Interactive comment on "Quantifying Hail Size Distributions from the Sky: Application of Drone Aerial Photogrammetry" by J. S. Soderholm et al.

Review 1

Anonymous Referee #2

Received and published: 19 September 2019

The authors present a new technique for measuring hail size distribution using aerial photography and a machine learning approach to detecting and measuring individual hailstones. This is a novel method using new technologies and would provide better sampling with more automation. There are, however, a number of requirements for this technique and I would like to see more discussion on these limitations and in what kind of scenarios this method would be practical.

The authors wish to thank the reviewer for taking the time to read this paper and the valuable comments regarding improvements to the discussion of requirements and limitations.

Further concerns:

Pg 2 Ln 22-25: Can you quantify these requirements better? It looks like, in this case, the hailstones must be larger than 20mm. What kind of spacing must there be between them on the ground? How much is the technique affected by the presence of smaller hailstones as well?

This is an important point, thank for for highlighting it. The hail size requirements are a dependent on the effective ground resolution of the orthomoasic, and thus are best discussed after the sensor is introduced (section 2). Two additional sentences have been added to section 4 (p. 6 lines 22-25) that provide recommended minimum hail size and ground coverage for successful HailPixel surveys. In summary this is a minimum value of 20 mm for the major axis length and a maximum of 50 % ground coverage. Further testing would be required to determine the effect of smaller hailstones on the measurement accuracy.

Pg 3 Ln 6: Is there a restriction on the 10m wind speed in order to fly the drone at such a low speed? Does accuracy fall off with a large or variable wind?

Thank you for raising this point. The wind speed at the time of capture was a gentle breeze (3.5 – 5.5 m/s) and this has been added to page 3, lines 15. The authors recommend that hailpixel surveys are conducted with near surface below 8 m/s to reduce the likelihood of motion blurring. This recommendation has been added to the text in page 6, lines 26-27. The authors believe the accuracy of the hail size retrieval would be effected without a doubt in stronger winds.

Pg 4 Ln 8: How were the 12 training tiles chosen? Does the number of tiles necessary for training depend on the concentration of hail, or on the type of background?

Twelve tiles were chosen to ensure the total number of stones was approximately 1000. A smaller sample size may have also been appropriate but the sensitivity of the model performance to training size was not explored. Further, pre-training weights from the COCO dataset were used to initialise the model, limiting the need for very large training datasets. The specific tiles used for model training were manually selected to represent the different background types, as this is also an important part of training an RCNN model. The manuscript has been amended to clarify that the 12 tiles used used to achieve a sufficient sample size and were selected manually to sample the different background types (page 4, lines 14-16).

Pg 4 Ln 26: Do the tiles have similar distributions or concentrations?

The hail concentration of individual tiles varied between 15 and 91 stones per m², with a mean of 47 stones per m². Inspection of the spatial distribution indicated that vegetation density and slope plays an important role in the concentration of hail when it finally comes to rest. Higher concentrations appear in regions of denser grass, likely due to the grassy areas acting to dampen bouncing and rolling of stones, increasing collection. The lowest concentrations appear on unvegetated areas where hail can readily bounce and roll. An analysis of individual tile concentration has not been included in the text because it doesn't reflect the true distribution of hailstones for these reasons. Additional information on the concentration has been added to p5 L27-30.

Pg 5 Ln 5: What kind of range does the 'lightness value' have? If the lightness value of a hailstone must be >50 more than that of the edge, how much does this restrict the type of background against which hail can be measured? How sensitive is the lightness value or its variability to the overhead light (sky conditions, sun angle)?

The lightness value is a 8 bit index that has a range of 256 values (added to the text in the foot note on page 4). The authors agree that the minimum difference in lightness between the hailstone centre and hailstone edge is definite limitation for the technique. However, even for light-coloured soils, the lightness difference still remained well above 50 units. A comments has been added to the text (page 5, lines 13-14) regarding the performance of this lightness difference threshold for different background types. Regarding the sensitivity of lightness values to the sky conditions, the November 2018 survey was conducted during cloudy conditions. The automatic exposure and white balance control on the UAV camera was able to compensate for the low lighting conditions. Comparable lightness values for hail and background types would be expected for full-sun conditions with the correct exposure and white balance adjustments.

Pg 5 Ln 10: How many hailstones were counted on the pad? How does the total concentration compare? Were there no hailstones <20 mm measured on the hail pad or were they just not considered for this comparison?

Thank you for raising these important points. A total of 17 impacts with major and minor axis measurements were sampled from the hail pad, with a concentration of 141 stones per m². This information has been added to the text on page 5 lines 20 and lines 27-28 respectively. The mean concentration observed by the HailPixel survey was 47 stones per m², significantly less than the hail pad. This is possibly due to the fact the two samples were not co-located, and the hail pad experienced a longer duration of hail fall. Further, bouncing hail stones may have introduced

secondary impacts on the hail pad that were indistinguishable. This justification has been added on page 5 lines 23-25. No stones less than 24 mm major axis length were sampled by the hail pad.

Review 2

Referee: Tomeu Rigo

Received and published: 11 December 2019

Dear authors, your manuscript presents a novelty in hail size research, taking profit of capabilities of drones and photography treatment techniques. Having in mind the review criteria of the journal, I'm considering that it has an excellent scientific significance and good scientific and presentation qualities. Besides, the paper answers positively to all the questions made in the same web. However, I think that there are some points that you must solve before the acceptance of the paper

Hello Dr Tomeu Rigo, Thank you for the kind words and constructive review! Your feedback is much appreciated.

- There are some typos: L1: "HailPixel." (the dot must be placed after the ") L6 of the page 3: ((Fig. 1a) - remove one (

After some investigation we have confirmed that the comma is to be placed inside the quotes on L1, page 1 (https://www.grammarly.com/blog/quotation-marks/). The additional bracket from L6, page 3 has been removed.

- I think that your technique can be useful for more aspects that the cited in the text:for instance, for identifying the whole area affected by hail. I understand that your technique can discriminate between hail/non-hail pixels and then, you can delimitate the hail path. In the same way, do you think that this is applicable in real-time? If your answer is positive, explain it in the text, because this could help in many fields, in those areas commonly affected by hail events, such identification of damaged agriculture production or for insurance interests, among others.

This is a very interesting idea to sample the hail swath extent/coverage rather than the hail size distribution. To sample at the resolution required for retrieving hail size, we had to fly the drone very low to the ground (10m), limiting the sample area to a few hundred meters. To collect aerial imagery that covers a significant portion of a hail swath (or say a farm or suburb), you would probably need to use a fixed wing drone, which provides greater endurance, flight speed and altitude performance than a quadcopter. Real-time extraction of information from the imagery would be more challenging, but not impossible. The authors thank the reviewer for this suggestion but believe this is outside the scope of this paper and expect it will be attempted at some point by hail researchers.

- When you introduce hail-pads, you forget to mention automatic hail-pads (see, e.g.Martin Löffler-Mang, Dominik Schön, Markus Landry, Characteristics of a new automatic hail recorder, Atmospheric Research, V. 100, Issue 4, 2011, Pp. 439-446, ISSN0169-8095, https://doi.org/10.1016/j.atmosres.2010.10.026.)

The reference to Löffler-Mang 2011 description of an impact disdrometer design has been added to L3, p2

- In data and approach, please provide numbers (L25 pg 2): which size and density can be considered as thresholds?

A lower threhold for the hail diameter is now suggested in L33 p2. Providing an estimate for a maximum concentration threshold is not feasible as it depends on the hail size distribution. The author's suggestion of non-accumulating hail should provide sufficient guidance for the concentration limits.

- Where is the hail-pad used for the comparison located? You should indicate in a figure.

The authors believe showing the location of the hail pad and hail survey location on a map would have little value to the article, as the focus is not on the physical setting or spatial variability, but the technique. The exact location of the hail pad has been added in L5 p3, and the location of the aerial survey in L14 p3 for reference.

- It results difficult to me understand which is the final size of the pixel, the one you use in fig 2c

The final pixel size of the orthomoasic has been added to the paper in L24 p3 to clarify the pixel size used in the analysis.

- Those parts of the manuscript that are not referring to your work should be moved to the introduction, where the state-of-art is presented: e.g. L 1-5 of page 4, or some previous results used in your discussion.

Thank you for this suggestion. L1-5 of p4 have been moved to the introduction in p2 L16-20. For the discussion, the citations to other work are an essential part of this section and the author feels that duplicating or moving these citations in the introduction wouldn't add much value for the reader.

- León is not placed in France (L18 Page 5)

Thank you, this has been corrected to Spain:)

- I think that you could do an effort and give more applicabilities to your research, such the cited previously in my repport, including some references about this point Botzen, W. J. W., Bouwer, L. M., & Van den Bergh, J. C. J. M. (2010). Climate change and hailstorm damage: Empirical evidence and implications for agriculture and insurance. Resource and Energy Economics, 32(3), 341-362.

Changnon, S. A., Changnon, D., Fosse, E. R., Hoganson, D. C., Roth Sr, R. J., & Totsch, J. M. (1997). Effects of recent weather extremes on the insurance industry: major implications for the atmospheric sciences. Bulletin of the American Meteorological Society, 78(3), 425-436.

Sánchez, J. L., Fraile, R., De La Madrid, J. L., De La Fuente, M. T., Rodríguez, P., & Castro, A. (1996). Crop damage: The hail size factor. Journal of Applied Meteorology, 35(9),1535-1541.

Hohl, R., Schiesser, H. H., & Aller, D. (2002). Hailfall: the relationship between radar-derived hail kinetic energy and hail damage to buildings. Atmospheric Research, 63(3-4), 177-207.)

Thank you for this suggestion. An additional sentence has been added to the introduction to highlight the application of improving hail size distribution knowledge. Citations to Changnon et al. 1997, Sanchez et al. 1996 and Hohl et al. 2002 has been added to support this. P1 L15-17.

- In my opinion, you need to separate more clearly the part of your work from other previous techniques, and, besides, to present, maybe in a table, the technical characteristics of the analysed imagery

Thank you for making this point. Two additional sentences have been added to the summary to highlight that this technique provides a significantly larger hailstone sample size that leads to more robust statistics (p7 L7-9). These sentences include the sample area and sample size of the analysed imagery.

Quantifying Hail Size Distributions from the Sky: Application of Drone Aerial Photogrammetry

Joshua S. Soderholm¹, Matthew R. Kumjian², Nicholas McCarthy³, Paula Maldonado⁴, and Minzheng Wang⁵

Correspondence: Joshua Soderholm (joshua.soderholm@bom.gov.au)

Abstract. A new technique, named "HailPixel," is introduced for measuring the maximum dimension and intermediate dimension of hailstones from aerial imagery. The photogrammetry procedure applies a convolutional neural network for robust detection of hailstones against complex backgrounds and an edge detection method for measuring the shape of identified hailstones. This semi-automated technique is capable of measuring many thousands of hailstones within a single survey, which is several orders of magnitude larger (e.g., 10,000 or more hailstones) than population sizes from existing sensors (e.g., a hail pad). Comparison with a co-located hail pad for an Argentinan hailstorm event during the RELAMPAGO project demonstrates the larger population size of the HailPixel survey significantly improves the shape and tails of the observed hail size distribution. When hailfall is sparse, such as during large and giant hail events, the large survey area of this technique is especially advantageous for resolving the hail size distribution.

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1 Introduction

Measurements of the hail size distribution (HSD) are challenging to collect owing to the infrequent and hostile nature of hailstorms. Because of these constraints, HSD measurements are uncommon, especially for larger hail (> 25 mm). Such observations are necessary to constrain hail microphysics parameterization schemes used in weather and climate models, and for hail detection and sizing algorithms from weather radar. Improvements to hail retrievals and modelling are an important step towards mitigating the increasingly significant hail-related losses to agriculture, motor vehicles and buildings (Sánchez et al., 1996; Changnon et al.

Ground sensors for measuring the size distribution of large hail can be separated into those that provide time-recording (e.g., hail disdrometer) and those that provide time-integrated measurements (e.g., hail pad). Time-recording instruments such as im-

¹University of Bonn, Meteorological Institute, Germany

²The Pennsylvania State University, Department of Meteorology and Atmospheric Science, United States

³The University of Queensland, Australia

⁴University of Buenos Aires, Argentina

⁵Northraine PTY, LTD.

pact or optical disdrometers provide valuable information on the temporal variability of the HSD within a given storm, but are often expensive to fabricate and maintain, and difficult to deploy. Thus, such instruments typically are only deployed as smaller networks or for field campaigns (e.g., Federer and Waldvogel, 1975; Brown et al., 2014)(e.g., Federer and Waldvogel, 1975; Löffler-Mang. In contrast, time-integrated instruments often are cheaper to fabricate, maintain, and deploy, making them attractive options for longer-term monitoring of hailfall. The most commonly used time-integrated instrument for measuring HSDs is a hail pad, consisting of a foil covered styrofoam pad that preserves dents of hail impact (Long et al., 1979). This sensor is cost effective and has seen extensive use by previous and ongoing campaigns in the US and Europe over the last 50 years (Cheng and English, 1983; Fraile et al., 1992; Cifelli et al., 2005; Kalina et al., 2014). Both hail pads and hail disdrometers provide reasonable estimates of hail size, but are subject to significant limitations even with careful calibration (e.g., Palencia et al., 2011). Further, both time-recording and time-integrated instruments for measuring the HSD utilize a small sample area on the order of 0.1 to 0.3 m². Towery et al. (1976) suggests this small sample area is likely to underrepresent the HSD, particularly for larger hail, and recommends deployment of multiple sensors to minimize this effect.

The concentration of large hail, and particularly giant hail (> 100 mm) can be very sparse (Witt et al., 2018), severely limiting the effectiveness of these small ground sensors even with multiple units. To overcome these sampling limitations, we describe a new time-integrated technique for measuring the HSD by combining aerial imagery captured from a small unmanned aircraft with deep learning and computer vision feature extraction. Methods involving Machine Learning and Artificial Intelligence deep learning have seen increased utilization in the atmospheric sciences community, including for the application of severe weather (e.g., McGovern et al., 2017; Gagne et al., 2019); however, there has been limited usage of such methods in targeted field observation datasets and in situ data. Over the last two decades, convolutional neural network's (CNN) have become a rapidly developing deep learning research tool that excels at image feature recognition (e.g., Razavian et al., 2014; Krizhevsky et al., 2012). This is achieved by developing complex feature recognition filters independent of prior knowledge, inspired by processes within the animal visual cortex (Hubel and Wiesel, 1968).

The new technique described here, named "HailPixel," enables the capture of very large areas (> 1500 m², equivalent area to several thousand hail pads) immediately following a hailstorm. This paper describes the methods of imagery capture and semi-automated extraction of the HSD using a combination of CNN and computer vision techniques. Results from a HailPixel survey of a hailstorm on 26 November 2018 in San Rafael (Argentina) are discussed in the context of existing studies and potential improvements for future surveys.

2 Data and Approach

Acrial imagery surveys of hail coverage were conducted in the Mendoza Provence of Argentina after hailstorms on 25 and 26 November 2018 during the RELAMPAGO field campaign (Nesbitt, 2019). To effectively extract the HSD from aerial imagery, hailstone size must of significantly larger than the effective ground resolution of the sensor and the concentration of hailstones must be sufficiently low so that overlapping stones are minimised. These requirements were satisfied only for Aerial imagery surveys of hail coverage were conducted in the Mendoza Provence of Argentina after hailstorms on 25 and 26 November 2018

during the RELAMPAGO field campaign (Nesbitt, 2019). Only the 26 November event produced non-accumulating hail of sufficiently large diameter (> 20 mm), and imagery from this event will be used throughout the paper. The hail swath observed from the 26 November event was produced by a marginally supercellular storm that developed in an environment of moderate instability and deep-layer shear. The storm initiated on the Andean Mountains and tracked approximately 120 km east-northeast towards the city of San Rafael before observations were made. A single hail pad (300 × 400 × 30 mm polystyrene foam block covered in aluminium foil) was also deployed 2 km southwest of the aerial survey site for the San Rafael hailstorm (34.6533° S 68.5030° W), providing a secondary measure of the HSD. To estimate hail size from hail pad indentations, the major and minor axis length of individual dents was measured with digital calipers and transformed into hail major and minor axis size using a relationship developed by the Community Collaborative Rain, Hail, and Snow Network (N. Doesken, personal communication, April 17, 2019).

2.1 Imagery

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A DJI Phantom 4 Pro V2 aircraft and Pix4DCapture flight-control software was used for image acquisition. The integrated, gimbal-mounted aircraft camera uses a 13.2×8.8 mm CMOS sensor which provides 20M effective pixels, automatic exposure, and an auto-focus lens with a focal length of 8.8 - 24 mm and maximum field of view of 84° . For the 26 November event, the aircraft was flown at an altitude of 10 m (relative to the take-off location) over a rectangular survey area of 1290 m^2 centred on 34.6459° S 68.4814° W, providing a \sim 2.7-mm ground sampling distance ((Fig. 1a). Near surface wind speed at the time of capture was noted by the authors to be a gentle breeze $(3.5 - 5.5 \text{ ms}^{-1})$, reducing the likelihood wind-induced motion blur. Images were captured with a 70% overlap laterally and medially at a flight speed of 1 m s⁻¹ over a surface consisting of sparse grasses, small shrubs, gravel, and dirt. A large image overlap and slow flight speed was selected to increasing the number of quality matching points during orthomosaic construction and reduce motion blur (Bemis et al., 2014). The survey was initialized immediately once hail fall concluded and required approximately 4 minutes to complete. The location of images was measured using the integrated GPS receiver, which has an accuracy of ± 1.5 m. Precise location measurements (e.g., real-time kinematic positioning) are not essential for improving the pixel size accuracy during photogrammetry processing (Strecha, 2012).

The Pix4DMapping software package was used to generate orthomosaic imagery and a digital elevation model (DEM) from the survey photos (Strecha, 2012) with a ground sampling distance of 2.7 mm. The software is based on the Structure from Motion photogrammetry technique and uses the following automated steps:

- 1. Tie points between the survey images are identified. Each tie point must be matched in at least 3 images
- 2. Tie points are combined with positioning and orientation information from the aircraft autopilot to reconstruct the camera perspective and position for each survey image. This information is used to verify the quality of matching points and calculate the 3D coordinates of tie points.
- 3. The sparse point cloud of 3D coordinates is interpolated to obtain a gridded DEM.

4. The DEM is used to project every image pixel and to calculate a orthomosaic.

An average of 181,081 matched tie points were found per m³ with a mean geolocation error of less than 1 mm. Analysis of the DEM indicates a gradual slope was present across the survey area with a total change in elevation of approximately 2.3 m (not shown). Two scale markers consisting of 300×300 mm black and white vinyl tiles were also placed into the aerial survey area to provide a secondary check of pixel size within the orthomosaic.

2.2 Hail Detection

To efficiently identify the many thousands of hailstones captured in the aerial imagery after the San Rafael hailstorm, automated feature detection techniques were explored. Simple thresholding of pixel luminosity for detecting hailstones performed poorly owing to similar luminosity from sparse grasses, light dirt patches, pale colored rocks and leaf debris, and for instances where hailstones were in contact. Despite the low contrast, hailstones were easily identifiable in the imagery by human observers, motivating the application of a convolutional neural network (CNN) model. This class of machine learning algorithms develops complex feature recognition filters independent of prior knowledge, inspired by processes within the animal visual cortex (Hubel and Wiesel, 1968). Over the last two decades, CNN's have become a rapidly developing research tool that excels at image recognition (e.g., Razavian et al., 2014; Krizhevsky et al., 2012). For hail detection, the state-of-art Mask Region-CNN mask region-CNN (R-CNN) model was implemented (He et al., 2017). This technique combines the optimised selection and parallel processing of proposed feature regions (Fast R-CNN) with semantic segmentation, whereby each pixel is classified. Mask R-CCN architecture and implementation used is described in detail by He et al. (2017).

To reduce memory requirements, the 489-megapixel aerial survey orthomosaic was divided into 1961 tiles of size 600×600 pixels, including a 50-pixel overlap along edges with neighbour tiles to avoid cropped hailstones (Fig. 1b, 2a,b). For To provide a sufficiently large sample of hailstones for training the Mask R-CNN model, 12 tiles were selected to represent the varying background conditions and were manually manually selected that in total contained more than one thousand stones and were annotated using the VGG Image Annotator (VIA) tool (Dutta and Zisserman, 2019). These tiles were also selected to sample the varying background types across the orthomosaic. Nine tiles were randomly selected for training (containing 729 annotated hailstones) and the remaining three for validation (Fig. 1c). The Mask-RCNN training was initialized with the pre-trained weights from the Microsoft COCO dataset (set of $> 2 \times 10^5$ labelled images; Lin et al. (2014)), which capture many features in natural images. Utilizing these weights greatly reduces the training time required to recognize hail. The default learning and weighting configuration described by He et al. (2017) were applied and training was performed on 8 GPUs with 1 image per GPU for 3,000 iterations (\sim 43 minutes of computation time). The trained model detected more than 94% of hailstones in each validation tile with a false alarm rate of <1%. When applied to all tiles, the trained Mask R-CNN model detected a total of 46.871 hailstones.

2.3 Hail Size Measurement

The segmentation mask generated by the Mask R-CNN model was initially tested for hail size measurement, but found to contain small errors that rendered it unsuitable. To provide the pixel-level accuracy required for measuring hailstones, an edge detection algorithm was developed to find the steep "lightness" gradient at the hailstone edge that occurs radially from the hailstone centroid (Figure 1d). The HSL color space is an alternative to Red-Green-Blue (developed for color displays) that is commonly used in computer vision applications for reducing the correlation between colors (Cheng et al., 2001). The hailstone centroids required to initialize the edge detection technique are derived from the segmentation mask. Two additional quality control steps are also applied to the centroids and image tiles:

- 1. Tiles where hail was obscured (e.g., under long grass or shrubs) or water had accumulated were removed, leaving 188 "clean" image tiles containing 15,983 hailstones over a total area of 342.6 m² for hail size measurement.
 - 2. Hail centroids for the 188 clean image tiles were manually assessed and amended if required using the VIA annotation tool.

The clean image tiles are next transformed into the HSL color space and the hailstone size is measured for every centroid using the following procedure. First, coordinates of 12 equally spaced radials of length 20 pixels from the centroid (p_0) are calculated (Fig. 2c), denoted as p_i^k Where i is the pixel index (i = 1,...20) and k is the radial index (k = 1,...,12). For all points along a radial, the lightness values $L(p_i^k)$ are extracted. Then, the gradient of lightness values $L'(p_i^k)$ along each radial are calculated. Starting from the centroid of each radial, the edge point is found at coordinate p_i^k when the following criteria are met:

$$L'(p_i^k)<0.75\times L'(p_{i-1}^k)$$

and

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$$L(p_0) - L(p_i^k) > 50$$

The required minimum lightness difference between the hailstone centroid and background (50) was found to perform well across all background types, including light colored soils. Once all edge points are found along the radials (Figure 2d), the median distance \tilde{d} from the centroid is calculated for each edge point. If an edge point falls outside the range $[\tilde{d} \times 0.5]$ to $\tilde{d} \times 1.5$, it is replaced by \tilde{d} . Finally, to measure the major and minor axis length of the hailstone, the minimum bounding box (allowing for rotation) is calculated for the set of edge points.

3 Results and Discussion

The resulting distribution of major axis length and axis ratio for the San Rafael hailstorm is shown in Figure 4, along with the distributions obtained from the hail pad (total of 17 impacts). Comparison of the major axis length distribution from the

¹lightness here is from the Hue-Saturation-Lightness (HSL) color space with a range 0-255

HailPixel and hail pad techniques clearly demonstrate the value of aerial photogrammetry: the large population size (n = 15983) of the aerial survey provides a defined distribution shape and tails (Figure 4). The HailPixel distribution peak is 2.5 mm lower than the hail pad peak, possibly due to melting of hail on the ground before it was photographed or uncertainty in hail size retrievals. The distribution shape is well approximated by a Gamma probability distribution function (PDF) with a mostly absent lower quartile and long upper tail. The gamma PDF was also found to be most suited for major axis length in a number of other case studies studies, including Ziegler et al. (1983) for Okalahoma (US), Wong et al. (1988) for the Albert (Canada), and Fraile et al. (1992) for León (France). Spain). Impact concentration observed by the hail pad was 141 m⁻², significantly higher than the mean hail concentration observed by the aerial survey (47 m⁻²). This difference is speculated to be the product of longer hail fall duration at the hail pad location (2 km south west) and possible secondary impacts on the hail pad from bouncing stones.

The distribution of axis ratios from the HailPixel survey is well approximated by an exponentially increasing function (not shown); however, a local maximum at 0.8-0.85 suggests a more complex underlying distribution. Note that oblate hailstones are most likely to rest on the surface with their minor axis orientated vertically; thus, the HailPixel technique would not measure the true minor axis length in this scenario, but rather provides estimates of the intermediate axis ratio (assuming ellipsoidal geometry for hailstones). This limitation is likely to also effect hail pad measurements when tumbling motions of oblate hailstones are not too extreme. Comparing the HailPixel intermediate/major axis distribution with Giammanco et al. (2014) major/minor axis distributions demonstrates the expected skew towards higher axis ratios in the HailPixel dataset.

Both HailPixel and hail pad data demonstrate decreasing axis ratio with increasing hail size. This shape of this relationship becomes apparent when the highly variable HailPixel data are binned into 5-mm intervals. For the 20-30-mm hail size range, axis ratio remains constant and close to 0.9. For sizes > 30 mm, axis ratio decreases by 0.4-0.5% mm⁻¹. Despite the potential bias in axis ratio measurements, the shape of this trend is comparable to observations by Knight (1986) for an Alberta (Canada) hailstorm. Another study using a 3-year database of hailstones collected from the Great Plains by Giammanco et al. (2014) demonstrates a less significant decreasing trend between hail size and axis ratio, and less spherical stones for smaller sizes. It is likely that this relationship is also highly variable between hailstorm cases and within hailstorms (Federer and Waldvogel, 1975; Ziegler et al., 1983; Knight, 1986).

It is also important to highlight the optimal conditions and configuration for future HailPixel surveys. We recommend avoiding inhomogeneous background surfaces if possible, with cut or grazed turf grasses being most ideal. A uniform and contrasting background will likely permit the use of less complex hail detection and sizing techniques. Large survey areas $(>1000 \text{ m}^2)$ are only necessary when very sparse giant (>100 mm) hailstones are present. Assuming a normally distributed sample mean, a sample size of 2088 hailstones is required to represent the population mean (from 15,983 hailstones) within a 2% confidence level at the 95% significance level. This sample size equates to a sample area of 40.1 m^2 for the 26 November survey; however, an area of at least $\sim 250 \text{ m}^2$ is recommended to adequately resolve the tails of the distribution. For the DII Phantom 4 Pro V2 aircraft flown at a 10 m altitude, the authors recommend a minimum hail diameter of 20 mm (major axis length) and less than 30% total ground coverage. The minimum size limit is particularly critical when separating multiple stones in contact. Higher resolution imagery would allow for small (< 20 mm) hailstones to be measured, but the increased

susceptibility to motion blur would likely require the aircraft to remain stationary during image capture. Further, 10 m winds exceeding a moderate breeze ($> 8 \text{ ms}^1$) would increase the likelihood of motion blur. To quantify the measurement uncertainty from the HailPixel technique, we recommend that hail within a 1 m^2 area of the aerial survey is manually measured for three orthogonal axes immediately following aerial capture. Finally, minimizing the melting of hailstones is critical. Where possible, avoid aerial surveys of areas where water may flow or accumulate and conduct surveys immediately after hail fall ceases.

4 Summary

This paper describes the novel HailPixel aerial photogrammetry technique for measuring time-integrated HSD's after cessation of hail fall. The workflow for collecting imagery, detecting hail stones, and measuring hail size is described, including the use of state-of-art mask R-CNN image segmentation algorithm. Results from a HailPixel survey after the 26 November 2018 San Rafael (Argentina) hailstorm are compared with observations from a co-located hail pad. Despite potential bias of axis ratio measurements, the HSD distributions and relationships observed for the San Rafael hailstorm are comparable to previous studies, although with substantially larger population sizes. Further HailPixel surveys are encouraged to quantify the variability of these distributions, particularly for hailstorms producing sparse giant hail. In contrast to hail impact sensors, the use of aerial imagery provided a sample area that is several orders of magnitude larger (341.6 m²). As a result, the hailstone population size from the 26 November 2018 imagery was substantially larger than previous studies (15,983 hailstones), providing a more robust analysis of the HSD. Ongoing work to relate HailPixel results with mobile polarmetric radar observations from the 26 November 2018 San Rafael hailstorm will explore the signatures of hail size and swath extent for this event. Future HailPixel surveys are encouraged to quantify the variability of these distributions, particularly for hailstorms producing sparse giant hail.

20 Data availability. All imagery, hail pad and hail retrieval data used for this work are publicly available: Soderholm. (2019). HailPixel Survey Data and Analysis from 26 November 2018, San Rafael, Argentina [Data set]. Zenodo. http://doi.org/10.5281/zenodo.3383227

Author contributions. Soderholm designed the HailPixel technique and Soderholm, Kumjian, McCarthy and Maldonado applied the technique following hailstorms during the RELAMPAGO project. Soderholm developed the hail identification and sizing algorithms with assistance from Wang. Soderholm prepared the manuscript with contributions from all co-authors.

25 Competing interests. The authors declare that they have no conflict of interest.

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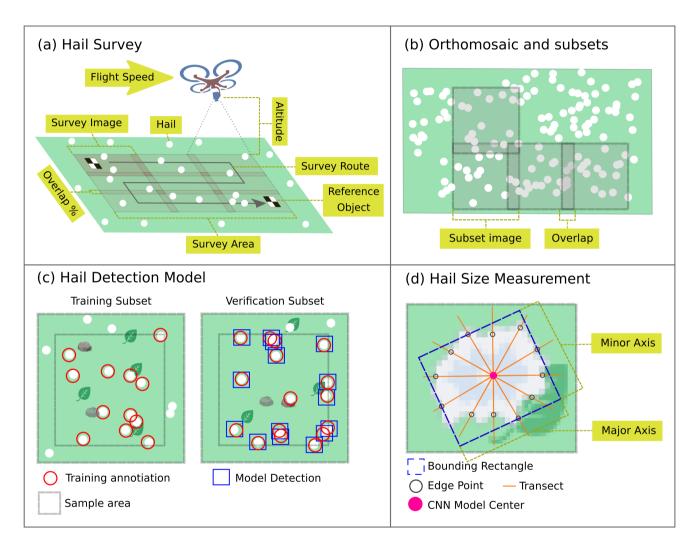


Figure 1. Workflow of (a) data collection, (b) tiling of orthomosaic, (c) hail detection using the Mask-RCNN technique and (d) hail size measurement using radial transects.

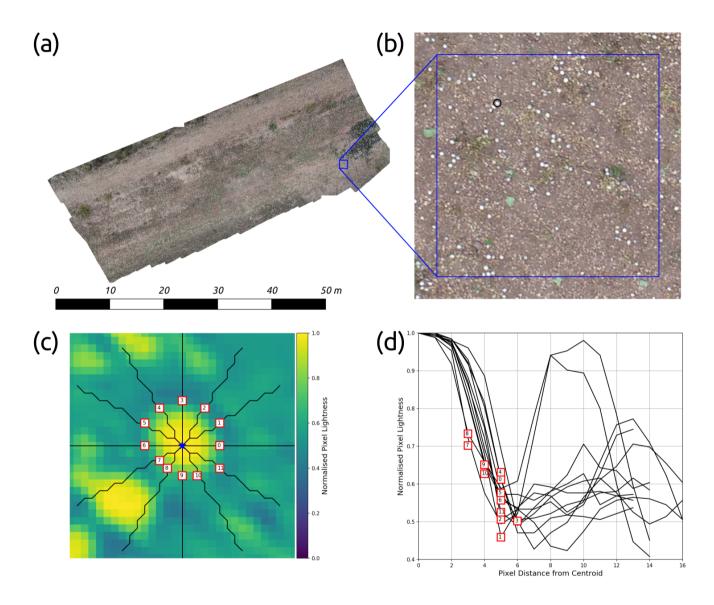


Figure 2. Demonstration of hail size extraction from the 26 November 2018 survey orthomosaic (a) from a single tile (b; blue bounding box) for a single hail stone (c, black circle in b). Radial transects for extracting imagery lightness are shown in (c) as black lines radiating from hailstone centroid (blue marker) and hailstone edge pixels along transects are numbered. Subplot (d) shows the normalised pixel lightness along the 12 transects shown in (c) with the corresponding edge pixels marked.

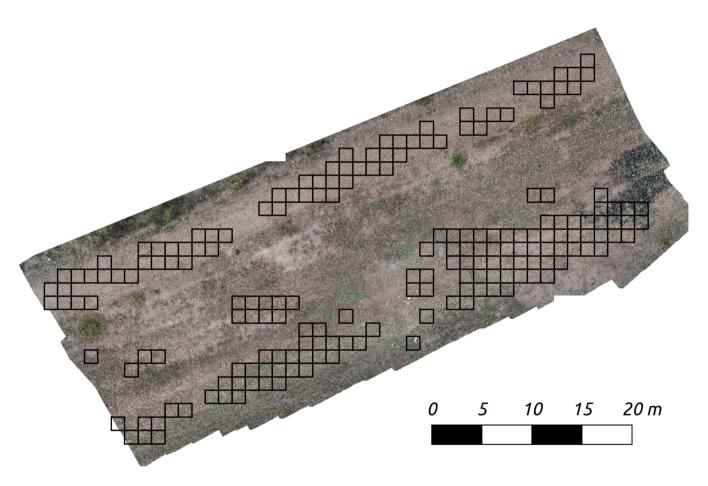


Figure 3. Orthomosaic RGB imagery from the 26 November 2018 survey overlaid with the outlines of tiles used for the extraction of hail size distribution statistics

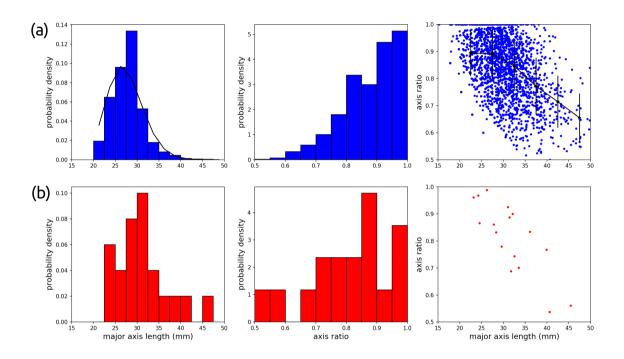


Figure 4. Distribution of hail major axis length, minor axis ratio and scatter plot of axis ratio and major axis length from the (a) photogrammtery and (b) pad hail size retrievals for the 26 November 2018 survey. A fitted Gamma distribution probability density function for photogrammtery major axis size distribution is shown (black line). Photogrammtery scatter plot observations are binned using 5-mm bin sizes and error bars represent ± 1 standard deviation.