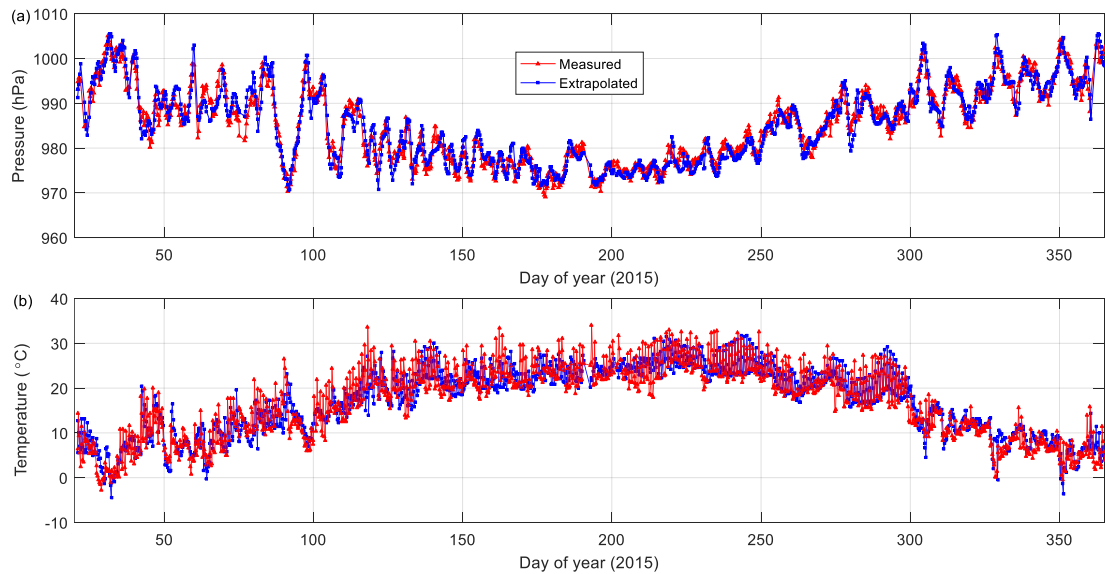


Figure 1 Time series of measured and extrapolated (a) surface pressure and (b) temperature at a sampling synoptic station.

Table 1 Comparison of measured and extrapolated surface pressure and temperature at the 20 synoptic stations

Comparison	Pressure (hPa)	Temperature (°C)
Bias	0.05	0.25
RMS	2.43	2.31
Max	5.64	5.32
Min	-6.52	-4.56

Line 28 pag.3 What kind of radiosonde are used?

R. The radiosonde data are provided by the Integrated Global Radiosonde Archive (IGRA). We thus modified the sentence ‘In addition, atmospheric profiles observed by three radiosonde sites (marked with cyan diamonds in Figure 1) will be used to evaluate the meteorological data and PWV measurements’ to ‘In addition, quality-assured atmospheric profiles observed by three radiosonde sites (marked with black diamonds in 错误!未找到引用源。) from the Integrated Global Radiosonde Archive (IGRA) (Durre et al., 2006) will be used to evaluate the meteorological data and PWV measurements’.

Thanks very much!

Line 12 pag.5 Different set of refractivity coefficients are available in literature, please add the reference about the used ones.

R. Thanks for this suggest. The refractivity coefficients are available by Rüeger (2002). We have added this reference as you indicated.

Rüeger, J. M.: Refractive index formulae for radio waves, in Integration of Techniques and Corrections to Achieve Accurate Engineerin, p. 13, Washington, D.C. USA., 2002.

Line 18 pag.5 The empirical model of eq.4 suffers from diurnal and seasonal biases, are such biases acceptable for the considered application?

R. Thanks for this interesting question. The following table is a part of the Table 2 in the manuscript. We can observe the max and min biases between radiosonde measured and interpolated Tm are 9.69 K and -6.4 K, respectively. If we use an extreme error of 9.69 K for the Tm, in this case the diurnal and seasonal biases are included, the relative error of Tm will be about 3.3%. This will result in a PWV error of about 2.64 mm in for an extreme PWV value of 80 mm. Such error is still less than the accuracy threshold of 3 mm for the application of weather nowcasting. Therefore, we think that the diurnal and seasonal biases in Tm model are acceptable.

Comparison of P_s and T_m for Radiosonde-Synoptic at the three radiosonde stations

Comparison		RSCS		RSCZ		RSHH	
		Ps (hPa)	Tm (K)	Ps (hPa)	Tm (K)	Ps (hPa)	Tm (K)
Radiosonde vs Synoptic	Bias	2.91	1.47	-1.66	1.14	-2.58	1.49
	RMS	2.97	2.92	1.74	2.58	2.61	2.76
	Max	5.04	9.69	0.48	7.42	-1.33	9.17
	Min	1.01	-6.40	-3.82	-5.18	-3.82	-5.40

Line 22 pag. 5 Could the authors explain on which ground they chose 100 km as the radius of the circumference centred on the GNSS site? On average, how many synoptic stations fall into that area for each GNSS sites?

R. Thanks for this question. The reason why we chose 100 km as the radius is to make sure at least one synoptic site being located within the circumference centred on the GNSS station. For each GNSS site, on average, two synoptic stations fall into that circumference. We have added the above clarifications in the revised manuscript.

Line 7 pag.6. Considering eq.7, what is the interpolation error?

R. Thanks for this comment. Actually, we have shown the interpolation errors for Ps and Tm in section 4.1. The interpolated Ps and Tm data were evaluated by three radiosonde sites' data over the whole year of 2015. RMS errors of Ps and Tm derived from synoptic interpolation vary in the range of 1.7-3.0 hPa and 2.5-3.0 K,

respectively. The obtained RMS errors include both the interpolation errors in space and time.

Line 10 pag. 7 Why in this error analysis the authors are not considering the ZTD error? The ZTD error is of course the same in both models the authors are evaluating but I think has to be considered in the total error budget.

R. Thanks for this comment. Following your suggestions, we have considered the ZTD error in the error analysis in section 4.1. Our comparison with radiosonde shows that the estimated ZTDs have an accuracy of about 9 mm (~ 1.45 mm in PWV). By taking into account the ZTD error, the accuracy of PWV retrieved from GNSS-ZTD using P_s and T_m from synoptic interpolation is better than 3.4 mm. The uncertainty of PWV caused by GPT2w model is about 4.6 mm. However, in section 4.2, the actual accuracies of interpolated PWVs by radiosonde are still better than 3 mm. Thus, the GNSS-derived PWVs still meet the accuracy requirement for the application of weather nowcasting.

Line 1 pag. 10 Replace ‘measured’ with ‘estimated’

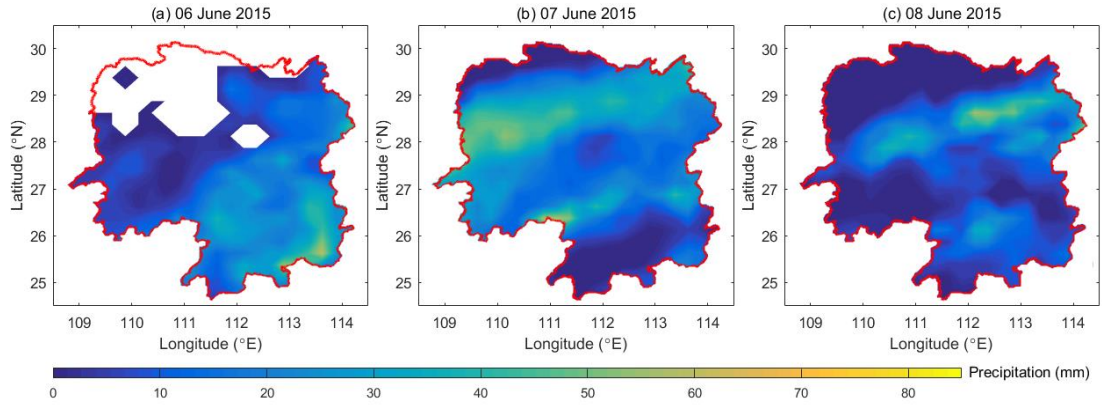
R. Thanks for this comment. We have revised the ‘GNSS measured ZTDs’ to ‘GNSS-derived ZTDs’.

In the manuscript several times, ECMWF should be replaced with ERA-Interim. The quality of the maps should be improved. Fig. 8a check the white spot

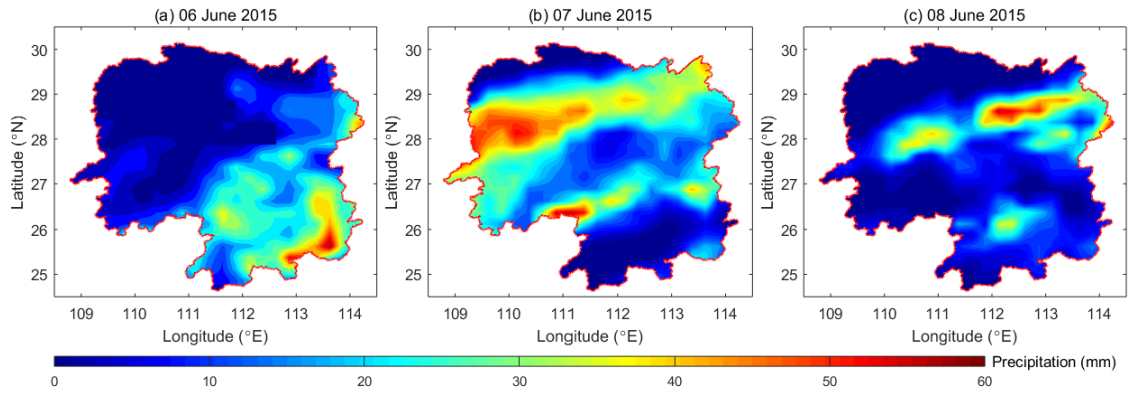
R. Thank very much for your invaluable comments. We have revised the ECMWF accordingly. For a sake of simplicity, we abbreviated the ‘ERA-Interim’ to ‘ERA-Interim’.

There are some mistakes in the original Figure 8. Thanks for your correction. We replotted the Figure 8. Below shows the comparison between original and modified Figure 8. In addition, we also replotted Figure 1, Figure 6, Figure 9 and Figure 11 to improve their quality.

Original Figure 8



Modified Figure 8



Thank you again for your invaluable comments. Your constructive comments and suggestions have improved the quality of our manuscript greatly.

Constructing Precipitable Water Vapor Map from Regional GNSS Network Observations without Collocated Meteorological Data for Weather Forecasting

Biyan Chen^{1,2,3}, Wujiao Dai^{1,2,3}, Zhizhao Liu⁴, Lixin Wu^{1,3}, Cuilin Kuang^{1,2,3}, Minsi Ao⁵

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²Key Laboratory of Precise Engineering Surveying & Deformation Disaster Monitoring of Hunan Province, Changsha, Hunan, China

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Abstract. Surface pressure (P_s) and weighted mean temperature (T_m) are two necessary variables for the accurate retrieval of precipitable water vapor (PWV) from global navigation satellite system (GNSS) zenith total delay (ZTD) estimates. The lack of P_s or T_m information is a concern for those GNSS sites that are not collocated with meteorological sensors. This paper investigates an alternative method of inferring accurate P_s and T_m at the GNSS station using nearby synoptic observations. P_s and T_m obtained at the nearby synoptic sites are interpolated onto the location of GNSS station by performing both vertical and horizontal adjustments, in which the parameters involved in P_s and T_m calculation are estimated from ERA-Interim reanalysis profiles. In addition, we present a method of constructing high quality PWV maps through vertical reduction and horizontal interpolation of the retrieved GNSS PWVs. To evaluate the performances of the P_s and T_m retrieval, and the PWV map construction, GNSS data collected from 58 stations of the Hunan GNSS network and synoptic observations from 20 nearby sites in 2015 were processed to extract the PWV so as to subsequently generate the PWV maps. The retrieved P_s and T_m , and constructed PWV maps were assessed by the results derived from radiosonde and the ERA-Interim reanalysis. The results show that (1) accuracies of P_s and T_m derived by synoptic interpolation are within the range of 1.7-3.0 hPa and 2.5-3.0 K, respectively, which are much better than the GPT2w model; (2) the constructed PWV maps have good agreements with radiosonde and ERA-Interim reanalysis data with the overall accuracy being better than 3 mm; and (3) PWV maps can well reveal the moisture advection, transportation and convergence during heavy rainfall.

1 Introduction

Water vapor is an important meteorological parameter, which plays a crucial role in the formation of various weather phenomenon such as cloud, rain and snow (Ahrens and Samson, 2011). Water vapor accounts for only 0.1-3% of the total atmosphere, however due to the latent heat release, a small amount of water vapor may cause severe weather changes

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(Mohanakumar, 2008). The monitoring of atmospheric water vapor variation is thus of significant value for short-term severe weather forecasting (Brenot et al., 2013; Labbouz et al., 2013; Van Baelen et al., 2011; Zhang et al., 2015). Among the various atmosphere sensing techniques, GNSS (Global Navigation Satellite System (GNSS)) is regarded as a uniquely powerful means to estimating the water vapor with advantages of all-weather capability, high accuracy, and low-operating expense (Bevis et al., 1992; Guerova et al., 2016; Yao et al., 2017)(Bevis et al., 1992; Yao et al., 2017).

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While GNSS signals are transmitted from satellites to ground receivers, they are delayed by the terrestrial troposphere. In GNSS data processing, the tropospheric delay is usually expressed as the zenith tropospheric delay (ZTD) multiplied by a mapping function, and sometimes plus horizontal gradients for a better GNSS positioning performance (Lu et al., 2016). The accuracy of the GNSS ZTD estimates depends on the data processing strategies and on the global products used in the processing. At present, ZTDs are likely to be determined with accuracies up to several millimeters by a wide range of GNSS processing software programs (Pacione et al., 2017; Yuan et al., 2014)(Yuan et al., 2014). ZTD is normally divided into two parts: the zenith hydrostatic delay (ZHD) which is caused by the dry gases of the troposphere and the zenith wet delay (ZWD) by which stems from the water vapor. The ZHD can be accurately calculated using empirical models with surface pressure (P_s) measured by meteorological sensors (Saastamoinen, 1972). ZWD is readily obtained with the subtraction of ZHD from ZTD. The precipitable water vapor (PWV) can then be retrieved from ZWD with a conversion factor which is a function of the weighted mean temperature (T_m). T_m can be calculated by numerical integration from the vertical profiles of atmospheric temperature and humidity (Davis et al., 1985). PWV is a key parameter in studying water vapor variations during severe weather phenomena, since it can reflect the inflow and outflow of water vapor in a vertical air column above a certain area (Yao et al., 2017).

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As stated above, the retrieval of PWV from GNSS-ZTD needs two key meteorological parameters P_s and T_m . The first choice is to measure the P_s by barometer collocated at the GNSS station. However, a large number of GNSS stations have been deployed for positioning purposes and not equipped with collocated meteorological sensors. In this case, one may use pressure derived from a global atmospheric reanalysis (Dee et al., 2011; Zhang et al., 2017) or interpolated from nearby meteorological observations (Alshawaf et al., 2015; Musa et al., 2011; Wang et al., 2007) or predicted by a blind model (Böhm et al., 2015; Wang et al., 2017). For T_m , since the temperature and humidity profiles are very difficult to obtain, particularly in a near-real-time mode, T_m has to be calculated from a model. An empirical T_m model dependent on surface temperature (T_s) (Bevis et al., 1994; Li et al., 2018)(Bevis et al., 1994) or a blind model developed from atmospheric reanalysis products (Böhm et al., 2015; Yao et al., 2013; Zhang et al., 2017) is are often employed.

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The work presented in this paper is carried out for constructing high quality PWV maps by a regional GNSS network in the Hunan Province, China for precipitation forecasts and analysis. The constructed high quality PWV maps will also be of significant values for monitoring and early warning of geological disasters, such as landslides and mud-rock flows. In such a near-real-time application, the use of reanalysis products is not feasible. P_s and T_m have to be determined only using a blind model or nearby surface synoptic stations. The use of blind models is a very convenient means, however, most blind models (e.g. Global Pressure and Temperature 2 wet, GPT2w; (Böhm et al., 2015)) are developed at a global scale and are not likely

to capture regional small-scale variations. More accurate P_s and T_m could be achieved by interpolation from nearby meteorological observations if they can be accessed simultaneously. In this study, we investigate the construction of PWV maps from GNSS observations over the Hunan Province by performing the following five tasks: (1) $T_m - T_s$ relationship and vertical reduction models for P_s and T_m are developed for each synoptic station; (2) P_s and T_m data interpolated by nearby meteorological observations are compared with those derived from radiosonde and GPT2w models; (3) PWV vertical reduction model is developed for each GNSS station; (4) PWV interpolation is performed over the whole Hunan region and evaluated by radiosonde and European Centre for Medium Range Weather Forecasts Reanalysis (ECMWF) ERA-Interim reanalysis (hereafter short as ERA-I); and (5) the water vapor variation during a heavy rain event that occurred over a wide range of Hunan is examined based on PWV maps.

This paper is organized as follows. Section 2 presents the study area and the datasets used in the study. Section 3 describes the methodology to retrieve PWV from GNSS data. The strategy for meteorological data interpolation, T_m modeling and PWV interpolation is also presented in this section. The assessment of P_s and T_m interpolated by nearby synoptic observations is described in section 4. The PWV maps constructed by GNSS data and PWV evolution during a heavy rain event ~~is~~ are also presented in section 4. The summary and conclusions are given in section 5.

2 Study area and data description

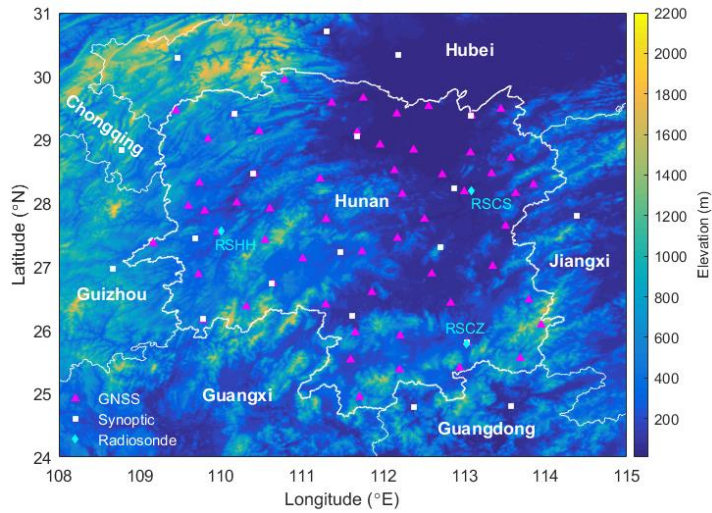
The Hunan Province is located in the middle reaches of the Yangtze watershed in ~~the~~ South Central China, with a territory of about 211,800 km². Hunan enjoys a subtropical humid monsoon climate bearing obvious continental climate features ~~of continental climate~~. The average annual rainfall varies between 1200-1700 mm, with 50%-60% concentrating in the months from April to August. Heavy showers and thunderstorms frequently occur in summer, causing catastrophic casualties ~~conditions~~ as well as significant damages to urban infrastructure and agricultural production. The monitoring of water vapor variations using the GNSS network has a great potential to improve the capacity of extreme weather forecasting in the Hunan region.

2.1 GNSS, synoptic and radiosonde stations in Hunan

In 2015, 58 GNSS stations were deployed in the Hunan GNSS network and new stations ~~are continually~~ have subsequently ~~been~~ added (see Figure 1 ~~Figure 4~~). At present, the GNSS network consists of more than 90 stations and the number is still ~~in~~ increasing (Li et al., 2017). Most of the GNSS stations are equipped with Trimble or Leica receivers and have a typical sampling interval of 30 s. In this study, the ZTDs are estimated using GNSS precise point positioning (PPP) technique with the Bernese 5.2 software (Dach et al., 2015). To examine their performance in real-time applications, IGS (International GNSS Service) ultra-rapid satellite orbit data and clock corrections are adopted in PPP processing. The ZTDs are estimated with an interval of 30 min, whilst the horizontal gradients are estimated every 12 h. The global mapping function (GMF) is used (Boehm et al., 2006) in the estimation, and GNSS observations with elevation angles below 5° are rejected. Evaluation

results show that our estimated ZTDs have an accuracy of ~9 mm in the comparison by radiosonde measured ones. ~~None of them is however~~However, ~~some stations in the Hunan GNSS network are not~~collocated with meteorological sensors, thus they cannot be directly used for water vapor monitoring. ~~Except the Hunan GNSS network, there are many GNSS stations without meteorological observations distributed across the province, which could be included for enhancing the quality of constructed PWV maps in the future.~~ Therefore, a strategy of using nearby synoptic observations is needed to acquire the necessary meteorological parameters for GNSS-PWV retrieval. As shown in ~~Figure 1~~Figure 1, a total of 20 synoptic sites situated in Hunan and surrounding provinces can be used for this study. ~~The average distance between a synoptic station and a GNSS station is about 41 km.~~ The 6-hourly pressure and temperature data measured at the synoptic sites can be retrieved from the National Center for Atmospheric Research (NCAR) (<http://rda.ucar.edu/datasets/ds336.0/>). ~~For real-time applications, pressure and temperature data at a given epoch are extrapolated from empirical models established using the past 20-day data. Here, the 4-order Fourier function is adopted for the empirical models. In addition, quality-assured atmospheric profiles observed by three radiosonde sites (marked with cyan-black diamonds in Figure 1) from the Integrated Global Radiosonde Archive (IGRA) (Durre et al., 2006) will be used to evaluate the meteorological data and PWV measurements.~~

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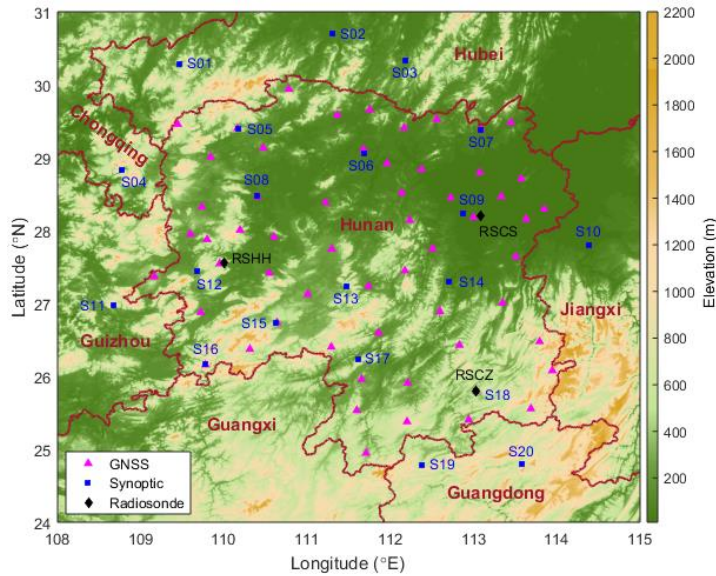


Figure 1: Geographic distribution of GNSS, synoptic and radiosonde stations in Hunan and surrounding provinces.

2.2 ECMWF reanalysis

ECMWF ERA-Interim is a global atmospheric reanalysis from 1979, continuously updated in near real time. In the reanalysis data generation, meteorological observations from in situ platforms (e.g., surface weather stations, ships, buoys, radiosonde stations, and aircraft) and remote sensing satellites are assimilated into atmospheric physical models to recreate the past atmospheric conditions (Dee et al., 2011). Due to its high quality and global coverage, the ERA-Interim reanalysis has been exploited in various fields, e.g. GNSS meteorology (Wang et al., 2017; Zhang et al., 2017) and climate change research (Chen and Liu, 2016b; Lu et al., 2015). The ERA-Interim reanalysis provides pressure, temperature, humidity and many other meteorological variables at 37 isobaric levels from 1000 hPa to 1 hPa with a 6 h interval. The reanalysis contains grid products with 11 different scales from 0.125°×0.125° to 3°×3°, and the horizontal resolution of 0.25°×0.25° is selected for this study, which equates to about 26 km in Hunan.

2.3 GPT2w model

The ~~blind-model~~ global pressure and temperature (GPT) model, which is developed using spherical harmonics (Boehm et al., 2007), can provide pressure and temperature at any site in the vicinity of the Earth's surface. Lagler et al., (2013)

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significantly improved the GPT model, especially for its spatial and temporal variability, and named this new version as GPT2. An extension version called GPT2w was developed by Böhme et al., (2015) with improved capability to determine ZWD in blind mode. Besides the pressure and temperature, the refined GPT2w model also provides various parameters such as water vapor pressure, weighted mean temperature and the temperature lapse rate.

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5 3 PWV map construction with GNSS network observations

3.1 Retrieval of PWV from GNSS-ZTD

To retrieve PWV from GNSS inferred ZTD, ZHD should be determined first-determined. The ZHD calculation formula is theoretically derived based on the assumption that the air is an ideal gas and that the troposphere satisfies the hydrostatic equilibrium (Davis et al., 1985)(Davis et al., 1985). Saastamoinen (1972) derived the most widely used ZHD model as follows (Chen and Liu, 2016a)(Chen and Liu, 2016b):

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$$ZHD = 2.2793 P_s / (1 - 0.0026 \cos 2\varphi - 0.00028h), \quad (1)$$

where φ is the station latitude (unit: radians) and h is the height of the station above sea level (unit: km). By subtracting ZHD from ZTD, the remainder ZWD can then be converted to PWV by using the formula below (Askne and Nordius, 1987):

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$$PWV = \frac{10^5}{(k_3/T_m + k'_2)R_v} ZWD, \quad (2)$$

15 where $k_3 = 3.776 \times 10^5 \text{ K}^2/\text{hPa}$, $k'_2 = 16.52 \text{ K/hPa}$, and $R_v = 461.495 \text{ J/K/kg}$ are physical constants (Rüeger, 2002). The weighted mean temperature T_m is defined as (Davis et al., 1985)(Davis et al., 1985):

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$$T_m = \frac{\int \frac{e(h)}{T(h)} dh}{\int \frac{e(h)}{T(h)^2} dh}, \quad (3)$$

where $e(h)$ and $T(h)$ are the water vapor pressure (hPa) and temperature (K) at height h , respectively. Since the humidity and temperature profiles are usually unavailable, a linear relationship between surface temperature T_s and T_m is often adopted to determine the T_m :

20
$$T_m = a + bT_s, \quad (4)$$

where a and b are coefficients that need to be fitted locally using radiosonde or reanalysis profiles.

3.2 Spatial adjustments for P_s and T_m

25 Because no-some stations in the Hunan GNSS network are not equipped with meteorological sensors, a method of spatially adjusting nearby meteorological observations to the GNSS stations is-was developed. Adjacent synoptic sites within the 100 km radius of a given GNSS station are employed in the adjustments. This ensures at least one synoptic site being located

within the circumference centred on the GNSS station. For each GNSS site, on average, two synoptic stations fall into that circumference. First, surface pressure and mean weighted temperature data at the synoptic sites are adjusted to the height H_s of the given GNSS station (Zhang et al., 2017):

$$P_s = P_r e^{\mu(H_s - H_r)}, \quad (5)$$

$$5 \quad T_m = T_{mr} + \alpha(H_s - H_r), \quad (6)$$

where P_r , T_{mr} , and H_r are the pressure (hPa), weighted mean temperature (K), and height (km) at the synoptic site, respectively. Here, T_{mr} is calculated by equation (4) using the surface temperature (K) measured on site. P_s and T_m are the pressure and weighted mean temperature corresponding to the height H_s at the synoptic site. μ and α are parameters needed to be estimated at the synoptic site.

10 Then the vertically adjusted meteorological data are interpolated to the location of the GNSS station according to:

$$y_G = \frac{\sum_{i=1}^n \exp(-d_i^2) y_i}{\sum_{i=1}^n \exp(-d_i^2)}, \quad (7)$$

where n is the number of synoptic sites with a distance less than 100 km to the given GNSS site; y_G is the interpolated value; y_i is the adjusted meteorological data at synoptic site i ; and d_i is the distance between synoptic site i and the GNSS station.

3.3 PWV interpolation from GNSS stations

15 With the use of interpolated P_s and T_m , PWV data at the GNSS stations could be obtained in near-real-time. In order to construct the PWV map, GNSS PWV data are used to interpolate at a $0.25^\circ \times 0.25^\circ$ grid. Similar to the meteorological data, PWVs at nearby GNSS stations are interpolated to the given height H_p of the grid point as follows (Dousa and Elias, 2014):

$$\text{PWV} = \text{PWV}_r \left[1 - \frac{\beta(H_s - H_r)}{T_s} \right]^{\frac{\theta \cdot g}{\beta \cdot R_d}}, \quad (8)$$

20 where PWV_r is the PWV estimated at the GNSS station; β refers to the temperature lapse rate (unit: K/km); θ a numerical coefficient; g is gravity acceleration (unit: $m \cdot s^{-2}$); and $R_d = 287.053 J \cdot K^{-1} \cdot kg^{-1}$ is the gas constant for dry air. Both β and θ are required to be determined from local observations for a better performance. The PWV at the grid point can then be acquired by interpolation using equation (7). In this study, the height of each grid point is derived from the global topography/bathymetry grid that has a 30-arc second resolution (SRTM30 PLUS) (Becker et al., 2009).

4.1 Evaluation of P_s and T_m interpolated by synoptic data

5 **ECMWF-ERA-I** atmospheric profiles over the whole year of 2014. With the use of the estimated parameters, spatial adjustments for P_s and T_m to radiosonde stations are performed throughout the year of 2015. Then the interpolated meteorological data are directly compared with the radiosonde observations.

Table 1 Estimated values of a , b , μ and α for the 20 synoptic sites using ERA-I atmospheric profiles over the whole year of 2014[illegible]

Figure 2 shows the time series of P_s provided by radiosonde, synoptic interpolation and GPT2w model at three radiosonde stations over 2015. Surface pressures interpolated from synoptic observations have a very good agreement with radiosonde measured ones. The GPT2w model basically reflects the variation trend of P_s throughout the year, however, it is unable to capture the fluctuations which are especially obvious in winter and spring months. Similar results can be observed in Figure 3 for T_m comparison. Detailed statistics of the comparison results are given in Table 2. RMS (root mean squares) errors of P_s and T_m derived from synoptic interpolation vary in the range of 1.7-3.0 hPa and 2.5-3.0 K, respectively. In comparison, the RMS errors from the GPT2w model are 4.7-5.6 hPa and 3.8-4.2 K, respectively, for P_s and T_m , which are much larger than the synoptic interpolation method. In terms of maximum and minimum differences, GPT2w derived values are significantly larger than those derived from synoptic interpolation, further indicating the GPT2w model is less accurate. According to equation (1), 1 hPa error in surface pressure would cause about 2.3 mm error in ZHD. Therefore, 3 hPa error in P_s will result in an error of about 6.9 mm in ZHD (~1.15 mm in PWV). In addition, the relative error of the PWV caused by the T_m error is approximately equal to the relative error of the T_m (Zhang et al., 2017). Derived from Figure 3 and Table 2, the relative error of synoptic data interpolated T_m is about 1%. In the study region, the PWV value is usually less than 80 mm, meaning the PWV error caused by T_m error is within 0.8 mm. In addition, as mentioned in section 2.1, our estimated ZTDs have an accuracy of about 9 mm (~1.45 mm in PWV). On the whole, the accuracy of PWV retrieved from GNSS-ZTD using P_s and T_m from synoptic interpolation is better than 2-3.4 mm. Following the same error analysis, the uncertainty of PWV caused by GPT2w model is about 4.6-5.4 mm. It is notable that for weather nowcasting, the accuracy threshold is 3 mm (Yuan et al., 2014). This indicates that the PWVs retrieved from synoptic interpolation are accurate enough for weather nowcasting, while the PWV from GPT2w model fails to meet this requirement.

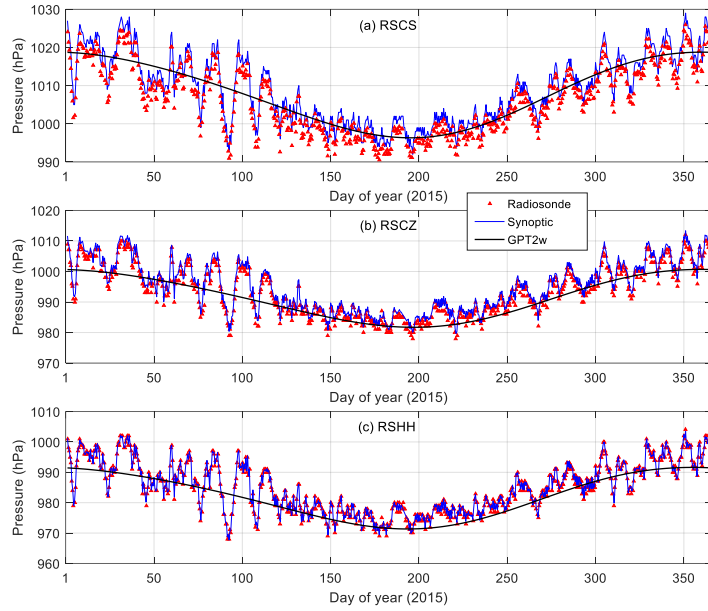


Figure 2: Time series of surface pressure provided by radiosonde, synoptic adjustment and GPT2w model over the whole year of 2015 at three radiosonde stations: (a) RSCS, (b) RSCZ, and (c) RSHH, all of which are located in the Hunan Province, China.

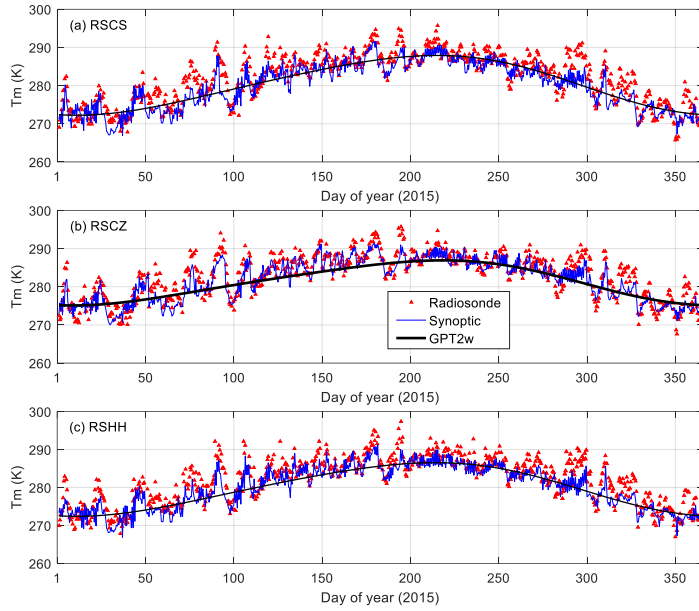


Figure 3: Time series of weighted mean temperature provided by radiosonde, synoptic adjustment and GPT2w model over the whole year of 2015 at three radiosonde stations: (a) RSCS, (b) RSCZ, and (c) RSHH, all of which are located in the Hunan Province, China.

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Table 2: Comparison of P_s and T_m for Radiosonde-Synoptic and Radiosonde-GPT2w at the three radiosonde stations

Comparison		RSCS		RSCZ		RSHH	
		Ps (hPa)	Tm (K)	Ps (hPa)	Tm (K)	Ps (hPa)	Tm (K)
Radiosonde vs Synoptic	Bias	2.91	1.47	-1.66	1.14	-2.58	1.49
	RMS	2.97	2.92	1.74	2.58	2.61	2.76
	Max	5.04	9.69	0.48	7.42	-1.33	9.17
	Min	1.01	-6.40	-3.82	-5.18	-3.82	-5.40
Radiosonde vs GPT2w	Bias	1.23	1.59	2.02	1.68	3.06	2.23
	RMS	4.70	3.84	4.76	4.02	5.56	4.16
	Max	13.75	13.29	12.26	14.21	14.96	14.48
	Min	-16.13	-7.44	-13.46	-8.08	-14.94	-6.34

4.2 Evaluation of PWV by radiosonde

At each GNSS station, GNSS-derived -measured-ZTDs are converted to PWVs with the use of meteorological parameters interpolated from synoptic data. In order to evaluate the GNSS-GNSS-derived PWV, the GNSS PWVs are interpolated onto the radiosonde stations according to equations (7) and (8) for a direct comparison with radiosonde measured ones. As displayed in Figure 4, GNSS interpolated PWVs agree well with the radiosonde measured ones at all the three radiosonde stations. Mean biases of the PWV differences at RSCS, RSCZ and RSHH station are -0.59 mm, 1.04 mm and 1.40 mm, respectively (see Table 3). In terms of the RMS error, they are 2.47 mm, 2.94 mm and 2.69 mm for RSCS, RSCZ and RSHH stations, respectively. The accuracy of GNSS derived PWV is better than 3 mm, which is good enough for the application of weather nowcasting. It is notable that for weather nowcasting, the accuracy threshold is 3 mm (Yuan et al., 2014). This indicates that the GNSS-derived PWVs are accurate enough for the application of weather nowcasting. Additionally, the probability density function (PDF) of PWV differences and the fractional error as percent by radiosonde 5 mm PWV bins are exhibited in Figure 5. As shown in Figure 5(a), about 83% PWV differences are within the range of -5~5 mm. The fractional errors vary from about -15% to 6% as radiosonde PWV increases from 0 mm to 75 mm. When PWV values are less than 10 mm, there is an obvious wet bias relative to the radiosonde. This is probably related to the dry bias of radiosonde sensors caused by solar heating (Moradi et al., 2013). Whereas, an obvious dry bias can be observed for PWV values larger than 65 mm. The dry bias is likely due to the overestimation of water vapor by radiosonde as the humidity sensors suffer contamination from rain and clouds during radiosonde ascents (Bock et al., 2005).

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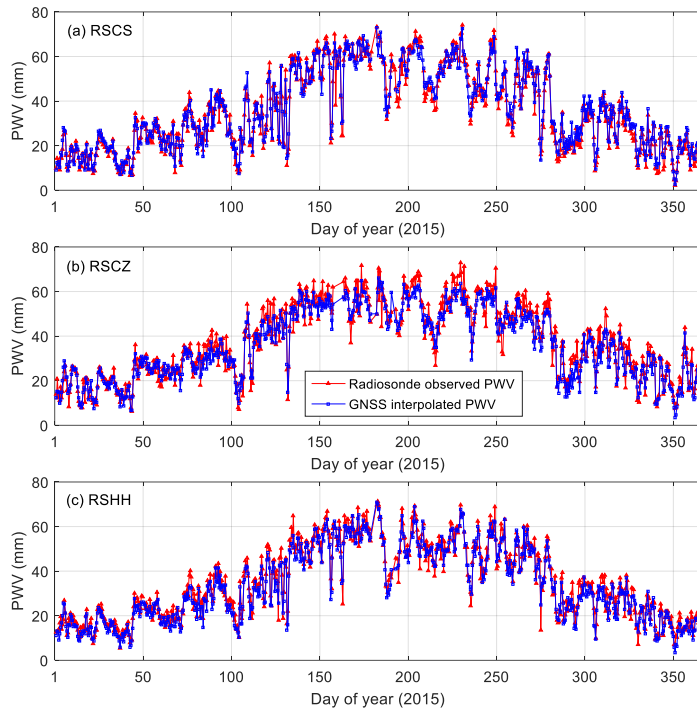


Figure 4: Time series of PWV measured by radiosonde and interpolated by GNSS over the whole year of 2015 at three radiosonde stations: (a) RSCS, (b) RSCZ, and (c) RSHH.

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Table 3: Comparison between radiosonde observed and GNSS interpolated PWV at the three radiosonde stations

Radiosonde station	Bias (mm)	RMS (mm)	Max (mm)	Min (mm)
RSCS	-0.59	2.47	7.46	-7.96
RSCZ	1.04	2.94	10.44	-11.15
RSHH	1.40	2.69	9.27	-8.59

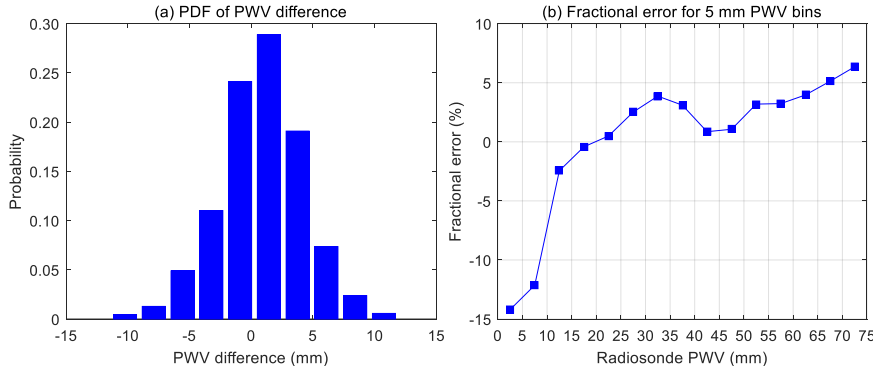


Figure 5: (a) PDF of PWV difference and (b) fractional error as percent by radiosonde 5 mm PWV bins. All the three radiosonde stations are used in the statistics.

4.3 PWV comparison between ERA-IECMWF reanalysis products and GNSS interpolation data

The ERA-IECMWF reanalysis products are used to further assess the performance of PWV maps constructed by the GNSS network data. For the comparison, the GNSS PWVs are interpolated onto grid points with a spatial resolution of 0.25° in both latitude and longitude directions to match the ERA-IECMWF-PWV data. Figure 6 presents the spatial distribution of the bias and RMS error of the PWV differences between ERA-IECMWF and GNSS over the Hunan region. As seen in Figure 6(a), the bias varies from -8 mm to 6 mm depending upon the location. In general, mountainous regions have a larger bias than plain regions. In terms of RMS error, as shown in Figure 6(b), its values vary in the range of 2–8 mm. Large parts of the studied region are populated with RMS errors less than 3 mm. However, relatively large RMS errors of more than 6 mm are obtained for some mountainous regions.

In addition, the PDF of the PWV differences shown in Figure 7(a) indicates that there is a higher probability of negative PWV difference. Negative values account for about 64% of the total PWV difference. The fractional error as percent by ERA-IECMWF-5 mm PWV bins varies greatly from about -65% to 10%. When PWV values are smaller than 10 mm, there is an obvious wet bias relative to the ERA-IECMWF. The largest negative fractional error occurs at the extremely low (less than 5 mm) PWV values. When PWV values are larger than 60 mm, dry bias relative to ERA-IECMWF can be observed for PWV values. Figure 7(c) exhibits the relationship between RMS error and elevation. It is clearly seen that the RMS error increases generally with increase in elevation. A high correlation coefficient of 0.73 is achieved between RMS error and elevation. This is consistent with the bias and RMS error maps in Figure 6. The high correlation coefficient is probably due to reasons: 1) vertical adjustment for PWV according to equation (8) is unable to accurately capture the highly dynamic water vapor variation in the vertical direction; and (2) the performance of the high-resolution ($0.25^\circ \times 0.25^\circ$) ERA-IECMWF-PWV product degrades with increased elevation.

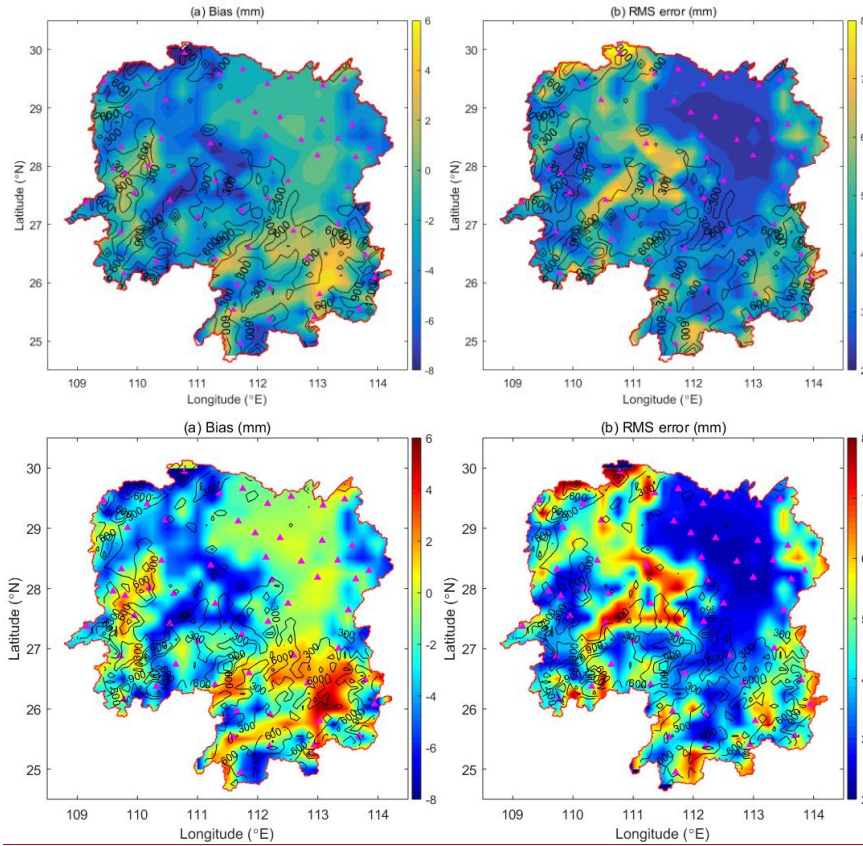


Figure 6: Map of (a) bias and (b) RMS error of the differences between ~~ERA-I~~ ECMWF PWV and GNSS interpolated PWV over the Hunan Province for the year 2015. Black contours represent the elevation (unit: m).

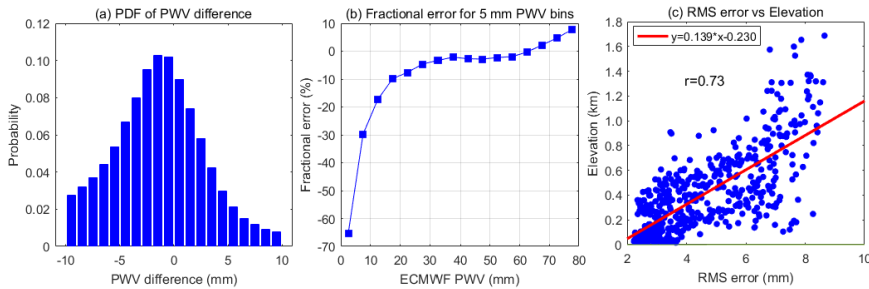


Figure 7: (a) PDF of PWV difference, (b) fractional error as percent by ERA-Interim 5 mm PWV bins, and (c) relationship between RMS error and elevation for the comparison between ERA-Interim and GNSS.

4.4 Monitoring water vapor variations using GNSS derived PWV maps

5 The ultimate goal of this study is to apply the constructed PWV maps for the study of weather forecasting. We further investigated the water vapor variations during a large-scale heavy precipitation event using the PWV maps derived from GNSS observations. In June 2015, Hunan Province suffered several large-scale regional torrential rains, which caused major floods and massive landslides in some places. An average rainfall of 236 mm over the whole province was recorded in that month, and the accumulated rainfall exceeded 500 mm in many areas. In this study, we focused on a heavy rainfall process

10 occurring during 6-8 June 2015. Figure 8Figure-8 exhibits the geographic distribution of the daily accumulated precipitation over the Hunan Province for 6, 7 and 8 June 2015. The precipitation data are retrieved from the Tropical Rainfall Measuring Mission (TRMM), a joint mission of NASA (National Aeronautics and Space Administration) and the Japan Aerospace Exploration Agency to measure rainfall for weather and climate research (Kummerow et al., 1998; Lau and Wu, 2011)(Kummerow et al., 1998; Lau and Wu, 2011). As shown in Figure 8Figure-8(a), the accumulated precipitation on 6

15 June decreased from about 60 mm at the southeast to 0 mm at the northwest. On 7 June, rainfalls were observed over most parts of the whole-province with heavy precipitation mainly occurring in the northern Hunan Province. Afterwards, on 8 June, the precipitation weakened on most of the Hunan province except for an increase in the northeast.

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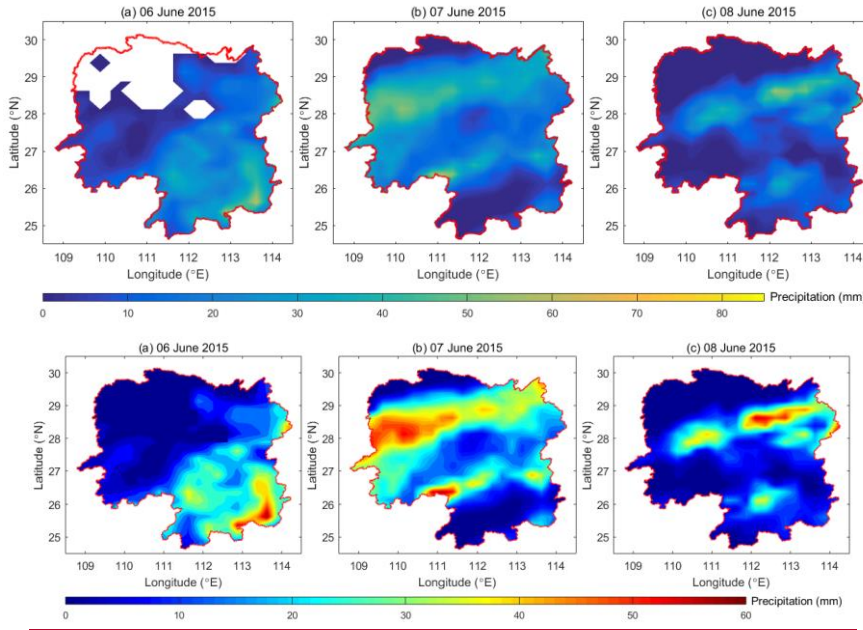
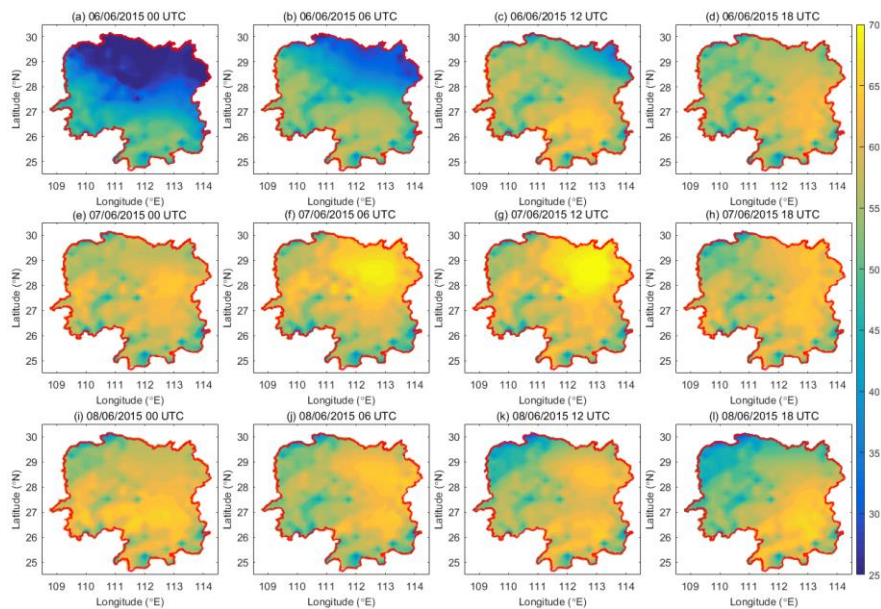


Figure 8: Map of daily accumulated precipitation in the Hunan Province on (a) 6 June 2015, (b) 7 June 2015, and (c) 8 June 2015. The precipitation data were retrieved from the TRMM with a spatial resolution of $0.25^{\circ} \times 0.25^{\circ}$.

- 5 Figure 9Figure-9 presents the evolution of PWV derived from GNSS observations for the Hunan Province during the period of 6-8 June 2015 with a time interval of 6 h. In addition, the TRMM derived rain rates over Hunan for the same epochs are displayed in Figure 10. On 6 June (see Figure 9Figure-9(a-d)), the whole province experienced an obvious increase in PWV from south to north, indicating that a large amount of moisture from the south flowed into Hunan. This is consistent with the precipitation pattern displayed in Figure 8Figure-8(a) in that the rainfall gradually decreased from south to north. On 7 June,
- 10 significant PWV changes mainly concentrated in regions above-north of 238°N . Especially in the northeast, PWV experienced an increase of 10-15 mm from UTC 00 to UTC 12 of 7 June and then dissipated quickly. On 8 June, obvious PWV decreases were observed in the northwest, whereas the southeast experienced a slight increase in PWV. The precipitation maps shown in Figure 8Figure-8(b) and (c) also agree well with the PWV variations. From 7 June to 8 June, the precipitation areas shrank-greatlylargely decreased in the north whilst slightly expanded in the south. Refer to the rain rates
- 15 at the corresponding epochs, as shown in Figure 10, we cannot observe close correlations between the PWV and the rain rate. Larger moisture convergence is not necessarily linked with higher rain rate occurrence. This is because the moisture convergence is not the only cause of precipitation, whilst also controlled by many other factors such as wind, temperature

and terrain. However, the GNSS-derived PWV maps are able to reveal the moisture advection, transportation and convergence during the heavy precipitation event.



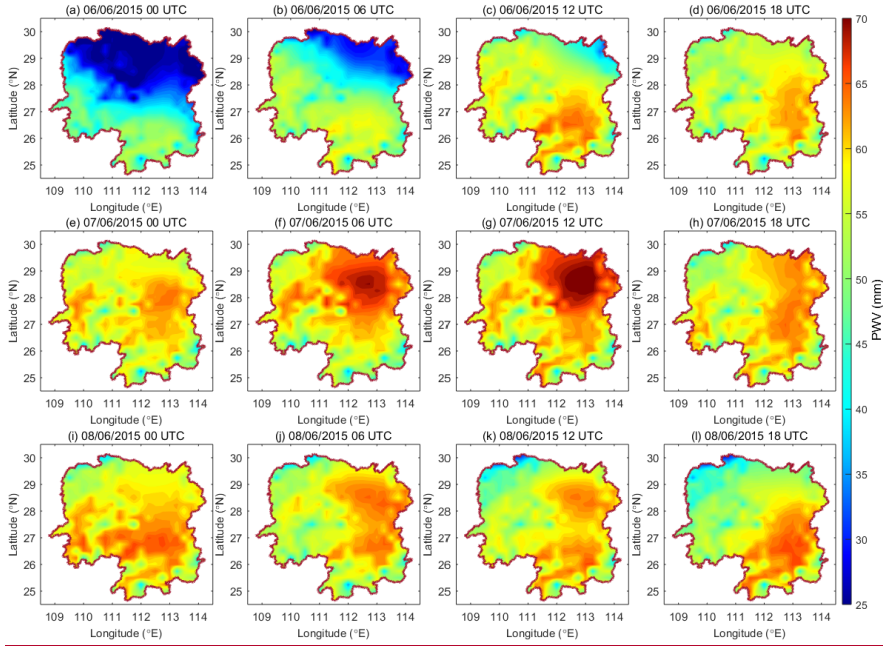


Figure 9: Evolution of GNSS-derived PWV maps for the Hunan province every 6 h from UTC 00, 6 June 2015 to UTC 18, 8 June 2015.

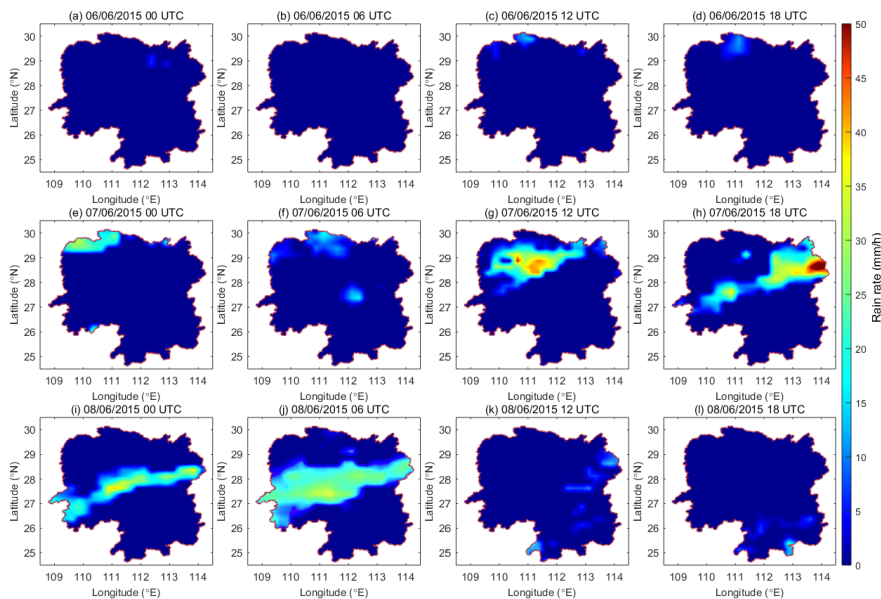


Figure 10: Evolution of rain rate maps for the Hunan province every 6 h from UTC 00, 6 June 2015 to UTC 18, 8 June 2015. The rain rate data were retrieved from the TRMM with a spatial resolution of $0.25^{\circ} \times 0.25^{\circ}$.

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- 5 In addition, ~~Figure 11~~Figure 10 further exhibits the geographic distribution of the correlation coefficient between precipitation and PWV. The correlation coefficients vary greatly from -0.9 to 0.8 depending upon the location. High positive correlation coefficients are present in western regions between 27°N and 27.5°N . Precipitation and PWV shows a high negative relationship in regions between 26°N and 27°N . It can be found that high positive/negative correlation coefficients mainly occur in mountainous regions, especially in hillsides and valleys. This is because the meso-scale orography creates
- 10 favorable conditions for precipitation formation by generating moisture convergence and the small scale orography plays an important role by triggering convective initiation and enhancement (Labbouz et al., 2013)(Labbouz et al., 2013). Therefore, precipitation and PWV correlate more closely in mountainous regions than flat terrains, and mountainous regions are often sensitive areas prone to high frequency of heavy precipitation.

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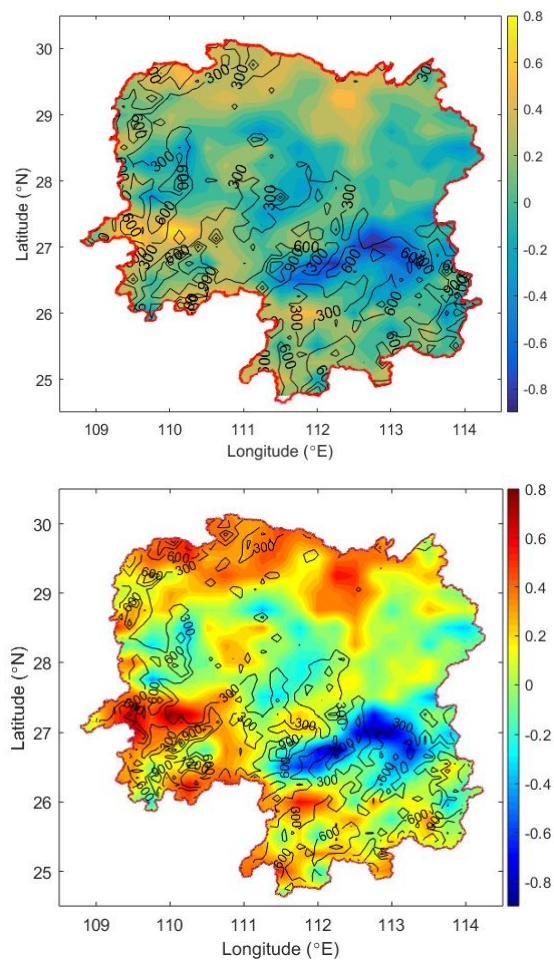


Figure 11: Map of the correlation coefficient between precipitation and PWV for the heavy rainfall process of 6-8 June 2015 over the Hunan Province. Black contours represent the elevation (unit: m).

5 Summary and conclusions

The lack of collocated meteorological data at GNSS station makes it difficult to take full advantage of GNSS observations for weather studies. This paper investigates an alternative method for accurate determination of PWV for near-real-time applications using GNSS data and nearby synoptic observations. Moreover, we present a method to construct PWV maps with the use of [a](#) GNSS network, which is critical for improving the forecasting capability of extreme weathers, e.g. heavy rainfall.

The proposed approach for PWV map construction consists of two main steps: 1) the P_s and T_m derived at the nearby synoptic sites are interpolated onto the location of the GNSS stations through both vertical and horizontal adjustments; and 2) vertical reduction and horizontal interpolation are performed to construct PWV map using the retrieved GNSS PWV. In this study, ~~ERA-IECMWF reanalysis~~ data over the whole year of 2014 were employed to estimate all the parameters involved in the above two steps. The accuracies of the synoptic interpolated and GPT2w derived P_s and T_m have been evaluated by comparing them against the observed values at 3 radiosonde sites in 2015. RMS errors of P_s and T_m derived from ~~the~~ GPT2w model vary in the range of 4.7-5.6 hPa and 3.8-4.2 K, respectively. The RMS errors from synoptic interpolation are 1.7-3.0 hPa and 2.5-3.0 K, respectively, which are much better than the GPT2w model.

In addition, GNSS interpolated PWVs are assessed with respect to reference PWV values from radiosonde and ~~ERA-IECMWF reanalysis~~. GNSS interpolated PWVs show a good agreement with the radiosonde measured ones with RMS errors varying in the range of 2.4-3.0 mm. In the comparison with ~~ERA-IECMWF~~, the biases of their differences vary from -8 mm to 6 mm over the Hunan Province and mountainous regions have a larger bias than flat regions in general. RMS errors are within the range of 2-8 mm with ~~those for~~ most regions ~~being~~ less than 3 mm. For PWV values less than 10 mm or more than 60 mm, there is an obvious wet or dry bias relative to ~~the ERA-IECMWF~~. Furthermore, the RMS errors are found to increase with increased elevation in general and a high correlation coefficient of 0.73 is obtained between RMS error and elevation.

We further apply the constructed PWV maps to monitor the water vapor variability during a large-scale heavy precipitation event that occurred during 6-8 June 2015 in the Hunan Province. Results demonstrate that it is possible to reveal the moisture advection, transportation and convergence during the heavy rainfall using PWV maps. Since the orography provides favorable conditions for precipitation formation, we also find that the precipitation and PWV correlate more closely in mountainous regions, especially in hillsides and valleys.

This research demonstrates the potentials of retrieving accurate PWV from GNSS observations using adjacent synoptic data and generating high-quality PWV maps from ~~the~~ GNSS network for weather prediction in near-real-time. Future work will focus on the three following issues: (1) examining the reliability of the PWV map construction in other areas with highly dynamic water vapor; (2) assessing the performance ~~of~~ the constructed PWV maps with higher spatial and temporal resolutions; and (3) assimilating the PWV maps into ~~a~~ numerical prediction model to enhance the capability of extreme weather forecasting.

Data availability. The ECMWF ERA-Interim reanalysis products are available online (<http://apps.ecmwf.int/datasets/>). The radiosonde data were obtained from <http://weather.uwyo.edu/upperair/sounding.html>. The TRMM rainfall data were provided by <https://pmm.nasa.gov/data-access/downloads/trmm>. The synoptic observations were provided by <http://rda.ucar.edu/datasets/ds336.0/>. The SRTM30 PLUS data were provided by <http://topex.ucsd.edu/index.html>. The radiosonde data of Hong Kong were obtained from <http://weather.uwyo.edu/upperair/sounding.html>. The GNSS observations of the Hunan GNSS network presented in this study are available from the authors upon request (yeary124@csu.edu.cn).

10 *Competing interests.* The authors declare that they have no conflict of interest.

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