

**Determining the Daytime Earth Radiative Flux from National  
Institute of Standards and Technology Advanced Radiometer  
(NISTAR) Measurements**

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6 The National Institute of Standards and Technology Advanced Radiometer (NISTAR) on-  
 7 board Deep Space Climate Observatory (DSCOVR) provides continuous full disc global  
 8 broadband irradiance measurements over most of the sunlit side of the Earth. The three ac-  
 9 tive cavity radiometers measures the total radiant energy from the sun-lit side of the Earth in  
 10 shortwave (SW, 0.2-4  $\mu\text{m}$ ), total (0.4-100  $\mu\text{m}$ ), and near-infrared (NIR, 0.7-4  $\mu\text{m}$ ) channels.  
 11 The Level 1 NISTAR dataset provides the filtered radiances (the ratio between irradiance  
 12 and solid angle). To determine the daytime top-of-atmosphere (TOA) shortwave and long-  
 13 wave radiative fluxes, the NISTAR measured shortwave radiances must be unfiltered first.  
 14 An unfiltering algorithm was developed for the NISTAR SW and NIR channels using a spec-  
 15 tral radiance data base calculated for typical Earth scenes. The resulting unfiltered NISTAR  
 16 radiances are then converted to full disk daytime SW and LW flux, by accounting for the  
 17 anisotropic characteristics of the Earth-reflected and emitted radiances. The anisotropy fac-  
 18 tors are determined using scene identifications determined from multiple low Earth orbit and  
 19 geostationary satellites and the angular distribution models (ADMs) developed using data  
 20 collected by the Clouds and the Earth's Radiant Energy System (CERES). Global annual  
 21 daytime mean SW fluxes from NISTAR are about 6% greater than those from CERES, and  
 22 both show strong diurnal variations with daily maximum-minimum differences as great as  
 23 20  $\text{Wm}^{-2}$  depending on the conditions of the sunlit portion of the Earth. They are also  
 24 highly correlated, having correlation coefficients of 0.89, indicating that they both capture  
 25 the diurnal variation. Global annual daytime mean LW fluxes from NISTAR are 3% greater  
 26 than those from CERES, but the correlation between them is only about 0.38.

# 1. Introduction

The Earth’s climate is determined by the amount and distribution of the incoming solar radiation absorbed and the outgoing longwave radiation (OLR) emitted by the Earth. Satellite observations of Earth Radiation Budget (ERB) provide critical information needed to better understand the driving mechanisms of climate change; the ERB has been monitored from space since the early satellite missions of the late 1950s and the 1960s (House et al. 1986). Currently, the Clouds and the Earth’s Radiant Energy System (CERES) instruments (Wielicki et al. 1996; Loeb et al. 2016) have been providing continuous global top-of-atmosphere (TOA) reflected shortwave radiation and OLR since 2000. CERES data have been crucial to advance our understanding of the Earth’s energy balance (e.g., Trenberth et al. 2009; Kato et al. 2011; Loeb et al. 2012; Stephens et al. 2012), aerosol direct radiative effects (e.g., Satheesh and Ramanathan 2000; Zhang et al. 2005; Loeb and Manalo-Smith 2005; Su et al. 2013), aerosol-cloud interactions (e.g., Loeb and Schuster 2008; Quaas et al. 2008; Su et al. 2010b), and to evaluate global general circulation models (e.g., Pincus et al. 2008; Su et al. 2010a; Wang and Su 2013; Wild et al. 2013).

The Earth’s radiative flux data record is augmented by the launch of the Deep Space Climate Observatory (DSCOVR) on February 11, 2015. DSCOVR is designed to continuously monitor the sunlit side of the Earth, being the first Earth-observing satellite at the Lagrange-1 (L1) point,  $\sim 1.5$  million km from Earth, where it orbits the Sun at the same rate as the Earth (see Figure 1a). DSCOVR is in an elliptical Lissajous orbit around the L1 point and is not positioned exactly on the Earth-sun line, therefore only about 92~97% of the sunlit Earth is visible to DSCOVR. As illustrated in Figure 1b, the daytime portion ( $A_h$ ) is not visible to the DSCOVR. Strictly speaking, the measurements from DSCOVR are not truly ‘global’ daytime measurements. However, for simplicity we refer to them as global daytime measurements. Onboard DSCOVR, the National Institute of Standards and Technology Advanced Radiometer (NISTAR) provides continuous full disc global broadband irradiance measurements over most of the sunlit side of the Earth (*viewing the sunlit side*

of the Earth as one pixel). Besides NISTAR, DSCOVR also carries the Earth Polychromatic Imaging Camera (EPIC) which provides 2048 by 2048 pixel imagery 10 to 22 times per day in 10 spectral bands from 317 to 780 nm. On June 8, 2015, more than 100 days after launch, DSCOVR started orbiting around the L1 point.

The NISTAR instrument was designed to measure the global daytime shortwave (SW) and longwave (LW) radiative fluxes. *The original objective of NISTAR was to monitor the energy from the sunlit side of the Earth continuously, and to understand the effects of weather systems and clouds on the daytime energy. However, one limitation of NISTAR is its relatively low signal-to-noise ratios, which necessitates averaging significant time periods to adequately reduce the instrument noise levels. This constrains the temporal resolution of meaningful results to about 4 hours, thus prevent us from “continuously” monitoring the sunlit side of the Earth. Nevertheless, NISTAR measurements can still be useful for assessing the hourly fluxes produced by combining the observations from multiple low-Earth orbit and geostationary satellites (Doelling et al. 2013) and for model evaluation using the spectral ratio information (Carlson et al. 2019).* NISTAR measures an irradiance at the L1 point at a small relative azimuth angle,  $\phi_o$ , which varies from  $4^\circ$  to  $15^\circ$ , as shown in Figure 1a. As such, the radiation it measures comes from the near-backscatter position, which is different from that seen at other satellite positions as indicated in Figure 1a by the varying arrow lengths corresponding to scattering angles,  $\Theta_1 - \Theta_3$ . Other types of Earth-orbiting satellites view a given spot on the Earth from various scattering angles that vary as a function of local time (e.g., geostationary) or overpass time (e.g., Sun-synchronous). When averaged over the globe, the uncertainties in the anisotropy corrections are mitigated by compensation. That is, any small biases at particular angles are balanced by observations taken at other angles. In contrast, instruments on DSCOVR view every spot on the Earth from a single scattering angle that varies slowly within a small range over the course of the Lissajous orbit. Thus, the correction for anisotropy is critical. The biases in the anisotropy correction for the DSCOVR scattering angle are mitigated and potentially minimized by the wide range of different scene

types viewed in a given NISTAR measurement (Su et al. 2018).

Su et al. (2018) described the methodology to derive the global mean daytime shortwave (SW) anisotropic factors by using the CERES angular distribution models (ADMs) and a cloud property composite based on lower Earth orbiting satellite imager retrievals. These SW anisotropic factors were applied to EPIC broadband SW radiances, that were estimated from EPIC narrowband observations based upon narrowband-to-broadband regressions, to derive the global daytime SW fluxes. Daily mean EPIC and CERES SW fluxes calculated using concurrent hours agree with each other to within 2%. They concluded that the SW flux agreement is within the calibration and algorithm uncertainties, which indicates that the method developed to calculate the global anisotropic factors from the CERES ADMs is robust and that the CERES ADMs accurately account for the Earth’s anisotropy in the near-backscatter direction.

In this paper, the same global daytime mean anisotropic factors developed by Su et al. (2018) are applied to the NISTAR measurements to derive the global daytime mean SW and longwave (LW) fluxes. The NISTAR data and the unfiltering algorithms developed for the NISTAR shortwave and near-infrared channels are detailed in section 2. The data and methodology used to derive the global daytime mean anisotropic factors are presented in section 3. Hourly daytime SW and LW fluxes calculated from NISTAR measurements and comparisons with the CERES Synoptic flux products (SYN1deg, Doelling et al. 2013) are detailed in section 4, followed by conclusions and discussions in section 5.

## 2. NISTAR observation

The NISTAR instrument measures Earth irradiance data for an entire hemisphere using cavity electrical substitution radiometers (ESRs) and filters covering three channels: short-wave (SW, 0.2-4.0  $\mu\text{m}$ ), near-infrared (NIR, 0.7-4.0  $\mu\text{m}$ ), and total (0.2-100  $\mu\text{m}$ ). *Each channel has a dedicated ESR, that by itself is sensitive to radiation from 0.2-100  $\mu\text{m}$ . For*

the NIR and SW channels, filters are positioned in front of each ESR to limit the incident radiation to spectral bands. The filters reside in a filter wheel that, during normal operation, configures each ESR to measure contemporaneously in a different band. Additionally, each ESR has a shutter that modulates the Earth signal by cycling between open and closed states continually with a 50% duty cycle and a period of 4 minutes. The modulation is necessary as the ESRs only measure changes in the incident optical power and, being thermal detectors, they have large offsets (background signals) which drift over relatively short time frames (hours) but not significantly over a shutter cycle. Demodulating the resulting signal removes those offsets and the associated drifts/noise. What remains is a much more stable shutter modulated background that is measured during periodic views of dark space and subsequently subtracted from the signal. The shutter modulated background is largest for the total channel and much smaller for the SW and NIR channels.

The NISTAR calibrated Level 1B data products are derived from pre-launch system level optical calibration and on-orbit offset measurements. The former involved optical response measurements of each active cavity radiometer without a filter in place using a narrow band calibration source whose irradiance was measured with a NIST calibrated reference detector. Those measurements establish the irradiance responsivity of each spectrally flat broadband radiometer. Additionally, measurements of the transmittance of the SW and NIR filters were made. This was done at NIST prior to installation into the NISTAR filter wheel at wavelengths ranging from 200 nm to approximately 18 micrometers. Further, system-level filter transmittance measurements at discrete visible and near-infrared wavelengths were made using the external light source and the NISTAR photodiode channel as a detector. The two transmittance measurements agreed to within a few tenths of a percent. Radiometric offsets are measured on-orbit monthly when NISTAR briefly views dark space. The offset measurement uncertainty is determined by the instrument noise level and the relatively short time allotted to the space-views.

NISTAR Level 1B radiometric products are derived by first subtracting the offsets from

Earth-view measurements and then dividing by the laboratory measured responsivity. The result is irradiance measured at the instrument aperture. Radiance ( $I$ ) is then calculated from the irradiance data and the solid angle ( $\Theta$ ) determined from the DSCOVR-to-Earth distance and the Earth dimensions. When averaging over a 4-hour period, the NISTAR total and SW channel uncertainties ( $k=1$ ) are 1.5% and 2.1%, respectively. As the LW is derived from the difference between the total and unfiltered-SW channels, it contains noise contributions from both. The LW uncertainty is about 3.3% ( $8 \text{ Wm}^{-2}$ ) given that the daytime mean LW and SW fluxes are approximately  $210 \text{ Wm}^{-2}$  and  $240 \text{ Wm}^{-2}$ , respectively, and that the uncertainties between the Total and SW channels are largely uncorrelated.

As mentioned before, Filters are placed in front of the radiometers to measure the energies from the SW and NIR portions of the spectrum. Since no corrections for the impact of filter transmission were applied to the NISTAR L1B data, the SW and NIR radiances from NISTAR must first be unfiltered before they can be used to derive daytime Earth's radiative flux. Here we follow the algorithm developed by (Loeb et al. 2001) to convert measured NISTAR filtered radiances to unfiltered radiances.

Unfiltered SW and NIR radiances are defined as follows:

$$I_u^{band} = \int_{\lambda_1}^{\lambda_2} I_\lambda d\lambda, \quad (1)$$

where 'band' represent either SW or NIR,  $\lambda(\mu\text{m})$  is the wavelength, and  $I_\lambda$  ( $\text{Wm}^{-2} \text{ sr}^{-1} \mu\text{m}^{-1}$ ) is the spectral SW radiance. The filtered radiance is the radiation that passes through the spectral filter and is measured by the detector:

$$I_f^{band} = \int_{\lambda_1}^{\lambda_2} S_\lambda^{band} I_\lambda d\lambda, \quad (2)$$

where  $S_\lambda^{band}$  is the spectral transmission function. Figure 2 shows the NISTAR SW and NIR spectral transmission functions. These functions are determined from ground testing done in 1999 and 2010 at the National Institute of Standards and Technology (NIST).

The spectral radiance database is calculated using high-spectral-resolution radiative transfer model (Kato et al. 2002). Unfiltered radiances are determined by integrating spectral

157 radiances over the appropriate wavelength intervals using Gaussian quadrature. Similarly,  
 158 filtered radiances are computed by integrating over the product of spectral radiance and  
 159 spectral transmission function. The regression coefficients are derived at 480 angles: 6 so-  
 160 lar zenith angles (0.0, 29.0, 41.4, 60.0, 75.5, 85.0 degrees), 8 viewing zenith angles (0, 12,  
 161 24, 36, 48, 60, 72, 84 degrees), and 10 relative azimuth angles (0 to 180, at every 20 de-  
 162 grees). For angles between those given above, the regression coefficients are derived by linear  
 163 interpolation.

164 The database includes spectral radiances calculated over ocean, land/desert, snow/ice  
 165 surfaces for clear and cloudy conditions. Table 1 summarizes the number of each variable  
 166 that are included in the database, there are a total of 722 clear-sky cases and a total of 1519  
 167 cloudy-sky cases for each Sun-viewing geometry. *This is a much larger database comparing*  
 168 *with that used by Loeb et al. (2001).*

169 *For CERES unfiltering, regression coefficients between filtered and unfiltered radiances*  
 170 *were derived as functions of scene type and Sun-viewing geometry (Loeb et al. 2001). Given*  
 171 *that NISTAR views the Earth as a single pixel, a mix of scenes and many Sun-viewing*  
 172 *geometries are observed at the same time. The method used for CERES is not feasible for*  
 173 *unfiltering NISTAR observation. We instead investigated the feasibility of using the ratio,*  
 174  *$\kappa$ , between filtered and unfiltered radiances for unfiltering the NISTAR observations.* Table 2  
 175 lists the mean and the standard deviations of the ratios at different solar zenith angles. *The*  
 176 *ratios for the SW band are extremely stable, varying less than 0.3% among the scenes and*  
 177 *Sun-viewing geometries considered (the smallest ratio, 0.8659, occurs for clear ocean under*  
 178 *overhead sun and the largest ratio, 0.8694, occurs for clear/cloudy land under overhead*  
 179 *sun). As the ratio is not sensitive to the scene type and the Sun-viewing geometry, the SW*  
 180 *unfiltering for NISTAR can be accomplished by:*

$$I_u^{sw} = \frac{I_f^{sw}}{\kappa^{sw}}, \quad (3)$$

181 Here  $I_f^{sw}$  is the filtered radiances directly from the NISTAR L1B data. As the NISTAR view  
 182 always contains clouds, we choose to use the mean ratios of the cloudy ocean and land cases



in Table 2, which is 0.8690 for the SW band. The estimated uncertainty of using this single ratio for unfiltering the SW band is less than 0.3%.

On the other hand, the variability in the ratios of the NIR band can be as large as 6%. Fortunately, the large variability only occurs between clear ocean and clear land. As mentioned earlier, NISTAR view always contains clouds and the mean ratios of the cloudy ocean and land cases, which is 0.8583, is used to unfilter the NISTAR NIR observations. This mean ratio can differ with the individual ratios for different solar zenith angles under cloudy conditions by about 1~2%. The mean ratio of the NIR bands is used to convert the filtered radiances to unfiltered radiances:

$$I_u^{nir} = \frac{I_f^{nir}}{\kappa^{nir}}. \quad (4)$$

In this paper, the measurements from NISTAR NIR channel are not used. The unfiltering of NIR channel is reported here for readers who intend to use this channel.

As there is no filter placed in front of the total channel, the radiance from the total channel does not need to be unfiltered. The LW (4-100  $\mu\text{m}$ ) radiance can be derived by subtracting the unfiltered SW radiance from the total:

$$I_u^{lw} = I^{tot} - I_u^{sw}, \quad (5)$$

The unfiltered radiances ( $I_u^{sw}$  and  $I_u^{lw}$ ) will be used hereafter to derive the daytime mean radiative flux. Although NISTAR L1B data provide observations every second, hourly data (smoothed with 4-hour running mean) are used to derive fluxes because of the level of noise presented in the measurements (DSCOVN NISTAR data quality report v02).

### 3. Global daytime shortwave and longwave anisotropic factors

To derive the global daytime mean SW and LW fluxes from the NISTAR unfiltered radiances, the anisotropy of the TOA radiance field must be considered. The CERES Edition

4 empirical ADMs and a cloud property composite based upon lower Earth orbit satellite retrievals are used here to estimate the global mean shortwave and longwave anisotropic factors.

#### *a. CERES ADMs*

The Edition 4 CERES ADMs (Su et al. 2015) are constructed using the CERES observations taken during the rotating azimuth plane (RAP) scan mode. In this mode, the instrument scans in elevation as it rotates in azimuth, thus acquiring radiance measurements from a wide range of viewing combinations. The CERES ADMs are derived for various scene types, which are defined using a combination of variables (e.g., surface type, cloud fraction, cloud optical depth, cloud phase, aerosol optical depth, precipitable water, lapse rate, etc). To provide accurate scene type information within CERES footprints, imager (Moderate Resolution Imaging Spectroradiometer (MODIS) on *Terra* and *Aqua*) cloud and aerosol retrievals (Minnis et al. 2010, 2011) are averaged over CERES footprints by accounting for the CERES point spread function (PSF, Smith 1994) and are used for scene type classification. Over a given scene type ( $\chi$ ), the CERES measured radiances are sorted into discrete angular bins. Averaged radiances ( $\hat{I}$ ) in all angular bins are calculated and all radiances in the upwelling directions are integrated to provide the ADM flux ( $\hat{F}$ ). The ADM anisotropic factors ( $R$ ) for scene type  $\chi$  are then calculated as:

$$R(\theta_0, \theta, \phi, \chi) = \frac{\pi \hat{I}(\theta_0, \theta, \phi, \chi)}{\int_0^{2\pi} \int_0^{\frac{\pi}{2}} \hat{I}(\theta_0, \theta, \phi, \chi) \cos\theta \sin\theta d\theta d\phi} = \frac{\pi \hat{I}(\theta_0, \theta, \phi, \chi)}{\hat{F}(\theta_0, \chi)}, \quad (6)$$

where  $\theta_0$  is the solar zenith angle,  $\theta$  is the CERES viewing zenith angle, and  $\phi$  is the relative azimuth angle between CERES and the solar plane.

#### *b. EPIC composite data*

As stated in the section above, anisotropy of the radiation field at the TOA was constructed for different scene types, which were defined using many variables including cloud

properties such as cloud fraction, cloud optical depth, and cloud phase (Loeb et al. 2005; Su et al. 2015). Although the EPIC L2 cloud product includes threshold-based cloud mask, which identifies the EPIC pixels as high confident clear, low confident clear, high confident cloudy, and low confident cloudy (Yang et al. 2018), the low resolution of EPIC imagery ( $24 \times 24 \text{ km}^2$ ) and its lack of infrared channels diminish its capability to identify clouds and to accurately retrieve cloud properties. As EPIC lacks the channels that are suitable for cloud size and phase retrievals (Meyer et al. 2016), two cloud optical depths are determined assuming the cloud phase is liquid or ice using constant cloud effective radius ( $14 \mu\text{m}$  for liquid and  $30 \mu\text{m}$  for ice) for cloudy EPIC pixels. These cloud properties are not sufficient to provide the scene type information necessary for ADM selections. Therefore, more accurate cloud property retrievals are needed to provide anisotropy characterizations to convert radiances to fluxes.

To accomplish this, we take advantage of the cloud property retrievals from multiple imagers on low Earth orbit (LEO) satellites and geostationary (GEO) satellites. The LEO satellite imagers include the MODerate-resolution Imaging Spectroradiometer (MODIS) on the Terra and Aqua satellites, the Visible Infrared Imaging Suite (VIIRS) on the Suomi-National Polar-orbiting Partnership satellite, and the Advanced Very High Resolution Radiometer (AVHRR) on the NOAA and MetOps platforms. The GEO imagers are on the Geostationary Operational Environmental Satellites (GOES), the Meteosat series, and Himawari-8 to provide semi-global coverage. All cloud properties were determined using a common set of algorithms, the Satellite Cloud and Radiation Property retrieval System (SatCORPS, Minnis et al. 2008b, 2016), based on the CERES cloud detection and retrieval system (Minnis et al. 2008a, 2010, 2011). Cloud properties from these LEO/GEO imagers are optimally merged together to provide a seamless global composite product at 5-km resolution by using an aggregated rating that considers five parameters (nominal satellite resolution, pixel time relative to the EPIC observation time, viewing zenith angle, distance from day/night terminator, and sun glint factor to minimize the usage of data taken in the glint region) and

selects the best observation at the time nearest to the EPIC measurements. About 80% of the LEO/GEO satellite overpass times are within 40 minutes of the EPIC measurements, while 96% are within two hours of the EPIC measurements. Most of the regions covered by GEO satellites (between around 50°S and 50°N) have a very small time difference, in the range of  $\pm 30$  minutes, because the availability of hourly GEO observations. The polar regions are also covered very well by polar orbiters. Thus, larger time differences are generally occurred over the 50° to 70° latitude regions. Given the temporal resolution of the currently available GEO/LEO satellites, this is the best collocation possible for those latitudes.

The global composite data are then remapped into the EPIC FOV by convolving the high-resolution cloud properties with the EPIC point spread function (PSF) defined with a half-pixel accuracy to produce the EPIC composite. As the PSF is sampled with half-pixel accuracy, the nominal spacing of the PSF grid is about the same size as in the global composite data. Thus, the accuracy of the cloud fraction in the EPIC composite is not degraded compared to the global composite (Khlopenkov et al. 2017). PSF-weighted averages of radiances and cloud properties are computed separately for each cloud phase, because the LEO/GEO cloud products are retrieved separately for liquid and ice clouds (Minnis et al. 2008b). Ancillary data (i.e. surface type, snow and ice map, skin temperature, precipitable water, etc.) needed for anisotropic factor selections are also included in the EPIC composite. These composite images are produced for each observation time of the EPIC instrument (typically 300 to 600 composites per month). Detailed descriptions of the method and the input used to generate the global and EPIC composites are provided in Khlopenkov et al. (2017).

Figure 3(a) shows an image from EPIC taken on May 15, 2017 at 12:17 UTC, the corresponding total cloud fraction (the sum of liquid and ice cloud fractions) from the EPIC composite is shown in 3(b). The liquid and ice cloud fraction, optical depth, and effective height are shown in Figure 3(c-h). For this case, most of the clouds are in the liquid phase. Optically thick liquid clouds with effective heights of 2 to 4 km are observed in the northern

Atlantic ocean and in the Arctic. Ice clouds with effective heights of 8 to 10 km are observed off the west coast of Africa and Europe.

*c. Calculating global daytime anisotropic factors*

To determine the global daytime mean anisotropic factors, we use the anisotropies characterized in the CERES ADMs and they are selected based upon the scene type information provided by the EPIC composite for every EPIC FOV. For a given EPIC FOV ( $j$ ), its anisotropic factor is determined based upon the Sun-EPIC viewing geometry and the scene identification information provided by the EPIC composite:

$$R_j(\theta_0, \theta^e, \phi^e, \chi^e) = \frac{\pi \hat{I}_j(\theta_0, \theta^e, \phi^e, \chi^e)}{\hat{F}_j(\theta_0, \chi^e)}, \quad (7)$$

where  $\theta^e$  is the EPIC viewing zenith angle,  $\phi^e$  is the relative azimuth angle between EPIC and the solar plane, and  $\chi^e$  is the scene identification from the EPIC composite. Here  $\hat{I}_j$  is the radiance from CERES ADMs and  $\hat{F}_j$  is the flux from CERES ADMs (see Eq. 6). To derive the global mean anisotropic factor, we follow the method developed by Su et al. (2018) and calculate the global daytime mean ADM radiance as:

$$\bar{\hat{I}} = \frac{\sum_{j=1}^N \hat{I}_j(\theta_0, \theta^e, \phi^e, \chi^e)}{N}. \quad (8)$$

To calculate the global mean ADM flux, we first grid the ADM flux ( $\hat{F}$ ) for each EPIC pixel into  $1^\circ$  latitude by  $1^\circ$  longitude bins ( $\hat{F}(\text{lat}, \text{lon})$ ). These gridded ADM fluxes are then weighted by *cosine* of latitude to provide the global daytime mean ADM flux:

$$\bar{\hat{F}} = \frac{\sum_{j=1}^M \hat{F}_j(\text{lat}, \text{lon}) \cos(\text{lat}_j)}{\sum \cos(\text{lat}_j)}. \quad (9)$$

The global mean anisotropic factor is calculated as:

$$\bar{R} = \frac{\pi \bar{\hat{I}}}{\bar{\hat{F}}}. \quad (10)$$

299 We use  $\overline{R_{sw}}$  and  $\overline{R_{lw}}$  to denote the mean SW and LW anisotropic factors. The mean SW  
 300 anisotropic factor is then used to convert the NISTAR SW unfiltered radiance to flux:

$$F_n^{sw} = \frac{\pi I_u^{sw}}{\overline{R_{sw}}}. \quad (11)$$

301 The LW flux is similarly derived from the following:

$$F_n^{lw} = \frac{\pi I_u^{lw}}{\overline{R_{lw}}}. \quad (12)$$

302 Figure 4 shows an example of SW and LW anisotropic factors for every EPIC FOV. The  
 303 SW anisotropic factors are generally smaller over clear than over cloudy oceanic regions.  
 304 Over land, however, the SW anisotropic factors are larger over clear regions than over cloudy  
 305 regions because of the hot spot effect, which leads to anisotropic factors greater than 1.6  
 306 over clear land regions at large viewing zenith angles. The LW anisotropic factors show  
 307 much less variability compared to the SW anisotropic factors, with limb darkening being  
 308 the dominant feature. The mean SW and LW anisotropic factors for this case are 1.275 and  
 309 1.041, respectively.

## 310 4. NISTAR shortwave and longwave flux

311 The temporal resolution of the NISTAR Level 1B data is one second, however, meaning-  
 312 ful changes in the data only occur over many shutter cycles (*each shutter cycle is 4 minutes*)  
 313 *due to the demodulation algorithm, which includes a box car filter having the width of a shut-*  
 314 *ter period. The filter reduces noise and rejects higher harmonics of the shutter frequency.*  
 315 *Following demodulation, significant instrument noise remains. Therefore, further averag-*  
 316 *ing in time over a minimum of 2 hours is recommended* to further reduce the noise levels  
 317 ([https://eosweb.larc.nasa.gov/project/dscovr/DSCOVNISTARDataQualityReport\\_V02.pdf](https://eosweb.larc.nasa.gov/project/dscovr/DSCOVNISTARDataQualityReport_V02.pdf)).  
 318 In this study, we use hourly radiances averaged from 4-hour running means as suggested by  
 319 the NISTAR instrument team. The hours that are coincident with the EPIC image times

are converted to fluxes using the global anisotropic factors calculated using the EPIC composites for scene identification. Figure 5 shows the hourly SW and LW fluxes derived from NISTAR for April (a) and July (b) 2017. For both months, the SW fluxes fluctuate around  $210 \text{ W m}^{-2}$ , with the difference between daily maximum and minimum as large as  $30 \text{ W m}^{-2}$ . The LW fluxes fluctuate around  $260 \text{ W m}^{-2}$ , and exhibit surprisingly large diurnal variations.

These NISTAR fluxes are compared to the CERES Synoptic radiative fluxes and clouds product (SYN1deg, Doelling et al. 2013), which provides hourly cloud properties and fluxes for each  $1^\circ$  latitude by  $1^\circ$  longitude. Within the SYN1deg data product, fluxes between CERES observations are inferred from hourly GEO visible and infrared imager measurements between  $60^\circ\text{S}$  and  $60^\circ\text{N}$  using observation-based narrowband-to-broadband radiance and radiance-to-flux conversion algorithms. However, the GEO narrowband channels have a greater calibration uncertainty than MODIS and CERES. Several procedures are implemented to ensure the consistency between the MODIS-derived and GEO-derived cloud properties, and between the CERES fluxes and the GEO-based fluxes. These include calibrating GEO visible radiances against the well-calibrated MODIS  $0.65 \mu\text{m}$  radiances by ray-matching MODIS and GEO radiances; applying similar cloud retrieval algorithms to derive cloud properties from MODIS and GEO observations; and normalizing GEO-based broadband fluxes to CERES fluxes using coincident measurements. Comparisons with broadband fluxes from Geostationary Earth Radiation Budget (GERB, Harries et al. 2005) indicate that SYN1deg hourly fluxes are able to capture the subtle diurnal flux variations. Comparing with the GERB fluxes, the bias of the SYN SW fluxes is  $1.3 \text{ W m}^{-2}$ , the monthly regional all-sky SW flux RMS error is  $3.5 \text{ W m}^{-2}$ , and the daily regional all-sky SW flux RMS error is  $7.8 \text{ W m}^{-2}$  (Doelling et al. 2013). These uncertainties could be overestimated, as the GERB domain has a disproportionate number of strong diurnal cycle regions as compared with the globe.

To account for the missing energy from the daytime portion that is not observed by the NISTAR ( $A_h$  in Figure 1b), *and the energy from the nighttime sliver that are within the DSCOVR view ( $A_d$  in Figure 1b, only applicable to LW flux)*, the hourly gridded SYN fluxes

are integrated by considering only the grid boxes that are visible to NISTAR to produce the global mean daytime fluxes that are comparable to those from the NISTAR measurements:

$$\overline{F_{syn}} = \frac{\sum F_j \cos(lat_j) \omega_j}{\sum \cos(lat_j) \omega_j}. \quad (13)$$

Here  $F_j$  is the gridded hourly CERES SYN fluxes,  $lat$  is the latitude, and  $\omega$  indicates whether a grid box is visible to NISTAR (=1 when visible, =0 when not visible). Figure 6a) shows an example of the gridded SYN SW fluxes at 13 UTC on February 1, 2017. SW fluxes for the daytime grid boxes are shown in color, while all nighttime grid boxes are shown in white. Figure 6b) shows the daytime areas (in red) and the nighttime areas (in grey) visible to the NISTAR view. Daytime areas of northern high latitude and North America are not within the NISTAR view and are therefore not included in the comparison with the NISTAR fluxes, and the nighttime slivers in the southern high latitude of Indian Ocean and Pacific Ocean are included in the LW flux comparison with the NISTAR.

Figure 7 compares the SW fluxes from NISTAR with those from CERES SYN1deg product integrated for the NISTAR view (Eq. 19) for April (a) and July (b) 2017. The CERES SW fluxes oscillate around 200 Wm<sup>-2</sup> and 195 Wm<sup>-2</sup> for April and July, whereas the NISTAR counterparts are about 10 to 20 Wm<sup>-2</sup> greater. The maxima and minima of SW fluxes from NISTAR align well with those from CERES, though the differences between daily maximum and minimum from NISTAR appear to be larger than those from CERES. The diurnal variations of SW flux derived from EPIC showed a much better agreement with those from CERES (Su et al. 2018). The exact cause for these larger diurnal variations from NISTAR SW flux is not known. LW flux comparisons are shown in Figure 8. The daily maximum-minimum LW differences from CERES are typically less than 15 Wm<sup>-2</sup> and exhibit small day-to-day and month-to-month variation. However, the daily maximum-minimum LW differences from NISTAR can vary from 10 Wm<sup>-2</sup> to 50 Wm<sup>-2</sup>. These larger than expected variability of NISTAR LW fluxes are due to the fact that noise and offset variabilities from both the NISTAR total and SW channel are present in the NISTAR LW radiances. The NISTAR LW fluxes are consistently greater than CERES LW fluxes by about



373 10 to 20  $\text{Wm}^{-2}$  in April. The LW fluxes agree better for July, but the NISTAR LW fluxes  
374 show larger diurnal variations than the CERES fluxes.

375 Figure 9 compares the SW and LW fluxes from CERES SYN1deg product with those  
376 from NISTAR at all coincident hours of 2017. The mean SW fluxes are 203.7  $\text{Wm}^{-2}$  and  
377 217.0  $\text{Wm}^{-2}$ , respectively, for CERES and NISTAR, and the RMS error is 14.6  $\text{Wm}^{-2}$  (Fig-  
378 ure 9a). The mean LW fluxes are 246.0  $\text{Wm}^{-2}$  and 252.8  $\text{Wm}^{-2}$  for CERES and NISTAR,  
379 and the RMS error is 10.5  $\text{Wm}^{-2}$  (Figure 9b). Tables 3 and 4 summarize the flux com-  
380 parisons between NISTAR and CERES for all months of 2017. The NISTAR SW fluxes  
381 are consistently greater than those from CERES SYN1deg by about 3.4% to 7.8%, and the  
382 NISTAR LW fluxes are also greater than those from CERES SYN1deg by 1.0% to 5.0%.  
383 Furthermore, the SW fluxes from NISTAR are highly correlated (correlation coefficient of  
384 about 0.89) with those from CERES SYN1deg, but the correlation for the LW fluxes are  
385 rather low (correlation coefficient is about 0.38).

386 NISTAR fluxes derived at the EPIC image times are averaged into daily means and are  
387 compared with the daily means from CERES SYN1deg using concurrent hours (Figure 10).  
388 The NISTAR SW fluxes are consistently higher than those from CERES by about 10 to 15  
389  $\text{Wm}^{-2}$ . CERES SW fluxes show a strong annual cycle, which is driven by the incident solar  
390 radiation that is affected by the Earth-Sun distance. This annual cycle is also evident in the  
391 NISTAR SW fluxes, albeit the fluxes during the period from April to August are flatter than  
392 those from CERES. The NISTAR LW fluxes are greater than those from CERES except  
393 during the boreal summer months, with the largest difference of 10  $\text{Wm}^{-2}$  in February and  
394 the smallest difference of a few  $\text{Wm}^{-2}$  during the boreal summer months. The CERES LW  
395 fluxes show an annual cycle of about 10  $\text{Wm}^{-2}$ , with the largest LW fluxes occurring during  
396 the boreal summer when the vast land masses of the northern hemisphere are warmer than  
397 during the other seasons. The annual cycle of the NISTAR LW fluxes shows less seasonal  
398 variation. From April to October, the NISTAR LW fluxes oscillate around 255  $\text{Wm}^{-2}$ , and  
399 oscillate around 250  $\text{Wm}^{-2}$  for other months. Additionally, the CERES LW fluxes exhibit

much smaller day-to-day variations than their NISTAR counterparts. Note some of the variations of daily mean fluxes shown in Figure 10 are due to temporal sampling changes when data transmissions encountered difficulties and/or during spacecraft maneuvers.

## 5. Conclusions and discussions

The SW radiances included in the NISTAR L1B data are filtered radiances and the effect of the filter transmission must be addressed before these measurements can be used to derive any meaningful fluxes. A comprehensive spectral radiance database has been developed to investigate the relationship between filtered and unfiltered radiances using theoretically derived values simulated for typical Earth scenes and the NISTAR spectral transmission functions. The ratio between filtered and unfiltered SW radiances is very stable, varying less than 0.3% for the scenes and the Sun-viewing geometries included in the database. The mean ratio of 0.8690 is used to derive the unfiltered SW radiance from the NISTAR L1B filtered SW radiance measurements.

To convert these unfiltered radiances into fluxes, the anisotropy of the radiance field must be taken into account. We use the scene-type dependent CERES angular distribution models to characterize the global SW and LW anisotropy. These global anisotropies are calculated based upon the anisotropies for each EPIC pixel. To accurately account for the anisotropy for each EPIC pixel, an EPIC composite was developed which includes all information needed for angular distribution model selections. The EPIC composite includes cloud property retrievals from multiple imagers on LEO and GEO satellites. Cloud properties from these LEO and GEO imagers are optimally merged together to provide a global composite product at 5-km resolution by using an aggregated rating that considers several factors and selects the best observation at the time nearest to the EPIC measurements. The global composite data are then remapped into the EPIC FOV by convolving the high-resolution cloud properties with the EPIC PSF to produce the EPIC composite. PSF-weighted averages of radiances

and cloud properties are computed separately for each cloud phase, and ancillary data needed for anisotropic factor selections are also included in the EPIC composite.

These global anisotropies are applied to the NISTAR radiances to produce the global daytime SW and LW fluxes and they are validated against the CERES Synoptic  $1^\circ$  latitude by  $1^\circ$  longitude flux product. Only the grid boxes that are visible to the NISTAR view are integrated to produce the global mean daytime fluxes that are comparable to the fluxes from the NISTAR measurements. The NISTAR SW fluxes are consistently greater than those from CERES SYN1deg by  $10 \text{ Wm}^{-2}$  to  $15 \text{ Wm}^{-2}$  (3.3% to 7.8%), but these two SW flux datasets are highly correlated indicating that the diurnal and seasonal variations of the SW fluxes are fairly similar for both of them. The NISTAR LW fluxes are also greater than those from CERES SYN1deg, but the magnitude of the difference has larger month-to-month variations than that for the SW fluxes. The largest difference of about  $14 \text{ Wm}^{-2}$  ( $\sim 5.5\%$ ) occurred in April 2017 and the smallest difference of about  $\sim 4 \text{ Wm}^{-2}$  ( $\sim 1.6\%$ ) occurred during July. Furthermore, the NISTAR LW fluxes have very low correlations with the CERES LW fluxes. NISTAR LW fluxes exhibit a nearly flat annual variation, whereas the CERES LW fluxes exhibit a distinct annual cycle with the highest LW flux occurs in July when the vast northern hemisphere land masses are warmest. The NISTAR LW fluxes also exhibit unrealistically large day-to-day variations.

The SW flux discrepancy between NISTAR and CERES is caused by: 1) CERES instrument calibration uncertainty, 2) CERES flux algorithm uncertainty, 3) NISTAR instrument measurement uncertainty, and 4) NISTAR flux algorithm uncertainty. The CERES SW channel calibration uncertainty is 1% ( $1\sigma$ , McCarthy et al. 2011; Priestley et al. 2011; Loeb et al. 2018), which corresponds to about  $2.1 \text{ Wm}^{-2}$  for daytime mean SW fluxes. The CERES algorithm uncertainty includes radiance-to-flux conversion error, which is  $1.0 \text{ Wm}^{-2}$  according to Su et al. (2015), and diurnal correction uncertainty, which is estimated to be  $1.9 \text{ Wm}^{-2}$  when Terra and Aqua are combined (Loeb et al. 2018). The NISTAR SW channel measurement uncertainty is 2.1%, which corresponds to  $4.4 \text{ Wm}^{-2}$ . The NISTAR algorithm

uncertainty is essentially the radiance-to-flux conversion error. The estimation of this error source is not readily available given the unique NISTAR viewing perspective. However, if we assume the discrepancy between EPIC derived SW flux and CERES SW flux (Su et al. 2018) is also from uncertainty sources 1) and 2) listed above, plus the EPIC calibration, narrowband-to-broadband conversion, and radiance-to-flux conversion for EPIC, then we can deduce that the radiance-to-flux conversion uncertainty for the NISTAR viewing geometry should be less than  $2 \text{ Wm}^{-2}$ . Thus the total difference expected from these uncertainty sources should be  $(2.1^2 + 1.9^2 + 1.0^2 + 4.4^2 + 2.0^2)^{1/2} = 5.7 \text{ Wm}^{-2}$ .

Similarly, the LW flux discrepancy between NISTAR and CERES is due to the same sources of error. *The daytime CERES LW flux uncertainty from calibration is  $2.5 \text{ Wm}^{-2}$  ( $1\sigma$ , Loeb et al. 2009).* The CERES LW radiance-to-flux conversion error is about  $0.75 \text{ Wm}^{-2}$  (Su et al. 2015), and diurnal correction uncertainty is estimated to be  $2.2 \text{ Wm}^{-2}$  (Loeb et al. 2018). However, the CERES LW ADMs were developed without taking the relative azimuth angle into consideration, which has little impact on the CERES LW flux accuracy because of its Sun-synchronous orbit. Given that the NISTAR only views the Earth from the backscattering angles, the LW flux uncertainty due to radiance-to-flux conversion could be larger for the clear-sky footprints (Minnis et al. 2004). As the clear-sky occurrences are small at the EPIC footprint size level, our best *estimate* of this uncertainty is no more than  $0.4 \text{ Wm}^{-2}$ . The calibration uncertainty for NISTAR LW is deduced from the calibration uncertainties of total and SW channels. The total channel calibration uncertainty is 1.5%, which is about  $6.8 \text{ Wm}^{-2}$  assuming the total radiative energy of  $450 \text{ Wm}^{-2}$ . The SW channel measurement uncertainty is  $4.4 \text{ Wm}^{-2}$ . The resulting LW channel measurement uncertainty is thus equal to  $(6.8^2 + 4.4^2)^{1/2} = 8.1 \text{ Wm}^{-2}$ . Although no direct estimation of the radiance-to-flux conversion uncertainty for LW is available, we do not expect that it exceeds its SW counterpart of  $2.0 \text{ Wm}^{-2}$ . *Thus the total difference expected from these uncertainty sources should be  $(2.5^2 + 0.75^2 + 0.4^2 + 2.2^2 + 8.1^2 + 2.0^2)^{1/2} = 9.1 \text{ Wm}^{-2}$ .*

The uncertainty sources listed above can explain part of the SW flux differences and

all of the LW flux differences between CERES and NISTAR. The error sources related to NISTAR are preliminary and are under careful evaluation. Although the LW flux differences between CERES and NISTAR are within the uncertainty estimation, the correlation between NISTAR and CERES is rather low, about 0.38. This is because the NISTAR LW radiance is derived as the difference between total channel radiance and SW channel radiance, thus noise and offset variability of both the NISTAR total and SW channels are present in the NISTAR LW fluxes. As a result, more variability is expected in the LW data which leads to the low correlation. *Although the noise level present in the NISTAR measurements prevent the production of high frequency SW flux, the current 4-hour running mean fluxes are highly correlated with the CERES product. The NISTAR SW flux can be used to test the diurnal variations of SW flux in the high-temporal resolution model outputs from the Coupled Model Intercomparison Project. Furthermore, the spectral ratio information from NISTAR presents a new way to evaluate the models and opens a new perspective on exoplanet observations (Carlson et al. 2019).*

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## REFERENCES

- 501 Carlson, B. E., A. A. Lacis, C. Colose, A. Marshak, W. Su, and S. Lorentz, 2019: Spectral  
502 signature of the biosphere: NISTAR finds it in our solar system from the Lagrangia L-1  
503 point. *Geophys. Res. Lett.*, doi:10.1029/2019GL083736.
- 504 Doelling, D. R., N. G. Loeb, D. F. Keyes, M. L. Nordeen, D. Morstad, B. A. Wielicki, D. F.  
505 Young, and M. Sun, 2013: Geostationary enhanced temporal interpolation for CERES flux  
506 products. *J. Atmos. Oceanic Technol.*, **30**, 1072–1090, doi:10.1175/JTECH-D-12-00136.1.
- 507 Harries, J. E., J. E. Russell, J. A. Hanafin, H. Brindley, J. Futyan, J. Rufus, and coauthors,  
508 2005: The geostationary earth radiation budget project. *Bull. Am. Meteor. Soc.*, **86**, 945–  
509 960.
- 510 House, F. B., A. Gruber, G. E. Hunt, and A. T. Mecherikunnel, 1986: History of satellite  
511 missions and measurements of the Earth radiation budget (1957-1984). *Rev. Geophys.*, **24**,  
512 357–377.
- 513 Kato, S., N. G. Loeb, and K. Rutledge, 2002: Estimate of top-of-atmosphere albedo for a  
514 molecular atmosphere over ocean using Clouds and the Earth’s Radiant Energy System  
515 measurements. *J. Geophys. Res.*, **107** (D19), doi:10.1029/2001JD001309.
- 516 Kato, S., et al., 2011: Improvements of top-of-atmosphere and surface irradiance computa-  
517 tion with CALIPSO-, and MODIS-derived cloud and aerosol properties. *J. Geophys. Res.*,  
518 **116** (D19209), D19 209, doi:10.1029/2011JD016050.
- 519 Khlopenkov, K., D. Duda, M. Thieman, P. Minnis, W. Su, and K. Bedka, 2017: Devel-  
520 opment of multi-sensor global cloud and radiance composites for Earth radiation budget

monitoring from DSCOVR. *Remote sensing of clouds and the atmosphere XXII*, A. Comeron, E. I. Kassianov, K. Schafer, R. H. Picard, and K. Weber, Eds., Warsaw, Poland, Proc. SPIE 10424, Vol. 10424K (2 October 2017), doi:10.1117/12.2278645.

Loeb, N. G., S. Kato, K. Loukachine, and N. Manalo-Smith, 2005: Angular distribution models for top-of-atmosphere radiative flux estimation from the clouds and the earth’s radiant energy system instrument on the terra satellite. part I: Methodology. *J. Atmos. Oceanic Technol.*, **22**, 338–351.

Loeb, N. G., J. M. Lyman, G. C. Johnson, R. P. Allan, D. R. Doelling, T. Wong, B. J. Soden, and G. L. Stephens, 2012: Observed changes in top-of-the-atmosphere radiation and upper-ocean heating consistent within uncertainty. *Nature Geosci.*, **5**, 110–113, doi:10.1038/NGEO1375.

Loeb, N. G. and N. Manalo-Smith, 2005: Top-of-atmosphere direct radiative effect of aerosols over global oceans from merged CERES and MODIS observations. *J. Climate*, **18**, 3506–3526.

Loeb, N. G., N. Manalo-Smith, W. Su, M. Shankar, and S. Thomas, 2016: CERES top-of-atmosphere Earth radiation budget climate data record: Accounting for in-orbit changes in instrument calibration. *Remote Sens.*, **8** (182), doi:10.3390/rs8030182.

Loeb, N. G., K. J. Priestley, D. P. Kratz, E. B. Geier, R. N. Green, B. A. Wielicki, P. O. Hinton, and S. K. Nolan, 2001: Determination of unfiltered radiances from the Clouds and the Earth’s Radiant Energy System instrument. *J. Appl. Meteor.*, **40**, 822–835.

Loeb, N. G. and G. L. Schuster, 2008: An observational study of the relationship between cloud, aerosol and meteorology in broken low-level cloud conditions. *J. Geophys. Res.*, **113** (D14214), D14 214, doi:10.1029/2007JD009763.

Loeb, N. G., B. A. Wielicki, D. R. Doelling, G. L. Smith, D. F. Keyes, S. Kato, N. Manalo-

Smith, and T. Wong, 2009: Towards optimal closure of the Earth’s top-of-atmosphere radiation budget. *J. Climate*, **22**, 748–766, doi:10.1175/2008JCLI2637.1.

Loeb, N. G., et al., 2018: Clouds and the Earth’s Radiant Energy System (CERES) Energy Balanced and Filled (EBAF) Top-of-Atmosphere (TOA) Edition-4.0 Data Product. *J. Climate*, **31**, 895–918, doi:10.1175/JCLI-D-17-0208.1.

McCarthy, J. M., H. Bitting, T. A. Evert, M. E. Frink, T. R. Hedman, P. Skaguchi, and M. folkman, 2011: A summary of the performance and long-term stability of the pre-launch radiometric calibration facility for the Clouds and the Earth’s Radiant Energy System (CERES) instruments. *2011 IEEE International Geoscience and Remote Sensing Symposium*, 1009–1012, doi:10.1109/IGARSS.2011.6049304.

Meyer, K., Y. Yang, and S. Platnick, 2016: Uncertainties in cloud phase and optical thickness retrievals from the Earth Polychromatic Imaging Camera (EPIC). *Atmos. Meas. Tech.*, **9**, 1785–1797, doi:10.5194/amt-9-1785-2016.

Minnis, P., K. Bedka, Q. Z. Trepte, C. R. Yost, S. T. Bedka, B. Scarino, K. V. Khlopenkov, and M. M. Khaiyer, 2016: *A consistent long-term cloud and clear-sky radiation property dataset from the Advanced Very High Resolution Radiometer (AVHRR). Climate Algorithm Theoretical Basis Document (C-ATBD)*. CDRP-ATBD-0826 Rev 1–NASA,NOAA CDR Program, doi:10.7289/V5HT2M8T.

Minnis, P., A. V. Gambheer, and D. R. Doelling, 2004: Azimuthal anisotropy of longwave and infrared window radiances from the Clouds and the Earth’s Radiant Energy System on the Tropical Rainfall Measuring Mission on Terra satellites. *J. Geophys. Res.*, **109** (D08202), doi:10.1029/2003JD004471.

Minnis, P., et al., 2008a: Cloud detection in nonpolar regions for CERES using TRMM VIRS and TERRA and AQUA MODIS data. *IEEE Trans. Geosci. Remote Sensing*, **46** (11), 3857–3884.



570 Minnis, P., et al., 2008b: Near-real time cloud retrievals from operational and research  
571 meteorological satellites. *Proc. SPIE 7108, Remote Sens. Clouds Atmos. XIII*, Cardiff,  
572 Wales, UK, doi:10.1117/12.800344.

573 Minnis, P., et al., 2010: CERES Edition 3 cloud retrievals. *13th Conference on Atmospheric*  
574 *Radiation*, Am. Meteorol. Soc., Oregon, Portland.

575 Minnis, P., et al., 2011: CERES Edition-2 cloud property retrievals using TRMM VIRS  
576 and TERRA and AQUA MODIS data, Part I: Algorithms. *IEEE Trans. Geosci. Remote*  
577 *Sensing*, **49** (11), 4374–4400, doi:10.1109/TGRS.2011.2144601.

578 Pincus, R., C. P. Batstone, R. J. P. Hofmann, K. E. Taylor, and P. J. Glecker, 2008: Evalu-  
579 ating the present-day simulation of clouds, precipitation, and radiation in climate models.  
580 *J. Geophys. Res.*, **113** (D14209), D14 209, doi:10.1029/2007JD009334.

581 Priestley, K. J., et al., 2011: Radiometric performance of the CERES Earth radiation budget  
582 climate record sensors on the EOS Aqua and Terra spacecraft through April 2007. *J.*  
583 *Atmos. Oceanic Technol.*, **28**, 3–21, doi:10.1175/2010JTECHA1521.1.

584 Quaas, J., O. Boucher, N. Bellouin, and S. Kinne, 2008: Satellite-based estimate of the  
585 direct and indirect aerosol climate forcing. *J. Geophys. Res.*, **113** (D05204), D05 204,  
586 doi:10.1029/2007JD008962.

587 Satheesh, S. K. and V. Ramanathan, 2000: Large differences in tropcial aerosol forcing at  
588 the top of the atmosphere and earth’s surface. *Nature*, **405**, 60–63.

589 Smith, G. L., 1994: Effects of time response on the point spread function of a scanning  
590 radiometer. *Appl. Opt.*, **33**, 7031–7037.

591 Stephens, G. L., et al., 2012: An update on Earth’s energy balance in light of the latest  
592 global observations. *Nature Geosci.*, **5**, 691–696, doi:10.1038/NGEO1580.

- 593 Su, W., A. Bodas-Salcedo, K.-M. Xu, and T. P. Charlock, 2010a: Comparison of the trop-  
594 ical radiative flux and cloud radiative effect profiles in a climate model with Clouds and  
595 the Earth’s Radiant Energy System (CERES) data. *J. Geophys. Res.*, **115** (D01105),  
596 D01 105, doi:10.1029/2009JD012490.
- 597 Su, W., J. Corbett, Z. A. Eitzen, and L. Liang, 2015: Next-generation angular distribution  
598 models for top-of-atmosphere radiative flux calculation from the CERES instruments:  
599 Methodology. *Atmos. Meas. Tech.*, **8**, 611–632, doi:10.5194/amt-8-611-2015.
- 600 Su, W., N. G. Loeb, G. L. Schuster, M. Chin, and F. G. Rose, 2013: Global all-sky shortwave  
601 direct radiative forcing of anthropogenic aerosols from combined satellite observations and  
602 GOCART simulations. *J. Geophys. Res.*, **118**, 1–15, doi:10.1029/2012JD018294.
- 603 Su, W., N. G. Loeb, K. Xu, G. L. Schuster, and Z. A. Eitzen, 2010b: An estimate of aerosol  
604 indirect effect from satellite measurements with concurrent meteorological analysis. *J.*  
605 *Geophys. Res.*, **115** (D18219), D18 219, doi:10.1029/2010JD013948.
- 606 Su, W., et al., 2018: Determining the shortwave radiative flux from Earth Polychromatic  
607 Imaging Camera. *J. Geophys. Res.*, **123**, doi:10.1029/2018JD029390.
- 608 Trenberth, K. E., J. T. Fasullo, and J. Kiehl, 2009: Earth’s global energy budget. *Bull. Am.*  
609 *Meteor. Soc.*, **90**, 311–323, doi:10.1175/2008BAMS2634.1.
- 610 Wang, H. and W. Su, 2013: Evaluating and understanding top of the atmosphere cloud  
611 radiative effects in Intergovernmental Panel on Climate Change (IPCC) fifth assessment  
612 report (AR5) clouded model intercomparison project phase 5 (CMIP5) models using  
613 satellite observations. *J. Geophys. Res.*, **118**, 1–17, doi:10.1029/2012JD018619.
- 614 Wielicki, B. A., B. R. Barkstrom, E. F. Harrison, R. B. Lee, G. L. Smith, and J. E. Cooper,  
615 1996: Clouds and the Earth’s Radiant Energy System (CERES): An Earth Observing  
616 System experiment. *Bull. Amer. Meteor. Soc.*, **77**, 853–868.

- 617 Wild, M., D. Folini, C. Schar, N. G. Loeb, E. G. Dutton, and G. Konig-Langlo, 2013:  
618 The global energy balance from a surface perspective. *Clim. Dyn.*, **40**, 3107–3134, doi:  
619 10.1007/s00382-012-1569-8.
- 620 Yang, Y., et al., 2018: Cloud Products from the Earth Polychromatic Imaging Camera  
621 (EPIC): Algorithms and Initial Evaluaiton. *Atmos. Meas. Tech. Discussions*, 1–23, doi:  
622 10.5194/amt-2018-316.
- 623 Zhang, J., S. A. Christopher, L. A. Remer, and Y. J. Kaufman, 2005: Shortwave aerosol  
624 radiative forcing over cloud-free oceans from Terra: 2. Seasonal and global distributions.  
625 *J. Geophys. Res.*, **110 (D10S24)**, D10S24, doi:10.1029/2004JD005009.

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TABLE 1. Summary of the cases included in the spectral radiance database. AOD is for aerosol optical depth, COD is for cloud optical depth.

Clear				
	AOD	Aerosol type	Surface	
Ocean	8	6	4	
Land	8	4	15	
Snow	5	2	5	
Cloudy				
	COD	Cloud type	Surface	Atmosphere
Ocean	7	4 liquid and 3 ice	4	4
Land	7	4 liquid and 3 ice	15	1

TABLE 2. Mean ratio and standard deviation (in parenthesis) of filtered radiance to unfiltered radiance for SW and NIR bands over different scene types.

	SW ratio (standard deviation $\times$ 1000)					
	0.0	29.0	41.4	60.0	75.5	85.0
Clear Ocean	0.8659(1.0)	0.8660(1.0)	0.8661(1.1)	0.8664(1.2)	0.8669(1.0)	0.8674(0.8)
Clear Land	0.8694(0.6)	0.8693(0.6)	0.8692(0.6)	0.8690(0.5)	0.8687(0.5)	0.8685(0.8)
Clear Snow	0.8689(0.1)	0.8689(0.1)	0.8689(0.2)	0.8688(0.2)	0.8688(0.3)	0.8687(0.4)
Cld Ocean	0.8687(1.0)	0.8687(1.0)	0.8688(0.9)	0.8688(0.8)	0.8688(0.7)	0.8687(0.6)
Cld Land	0.8694(0.4)	0.8693(0.3)	0.8693(0.3)	0.8692(0.3)	0.8690(0.4)	0.8689(0.5)
	NIR ratio (standard deviation $\times$ 1000)					
	0.0	29.0	41.4	60.0	75.5	85.0
Clear Ocean	0.8293(23.1)	0.8270(24.0)	0.8253(25.5)	0.8235(28.3)	0.8238(28.4)	0.8229(26.4)
Clear Land	0.8790(9.6)	0.8777(10.4)	0.8764(10.7)	0.8730(10.8)	0.8663(10.1)	0.8501(12.4)
Clear Snow	0.8360(1.7)	0.8360(1.8)	0.8361(1.9)	0.8363(2.1)	0.8370(2.8)	0.8365(6.0)
Cld Ocean	0.8557(3.2)	0.8555(2.6)	0.8562(2.4)	0.8567(3.1)	0.8565(4.4)	0.8539(7.9)
Cld Land	0.8627(8.2)	0.8624(7.8)	0.8621(7.3)	0.8613(6.2)	0.8598(4.8)	0.8566(6.2)

TABLE 3. SW flux comparisons between NISTAR and CERES SYN1deg for all coincident observations of 2017.  $F_n$  is the NISTAR flux (in  $\text{Wm}^{-2}$ ),  $F_s$  is the SYN flux (in  $\text{Wm}^{-2}$ ), and the root mean square (RMS) error between them (in  $\text{Wm}^{-2}$ ).

	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec
$F_s$	—	208.1	203.4	199.8	201.0	200.2	194.4	193.0	198.7	208.9	221.6	228.2
$F_n$	—	218.5	215.4	211.5	214.1	213.5	209.2	208.7	211.2	222.8	235.1	240.0
$RMS$	—	11.9	14.0	12.9	14.0	14.6	16.0	16.8	13.9	15.5	14.5	14.0

TABLE 4. LW flux comparisons between NISTAR and CERES SYN1deg for all coincident observations of 2017.  $F_n$  is the NISTAR flux (in  $\text{Wm}^{-2}$ ),  $F_s$  is the SYN flux (in  $\text{Wm}^{-2}$ ), and the root mean square (RMS) error between them (in  $\text{Wm}^{-2}$ ).

	Jan	Feb	Mar	Apr	May	Jun	Jul	Aug	Sep	Oct	Nov	Dec
$F_s$	—	242.0	241.1	243.0	246.3	249.1	251.5	248.9	245.5	242.9	239.8	240.6
$F_n$	—	253.1	248.1	257.7	255.8	255.2	255.6	253.2	255.5	253.5	250.4	253.3
$RMS$	—	13.4	10.0	16.0	11.5	10.3	8.7	10.0	12.2	12.5	12.4	14.4



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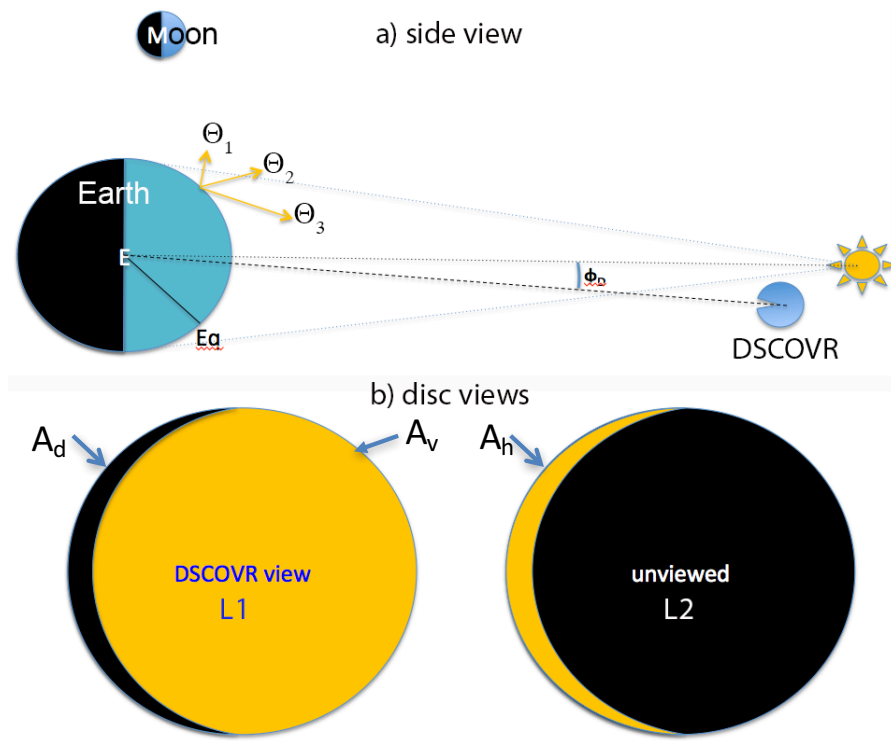


FIG. 1. Schematic of a) Earth-Sun-DSCOVR geometry and b) Earth disc that are visible to the L1 DSCOVR view (left with an area fraction of  $A_t$ ) and to the L2 view (right). The golden area on the left shows the daytime area fraction ( $A_v$ ) that are visible to DSCOVR, the black area on the left shows the night portion ( $A_d$ ) that are within the DSCOVR view, and the golden area on the right is the daytime portion ( $A_h$ ) missed by the DSCOVR. Not to scale.

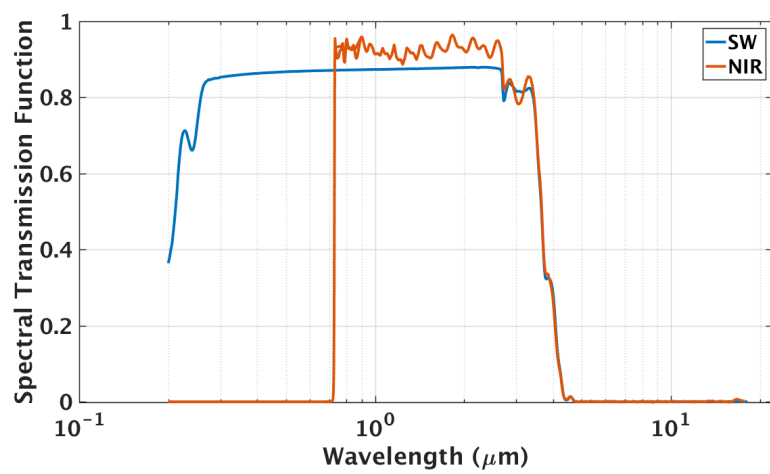


FIG. 2. NISTAR SW and NIR spectral transmission function.

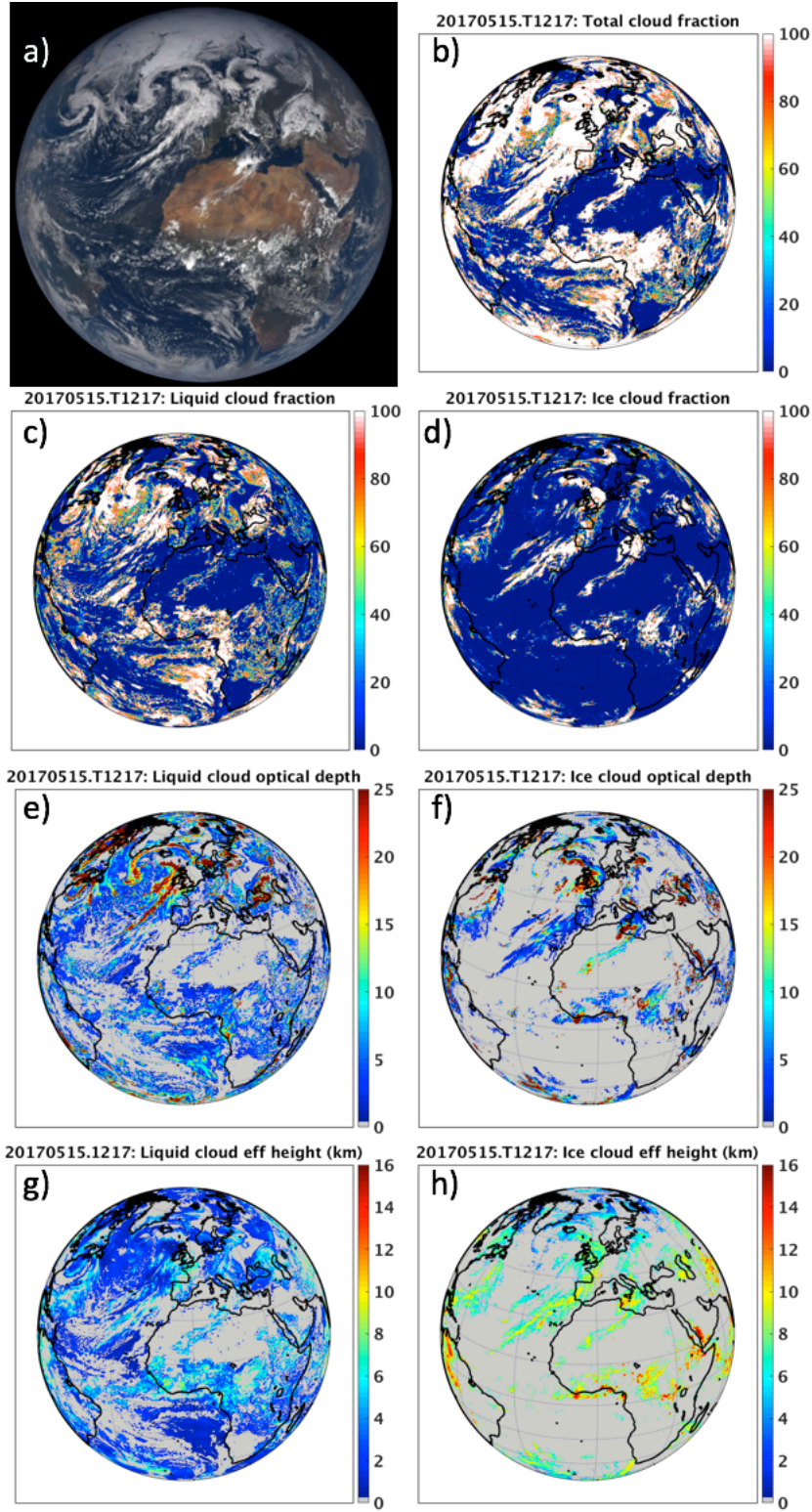


FIG. 3. EPIC RGB image for May 15, 2017 at 12:17 UTC (a), and the corresponding total cloud fraction (b, in %). Liquid and ice cloud fractions are shown in (c) and (d), liquid and ice cloud optical depths are shown in (e) and (f), and liquid and ice cloud effective height (in  $km$ ) are shown in (g) and (h). (b) to (h) are all derived from the EPIC composite.

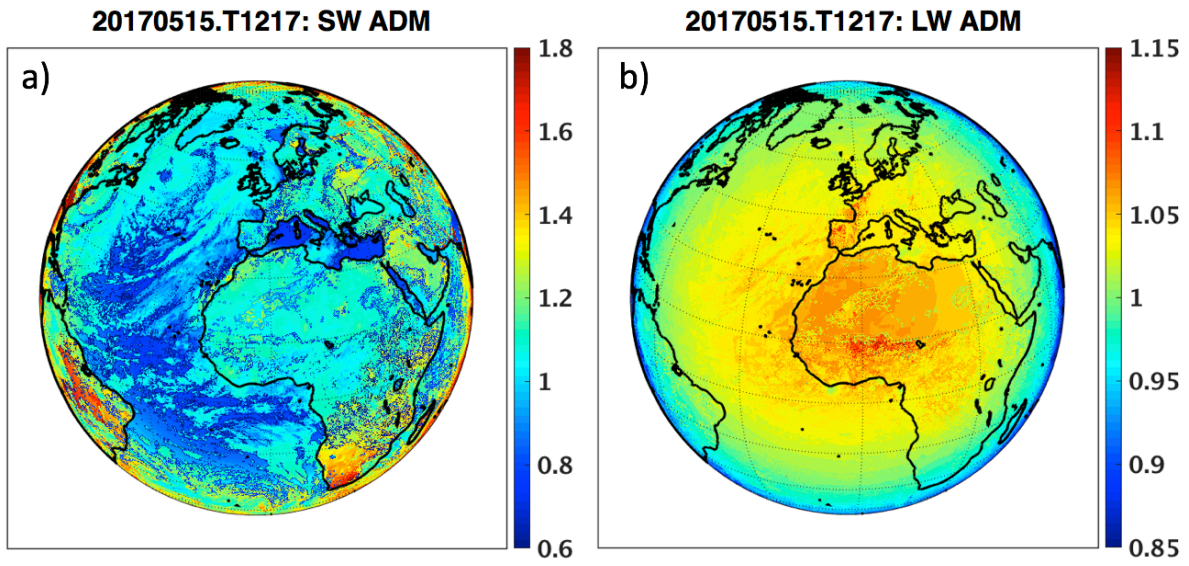


FIG. 4. SW anisotropic factors (a) and LW anisotropic factors (b) derived from the CERES ADMs using the EPIC composite for scene identification for May 15, 2017 at 12:17 UTC.

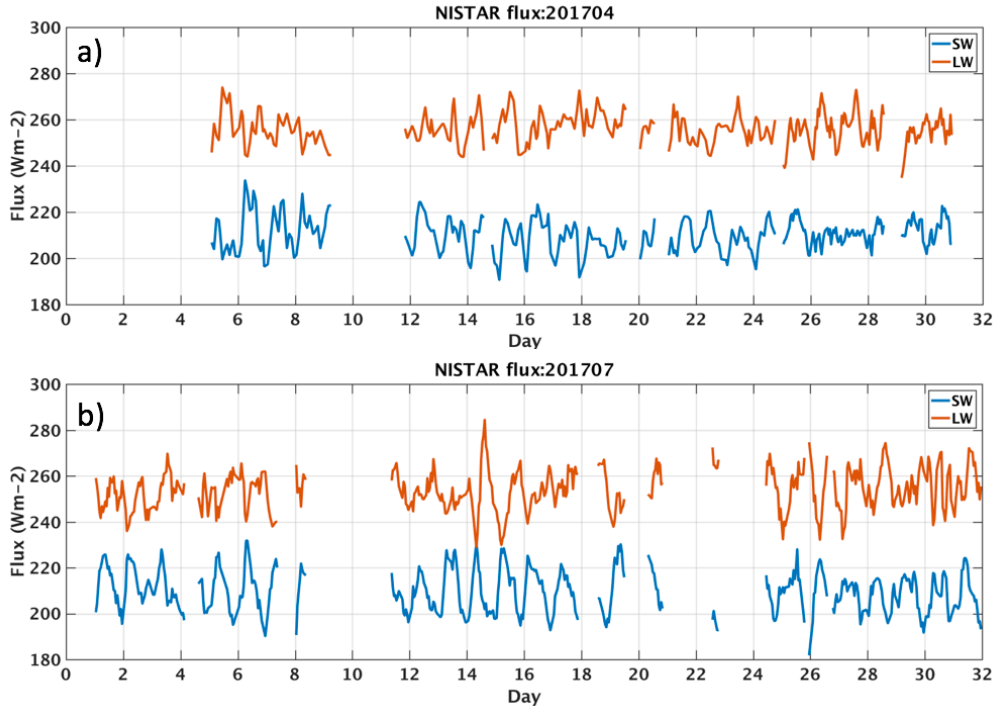


FIG. 5. SW flux (blue) and LW flux (red) derived from NISTAR measurements for April (a) and July (b), 2017.

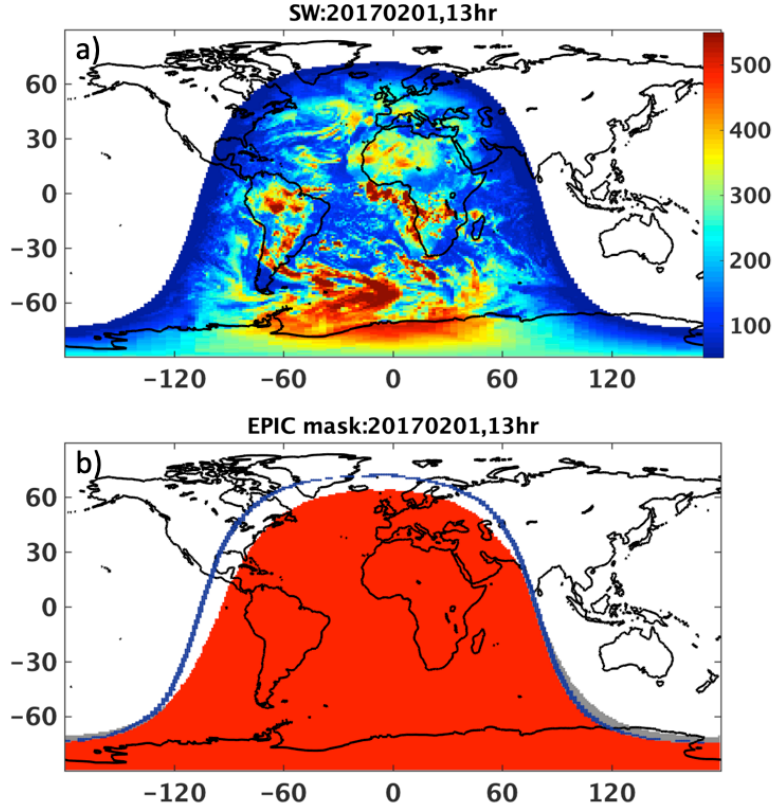


FIG. 6. An example of the daytime SW flux distributions from CERES SYN1deg product at 13 UTC on February 1, 2017 (a), and the corresponding daytime areas (in red) and nighttime areas (in grey) that are visible to NISTAR and the terminator boundary (in blue) (b).



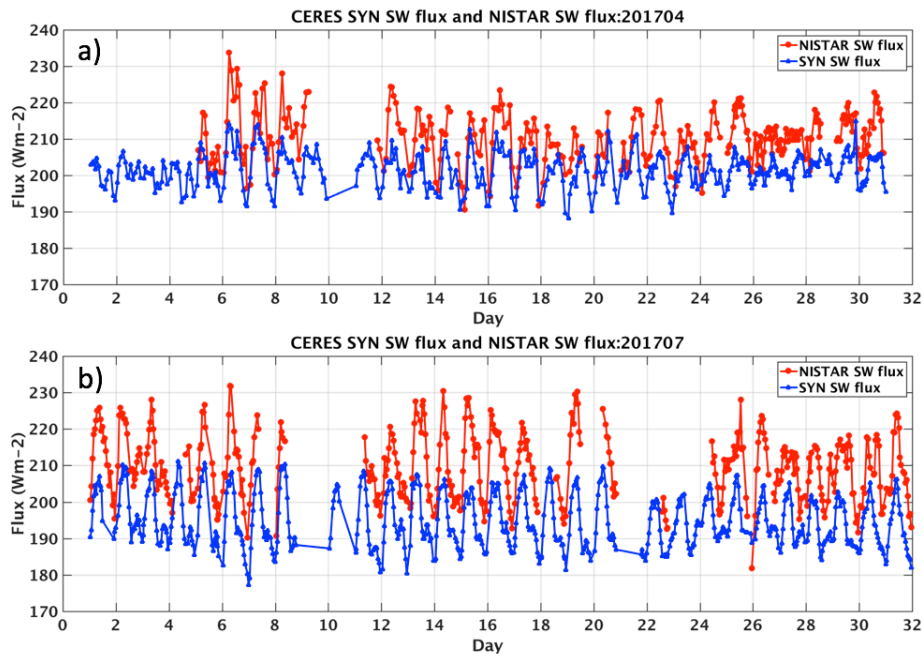


FIG. 7. SW flux (in  $\text{Wm}^{-2}$ ) comparisons between NISTAR and CERES SYN for April (a) and July (b) 2017.

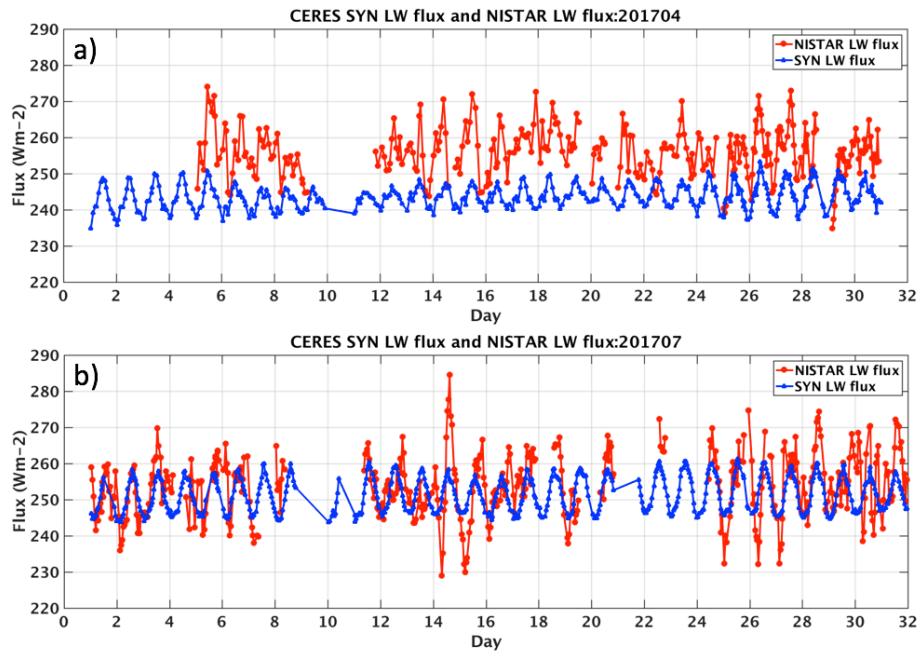


FIG. 8. LW flux (in  $\text{Wm}^{-2}$ ) comparisons between NISTAR and CERES SYN for April (a) and July (b) 2017.

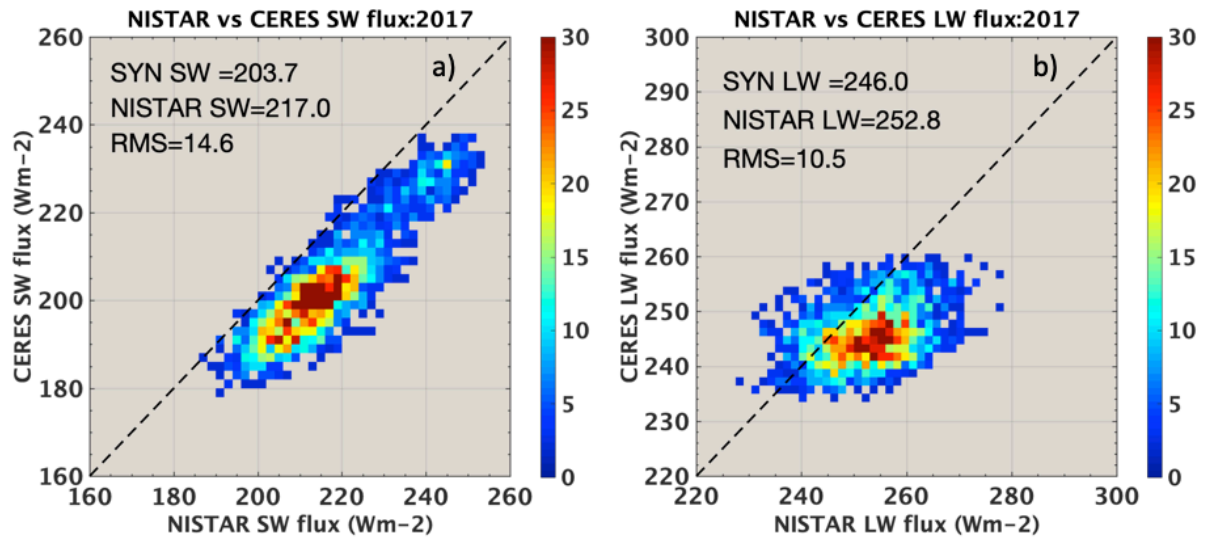


FIG. 9. Comparison of coincident hourly SW and LW fluxes from NISTAR and CERES SYN1deg for 2017. Color bar indicates the number of occurrence.

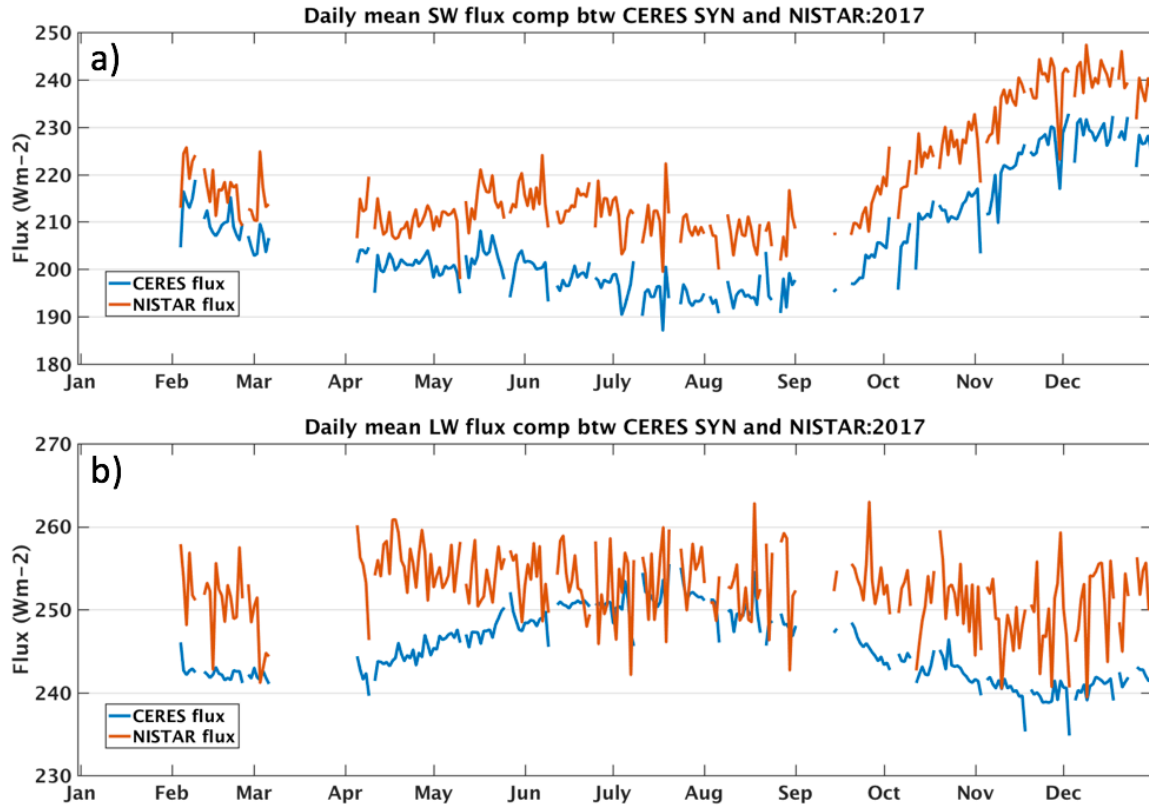


FIG. 10. Daily mean SW flux (a) and LW flux (b) comparisons between CERES SYN1deg (blue) and NISTAR (red) for 2017.

Interactive comment on “Determining the Daytime Earth Radiative Flux from National Institute of Standards and Technology Advanced Radiometer (NISTAR) Measurements” by Wenying Su et al.

Anonymous Referee #1

Received and published: 18 August 2019

11. In ERB calibration your definition of filtered radiance as IRRADIANCE/SOLIDANGLE is only true if the instrument has a completely flat spectral response, which from fig 2 is certainly not the case for NISTAR SW & NIR channels.

*We are not sure what the reviewer means here. NISTAR is a broadband instrument, it measures the energy from the spectral ranges defined in Fig. 2. The relationship between radiance and irradiance should not change with the spectral response function.*

113-132. This is based on CERES unfiltering of Loeb et al 2001 I assume (where they are also labelled Eqns 3 & 4 at <https://journals.ametsoc.org/doi/pdf/10.1175/1520-0450%282001%29040%3C0822%3ADOURFT%3E2.0.CO%3B2> ). Is it completely identical to CERES using the same decades old CERES radiative transfer database? This might be important to briefly mention as it could help eliminate mere inversion biases when you compare to CERES later in the paper.

*The concept of unfiltering used here is the same as that by Loeb et al. (2001), but the database used here is different and was calculated specifically for this study. The current database contains 722 clear-sky cases and 1519 cloudy-sky cases (line 166), whereas the total number of cases used by Loeb et al. (2001) was 272. This information is added to the manuscript.*

142. This is confusing, although I accept it probably amounts to same thing, are look up tables of a & b values (Eqns 3 & 4) or a table of kappa ratios actually used? Only if both techniques are used separately should there be 4 rather than just 2 eqns?

*Thank you for catching this. The Equations 3 and 4 are the original method used to unfilter the CERES observations. As you know, we have scene-type information and Sun-viewing geometry for each CERES footprint, thus the regression can be applied based upon the scene type and Sun-viewing geometry of the CERES footprint. NISTAR views the entire Earth as a single pixel, and the cloud fraction, cloud type, and land/ocean portions differ from time to time. Luckily, the NISTAR SW spectral response function is such that the ratio between filtered and unfiltered radiances exhibit very little sensitivity to the scene types and Sun-viewing geometry. We rewrote the section on page 7 and 8 to correct this.*

143. How are spectrally dependent changes to the transmission of the quartz filter due to outgassing contamination measured after launch and throughout the mission?

*On-orbit measurements indicated that the filters have not degraded significantly since they were measured on the ground during calibration. On orbit measurements of the broadband transmission of the filter stack are continually made every three months using the earth as a source and the photodiode as a detector. The ratios of the on-orbit transmittances amongst each*

*of the two sets of 3 nominally identical filters of each type (SW and NIR) are within 0.2% of each other (as expected from ground measurements) and have remained stable to less than 0.1% throughout the mission. In the case of the SW filter (quartz), the on-orbit broadband transmittance is within 1% of the spectral transmittance of the filter stack over the wavelength range from 500 nm to 2500 nm.*

144. What about quartz filter leakage? Are you using the NIR channel for that somehow (similar to Loeb et al 2001 above)?

*A thin quartz filter can transmit significantly at wavelengths greater than many tens of micrometers, however, the NISTAR filter stacks consists of a pair of 3 mm thick quartz substrates—one is a bare uncoated substrate and the other has dielectric coatings to block light below about 700nm. At 3 mm thickness per substrate the transmittance below 100 micrometers is negligible. Loeb et al (2001) did not use any NIR channel.*

146. Are you sure no unfiltering of the total channel is required, if so how? Was its spectral response measured to be certain? How are you certain no changes to the effective gains of the cavity channels due to electronics radiation exposure are occurring? I'm assuming you do not have onboard blackbodies?

*We do not unfilter the total channel. The total channel spectral response is determined by the spectral absorptance of its cavity absorber, which, like a blackbody, relies on multiple reflections to achieve a high degree of absorptance (emissivity). Each cavity is conical in shape to trap light and is painted with a specular black paint, Z302, which has a very small component of diffuse reflectance. Measurements of the cavity absorptance made on the ground at wavelengths of 488 nm, 514 nm and 632 nm confirmed that the cavity absorbed more than 0.9997 of the incident light. Given the known spectral reflectance of Z302 to long wavelengths and the cavity design (verified at visible wavelengths), un-filtering of the total channel is not required.*

*The only electronics that affect the cavity channel gains are those that measure the electric power applied to the cavity heaters. Those electronics were chosen for their radiation tolerance and long term stability. Given the on-orbit radiation exposure levels, such degradation is not expected to significantly affect heater power measurements. Similar techniques and electronics are used to measure the total solar irradiance from space with a stability of less than 0.1%, which is sufficient to resolve the 11 year solar cycle. Unlike those measurements, degradation from UV exposure is not an issue here. You are correct, there aren't any on-board blackbodies to use as a references. Such blackbodies would have to have phase transition temperature references to be less sensitive to radiation exposure than the radiometers. This is because electronic temperature measurements are much more challenging than measurement of the power applied to the cavity heaters.*

147. How are the SW and Total channels balanced in the solar region as in Kratz et al 2002? (<https://agupubs.onlinelibrary.wiley.com/doi/epdf/10.1029/2001JD001170>)

*Since the NISTAR instrument only views the sunlit side of the Earth, there are no measurements taking during nighttime that can be used in the same manner as Kratz et. al. (2002), in which they looked at the correlation between nighttime total channel and window channel.*

261. what is a shutter cycle? Why is a boxcar filter used in the demodulation algorithm? Are details of these processes important?

*NISTAR utilizes a shutter to modulate light from the Earth just as a chopper wheel is used in the laboratory to modulate a light source. The shutter is opened and closed continuously with a 50% duty cycle with a period of nominally 4 minutes. Each 4 minute period is a shutter cycle. The demodulation algorithm is analogous to what is performed in a digital lock-in amplifier. Use of a boxcar filter having the width of a shutter period strongly rejects higher harmonics of the shutter frequency. Other low pass filters could be used. Note that additional filtering at lower frequencies, e.g., 4 hour running averages, are used to further reduce noise levels. Description is added on page 5.*

264. Recommended based on what, the URL does not work?

*Based on the noise level. The URL was temporally unavailable due to internal web maintenance, it should be available now. Sorry about that.*

266. Why 4 hour “running means” and how is this different from a 4-hour wide boxcar filter from the terminology you used earlier? Does this mean a 4-hour running mean is taken of the boxcar filtered then 2-hour averaged data? Why are the 4 hour means suggested by the NISTAR instrument team?

*A running mean of 4 hours is conceptually the same as a boxcar filter. The 4 hour averages are additional filtering that occurs after the 4 minute wide boxcar filter to reduce noise levels. A four hour compromise is proposed as a trade-off between reducing noise and attenuating the signal of interest, however, the data is also provided without the additional filtering so the user may apply their own filter.*

286. These GERB comparisons need a reference.

*Reference "Doelling et al. (2013)" was provided on line 342, immediately after summarizing the comparison results.*

311. Why does the onboard data processing cause this?

*We removed this sentence in the revised version.*

315. How are the offsets countered, space looks?

*Yes. The shutter removes some, but not all offsets. Those that remain are removed with monthly space looks. Description is added on page 5.*

332. With as few as 10 EPIC results per day are these always equally spaced in time? If not, could this not lead to biases?

*When EPIC is in normal operations, it receives about 10 images daily during the winter cadence. They are normally spaced about 2 hours apart. EPIC receives about 20 images a day during the summer cadence and they are about 1 hour apart. If we simply compare the daily mean fluxes averaged using the EPIC image times with those averaged over the 24 hours, that would lead to biases. In this study, we only averaged the CERES SYN1deg using the hours that coincide with the EPIC times (line 376). Thus ensure both daily means are calculated using same number of hours.*

338. So it seems the LW difference is greatest in Northern Hemisphere Winter, when more ocean is observed? This may be a calibration artifact or error in knowledge of the NISTAR SW channel for the UV region. As per the point above for line 147, how are you balancing SW and Total channels to assure accurate LW in daylight?

*Preliminary analysis of the 2018 measurements does not show the same difference pattern (i.e. larger difference over the boreal summer months than the winter months), thus not supporting the hypothesis of the reviewer. As we mentioned earlier, NISTAR only views the sunlit side of the Earth and the same method used by Kratz et al (2002) cannot be applied here.*

352. “A comprehensive spectral database has been developed”, so is it different from that used by CERES?

*Yes, and more details are added on page 7-8.*

355. So is a constant of 0.8690 used for all NISTAR unfiltering? Unfiltering of LEO scenes varies greatly by several percent especially for ocean scenes etc. So, it seems a value of 0.3% difference for primarily land vs Pacific Ocean scenes would vary more (and maybe adds to your seasonal cycle). What results lead to the 0.3% conclusion and did you try a scene by scene unfiltering?

*Based on the simulated filtered and unfiltered radiances for 722 clear-sky cases and 1519 cloudy-sky cases for each Sun-viewing geometry, the ratio between filtered and unfiltered radiances is extremely stable (see Table 2). Table 2 summarized the ratios and their standard deviation for each solar zenith angle bin for each scene type. For clear-sky case, each solar zenith angle bin contains over 57,000 simulations; and for cloudy-sky case, each solar zenith angle bin contains over 120,000 simulations. The largest ratio difference over different scene types happens under overhead sun, where the ratio for clear ocean is 0.8659 and is 0.8694 for clear land. Using constant unfiltering ratio of 0.8690, it could cause up to 0.3% unfiltering uncertainty if a clear ocean scene is encountered. However, NISTAR views the sunlit side of the Earth as a single pixel. There are always clouds and land mixed in. Thus we state the unfiltering*



*uncertainty should be less than 0.3%. We rewrote the unfiltering portion of the paper on page 7 and 8 to clarify the reviewer's concerns.*

370. Is this the PSF of the EPIC telescope separate from its array of detectors? How was it measured?

*We are not sure we understand the reviewer's question. The PSF tells us where does the light measured in one pixel come from. It's a function of the instrument's entire optical system, telescope and detector. The EPIC PSF was measured in the laboratory before launch, nominal PSF is given in Khlopenkov et al. (2017, SPIE).*

388. Again, this could be due to a constant unfiltering factor?

*Please see response above.*

392. Loeb et al 2018 only quotes the 1% accuracy figure as do you, please provide a peer reviewed SI traceable reference.

*The following CERES calibration references are added:*

*J. M. McCarthy, H. Bitting, T. A. Evert, M. E. Frink, T. R. Hedman, P. Skaguchi, and M. folkman. A summary of the performance and long-term stability of the pre-launch radiometric calibration facility for the Clouds and the Earth's Radiant Energy System (CERES) instruments. In 2011 IEEE International Geoscience and Remote Sensing Symposium, pages 1009–1012, 2011.*

*K. J. Priestley, G. L. Smith, S. Thomas, D. Cooper, R. B. Lee, D. Walikainen, P. Hess, Z. P. Szewczyk, and R. Wilson. Radiometric performance of the CERES Earth radiation budget climate record sensors on the EOS Aqua and Terra spacecraft through April 2007. J. Atmos. Oceanic Technol., 28:3–21, 2011.*

396. Please give a peer reviewed reference for the 2.1% NISTAR SW accuracy figure.

*NISTAR is a relatively new instrument and so far no peer reviewed publication describing the calibration is available. The presentation describing the NISTAR calibration is available at: [https://avdc.gsfc.nasa.gov/pub/DSCOV/Science\\_Team\\_Meeting\\_Sept\\_2019/L1/NISTAR\\_Godda rd%20Science%20Team%2020190917.pdf](https://avdc.gsfc.nasa.gov/pub/DSCOV/Science_Team_Meeting_Sept_2019/L1/NISTAR_Godda rd%20Science%20Team%2020190917.pdf)*

404. With so many error sources not well known it is wrong to simply add them all in quadrature, which assumes they are all random and independent. A more sophisticated error analysis is needed.

*The reviewer is correct that the error sources considered here were simply added to approximate the uncertainty. I would say this is a simplified estimate of the uncertainty, but not the wrong estimate. We know the sources of the uncertainty, but don't know the correlation of all the error sources and therefore unable to estimate the covariances of the sources considered here. The*

*uncertainty given here can be regarded as the upper bound, and this method has been used by Loeb et al. (2009) and Loeb et al. (2018).*

407. The  $1.8 \text{ Wm}^{-2}$  accuracy for CERES LW applies for nighttime LW only. During the day which is always the case for NISTAR it is less accurate. This is because it requires the earlier discussed balancing of the SW and Total channel which if done wrong can result in measuring the Earth warmer at night than during the day for example (see Fig11b, Page 14 at <https://journals.ametsoc.org/doi/pdf/10.1175/2010JTECHA1521.1>). Hence for NISTAR which only views day LW, this is an important consideration.

*The reviewer is correct that the accuracy of the daytime and nighttime LW is different. The daytime LW uncertainty due to calibration is  $2.5 \text{ Wm}^{-2}$  (1 sigma). The combined uncertainty is updated based on the daytime LW flux uncertainty (line 461).*

415. Guesstimate? This is most unsatisfactory for any science paper, let alone one on climate measurements. Please do better.

*Changed to "estimate".*

423. Again, adding in quadrature for so many uncertain, often modelling terms is not acceptable. For example, consider how the error in knowledge of SW vs Total solar response could be systematic because of an error in the ground lab, it will partly cancel in the Total – SW subtraction.

*Please see our response above regarding the uncertainty estimation. The daytime LW flux uncertainty due to calibration is estimated by accounting for the calibration uncertainty in both total channel and SW channel, and the correlations between these two channels.*

428. This is true, in addition to the above-mentioned systematic nature of SW and Total errors not considered in your quadrature additions. A more sophisticated analysis is needed.

*As we stated above, the error analysis considered both SW and total channel. However, changes in error analysis won't affect the correlation between the LW flux from CERES and from NISTAR.*

Overall this paper has merit but needs work to fill in the blanks on some of the processes/references used. The large differences of NISTAR from CERES appears strange and would seem at first look to be largely from algorithm errors. I feel this could be acceptable being a new measurement, but needs to be stated more clearly in the paper as such.

*The NISTAR instrument is the first ever cavity radiometer placed at the L-1 point to measure the Earth's radiation. EPIC on board the DSCOVR also provides 10 narrowband observations from the same Sun-viewing geometry and the visible channels of EPIC are calibration against MODIS. When the global SW anisotropic factors were applied to the EPIC broadband radiance*

*(derived by applying narrowband-to-broadband regressions to EPIC blue, green, and red measurements), the EPIC SW flux agrees with the CERES SYN SW flux to within 2%. The good agreement indicates that the algorithm that we developed is accurate and is not the cause for the large discrepancy between NISTAR and CERES SYN. Even though there are discrepancies between the NISTAR fluxes and CERES SYN fluxes, we feel it is important to document the measurement, the algorithm, and the validation for future reference.*

The use of constant SW unfiltering also raises concern and leads to the possibility it is a cause of the larger than expected seasonal cycles, but more investigation is needed. Also some insight in the introduction into the purpose of NISTAR would be good, such as giving illustration if and how it complements the climate observing system discussed by Weilicki et al 2013 (<https://journals.ametsoc.org/doi/pdf/10.1175/BAMS-D-12-00149.1> ).

In summary, this paper could become suitable for publication, given more work, research and additions that address the points above. It should then be re-considered under peer review.

*Based on the simulated filtered and unfiltered radiances for 722 clear-sky cases and 1519 cloudy-sky cases for each Sun-viewing geometry, the ratio between filtered and unfiltered radiances is extremely stable (see Table 2). Table 2 summarized the ratios and their standard deviation for each solar zenith angle bin for each scene type. For clear-sky case, each solar zenith angle bin contains over 57,000 simulations; and for cloudy-sky case, each solar zenith angle bin contains over 120,000 simulations. The largest ratio difference over different scene types happens under overhead sun, where the ratio for clear ocean is 0.8659 and is 0.8694 for clear land. Using constant unfiltering ratio of 0.8690, it could cause up to 0.3% unfiltering uncertainty if a clear ocean scene is encountered. However, NISTAR views the sunlit side of the Earth as a single pixel. There are always clouds and land mixed in. Thus we state the unfiltering uncertainty should be less than 0.3%. We rewrote the unfiltering portion of the paper on page 7 and 8 to clarify the reviewer's concerns.*

Interactive comment on “Determining the Daytime Earth Radiative Flux from National Institute of Standards and Technology Advanced Radiometer (NISTAR) Measurements” by Wenying Su et al.

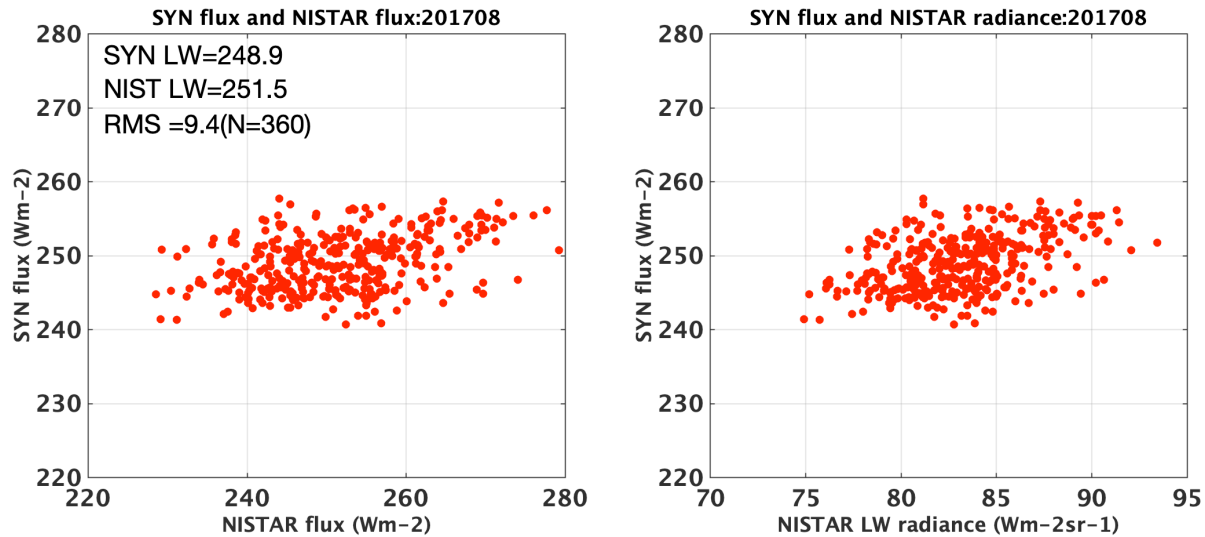
Anonymous Referee #3

Received and published: 19 August 2019

General comments:

This paper presents the scheme and algorithm of deriving TOA SW/LW flux from NISTAR measurements and comparison also made with the corresponding results derived from CERES. I am impressed by the detailed and clear description of the algorithms. The paper is very well written and relevant to the community. I recommend publication after addressing the minor issues listed below. It doesn't seem that the uncertainties in the algorithms would give a consistent bias seeing in the differences between NISTAR and CERES. Has there been analysis with the NISTAR instrument measurements and calibration? The low correlation between NISTAR LW flux and that of CERES is puzzling. To bypass the potential uncertainties in part of the algorithms, it may be useful to look at the correlation between the NISTAR LW radiances and the CERES flux to see if they are correlated at all.

*The NISTAR instrument team (who produces the LI data) is responsible for the instrument calibration and the team has presented their calibration at the DSCOVR science team meetings([https://avdc.gsfc.nasa.gov/pub/DSCOVR/Science\\_Team\\_Meeting\\_Sept\\_2019/LI/NISTAR\\_Goddard%20Science%20Team%2020190917.pdf](https://avdc.gsfc.nasa.gov/pub/DSCOVR/Science_Team_Meeting_Sept_2019/LI/NISTAR_Goddard%20Science%20Team%2020190917.pdf)). So far their analysis are mainly focused on the SW channel. NISTAR has three broadband electrical substitution radiometers (ESRs). All ESRs have a large background noise as they measure the change in incident optical power. Two steps are utilized to remove the background noise: first using a shutter to modulate the source which removes most of the background noise then using dark space view to remove the residual shutter-modulated background. The shutter modulated background is largest for the total channel and is smaller for the SW channel. As the LW is derived from the difference between total and SW channels, both total channel and SW channel background noises contribute to the LW uncertainty. The NISTAR total channel uncertainty is 1.5% and the SW channel uncertainty is 2.1%. Assuming the SW flux is  $210 \text{ Wm}^{-2}$  and the LW flux is  $240 \text{ Wm}^{-2}$ , thus gives the total flux uncertainty as  $450 \times 1.5\% = 6.8 \text{ Wm}^{-2}$ , and the SW flux uncertainty as  $210 \times 2.1\% = 4.4 \text{ Wm}^{-2}$ . The resulted uncertainty in LW flux is  $8.1 \text{ Wm}^{-2}$ , which can explain most of the LW differences between NISTAR and CERES SYN shown in Table 4. See added description on page 6. The low correlation is also caused by the background noise in both the total and SW channels. Details on NISTAR calibration are added on pages 5 and 6. Below is an example of August 2017, where the correlation between CERES SYN LW and NISTAR LW flux is about 0.38, and the correlation between CERES SYN LW flux and the NISTAR LW radiance is about 0.41. It is obvious that the low correlation is mainly from the instrument calibration.*



Specific comments:

Page 5 and 6: The authors have derived the regression equations for the unfiltered radiances (Eq 3 and 4); what is the reason for using the less accurate ratio method (Eq. 5 and 6)?

*The Equations 3 and 4 are the original method we planned to use for the NISTAR unfiltering. But unlike other LEO instruments that have scene-type information and Sun-viewing geometry for each footprint, and the regression can be applied based upon the scene type and Sun-viewing geometry of each footprint. NISTAR views the entire Earth as a single pixel, and the cloud fraction, cloud type, and land/ocean portions differ from time to time. Luckily, the NISTAR SW spectral response function is such that the ratio between filtered and unfiltered radiances exhibit very little sensitivity to the scene types and Sun-viewing geometry. We rewrote the sections on page 7 and 8 to correct this.*

Page 14: How are the portion of the Earth not visible to NISTAR decided? Also, similar to NISTAR missing some of the daytime portion of the Earth, it must be seeing part of the night time side of the Earth. Are these taking into account for the longwave calculations?

*The mask is calculated based upon the solar zenith angle and the EPIC viewing zenith angle and each EPIC pixel is identified as nighttime hidden to EPIC, or nighttime visible to EPIC, or daytime hidden to EPIC, or daytime visible to EPIC. Both the daytime and nighttime visible to EPIC are considered for the CERES SYN product to compare with the NISTAR LW measurements. Some clarification is added on page 14 and Figure 6b) is modified accordingly.*

Review on “*Determining the Daytime Earth Radiative Flux from National Institute of Standards and Technology Advanced Radiometer (NISTAR) Measurements*” by Su et al.

This paper documents the methodology to derive the broadband radiative flux from the measurements of the NISTAR instrument onboard of the DSCOVR mission. Some preliminary results based on this method are compared with the well-developed CERES data. The SW fluxes derived from the NISTAR compares reasonably well with CERES, but the LW fluxes from NISTAR have a systematic bias and low correlation coefficient when benchmarked with CERES.

The topic of this paper is important and suitable for AMT. The paper is well organized. However, the paper lacks some important technical details about the instrument and the methodology, as well as the author’s opinion about the usefulness of the NISTAR product. In my view, some significant revisions are needed before the paper can be accepted for publication. Below is a list of questions and concerns I have.

1) The parameterization scheme described in Section 2 to obtain unfiltered radiance from observed filtered radiance is confusing. Up to line 132, the method seems to be based on the polynomial parameterization scheme in Eqs (3) and (4). But then it suddenly changed to the simply ratio-based parameterization in Eqs. (5) and (6). Why are there two types of parameterization? Which one is used?

*The Equations 3 and 4 are the original method we planned to use for the NISTAR unfiltering. But unlike other LEO instruments that have scene-type information and Sun-viewing geometry for each footprint, and the regression can be applied based upon the scene type and Sun-viewing geometry of each footprint. NISTAR views the entire Earth as a single pixel, and the cloud fraction, cloud type, and land/ocean portions differ from time to time. Luckily, the NISTAR SW spectral response function is such that the ratio between filtered and unfiltered radiances exhibit very little sensitivity to the scene types and Sun-viewing geometry. We rewrote the sections on page 7 and 8 to correct this.*

2) What is the FOV size of the NISTAR instrument? Does it observe the earth pixel by pixel (similar to EPIC) or as a whole? Does its FOV include some cosmic background and, if so, how is that treated?

*NISTAR observes the entire sunlit side of the Earth as one pixel. We specifically mentioned this on lines 53-54.*

3) Within its FOV, does the NISTAR instrument response to the radiance from different locations and angles equally? In other words, do the radiances from the edge of the earth disc have the same weighting as those from the center of the disc?

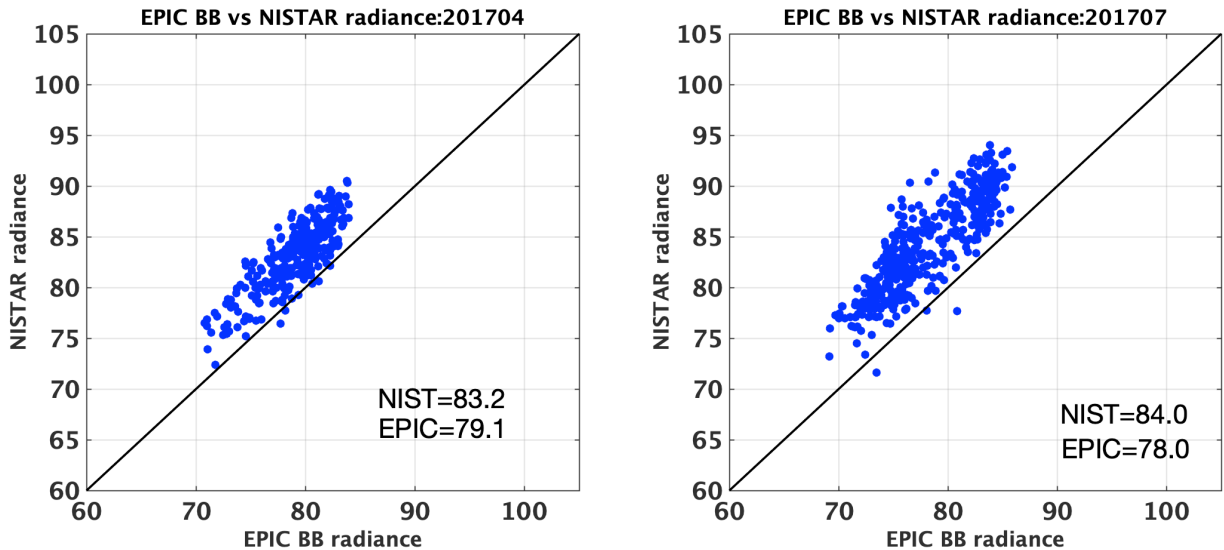
*Yes, NISTAR response to the radiance from different locations and angles equally. Optically the instrument is very simple—there aren’t any lenses or mirrors, just filters, and a pair of apertures, and incident light is nearly perpendicular to the filters and apertures. The Earth subtends an angle of less than 1 degree from DSCOVR.*

4) It is stated that “The biases in the anisotropy correction for the DSCOVR scattering angle are mitigated and potentially minimized by the wide range of different scene 71 types viewed in a given NISTAR measurement.” Some references are needed to support it.

We referenced the Su et al. (2018) paper here. As this is a very new way to measure the Earth radiative flux, no other references are available.

5) In Su et al. (2018), a similar method is used to derive the fluxes from EPIC measurements. One of the byproduct from this EPIC-based method is the “global day-time mean SW radiance”  $I_{bb}$ . Is it something directly comparable to the observation of NISTAR instrument? If so, some comparisons should be made because both EPIC and NISTAR have the similar sun-satellite geometry.

Indeed, we have derived the "global daytime mean SW radiances from EPIC". They are consistently lower than the radiances from NISTAR. Below are the comparison between NISTAR and EPIC radiances for April and July 2017. The mean differences are between 4 to 6  $Wm^{-2}sr^{-1}$ . We chose not to include these results in this paper to avoid any confusions and the EPIC and CERES comparisons were provided in Su et al. (2018).



6) I have several questions about the method described in Section 3c. First of all, what is the theoretical based for Eqs 9~ 11? If my understanding is correct, the global mean SW flux is

$$F = \iint_{sunlit} \frac{I[\theta_0(r), \theta^e(r), \phi^e(r), \chi(r)]}{R(\theta_0, \theta^e, \phi^e, \chi) d^2r} d^2r$$

Where r denotes a point on earth. But this is not equal to

$$\frac{\iint_{\text{sunlit}} I[\theta_0(r), \theta^e(r), \phi^e(r), \chi(r)] d^2r}{\iint_{\text{sunlit}} R(\theta_0, \theta^e, \phi^e, \chi) d^2r}.$$

More detailed mathematical derivations are needed here. Secondly, one might ask if a global mean anisotropic factor is even physically meaningful? The average is over a large range of viewing angles and scene types. Does the result have any physical meaning? Moreover, are the angular and spectral averaging independent and can be treated independently? The derivations in Section 3c seem to suggest they are independent, but this is not obvious to me. Some clarification is needed.

*We agree with the reviewer that the above two equations are not the same. The first equation is how we calculate global mean flux from low-Earth orbit satellites (i.e. CERES) using the footprint level data (resolution on the order of 20 km) by first grid the data then area weight to calculate the global mean. We did not use the second equation in our study to derive the fluxes from NISTAR radiance measurements.*

*To derive the global mean flux from NISTAR measurements, a corresponding anisotropic factor to characterize the sunlit portion of the Earth as a whole is needed, and this is the definition of the global mean anisotropic factor we used in the paper. The global mean anisotropic factor is derived by using the radiances and fluxes defined in the CERES angular distribution models (ADMs). The global mean radiance and flux from CERES ADMs were calculated independently (see Equations 8 and 9 in the revised version). They are used to derive the global anisotropic factor (Equation 10) and subsequently to convert the NISTAR radiance to flux (Equation 11). The deviation of the NISTAR flux used here is not the same as illustrated by the second equation above. This method has been tested for both the NISTAR and EPIC measurements.*

7) This paper only shows “how to do it” but does not explain “why to do it” other than it can be done. I understand that this paper is to document the method used to derive the flux from the radiance observations of NISTAR. But I think in addition to the technical details the reader would appreciate some insights and opinions from the authors about the usefulness of the product. We already have the state-of-the-art CERES flux product and in Su et al. (2018) flux product has also been developed. What is new/novel/important about the NISTAR flux product other than the fact it can be done? What kind of applications can this product be used for? Some discussions about these important questions should be added to the abstract and conclusion parts.

*We added some information on NISTAR measurement and its utility in the introduction (lines 59-68). We also added some perspective on the utility of NISTAR SW fluxes in the conclusion section (lines 486-492).*



Review of “Determining the Daytime Earth Radiative Flux from National Institute of Standards and Technology Advanced Radiometer (NISTAR) Measurements” by Su et al. 2018

General comments:

This manuscript derives sunlit side of the Earth’s radiation budget (SW and LW) from a single pixel measurement of NISTAR instrument on board the DSCOVR mission and compares with the radiation fluxes derived from the CERES measurements. This is a very interesting and important work as the Earth’s radiation budget has been so far solely measured by the ERBE/CERES project and there are very little independent and direct measurements of these important quantities. This work builds upon many previous works the team has been working on for many years including narrowband broadband conversion, ADM, GEO/LEO composite cloud products etc. The paper is well written and structured. I do have some questions and suggestions regarding the derivation of global ADM and evaluation of each components of the fluxes.

Specific comments:

Line 98: What is the uncertainty level of NISTAR L1B radiance? What kind of calibration procedures have been used to produce the L1B radiance? You have discussed some of the issues later in the paper but it’s worthwhile to have a paragraph to discuss the NISTAR at the beginning of the paper. NISTAR provides a completely different methodology of estimating the earth’s radiation budget and independent check of Earth’s radiation budget created from CERES measurements, the difference found in this article is very serious and should be adequately explained. NISTAR’s absolute calibration and uncertainty is of fundamental importance, otherwise the readers would question the well-established CERES products.

*The NISTAR instrument team (who produce the L1 data) is responsible for the instrument calibration and the team has presented their calibration at the DSCOVR science team meetings ([https://avdc.gsfc.nasa.gov/pub/DSCOVR/Science\\_Team\\_Meeting\\_Sept\\_2019/L1/NISTAR\\_Goddard%20Science%20Team%2020190917.pdf](https://avdc.gsfc.nasa.gov/pub/DSCOVR/Science_Team_Meeting_Sept_2019/L1/NISTAR_Goddard%20Science%20Team%2020190917.pdf)). So far their analysis are mainly focused on the SW channel. NISTAR has three broadband electrical substitution radiometers (ESRs). All ESRs have a large background noise as they measure the change in incident optical power. Two steps are utilized to remove the background noise: first using a shutter to modulate the source which removes most of the background noise then using dark space view to remove the residual shutter-modulated background. The shutter modulated background is largest for the total channel and is smaller for the SW channel. As the LW is derived from the difference between total and SW channels, both total channel and SW channel background noises contribute to the LW uncertainty. The NISTAR total channel uncertainty is 1.5% and the SW channel uncertainty is 2.1%. More details on NISTAR calibration is added on page 4-6.*

Line 147. The conversion from filtered to unfiltered radiances used the ratio derived from model simulation data using eq 5 and 6. Why not using the regression (3) and (4)? The regression indicates the ratio could not be constant because it’s a quadratic function and has an offset. It’s justified to use a constant ratio between the two if the ratio varies little as for the SW band, but a constant ratio for NIR would introduce an unnecessary source of error (1\_2%) for the NIR and I don’t see why you should

abandon the regression.

*The Equations 3 and 4 are the original method we planned to use for the NISTAR unfiltering. But unlike other LEO instruments that have scene-type information and Sun-viewing geometry for each footprint, and the regression can be applied based upon the scene type and Sun-viewing geometry of each footprint. NISTAR views the entire Earth as a single pixel, and the cloud fraction, cloud type, and land/ocean portions differ from time to time. Luckily, the NISTAR SW spectral response function is such that the ratio between filtered and unfiltered radiances exhibit very little sensitivity to the scene types and Sun-viewing geometry. As we don't have the scene-type information, the regression method can't be applied to NIR either. We rewrote the sections on pages 7 and 8 to correct this.*

Line 152: Did you use NIR in this work? If not, could you explain why NISTAR takes the NIR measurement?

*We did not use the NIR channel in our work. When NISTAR was design in the 1990s, the primary utility for the NIR channel is to study the enhanced absorption of SW radiation by clouds (Collins 1998) and more recently by Carlson et al. (2019) to look at the spectral ratio of the sunlit side of the Earth and the potential of using the ratio for model evaluation (see introduction on page 3).*

Line 187: EPIC images have 8x8 km<sup>2</sup> resolution at nadir and are  $1/\cos(\text{vza})$  larger at larger view zenith angles. The EPIC cloud products are retrieved at its native resolution with (2014x2014) pixels in a granule. Some channels have degraded into 1024x1024 for downlink but reversed to 2014x2014 afterwards.

Equation (9) and (11),  $I_j$  and  $F_j$  seem to refer to radiance and flux in each EPIC composite pixel. Do you actually use those in the mean ADM calculations? If yes, did you use the EPIC measured narrowband radiances to compute the broadband radiance and flux for each pixel?

*In Equations (9) to (11) (now Equations 7-9), the  $I$  and  $F$  refer to the radiances and fluxes from the CERES ADMs (same symbols are used in Equation 6). To clarify the confusion, we added a sentence on page 12 (lines 291-292). They are not from EPIC.*

Why did you grid the fluxes into 1x1 grid boxes and not the radiances? The global mean flux is computed from Eq. 11 to take care of different sizes of grids in each latitude. If you grid the radiance, then you would compute the mean radiance the same fashion as the flux. Otherwise, if you average the radiance from each pixel directly, then you would also have to consider the pixel size differences and the radiance average has to be a pixel-size weighted average.

*Radiance and flux are fundamentally different physical quantities. Radiance is the total amount of energy confined to a given direction per unit surface area (in  $\text{Wm}^{-2}\text{sr}^{-1}$ ). One essential property of radiance is that it is additive, meaning if several sources contribute to the radiance at a particular point and in a particular direction, the total radiance is the sum of the radiances from each source as if it were acting alone (Bohren and Clothiaux, 2006). On the other hand,*

*flux is the energy per unit surface area ( $Wm^{-2}$ ) and need to be area-weighted when compute global means.*

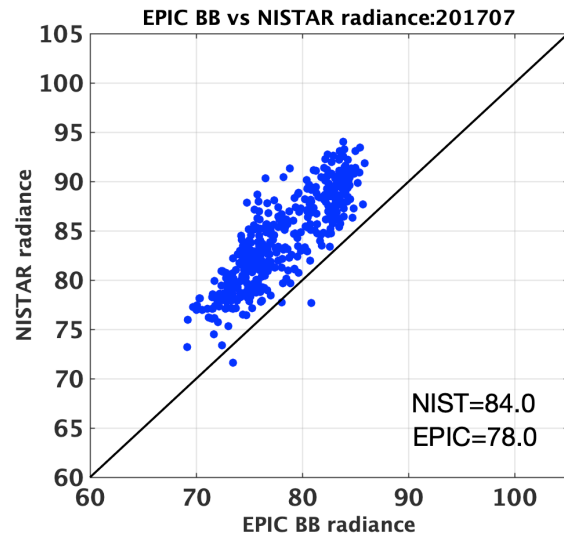
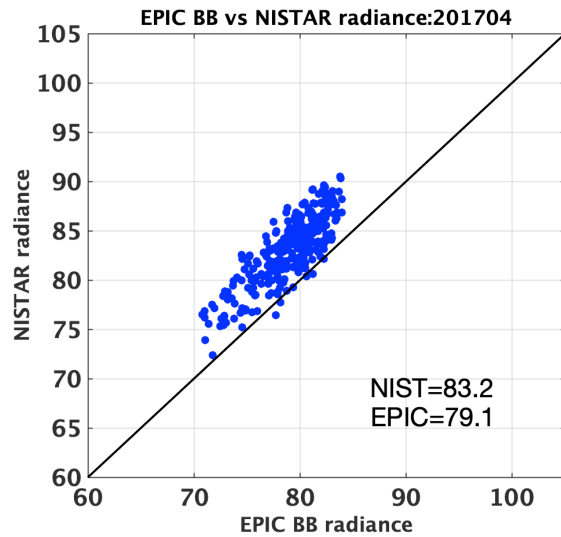
If my understanding is correct, then the global ADM not only rely on composite product's scene identification, CERES ADM for each pixel, but also on EPIC's radiances measurements (which rely on CERES-MODIS collocation and narrowband to broadband conversion) to derive the global mean ADM. The EPIC-based sunlit global SW flux (Su et al. 2018) has used EPIC radiances and CERES ADMs and does not really need global ADM and thus global ADM is essentially untested. From EPIC radiance to flux, it relies on CERES derived narrowband-broadband conversion and CERES ADM, therefore the EPIC global flux provides some consistent check but not absolute validation in my opinion.

*The global ADM is derived using the scene identifications (surface type, cloud fraction, cloud optical depth, etc. ) from EPIC composite to select the anisotropic factors from the CERES ADMs. EPIC radiances are not used.*

*In Su et al. (2018), we used the same methodology to derive the global SW anisotropy factors (same as Equations 7-10 in the revised version). They are then applied to the EPIC global daytime mean "broadband" SW radiances, which were derived by using narrowband-to-broadband regressions. The EPIC global daytime mean "broadband" SW radiances are analogous to the NISTAR SW measurements, and the same SW anisotropy factors were applied to both NISTAR and EPIC SW radiances to derive SW fluxes. As demonstrated in Su et al. (2018), the SW flux from EPIC agree with CERES SYN to within 2%, which means that the method that we developed to derived the global mean anisotropic factors are robust.*

Eq. 13 and 14. From these equations, we know that the NISTAR flux depends on unfiltered radiances from NISTAR and the global ADM derived from EPIC (which itself depend on many other instruments and procedures). I would strongly suggest the authors examine the global ADM and NISTAR's radiance measurements separately to understand the variability and trends from each of these components. The computation of global ADM can be refined as mean radiance could be computed with pixel-size weighted average. The NISTAR total radiance and NIR radiance are also worth looking at especially when LW is derived from total subtract the SW.

*Again, EPIC composite provides the cloud properties that needed to select the anisotropy factors. We don't use any EPIC measurements in this study. The global mean anisotropy factors are calculated by deriving the anisotropic factors for each EPIC pixel. We did examine the NISTAR radiance against the "global daytime mean SW radiances from EPIC", derived by using narrowband-to-broadband regressions (Su et al. 2018). The NISTAR radiances are consistently greater than the EPIC "broadband SW" radiances. Below are the comparison between NISTAR and EPIC radiances for April and July 2017. The mean differences are between 4 to 6  $Wm^{-2}sr^{-1}$ . We chose not to include these results in this paper to avoid any confusions and the EPIC and CERES comparisons were provided in Su et al. (2018).*



*As noted by the reviewer, the LW radiance is derived by subtracting SW from the total. Thus the LW contains information of the total channel. As the focus of this paper is to derive SW and LW fluxes from NISTAR and validate the product with CERES product, and there aren't any global daytime total and NIR measurements that can be used to compare with the NISTAR measurements, simply looking at the NISTAR total and NIR channel measurement won't add value to the paper.*