#### Associate Editor Decision: Publish subject to technical corrections (16 Jan 2020) by Wiebke Frey

#### **Comments to the Author:**

Dear Fan Mei et al., I found some remaining very minor issues in the supplement:

Response: We sincerely appreciate the comments and suggestions from our associate editor Wiebke Frey. Thank you very much for considering the publication of our manuscript. We address your comments below (also in blue).

line 45: 'blew' -> below

**Response: Corrected in the line 46.** 

line 50: missing white space: '(RIE)were'

**Response: Corrected in the current line 51.** 

line 76: You say G-1 here, but G1 elsewhere, please be consistent.

**Response: Corrected in the current line 77.** 

line 82: 'October 01' here, October 1 elsewhere, please be consistent.

#### **Response: Corrected in the current line 83.**

Sections: Some Figure do not really belong to the sections they appear in, for example in Section 1 Figure S2-S5 are not related to 'Pilot preparation'. Maybe make the 'pilot preparation' part a subsection of a section that could be called something like 'Additional Observations'. Figures S11-S13 are not related to 'CCN closure' and could thus be in a new Section etc.

Response: We changed the section 1 to "Additional observations" and added additional section titles to the related figures: section 4 "Cloud probe observations" (in the current line 135) before the figure S11-S13 and section 5 "Radiation measurements" (in the current line 155) before the figure S14.

For final supplement: Please check that figure captions are not on another page than the respective figure itself!

Response: We revised the figure and captions accordingly.

Best wishes, Wlebke Frey

### 1 Supplemental material:

#### 2 1. Additional observations

3 Both aircraft cannot appear at the same location at the same time due to safety concerns. Thus, the approval of a formation (inter-comparison) flight was acquired six months before the 4 campaign through DOE Pacific Northwest Site Office (PNSO) and the Office of Aviation 5 Management (OAM). Essential risk mitigation was also discussed and approved by the Pacific 6 Northwest National Laboratory Aviation Risk Management Committee (PNNL ARMC). During 7 the IOP, both aircraft crew and scientists teams set up a meeting to discuss the potential flight plan. 8 9 After the flight plan was formed, both pilots briefed the plan to the Brazilian Air Force (BAF) and Airport Traffic Control (ATC). The clear-sky flight would be under Visual Flight Rules (VFR), 10 11 which means good weather and no cloud, and pilots communicate with each other using an air-to-12 air frequency. For coordinated flights in cloudy conditions, the G1 and the HALO were both on Instrument Flight Rules (IFR) flight plan. 13

The coordinated flight on October 1, 2014, was initially designed to be a coordinated flight under a cloudy condition, which means the G1 and the HALO flew the same flight leg with at least 300 m altitude offset and at least 5 minutes apart. However, the coordinated two flight legs (~900 m and ~1200 m) are all below the cloud. Thus, the comparison focus on the correlation between two aircraft measurements, not vertical profiling.

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Figure S1. Time colored flight track of the G1 (circle) and the HALO (triangle) on October 1,
2014, during a cloudless coordinated flight (This figure was created using Mapping Toolbox<sup>TM</sup>

24 © COPYRIGHT 1997–2019 by The MathWorks, Inc).





Figure S2, Atmospheric parameters observed by the G1 and the HALO on October 1, 2014.



Figure S3. Horizontal wind speed between 2000-3000 m altitude on September 21, 2014.











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44 2. Additional information for AMS

Most of the details for the AMS measurements have been included in the separate AMS papers
(Schulz et al., 2018; Shilling et al., 2018). Brief summaries are provided below.

The G1 AMS was operated with a constant pressure inlet (CPI), which was set to a constant pressure during the campaign. The G1 AMS was calibrated once a week during the deployment. One additional calibration was performed after the flight day, and all the calibrations were in agreement with each other. Based on five calibrations, the averaged parameters such as the airbeam signal (AB), the ionization efficiency (IE), and the relative ionization efficiency (RIE) were applied to all of the data. A real-time correction was made to account for the variations in the AB changes to improve the instrument sensitivity. Typically, this correction is small (<20%) in absolute magnitude. The particle collection efficiency (CE) was determined by comparing AMS data to UHSAS and FIMS data. We also confirmed the CE=0.5 by comparing mass loadings observed at the T3 site to the G1 data.

The HALO AMS was calibrated before, during (twice), and after the campaign for (relative) ionization efficiencies of nitrate, ammonium, and sulfate (Schulz et al., 2018). For organics, the default relative ionization efficiency of 1.4 was assumed. The inlet flow was kept constant by the CPI and was measured before and during the campaign. Collection efficiency of 0.5 was applied, as recommended by Middlebrook et al. (2012) for low nitrate conditions. Further details on the operation of the C-ToF-AMS are given in Schulz et al. (2018).

63 Figure S6(a) shows vertical profiles of the total mass concentrations measured by the two AMS instruments on September 21. Above 2500 m altitude, the agreement between the two 64 instruments is excellent (mean difference less than 5%). Between 2000 and 2500 m, the agreement 65 is within the uncertainty range. Below 2000 m altitude, however, the aerosol particle mass 66 concentrations measured by the AMS operated on HALO are lower than the concentrations 67 measured by the AMS on the G1. To compare AMS data to UHSAS data, the aerosol mass 68 concentrations of the G1 AMS were converted to the aerosol volume concentration assuming an 69 organic compound density of 1.5 g cm<sup>-3</sup> (Pöschl et al., 2010). The converted aerosol volume 70 concentration agreed well with the volume concentration calculated based on UHSAS data, 71 especially below 2500 m, as shown in Figure S6(b). The agreement at lower altitudes suggests that 72 73 the lower concentration measured by the HALO AMS is due to the transmission efficiency issue 74 in the constant pressure inlet used by the HALO AMS. This inlet was a prototype, designed and 75 built at MPIC Mainz, and works by changing the size of the critical orifice that regulates the flow 76 into the aerodynamic lens. The design and transmission characteristics will be described in an 77 upcoming publication (Molleker, S., in prep.). The AMS aboard the G1 used a constant pressure 78 inlet based on the design in Bahreini et al., 2008. Thus, we conclude that data above 2500 m 79 altitude measured by the AMS aboard HALO in 2014 are valid, while data below 2500 m need to be corrected using correction factors derived from laboratory characterization before further study. 80

After 2014, the HALO AMS inlet design was improved to address the inlet transmission issues
specific to this field campaign.

The second comparison between the two AMS conducted on October 1 is shown in Figures S6 and S7. The findings are basically in agreement with those of September 21, although the underestimation of aerosol mass concentration due to the inlet in the HALO AMS appears here to be restricted to altitudes lower than 1500 m.



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Figure S6. (a) Comparison of aerosol mass loading measured by the G1 and HALO AMS on
September 21; (b) aerosol volume concentration comparison from AMS and the integrated
UHSAS on the G1.





Figure S8. The vertical profiling of the relative fractions for the chemical species observed by the
 G1 and HALO during October 1.

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#### 102 **3.** CCN closure

To further examine the relative importance of mixing state and chemical composition, the 103 CCN concentrations were calculated from aerosol particle size distribution, and chemical 104 105 composition measured onboard the G1. The calculation was based on *k*-Köhler parameterization, (Kohler, 1936; Petters and Kreidenweis, 2007, 2008, 2013) and the detail of the approach was 106 described by Mei et al. (2013b). For the flight on September 9, 2017, the CCN number 107 concentration calculated from the G1 UHSAS size distribution and chemical composition exhibits 108 109 underestimation at a supersaturation of 0.5% (Fig, S9(a)) and when the altitude is below 1000 m (Fig S9(b)). This underestimation suggests that the UHSAS size range (90-500 nm) did not fully 110 cover the aerosols with the critical activation diameter  $(D_{p,50})$  at high supersaturation. Thus, the 111 FIMS measurements onboard the G1 was the more appropriate size distribution for both the CCN 112 closure study. The CCN concentration calculated using the size distribution from FIMS agrees 113

- well with the measurement (Fig. S10). The scattering of the comparison data in Figure 15 is likelydue to the chemical composition and mixing state effect on aerosol hygroscopicity.
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Figure S9. Comparison of calculated CCN with measured CCN using the averaged 1 min measurements from the G1: (a) colored by different supersaturations. (b) colored by different altitudes. (Note that both plots used the calculated CCN number concentration from UHSAS size distribution.)



## **4. Cloud probe observations**







Figure S12. The cloud droplet number concentration from HALO on September 21.











Figure S14. Time series comparison of the G1 (SPN-1) and HALO (SMART-Albedometer) 157 radiation measurements on September 9.

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#### Table S1. Calibration and maintenance for the instruments deployed on G1 160

Measurement Variables	Instruments deployed on the G1 (Martin et al., 2016; Schmid et al., 2014)	Calibration/Maintenance
Static Pressure	Rosemount (1201F1), 0-1400 hPa	Calibrated before/after each field campaign
Static air temperature	Rosemount E102AL/510BF -50 to +50 °C	Calibrated before/after each field campaign
Dewpoint temperature	Chilled mirror hygrometer 1011B -40 to +50 °C	Calibrated before/after each field campaign
3-D wind	Aircraft Integrated Meteorological Measurement System 20 (AIMMS- 20)	Calibrated with Special flight pattern before each field campaign. Inter- comparison with other GPS/INS during deployment.
Particle number concentration	CPC, cut off size $(D_p) = 10 \text{ nm}$	Calibrated before/after each field campaign. Weekly calibration of sample and sheath flow rates and inter-comparisons with similar counters during deployment.
Size distribution*	UHSAS-A, 60-1000 nm.	Calibrated before/after each field campaign. Weekly check of sizing with PSL
FIMS	10 nm – 500 nm	Calibrated before/after each field campaign. Weekly calibration of sample and sheath flow rates and checks with one size PSL

Non-Refractory particle chemical composition	HR-ToF-AMS: Organics, Sulfate, Nitrate, Ammonium, Chloride, 60- 1000 nm	Weekly calibrations.
CCN concentration	CCN-200, SS= 0.25, 0.5%	Calibrated before/after each field campaign. Biweekly calibration with ammonium sulfate particles.
Gas phase concentration	N2O/CO and Ozone Analyzer, CO, O <sub>3</sub> concentration, precision 2 ppb	Calibrated before/after each field campaign with calibration gas mixture.
CDP	2-50 μm, Δ <i>D</i> <sub>p</sub> =1-2 μm	Calibrated before/after each field campaign by the vendor. Weekly check of sizing with glass beads of several sizes
FCDP	2-50 μm, Δ <i>D</i> <sub>p</sub> =1-2 μm	Calibrated before/after each field campaign by the vendor. Weekly check of sizing with glass beads of several sizes
2DS	10-1000 μm	Calibrated before/after each field campaign by the vendor.
Radiation	SPN1 downward irradiance, 400- 2700 nm	Calibrated before/after each field campaign

# 162 Table S2. Calibration and maintenance for the instruments deployed on HALO

Measurement Variables	Instruments deployed on HALO	Calibration/Maintenance
	(Wendisch et al., 2016)	
Static Pressure	Instrumented nose boom tray (DLR	Calibrated before/after each field
	development), 0-1400 hPa	campaign
Static air temperature	Total Air Temperature (TAT) inlet	Calibrated before/after each field
	(Goodrich/Rosemount type 102) with	campaign
	an open wire resistance temperature	
	sensor (PT100),	
	-70 to +50 °C	
Dewpoint temperature	Derived from the water-vapor mixing	Calibrated before/after each field
	ratio, which is measured by a tunable	campaign
	diode laser (TDL) system (DLR	
	development), 5-40000 ppmv	
3-D wind	Instrumented nose boom tray (DLR	Calibrated before/after each field
	development) with an air data probe	campaign
	(Goodrich/Rosemount) 858AJ and	
	high-precision Inertial Reference	
	System (IGI IMU-IIe)	
Particle number concentration	CPC, cut off size $(D_p) = 10 \text{ nm}$	Calibrated before/after each field
		campaign. Weekly inter-comparisons
		with similar counters during
		deployment.
Size distribution*	UHSAS-A, 60-1000 nm.	Calibrated before/after each field
		campaign. weekly check of sizing
		with PSL
Non-Refractory particle chemical	C-ToF-AMS: Organics, Sulfate,	Calibrated before and after the
composition	Nitrate, Ammonium, Chloride, 60-	campaign and twice during the
	1000 nm	campaign

CCN concentration	CCN-200, SS= 0.13-0.53%	Calibrated before/after each field
		campaign. Weekly calibration with
		ammonium sulfate particles.
Gas phase concentration	N2O/CO and Ozone Analyzer, CO,	Calibrated before/after each field
	O <sub>3</sub> concentration, precision 2 ppb	campaign with calibration gas
		mixture.
Cloud properties*	CCP-CDP, 2.5-46 mm, $\Delta D_p=1-2 \mu m$	Calibrated before/after each field
		campaign. Weekly check of sizing
		with glass beads of several sizes
	NIXE-CAS: 0.61 -52.5 μm	Calibrated before/after each field
		campaign. Weekly check of sizing
		with glass beads of several sizes
	NIXE-CIPgs, 15-960 µm	Calibrated before/after each flight
		with a spinning disk.
	ССР-СІРдs: 15-960 µm	Calibrated before/after each flight
		with a spinning disk.
Radiation	SMART Albedometer, downward	Weekly calibrations.
	spectral irradiance 300-2200 nm	

- 164 Table S3. List of compared measurement ranges and measurement variances caused by the spatial
- 165 variation during the field campaign.

Measurement	Measured Range during the Field	Measurement Variances between the
Variables	Campaign	Two Aircraft
Static Pressure	500 – 1010 hPa	< 1 %
Static air	272 – 310 K	< 1%
temperature		
Dewpoint	230 -300 K	Without clouds, <1%
temperature		With clouds, the measurement from the G1
		can be up to 5% lower than that of HALO
3-D wind	1-15 m/s	< 40%
Particle number	$500 - 15,000 \text{ cm}^{-3}$	< 20% for CPC, <50% for UHSAS (size
concentration		dependent)
Non-Refractory	$< 10 \ \mu g \cdot m^{-3}$	< 10% above 2500 m
particle chemical		
composition		Up to 50% below 2500 m
CCN	$SS=0.25\%$ , $100 - 2000 \text{ cm}^{-3}$	< 10% above 2500 m
concentration		
		Up to 50% below 2500 m
Gas phase	Ozone: 15-75 ppb	Ozone: < 25%
concentration	CO: 50-200 ppb	CO: < 15%
Cloud droplet	3- 20 μm	<50 %
number		
concentration		
Downward	200 -1500 W·m <sup>-2</sup>	< 10%
irradiance		

#### 169 **Reference**

- 170 Fan, J., Rosenfeld, D., Zhang, Y., Giangrande, S. E., Li, Z., Machado, L. A., Martin, S. T., Yang, Y., Wang, J.,
- and Artaxo, P.: Substantial convection and precipitation enhancements by ultrafine aerosol particles,
  Science, 359, 411-418, 2018.
- 173 Kohler, H.: The nucleus in and the growth of hygroscopic droplets, Transactions of the Faraday Society,
- 174 32, 1152-1161, 1936.
- Kotchenruther, R. A. and Hobbs, P. V.: Humidification factors of aerosols from biomass burning in Brazil, J
   Geophys Res-Atmos, 103, 32081-32089, 1998.
- 177 Moran-Zuloaga, D., Ditas, F., Walters, D., Saturno, J., Brito, J., Carbone, S., Chi, X. G., de Angelis, I. H., Baars,
- 178 H., Godoi, R. H. M., Heese, B., Holanda, B. A., Lavric, J. V., Martin, S. T., Ming, J., Pohlker, M. L.,
- 179 Ruckteschler, N., Su, H., Wang, Y. Q., Wang, Q. Q., Wang, Z. B., Weber, B., Wolff, S., Artaxo, P., Poschl, U.,
- Andreae, M. O., and Pohlker, C.: Long-term study on coarse mode aerosols in the Amazon rain forest with
   the frequent intrusion of Saharan dust plumes, Atmos Chem Phys, 18, 10055-10088, 2018.
- Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud
- 183 condensation nucleus activity, Atmos Chem Phys, 7, 1961-1971, 2007.
- 184 Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud
- 185 condensation nucleus activity Part 2: Including solubility, Atmos Chem Phys, 8, 6273-6279, 2008.
- 186 Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud
- condensation nucleus activity Part 3: Including surfactant partitioning, Atmos Chem Phys, 13, 1081-1091,
  2013.
- 189 Pöschl, U., Martin, S., Sinha, B., Chen, Q., Gunthe, S., Huffman, J., Borrmann, S., Farmer, D., Garland, R.,
- and Helas, G.: Rainforest aerosols as biogenic nuclei of clouds and precipitation in the Amazon, Science,
  329, 1513-1516, 2010.
- 192 Schulz, C., Schneider, J., Amorim Holanda, B., Appel, O., Costa, A., de Sá, S. S., Dreiling, V., Fütterer, D.,
- Jurkat-Witschas, T., Klimach, T., Krämer, M., Martin, S. T., Mertes, S., Pöhlker, M. L., Sauer, D., Voigt, C.,
- Weinzierl, B., Ziereis, H., Zöger, M., Andreae, M. O., Artaxo, P., Machado, L. A. T., Pöschl, U., Wendisch,
- 195 M., and Borrmann, S.: Aircraft-based observations of isoprene epoxydiol-derived secondary organic 196 aerosol (IEPOX-SOA) in the tropical upper troposphere over the Amazon region, Atmos. Chem. Phys.
- 197 Discuss., 2018, 1-32, 2018.
- Shilling, J. E., Pekour, M. S., Fortner, E. C., Artaxo, P., Sá, S. d., Hubbe, J. M., Longo, K. M., Machado, L. A.,
  Martin, S. T., and Springston, S. R.: Aircraft observations of the chemical composition and aging of aerosol
  in the Manaus urban plume during GoAmazon 2014/5, Atmos Chem Phys, 18, 10773-10797, 2018.
- 201 Wang, J., Krejci, R., Giangrande, S., Kuang, C., Barbosa, H. M., Brito, J., Carbone, S., Chi, X., Comstock, J.,
- Ditas, F., Lavric, J., Manninen, H. E., Mei, F., Moran-Zuloaga, D., Pohlker, C., Pohlker, M. L., Saturno, J.,
- Schmid, B., Souza, R. A., Springston, S. R., Tomlinson, J. M., Toto, T., Walter, D., Wimmer, D., Smith, J. N.,
- Kulmala, M., Machado, L. A., Artaxo, P., Andreae, M. O., Petaja, T., and Martin, S. T.: Amazon boundary
- 205 layer aerosol concentration sustained by vertical transport during rainfall, Nature, 539, 416-419, 2016.
- 206 Williamson, C., Kupc, A., Wilson, J., Gesler, D. W., Reeves, J. M., Erdesz, F., McLaughlin, R., and Brock, C.
- A.: Fast time response measurements of particle size distributions in the 3-60 nm size range with the
- 208 nucleation mode aerosol size spectrometer, Atmos Meas Tech, 11, 3491-3509, 2018.