

Some further comments that should be looked at in a revised manuscript:

1) Table 1 is potentially useful but could be improved. Some of the points here also relate to broader discussion in the text on uncertainty.

- “Sampling flow rate” instead of “flow rate”

It has been changed

- Detection limit for ANSTO is 0.03 but an approximation is given for the others. Why is this? A footnote explaining, e.g. differences in definitions, would be good.

The ARMON’s detection limit was calculated following Gilmore, 2008, as reported in Grossi et al., 2012 with a confidence level of 95%. The Detection Limits of the others instruments were calculated as presented in their reference papers (reported in the last column of the table for more details) as the ambient radon concentration at which the estimated counting error reaches 10% (Levin et al., 2002) and 30% (Chambers et al., 2016). We have now used this latter to harmonize the table column of the detection limits for all instruments.

It is important to underline that in the present study we did not measure the radon concentration background of each instrument for harmonizing the calculation of the detection limits properly because the inter-comparison campaign was carried out in field conditions and it was out of the scope of the study.

- Over what range is the stated uncertainty relevant? This is important for matching the right instrument to the right application or measurement location. A lot of the measurements in the paper’s time series border on the detection limit of the ARMON. Is a 20% uncertainty the case at 0.3 Bqm-3, for instance? For the ANSTO there is a discussion of some of this on page 5 “a counting uncertainty of around 2% for radon concentrations ≥ 1 Bq m-3”, and a discussion for the HRM at the bottom of page 6. The discussion of uncertainties and what is stated in the table needs to be completely transparent for comparison between instruments.

The total expanded uncertainties of all monitors have been now presented coherently with $k = 2$. The ARMON was calibrated within the INTE-UPC chamber at a range of hundreds Bq m⁻³ because, so far, European radon chamber facilities are not able to create low level radon reference air. An uncertainty of <35% ($k=2$) was estimated for atmospheric radon levels of few Bq m⁻³. A sentence has been included to explain this. In addition in the next future we want to carry out a long term intercomparison campaign, in the mark of the new EMPIR project traceRadon, in order to correctly harmonize and calculate the uncertainties of HRM, ANSTO and ARMON monitors using the same reference radon air and background.

- The portability column could be improved. A grading such as low/high might not be useful. Instead call this “portability considerations” and let the potential user decide based on their specific circumstances. Please state the three measured dimensions of each instrument in the description rather than a volume (which is difficult to physically relate to), and add the mass of the instruments – this is obviously very important too in terms of transportation and handling.

As suggested by the reviewer we have reported these values within the column and changed the column names. However each instrument is composed from different parts, not only the main detection volume. There are also pump, filters or electronics components. Depending on the instrument. All such details are already reported in the reference papers of each instrument.

- Alongside portability is “deployability” i.e. level of automation, consumables required, energy consumption, which might be of even greater interest than portability. The basic monitor also needs peripherals e.g. large pumps, cryocoolers etc.

A new column has been created in the table where we have reported the main needs of each instrument (dry air, possibility to check the spectrum, remote connection, etc.). Other aspects such as filter, maintenance, etc. are interesting and they have been reported in more detail in the text because within the table columns there is not enough space. None of the instruments consume a large amount of energy, so this does not make a significant difference between them.

2) The conclusions and abstract need rephrasing and tightening up. Some things below but not exhaustive.

We have worked on improving the conclusions and the abstract. We have also applied the suggested changes.

- The last sentence on page 17 is very confusing. What is “close to one” – the regression line? But that is not referred to in the sentence.

We have changed the sentence.

- “last behaviour” change to “the latter”

The change has been applied

- Line 463 “very good” to “significant”

The change has been applied

- Line 464 “slope of this correlation”. This correlation discussion is confusing given the stated small uncertainties on the slopes stated alongside “within uncertainties well comparable”. Please explain.

Here we were comparing the slope of the regression fit calculated between the LSCE and the HRM monitors at ODM station during this study with the same slope calculate by Schmithüsen et al. (2017). We have changed the sentence.

- “underlines that to assure” .. “is important” – revise sentence structure.

The sentence has been changed.

- So does the ARMON help to meet the requirements on lines 476-480? It is stated that the ARMON has great potential but not why specifically in relation to what is needed in networks. Can you explain why further inter-comparison with the ANSTO is needed?

The text has been improved in order to clarify these points.

- Line 34 “daily basis”. Not sure what this means – daily averages or within days?

We mean that the monitors were all able to observe the changes of the atmospheric radon concentration during the day: the nocturnal accumulation of the radon concentration during the night due to the shallow planetary boundary layer and the diurnal dilution of the concentration due to increase of the turbulence.

- Lines 42 to 44 refer to the same points made at the end of the conclusion. This leaves the reader unclear as to what has been advanced in this work and what is needed next.

The sentence has been changed to better differentiate between what has been done here and what still needs to be done.

Specifics:

“close to and further up” change to “when sampling at 2 and 100 magl”

It has been changed

Minor corrections/explanations needed:

Page 5 line 164 “measurement uncertainty”

Corrected

Page 5 mentions “detection limit”, page 6 mentions “minimum detectable activity”. This should be consistent throughout if these are referring to the same thing.

This has been clarified in the beginning of this document and it has been corrected in the text.

Table 1 “Need of .. height of inlet” could just be “Sampling inlet height correction”

It has been changed accordingly

Table 1 Uncertainty of HRM is 15-20% but in text <20%. Just be consistent with these reported values throughout the text so that the instruments can really be compared.

Values have been reported coherently in Table 1 now.

Page 6 line 190 – give details of the cryocooler

A sentence on this has been added within the text with more reference to past studies.

Page 17 – make space between number and unit.. 100 m.. 2 m etc

It has been done

Page 13 – what is the approximated response time correction?

For the deconvolution routine of Griffiths et al (2016) to be run in its intended form it is necessary to perform a source “spike test” at the sampling inlet (so that the combined characteristics of the whole intake system, delay volume and detector can be taken into account by the model). Unfortunately, we were not able to perform a “spike test” on the detectors installed for this inter-comparison campaign, so we estimated the characteristics based on what we knew of the setup, and performance of similar detectors. A small paragraph has been added to clarify this point.

Inter-comparison study of atmospheric ^{222}Rn and ^{222}Rn progeny monitors

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Abstract.

The use of the noble gas radon (^{222}Rn) as tracer for different research studies, for example observation-based estimation of greenhouse gas (GHG) fluxes, has led to the need of high-quality ^{222}Rn activity concentration observations with high spatial and temporal resolution. So far a robust metrology chain for these measurements is not yet available.

A portable direct Atmospheric Radon MONitor (ARMON), based on electrostatic collection of ^{218}Po , is nowadays running at Spanish stations. This monitor has not yet been compared with other ^{222}Rn and ^{222}Rn progeny monitors commonly used at atmospheric stations.

A 3-month inter-comparison campaign of atmospheric ^{222}Rn and ^{222}Rn progeny monitors based on different measurement techniques was realized during the fall and winter of 2016-2017 to evaluate: i)

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calibration and correction factors between monitors necessary to harmonize the atmospheric radon observations; and ii) the dependence of each monitor's response in relation to the sampling height, meteorological and atmospheric aerosol conditions.

Results of this study have shown that: i) all monitors were able to reproduce the atmospheric radon variability on daily basis; ii) linear regression fits between the monitors exhibited slopes, representing the correction factors, between 0.62 and 1.17 and offsets ranging between -0.85 Bq m^{-3} and -0.23 Bq m^{-3} when sampling 2 m above ground level (a.g.l.). Corresponding results at 100 m a.g.l. exhibited slopes of 0.94 and 1.03 with offsets of -0.13 Bq m^{-3} and 0.01 Bq m^{-3} , respectively; iii) no influence of atmospheric temperature and relative humidity on monitor responses was observed for unsaturated conditions at 100 m a.g.l. whereas slight influences (order of 10^{-2}) of ambient temperature were observed at 2 m a.g.l.; iv) changes of the ratio between ^{222}Rn progeny and ^{222}Rn monitor responses were observed under very low atmospheric aerosol concentrations.

~~A more statistically robust evaluation of these last influences based on a longer dataset should be conducted to improve the harmonization of the data. Results also show that the new ARMON has a great potential to be used in atmospheric radon monitoring stations with space restrictions or as a mobile reference instrument to calibrate in situ ^{222}Rn progeny monitors and fixed radon monitors networks. However, its qualities and faults should be deeply investigated. In the near future a long-term comparison study between an ARMON, HRM and an ANSTO monitors will be useful to better evaluate: i) the radon uncertainties of the measurements of the radon concentration in the range of a few hundreds mBq m^{-3} and to a few Bq m^{-3} ; and ii) evaluate the response time correction of the ANSTO monitor for representing fast changes in the atmospheric ambient radon concentrations.~~

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1 Introduction

Over continents, the natural radioactive noble gas radon (^{222}Rn) (half-life $T_{1/2} = 3.8$ days) is continuously generated within the soil from the decay of radium (^{226}Ra) (Nazaroff and Nero, 1988; Porstendorfer, 1994) and it can then escape into the atmosphere by diffusion, depending on soil characteristics and meteorological conditions (Grossi et al., 2011; Lopez-Coto et al., 2013; Karstens et al., 2015). The global ^{222}Rn source into the atmosphere is mainly restricted to land surfaces (Szegevary et al., 2009; Karstens et al., 2015), with the ^{222}Rn flux from water surfaces considered negligible for most applications (Schery and Huang, 2004).

In recent decades the atmospheric scientific community has been addressing different research topics using ^{222}Rn as a tracer. Examples of such applications include: the improvement of inverse transport models (Hirao et al., 2010), the improvement of chemical transport models (Jacob and Prather, 1990; Chambers et al. 2019a), the study of atmospheric transport and mixing processes within the planetary boundary layer (Zahorowski et al., 2004; Galmarini, 2006; Baskaran, 2011; Chambers et al., 2011, 2019b; Williams et al., 2011, 2013; Vogel et al. 2013; Vargas et al., 2015; Baskaran, 2016), the experimental estimation of greenhouse gas (GHG) fluxes (Levin et al., 1999; 2011; Vogel et al., 2012; Wada et al., 2013; Grossi et al., 2018), and others listed in Grossi et al. (2016).

In light of this, atmospheric ^{222}Rn measurements are being carried out at numerous monitoring stations of GHG concentrations and air quality using three fundamentally different measurement principles: one filter, two filters, and electrostatic deposition (Stockburger and Sittkus, 1966; Polian, 1986; Hopke, 1989; Whittlestone and Zahorowski, 1998; Paatero et al., 1998; Levin et al., 2002). The two most commonly employed measurement systems at European ^{222}Rn monitoring stations are: the dual-flow-loop two-filter monitor (Whittlestone and Zahorowski, 1998; Zahorowski et al. 2004; Chambers et al., 2011, 2014, 2018; Griffith et al., 2016), which samples and measures radon directly, and the one-filter monitors, of which several kinds are in use (e.g. Stockburger and Sittkus, 1966; Polian, 1986; Paatero et al., 1998; Levin et al., 2002), which sample and measure aerosol-bound radon progeny. Finally, a third method is being used at several Spanish atmospheric stations (Vargas et al., 2015; Hernández-Ceballos et al., 2015; Grossi et al., 2016; Frank et al., 2016; Grossi et al., 2018; Gutiérrez-Álvarez et al., 2019). This type of instrument performs a direct measurement of ^{222}Rn and ^{220}Rn (thoron) activity concentrations using the already existent method based on the electrostatic deposition of ^{218}Po and ^{216}Po , respectively (Hopke, 1989; Tositti et al., 2002; Grossi et al., 2012).

The diversity of these three aforementioned measurement techniques could introduce biases or compatibility issues that would limit the comparability of the results obtained by independent studies and the subsequent application of atmospheric radon data for regional-to-global investigations (e.g. Schmithüsen et al., 2017). Thus, a comparative assessment of all the experimental techniques applied for atmospheric ^{222}Rn activity concentration measurements and a harmonization of their datasets is needed, as suggested by the International Atomic Energy Agency (IAEA, 2012).

Xia et al. (2010) carried out a comparison of the response of a dual-flow-loop two-filter detector from the Australian Nuclear Science and Technology Organisation (ANSTO, Whittlestone and Zahorowski 1998) and a one-filter monitor (α/β Monitor P3) manufactured by the Bundesamt für Strahlenschutz, Germany (BfS) (Stockburger and Sittkus, 1966), for atmospheric ^{222}Rn measurements under various meteorological conditions at 2.5 m above ground level (a.g.l.) over one year. Their results showed that both systems followed the same patterns and produced very similar results most of the time, except under specific meteorological conditions such as when precipitation or the proximity of the forest canopy could remove short-lived progeny from the air mass to be measured by the one-filter monitor. However, Xia et al. (2010) did not find a clear relationship between precipitation intensity and the ratio between progeny-derived ^{222}Rn and ^{222}Rn activity concentration to convert the progeny signal to ^{222}Rn activity concentration.

Grossi et al. (2016) presented results from two short (about 7-9 days) comparisons between a one-filter monitor from Heidelberg University (HRM; Levin et al., 2002), and an Atmospheric Radon MONitor (ARMON, Grossi et al., 2012), an electrostatic deposition monitor from the Universitat Politècnica de Catalunya (UPC). The two comparison campaigns were carried out at a coastal and a mountain site, with sampling in both cases from 10 m a.g.l. These comparisons revealed that the responses of both monitors were in agreement except for water saturated atmospheric conditions or periods of rainfall. Again, the quantity of comparison data was not sufficient to confirm any statistical correlation.

Loss of aerosols in the air intake systems can also complicate the derivation of ^{222}Rn activity concentrations from one-filter systems such as the HRM. Levin et al. (2017) carried out an assessment of ^{222}Rn progeny loss in long tubing by laboratory and field experiments. Results of these experiments, for 8.2 mm inner diameter (ID) Decabon tubing, gave an empirical correction function for ^{222}Rn progeny measurements, which enables the correction of measurements for this specific experimental setup (e.g. tubing type and diameter, flow rate, aerosol size distribution).

Finally, Schmithüsen et al. (2017) conducted an extensive European-wide $^{222}\text{Rn}/^{222}\text{Rn}$ progeny comparison study in order to evaluate the comparative performance of one-filter and two-filter measurement systems, determining potential systematic biases between them, and estimating correction factors that could be applied to harmonize ^{222}Rn activity concentration estimates for their use as a tracer in various atmospheric applications. In this case, the authors employed a -HRM monitor as the reference device. It was taken to nine European measurement stations to run for at least one month at each of them. This -monitor was run in parallel to other one-filter and two-filter radon monitors operating at each station of interest.

Although several inter-comparison campaigns have been carried out in the past, none of them has included simultaneous observations from one-filter, two-filter and electrostatic deposition methods. Here, we present the results of a three-month inter-comparison campaign carried out in the fall and winter of 2016-2017 in Gif Sur Yvette (France) where, for the first time, co-located measurements from monitors based on the three measurement principles were included. Two two-filter ^{222}Rn monitors, two single-filter ^{222}Rn progeny monitors and an electrodeposition monitor were run simultaneously under different meteorological and aerosol conditions sampling from heights of 2 and 100 m a.g.l.

The main objectives of the present study were to: i) compare the calibration and correction factors between all monitors required to derive harmonized atmospheric radon activity concentrations; and ii) analyze the influence that meteorological and environmental parameters, as well as sampling height, can have on the finally determined ^{222}Rn activity concentration.

In the present manuscript the applied methodology is reported, including a short presentation of the $^{222}\text{Rn}/^{222}\text{Rn}$ progeny monitors participating in the campaigns, the sampling sites and the statistical analysis carried out. Finally, the outcomes of the present study are discussed and compared with the ones from Schmithüsen et al. (2017).

2 Methods

In section 2.1 a short description is given of the monitors compared in the experiment, mainly focusing on measurement techniques, instrument calibration and maintenance. The main characteristics of these monitors are then summarized in Table 1. Section 2.2 presents the French atmospheric stations of Orme de Mérisiers (ODM) and Saclay (SAC) where the two phases of the inter-comparison campaign were realized. Section 2.3 briefly describes the devices used to measure the environmental parameters and the atmospheric aerosol concentration at the above sites during the experiments. Finally, the statistical analysis applied is described in section 2.4.

2.1 ²²²Rn and ²²²Rn progeny monitors

2.1.1 Direct methods

Dual-flow-loop two-filter detectors

The two 1500 L dual-flow-loop two-filter detectors included in this exercise were designed and built at the Australian Nuclear Science and Technology Organisation (ANSTO). This model of detector, which will henceforth be named ANSTO, is based on a previous design by Thomas and Leclare (1970), with some early iterations of the modified design being described by Whittlestone and Zahorowski (1998) and Brunke et al. (2002). The subsequent evolution of two-filter detectors in recent decades, and the current principle of operation, has been described in detail by Williams and Chambers (2016) and Griffiths et al. (2016).

During the measurement campaign ambient air was sampled continuously at a rate of about 83 L min⁻¹ through a 50 mm ID HDPE inlet tube and a 400 L delay volume to allow decay of the short-lived ²²⁰Rn ($T_{1/2} = 56$ s). The air stream then passes through the first filter, which removes all ambient aerosols as well as ²²²Rn and ²²⁰Rn progeny. The filtered sample, now containing only aerosol-free air and ²²²Rn gas, enters the main delay volume (1500 L) where ²²²Rn decay produces new progeny. The newly formed ²¹⁸Po and ²¹⁴Po are then collected on a second filter and their subsequent α decays are counted with a ZnS photomultiplier system. Atmospheric ²²²Rn activity concentrations are then calculated from the α count rate and the flow rate through the chamber.

The detection limit (L_D) of two-filter detectors is directly related to the volume of the main delay chamber. Here, L_D is understood to represent the ambient radon concentration at which the estimated counting error of the instrument reaches 30%. The ~~L_D lower limit of detection~~ of the 1500 L model used in this study was around 0.03 Bq m⁻³. Under normal operation ANSTO monitors are automatically calibrated in situ every month by injecting radon into the sampling air stream from a well-characterized Pylon ²²⁶Ra source (ca. 41 kBq radium at SAC station) for 5 hours at a fixed flow rate of ~100 cc min⁻¹. Automatic instrumental background checks, each lasting 24 hours, are also performed every 3 months to keep track of long-lived ²¹⁰Pb accumulation on the detectors second filter (which should be changed every 5 years). Based on a calibration source uncertainty of 4%, coefficient of variability of valid monthly calibrations of 2-6%, and a counting uncertainty of around 2% for radon concentrations ≥ 1 Bq m⁻³, the total measurement uncertainty of 1500 L ANSTO radon detectors is typically between 8%- and 12% ($k = 2$). The ANSTO monitors have low-maintenance requirements but, due to their dimensions (2.5 – 3m long) it can be challenging to install them at stations with space restrictions. As an alternative to the 1500 L detectors, a 700 L model is also available, which is more portable and has a ~~L_D detection limit~~ of around 0.04-0.05 Bq m⁻³. The combination of detector volume, operating flow rate, and radon decay chain result in ANSTO monitors having a response time of ~45 minutes, which can be corrected for in post processing (Griffiths et al. 2016).

Two ANSTO monitors were used during this study. As explained later in the text these monitors are permanently running at SAC and ODM stations. No calibration source was available when the ANSTO

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monitor was installed at the ODM site, so calibration and background information derived prior to transport have been used.

Electrostatic deposition monitor

The Atmospheric Radon Monitor (ARMON) used in this experiment was designed and built at the Institut de Tècniques Energètiques (INTE) of the UPC. The ARMON is a portable instrument based on electrostatic deposition method, consisting of alpha spectrometry of positive ions of ^{218}Po electrostatically collected on a detector (Hopke, 1989; Pereira and da Silva, 1989; Tositti et al., 2002). The ARMON is described in detail in Grossi et al. (2012).

Sampled air with a flow rate between 1-2 L min⁻¹, is first filtered to remove ambient ^{222}Rn and ^{220}Rn progeny and then pumped through a ~20 L spherical detection volume uniformly covered internally with silver. Within this volume the newly formed ^{222}Rn and ^{220}Rn progeny, i.e. positive ^{218}Po and ^{216}Po ions, respectively, are electrostatically collected on a Passivated Implanted Planar Silicon (PIPS) detector surface by an electrostatic field inside the spherical volume. An 8 kV potential is applied between the PIPS detector base and the sphere walls. As for the ANSTO detector, the sensitivity of this instrument type depends on the detector volume. The design of the monitor employed in this study ~~has allows as a~~ minimum detectable activity concentration of about 0.072 Bq m⁻³ in agreement with definition given above. (Grossi et al., (2012) reported a minimum detection limit for this instrument of around 0.2 Bq m⁻³ in agreement with the definition of Gilmore, (2008). The measurement efficiency of the electrodeposition method is reduced due to neutralization of the positive ^{218}Po in recombination with OH⁻ ions in the sampled air (Hopke, 1989). Consequently, it is necessary to dry the sampled air as much as possible before it enters the detection volume. To this end, a dew point of < -40°C was maintained at both inter-comparison sites using a cryocooler, ~~consisting of a vessel tube where sampling air was passing through before reaching the radon monitor~~ (Grossi et al., 2018).

Each ARMON is calibrated at the INTE-UPC ^{222}Rn chamber (Vargas et al., 2004) under different ^{222}Rn and relative humidity conditions (Grossi et al., 2012). The radon chamber of the INTE-UPC is a 20 m³ installation, which allows control of the exhalation rate (0-256 Bq min⁻¹) and the ventilation air flow rate (0-100 L min⁻¹). The ^{222}Rn source is a dry powder material containing 2100 kBq ^{226}Ra activity enclosed in the source container (RN-1025 model manufactured by Pylon Electronics). The calibration factor F_{cal} of the ARMON used in this study was of 0.39 counts per minute (cpm) per Bq m⁻³ with an uncertainty of 10% (k=2). The correction factor for the humidity influence inside the sphere was of 6.5·10⁻⁵ per part per million H₂O (ppm) with a maximum uncertainty of 10% (k=2). The total uncertainty of the atmospheric radon activity concentration measured by the ARMON is of about 20% (k=2) for atmospheric ^{222}Rn levels in the range of a few hundred Bq m⁻³ but could increase up to 35% (k = 2) when atmospheric ^{222}Rn levels decrease to a few Bq m⁻³ due to the increase of the error of the alpha counts. The total uncertainty where it is including the calibration factor F_{cal} , the background due to the presence of ^{212}Po from ^{220}Rn , the net ^{218}Po counts and the humidity correction factor (Grossi et al., 2012; Vargas et al., 2015). Every 1-2 years the progeny filter at the ARMON inlet should be changed. The detection volume of the ARMON is safety isolated because it is located within an external wooden cube of 0.18 m³.

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2.1.2 Non direct methods

One-filter monitors

One-filter detectors measure the decay rates of aerosol-bound ^{222}Rn progeny directly accumulated by air filtration (Schmithüsen et al., 2017). The ^{222}Rn activity concentration is then calculated assuming a constant disequilibrium factor (F_{eq}) for a given site and sampling height between ^{222}Rn and the measured progeny in the sampled air.

In the present study two monitors based on this method were used. One, named here as HRM, was developed at the Institute of Environmental Physics of Heidelberg University, Germany, and is described in detail by Levin et al. (2002). Rosenfeld (2010) describe the most recent version of this monitor for which the electronics, data acquisition, and evaluation hardware and software were modernized. The HRM measurement is based on α spectrometry of ^{222}Rn daughters attached to atmospheric aerosols collected on a static quartz fiber filter (QMA Ø 47 mm) using a surface barrier detector (Canberra CAM 900 mm² active surface). The ~~L_d detection limit~~ of the HRM is about ~~0.07-07~~ Bq m⁻³ at a flow rate of about 20 L min⁻¹ with an uncertainty ~~smaller than 15% (k=2) for atmospheric ^{222}Rn levels above 2 Bq m⁻³~~³. This includes the uncertainty of the line loss correction (see below) ~~below 20% for atmospheric ^{222}Rn levels above 1 Bq m⁻³~~. Since one-filter detectors have no need for any delay chambers but use only a compact filter holder with integrated detector and pre-amplifier, the HRM is a small instrument with high portability. Regarding maintenance requirements, the quartz fiber filter should be changed monthly.

During the measurement campaign carried out at the Saclay station, where air samples were collected via a 100m Decabon tubing (see below), the line loss correction of Levin et al. (2017) was applied to all data of the HRM. No loss of aerosol was assumed in the short tubing used at Orme de Mérisiers station. Here we report for both sites ^{214}Po activity concentrations. However, for the 100 m intake height at Saclay we would not expect any disequilibrium, meaning that, based on the results from Schmithüsen et al. (2017), the reported ^{214}Po activity concentrations directly correspond to ^{222}Rn activity concentrations. By contrast, for the 2 m intake height at ODM we expect a $^{214}\text{Po}/^{222}\text{Rn}$ disequilibrium of about 0.85 to 0.9.

The second type of one-filter monitor participating in this study was built at the Laboratoire des Sciences du Climat et de l'Environnement, LSCE, France (Polian, 1986; Biraud, 2000; Schmithüsen et al., 2017). Within this manuscript this monitor will be called the LSCE monitor. This monitor uses a moving filter band system, which allows the determination of atmospheric ^{222}Rn activity concentration based on measurements of its progeny ^{218}Po and ^{214}Po . Attached ^{222}Rn progeny are collected on a cellulose filter (Pöllman-Schneider) over a one-hour period at a flow rate of 160 L min⁻¹ and after this aerosol sampling period, the loaded filter is moved to the α spectrometry for a one hour measurement period by a scintillator from Harshaw Company and photomultiplier from EMI, Electronics Ltd (Biraud, 2000). The ~~L_d minimum detection activity~~ is about 0.01 Bq m⁻³ with an uncertainty of about 20%.

Regarding maintenance on regular basis, the LSCE monitor's filter roll has to be changed every three weeks. Automatic detector background is performed every three weeks and counting efficiency is manually tested with an americium source. The instrument is designed to measure radioactive aerosols a few meters above the ground level. An inlet filter is installed to block black carbon or dirt deposition

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when the instrument is installed in urban areas as the flow rate drops below $9 \text{ m}^3 \text{ h}^{-1}$. The instrument size is about 25 cm high, 40 cm long and 25 cm deep, and it can be easily deployed at a station.

260

Monitor	Method	Sampling Flow Rate (L min^{-1})	Detection Limit (Bq m^{-3})	Typical uncertainty ($k=2$)	Portability considerations Level-and monitor-size Dimensions (cmxcmxcm) and weight (kg)	Deployability	References
ANSTO	Dual-flow-loop two-filter	~83	~0.03	~8-12%	Low- +92- m^3 300x80x80 ~120	<ul style="list-style-type: none"> Remote control Time response correction Need of large pump if the simple intake line is more than ~10m in length 	Whittlestone and Zahorowski (1998); Brunke et al. (2002); Chambers et al. [2018]
ARMON	Electrostatic deposition	~1-2	~0.076	~35-50%	Medium; 0.18- m^3 90x80x80 ~10	<ul style="list-style-type: none"> Spectrum Remote control Need of dry air simple 	Grossi et al. (2012); Vargas et al. (2015)
HRM	One-filter	~20	~0.07	~5-~20%	High; 0.08- m^3 35x30x15 ~8	<ul style="list-style-type: none"> Spectrum Remote control Sampling inlet height correction 	Levin et al. (2002)
LSCE	One-filter	~160	~0.01	~20%	High; 0.03- m^3 25x25x40 ~8	<ul style="list-style-type: none"> Spectrum Remote control Sampling inlet height correction Need of large pump 	Polian, (1986); Biraud, (2000)

Table 1. Summary of principal characteristics of the ^{222}Rn and ^{222}Rn progeny monitors compared in the present study.

2.2 Sites

The present inter-comparison study was carried out at two stations located 30 km southwest of Paris in the fall and winter of 2016-2017 (Figure 1). Both stations, 3.5 km apart, belong to the LSCE and are located in a region with a radon flux of ca. $5\text{-}10 \text{ mBq m}^{-2} \text{ s}^{-1}$ in winter, according to output of the Karsten et al. (2015) model.

Phase I of the measurements started at Orme des Mérisiers (ODM, latitude 48.698, longitude 2.146, 167 m above sea level) and ran between 25 November 2016 and 23 January 2017. Here, LSCE and ANSTO (for convenience named here as ANSTO_ODM) monitors are routinely running. During Phase I of the

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inter-comparison exercise these two monitors were operated in parallel with a HRM and an ARMON. The sampling height for all radon detectors at ODM was 2 m a.g.l.

Phase II of the exercise was realized at Saclay (SAC, latitude 48.730, longitude 2.180, Figure 1) between 25 January 2017 and 13 February 2017. At this location the sampling inlet height was at 100 m a.g.l. At SAC station another ANSTO monitor (from now on labelled as ANSTO_SAC) was already running. In addition, during Phase II this detector was running in parallel with the portable ARMON and HRM detectors. The LSCE monitor did not participate in Phase II of the experiment.

Meteorological parameters were also available at both stations during the inter-comparison periods at heights corresponding to the radon measurements (2 m and 100 m a.g.l.). In the case of the ODM site, atmospheric aerosol concentrations were also measured for this period.

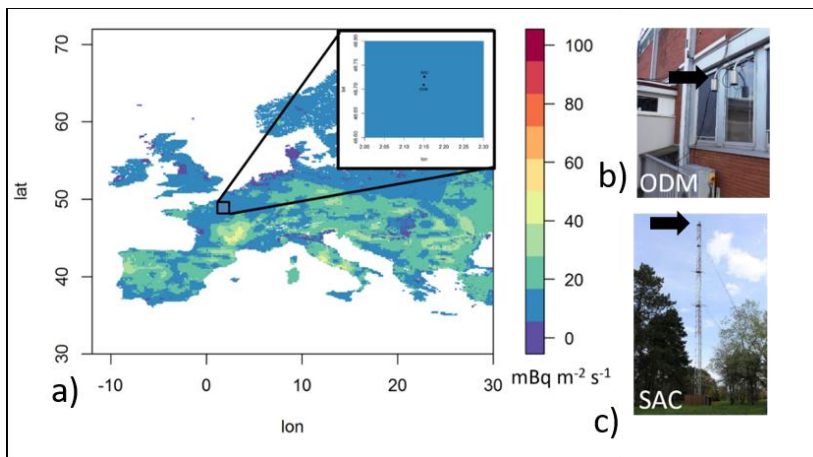


Figure 1. The INGOSv2.0 ^{222}Rn flux map (Karstens et al., 2015) is shown for a typical winter month (December), with locations of the ODM and SAC sites shown in inset (a). The radon sampling inlets are indicated both for ODM (b) and SAC (c) by the black arrows.

2.3 Environmental parameters and atmospheric aerosol concentration

Meteorological data used within this study were available from continuous measurements carried out at the SAC and ODM stations at 100 m and at 10 m a.g.l. respectively. The measurements were performed with a Vaisala Weather Transmitter WXT520 (Campbell Scientific) for: (1) wind speed and direction (accuracies of $\pm 3\%$ and $\pm 3^\circ\text{C}$, respectively); (2) Humidity and temperature (accuracies of $\pm 3\%$ and $\pm 0.3^\circ\text{C}$, respectively). In addition, the atmospheric aerosol concentration was measured at ODM site using a fine dust measurement device Fidas® 200 S (Palas) at 10 m a.g.l.. The measurement range is between 0 and 20.000 particles cm^{-3} . All the accuracies refer to the manufacturer's specifications.

2.4 Data Analysis

2.4.1 Correlation factors between monitors

To study the correlation between responses of the different detectors, linear regression models were calculated using hourly atmospheric radon activity concentrations from each monitor. The linear regression fits were calculated following Krystek and Anton (2007), relative to the two portable detectors, ARMON and HRM, because they both were measuring at SAC and at ODM.

2.4.2 Analysis of the influence of the environmental and meteorological parameters on detector response

The present study intended to build upon the findings of Xia et al. (2010) and Schmithüsen et al., (2017) regarding the possible influence of meteorological conditions on the response of radon and radon progeny monitors.

With this in mind, the ratio between hourly atmospheric ^{222}Rn activity concentrations measured and/or obtained by the HRM, LSCE and ANSTO monitors, and that measured by the ARMON were calculated, and their variability analyzed in relation to hourly atmospheric temperature, relative humidity and atmospheric aerosol concentration measured at ODM and at SAC, respectively. Not enough rain data were available to be used in this study. For this part of the study, the ARMON was used as reference being the only direct radon monitor running at both sites.

3 Results

Hourly time series of atmospheric ^{222}Rn , in the case of ARMON and ANSTO monitors, and ^{222}Rn progeny (^{214}Po activity concentration) for the HRM and LSCE monitors, measured at ODM and SAC during Phase I and Phase II of the inter-comparison experiment are presented in Figures 2 and 3, respectively. In each of the previous Figures, a zoom plot has been also reported as example to look at the response of each monitor to the sub-diurnal atmospheric radon variability. As shown, all monitors running at both sites follow this variability, with ^{222}Rn and ^{222}Rn progeny data measured or estimated by the three different measurement techniques showing the same general patterns. Table 2 summarizes the means, minima and maxima hourly atmospheric radon or radon progeny activity concentrations measured by each monitor for both campaigns. For further information, Figures S1 and S2 of the supplementary material show the time series of the differences (absolute) and of the ratios (relative) between the hourly ^{214}Po or ^{222}Rn activity concentrations measured by HRM, LSCE and ANSTO monitors and those measured by the ARMON.

3.1 Phase I: ODM site

During Phase I the LSCE, HRM, ARMON and ANSTO_ODM monitors were operating in parallel, sampling air from the same height (2 m a.g.l.). The mean temperature over Phase I of the campaign was 2.9 °C with an interquartile range of 0.10 °C to 5.8 °C. The mean relative humidity was 80% with an interquartile range of 73% to 89%. An average accumulated rain per day of 13 mm was recorded. The main wind patterns during Phase I were from northeast and southwest, with speeds typically between 1 and 7 m s⁻¹. The mean atmospheric aerosol concentration observed at ODM during Phase I was 505 particles cm⁻³ with an interquartile range of 233 cm⁻³ to 660 cm⁻³.

The means of the atmospheric ^{222}Rn activity concentration measured by the ARMON and the ANSTO_ODM are in the same order (Table 2). The means of the atmospheric ^{214}Po activity concentrations measured by LSCE monitor were ca. 50% lower and by the HRM ca. 30% lower than the atmospheric ^{222}Rn activity concentration.

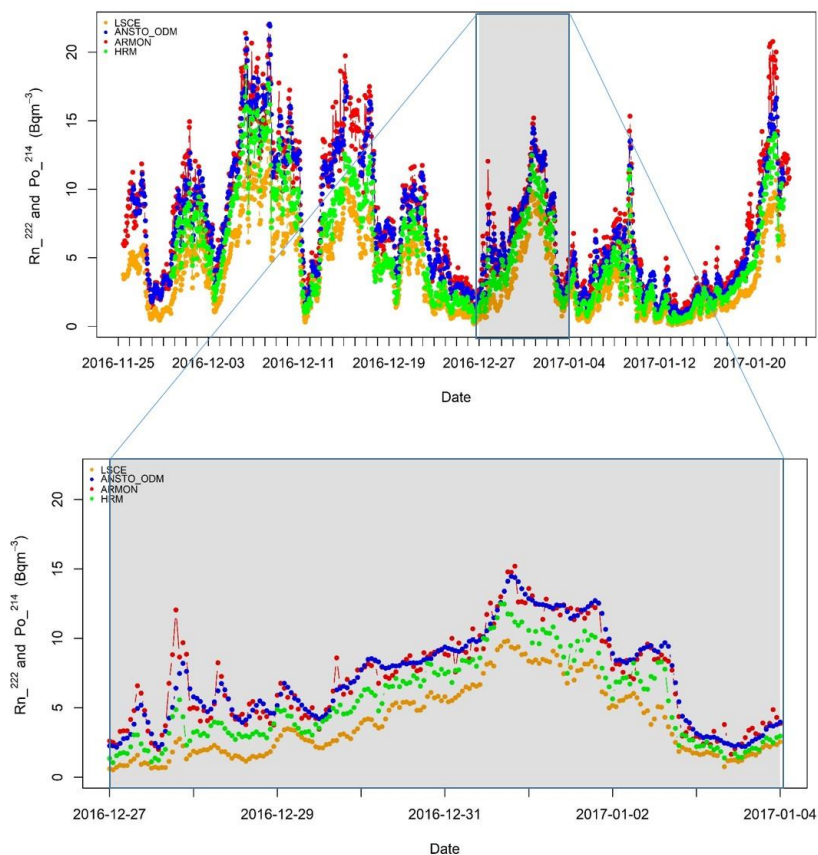


Figure 2. Main panel: Hourly time series of the atmospheric ^{222}Rn and, in the case of LSCE and HRM data ^{214}Po activity concentration, measured at Orme de Merisiers (ODM) station during Phase I (between 25 November 2016 and 23 January 2017) by: ARMON (red circles), ANSTO_ODM (blue circles), HRM (green circles) and LSCE (orange circles) monitors. Zoomed panel: Hourly time series of the atmospheric ^{222}Rn and ^{214}Po measured between 27th December 2016 and 04th January 2017.

Table 2 shows the slopes (b) and intercepts (a) of the linear regression fits calculated between the hourly atmospheric ^{222}Rn and ^{214}Po activity concentrations measured by the ARMON and/or the HRM and the other ^{222}Rn and ^{222}Rn progeny monitors deployed in Phase I. The calculated slopes were in the range of

0.62 to 1.17 and the R^2 values varied between 0.90 and 0.96. The slope closest to unity was calculated between the ARMON and ANSTO_ODM monitors, and was 0.96 ± 0.01 , while the lowest slope was observed between the ARMON and LSCE monitors, and was 0.62 ± 0.01 . The highest correlation ($R^2=0.96$) was found between the HRM and LSCE monitors. The plots of the linear regression fits of the Phase I are shown in the left panels of the Figures S3, S4 and S5 of the supplementary material. Notably, the offset (a value) of the regression between the ANSTO and ARMON detectors at ODM is considerably greater than that at SAC. The regression slopes are also slightly different. These differences are likely related to the limited calibration and background information available for the ANSTO_ODM detector for this inter-comparison project. In particular, a substantial component of the instrumental background signal is site specific. This is likely responsible for much of the change in offset value.

	Monitors Phase I	Mean (Bq m ⁻³)	Min/Max (Bq m ⁻³)	x					
				b (ARMON)	a (ARMON)	R^2 (ARMON)	b (HRM)	a (HRM)	R^2 (HRM)
y	ANSTO_ODM	7.02	0.73/22.04	0.96 ± 0.01	-0.23 ± 0.03	0.94	1.17 ± 0.01	0.63 ± 0.03	0.93
	HRM	5.45	0.26/18.91	0.82 ± 0.01	-0.71 ± 0.03	0.93	-	-	-
	ARMON	7.55	0.50/21.98	-	-	-	-	-	-
	LSCE	3.84	0.10/14.93	0.62 ± 0.01	-0.85 ± 0.03	0.90	0.76 ± 0.004	-0.29 ± 0.03	0.96
	Monitors Phase II	Mean (Bq m ⁻³)	Min/Max (Bq m ⁻³)	Slope (ARMON)	Intercept (ARMON)	R^2 (ARMON)	Slope (HRM)	Intercept (HRM)	R^2 (HRM)
	ANSTO_SAC	3.50	0.43/10.71	0.97 ± 0.01	0.01 ± 0.06	0.95	1.03 ± 0.01	0.15 ± 0.06	0.90
	HRM	3.26	0.26/11.15	0.94 ± 0.01	-0.13 ± 0.06	0.91	-	-	-
	ARMON	3.60	0.17/11.51	-	-	-	-	-	-

Table 2. The means, maxima, and minima of the atmospheric ^{222}Rn and ^{214}Po activity concentration observed by each monitor participating in the Phase I and II of the inter-comparison campaigns. The slopes (b) and intercepts (a) of the linear regression fits calculated between the hourly atmospheric ^{222}Rn and ^{214}Po activity concentrations measured by the ARMON and/or the HRM and the other ^{222}Rn and ^{222}Rn progeny monitors deployed in both phases are also reported.

3.2 Phase II: SAC station

Phase II lasted 18 days. The mean temperature during this period was 5 °C with an interquartile range of 2 °C to 8 °C. The mean relative humidity was 86% with an interquartile range of 80% to 94%. An average accumulated rain per day of 3 mm was recorded. The main wind patterns during this phase at 100 m a.g.l. were from the south and southwest with speeds typically between 3 and 10 m s⁻¹.

Figure 3 shows the hourly atmospheric ^{222}Rn and ^{214}Po activity concentrations observed at SAC during Phase II by the ARMON, HRM and ANSTO_SAC instruments.

Table 2 reports the means, minima, and maxima of the atmospheric data measured during Phase II by all participating monitors. In this case, the mean atmospheric ^{222}Rn and ^{214}Po activity concentrations measured by all monitors agreed within the instrumental errors. At 100 m a.g.l. the slopes of the hourly fits of the monitor's response in this case were all close to unity. The calculated offsets also decreased at 100 m a.g.l. relative to 2 m a.g.l. The plots of the linear regression fits of Phase II are shown in the right panel of Figures S5 and S6 of the supplementary material. During the period of Jan 30 – February 1, 2019, the HRM shows significantly lower values than the ANSTO and ARMON. This period coincides with saturated air humidity conditions.

Figure S7 of the supplementary material presents two plots to summarize the results of the slopes and offsets calculated both at ODM and SAC stations relative to the ARMON.

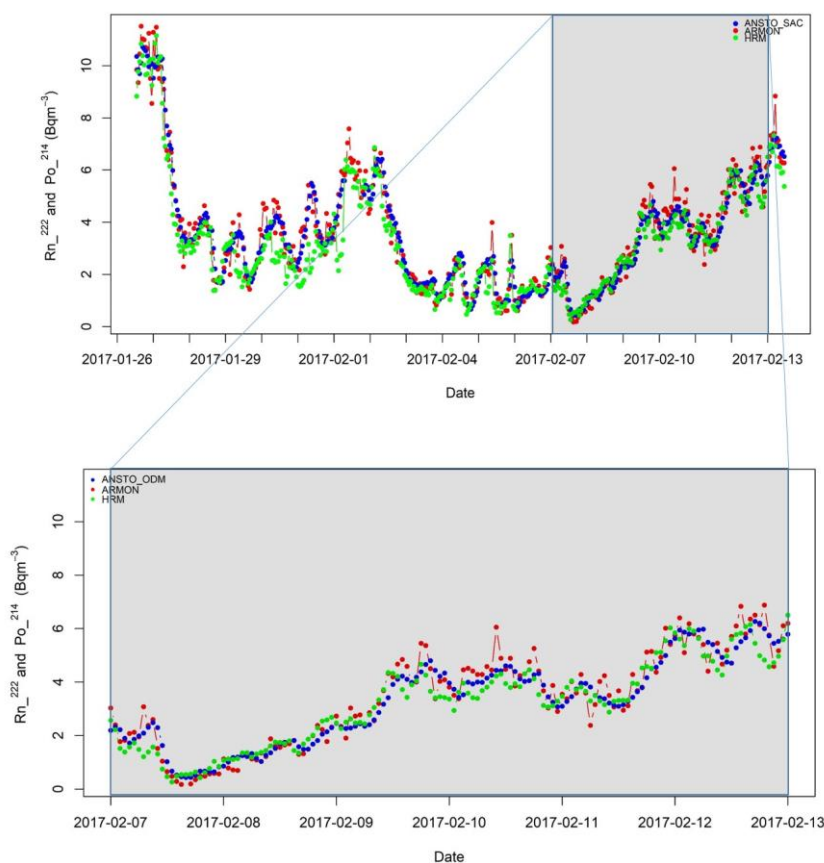


Figure 3. Main panel: Hourly time series of the atmospheric ^{222}Rn and ^{214}Po (HRM) activity concentration measured at Saclay (SAC) station between 25 January 2017 and 13 February 2017 by: ARMON (red circles), ANSTO_SAC (blue circles) and HRM (green circles) monitors. Zoomed panel: Hourly time series of the atmospheric ^{222}Rn and ^{214}Po measured between 7 February 2017 and 13 February 2017.

Figure 2 and 3 show a larger hourly variability of the HRM and ARMON signals compared with the ANSTO ones. This difference in variability is likely due to a larger uncertainty of the HRM and ARMON detectors for atmospheric ^{222}Rn levels of around 1 Bq m^{-3} . In addition, and that it has to be taken into account that only an approximated form of the Griffiths et al. (2016) response time correction at the top of the ANSTO has been used to improve the information in the ANSTO data. The main reason for this is that the ANSTO data are not corrected for the response time of the detectors. This is a limitation of the ANSTO data and it should be taken into account in future investigations.

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exactly quantify the detectors uncertainties for the low ^{222}Rn concentrations typical for outdoor environmental monitoring at or above 100 m a.g.l.

3.2 Comparison with past studies

The results obtained in the present study of the slopes (b) and of the offsets (a) of the regression lines calculated between ANSTO or LSCE monitors against the HRM are here compared with the ones presented by Schmithüsen et. al., 2017. Table 3 shows a summary of this comparison. All slopes (correction factors) are defined as (routine station monitor) / HRM because this last was used as reference instrument by Schmithüsen et. al., 2017.

Site Input Height	Schmithüsen et al., 2017			Present study		
ANSTO/HRM	Activity Range (Bq m^{-3})	b	a	Activity Range (Bq m^{-3})	b	a
Cabauw: 200/180 m	0-8	1.11 ± 0.04	0.11 ± 0.06			
Saclay: 100 m				0-11	1.03 ± 0.01	0.15 ± 0.06
Lutjewad: 60 m	0-6	1.11 ± 0.02	0.11 ± 0.02			
Heidelberg: 35 m	0-15	1.22 ± 0.01	0.42 ± 0.04			
Cabauw: 20 m	0-12	1.30 ± 0.01	0.21 ± 0.03			
Orme des Mérisiers: 2 m				0-22	1.17 ± 0.01	0.63 ± 0.03
LSCE/HRM	Activity Range (Bq m^{-3})	b	a	Activity Range (Bq m^{-3})	b	a
Orme des Mérisiers: 2 m	0-9	0.68 ± 0.03	-0.18 ± 0.09	0-15	0.76 ± 0.01	-0.29 ± 0.03

Table 3. Offsets and slopes of the regression lines calculated between ANSTO or LSCE monitors against the HRM in the present study and by Schmithüsen et. al., 2017.

Data in Table 3 need to be analysed taking into account that a unique traceability chain is not yet available for atmospheric radon measurements and the different monitors routinely running at the different stations could have different calibration chains (e.g. radon source, primary standard, etc.). Generally speaking, for both studies, it can be observed that the correction factor between the atmospheric ^{214}Po activity concentration measured by HRM and the atmospheric ^{222}Rn activity concentration measured by ANSTO at each station approaches unity with the increase of the height of the sampling input. By contrast, the offsets of the regression fits decrease with the increase of the input height.

The only case where the compared instruments were exactly the same and at the same height is for Orme des Mérisiers station. Here the slope between the atmospheric ^{214}Po activity concentration measured by LSCE and HRM is equal to 0.76 ± 0.01 . This number is slightly larger but within uncertainties well comparable to the number reported by Schmithüsen et al. (2017) of 0.68 ± 0.03 (see Table 3).

3.4 Influence of the weather conditions on the ratio between ^{214}Po and ^{222}Rn measurements

Figure 4 shows the variability of the ratio between hourly atmospheric ^{214}Po and/or ^{222}Rn activity concentration measured by each monitor relative to those measured by the ARMON versus the hourly means of ambient temperature and relative humidity. Analysis was carried out at ODM (Figure 4, upper panels) and at SAC (Figure 4, bottom panels) versus ambient temperature (Figures 4, left panels) and relative humidity (Figures 4, right panels) measured at the corresponding stations.

Figure 5 shows the same variability plotted in relation to the ANSTO_ODM at ODM (Figure 5, upper panels) and to the ANSTO_SAC at SAC (Figure 5, bottom panels) versus the hourly means of ambient temperature (Figures 5, left panels) and relative humidity (Figures 5, right panels).

Data does not show any evident patterns at 100 m a.g.l. (SAC station), which could indicate that there is any impact on ^{222}Rn or ^{222}Rn progeny measurements due to change of ambient temperature and relative humidity, at least not until saturated conditions are achieved. By contrast, a small decrease, of about $10^{-2} \text{ }^{\circ}\text{C}^{-1}$, is observed in the ratio between the ^{214}Po activity concentration (measured by HRM and LSCE monitors) and the ^{222}Rn activity concentration (measured by ANSTO_ODM and ARMON monitors) with the increase of the ambient temperature (Figure S8 of the supplementary material) at 2 m a.g.l. (ODM station). This temperature dependency may be rather due to the effect of atmospheric activity concentrations, increasing during nighttime, on the disequilibrium between radon and its progeny. However, this influence on measured $^{214}\text{Po}/^{222}\text{Rn}$ ratios is really small compared with others observed effects (e.g.: loss of progeny within the sample tube (Levin et al., (2017)), atmospheric aerosol concentration (see below)). Looking at Figure 5, there appears to be less scatter in the point clouds (particularly at SAC) when the ANSTO_SAC monitor is used as the reference, likely attributable to the lower measurement uncertainty of the ANSTO monitor used at this station.

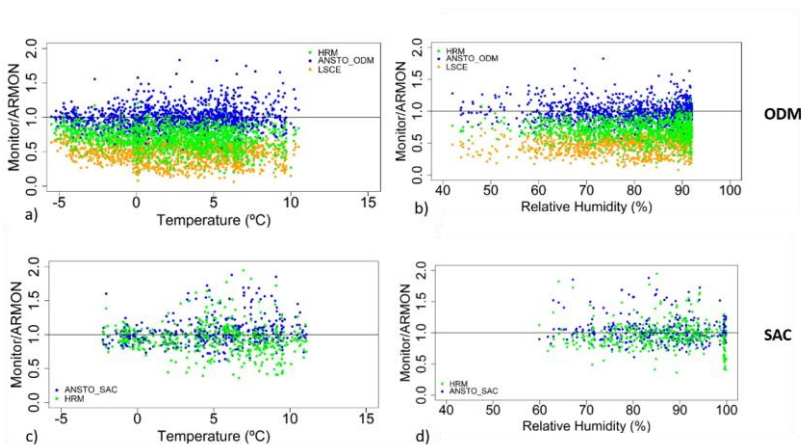


Figure 4. Hourly atmospheric ^{222}Rn or ^{214}Po activity concentration obtained by HRM, LSCE and ANSTO monitors divided by the ^{222}Rn activity concentration measured by the ARMON detector as function of the hourly measured atmospheric temperature and relative humidity at ODM (a and b) and at SAC (c and d), respectively.

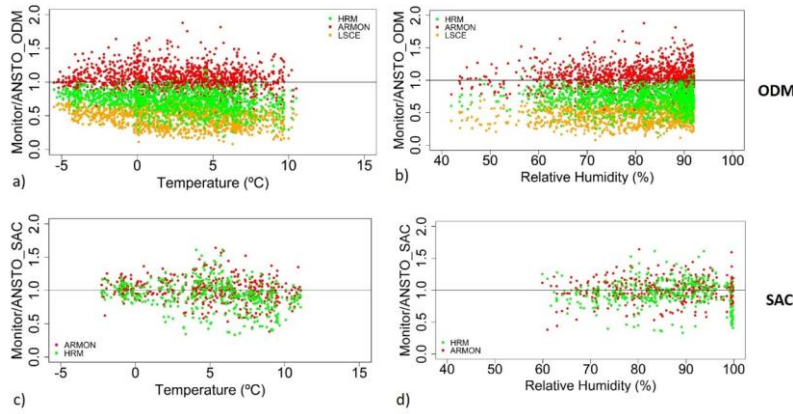


Figure 5. Hourly atmospheric ^{222}Rn or ^{214}Po activity concentration obtained by ARMON, HRM and LSCE monitors divided by the ^{222}Rn activity concentration measured by the ANSTO detectors as function of the hourly measured atmospheric temperature and relative humidity at ODM (a and b) and at SAC (c and d), respectively.

In Figure 6 the ratio of the hourly atmospheric ^{222}Rn or ^{222}Rn progeny activity concentration measured by the HRM (^{214}Po in Figure 6a), the LSCE (^{214}Po in Figure 6b) and the ANSTO_ODM (^{222}Rn in Figure 6c) monitor and the ^{222}Rn activity concentration measured with ARMON (^{222}Rn) are plotted against the logarithm of the hourly aerosol concentration data. Data indicate the existence of a linear relationship between these variables, i.e. of the form:

$$\frac{^{222}\text{Rn}(\text{Monitor}_i)}{^{222}\text{Rn}(\text{ARMON})} = a + b \cdot \text{Log}_{10}(\text{Aerosol Conc.}). \quad (1)$$

Here $^{222}\text{Rn}(\text{Monitor}_i)$ is the hourly atmospheric ^{222}Rn or ^{214}Po activity concentration measured by individual monitors HRM (^{214}Po), LSCE (^{214}Po) and ANSTO_ODM (^{222}Rn), $^{222}\text{Rn}(\text{ARMON})$ is the one measured by the ARMON monitor and *Aerosol Conc.* is the hourly atmospheric aerosol concentration measured at ODM during Phase I. The results of the linear regression fits are reported in Table 4. The slope of the ratio between the ANSTO_ODM and ARMON monitors in relation to the variability of the logarithm of the hourly atmospheric aerosol concentration is close to zero and the intercept is close to one. The ratio between the hourly atmospheric aerosol-bound radon progeny data measured by the two one-filter radon progeny monitors and the one measured by the ARMON seems to decrease with decreasing aerosol concentration (Figures 6a and 6b). However, this effect becomes only evident when atmospheric aerosol concentration is lower than 300 particles cm^3 .

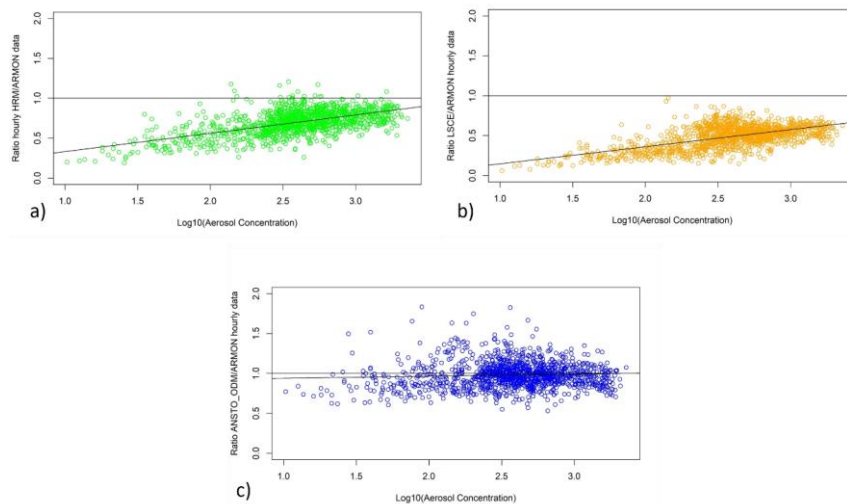


Figure 6. Ratio of the atmospheric ^{222}Rn or ^{214}Po activity concentration measured by the HRM (green dots), LSCE (orange dots) and ANSTO_ODM (blue dots) monitors and those measured by the reference ARMON monitor against the logarithm of the atmospheric aerosol concentration measured at ODM station.

Monitor	a	b	R^2
HRM	0.10 ± 0.02	0.23 ± 0.01	0.34
LSCE	-0.07 ± 0.02	0.21 ± 0.01	0.34
ANSTO_ODM	0.91 ± 0.03	0.03 ± 0.01	$0.04 \cdot 10^{-1}$

Table 4. Intercepts and slopes of the linear regression fits of the Equation 1

Conclusions

In order to confirm and build upon the results obtained by Xia et al. (2010), Grossi et al. (2016) and Schmithüsen et al. (2017) a three month inter-comparison campaign was carried out in the south of Paris, France, in the fall-winter period of 2016-2017. For the first time, three fundamentally distinct radon and radon progeny measurement approaches were compared side-by-side at two measurement heights: 2 and 100 m a.g.l., under a range of environmental conditions with the aim to compare their responses.

The results of this study show that ^{222}Rn and ^{222}Rn progeny measurements follow the same general patterns of diurnal variability, both close to and further up from the surface. The slopes and intercepts of the linear regression fits between the radon and the radon-progeny measurements, which represent the calibration factors, are not significantly different from one at 100m height above ground (SAC), but they differ at the 2_m level (ODM). The last behavior latter is attributable to the disequilibrium known to exist between ^{222}Rn freshly emitted from the ground and its short-lived progeny in the lowest 10s of meters of the atmosphere, the magnitude of which is known to decrease with distance from the surface, as shown in earlier work, and to be close to one at a height of 100_m and above (e.g. Jacobi and André, 1963; Schmithüsen et al., 2017).

For the 2 m level, we found a ~~very good~~ significant correlation of radon progeny activity concentrations between LSCE and HRM measurements (see Figure S3 in the Supplement). The slope, however, is only equal to 0.76 ± 0.01 . This ~~result number is slightly larger but is comparable, considering its within~~ uncertainties, ~~with the result well comparable to the number~~ reported by Schmithüsen et al. (2017) of 0.68 ± 0.03 (see Table 3) ~~based on the comparison of the same two monitors (HRM and LSCE) and at the same station (ODM) -based on a comparison campaign conducted at ODM-~~ in March and April 2014.

Observations of the total atmospheric aerosol concentration available at ODM station during the first two months of the experiment were used to investigate the influence of changing atmospheric aerosol concentrations on the response of the radon/radon progeny measurements. Under very low atmospheric aerosol loading (< 300 particles cm^{-3}), the ^{222}Rn progeny monitors seem to underestimate the atmospheric ^{214}Po activity concentrations by up to 50%. This effect may be attributable to loss of un-attached ^{218}Po and ^{214}Po . Particle number concentrations below 300 particles cm^{-3} at continental stations are, however, very rare and even during winter at Alpine stations like Schneefernerhaus such low particle concentrations are only occasionally observed (Birmili et al., 2009).

The comparison of the results obtained in the present study with the ~~ones those~~ reported in Schmithüsen et al. (2017) ~~put in evidence demonstrate that in order to underlines that to assure the harmonization of the~~ atmospheric ^{222}Rn activity concentrations measured at ~~different~~ atmospheric networks ~~it will be important is important~~ to: i) have a well-established metrological chain; ii) ~~have traceable methods for measuring low-level atmospheric radon activity concentrations; iii) harmonize the calculation of the total expanded uncertainty of in~~ atmospheric ^{222}Rn concentrations measured by all monitors when ambient radon is only a few Bq m^{-3} or less; ~~iiiiv) use as a direct radon monitor as a mobile reference instrument, a direct radon monitor which the~~ response of which is not influenced by meteorological conditions or inlet tube dimensions and length.

Finally, the new portable ARMON seems to have a great potential for being used ~~within atmospheric radon networks at atmospheric radon stations with space restrictions. It could be also be useful as mobile reference instrument to calibrate~~ ^{222}Rn progeny instruments or fixed radon monitors. However the ~~n-order to deeply evaluate the total expanded uncertainty of the ARMON could increase response for really low radon activity concentration (< 1 Bq m^{-3}) and when atmospheric ^{220}Rn is also present. This should be better investigated in the near future. the As should being investigated the uncertainties related to the ANSTO detector response time correction when characteristics of the entire intake system have not been directly measured. response time correction periods of rapid also furtherd qualities and faults of this new instrument a long term inter-comparison study should be carried out using a direct ANSTO instrument.~~

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